## The Shipboard Scientific Party<sup>1</sup>

### SYNOPSIS

Area: Carnegie Ridge, eastern equatorial Pacific

Date Occupied: 12-14 February, 1971

Position:

Lat. 01°45.70'S Long. 85°54.17'W

Water Depth: 1591 meters (corrected)

Penetration:

157: 437 meters 157A: 27 meters

Number of Holes: 2

Number of Cores: 157: 49 157A: 3

Core Recovery: 157: 273.6 meters

157A: 19.3 meters

#### Acoustic Basement:

Depth: 0.45 second

Nature: Chert

Inferred acoustic velocity for sedimentary section: 1530 meters

Age of Oldest Sediment: Late Miocene

#### Basement: Basalt

Hole DSDP 157 was continuously cored from the surface to 3 meters into basalt; DSDP 157A recored the upper 27 meters. DSDP 157 is lithologically uniform and all changes are gradual.

0-345 meters – Siliceous-calcareous ooze to approximately 240 meters, gradually increasing in lithification to semiindurated chalk; chalk below 284 meters. Calcareous nannofossils abundant but poorly preserved; radiolarians abundant, poorly preserved below 324 meters. Age: Quaternary to late Miocene.

345-431 meters – Interbedded cherts and chalk; chalk similar to overlying unit except for absence of siliceous microfossils.

Age: Late Miocene.

431-437 meters – Basalt, flow structures near upper part, rest massive and holocrystalline, moderately altered, extrusive.

The sedimentation rate is about 47 m/m.y. from the surface to 75 meters, changing to about 63 m/m.y. between 75 and 326 meters. Between 325 and 330 meters, the rate is sharply reduced or a hiatus of approximately 1 m.y. is present. Between 330 meters and basement, the rate is at least 29 m/m.y.

# **REGIONAL SETTING AND OBJECTIVES**

Site DSDP 157 is located at the southern edge of the Panama Basin on the south flank of the Carnegie Ridge (Figure 1). The Panama Basin is bordered by the continental margins of eastern Central America and northern South America and by two isostatically compensated ridges, the Carnegie and Cocos ridges, which enclose a central basin traversed by several north-south trending fracture zones. As shown in Chapter 2 (this volume) and in van Andel et al. (1971), these ridges and several high blocks within the eastern part of the Panama Basin exhibit considerable similarity in the nature of the acoustic basement and the sediment cover. On these grounds and with the aid of a detailed tectonic analysis of the basin, van Andel et al. (1971) have suggested that all high blocks of the basin originally formed part of a single, east-west trending ancestral Carnegie Ridge located approximately at the position of the present one. This ridge was split by the formation of the Galapagos Rift Zone. Opening of the rift began in the east and proceeded westward in a stepwise fashion, affecting the block west of the Coiba Fracture Zone approximately 10 m.y. ago, and subsequent more westerly blocks at later dates. While the southern half of the split ridge remained stationary, successive northern blocks migrated northward until, one by one, they reached and sealed the eastern extension of the Middle Americas Trench.

Three sites in the Panama Basin (one each on the Coiba, Carnegie, and Cocos ridges) were selected by the Pacific Site Selection Panel to test this hypothesis through comparison of sedimentary sections, basement ages, and depositional histories of the three ridges. In addition, the sites were chosen to determine the late Cenozoic biostratigraphy of the eastern equatorial Pacific in locations shallow enough to insure preservation of calcareous microfossils. Additionally, DSDP 157 was selected to examine temporal changes in the location and intensity of the Peru Current as reflected in the biogenous deposits at the site, especially those of Quaternary age. The site is located in the area of interaction between the Peru Current and the equatorial current system.

## TOPOGRAPHIC AND GEOLOGIC SETTING

DSDP 157 is located on the south flank of the Carnegie Ridge slightly less than halfway between the crest of this ridge and the deep Pacific basin and well east of the

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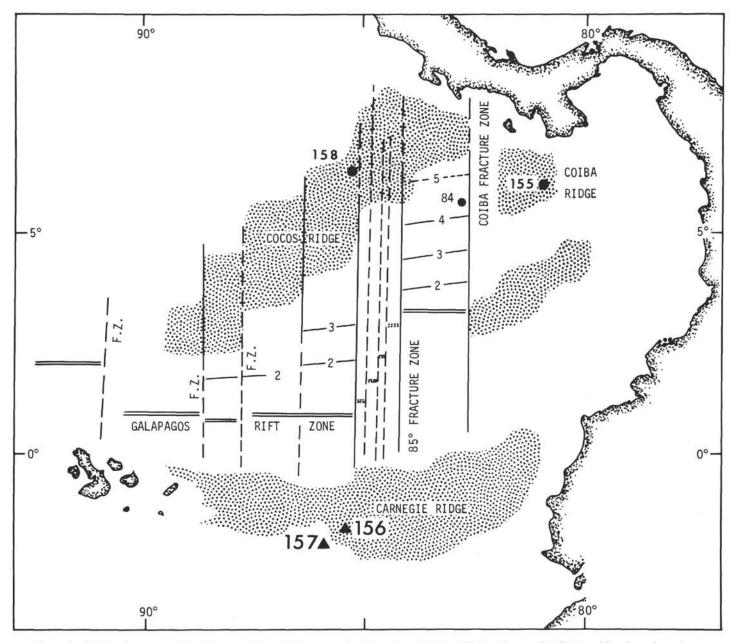


Figure 1. Tectonic map of the Panama Basin (after van Andel et al., 1971) with locations of drill sites. Numbered contours are magnetic anomalies.

pedestal of the Galapagos Islands (Figure 2). The Carnegie Ridge has been uplifted along a series of east-west trending normal faults resulting in the formation of several structural terraces on each flank (Figure 3). The terrace surfaces are level and covered with a thick sedimentary blanket. On the crestal side, each is bordered by a steep scarp leading to the next step and on the downward side by a protruding basement ridge or a series of basement pinnacles (Figure 3). The terraces are dissected by numerous gullies or canyons of erosional origin that extend to the acoustic basement. Few of these gullies appear on north-south traverses, but they are numerous on east-west traverses (see Chapter 3, Figure 3, this volume) indicating that they have a general north-south trend down the ridge slope. The relatively flat crest of the Carnegie Ridge, approximately 70 km north of the drill site, was eroded to acoustic basement at an unknown time after most of the sedimentary cover of the ridge had been deposited (van Andel et al. 1971). Regional sediment isopach maps (van Andel et al., 1971) show that a band of anomalously thick sediment, which presumably contains part or all of the eroded material, occurs north of the ridge crest. No such overthickened zone can be distinguished south of the crest, but the numerous canyons suggest that at least part of the material is transported southward into the deep Pacific basin to the south.

