# 2. SITE 135

#### The Shipboard Scientific Party<sup>1</sup>

#### ABSTRACT

Site 135 lies about 35 km southeast of the southern edge of the Horseshoe Abyssal Plain on a topographic high 750 meters above the abyssal plain. The top unit, 325 meters thick, is comprised of nannoplankton chalk ooze of Pleistocene, Pliocene, and Miocene ages. Below this, there are 364 meters of mostly terrigenous sediments with some silicified intervals and marl or limestone at the base which range in age from Early Eocene to Early Aptian. It is not possible to reasonably estimate the thickness of individual lithologic units within this lower interval.

A prominent reflecting horizon at  $\sim 0.4$  second can be traced from beneath the abyssal plain on to the topographic high. In the drilled section this corresponds to a major unconformity which marks the abrupt change from terrigenous to pelagic sedimentation that occurred following post-early Eocene and pre-late Oligocene uplift and faulting.

An estimated 350+ meters of sediment are present between the lowest sample obtained and oceanic basement.

#### SITE DATA

Time: 0053 October 10, 1970 0146 October 14, 1970

Position: 35° 20.80'N 10° 25.46'W

- Water Depth: 13,621 feet
  - 2,270 nominal fathoms 4,152 meters

#### Total Penetration: 689 meters

Cores Taken: Nine cores and two sidewall samples in Hole 135; one core (no recovery) in Hole 135A.

#### **BACKGROUND AND OPERATIONS**

Because there was no detailed survey data available, the *Glomar Challenger* approached the site from the northnorthwest along the existing *Jean Charcot* track (see Figure 1). Seismic reflection profiles from the *Jean Charcot* (Figure 2) provided the primary control for site selection.

The initial objectives at Site 135 were threefold: (1) to sample the Tertiary biostratigraphic sequence for comparison with European type-sections; (2) to date the major intermediate reflecting horizon at about 400 meters subbottom; and (3) to sample the basement (at a subbottom depth of  $\sim 1000$  m) and the sediments immediately overlying it in order to date the oceanic crust near the eastern North Atlantic margin.

The site lies about 35 km southeast of the Horseshoe Abyssal Plain on a topographic high about 750 meters (400 fm) above the abyssal plain, and south of the eastward extension of the seismically active Azores-Gibraltar Fracture Zone (see Figure 1 and map at the back of this volume).

Acoustic basement reaches a minimum depth of burial of about 1000 m near this location (see Figures 2 and 3). A widespread reflecting horizon, generally conformable with the sea floor, is offset by high angle faults and interrupted by piercement structures. The piercement structures have been interpreted as salt diapirs (Pautot *et* al., 1970).

Site 120 of Leg 13 (DSDP) was located about 170 km northwest of Site 135, on the north slope of Gorringe Bank and north of the Azores–Gibraltar fracture zone (see Figure 1). Two major unconformities, one between Pleistocene and Miocene and the other between Miocene and Cretaceous, have been inferred at Site 120 by the scientific staff Leg 13 (Ryan and Hsu *et al.*, 1972).

| Seismic Reflection Data: | Jean Charcot | Glomar Challenger |
|--------------------------|--------------|-------------------|
| First Reflector          | 0.40 sec     | 0.38 sec          |
| Second Reflector         | 0.75 sec     | 0.71 sec          |
| Basement Reflector       | 1.03 sec     | 0.85 sec (not     |
|                          |              | clearly recorded) |

A summary of the drilling and coring record at Hole 135 is given in Table 1 and Figure 4. Changes in drilling rates agree well with the lithostratigraphy as deduced from cored intervals (Figure 4).

Drilling was terminated at a total depth of 689 meters subbottom before acoustic basement was reached. The exact depth to basement, while unknown, is estimated to be over 1000 meters judging from the *Jean Charcot* records and sonic velocity measurements on core samples.

Two sidewall samples were taken to supplement the previously cored intervals. Although partial samples were

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Figure 1. Location map for Site 135. Contours are in nominal fathoms from unpublished maps of A. S. Laughton (pers. comm.). Letters key profiles in Figure 2. Site 120 was drilled on Leg 13, Deep Sea Drilling Project.



Figure 2. Seismic profiles in vicinity of Site 135. Location of profiles is shown on Figure 1. Jean Charcot seismic record (vertical exaggeration is 20:1) is from X. LePichon et al., 1970.

TABLE 1 Drilling and Coring Record for Site 135

| Description | Interval Below<br>Sea Floor<br>(m)   | Core Recovery<br>(m) | Drilling Rate<br>(m/min)   |
|-------------|--|----------------------|--|
| Core 1      | 0-4  | 4.0                  |  |
| Drill       | 4-8<br>8-27<br>27-46<br>46-65  |                      | 0.7<br>2.0<br>1.6<br>0.9   |
|             | 65-80  |                      | 2.5  |
| Core 2      | 80-89  | 8.5                  |  |
| Drill       | 89-104<br>104-123<br>123-132<br>132-151<br>151-161<br>161-173  |                      | 1.9<br>2.0<br>2.25<br>1.5<br>1.7<br>1.2  |
| Core 3      | 173-182  | 2.5                  |  |
| Drill       | 182-199<br>199-209<br>209-228<br>228-237<br>237-259  |                      | 1.4<br>1.7<br>1.9<br>2.25<br>3.7   |
| Core 4      | 259-268  | 2.1                  |  |
| Drill       | 268-285<br>285-294<br>294-313<br>313-323<br>323-335  |                      | 0.5<br>0.45<br>0.5<br>0.7<br>0.5   |
| Core 5      | 335-341  | 0.7                  |  |
| Core 6      | 341-350  | 0.1                  |  |
| Drill       | 350-352<br>352-361<br>361-370<br>370-379<br>379-389<br>389-399<br>399-418<br>418-431                       |                      | $\begin{array}{c} 0.5 \\ 1.5 \\ 1.5 \\ 1.1 \\ 0.8 \\ 1.7 \\ 1.2 \\ 1.1 \end{array}$                |
| Core 7      | 431-435  | 3.8                  |  |
| Drill       | 435-437<br>437-457<br>457-476<br>476-485<br>485-495<br>495-514<br>514-524<br>524-543<br>543-552<br>552-564 |                      | 0.3<br>0.5<br>0.5<br>0.45<br>0.5<br>0.6<br>0.5<br>0.6<br>0.7<br>1.0                                |
| Core 8      | 564-569  | 1.0                  |  |
| Drill       | 569-571<br>571-580<br>580-589<br>599-617<br>617-627<br>627-646<br>646-656<br>656-676<br>676-685            |                      | $\begin{array}{c} 0.2 \\ 0.4 \\ 0.45 \\ 0.6 \\ 0.4 \\ 0.6 \\ 0.3 \\ 0.2 \\ 0.2 \\ 0.3 \end{array}$ |
| Core 9      | 685-689  | 1.8                  |  |
|             | Water Depth  | 4,152 meters         |  |

obtained, minor difficulties were encountered in recovering the corer. These problems were attributed to the failure of the sampler to open fully after passing through the drill bit.

