

8. SITE 153

The Shipboard Scientific Party¹

Position:

13°58.33'N
72°26.08'W.

Water Depth: 3932 meters.

Penetration: 776 meters.

ABSTRACT

This site is located in the Aruba Gap, between the Beata Ridge and Venezuela where processed reflection data indicate stratified layers below Horizon B''.

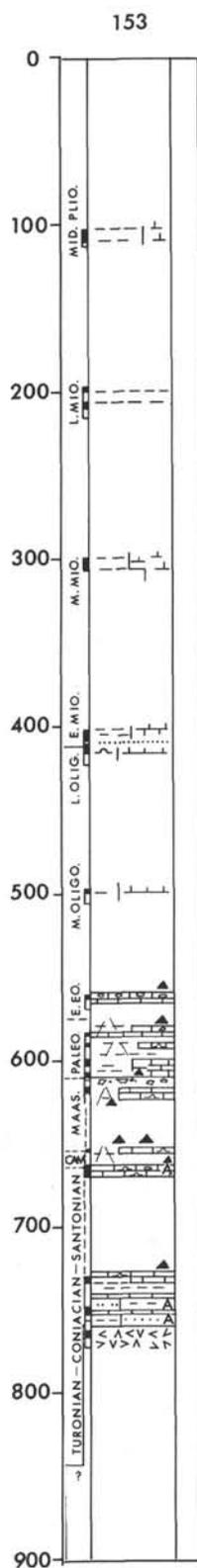
Thick soft sediments of Cenozoic age overlie hard, siliceous limestones and cherts of Early Eocene to Cretaceous (Coniacian) age. Horizon A'' marks the interface between these two units. Horizon B'' corresponds to a very fine-grained basalt suggestive of the top of a flow.

The limestone underlying Horizon A'' is irregularly silicified and contains interlayered cherts. Interbeds of carbonaceous, phosphatic, volcanic clay are found in the lower part.

BACKGROUND

Site 153 is located at the southern end of the Beata Ridge, between the ridge and South America (Aruba gap). A seismic refraction line (Edgar et al., 1971, profile 79) which lies over the continental margin and Venezuelan Basin to the east of this site shows a 3.1 and 3.6 km/sec seismic layer thinning and truncating against overlying lower velocity sediment that could be interpreted as pre-Horizon B'' layers and termination at about the Horizon B'' level (Figure 1). A processed seismic reflection line run over the area of the site by Esso Production Research Co. (Figure 2) also shows a wedge of sediment or rock below Horizon B'' truncating seaward. The location of these refraction and reflection lines are shown in Figure 3. The site had originally been selected if drilling to Horizon B'' in the Venezuelan Basin had failed, but the recovery of dolerites from sills made the objective of pre-Horizon B'' cores even more appealing.

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Seismic reflection records were supplied to us by Lamont-Doherty Geological Observatory on completion of a site survey by R/V *Vema*. *Glomar Challenger* reflection records approaching the site are shown in Figure 4; the track of the vessel is shown in Figure 5. There was a problem deciding on where to place the site—seaward of the oldest truncated layer where the oldest rock may lie close to Horizon B'' or landward where there was the possibility of not encountering igneous rock. The decision was made to drill a hole over the wedge on the assumption that Sites 146 and 150 had achieved essentially all that could be accomplished farther in the basin and we should attempt to sample the sediment wedge.

OBJECTIVES

The objective at this site was to penetrate Horizon B'' into layered rock that, according to seismic data, was truncated seaward and thickened rapidly beneath the continental slope of South America.

OPERATIONS

A recently received profile record from the R/V *Vema* precluded the necessity for detailed site surveying. The ship approached the site at 0800 hours on 23 January and passed near the site on a brief survey in order to find a location where both A'' and B'' were present. After spot coring five intervals of soft sediment above Horizon A'', continuous coring began at 563 meters where the first hard sediment was encountered (Horizon A''). Coring continued (with, however, one core representing 37 and another 64 meters of drilling) until basalt was reached at 766 meters. Poor recovery and loss of time retrieving and returning the barrel led us to core in excess of the normal 9-meter interval. The basalt was cored for an additional 10 meters in the hope of penetrating it, but when the drilling rate indicated that the bit was badly worn, the hole was terminated. The drill string was withdrawn and the ship was underway for Site 154 at 2130 hours on 27 January.

LITHOLOGY

The section above Horizon A'' was spot cored at five intervals. The uppermost core (102-111 m) is of a greenish gray nannoplankton calcareous clay. The next two cores (198-216 m) contain less than 10 percent calcium carbonate and the foraminifera are largely dissolved. The clay is soft and homogeneous. Minor volcanic material at the bottom of Core 1 includes phillipsite, quartz, hornblende, plagioclase, and biotite. In the core catcher of Core 2 is a dispersed turbidite with silt-sized quartz grains, manganese micronodules, bipyramidal quartz, gypsum, and pyritized foraminifera and Radiolaria. In this upper sequence the degree of solution of fossils is greater than in the lower interval.

The next lower core (300-309 m) consists of interbedded clay, similar to the first three cores, and light bluish gray marly calcareous clay, with layers 10 cm or more in thickness. The layering (Figure 6) is disrupted both by abundant burrows and by intense drilling disturbance. The carbonate content of the marly layers is about 50 percent. Foraminifera show the effects of solution. An unusual minor constituent of the clays of Core 4 are minute

(less than 10 microns) euhedral crystals of yellow to brown garnet. Dolomite (or siderite) is persistent, but minor.

The downward increase in carbonate marks the next cores which are light bluish gray marly foram-nanno chalk and bluish white nannofossil chalk (403-421 and 499-500 m). The carbonate content varies from 50 to 80 percent. Radiolaria and nannofossils are well preserved, but foraminifera show signs of solution. These sediments are compact but neither indurated nor lithified. Dolomite is scarce; volcanic debris is absent.

A sharp drop in the drilling rate at 555 meters probably identifies Horizon A''. Core 8 was taken at 563 meters and Early Eocene silicified hard limestone and chert were recovered.

Coring below Horizon A'' recovered a greenish gray limestone sequence with irregularly silicified zones and interlayered cherts. Burrowing is well developed. Minor volcanic ash layers consist of montmorillonite with clinoptilolite, biotite, apatite, and zircon. Barite is conspicuous in some of the lower limestones. Radiolaria occur in the upper three cores (9, 10 and 11) and foraminifera and nannofossils are found throughout, although they are generally poorly preserved. In the interval 586 to 609 meters, most of the foraminifera are benthonic forms.

At the Cretaceous-Tertiary boundary (Sample 153-12-1, 113; 611 m), a silicified breccia (Figure 7) somewhat resembling the hardground of Site 151, Core 11, occurs. The fragments of the breccia are separated by calcite-filled veins and have matching boundaries.

In the lower part of the hole (below 730 m), the limestone is interlayered with a carbonaceous, phosphatic, volcanic clay (Figure 8). This clay is olive black and has a variable carbonate content. Fish debris, including complete bones (some of them articulated), is abundant. Minor barite crystals, zeolites, and horizontal chevron burrows (*Zoophycos*) are conspicuous in some zones (see Warne et al. and Maurasse, this volume). Pea-sized nodules of sparry calcite, chalcedony, and significant (1-2%) reduced organic matter are found. Towards the base of the unit, the ash component increases and nearly pure, graded ash beds a few centimeters in thickness are found (Figure 9).

The lower limestones show small-scale, dominantly horizontal burrowing, much of which has been secondarily distorted. Pyrite is scattered throughout, and one concretion several centimeters in diameter was found (Figure 8; Frontispiece). Glauconite streaks are present.

Although the upper contact with the basalt was not recovered, the vitric groundmass and very fine grain size of the uppermost basalt sample suggests that it occurred near the flow top. About three meters of basalt were recovered; it is fresh, sparsely and finely amygdaloidal, and slightly fractured. There is no indication of the basalt being pillowed.

PHYSICAL PROPERTIES

Wet-bulk Density, Water Content, and Porosity

Wet-bulk density and porosity were measured by two methods aboard the *Glomar Challenger*: Gamma Ray Attenuation Porosity Evaluator (GRAPE) and individual

sample volume-weight measurements (the sample data are the enclosed dots in the hole and core plots). Water content was determined by weight-weight relationships. In general the precision of these data is ± 5 percent. Methods, errors, assumptions, corrections, sediment disturbances, and interpretation precautions are discussed in Appendix I.

Some stiff sediment or rock is cored without plastic flowage of the sample and, since the drill bit has a smaller diameter than the core liner (2.60 in. I.D.), the hard sediment samples also have that smaller diameter, with the remaining space being filled with a drilling slurry or highly disturbed sediment (in some cases air). A problem arises here because a 2.60 inch diameter is used in the density calculations.

Where necessary, the GRAPE data of Site 153 are adjusted (GRAPE corrected diameter [GCD] is the dotted line density and porosity in the core and hole plots) for the smaller diameter of the core. Diameters were generally smaller than 2.60 inches by 9% in Cores 4, 5, and 6; 10% in Cores 13, 16, and 17; 12% in 7, 8, 9, 12, 18, and 19; and 13.5% in Cores 10, 11, and 14. A typical slurry density of 1.1 g/cc was assumed in these calculations. See the Appendix for a discussion and for the equations used for correction.

Only the maximum density trends and minimum porosity trends were recalculated since lesser densities and greater porosities are suspect of being drill-disturbed sediments, smaller diameters, or rock fragments. These adjusted data are only approximations.

Results

Wet-bulk densities at Site 153 from Pliocene clay at 100 meters below bottom to Cretaceous micritic limestone at 760 meters vary from 1.50 g/cc to 2.45 g/cc (67 to 12% porosity) and increase irregularly with increasing depth. Specifically, wet-bulk densities of Pliocene nannofossil clay at 100 meters depth are 1.55 to 1.68 g/cc (68 to 60% porosity) but wet-bulk densities gradually increase to 1.70 to 1.90 g/cc (58 to 47% porosity) in similar Miocene sediments at 300 meters. Densities in Miocene-Oligocene marl and foraminiferal nannofossil chalk at 400 meters depth and Oligocene nannofossil chalk and clay at 500 meters depth are about the same as those sediments at 300 meters, although deeper and older. However, porosities appear to continue to decrease while the wet-bulk density remains constant. This is because of decreasing grain density which is related to the occurrences of radiolarians and sponge spicules.

Eocene-Paleocene micritic limestone (rare nannofossils) from 560 meters to 600 meters have wet-bulk densities of 2.2 g/cc to 1.8 g/cc (13 to 33% porosity). Paleocene-Cretaceous (?) micritic limestone (with minor marl and clay; silica in the lower section) have typical densities of 1.9 g/cc to 2.3 g/cc (31 to 14% porosity), with a range from 1.85 to 2.55 g/cc (45 to effectively 0% porosity?). The sand overlying the basalt has a density of 2.3 g/cc (33% porosity), while the density of the basalt is about 2.7 g/cc.

In Core 1, a high GRAPE density spike was caused by a pyrite nodule. The Core 3 graph also has a spike which may be a nodule. However, since these rocks were surrounded by sediment, the recorded density is not a true density of the nodules.