The acoustic section (Figure 3) contains numerous closely spaced and continuous reflectors. The character of these reflectors changes little with depth and is strongly reminiscent of the calcareous deposits of the equatorial zone of high biological productivity farther to the west

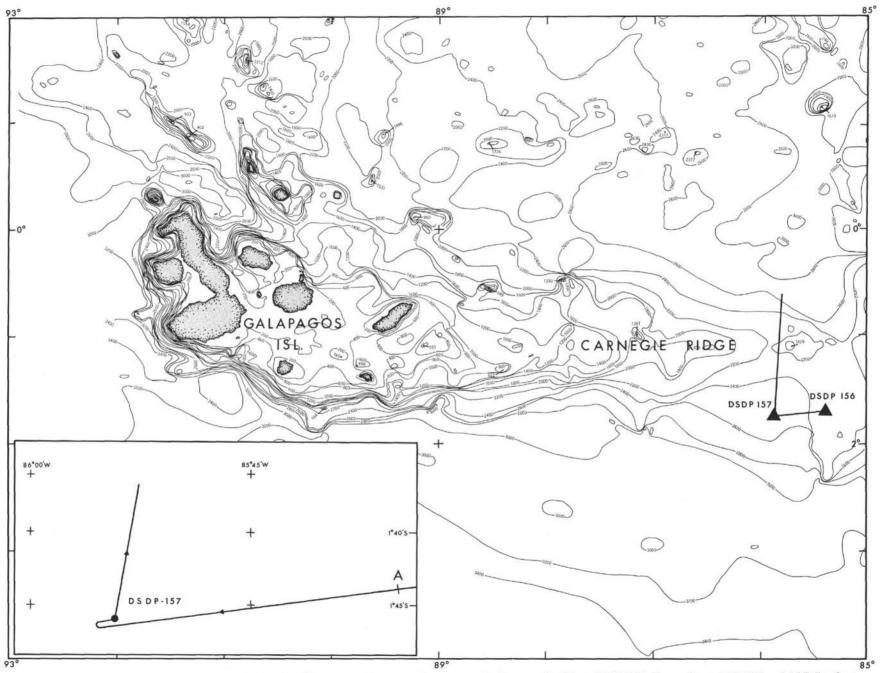


Figure 2. Location of DSDP 157 on Carnegie Ridge. Depth controus in meters (uncorrected after van Andel et al. (1971). Heavy lines 156-157 and 157-C refer to acoustic profiles of Figure 3. Insert: site approach track; A-157 is location of profile of Figure 3.

SITE 157

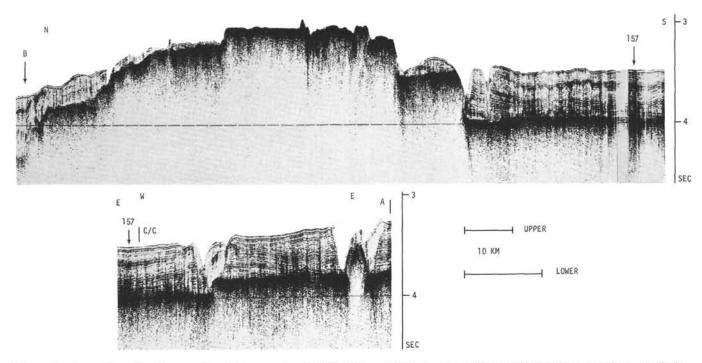


Figure 3. Acoustic reflection profiles between sites DSDP 156 and 157 (top) and from DSDP 157 across Carnegie Ridge (bottom). Locations on Figure 2. Depths in seconds 2-way travel time (1 sec = approx. 800 m of sediment). Horizontal scale approximate.

(Ewing et al., 1968). In drill holes on DSDP Legs 7, 8, and 9, the layering has been correlated with variations either of the carbonate content or of the degree of induration of the deposits (Winterer et al., 1971; Tracey et al., 1971; Hays et al., 1972).

Low-angle unconformities of limited horizontal extent are common in the sedimentary cover of the Carnegie Ridge. They are evident on north-south traverses (Figure 3) and are probably related to the regime of crestal erosion and flank redeposition described above. They are most common at a depth of approximately 110 to 125 meters below the sea floor and are generally restricted to the upper part of the section.

The acoustic basement in the site area is very smooth and shows strong internal reverberation. Such a smooth basement is widespread on the ridges of the Panama Basin area, but is absent from the deeper, younger crust. It generally terminates sharply against basement scarps and pinnacles which exhibit a quite different acoustic signature. The sedimentary origin for this basal reflector proposed by van Andel et al. (1971) was confirmed by the discovery of chert at this site and at DSDP 158. The diagenetic character of the reflector is responsible for the generally unconformable relation between the acoustic basement and the overlying beds, (Figure 3). The acoustic basement is affected by the numerous normal faults of the ridge flanks, both by the major ones that determine the ridge shape and by numerous small ones with random displacements. The acoustic section at the site is quite typical for the entire Carnegie Ridge, and shows little lateral variation except for minor thickening and thinning. In Table 1, the acoustic profile is compared with the results of drilling. The data support the correlation of acoustic basement with the top of the chert zone.

In view of the regional uniformity of the topography of the terrace and of the sedimentary section on this part of the Carnegie Ridge, no extensive site survey was carried out. The site was originally selected from seismic reflection profile record 1633 of R/V *Conrad*, Cruise 11, and lies on a flank terrace approximately 20 km wide (north-south). The final site was established after completing a short double pass in an east-west direction to insure that no local anomalies of small size were present. After completion of drilling, a profile across the beacon confirmed that the drilled site is typical for the area.