# BIOSTRATIGRAPHY

## General

Four widely spaced cores (1-4) indicate that the Quaternary to Late Oligocene section, in the upper 325 meters of this hole, consists of deep-water pelagic sediments with strong evidence of calcium-carbonate solution, particularly in the Miocene section. Within Core 2 there is a minor faunal break which may correspond to the Miocene/ Pliocene boundary. Below a possible hiatus at about 330 meters and down to total depth of 689 meters, the rocks are mostly clayey sand, with generally poorly preservedmicro-fossils. These fossils indicate an Eocene to Cretaceous age but give little precise data on the sedimentary history during this time. It is possible that the relatively rich Upper Cretaceous fossil assemblages of Core 7 and those obtained from a center bit sample near Core 9 were redeposited in Eocene time. The deepest samples at this Site (Core 9) contain a nannoflora of Early Aptian age. A poor sporomorph flora and microplankton of Late Cretaceous age (Senonian or younger) occurs in Core 7. Sporomorphs and microplankton from Core 9 are of Valanginian to Aptian age.

The age-diagnostic fossils from each core are shown in Table 2.

#### Foraminifera

Cores 1 to 4 contain normal pelagic deep-sea faunas consisting predominantly of planktonic foraminifera. Solution effects are visible throughout but are strongest in Cores 3 and 4. Below Core 4, there is little conclusive information that can be drawn from the foraminifera except that some (if not all) of the Eocene and Late Cretaceous faunas are probably transported and re-deposited. Core 5 contains a single specimen of a (Eocene?) nummulite. Another specimen of Nummulites? was found associated with the relatively rich Globotruncana assemblage of Core 7. Similarly, the large center bit sample recovered near Core 9 contained, apart from Late Cretaceous Orbitoids, a few specimens (Nummulites?, Operculina, Amphistegina) which are probably of Eocene age. These larger foraminifera were derived from a shallow-water environment since they are not normally found in pelagic sediments.

#### Nannoplankton

Rich and diversified assemblages of calcareous nannoplankton of Pleistocene to Late Miocene age were found in the uppermost 3 cores. Core 4 contains a fairly strongly etched Middle Miocene nannoflora. The two sidewall cores yielded fairly poorly preserved Upper Oligocene to Lower Miocene nannoplankton. Only the most solution resistant ortholiths were preserved in the Early Eocene of Core 5. The Campanian to Maestrichtian nannoplankton in Core 7 lack distinctive markers. Core 8 is barren of nannofossils. Core 9 contains a rich and well-preserved nannoflora of Early Aptian age.



Figure 3. Geological synthesis at Site 135.

TABLE 2

| ORE | DIAGNO   | DSTIC FOSSILS HOLE 135   |                   |
|-----|--|--|-------------------|
| ŏ   | FORAMINIFERA   | NANNOPLANKTON  | AGE               |
| 1   | Rich, predominantly planktonic faunas of<br>Quaternary age (Globorotalia truncatulinoides<br>Zone). The predominant species are:<br>Globorotalia truncatulinoides, Gr. inflata,<br>Gr. pseudopima and Globigerinoides ruber.   | Rich assemblage with Gephyrocapsa oceanica,<br>Coccolithus pelagicus<br>Preservation: G to El-Ol<br>Zone: Gephyrocapsa oceanica<br>Age: Late Pleistocene   | Pleisto-<br>cene  |
| 2   | Rich, predominantly planktonic faunas. Sections<br>1 to 5 belong to the Globorotalia margaritae<br>Zone (Early Pliocene) and contain Globorotalia<br>margaritae (small), Gr. orassaformis s.l.,<br>Gr. acostaensis (dextrally coiled),<br>Sphaeroidinellopsis seminulina and<br>Globigerinoides obliquus.<br>The bottom part of the Core (without<br>Globorotalia margaritae) contains<br>Globorotalia miocea and may already be Late<br>Miocene in age. | Sections 1-5: Diversified nannoflora with<br>Ceratolithus rugosus, Ceratolithus tricorni-<br>culatus, Reticulofenestra pseudoumbilica,<br>Discoaster broweri, D. surculus, D.<br>pentaradiatus<br>Preservation: E1-01<br>Zone: Ceratolithus rugosus<br>Age: Early Pliocene | Early<br>Pliocene |
| 3   | Planktonic and benthonic foraminifera are<br>fairly common. They indicate a Late Miocene<br>age (probably Globorotalia acostaensis Zone).<br>Typical species are Globorotalia acostaensis,<br>Gr. cultrata, Globoquadrina dehiscens,<br>Globigerina nepenthes, and Globigerinoides<br>obliquus. Rather strong solution effects.  | Rich and well-preserved nannoplankton with<br>Discoaster neohamatus, D. calcaris, D.<br>variabilis, Reticulofenestra pseudoumbilica<br>Preservation: El-Ol<br>Zone: Discoaster neohamatus<br>Age: Late Miocene   | Late<br>Miocene   |
| 4   | Fairly rich faunas, but distinctly affected by calcium carbonate solution. An early Middle Miocene age (Globorotalia fonsi peripheroronda Zone) is indicated by the presence of Gr. f. peripheroronda (with subacute periphery), together with Gr. archaeomenardii, Sphaeroidinellopsis seminulina and Globoquadrina dehiscens.  | Abundant but poorly preserved nannofossils.<br>Discoaster exilis, D. deflandrei, D.<br>divaricatus, Sphenolithus heteromorphus<br>Preservation: E1-02<br>Zone: Sphenolithus heteromorphus<br>Age: Middle Miocene   | Middle<br>Miocene |
| SW1 | A poor fauna with <i>Globigerina venezuelana</i> and<br><i>Catapsydrax?</i> sp.<br>Strong downhole contamination from the Pliocen <del>e</del><br>and Quaternary.  | Strongly etched nannoflora with Triquetrorhab-<br>dulus carinatus and Reticulofenestra abisecta<br>Preservation: E3-01<br>Zone: Triquetrorhabdulus carinatus<br>Age: Early Miocene   | Early<br>Miocene  |
| SW2 | A poor, badly contaminated sample. An Early<br>Miocene or possibly Oligocene age is indicated<br>by Catapsydrax dissimilis together with<br>Globigerina cf. cipercensis angustiumbilicata,<br>G. venezuelana, and Globorotalia opima nana.   | Same assemblage as in SW 1 plus Reticulofenestro<br>bisecta, rare Sphenolithus belemnos (small<br>specimen)<br>Zone: Triquetrorhabdulus carinatus<br>Age: Late Oligocene   | Late<br>Oligocene |
| 5   | No foraminifera except a badly preserved small <i>Nummulites</i> (?) sp.   | Core catcher:<br>Assemblage consisting of solution resistant<br>species only. Marthasterites tribrachiatus,<br>M. contortus, D. lodoensis<br>Zone: Marthasterites tribrachiatus  | Early<br>Eocene   |
|     |  | Age: Early Eocene  |                   |
| 6   | None   | None   |                   |
|     |  |  |                   |
|     |  |  |                   |