Water content samples were collected at a depth of 200 meters from clay, and at 300 meters from marl, with values of 33 percent and 31 percent, respectively.

Sound Velocity

Sound velocity through sediment and rock samples was measured by the Hamilton Frame technique, which is discussed in the Appendix. This method has a precision within ± 1.1 percent.

Sound velocity was measured on sediments and rocks which appeared to be physically undisturbed. These velocities were measured parallel to the bedding planes, unless otherwise noted in Table 1, and were measured at laboratory pressures and temperatures (25.4 to 28.5°C).

Results

At Site 153, sound velocities range from 1.57 to 5.63 km/sec, and irregularly increased with increasing depth. Specifically, between 200 and 500 meters, velocities increased from 1.57 to 2.11 km/sec in Miocene calcareous clay, marl, and chalk, and Oligocene nannofossil chalk. Between 609 and 610 meters, Paleocene silicified marl had higher velocities of 2.13 to 2.52 km/sec. Paleocene and Cretaceous limestone from 588 to 620 meters had velocities of 2.02 to 3.26 km/sec with typical velocities within 2.0 to 2.5 km/sec, while the Cretaceous limestones from 620 to 750 meters had greater velocities of 2.45 to 3.79 km/sec with 3.14 to 3.60 km/sec being typical. A similar trend of increasing velocities with increasing depth can be seen in the siliceous limestone velocities. Eocene and Paleocene siliceous limestone from 563 meters to 612 meters has sonic velocities of 2.23 to 3.42 km/sec, while Cretaceous siliceous limestone from 620 to 717 meters have velocities of 3.01 to 4.06 km/sec.

Eocene chert recovered at 563 and 572 meters has sound velocities of 3.90 to 4.50 km/sec, which was low compared to 5.20 to 5.63 km/sec measured through five Cretaceous conchoidally fractured chert fragments from 650 to 700 meters.

Volcanic shale and sandstone above (751 to 754 m) the basalt has sound velocities of 2.38 to 2.50 km/sec, which contrasted with the 4.78 km/sec of the basalt below these sediments.

Natural Gamma Radiation

Natural gamma ray emissions were counted for a period of 1.25 minutes at 7.62 cm (3 in.) intervals along the core, with a counting precision of about ± 100 counts. Methods, equipment, error porosity corrections, and sediment disturbance are discussed in the Appendix. The following data were not corrected for varying porosity.

Natural gamma radiation emitted from sediments recovered at Site 153 ranged from 0 to 3200 counts per 1.25 minutes at a 7.62 cm core interval. The Santonian-Coniacian micritic limestone from 728 meters to 770 meters below the sea floor, with organic- and phosphate-rich clastic beds, had the highest emissions of 300 to 3200 counts. The Pliocene-Miocene nannofossil clay, clay, and marl from 100 to 300 meters had intermediate counts of 1300 to 2600 while 200 to 1000 counts were emitted from the Miocene and Oligocene foraminiferal nannofossil chalk

TABLE 1
Hamilton Frame Sonic Velocities, Site 153

Core	Section	Upper Interval ^a (cm)	Depth in Hole (m)	Velocity ^b (m/sec)	Temperature (°C)	Remarks
3	1	32.0	207.32	1574	25.5	Clay.
3	1	114.0	208.14	1624	25.4	Clay.
4	1	0.0	300.00	1631	27.1	Clay, zeolitic.
4	2	40.0	301.90	2074	27.1	Clay, nanno-rich.
4	3	123.0	304.23	1951	27.1	Nanno marl ooze.
5	1	60.0	403.60	1669	26.7	Foram nanno marl.
5	2	85.0	405.35	1693	26.7	Foram nanno chalk, clay-rich.
5	3	85.0	406.85	1823	26.5	Foram nanno chalk, clay-rich.
5	4	52.0	408.02	1841	26.5	Nanno foram marl.
5	4	138.0	408.88	2110	26.5	Foram nanno chalk, clay-rich.
6	1	84.0	412.84	2031	26.5	Nanno chalk, clay-rich.
6	2	143.0	414.93	2102	26.5	Nanno chalk, clay-rich.
7	1	75.0	499.75	2093	25.5	Nanno chalk, clay-rich.
8	1	84.0	563.84	4498	28.0	Chert, calcareous.
8	1	92.0	563.92	2673	28.0	Limestone, silicified.
8	CC	0.0	572.00	3904	(Cold)	Limestone, silicified, velocity measured when core arrived at surface, thus "cold"
8	CC	0.0	572.00	3868	(Cold)	Limestone, silicified.
9	1	90.0	586.90	3204	27.7	Marlstone, silicified.
9	1	110.0	587.10	3175	27.7	Marlstone, silicified, dark.
9	2	33.0	587.83	2593	26.5	Marlstone, slightly silicified.
9	2	90.0	588.40	2496	26.5	Limestone, micritic.
10	1	145.0	592.45	2233	26.7	Marlstone, micritic.
10	2	84.0	593.34	2389	26.3	Claystone, silicified.
10	2	84.0	593.34	2396	26.3	Claystone, silicified.
10	2	143.0	593.93	2083	26.2	Micritic marl.
11	1	88.0	600.88	2165	26.2	Nanno marl.
11	2	96.0	602.46	2110	25.4	Nanno marl.
11	3	42.0	603.42	2285	27.6	Nanno marl.
11	3	139.0	604.39	2653	27.6	Claystone, zeolitic, calcareous.
12	1	70.0	609.70	2026	25.3	Nanno marl.
12	1	90.0	609.90	2830	25.3	Marl, silicified.
12	1	125.0	610.25	2218	25.3	Foram limestone, clay-rich, light.
12	1	125.0	610.25	2131	25.3	Foram limestone, clay-rich, dark.
12	2	27.0	610.77	2528	25.3	Foram limestone, clay-rich.
12	3	0.0	612.00	3425	27.2	Limestone, cherty.
12	3	140.0	613.40	3260	27.0	Limestone, clay-rich.
13	2	13.0	620.63	3602	26.3	Limestone, clay-rich.
13	2	92.0	621.42	3407	26.3	Limestone, clay-rich.
13	4	30.0	623.80	2910	28.5	Limestone, clay-rich.
13	4	83.0	624.33	2455	28.5	Limestone, clay-rich.
14	1	87.0	656.87	5634	27.2	Chert, brownish yellow.
14	1	97.0	656.97	5121	27.2	Chert, brownish gray.
14	1	103.0	657.03	3135	27.2	Limestone
14	1	109.0	657.09	5227	27.2	Chert, reddish brown.
15	1	27.0	667.27	5273	26.8	Chert, gray and tan layers.
15	1	84.0	667.84	5231	26.8	Chert, tan.
15	1	117.0	668.17	3011	26.8	Radiolarian foram limestone, "sandy."
15	1	145.0	668.45	3308	26.8	Radiolarian foram micritic limestone, partly silicified.
15	2	38.0	668.88	3429	26.0	Radiolarian foram micritic limestone, partly silicified.
15	2	100.0	669.50	3639	26.0	Radiolarian foram micritic limestone, partly silicified.
15	3	15.0	670.15	5198	25.8	Chert, gray.
15	3	70.0	670.70	3680	25.8	Radiolarian foram micritic limestone, to bedding, partly silicified.
15	3	70.0	670.70	3243	25.8	Radiolarian foram micritic limestone, ⊥ to bedding, partly silicified.
15	3	85.0	670.85	3417	25.8	Micritic limestone, partly silicified, some ash.
15	4	29.0	671.79	3200	25.5	Radiolarian foram micritic limestone, slightly silicified.
15	4	49.0	671.99	4397	25.5	Radiolarian foram micritic limestone, to bedding, silicified.
15	4	49.0	671.99	4497	25.5	Radiolarian foram micritic limestone, silicified, ⊥ to bedding.

TABLE 1 - Continued

Core	Section	Interval ^a (cm)	Depth in Hole (m)	Velocity (m/sec)	Temperature (°C)	Remarks
15	4	83.0	672.33	4067	25.5	Radiolarian foram micritic limestone, silicified.
16	1	36.0	731.36	3783	25.7	Foram limestone.
16	1	120.0	732.20	3780	25.7	Foram limestone, with burrows.
16	2	95.0	733.45	2720	25.7	Claystone, black, carbonaceous.
16	2	103.0	733.53	3449	25.7	Foram limestone.
18	2	13.0	750.63	3789	27.7	Micritic limestone.
18	2	50.0	751.00	2402	27.7	Shale, black, chloritic.
18	3	28.0	752.28	2434	26.5	Shale, black, chloritic.
18	3	90.0	752.90	2380	26.5	Shale, black, chloritic.
18	4	87.0	754.37	2503	27.4	Sandstone, volcanic, coarse grained.
18	4	102.0	754.52	2387	27.4	Sandstone, volcanic, fine-grained.
20	1	145.0	768.45	4785	27.4	Basalt, fractured.

^aUpper limit of a 3-cm sample interval.

^bVelocities were measured parallel to bedding unless noted otherwise.

and marl from 400 meters depth and foraminiferal nannofossil chalk and minor clay from 500 meters depth. The marl and clay emitted the higher counts, while the chalk emitted the lower counts of 200 to 600. Paleocene-Cretaceous micritic limestone with minor marl from 580 to 625 meters emitted counts from 100 to 700 while the Eocene micrite-sparite chalk, at 560 meters, and Cretaceous micritic limestone from 650 to 665 meters, and 770 meters, emitted very low counts of 100. The basalt only emitted 200 to 300 counts.

Penetrometer

Needle penetration tests were conducted at Site 153 with a 1 mm diameter needle. The methods, equipment, and sediment disturbance are discussed in the Appendix.

Penetration tests were performed on Cores 1, 3, and 4 from 100, 200, and 300 meters below the sea floor, where penetration decreased with increasing depth. Penetration into Pliocene nannofossil clay from 100 meters was 4 mm; penetration into Miocene clay at 200 meters depth was 1 to 2 mm; and Miocene clay and marl at 300 meters had penetrations of 1 to 2 mm and less than 1 mm, respectively. Below this depth, penetrations into Miocene foraminiferal chalk were not significant.

BIOSTRATIGRAPHY

Spot cores were taken through the upper part of the sedimentary sequence at this site.

Core 1 (102-111 m) contains rich calcareous planktonic fossils, but lacks radiolarians. It belongs to the *Globigerinoides trilobus fistulosus* Subzone of the *Globorotalia miocenica* Zone (Middle Pliocene). It contains calcareous nannofossils of the *Discoaster surculus* Zone.

Cores 2 and 3 (198-216 m) contain sparse planktonic foraminifera, strongly affected by solution, poor assemblages of calcareous nannofossils consisting almost exclusively of asterooliths, and no radiolarians. These cores probably belong to the *Globorotalia acostaensis* and *Discoaster quinqueramus* zones of the Late Miocene.

Core 4 (300-309 m) contains a good assemblage of the *Globorotalia fohsi fohsi* Zone and abundant but only moderately diverse calcareous nannofossils of the *Sphenolithus heteromorphus* Zone. Like the cores above it at this site, it also lacks radiolarians.