#### **OPERATIONS**

The beacon was dropped at 1015 hours on February 12 and the hole spudded using a Smith 10 1/8", 3-cone shaped insert bit which performed very well and was still in usable condition after completion of the hole. The hole was cored continuously, using a regular core barrel for Cores 1 through 7, and 13 through 49. Cores 8 through 12 were taken with an extended core barrel. A summary of coring at DSDP 157 is given in Table 2. Core recovery in the chert was not good, probably because of alternating hard and soft

**SITE 157** 

 TABLE 1

 Comparison of Acoustic Section and Drill Data, DSDP 157

Reflectors	Depth <sup>a</sup> (sec)	Drilling Results <sup>a</sup>	Interval (sec)	Velocity <sup>a</sup> (m/sec)	Calculated <sup>b</sup> Depth (m)
Top upper stratified zone	0.15				116
Top lower stratified	0.32 =	(345 m top chert)	0-0.32	(2160)	
zone	0.32 =	240 m top chalk	0-0.32	1500	248
	0.32 =	(265 m reduced drilling rate)	0-0.32	(1660)	
Top acoustic	0.45 =	345 m top chert	0-0.45	1530	345
basement	0.45 =	(431 m top basalt)	0-0.45	(1930)	

 $^{a}$  Underlined – accepted correlation; parenthesis (  $\,$  ) – correlation not accepted.

<sup>b</sup>Calculated assuming velocity of 1530 m/sec.

layers. A single sidewall core was attempted but the barrel sheared off, probably because the formation was too hard.

Upon completion of the first hole, DSDP 157A was drilled using the extended core barrel in the hope of reducing core disturbance. The first core showed significantly reduced disturbance, but the others were of the same quality as those of DSDP 157 and the attempt was abandoned after three cores.

## LITHOLOGY

Lithologically, the sediments of DSDP 157 and DSDP 157A are uniform. From the surface to the basement contact at 431 meters, they consist of nannofossil chalk ooze and chalk which gradually become more indurated with depth. The upper part of the section, above 342 meters, is rich in biogenous silica in the form of radiolarians and diatoms; below 324 meters, the state of preservation of the siliceous organisms deteriorates and below 342 meters all silica is in the form of chert stringers and nodules. At 431 meters, the sediment rest on basalt.

The chalk oozes and chalks vary in color. Alternating zones are greenish gray, pale olive, light to moderate olive brown, and pale greenish yellow. Mottling and burrowing occur but not excessively; mottles and diffuse laminations are commonly marked by concentrations of fine-grained pyrite and pyrite nodules. Gray, pyrite-rich intervals are common from 250 to 342 meters; above 250 meters, much pyrite occurs in large burrows up to 1 cm wide and 10 cm long.  $H_2S$  odor was noticeable throughout.

Near the surface, foraminifers are abundant. They decrease below 50 meters, but remain common to abundant to 250 meters, and common to 342 meters. The dominant components are nannofossils, followed in abundance by diatoms, which decrease markedly in abundance below 333 meters, perhaps as a result of solution. Radiolarians are common throughout the section to 324 meters, below 345 meters they disappear. Subsidiary components are silico-flagellates, pyrite nodules, volcanic glass, glauconite, and

14	ABLE 7	2		
Coring Summary,	DSDP	157	and	157A

	Depth Below Sea Level	Depth Below Sea Floor	Cored	Reco	vered
Core	(m)	(m)	(cm)	(cm)	(%)
1	2601-2610	10-19	900	850	94.4
2	2610-2619	19-28	900	900	100.0
3	2619-2628	28-37	900	893	99.2
2 3 4 5	2628-2637	37-46	900	900	100.0
5	2637-2646	46-55	900	900	100.0
6	2646-2655	55-64	900	900	100.0
7	2655-2664	64-73	900	828	92.0
8 9	2664-2673	73-82 82-91	900 900	593 822	65.9 91.3
10	2673-2682 2682-2690	91-98	800	600	75.0
11	2690-2699	99-108	900	534	59.3
12	2699-2708	108-117	900	900	100.0
13	2708-2717	117-126	900	555	61.7
14	2717-2726	126-135	900	830	92.2
15	2726-2735	135-144	900	890	98.9
16	2735-2744	144-153	900	862	95.8
17	2744-2753	153-162	900	900	100.0
18	2753-2762	162-171 171-180	900 900	877 900	97.4
19 20	2762-2771 2771-2780	180-189	900	170	18.9
21	2780-2789	189-198	900	890	98.9
22	2789-2798	198-207	900	895	99.4
23	2798-2807	207-216	900	900	100.0
24	2807-2816	216-225	900	480	53.3
25	2816-2825	225-234	900	750	83.3
26	2825-2834	234-243	900	895	88.3
27	2834-2843	243-252	900 900	896 860	99.6 95.6
28 29	2843-2852 2852-2861	252-261 261-270	900	356	39.6
30	2861-2870	270-279	900	900	100.0
31	2870-2879	279-288	900	885	98.3
32	2879-2888	288-297	900	900	100.0
33	2888-2897	297-306	900	900	100.0
34 35	2897-2906 3906-2915	306-315 315-324	900 900	0 150	0.0
36	2915-2924	324-333	900	362	40.2
37	2913-2924	333-342	900	18	2.0
38	2933-2936	342-345	300	73	24.3
39	2936-2941	345-350	500	100	20.0
40	2941-2950	350-359	900	0	0.0
41	2950-2959	359-368	900	51	5.7
42	2959-2968	368-377	900	150	16.7
43	2968-2977	377-386	900	45	5.0
44 45	2977-2986 2986-2995	386-395 395-404	900 900	0	0.0
	2986-2993	404-413	900	38	4.2
46 47	3004-3013	404-413	900	40	4.4
47	3013-3022	413-422	900	34	3.8
49	3022-3028	431-437	600	188	31.3
1A	2591-2600	0-9	900	900	100.0
2A	2600-2609	9-18	900	130	14.4
3A	2609-2618	18-27	900	900	100.0

sponge spicules, all of which are somewhat more common in the upper part of the section. The relatively rare occurrence of volcanic glass is noteworthy at this site, which is quite close to (albeit upwind from) the Galapagos volcanic complex. Severe disturbance of the sediments by the coring process partly masks the increase in consolidation with depth. However, steadily increasing induration is clearly indicated by reduced coring rates. From approximately 2.1 m/minute at the surface, the rate decreased to 1.6 m/minute at 110 meters. Between 110 and 260 meters, the rate remained approximately constant. From there to 342 meters, the rate again decreased from approximately 1 m/minute to 0.5 m/minute. Below 285 meters, the formation is arbitrarily logged as chalk rather than as chalk ooze, although the transition is very gradual.