 TABLE 2 - Continued

| ORE | DIAGNO   | DSTIC FOSSILS HOLE 135   |                    |  |
|-----|--|--|--------------------|--|
| ö   | FORAMINIFERA   | NANNOPLANKTON  | AGE                |  |
| 7   | The richest faunas seem to occur in the sandy<br>layers. Typical species are Globotruncana<br>stuartiformis, Gt. stuarti, Gt. contusa<br>Ft. caliciformis, Heterohelix ultimatumida,<br>Pseudoguembelina excolata, and Gyrcidina<br>florealis. They are of Late Cretaceous<br>(Maestrichtian) age. A single specimen of<br>Nummulites? sp. was found in Sec. 2, indicating<br>that the Cretaceous faunas were possibly re-<br>deposited in Eccene? time. (Similar, slightly<br>heterogeneous faunas, but with Cretaceous<br>orbitoids, were found in a center bit<br>sample in this hole). | Poor assemblages with Watznaueria barnesae,<br>Arkhangelskiella cymbiformis, Micula<br>staurophora, Tetralithus obscurus,<br>Eiffellithus turriseiffeli, Microrhabdulus<br>decoratus, Broinsonia bevleri.<br>Age: Campanian to Maestrichtian   | Maestrich-<br>tian |  |
| 8   | None   | None   | · · ·              |  |
| 9   | None   | Rich and well-preserved nannoflora including<br>Nannoconus colomi, Corollithion ellipticum,<br>Eiffellithus trabeculatus, Corollithion<br>achylosum, Cruciellipsis cuvillieri,<br>Micrantholithus obtusus, Cretarhabdus<br>coronadventis, Nannoconus bucheri, Micrantho-<br>lithus hoschulzi, Diazornatholithus lehmani,<br>Nannoconus steinmanni, Cyclagelosphaera<br>margareli.<br>Age: Early Aptian | Early<br>Aptian    |  |

#### **Organic Microfossils**

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Core 7, Section 2, 138-140 cm:

A poor sporomorph flora containing many Angiospermous pollen types such as: Subtriporopollenites anulatus, Multiporopollenites aff. maculosus, and several small tricolporate and triporate forms, together with Classopollis classoides, C. echinatus, Ephedripites sp., Trilobosporites cf. obsitus, Cicatricosisporites sp., and Lygodioisporites perverrucatus.

Age: Late Cretaceous.

Microplankton scarce and mainly indeterminate; the presence of *Aeroligera* sp. indicates "Senonian or younger."

#### Core 8, Section 1, 106-110 cm:

No sporomorphs were recovered and only a few indeterminate microplankton fragments.

### Core 9, Section 1, 147-149 cm:

A sporomorph flora containing: Appendicisporites sp., Classopollis echinatus, cf. Patellasporites distaverrucosus, Tsugaepollenites mesozoicus, Ephedripites spp., Zonalapollenites dampieri, Parvisaccites radiatus, Araucariacites australis, Eucommidites troedsonii, Spheripollenites scabratus, Cicatricosisporites dorogensis, Leptolepidites cf. psarosus, and Cicatricosisporites cf. cooksonae. Age: Hauterivian–Aptian.

Microplankton scarce, with Canningia hirtella, Hystrichosphaeridium asterigerum, and Microdinium sp.

Age: Early Cretaceous, probably Valanginium-Barremian to possibly Aptian.

Core 9, Section 2, 26-28 cm, and 89-91 cm:

A rich sporomorph flora with abundant Classopollis classoides, Ephedripites sp., common Cicatricosisporites sp., Circulina sp., Ephedripites spiralis, Eucommidites troedsonii and rare Zonalapollenites dampieri, Clavatipollenites hughesii, Sofrepites cf. legouxae, Araucariacites australis, Caytonipollenites pallidus, Trilites sp., (Jardiné & Magloire 1963, type S.CI.128), Verrutriletes sp., Appendicisporites tricornitatus, Klukisporites cf. variegatus, Spheripollenites scabratus, Syncolporites sp. and cf. Monocolpopollenites spheroidites. Age: Barremian-Aptian.

Some microplankton with Hystrichosphaeridium asterigerum, Pseudoceratium pelliferum, Canningia hirtella, Pareodinia ceratophora, Microdinium deflandrei and Dingodinium cerviculum.

Age: Early Cretaceous, probably Valanginium-Barremian to possibly Aptian.

Core 9, Section 2, 146-148 cm:

A poor sporomorph flora containing Classopollis echinatus, Cicatricosisporites cf. cooksonae, C. dorogensis, C. cf. mohrioides, Eucommidites troedsonii, Araucariacites

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Figure 4. Drilling and coring summary at Site 135.

australis, Spheripollenites scabratus, Ephedripites spp., cf. Patellasporites distaverrucosus, Appendicisporites tricornitatus and Tsugaepollenites mesozoicus. Age: Hauterivian-Aptian.

Microplankton scarce, with Microdinium deflandrei, and Hystrichosphaeridium asterigerum.

Age: Early Cretaceous, probably Valanginium-Barremian to possibly Aptian.

## LITHOSTRATIGRAPHY

At Site 135 one hole was drilled to a maximum depth of 689 meters below the sea floor; a second hole (135A) was drilled to about 59 meters, but no core was recovered.

Coring at intervals of approximately 80 meters enabled the following four lithostratigraphic units to be distinguished:

| Unit | Cores             | Lithology  | Depth Below<br>Sea Floor<br>(m) | Age                                 |  |
|------|-------------------|--|---------------------------------|-------------------------------------|--|
| 1    | 1,2,3,4<br>SW 1&2 | Light-gray nanno-<br>fossil chalk ooze                                 | 0-325                           | Pleistocene<br>to Late<br>Oligocene |  |
|      |                   | HIATUS   | 5                               |                                     |  |
| 2    | 5,6,7             | Olive gray to brown<br>silty mud(stone),<br>sand layers, brown<br>clay | 335-450                         | Early<br>Eocene to<br>Maestrichtian |  |
| 3    | 8                 | Black and green<br>shales with lime-<br>stone and chert<br>layers      | 450-560 <sup>a</sup>            | ? Cretaceous                        |  |
| 4    | 9                 | Nannofossil marl<br>and limestone                                      | 560-689a                        | Early Aptian                        |  |

<sup>a</sup>Unit thicknesses not well determined.

UNIT 1 – Light Gray Nannoplankton Chalk Ooze (Cores 1, 2, 3, 4), SW 1 and 2

This unit is about 325 meters thick and consists mainly of pelagic carbonates deposited since the Late Oligocene. The chalk ooze is largely calcareous nannoplankton, usually with less than 5 per cent planktonic foraminifera. The foraminifera show excellent preservation near the top of the unit but increasing signs of dissolution downward. In the marl near the base, preservation is very poor. Silt-size quartz occurs in traces; the amount is somewhat greater near both the top and base of the unit than elsewhere. Drilling rates usually varied between 1 and 2.5 m/min. Fast drilling near 250 meters suggests that sandy layers may be present. Below 270 meters, rates were reduced to about 0.5 m/min suggesting the presence of firm marls.