Core 5 (403-412 m) contains a very condensed Middle-Early Miocene sequence reminiscent of those in Core 4 at Site 150 and Core 4 at Site 151. A sample from Section 1 of Core 5 belongs to the *Globorotalia fohsi peripheroronda* Zone, samples from the lower part of Section 3 and upper part of Section 4 the *Globigerinatella insueta* Zone, and the catcher of Core 5 the *Globigerinoides primordius* Zone. The *Praeorbulina glomerosa*, *Globigerinita stainforthi*, and *Globigerinita dissimilis* zones were not detected but this may simply reflect the sampling interval used. Calcareous nannoplankton through the top of Section 4 of Core 5 belongs to the *Sphenolithus heteromorphus* Zone. A sample from the lower part of Section 4 belongs to the *Sphenolithus belemnus* Zone. The zones which should immediately overlie and underlie the *Sphenolithus belemnus* Zone, the *Helicopontosphaera ampliapertura* and *Discoaster druggi* zones, were not detected, but again this might be due to the sampling interval used. The lower part of Core 5 and Core 6 contain common, moderately to well-preserved (somewhat dissolved) radiolarians of the *Calocyclus virginis* Zone.

Core 6, which is immediately subjacent to Core 5 (412-421 m) contains planktonic foraminifera of the *Globorotalia kugleri* Zone of the Late Oligocene and calcareous nannoplankton of the *Triquetrorhabdulus carinatus* Zone.

Core 7 (499-508 m) contains assemblages of the *Globigerina ampliapertura* and *Sphenolithus predistentus* zones of the Middle Oligocene.

Sediment recovered from the center bit after drilling between Cores 7 and 8 contains *Globorotalia opima opima*, indicating that the *Globorotalia opima opima* Zone is present at this site; it must lie between Cores 6 and 7 (421-499 m).

Core 8 (563-572 m), lies in early Eocene deposits according to the planktonic foraminifera present in a

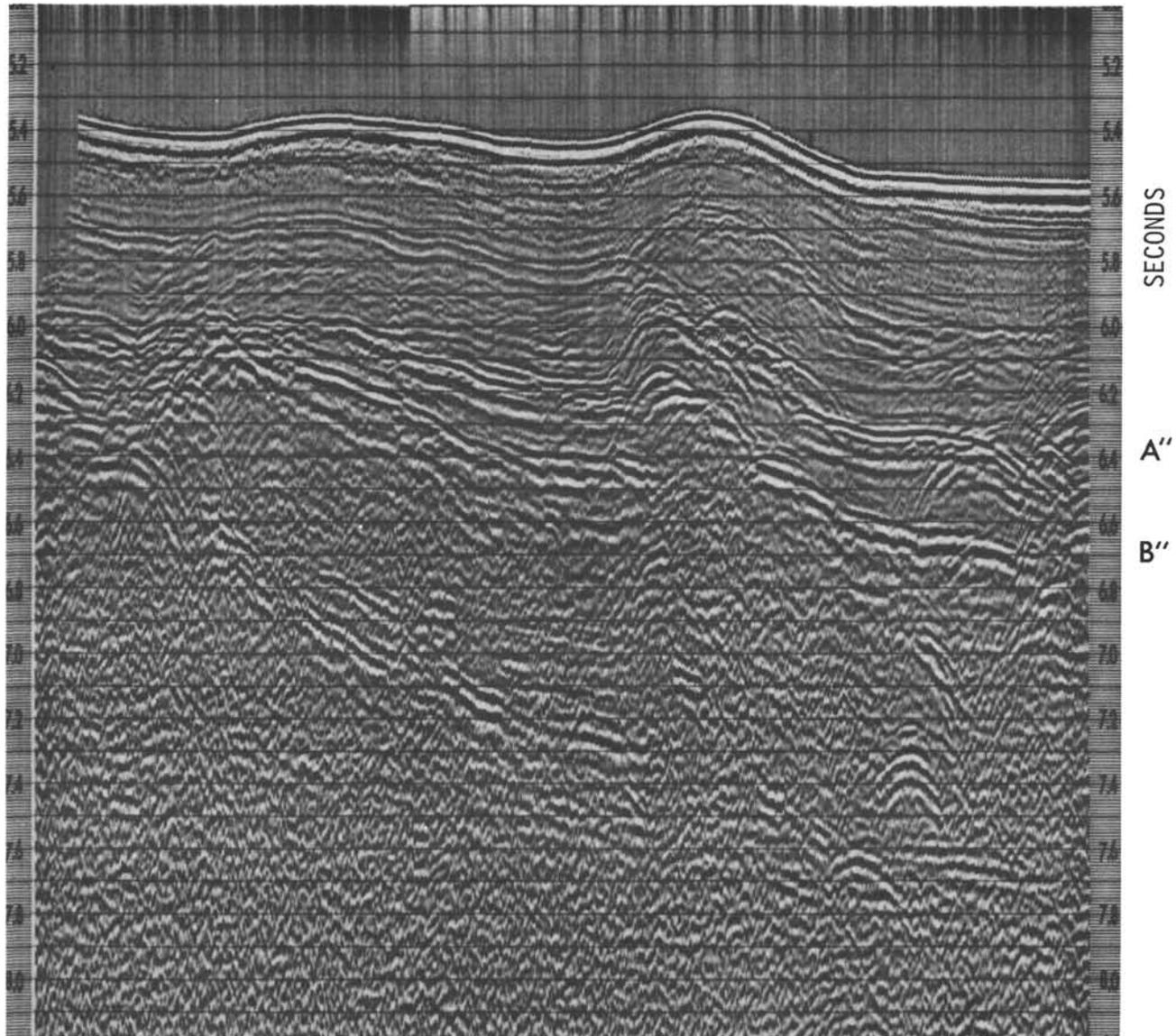


Figure 2. Processed seismic reflection line extending from the base of the continental slope of South America into the Caribbean Sea showing a thick wedge of possible sediment or volcanic rocks lying below Horizon B'' and thinning seaward. It terminates on the south edge of the Beata Ridge. Track shown on Figure 3. (Courtesy of Esso Production Research Company.)

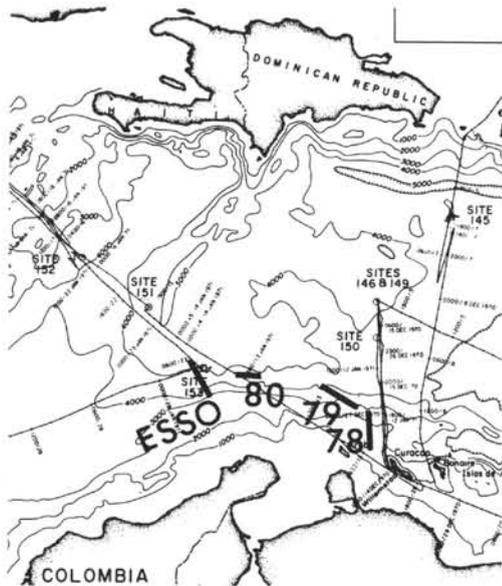


Figure 3. Chart showing location of refraction profiles of Figure 1 and reflection line of Figure 2.

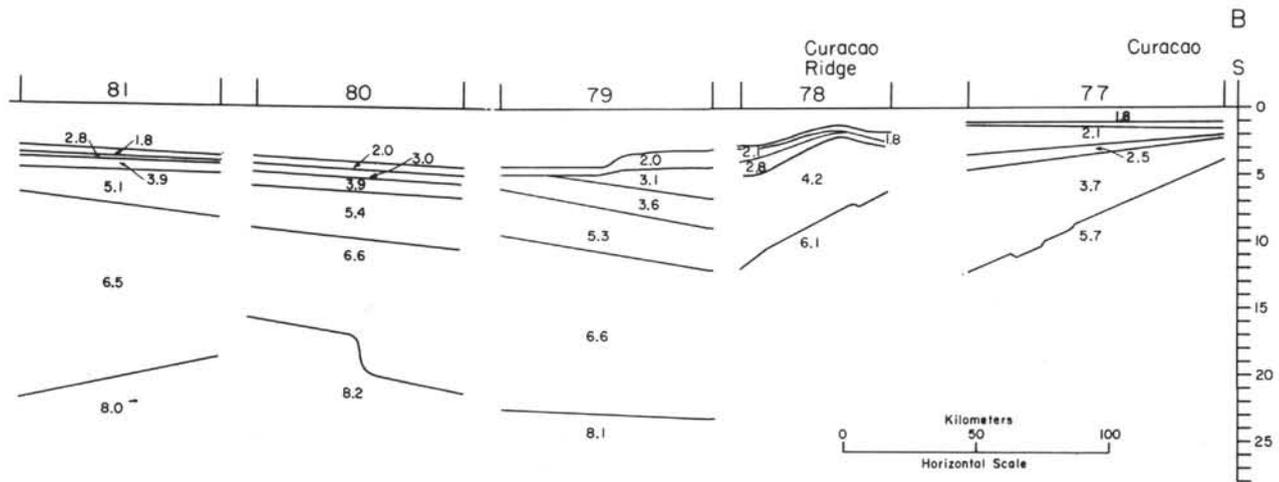


Figure 1. Seismic refraction profiles shot across the northern continental margin of South America showing the truncation of a wedge of 3.1 and 3.6 km/sec layers in the vicinity of Site 153. Track shown on Figure 3 (Edgar et al., 1971).

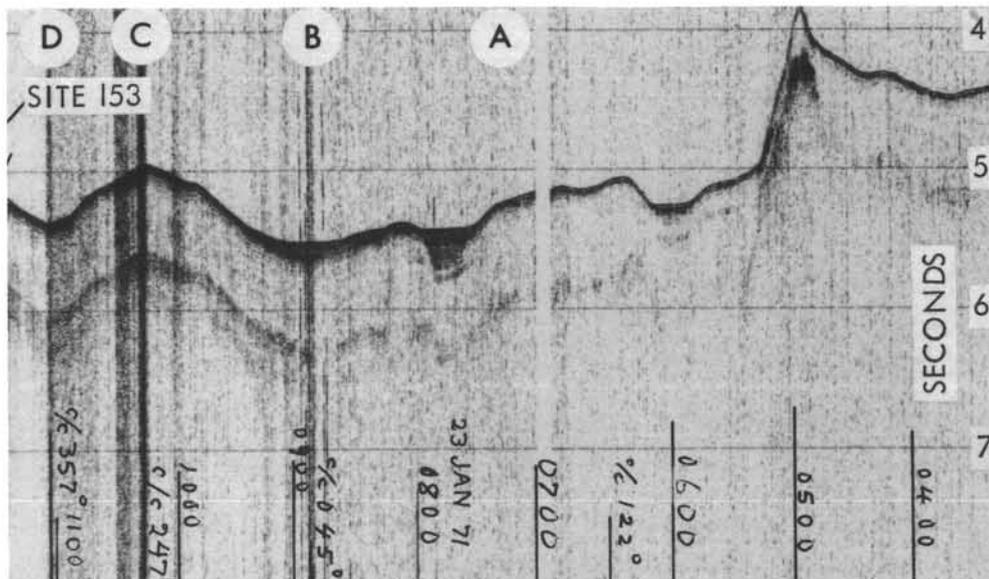


Figure 4. Profiler record made by Glomar Challenger on approach to Site 153. Track shown on Figure 5.