At 342 meters, a sharp decrease in the drilling rate marked the highest occurrence of chert and below this depth the formation consists of chert alternating with nannofossil chalk and chalk limestone. The highest chalk limestone appears near 350 meters. The chalk and chalk limestone are similar to the deposits higher in the section except for the absence of unaltered diatoms, radiolarians, or silicoflagellates. Thin sections contain very rare chalcedony and pyrite pseudomorphs after radiolarians. The beds show the same color range as the upper, non-chert-bearing section (yellowish gray, very light gray, pale olive and light olive gray, with yellowish gray predominating). Mottles and burrows are common; the burrows are generally compressed perpendicular to bedding. Fine-grained pyrite is common as darker gray laminae and streaks and as grayish purple to black concentrations.

The chert is commonly calcareous, particularly near the top of the chert section. Mottled bedding and bioturbation are well preserved. The chert is generally darker than the associated chalk and chalk limestone. Moderate to light olive brown and olive black colors predominate, but lighter colors such as pale greenish yellow, pale olive, yellowish gray, light olive gray, and olive gray are also present. The chert occurs as nodules, stringers and irregular crosscuting masses up to 10 mm thick. In thin section, it is seen to have replaced the limestone; microfossil tests are marked by coarser-grained aggregates of chalcedonic quartz. In limestone adjacent to a chert body, foraminiferal chambers are often filled with chalcedony.

Basalt was reached at 431 meters, based on a change in coring rate from 0.45 m/minute to 0.1 m/minute. Between 431 and 437 meters, 1.9 meters of basalt core was recovered. The top 20 cm consists of silicified crust-like masses of medium to light gray basalt with smeared vesicles and amygdules and associated gray fine-grained material which may be altered glass. Below, the rock is medium to dark gray massive basalt, very fine-grained with 0.2 to 1.0 mm labradorite laths and interstitial augite and 1 to 5 mm phenocryst clusters of oscillatory-zoned plagioclase showing synneusis. The groundmass contains plagioclase microlites, magnetite, chlorophaeite, and devitrified glass. The basalt is cut by thin, irregularly trending, mostly steeply dipping veins of calcite and serpentine with concentrations of pyrite and possibly chalcopyrite and pyrrhotite. The basaltsediment contact is depositional, based on the absence of evidence of metamorphism in overlying sediments and the presence at the top of the basalt of altered, crust-like masses and smeared amygdules which may be part of a basalt pillow.

## GEOCHEMISTRY

Interstitial water samples and shipboard observations for DSDP 157 are listed in Table 3.

 TABLE 3

 Interstitial Water Samples and Shipboard Observations, DSDP 157

Core	Section	Sampled Interval (cm)	pН	Eh (mv)	Lab. Temp. (°C)	Salinity (%)	Squeeze Pressure (psi)
1	2	0-10	7.50	-36	24.0	34.7	1240
1	2	0-10	7.73	-16	24.0	35.2	2436
2 3 4	4	0-10	7.46	-180	23.9	35.2	2436
3	1	0-10	7.38	-189	24.0	35.2	1523
4	2	140-150	7.38	-272	23.9	35.8	1015
5	3	0-9	7.31	-270	24.1	35.2	1523
6	1	142-150	6.40	-208	24.2	35.2	1523
7	4	0-9	7.36	-262	24.1	35.2	1015
8	1	0-3	7.36	-224	24.0	35.2	1015
9	6	0-8	7.43	-305	24.0	35.2	1523
10	1	0-9	7.34	-348	24.7	35.2	1015
12	5 3	0-10	7.25	-285	24.1	35.2	1015
14	3	0-8	7.37	-311	25.5	35.2	1015
15	5	0-9	7.31	-323	25.2	35.2	1523
16	6	0-7	7.23	-194	25.3	35.2	1015
17	6	0-6	7.27	-155	25.5	35.2	1015
18	4	0-6	7.20	-258	25.6	34.7	1015
19	6	0-8	7.24	-150	25.8	34.7	1015
21	5	0-5	7.13	-335	25.8	34.7	188
23	5	0-5	7.18	-301	25.0	35.2	-
27	3	0-6	7.22	-186	25.3	34.7	2436
28	5	0-6	7.19	-98	25.4	34.7	2436
29	2	0-7	7.48	-130	25.5	34.7	1015
30	6	0-6	7.27	8	25.1	34.7	1523
31	6	0-5	7.20	-124	25.3	34.7	1523
32	6	0-6	7.04	-240	25.4	34.7	2030
33	6	0-8	7.26	-88	25.0	34.7	1015
35	1	135-140	7.33	27	25.0	34.1	2436
36	3	0-6	7.11	-229	25.4	34.7	2436
39	1	145-150	7.08	-27	25.5	34.1	3000
1A	6	0-8	7.36	-175	24.5	34.7	1015
2A	1	145-150	7.38	-218	24.5	34.7	1523
3A	5	0-9	7.43	-256	24.5	35.2	1015

## BIOSTRATIGRAPHY

At Site DSDP 157 a thick, apparently complete upper Miocene to upper Pleistocene sequence was cored. Except for the cherty interval from 345 meters to 437 meters TD in which radiolarians are not preserved, foraminifera, coccoliths, and radiolarians occur jointly throughout. Foraminifera are moderately well to well-preserved in all samples examined, but increasing induration of the sediment from Core 39 downward made the obtaining of representative faunas difficult or impossible. Coccoliths are abundant in all 48 sediment cores, but preservation is generally poor and deteriorates downward. Upper Miocene discoasters are characterized by heavy secondary calcite overgrowth. For both calcareous fossil groups-foraminifera and coccoliths-species diversity is low, and many of the common guide fossils for the upper Miocene and Pliocene are scarce or absent.