### UNIT 2 – Olive Gray to Brown Mud (stone), Sand Layers and Brown Clay, Generally Quartz-rich (Cores 5, 6, 7)

This unit is between 115-235 meters thick (change in drilling rate suggests 120 m) and spans the time interval from Maestrichtian to Early Eocene.

Core recovery was poor except in Core 7. Terrigenous sediments low in carbonate (<25%), but carbonatecemented in places, exhibit various size-grades, colors, and degrees of induration. Silty and clayey mud (stones), usually quartzose and with detrital calcareous grains, are the chief components of the unit. Sand layers (3-25 cm thick) visibly graded in part, contain quartz grains up to 0.3 mm in diameter. The sand layers commonly are alternating with pelagic brown clay. Arkosic sands with highly altered feldspar grains are present. Iron oxide coated quartz grains are also present.

Coarse sand (common size 1-2 mm) consisting mainly of quartz and cert but having small amounts of mica, pyrite, limestone and other lithic fragments (but little feldspar) was recovered from the center bit following recovery of Core 9. Benthonic foraminifera are of Cretaceous age. This sand apparently belongs to either unit 2 or unit 3 (most probably unit 2) since the greater drilling rate suggests that much sand was penetrated).

The brown clay of this unit consists of 50-75 per cent clay minerals and iron oxides, 10-35 per cent quartz, and 10-20 per cent carbonate (some of it dolomite rhombs; the rest detrital). Alternating fast (1.5 m/min) and slow (0.8 m/min) drilling may indicate sand-rich and clay-rich sections in the unit.

# UNIT 3 – Black and Green Shale with Limestone and Chert Layers (Core 8)

The thickness of this unit is not known. Only 1 meter of core was recovered near the position where drilling slowed. The age is probably Late Cretaceous based on dates in Cores 7 and 9. The black shale and green siliceous mudstone contain only traces of carbonate; they are rich in zeolite and contain pyrite and 5-10 per cent terrigenous silt. A thin certy bed of silicified laminated silty mudstone contains silica (75%), quartz (29%), clay minerals, and opaque and organic materials. Recrystallized sponge spicules and radiolarians (?) were noted. A thin bed of non-fossiliferous recrystallized limestone consists of large (1-2 mm diameter) fibrous calcite spherulites separated by clay. Drilling rates vary between 0.4 and 1 m/min.

# UNIT 4 – Olive Gray and Black Nannofossil Marl and Limestone (Core 9)

The thickness of this unit is unknown; 1.8 meters were recovered. The age is Early Cretaceous. The main facies is dark (sapropelitic ?) shale, laminated in part and rich in calcareous nannofossils; some pyrite and authigenic carbonate is also present. In places, thin silty layers containing much quartz are intercalated with the shale. A slightly quartzose limestone bed, partly micritic, partly sparry, and with 5 per cent nannofossils, was noted. The core catcher contained sand-size material consisting of much quartz and detrital carbonate.

#### PHYSICAL AND CHEMICAL PROPERTIES

A few sediments recovered from Site 135 were disturbed during coring so physical property measurements may be unreliable.

Penetrometer measurements (units of  $mm \times 10^{-1}$ ) on the chalk oozes showed a decrease from a mean value of 60 near the surface to a value of 30 at about 175 meters; below this values were relatively constant. The terrigenous sediments (Early Tertiary and Cretaceous) in the lower part of the hole show a wide range of values (5 to 85) due to variations in lithology, degree of induration, and disturbance.

The bulk density (gm/cc) increases from 1.57 near the surface to 1.87 near the bottom of the ooze at 325 meters. Below this level, the density values have a wide scatter with individual readings correlating reasonably well with the different lithologies. The normal range of bulk densities is 1.5 to 1.8; in the less consolidated sands 1.6 to 1.8, and in the indurated beds of silicified mudstone and sandy limestone 1.8 to 2.1. Porosity and water content showed an inverse correlation with bulk density, decreasing downward through the ooze, with scattered values for the terrigenous sediments.

The porosity values determined from the GRAPE are, in general, higher than those calculated from measurement of the water content; the density values obtained by the two methods are in good agreement (Table 3).

Variations in the natural gamma radiation of the samples correlates closely with changes in lithology. Counts from the ooze are fairly uniform, averaging about 550. However, as is common in deep-sea sediments, counts in the top two meters of this ooze are relatively high (up to 1550 counts). Below the ooze, the terrigenous deposits have higher counts. The clays and silts range from about 1000 to 2500; the average being about 1500 counts. There was no significant difference between the black clays and the other clays. The sands give counts of 500 to 1000 except for the arkosic beds in the lower part of Core 7 which have counts of about 1500. The highest counts (up to 3500) were from the zeolitic clay at the bottom of Core 8.

Sonic velocity measurements were made on four samples of lithified sediment as follows: 135-8-1, 60 cm 3.28 km/sec, 110 cm 2.11 km/sec, 135 cm 3.42 km/sec; and 135-9-1, 105 cm 3.47 km/sec.

|      |      |         |                                 | GRAPE              |                       |                                 | Sediment 3              | Sample             |                 |
|------|------|---------|---------------------------------|--------------------|-----------------------|---------------------------------|-------------------------|--------------------|-----------------|
| Hole | Core | Section | Depth Below<br>Sea Floor<br>(m) | Density<br>(gm/cc) | Porosity<br>(%)       | Depth Below<br>Sea Floor<br>(m) | Water<br>Content<br>(%) | Density<br>(gm/cc) | Porosity<br>(%) |
| 135  | 1    | 1       | 0.75                            | 1.57               | 65                    | 0.30                            | 40                      | 1.56               | 62              |
| 135  | 1    | 2       | 2.25                            | 1.61               | 62                    | 1.51                            | 37                      | 1.65               | 60              |
| 135  | 1    | -       | -                               | -                  | <u> </u>              | 2.60                            | 33                      | 1.70               | 56              |
| 135  | 1    | 3       | 3.75                            | 1.61               | 62                    | 3.14                            | 34                      | 1.67               | 57              |
| 135  | 2    | 1       | 80.75                           | 1.76               | 50                    | 80.59                           | 31                      | 1.51               | 47              |
| 135  | 2    | 2       | 82.25                           | 1.79               | 48                    | 81.65                           | 32                      | 1.75               | 55              |
| 135  | 2    | 3       | 83.75                           | 1.74               | 51                    | 83.14                           | 30                      | 1.75               | 53              |
| 135  | 2    | 4       | 85.25                           | 1.80               | 47                    | 84.64                           | 32                      | 1.70               | 55              |
| 135  | 2    | 5       | 86.75                           | 1.73               | 52                    | 86.15                           | 31                      | 1.76               | 55              |
| 135  | 2    | 6       | 88.25                           | 1.80               | 47                    | 87.64                           | 32                      | 1.70               | 55              |
| 135  | 3    | 1       | 173.75                          | 1.85               | 36                    | 173.54                          | 26                      | 1.72               | 45              |
| 135  | 3    | 1       | -                               | 1.77               | -                     | 173.59                          | 20                      | 1.71               | 33              |
| 135  | 3    | 2       | 175.25                          | 1.87               | 34                    | 174.64                          | 29                      | 1.76               | 51              |
| 135  | 4    | 1       | 259.75                          | 1.79               | 37                    | -                               | -                       |                    | <u> 1917</u>    |
| 135  | 4    | 2       |                                 | 3000               | 777                   | 260.64                          | 26                      | 1.68               | 44              |
| 135  | 5    | 1       | 335.75                          | 1.46 <sup>a</sup>  | 50                    | 335.73                          | 54                      | 1.28               | 69              |
| 135  | 7    | 1       | -                               | -                  | -                     | 432.43                          | 23                      | 1.66               | 38              |
| 135  | 7    | 2       | 433.25                          | 1.89               | 29                    | 432.64                          | 21                      | 1.84               | 39              |
| 135  | 7    | 3       | 434.75                          | 1.88               | 30                    | 434.31                          | 21                      | 1.79               | 38              |
| 135  | 7    | 4       | 436.25                          | 1.86               | 31                    | 435.60                          | 22                      | 1.78               | 39              |
| 135  | 8    | 1       | 564.75                          | Variable 1         | ithology <sup>b</sup> |                                 |                         | -                  | -               |
| 135  | 9    | 1       | 685.75                          | 1.88               | 30                    | -                               | _                       |                    | 1.00            |
| 135  | 9    | 2       | 687.25                          | 1.65               | 49                    |                                 |                         |                    | -               |