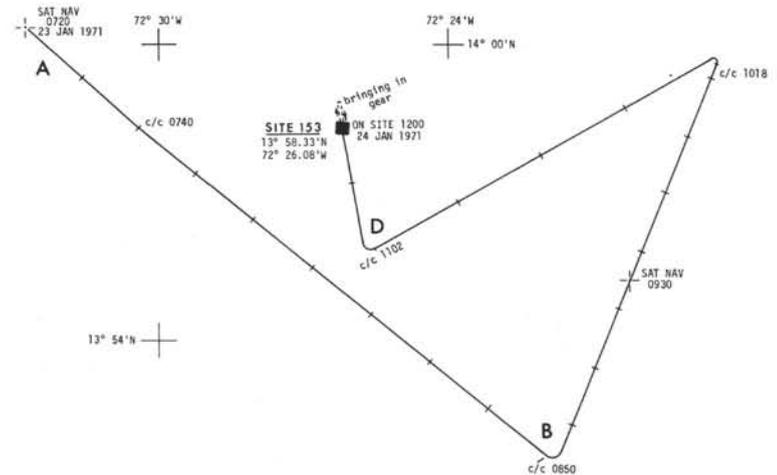


Figure 5. Track of Glomar Challenger on approach to Site 153.



Figure 6. Interbedded clay and marly chalk ooze showing intense burrowing (153-4-3, 50-70).

sample from Section 1; they are indicative of the *Globorotalia formosa formosa* Zone. Calcareous nannofossils are rare in this material and are questionably assigned to the Paleocene *Discoaster gemmeus* Zone, in which case they must be mixed or reworked into the younger Eocene sediment.

Continuous coring was carried out between 586 and 619 meters. Cores 9, 10, and 11 through Section 3 are generally poor in planktonic foraminifera although a few thin layers contain richer assemblages characteristic of the Paleocene *Globorotalia pusilla pusilla* Zone. These cores also contain sparse nannofossils belonging to the *Gasiculithus tympaniformis* Zone.

The catcher sample of Core 11 contains a planktonic foraminiferal assemblage of the *Globorotalia angulata* Zone.

Section 1 of Core 12 (609-619 m) is devoid of planktonic foraminifera to a depth of 90 cm, but does contain calcareous nannofossils of the *Chiasmolithus*



Figure 7. Limestone breccia occurring at a hiatus separating Cretaceous and Tertiary limestones (153-12-1, 95-118).

danicus Zone. Between 90 and 102 cm is a breccia containing fragments of micrite some of which contain planktonic foraminifera of the *Globigerina eugubina* Zone.

The breccia, which marks the base of the Cenozoic, is underlain by a few centimeters of sediment devoid of plankton fossils, but the lower parts of Core 12 contain globotruncanids of the Maestrichtian *Globotruncana gansseri* and *Globotruncana "tricarinata"* Zones and poor calcareous nannofossils.



Figure 8. Interbedded limestone and carbonaceous sediment. Note intense burrowing and deformation in some limestone layers. A large pyrite nodule is found in the upper part of the picture (see frontispiece). The carbonaceous layers are rich in fish remains and volcanic debris and contain conspicuous *Zoophycus* burrows (153-16-2, 87-132).

Core 13 was taken by drilling between 619 and 656 meters and contains assemblages of the *Globotruncana* "tricarinata" Zone.

Core 14 (656-667 m) contained only benthonic foraminifera and rare nannofossils.

Cores 15 and 16 (667-740 m) contain planktonic foraminifera of the *Globotruncana concavata concavata* Zone of the Coniacian.



Figure 9. Graded beds of basaltic ash immediately above basalt flow (153-19-1, 79-99).

Core 17 (740-749 m) yielded only a catcher sample with benthonic foraminifera.

Cores 18 and 19 (749-767 m) contain layers with good assemblages of the Late Turonian *Globotruncana schneegansi* Zone and layers devoid of foraminifera. Calcareous nannofossils are rare to absent.

CONCLUSIONS

Although only spot cores were taken above Horizon A'', we can draw several conclusions about this interval. The occurrence of clays in higher intervals and carbonate sediments in the early Miocene indicates that the fluctuation of the depth of calcium carbonate compensation seen at Sites 146/149 and adjacent sites affected this area also. The high accumulation rate of the clay-rich part of the section (around 25 m/m.y.) is noteworthy. Very minor, possibly turbiditic debris of obviously continental

origin was found. Very little of the section is volcanic and the high accumulation rates evidently reflect proximity to the South American continent. Possibly in the pre-Pliocene interval, before the Panamanian isthmus was raised, strong surface currents may have carried a large terrigenous clay sediment load to this site.

Beneath the cherts and lithified limestones of Horizon A", the lower limestone section is presently deeper than the contemporaneous, carbonate-free Paleocene of Site 146, suggesting that there may have been substantial differential vertical movement between the two sites since their deposition or a substantial difference in the type of sediment being deposited at the two areas.

The silicified breccia and chert at the Cretaceous-Tertiary boundary of Core 12 is similar to the ferruginous, siliceous, hardground of Core 11, Site 151. In both cases foraminiferal oozes have been silicified, and later compaction and volume changes may have accounted for local brecciation. Very probably, both are hardgrounds and

both represent temporally parallel (if not fully coincidental) epochs of relative shallowing, exposure to moderate bottom currents, and penecontemporaneous silicification.

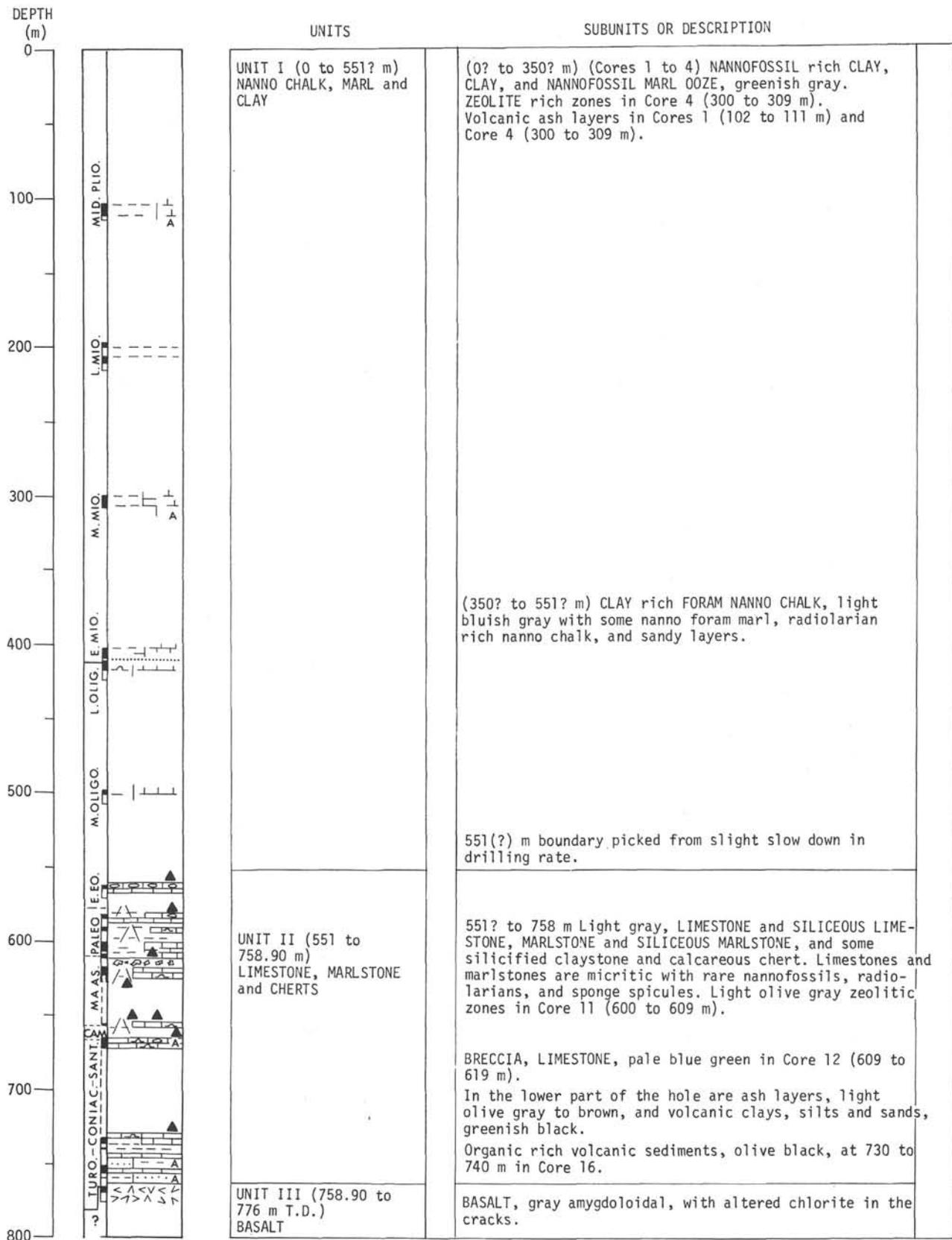
The interlayering of organic, phosphatic, and volcanic clays with the limestones at the lower part of the section indicates, as at Site 146, the presence of a topographically diverse sea floor, with stagnant low areas in which fish debris and other organic matter accumulated with incomplete decomposition. The occurrence of basaltic ash at this level suggests further that this topographic disruption of the sea floor may have accompanied widespread basaltic eruptions in the central Caribbean.