Coccoliths indicate that the sediment immediately above basalt is late Miocene in age (Discoaster neohamatus Zone). Foraminifera indicate a level within the Globorotalia plesiotumida Zone; however, compared to DSDP 158 this level should be already in the underlying Globorotalia acostaensis Zone. The discrepancey at DSDP 157 (because of a slightly more diverse foraminiferal fauna of its more tropical character, DSDP 158 is tentatively regarded as biostratigraphically more reliable). can be attributed to downhole displacement of G. plesiotumida, which would lower the zonal boundary; but because of the spotty occurrence of G. plesiotumida in the Panama Basin sequences, it is not unlikely that the lower occurrences of this species at DSDP 157 may be autochthonous and may have been missed at DSDP 158 because this species occurs rather sporadically.

Radiolarians are diverse and well-preserved down to 324 meters, but their preservation deteriorates downward to 345 meters below which depth they are absent.

In addition to the above fossil groups, silicoflagellates are conspicuous in nannofossil assemblages of Cores 1 through 38. They are discussed in detail by Bukry and Foster (Chapter 29, this volume).

Foraminifera, coccoliths, and radiolarians indicate that, except for the uppermost Pleistocene, the section at DSDP 157 is complete. Both coccoliths and radiolarians indicate that the uppermost Pleistocene is missing, even in Hole 157A, which cored the uppermost part of the section missed in Hole 157. Lack of resolution in this interval did not allow this assessment to be made on the basis of foraminiferal biostratigraphy.

### PHYSICAL PROPERTIES

Drilling in the upper 200 meters of DSDP 157 resulted in a relatively high degree of sediment core disturbance, rendering the GRAPE bulk density and porosity unreliable for this part of the section. Likewise, natural gamma measurements are reported only for the cores with the least apparent disturbance. During analysis of Core 25 and all subsequent cores, the GRAPE system developed technical problems reflected in unstable outputs from the calibration standards (water and aluminum). Bulk density and porosity data have been adjusted accordingly. However, GRAPE and natural gamma data should be regarded with some suspicion.

A relatively small number of sediment samples was collected for laboratory analysis. Consequently, the density of coverage is low, especially in the upper 230 meters of sediment. Likewise, below 340 meters, sediment disturbance and poor core recovery prohibited sampling, which was restricted to areas of least disturbance.

### **Bulk Density**

On the basis of data generated by the GRAPE system, large variations in the bulk density occur throughout the lithologic section. Values (average per section) range from 1.42 g/cc in the high-porosity upper Pliocene nannofossil chalk ooze to 1.94 g/cc in the low-porosity upper Miocene interbedded chert and chalk limestone which occurs just above the basalt. Laboratory bulk densities range from 1.28 g/cc at 61.5 meters in the nannofossil chalk ooze to 1.67 g/cc at 288 meters and 324 meters in the nannofossil chalk. A value of 1.67 g/cc also occurs in the chalk ooze at 258 meters. A low value of 1.34 g/cc at a depth of 251 meters in the nannofossil chalk ooze is due to the anomalous low grain density and high water content at this depth.

GRAPE and laboratory bulk densities generally decrease with depth, although variations occur throughout the sedimentary sequence. The lowest values occur predominantly in the chalk ooze. The highest bulk densities are common in the chalk limestone below a depth of 342 meters and increase sharply to a depth of about 370 meters where measurements terminated.

#### Porosity

GRAPE porosity (average value per section) ranges from 74 per cent at 200 to 220 meters in the nannofossil chalk ooze to 47 per cent at 370 meters in the chalk limestone. A pronounced overall decrease in porosity occurs with depth. Porosities in the chalk ooze range from 54 per cent to 74 per cent and decrease sharply with depth. Although porosities in the chalk range from 54 to 70 per cent, the general decrease with depth is not as clear as in the chalk ooze, and numerous variations are observed. The lowest values occur in the nannofossil chalk limestone where they range from 47 to 57 per cent and decrease sharply with depth within the sedimentary sequence.

A high of 84 per cent porosity at a depth of 61.5 meters contrasts with a low of 59 per cent at depths of 259 meters and 289 meters for laboratory measurements. The highest laboratory porosities with greatest relative variation occur in the chalk ooze. More consistent, lower values are common in the nannofossil chalk. Within the intervals sampled for laboratory measurements, porosities decrease slightly with depth. Unfortunately, the highest sampling densities occur in the interval where GRAPE porosities indicate the least decrease with depth.

## Water Content

Water content (per cent total weight) ranges from 64 per cent in the chalk ooze to 35 per cent in the nannofossil chalk. A value of 25 per cent is found in the lower chalk ooze but higher values are more common. A relatively high water content at 251 meters corresponds to the sample of low bulk density and low grain density. Low and fairly uniform water content in the nannofossil chalk contrasts with the overlying chalk ooze.

## Grain Density

Average grain densities range from 2.38 g/cc to 2.69 g/cc and vary less in the nannofossil chalk than in the overlying chalk ooze. A sharp decrease in grain density, from 2.74 g/cc to 2.40 g/cc, is observed in the upper 230 meters of the chalk ooze. A value of 2.38 g/cc, the lowest density observed, is found at 251.5 meters. Below this depth, the average grain density increases in the lower chalk ooze. The basal nannofossil chalk has grain densities of 2.62 g/cc to 2.66 g/cc.

## Sonic Velocity

Sonic velocities in the upper 330 meters were measured in split core sections confined to the core liner. Samples from greater depths were removed from the core liner and in several cases velocities were measured parallel as well as normal to bedding. In some instances, such as in the interval 345-385 meters, velocities were higher parallel to bedding. Velocities measured normal to bedding were generally higher than parallel measurements in the interval 400 to 430 meters. Differences in sonic velocity perpendicular and parallel to bedding ranged from as little as 0.01 km/sec to as much as 0.53 km/sec. The larger differences are probably due less to the anisotropic characteristics of the sediment than to inconsistencies in technique during testing. Extremely low values are almost certainly due to technique and are not representative of the overall sonic velocity characteristics of the sediment.