 TABLE 3

 Summary of Density, Porosity and Water Content Data for Site 135

<sup>a</sup>Questionable value as Core disturbed.

<sup>b</sup>Density varied from <1.5 to >2.0 gm/cc

Properties of the interstitial water are within the normal range for seawater with salinities ranging from 32.4 to 34.1 ppt. and pH values ranging from 7.24 to 7.80 (Table 4).

#### DISCUSSION AND CONCLUSIONS

Site 135 lies on a topographic high (SE of the Horseshoe Abyssal Plain) covered by about 325 meters of calcareous ooze. A thick section of predominantly quartz-rich terrigenous and pelagic sediments low in carbonates lies below the ooze. Calcareous sediments again characterize the lowermost part of the cored sequence below about 560 meters.

An important result obtained from this site is that the occurrence of the major contrast in lithologic type correlates very well in depth with that anticipated for a major reflecting horizon. The reflecting horizon at Site 135 lies at a depth of 0.38 second, and a compressional wave

| TABLE 4                                 |                         |
|---|-------------------------|
| <b>Chemical Property Measurements o</b> | n Samples from Site 135 |

|      |      |         | Sampl | e Interval<br>cm) |      |      | Salinity |
|------|------|---------|-------|-------------------|------|------|----------|
| Hole | Core | Section | Тор   | Bottom            | pH   | Eh   | (°/00)   |
| 135  | 1    | 2       | 141.0 | 150.0             | 7.44 | +69  | 34.1     |
|      | 2    | 6       | 0.0   | 8.0               | 7.24 | +116 | 33.6     |
|      | 3    | 2       | 0.0   | 7.0               | 7.36 | +119 | 33.0     |
|      | 4    | 2       | 0.0   | 5.0               | 7.26 | +123 | 33.6     |
|      | 7    | 3       | 140.0 | 150.0             | 7.80 | +130 | 32.4     |

velocity of about 1.77 km/sec for the pelagic sequence can be inferred from the above correlation. There are no other reflecting horizons observed on the seismic record that can plausibly correlate with the lithologic change. Similarly, there are no other observed major differences in lithologic type within the sampled section that could logically create the reflecting horizon.

This prominent reflecting horizon extends over many tens of miles as observed on the Jean Charcot and Challenger seismic records and can be traced from the site northward under the well-layered upper section of the Horseshoe Abyssal Plain. Southeast of Site 135, the reflector is interrupted by deep-seated piercement structures that may be salt diapirs (Pautot et al., 1970). The reflecting horizon is offset along apparent high-angle faults with displacements on the order of 300-500 meters. Although it was evident prior to drilling that tectonism followed the formation of the reflecting horizon, the age of this activity was unknown. The offset reflecting horizon is semi-conformable with the highs and lows of the basement reflector. The transparent (pelagic) section overlying the reflecting horizon is conformable with it and with the sea floor. The inference is that most, if not all, of the movement along these faults took place before Middle Miocene and prior to the deposition of the recovered pelagic carbonates.

The following sequence of events provides the best explanations of the geophysical data and drilling results at this site.

From Late Cretaceous to Early Eocene time, the area was a deep (though not so deep that all calcareous fossils were dissolved) flat-lying oceanic environment. The area was receiving olive and brown clays containing detrital carbonates and intercalated with coarse terrigenous sediments. The area was, in essence, an abyssal plain. The apparent sedimentation rate for this sequence is about 5 m/my. If some allowance is made for rapidly emplaced sands and silts, the true sedimentation for the clays is probably 2-3 m/my. Sediment reaching the area was varied: coarse sands (mostly quartz and chert and shallow-water foraminifera) were transported along the sea floor; and medium to fine sands were probably deposited by turbidity currents. Primary structures, such as crossbedding, indicate the presence of near bottom currents. The sands contain minerals that were derived from both granitic and metamorphic terrains; highly rounded quartz grains with iron oxide coatings are probably of eolian origin (?Saharasand); and detrital carbonate sands (later crystallized to coarse-grained limestones) occasionally reached the area.

During some intervals, this environment received little sand, and shales with a high content of organic material were deposited. Some calcareous pelagic components also reached the sea floor at that time, but their net contribution to the sequence was minor. The fertility of the upper waters possibly increased during the Late Cretaceous thus leading to the deposition of siliceous organisms (probably preserved as chert); calcareous organisms were, however, dissolved. An increasing influx of coarse-grained terrigenous sediments (turbidites) in the latest part of the Cretaceous may reflect uplift and erosion of the adjacent continent.

Following the Early Eocene, epeirogenic uplift accompanied by faulting created the topographic highs south of the Horseshoe Abyssal Plain, and the area of Site 135 was made inaccessible to terrigenous sediments. From at least as early as Late Miocene, the area was a site of rapid pelagic carbonate sedimentation at first close to the carbonate compensation depth ( $\sim$ 4500 meters) and subsequently well above it. Evidence for this is seen in the poor preservation of calcareous fossils in the Middle Miocene in contrast to the excellent preservation in the Pleistocene.

The unconformity between Early Eocene terrigenous sediments and Late Oligocene pelagic sediments is not yet clearly understood. It is possible that following uplift, the area remained for some time below the carbonate compensation depth so that during this period the site received neither terrigenous nor pelagic sediments. However, a similar stratigraphic hiatus is present at many other sites and an explanation involving more regional processes, such as deep sea erosion and slumping accompanying tectonism or a major change in the pattern of oceanic circulation, is probably more reasonable. Tectonic activity along the east Azores Fraction Zone beginning at least 45 million years ago has been inferred by Krause and Watkins (1970). The regional tectonism and deformation exhibited by the reflecting horizon may relate to these changes in plate tectonic activiity. It should be noted that Site 135 lies only 70 km south of the inferred Azores Fracture Zone, and therefore, could be subject to intense deformation associated along this boundary by interaction between two major crustal plates. The Early Tertiary also marked the initiation of the Alpine orogenic movements in the sub-Betic Chain, 400 km to the northeast of this site.