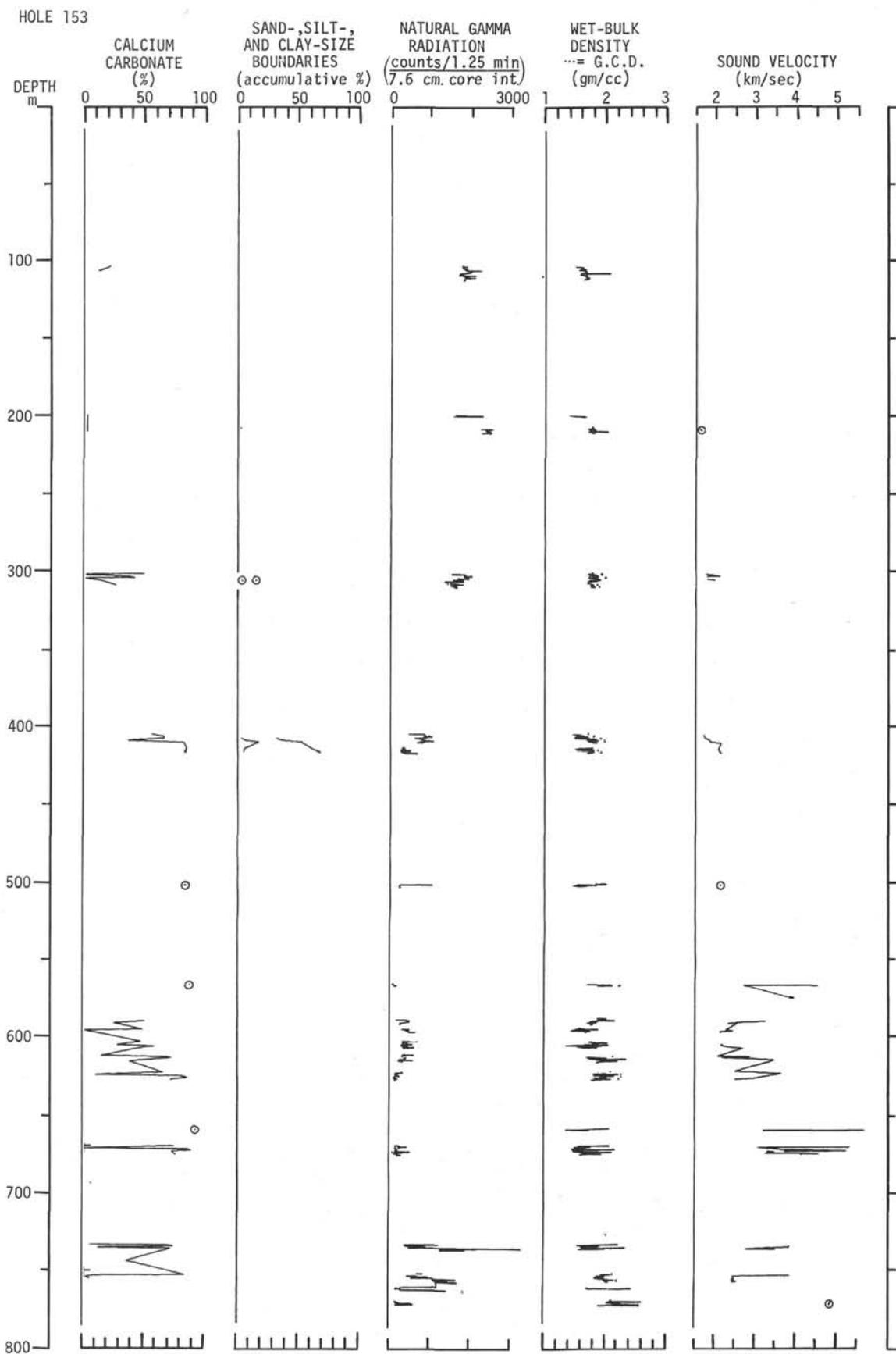
REFERENCE

- Edgar, N. T., Ewing, J. I. and Hennion, J., 1971. Seismic refraction and reflection in Caribbean Sea. *Am. Assoc. Petrol. Geologists Bull.* 55, 833.

SITE 153

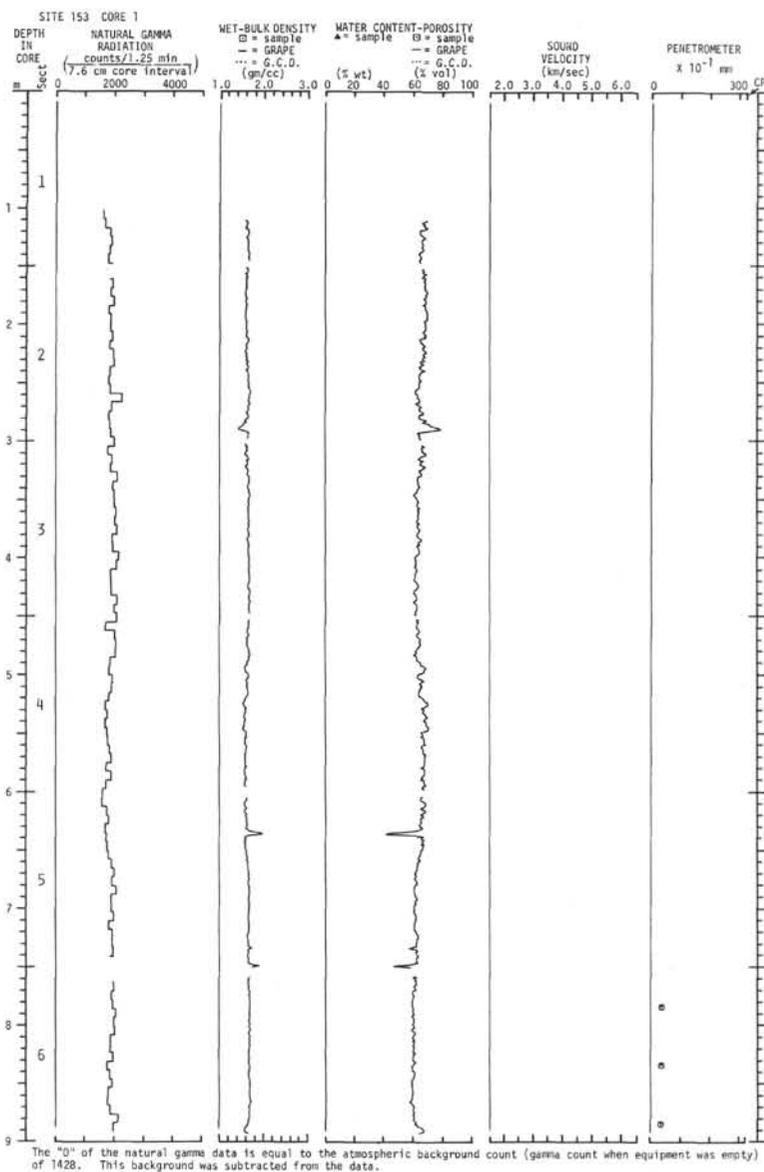
LITHOLOGY





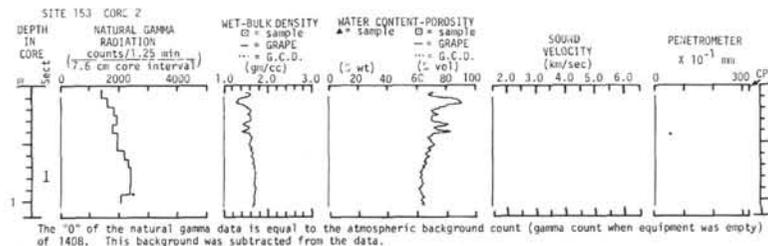
SITE 153 HOLE CORE 1 CORED INTERVAL (m) 102-111

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION			
	FORAM	NANNO	RAD												
MIDDLE PLIOCENE	Globorotalia miocenica (Globigermoides trilobus fistulosus sz)	Discoaster surculus		1	0.5	VOID			NANNOPLANKTON CALCAREOUS CLAY; varied in color from greenish gray to grayish olive.						
					1.0			F A W N C W	5GY5/1 5GY5/2	Iron sulfide staining is common and pyrite (indicated by "**") occurs as shell replacements. Pyrite increases slightly with depth. Some carbonate rhombs.	▲				
					0.0					5GY6/1 plagioclase	A light olive gray layer in Section 6 contains volcanic debris, rare quartz and chlorite, common zeolites and hornblende.		III		
				2	0.5					N C W	The entire core is disturbed.	▲			
					1.0					5GY4/1	Sediment is soft, plastic, and homogeneous.				
					0.0					10GY4/2 and 5GY6/1	III. Flow-in: intimate mixture.		II		
				3	0.5					N C W F A W		▲			
					1.0					5Y4/1 to 5Y5/2			·III		
				4	0.5						*(replacing a cylindrical burrow)		II		
					1.0					5GY4/1 and 5Y4/1					
				5	0.5					N F M F A W					
					1.0					N C W	*				
				6	0.5						5Y4/1			III	
					1.0					F A W		biotite, hornblende, plagioclase, phillipsite conspicuous			
										VOID					
										CORE CATCHER	F A W N C W				

¹For explanation of symbols, see Chapter 1

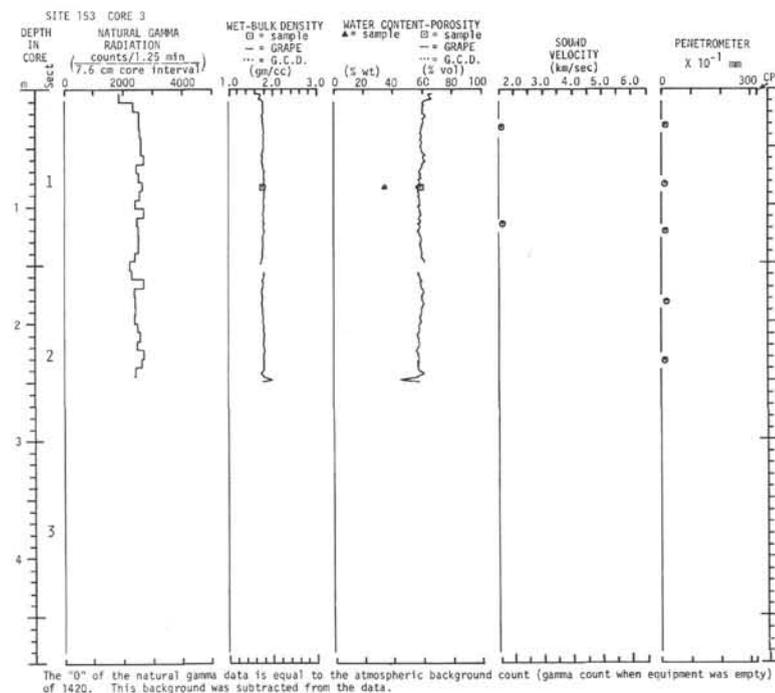
ZONE		SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%) SAND-SILT-CLAY (accumulative %)	DEFORMATION
AGE	FORAM								
LATE MIOCENE	Globorotalia acostrensis		0.5 1.0				5GY4/1		
	Discosaster quincranus								
		CORE CATCHER					<p>CLAY; drilling breccia (V) with iron sulfide streaking, becomes more compact below 67 cm. No horizontal structures observed.</p> <p>Core catcher contains detrital grains of quartz, plagioclase, orthoclase, zoisite, and amphibole. Fish teeth, gypsum micronodules, biotite, and apatite. Radiolaria replaced by pyrite also occur. In addition, X-ray diffraction shows K-feldspar.</p> <p>Analyses of pellets in the core catcher shows the following:</p> <p>FE₂O₃ 2.0% MnO 46 % Ba 22000 ppm Sr 2000 ppm Cu <200 ppm</p>		V