Sonic velocities are relatively constant in the upper 330 meters averaging 1.53 km/sec with extremes of 1.34 km/sec and 1.84 km/sec in the depth interval of 215 to 243 meters. A significant increase in velocity (to an average of 2.90 km/sec) is encountered in the interbedded chert and nannofossil chalk limestone between 345 meters and the basalt at a depth of 432 meters. Velocities of 1.85 km/sec to 2.44 km/sec occur at depths of approximately 359 meters, 422 meters, and 431 meters in chalk limestone with essentially no interbedded chert. Five measurements of sonic velocity on one sample of basalt recovered from DSDP 157 averaged 4.54 km/sec with a range of 4.46 km/sec to 4.62 km/sec.

# Natural Gamma Radiation

Relatively few core sections were suitable for natural gamma measurements. Since most of the data are limited to a small depth interval, no trends can be recognized in the sediment except that gamma activity is higher in the lower chalk ooze (260-270 m) than in the underlying nannofossil chalk. Within the short interval tested (260-340 m), the gamma activity is quite variable.

## Shear Strength

A few vane shear measurements were possible in selected intervals of DSDP 157 and 157A. The lowest values of 9.8 and 31.6 g/cm<sup>2</sup> are characteristic of the foraminiferal nannofossil chalk ooze in Cores 1 and 2 of DSDP 157A. High water contents and porosities and low bulk densities are typical of this sediment. Shear strenghts of 215.1 to 620.1 g/cm<sup>2</sup> occur in the lower chalk ooze from 225 meters to 265 meters. This deeper chalk ooze has lower water contents and porosities and higher bulk densities than the overlying sediment. Although a value of 620.1 g/cm<sup>2</sup> occurs at 252 meters, values of 200 to 400 g/cm<sup>2</sup> are most common. Shear strengths of 150.5 to 1474.4 g/cm<sup>2</sup> occur in the nannofossil chalk from the interval 288 to 333 meters, with the highest values most common in the deepest cores. Low values of 150.5 to 243.3 g/cm<sup>2</sup> characterize the upper chalk. Vane shear testing in the chert and chalk limestone was not possible for obvious reasons.

## DISCUSSION AND SUMMARY

The most striking feature of the section sampled by DSDP 157 is its great uniformity, disturbed only by

diagenetic changes. This must indicate very stable depositional conditions in the area, thereby strongly supporting a stable latitude for the Carnegie Ridge.

The sediment is siliceous-calcareous ooze to about 240 meters, then changes gradually over the interval 240 to 284 meters to a semi-indurated chalk. The transitional zone, probably consisting of alternating chalk and ooze beds but obscured by coring disturbance, corresponds to a prominent acoustic reflector. The sediment is dominated by nannofossils with abundant foraminifers in the Pleistocene deposits and diatoms abundant throughout. The diversity of the calcareous microfossils is low and the nannofossils are generally poorly preserved. Radiolaria are ubiquitous and of great diversity but become poorly preserved below 324 meters, and disappear altogether with all other siliceous fossils below 345 meters, and disappear altogether with all other siliceous fossils below 345 meters. All freshly opened cores gave off H<sub>2</sub>S, and pyrite is a ubiquitous component of the sediments. The only other nonbiogenous components are glauconite (generally rare and filling foraminiferal chambers) and rare, mostly light-colored, volcanic glass. The minor lithologic variation between 0 and 345 meters results from subtle changes in the dominant colors from greenish grays to olives to olive browns.

The paleontologic data suggest a sedimentation rate of  $60 \text{ m}/10^6$  yrs from the surface to about 300 meters. Below 330 meters, the rate is somewhat uncertain but may exceed 29 m/10<sup>6</sup> yrs. Between 300 and 330 meters, an interval straddling the Mio-Pliocene boundary, the rate is either markedly reduced or a hiatus of about 1 m.y. is present. Neither the lithologic nor the physical properties of the sediments give a clue to the change in sedimentation regime which this implies.

The average measured acoustic velocity above 345 meters is 1.63 km/sec, a value somewhat, and perhaps significantly higher than the surprisingly low value of 1.53 km/sec calculated from the profiler records. Since the number of reliable sonic measurements is small due to the disturbed condition of the cores, it is difficult to evaluate the meaning of this discrepancy. The porosity of the upper unit decreases irregularly from more than 75 per cent in the disturbed sediment above 200 meters to less than 50 per cent near the base of the section.

At 345 meters, the drilling rate decreased abruptly and the first of a series of cherts that continue to the base of the section at 434 meters was cored. The top of the chert chalk sequence can be confidently correlated with the smooth acoustic basement of the profiler records.

The chert in the calcareous deposits forms nodules, stringers, and irregular masses, rarely as much as 1 cm across. Volumetrically, at least 90 per cent of this unit is chalk or chalk limestone. The chert is mostly olive brown in color, has a conchoidal fracture and is largely chalcedonic quartz with up to 50 per cent relict calcite inclusions. Siliceous tests either near or in the chert masses are extremely rare.

The chalk enclosing and separating cherty beds is similar to the overlying unit except for the lack of siliceous microfossils and the absence of  $H_2S$  odor. The increased circulation and torque required to penetrate the chert

resulted in poor recovery of the softer intervening calcareous sediment.

Sonic velocities in the 345 to 434 meters interval were generally measured on the more indurated carbonate fragments which tended to contain chert masses. The large range of values, from 1.9 to 4.2 km/sec, reflects the variation in proportion of the two rock types, and is certainly too high for the interval as a whole.

Basement at DSDP 157 is basalt. The upper few centimeters are vesicular, have flow structures, and consist of strongly devitrified silicified glass. Below this, the basalt is fairly massive with plagioclase laths and interstitial augite in an altered groundmass. The basalt is cut by joints and fractures filled with calcite, serpentine, pyrite, and a few other minerals. The gradation from chilled margin to crystalline interior and the absence of metamorphism or metasomatism of the overlying sediments show that the basalt is a flow rather than an intrusion.