The unusually high content of nannofossils in the chalk ooze of Unit 1 is unexplained. Differential dissolution cannot be the main cause because the foraminifera in Core 1, although relatively rare, are quite well preserved.

Sediments of Early Aptian age were the oldest materials recovered at Site 135 at a subbottom depth of 689 meters. Although the seismic data from the *Challenger* are not definitive, records from the *Jean Charcot* indicate basement rocks (layer 2) lie an additional 350 meters below the Aptian sediments. These Aptian sediments are well lithified limestones and marl, and penetration by the drill slowed at this point. On both the *Challenger* (at 0.71 sec) and *Charcot* (0.75 sec) seismic reflection records, a second intermediate reflector could be correlated with this change in lithology.

Sedimentation rates below the unconformity are about 5-8 m/my. Extrapolating these rates from the depth of the Early Aptian sediments (~100 m.y.) to the inferred basement, the deduced age of the sediments directly above basement would be about 155-180 million years or Late to Early Jurassic. Site 135 lies about 200 km landward from the eastern Atlantic magnetic quiet zone boundary which has been interpreted as an isochron (Heirtzler and Hayes, 1967). At Site 105 of Leg 11 which lies close to the western Atlantic magnetic quiet zone boundary, sediments of Oxfordian–Callovian (early Late Jurassic; ~155 my) were recovered overlying an extrusive basalt (Hollister and Ewing *et al.*, 1972).

There is no evidence of salt or associated evaporite sediments in any of the 689 meters of section drilled and sampled. This is not inconsistent with the interpretation of the presence of salt diapirs nearby (Pautot, 1970) since seismic data indicates the source of the presumed salt to be much more deeply seated than the depth of the recovered cores (see Figure 2).

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# SITE 135-SUMMARY



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SITE 135 CORE 1

.

## DEPTH (m) 0-4

|                  | Z                             | ONE                   |     |                   |        | DE                    |  | NATURAL  |
|------------------|-------------------------------|-----------------------|-----|-------------------|--------|-----------------------|--|--|
| AGE              | FORAM                         | NANNO                 | RAD | SECTION           | METERS | LITHOLOGIC<br>SYMBOLS | LITHOLOGIC DESCRIPTION   | GAMMA<br>RADIATION<br>COUNTS/7.6 cm/1.5 min<br>1000 2000 |
| LATE PLEISTOCENE | Globorotalia truncatulinoides | Gephyrocapsa oceanica |     | 1<br>2<br>3<br>CC | 2      |                       | GZ 4-21-75 CaCO <sub>3</sub> 38<br>NANNO CHALK OOZE<br>Light gray (10YR 7/2) and<br>grayish brown (10YR 5/2)<br>Forams, mainly fragments 5-10%<br>Quartz silt, at top v10%<br>at base down to 1%<br>Clay 5-10% CaCO <sub>3</sub> 50<br>GZ 7-20-73<br>Some fine (~1 mm) lamination, especially<br>in Section 2<br>Coarse Fraction: CaCO <sub>3</sub> 62<br>Pelagic forams, moderately to well<br>preserved<br>sec.2, 125 cm GZ 5-13-82 CaCO <sub>3</sub> 72<br>GZ 2-13-85<br>GZ 1-18-81 |  |



**SITE 135** 

|        | Z      | ONE    |     |      |      |                       | Ш     |  | NATURAL                            |
|--------|--------|--------|-----|------|------|-----------------------|-------|--|------------------------------------|
|        | W      | NO     |     | NOI  | RS   |                       | R SLI | LITHOLOGIC DESCRIPTION                                       | GAMMA<br>RADIATION                 |
| AGE    | FORA   | NAN    | RAD | SECT | METE | LITHOLOGIC<br>SYMBOLS | SMEA  |  | COUNTS/7.6 cm/1.5 min<br>1000 2000 |
| Π      |        |        |     |      | 111  |                       |       | NANNO CHALK OOZE CaCO <sub>3</sub> 78                        | l                                  |
|        |        |        |     | 1    |      |                       |       | $CaCO_3$ 75<br>White (10YR 8/1) to light gray (5Y 7/1 to     | 3                                  |
|        |        |        |     |      | 1-   |                       | .00   | 5Y 6/1); dark specks and streaks locally<br>60 cm GZ 3-22-75 | j                                  |
|        |        |        |     |      | 111  |                       |       | Smear Slide Average:   | 1                                  |
|        |        |        |     |      | 1.62 |                       |       | Forams (mostly fragments) 2-10%<br>Clay Tr.2%                | 5                                  |
|        |        |        |     | 2    | 2    |                       |       | Quartz Tr.<br>30 cm GZ 2-22-76                               | 2                                  |
|        |        |        |     | 2    | 111  |                       | 90    | Coarse Fraction:   | 5                                  |
|        | tae    | 87     |     |      |      |                       |       | relagic forans, moderately well preserved                    | 3                                  |
| CENE   | argar  | 180BnJ |     |      | 11   |                       |       | GZ 1-20-79 CaCO3 70  | }                                  |
| PLIO   | ilia m | ithus  |     | 2    |      |                       | 80    |  | 5                                  |
| EARLY  | borote | ratol. |     | 5    | 4    |                       |       |  | 5                                  |
|        | 015    | Ce     |     |      |      |                       |       |  | 5                                  |
|        |        |        |     |      | -    |                       |       | GZ 5-24-71 CaCO <sub>3</sub> 85                              | Ş                                  |
|        |        |        |     | 4    | 5-   |                       |       |  | ξ                                  |
|        |        |        |     |      |      |                       | 00    |  | 5                                  |
|        |        |        |     |      | 6    |                       |       |  | I                                  |
|        |        |        |     |      |      |                       |       | CaCO <sub>3</sub> 71<br>GZ 1-21-78                           | 3                                  |
|        | -      | atus   |     | 5    |      |                       | 75    |  | ł                                  |
|        |        | micul  |     |      | 7-   |                       |       |  | 4                                  |
|        |        | tricor |     |      | -    |                       |       | CaCO 80  | ۲ <sup>7</sup>                     |
| OCENE  | ~      | thus 1 |     |      | -    |                       |       | GZ 4-22-74   | }                                  |
| ATE M. |        | atoli  |     | 6    |      |                       |       |  | ł                                  |
|        |        | Ce     |     |      | 1111 |                       | 00    |  | ſ                                  |
|        |        |        |     | сс   | _    |                       |       |  | ť                                  |