¹For explanation of symbols, see Chapter 1



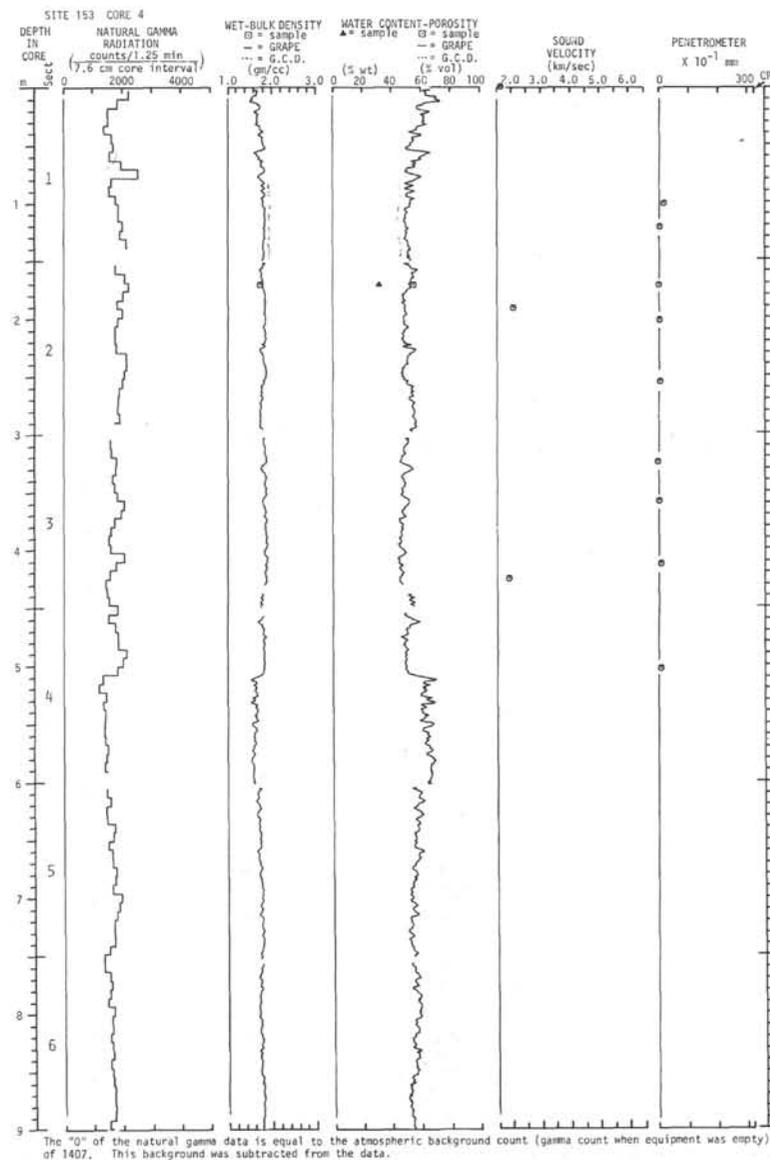
SITE 153 HOLE CORE 3 CORED INTERVAL (m) 207-216

AGE	ZONE		SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NANNO									
LATE MIOCENE	Globobrotalia acostaensis	Discoaster quinqueramus	1	0.5	CLAY; compact and homogeneous with abundant iron sulfide streaking. Slight mottling occurs. Sediment becomes less compact in Sections 3 and 4 due to disturbance. Pyrite is abundant to common in lower part of Section 1. Sparse plagioclase, scattered carbonate rhombs, and pyrite. In addition, X-ray shows K-feldspar.	N R P F C M		56Y4/1	0	50	100
				1.0							
LATE MIOCENE	Globobrotalia acostaensis	Discoaster quinqueramus	2	0.0	VOID	n			0	50	100
				0.5							
LATE MIOCENE	Globobrotalia acostaensis	Discoaster quinqueramus	3	0.0	VOID	N F P			0	50	100
				0.5							
			CORE CATCHER			R O F C M					

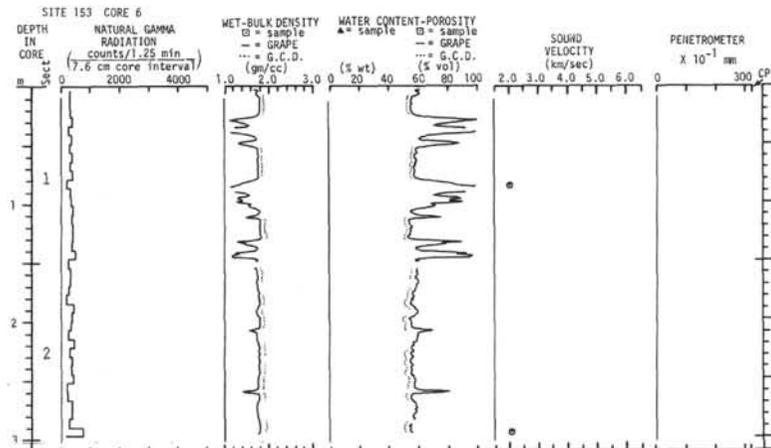
¹For explanation of symbols, see Chapter 1

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%) SAND-SILT-CLAY (accumulative %)	DEFORMATION	
	FORAM	NANNO	RAD									
MIDDLE MIOCENE	Globorotalia fohsi Sphenolithus heteromorphus			1	0.5	<p>Alternations of:</p> <p>NANNOPLANKTON MARL OOZE; light greenish gray (5G7/1) with irregular mottles of clay.</p> <p>CALCAREOUS CLAY; greenish gray (5G6/1) with common mottling and iron sulfide streaks.</p> <p>ZEOLITIC CLAY; dark greenish gray (5G4/1), displays no carbonate.</p> <p>Burrowing is common to intense throughout core. Contacts between alternating sediment types vary from sharp to gradational. Most laminae are horizontal but some are inclined. Carbonate rhombs throughout. Sparse high-refractive-index carbonate rhombs and fish debris. In addition, X-ray shows pyrite. Low-carbonate clays (<10%) appear silicified. Core 4 is more compact than previous cores based on penetrometer data.</p> <p>II and II. Disturbed zone begins near top of Section 4 and is a mixture of all three sediment types.</p> <p>** indicates plagioclase, alkali feldspar, phillipsite, zircon, chert micronodules, apatite, and anatase.</p>	0					
				1	1.0		N C W F A M					
				2	0.0							
				2	0.5							
				2	1.0		N C M	clay layer				
				3	0.0							
				3	0.5			gypsum garnet				
				3	1.0							
				4	0.0							
				4	0.5							
				4	1.0		N C M	chert micronodules				
				5	0.0							
5	0.5											
5	1.0	N C M					II to III					
6	0.0											
6	0.5		chert micronodules									
6	1.0	N C M										
				CORE CATCHER			F A M					

¹For explanation of symbols, see Chapter 1

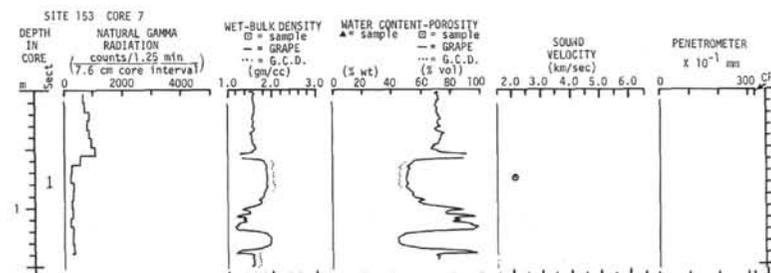


AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NANNO	RAD									
LATE OLILOCENE - EARLY MIOCENE	Globorotalia kugleri	Trifarina angulosa	Calocyclus virginis	1	0.5	VOID	N C P		NANNOPLANKTON CHALK; bluish white (589/1) to light gray (N7) uniform chalk containing nannofossils and micrite with foraminifers, radiolaria, and sponge spicules. Mottling is slight to common and is greenish black (5GY2/1). Core is very compacted and indurated. The acetate peel indicates that the components are unoriented and unsorted, and their chambers are either empty or micrite filled. Sparse plagioclase. X-ray also shows K-feldspar. One blackish red (5R2/2) layer occurs at 43-47 cm in Section 2. The lower boundary is sharp and the upper gradational. Composition is similar to remaining core except for a trace of dolomite.	0	50	100
				1.0	F A P							
				2	0.0		R A M					
					0.5		N C P					
					1.0		N C P					
							R C M					
				CORE CATCHER								



The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1396. This background was subtracted from the data.

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NANNO	RAD									
MIDDLE OLILOCENE	Globigerina amabilapertura	Sphenolithus predistentus	Theocyrtis tuberosa	1	0.5		N C M		NANNOPLANKTON CHALK; light bluish gray (5B7/1) mixed with clasts of dark greenish gray (5GY4/1) and greenish gray (5G6/1) noncalcareous clay, from 0-57 cm; the compact nannoplankton chalk (57-150 cm) displays slight to intense (57-76 cm) very dusky purple mottling (5P2/2) containing sulfides. Radiolaria and sponge spicules occur but are not common. Sparse plagioclase. X-ray also shows K-feldspar.	0	50	100
				1.0	F A M							
							R A M					
							R A G					
							F A M					
				CORE CATCHER								



The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1428. This background was subtracted from the data.

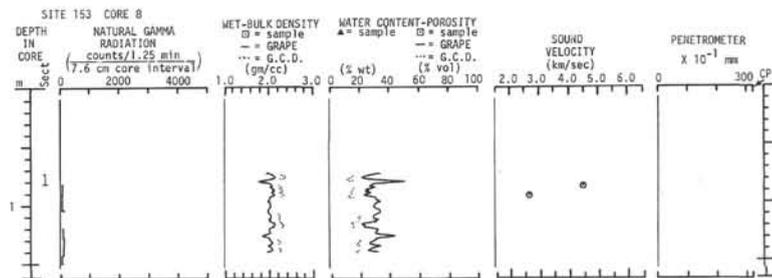
¹For explanation of symbols, see Chapter 1

SITE 153 HOLE CORE 8 CORED INTERVAL (m) 563-572

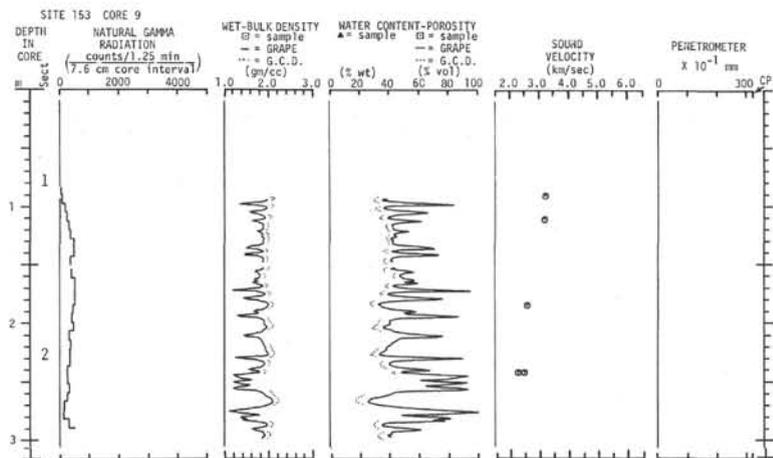
AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NANNO	RAD									
EARLY EOCENE	<i>Globobuccella formosa</i>	<i>Discoaster gemmeus</i>		1	0.5 1.0	VOID		N F P F A M	SILICEOUS LIMESTONE, light gray (N8) and laminated, interbedded with light olive gray (5Y6/1) CALCAREOUS CHERT.	0 50 100		
				CORE CATCHER				R O F A M	Foraminiferal test walls are preserved and chambers are empty or filled by matrix. Dolomite is observed. X-ray shows cristobalite.			
									Silicified limestone displays fossil walls all or partly silicified with outlines preserved and chambers filled with silica.			
									Silicified zones are concentrations of small silica lenses with both sharp and gradational contacts to the limestone matrix.			
									Staining indicates calcite content of core catcher material is irregular. Some chert occurs.			