As regards the hypothesis for the origin of the Panama Basin (van Andel et al., 1971), the data of DSDP 157 confirm that the Carnegie Ridge has remained within the depositional regime of the equatorial current system now prevailing at the site since the inception of deposition there. The lower sedimentary sequences at this site is very nearly identical with the lowermost deposits at DSDP 155 and 158. Magnetic anomaly data (van Andel et al., 1971; Heinrichs, in press) indicate that the Galapagos Rift north of DSDP 157 and west of the fracture zone at 85°W was activated somewhat before anomaly 3 time or around 4 to 5 m.y. ago. It is at this point that the hiatus or highly reduced sedimentation rate is observed at DSDP 157. Thus, this break may be at least partly the result of the tectonic disturbance produced by the onset of rifting.

In view of the middle Miocene age of the basement at DSDP 155 and 158, the much younger age of DSDP 157 is surprising and might be construed as an argument against the validity of the ancestral Carnegie Ridge concept. However, the approximate age of basement will be determined only if the oldest sediment in the area has been sampled. Since the acoustic basement here is chert, it effectively obscures the relief of the underlying volcanic basement. It is entirely possible, in fact even probable, that this volcanic basement. It is entirely possible, in fact even probable, that this volcanic basement possesses significant relief. Unless by coincidence the site is located in a position where the volcanic basement is deep and therefore was covered early by deposition, the chances are that the oldest sediment recovered was deposited significantly later than the time of formation of the basement. Alternatively, if the age of the basalt corresponds with that of the oldest sediment, late volcanism, either related to the Galapagos volcanic complex or to the process of rifting, may have produced an anomalously young basement. In the absence of an absolute age determination of the basalt itself,—which is probably not obtainable because of the alteration of the rock—the choice between these alternatives remains open. Furthermore, since the relief of the volcanic basement is not observable with ordinary geophysical techniques, there is no means of establishing whether the oldest and thickest sedimentary section in the area has been sampled.

The seismic reflection records of the site area show several internal unconformities in the upper portion of the sedimentary section to a depth of approximately 110 to 125 meters. This includes the entire Pleistocene portion of the sedimentary sequence. Since it is reasonable to assume that these unconformities are related to the processes causing erosion of the crest of the Carnegie Ridge and the formation of the canyons and gullies on its south flank, these erosion processes appear to be due to Pleistocene current conditions rather than to tectonic processes such as uplift of the ridge, which probably occurred considerably earlier during the onset of rifting.

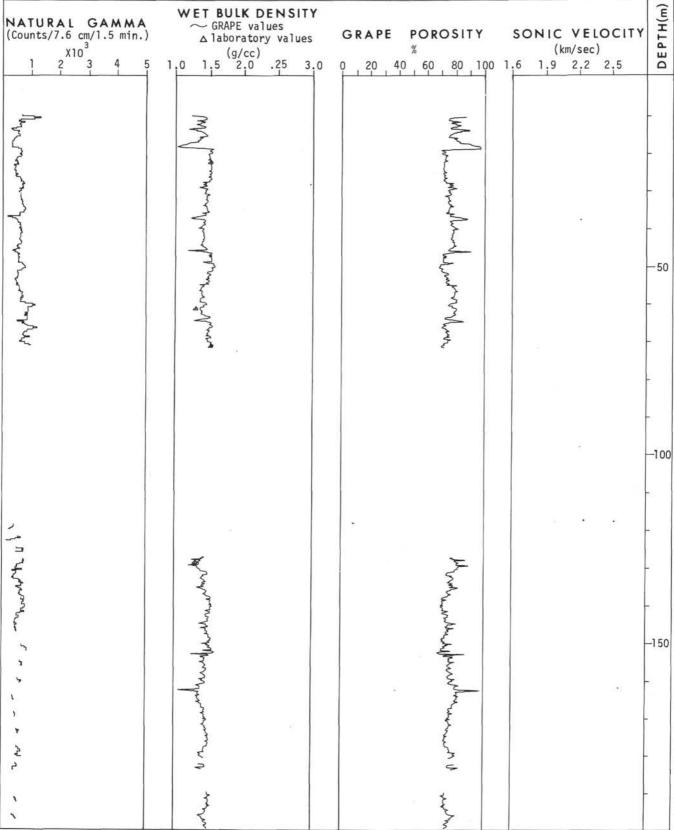
### REFERENCES

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- Winterer, E. L., Riedel, W. R. et al., 1971. Initial Reports of the Deep Sea Drilling Project, Volume VII. Washington (U.S. Government Printing Office).

1(m)			S OR ERIES	BI	OSTRATIGRAPH	IY
DEPTH(m)	CORES No./Depth	LITHOLOGY	SERIES OR SUBSERIES	FORAMINIFERA	NANNOFOSSILS	RADIOLARIA
		Diatomaceous foraminiferal nannofossil chalk ooze, light				
-	10		PLEISTOCENE			Collosphaera tuberosa
-	2 - 28 3 37 4 - 46	Image: Second	LATE PLEIS		Gephyrocapsa oceanica	Amphirhopalum ypsilon
- 50	5 - 55 6	mottling, silicoflagellates L L L Iocally common, pyritic, H <sub>2</sub> S dor L L L Iocally common, pyritic, H <sub>2</sub> S	PLEISTOCENE		Coccolithus	
-	- 64 7 - 73 8 8 82		EARLY	Pulleniatina obliquiloculata	doronicoides	Anthocyrtidium angulara
- 100	9 10 99 10 99 11 108 12	Image: Second				
	- 117 13 - 126 14 - 135 15 - 144		LATE PLIOCENE	Globorotalia limbata	Discoaster brouweri	Pterocanium prismatium
- 150 - -	16 	L L L Chalk ooze as above, but pale L L L Chalk ooze as above, but pale L L L J Chalk ooze as above, but pale L L L J Greenish gray to pale greenish yellow (10Y5-8/1-4), glau- conitic in part, less Foramin- ifera than units above.				
-	20 21 21 198			G. altispira		Spongaster pentas

Figure 4. DSDP 157, graphic hole summary. Vertical scale 1 cm = 10 m (1:1000).