|              | Z                          | ONE           |     |              |        | L L L L L L L L L L L L L L L L L L L |  |   |                                    |                                  |  |
|--------------|----------------------------|---------------|-----|--------------|--------|---------------------------------------|--|---|------------------------------------|----------------------------------|--|
| AGE          | FORAM                      | NANNO         | RAD | SECTION      | METERS |                                       | LITHOLOGIC DESCRIPTION   |   | GAM<br>RADIA<br>counts/7.6<br>1000 | MA<br>TION<br>cm/1.5 min<br>2000 |  |
| LATE MIOCENE | Globorotalia acostaensis ? | D. neohamatus |     | 1<br>2<br>CC | 2<br>3 |                                       | NANNO CHALK OOZE<br>Light gray (5Y 6/1 to 5Y 6/2,<br>5Y 7/1 to 5Y 7/2, 10YR 8/1)<br>Smear Slide Average:<br>Nannos 95%<br>Forams (mostly fragments) 1-2%<br>Clay, quartz, pyrite, Tr.<br>chlorite, Fe oxide<br>Alternating bands (15-30 cm) of lighter and<br>darker bands, composition the same.<br>Coarse Fraction:<br>Pelagic forams, poorly preserved; also,<br>benthonic forams, calcareous and arenaceous,<br>fish debris, ostracods, spicules<br>sec.1, 60 cm GZ 0-28-72<br>sec.2, 30 cm GZ 1-28-71 | CaCO <sub>3</sub> 71<br>CaCO <sub>3</sub> 7<br>CaCO <sub>3</sub> 85 |                                    |                                  |  |

|                |                                   |                            |     |              |        |                       |  | SITE 135   | CORE 4               | DEPTH (m) 259-268  |
|----------------|-----------------------------------|----------------------------|-----|--------------|--------|-----------------------|--|--|----------------------|--|
|                | Z                                 | ONE                        |     |              |        |                       | DE                                     |  |                      | NATURAL  |
| AGE            | FORAM                             | NANNO                      | RAD | SECTION      | METERS | LITHOLOGIC<br>SYMBOLS | SMEAR SLI                              | LITHOLOGIC DESCRIPTION   |                      | GAMMA<br>RADIATION<br>COUNTS/7.6 cm/1.5 min<br>1000 2000 |
| MIDDLE MIDCENE | Globorotalia foksi peripheroronda | Sphenolithus heteromorphus |     | 1<br>2<br>CC | 2      |                       | 110<br>121<br>139<br>_30<br>_60<br>100 | NANNO CHALK OOZE<br>Light gray (10YR 8/1) to light brownish<br>gray (10YR 6/2), colors alternate,<br>mottling intense<br>Smear Slide Average:<br>Nannos 90%<br>Carbonate flour 7%<br>Clay 3%<br>GZ 1-30-69<br>Coarse Fraction:<br>Pelagic forams, very poorly preserved;<br>benthonic forams, spicules, fish<br>debris, quartz | CaCO <sub>3</sub> 93 |  |

#### SITE 135 CORE 3

DEPTH (m) 173-182



| SITE | 135 | CORE | 5 | DEPTH (m) | 335-341 |
|------|-----|------|---|-----------|---------|
|      |     |      |   |           |         |

|              | Z     | ONE                              |     |         |        |                       | DE                     |  | NATURAL  |
|--------------|-------|----------------------------------|-----|---------|--------|-----------------------|------------------------|--|--|
| AGE          | FORAM | NANNO                            | RAD | SECTION | METERS | LITHOLOGIC<br>SYMBOLS | SMEAR SLI              | LITHOLOGIC DESCRIPTION   | GAMMA<br>RADIATION<br>COUNTS/7.6 cm/1.5 min<br>1000 2000 |
| EARLY EOCENE |       | Marthas terites<br>tribrachiatus |     | 1<br>CC |        | VOID                  | 73<br>94<br>102<br>142 | GZ 2-57-41<br>SILTY MUD<br>Banded colors as follows: CaCO <sub>3</sub> 9<br>69-78 cm, grayish yellow green (5GY 7/2);<br>78-91 cm, olive gray (5Y 5/2);<br>91-95 cm, reddish brown (5YR 5/3);<br>95-101 cm, olive gray (5Y 5/2);<br>101-108 cm, reddish brown (5YR 5/2);<br>then grayish yellow green (5GY 7/2)<br>grading downward to dark gray (5Y 4/1)<br>at base<br>Smear Slide Average:<br>Clay 75%<br>Quartz 10-15%<br>Carbonate rhombs 7-15%<br>Biotite, muscovite 38%<br>and chlorite<br>Fe oxides and pyrite 35%<br>Smear Slide, 102 cm:<br>(Quartz 25%<br>Clay 20%<br>Fe oxide and pyrite 15%<br>Biotite 15%<br>Muscovite 5%<br>Chlorite 5%<br>Chlorite 5%<br>Chlorite 3%<br>Carbonate flour 5%<br>Feldspar 1%<br>Heavies 3%<br>(Titanite, clinozoisite,<br>amphibole, tourmaline)<br>Coarse Fraction:<br>Flakes of indurated clay,<br>occasional chert fragment |  |



|                         |       |       |     |         |        |                       |           | SITE 135   | CORE 6 | DEPTH (m) 341-350   |
|-------------------------|-------|-------|-----|---------|--------|-----------------------|-----------|--|--------|---|
|                         | Z     | ZONE  |     |         |        |                       | DE        |  |        | NATURAL   |
| AGE                     | FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGIC<br>SYMBOLS | SMEAR SLI | LITHOLOGIC DESCRIPTION   |        | GAMMA<br>RADIATION<br>COUNTS/7.6 cm/1.5 min<br>1000 2000<br>I I |
| CRETACEOUS OR PALEOCENE |       |       |     |         |        |                       |           | QUARTZOSE SILTSTONE AND MUDSTONE<br>Black (5Y 2/2), semi-consolidated<br>Smear Slide Average:<br>Quartz 45%<br>Clay 20%<br>Biotite and chlorite 5%<br>Fe oxide and pyrite 5%<br>Feldspar 2%<br>Carbonate rhombs 15%<br>Carbonate flour 5%<br>Heavies 1%<br>(Hypersthene, clinozoisite)<br>Coarse Fraction:<br>Flakes of indurated clay,<br>polished quartz, (?) regrown<br>quartz, mica, lithic fragments<br>(schist, sandstone), glauconite |        |   |



|                                    |       |       |     |               |        |                       |             | SITE 135 CORE 7  | DEPTH (m) 431-435   |
|------------------------------------|-------|-------|-----|---------------|--------|-----------------------|-------------|--|---|
| AGE                                | FORAM | NANNO | RAD | SECTION       | METERS | LITHOLOGIC<br>SYMBOLS | SMEAR SLIDE | LITHOLOGIC DESCRIPTION   | NATURAL<br>GAMMA<br>RADIATION<br>COUNTS/7.6 cm/1.5 min<br>1000 2000 |
| LATE CAMPANIAN-EARLY MAESTRICHTIAN |       |       |     | 1<br>2*<br>3* | 2      | VOID                  |             | GZ 4-74-22SILTY MUDSTONE, SAND AND CLAYDark gray (10YR 4/1, 5Y 4/1), dark grayish<br>brown (10YR 4/2), greenish gray (5GY 6/1),<br>dark brown (10YR 3/3)Smear Slides show sands and silty muds are<br>mostly quartz, abundant feldspar, chlorite,<br>biotite, and authigenic carbonate. Some CaCO3 28<br>sand layers appear graded. Induration and<br>cementation variable throughout core. CaCO3 24Coarse Fraction:<br>Rounded to subangular quartz, lithic<br>fragments(gneiss, schist, limestone, sand-<br>stone), mica. Few pelagic forams, spicules,<br>ostracods, benthonic forams, glauconite* See Section Summary<br>GZ 76-15-9GZ 78-13-9<br>CaCO3 23 CaCO3 37GZ 1-82-17CaCO3 21 |   |