SITE 153 HOLE CORE 9 CORED INTERVAL (m) 586-591

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NANNO	RAD									
PALEOCENE	?	<i>Fasciculithus tympaniformis</i>		1	0.5 1.0	VOID		F F P F C P	SILICIFIED MARL grading into LIMESTONE. Silicified marl (5GY6/1) has numerous lenticular laminations (1 mm thick) due to burrowing. Several zones of olive gray CHERT (5Y4/1) occur where silicification increases. Grades to very light gray (N8) limestone near 60 cm in Section 2. The limestone displays rare nannofossils and grains of micrite. A trace of dolomite occurs. Carbonate rhombs throughout. X-ray shows cristobalite, clinoptilolite, and barite.	0 50 100		
MIDDLE PALEOCENE	<i>Globobuccella pusilla pusilla</i>		unzoned	2	0.5 1.0			F F P F F M				
				CORE CATCHER								

¹For explanation of symbols, see Chapter 1

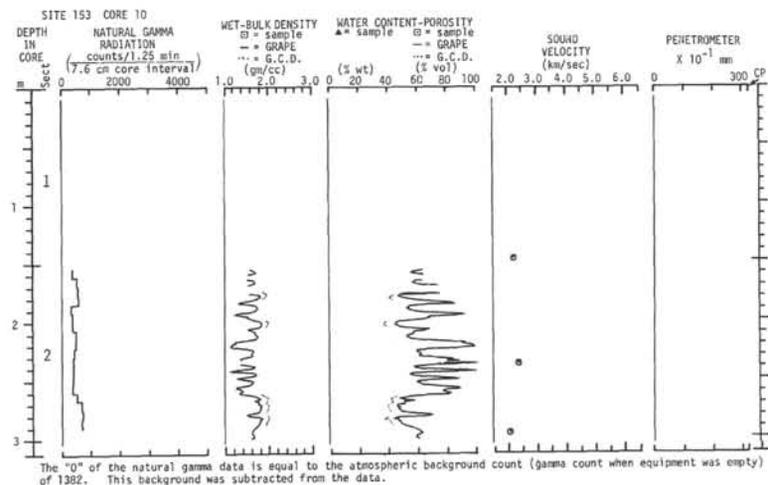
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The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1398. This background was subtracted from the data.

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	MAMMO	RAD									
MIDDLE PALEOCENE	Globobulimina pusilla pusilla (indirectly)	Fasciculithus tympaeniformis	unzoned	1	0.5 1.0	VOID			MICRITIC MARL; greenish gray (5GY6/1) to light olive (5Y6/1) with minor amounts of nannoplankton, radiolaria, and sponge spicules. Light colored glass, authigenic calcite, and plagioclase also occur. Marl is heavily bioturbated and has lenses (3 cm x 1 cm and smaller) displaying sharp and gradational contacts. Marl grades into underlying unit over ~5 cm at 60 cm in Section 2. Fractures abundant in lower part.	0	0	
				2	0.0 0.5 1.0		F F P F R P F R P N F P R F P	pyrite radiolarian molds	50	50	VI	
			CORE CATCHER					F F P R R M				

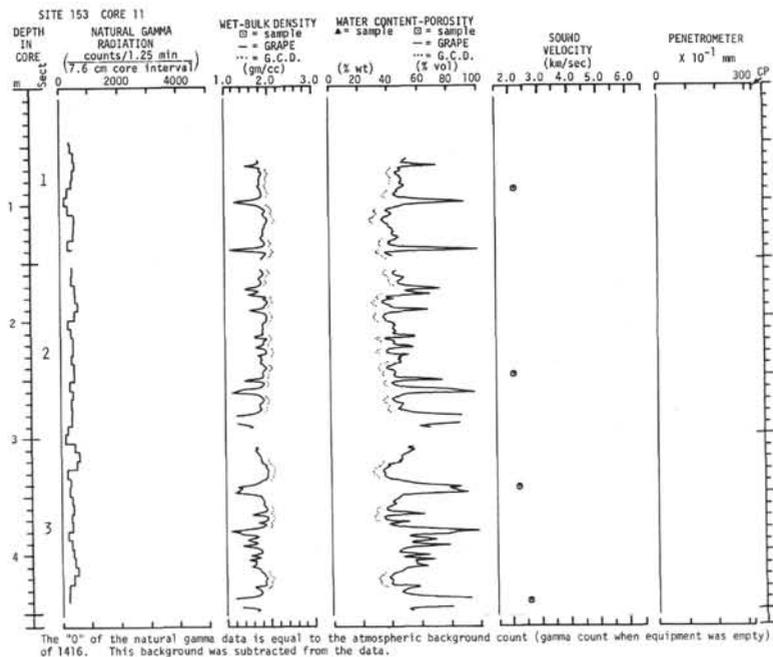
¹For explanation of symbols, see Chapter 1



SITE 153 HOLE CORE 11 CORED INTERVAL (m) 600-609

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION	
	FORAM	NANNO	RAD										
MIDDLE PALEOCENE	Globorotalia pusilla pusilla	Fasciculithus tymaniformis	unzoned	1	0.5	VOID			NANOPLANKTON MARL and ARGILLACEOUS LIMESTONE; light olive gray (5Y6/1) to light bluish gray (5B7/1) with darker zones of olive gray (5Y4/1). Laminae and lenses are common, parallel to bedding, and most have sharp contacts. Some plagioclase; sparse quartz and clinoptilolite. X-ray also shows cristobalite and barite. Light bluish gray zones seem associated with microfractures. Foraminifers are microspar and spar filled. Nannofossils appear to be partly recrystallized.	▲			
					1.0		R F M N F M						
					0.0		F C P						
					0.5		R F M						
					1.0		N F P F C P						
	2	unzoned	3	0.0		ZEOLITIC CALCAREOUS CLAY; olive gray (5Y4/1) to light olive gray (5Y6/1), occurs as zones in Section 3. Clinoptilolite is abundant and plagioclase minor. Fractures are green coated with black material.	▲						
				0.5		N F P F A M							
				1.0		F C P N T P							
						F C M							
						R F M							
			CORE CATCHER										

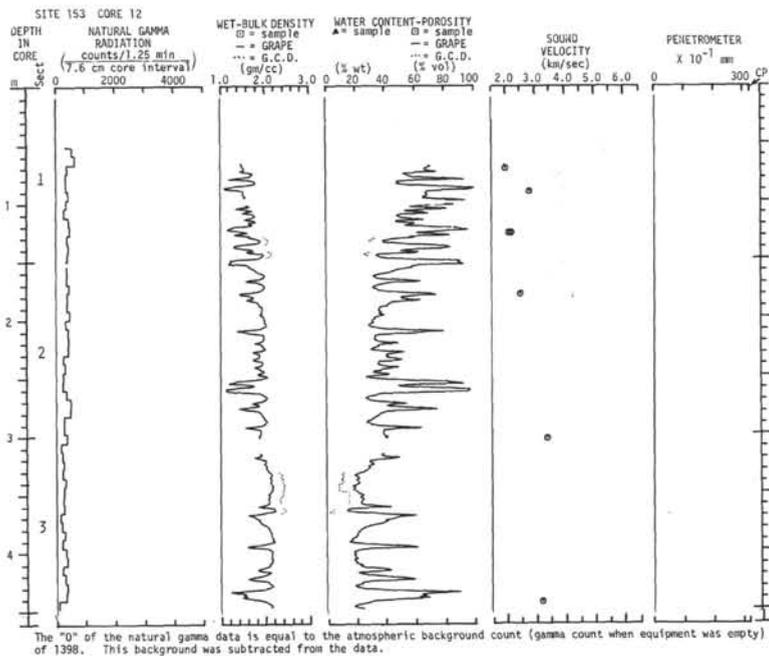
¹For explanation of symbols, see Chapter 1



AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%) SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NANNO	RAD								
2	7	5		1	0.5	VOID			Calcareous clay and nanno-plankton marl; olive gray (5Y4/1) to light olive gray (5Y6/1), intensively mottled and bioturbated. Sparse plagioclase, clinoptilolite, and chert micro-nodules. X-ray also shows cristobalite and barite. Burrows appear flattened. Clay-rich zones contain abundant clinoptilolite. Pale blue green (5BG7/2) zones around fractures. Sediment types alternate with no well-defined bedding.	0	
				2	1.0	Limestone breccia			Limestone breccia ("hard ground"); variable pale blue green (5BG7/2) to grayish blue green (5BG5/2) silicified marl fragments, micrites, and sparse biomicrites. Highly bioturbated; burrows contain packed foraminiferal biomicrite. The foraminifers are small and chalcedony filled. Also contains authigenic quartz and feldspar. The fractures contain breccia of the same composition cemented by drusy calcite and authigenic sand-sized feldspar. A chert pebble is found on top of the breccia.	50	
				3	1.0	Foraminiferal limestone			Foraminiferal limestone; light olive gray (N8) the color varying with the amount of calcium carbonate. Packed biomicrite, with foraminifers parallel oriented in some cases and unoriented in others. Rare radiolaria and benthonic foraminifers. Microspar and sparry filling, with a micritic and argillaceous matrix. Moderate to intense bioturbation. A cemented fracture occurs within the limestone (Section 3, 55-80 cm) and contains limestone fragments.	100	
				CORE CATCHER							

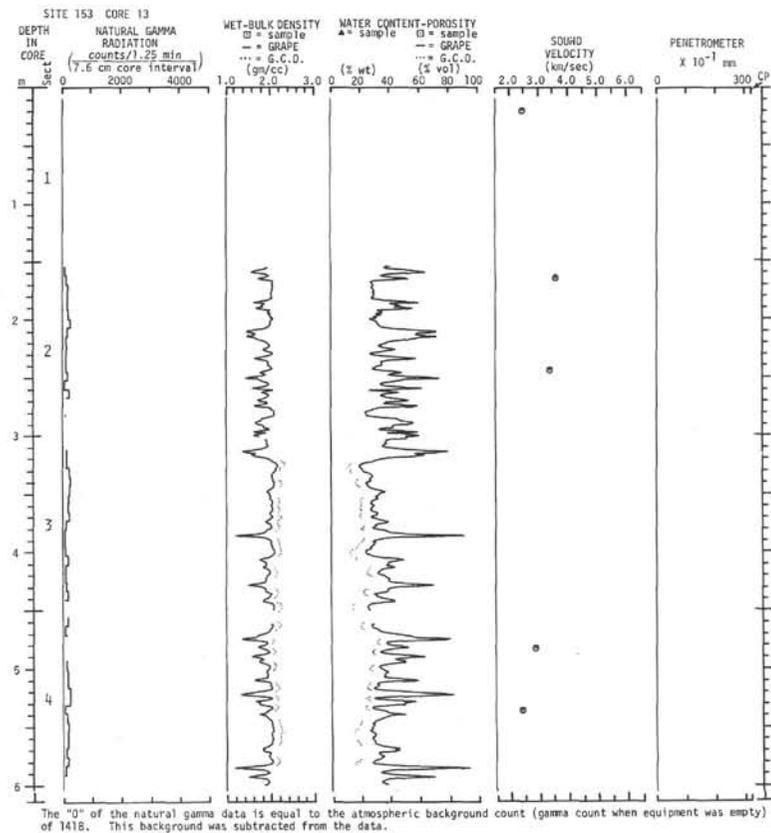
- 2. EARLY-MIDDLE PALEOCENE
- 3. EARLY PALEOCENE
- 4. Globigerina eugubina
- 5. Cruciplacolithus tenuis