DEPTH (m) 564-569

|            | 2     | ZONE  | E   |          |                   |        | DE     |  | NATURAL |        |                       |            |                        |  |
|------------|-------|-------|-----|----------|-------------------|--------|--------|--|---------|--------|-----------------------|------------|------------------------|--|
| AGE        | FORAM | NANNO | RAD | SECTION  | SECTION<br>METERS | METERS | METERS | METERS   | METERS  | METERS | LITHOLOGIC<br>SYMBOLS | SMEAR SLID | LITHOLOGIC DESCRIPTION | GAMMA<br>RADIATION<br>COUNTS/7.6 cm/1.5 min<br>1000 2000 |
| CRETACEOUS |       |       |     | 1*<br>CC | 1                 | VOID   |        | BLACK MUD AND GREENISH SILICEOUS MUDSTONE, WITH<br>LIMESTONE AND CHERT<br>Coarse Fraction:<br>Chips of light colored siltstone, pebbles of<br>dark laminated mudstone, traces of pyrite<br>and fish debris, fragments of coarsely<br>crystalline limestone * See Section Summary |         |        |                       |            |                        |  |

SITE 135 CORE 9

**SITE** 135

CORE 8

DEPTH (m) 685-687

| Γ            | ZONE                     |       |     |                |        | DE                    |           | NATURAL  |  |
|--------------|--------------------------|-------|-----|----------------|--------|-----------------------|-----------|--|--|
| AGE          | FORAM                    | NANNO | RAD | SECTION        | METERS | LITHOLOGIC<br>SYMBOLS | SMEAR SLI | LITHOLOGIC DESCRIPTION   | GAMMA<br>RADIATION<br>COUNTS/7.6 cm/1.5 min<br>1000 2000 |
| EARLY APTIAN | Chiastozygus litterarius |       |     | 1*<br>2*<br>CC | 2      | VOID                  |           | LIMESTONE<br>Greenish gray (5G 6/1)<br>GZ 1-34-65<br>QUARTZOSE NANNOFOSSIL MARL 00ZE<br>Olive gray (5Y 4/1) to black (5Y 2/2)<br>Quartz and other terrigenous components<br>Coarse Fraction:<br>Pyritized fecal pellets and burrow<br>fillings, mica, claystone fragments,<br>quartzose siltstone<br>* See Section Summary |  |



SITE 135 CORE SW 1 DEPTH (m) SW 1-315

|               | ZONE  |                                 |     |         |        | DE                    |           | NAT   | URAL                              |   |
|---------------|-------|---------------------------------|-----|---------|--------|-----------------------|-----------|---|-----------------------------------|---|
| AGE           | FORAM | NANNO                           | RAD | SECTION | METERS | LITHOLOGIC<br>SYMBOLS | SMEAR SLI | LITHOLOGIC DESCRIPTION  | GAN<br>RADI/<br>counts/7.<br>1000 | 1MA<br>ATION<br>6 cm/1.5 min<br>2000<br>I |
| EARLY MIOCENE |       | Triquetrorhabdulus<br>carinatus |     | 1       | 1      | NOT TO<br>SCALE       | _         | <pre>SW 1 NANNO CHALK OOZE Coarse Fraction:    Mixture of Mn/Fe crust, poorly preserved    benthonic forams, planktonic forams,    fish debris, angular quartz grains and    rounded quartz pebbles, pyritized    burrow fillings</pre> |                                   |   |

## SITE 135 CORE SW 2 DEPTH (m) SW 2-325

|                | ZONE                           |                                 | ONE |         |              |                       | DE        |  | NAT                               | JRAL                                      |
|----------------|--------------------------------|---------------------------------|-----|---------|--------------|-----------------------|-----------|--|-----------------------------------|---|
| AGE            | FORAM                          | NANNO                           | RAD | SECTION | METERS       | LITHOLOGIC<br>SYMBOLS | SMEAR SLI | LITHOLOGIC DESCRIPTION   | GAN<br>RADIA<br>COUNTS/7.<br>1000 | IMA<br>ATION<br>6 cm/1.5 min<br>2000<br>I |
| LATE OLIGOCENE | Globigerina<br>ciperoensis (?) | Triquetrorĥabdulus<br>carinatus |     | 1       | in fin fin i |                       |           | <pre>SW 2 NANNO CHALK OOZE Coarse Fraction:    Coarse, rounded quartz sand; also lithic    fragments (sandstone, limestone, pyrite,    chert, igneous)</pre> |                                   |   |

SITE 135 CORE 7 SECTION 2

| AGE             | SECTION<br>PHOTO | cm                            | LITHO | SMEAR       | DESCRIPTION  |
|-----------------|------------------|-------------------------------|-------|-------------|--|
| Y MAESTRICHTIAN |                  | 25 —<br><br><br><br>50 —      |       |             | Feldspathic, quartzose terrigenous SANDY SILT.         Dark gray (10YR 4/1) to dark grayish brown         (10YR 4/2). Soft at top to well indurated         at 45 cm.         Smear Slide (23 cm):         Quartz       60%         Feldspar       15%         Chlorite and biotite       10%         Carbonate rhombs       10%         Clay       5%         Heavies       Tr.         Sand and sandstone       Greenish gray (5GY 6/1). Cement is carbonate.         Well lithified 45-60 cm       Smear Slide Average: |
| AMPANIAN - EARL |                  | <br><br>75                    |       | <u>6</u> 7  | Quartz60-70%Clay20%Feldspar5-10%Carbonate cement5-20%Mica, heaviesTr.  |
| EARLY CAN       | -                | -<br>-<br>-<br>100-<br>-<br>- |       | <u>10</u> 0 | Colors dark grayish brown (10YR 4/2) and dark gray (5Y 4/1)<br>X-RAY (75.5 cm; mud layer under sand):<br>Quartz A<br>Feldspar, Kaolin, Calcite, Dolomite(?) C<br>Montmorillonite, Chlorite Tr.<br>Carbonate cemented sandstone, up to 70% cement.  |
|                 |                  |                               |       | 135         | Graded (?) sand layer.<br>Clay ball (1 cm. dia.) at 115 cm.<br><u>X-RAY (110-111 cm; top of sand layer):</u><br>Quartz A<br>Feldspar, Mica, Montmorillonite,<br>Mixed Layer (Palygorskite), C<br>Chlorite, Kaolin, Calcite, Dolomite.  |

## SITE 135 CORE 7 SECTION 3



# SITE 135 CORE 7 SECTION 4



# SITE 135 CORE 8 SECTION 1



SITE 135 CORE 9

SECTION 1



SITE 135 CORE 9 SECTION 2