¹For explanation of symbols, see Chapter 1



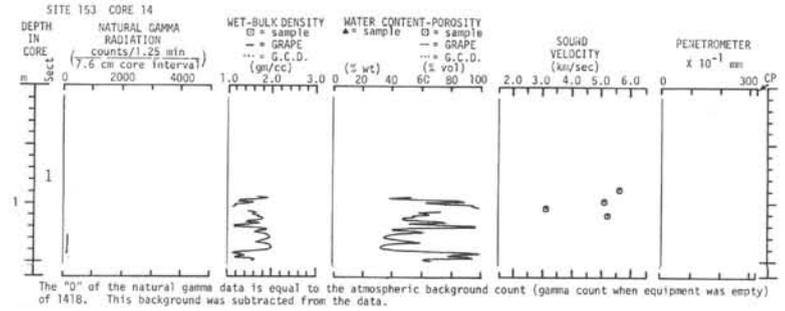
SITE 153 HOLE CORE 13 CORED INTERVAL (m) 619-656

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION	
	FORAM	NANO	RAD										
EARLY MAASTRICHTIAN	Globotruncana "tricarinata"	?		1	0.5	VOID							
				2	0.0	chert micro-nodules							
				3	0.0	chert micro-nodules	F A W						
				4	0.0	chert micro-nodules	F A W N F P						
				CORE CATCHER									

¹For explanation of symbols, see Chapter 1

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%) SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	MANNO	RAD								
CRETACEOUS	?	?		1	0.5 1.0	VOID		F N F P	LIMESTONE and CHERT fragments. Limestone is highly banded with argillaceous micrite. Fauna is oriented parallel to bedding and is filled by microspar and sparite. Argillaceous matrix bends around forams indicating compaction. Numerous laminae and lenses occur with both sharp and gradational contacts. Foraminifers, nannofossils, and radiolaria are rare. Some lenses are siliceous and display red molds and micronodules. Chert fragments are brownish gray (5YR6/1) with conchoidal fracture. Calcareous chert with irregular silicified areas of microquartz and ghosts of foraminifers and partially replaced sparite.	0 50 100	
				CORE CATCHER							

¹For explanation of symbols, see Chapter 1



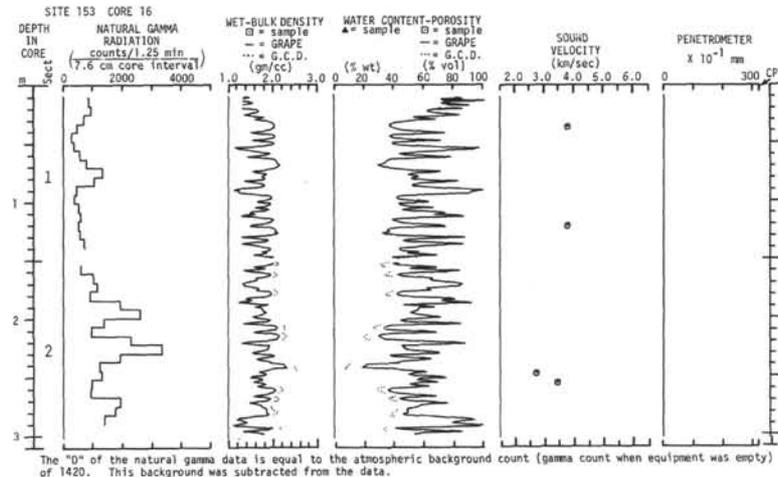
AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%) SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NANNO	RAD								
CONIACIAN	Globotruncana concavata	?		1	0.5	[Lithology diagram]			phosphorite	0	
					1.0				fish debris		
				2	0.0				organic matter	100	
					0.5				fish debris		
					1.0				glauconite		
									phosphorite		
									large pyrite nodules		
				CORE CATCHER							

FORAMINIFERAL LIMESTONE interbedded with phosphatic CAL-CAREOUS CLAYS and VOLCANIC IRON-RICH ORGANIC CLAY. Foraminiferal limestone is predominantly very light gray (N8) with variegated streaks and lenses of olive gray (5Y4/1) and greenish black (5GY2/1). Thin section shows a mottled planktonic foraminiferal biomicrite with argillaceous matrix. Radiolaria and benthic foraminifers are rare. Fossils are filled with sparite and fragments are often concentrated into lenses. Occasionally contains abundant fish debris, scattered plagioclase, barite and basaltic ash. X-ray shows clinoptilolite. Phosphatic calcareous clay is olive black (5Y2/1) and often grades into surrounding limestone. Nodules and fish debris are common. Other clays display altered basaltic ash, chlorite, glauconite. Chert fragments are black to medium dark gray (N4) often thin laminated argillaceous foraminiferal chert. Foraminifers are replaced by chalcedony. Silicified matrix. Nodules of phosphorite, conspicuous barite crystals and fish debris found at Section 2 (68-71 cm).

Analysis of the clay at 68-71 cm in Section 2:

SiO ₂	55.9%
Al ₂ O ₃	4.9%
TiO ₂	0.45%
Fe ₂ O ₃	2.5%
MgO	0.8%
CaO	8.6%
Na ₂ O	1.30%
K ₂ O	0.20%
MnO	0.04%
P ₂ O ₅	0.82%
Ba	1200 ppm
Sr	450 ppm
Cu	200 ppm

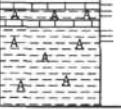
Laminated, bioturbated, RADIO-LARIAN FORAMINIFERAL LIMESTONE with barite, plagioclase, fish debris, glauconite and basaltic ash fragments.



The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1420. This background was subtracted from the data.

¹For explanation of symbols, see Chapter 1

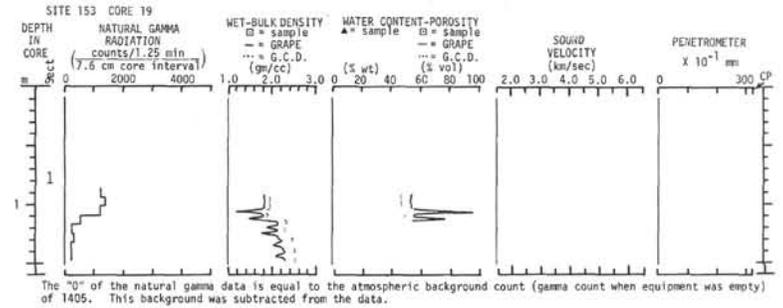
SITE 153 HOLE CORE 17 CORED INTERVAL (m) 740-749

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NAHNO	RAD									
CRETACEOUS				1	0.5	VOID			Interbedded MICRITIC LIMESTONE and SILTY VOLCANIC CLAY. Limestone is predominately light gray (N7-N8) and similar to Core 16 limestone. Volcanic clays are predominantly olive black (5Y2/1) and consist of altered basaltic ash with plagioclase, chert micronodules, chlorite, glauconite, clinoptilolite, abundant fish debris, and opal. Some alternation of coarse-fine lenses suggests reworking by currents.	0	50	100
					1.0		R F P					
				CORE CATCHER				F F P	Analysis of the clay at 121-127 cm in Section 1: SiO ₂ 48.7% Al ₂ O ₃ 12.6% TiO ₂ 0.85% Fe ₂ O ₃ 8.0% MgO 5.5% CaO 7.0% Na ₂ O 2.50% K ₂ O 0.70% MnO 0.04% P ₂ O ₅ 0.12% Ba 1900 ppm Sr 500 ppm Cu 110 ppm			

¹For explanation of symbols, Chapter 1

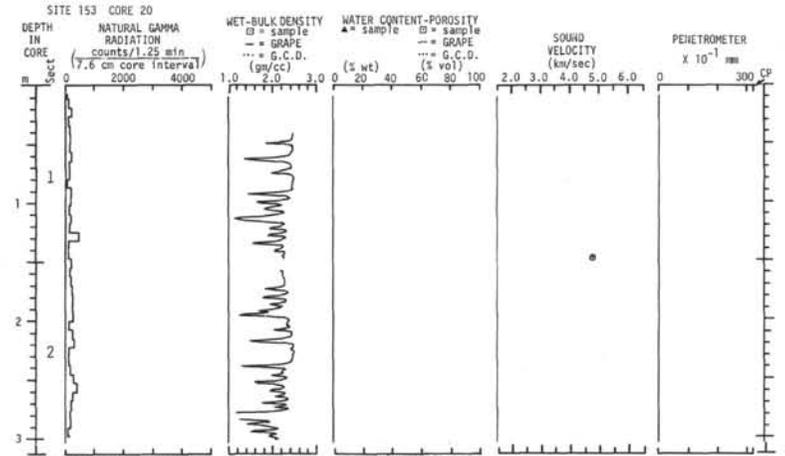
SITE 153 HOLE CORE 19 CORED INTERVAL (m) 758-767

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION	
	FORAM	NANNO	RAD										
LATE TURONIAN	Globotruncana			1	0.5 1.0	VOID			VOLCANIC SANDS and BASALT. Volcanic sands are medium light gray (N6) to dark gray (N3) and strongly bedded. Laminations are 5-50 mm thick and graded. Beds have sharp bottom contacts, grading upward from coarse-medium sand to silt, and burrowing near the upper contacts. Slides suggest an altered basaltic ash with abundant dark devitrified glass and clay. Plagioclase and chert micronodules are rare. Chlorite, glauconite, and radiolaria molds are present. Basalt is olive black (5Y3/1), fine grained, and sparsely amygdaloidal. Numerous fractures often filled with alteration material (chlorite). Mostly brown glass crowded with oxide and pyroxene microlite bundles. Some plagioclase microlites and pyroxene crystals occur. X-ray results at 101-107 cm: Diffuse scattered X-rays 62% Amorphous scattering 41% Calcite 58% Quartz 2% K-feldspar 6% Montmorillonite 33% Pyrite 2%	0	50	100	
	schneegansi?			CORE CATCHER			F A M N F M						

¹For explanation of symbols, see Chapter 1

AGE	ZONE			SECTION	METERS	LITHOLOGY	LITHO SAMPLE	PALEO SAMPLE ABUNDANCE PRESERVATION	LITHOLOGIC DESCRIPTION ¹	CaCO ₃ (%)	SAND-SILT-CLAY (accumulative %)	DEFORMATION
	FORAM	NANNO	RAD									
				1	0.5 1.0	VOID			BASALT: Same description as Core 19.	0 50 100		
				2	0.0 0.5 1.0							
				CORE CATCHER								

¹For explanation of symbols, see Chapter 1



The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1366. This background was subtracted from the data.