PHYSICAL PROPERTIES



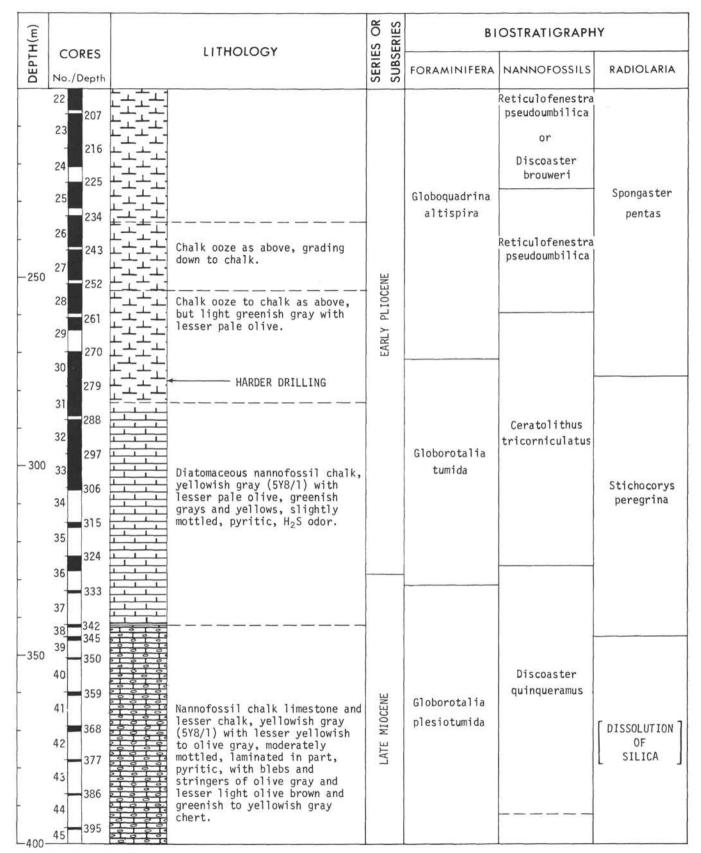
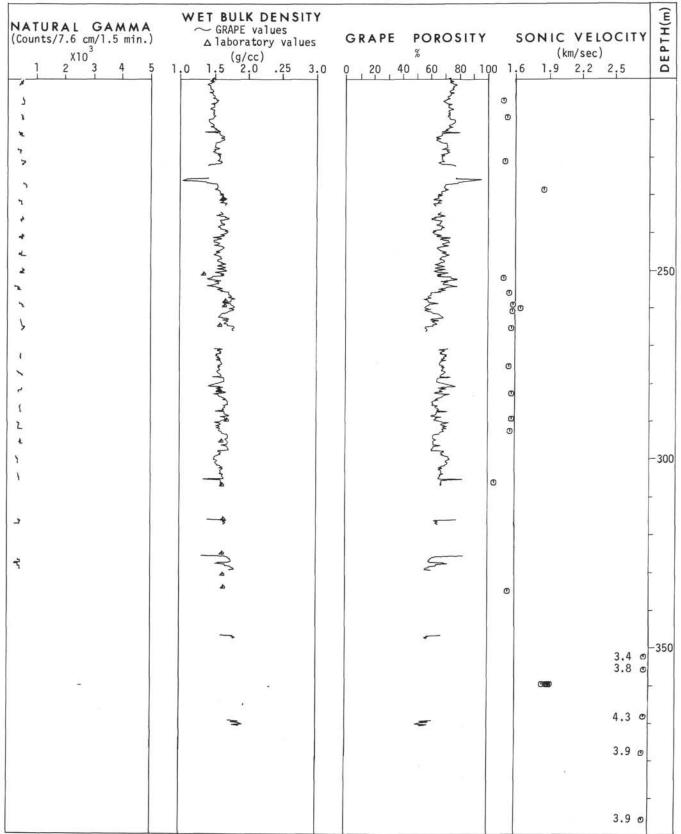


Figure 4. (Continued).

PHYSICAL PROPERTIES



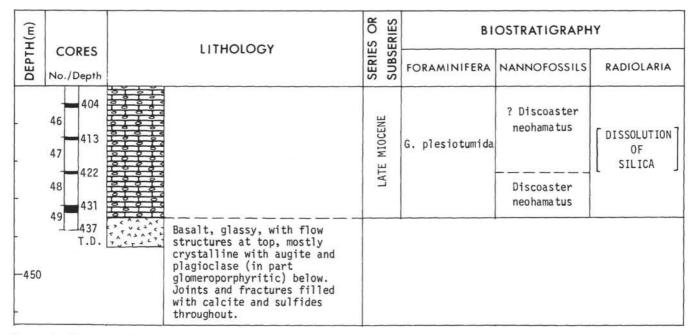
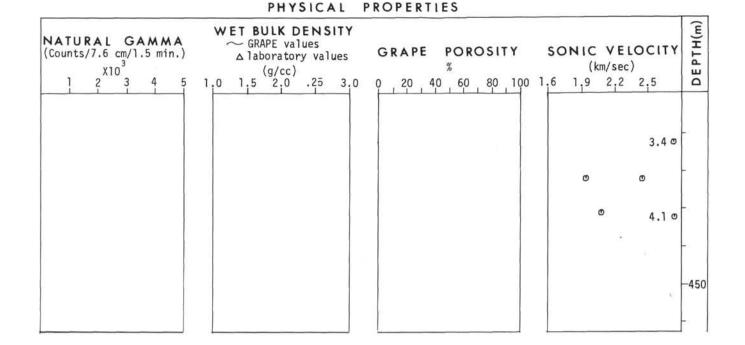


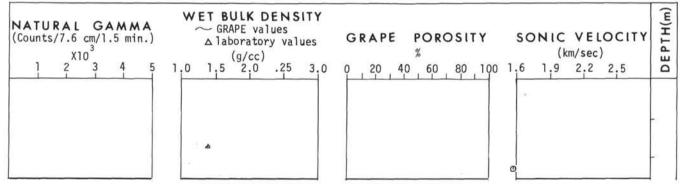
Figure 4. (Continued).

H(m)	CONFC	LITHOLOGY	S OR	SERIES	BI	OSTRATIGRAPH	Y
DEPTH(m)	CORES	timotogi	SERIES	SUBS	FORAMINIFERA	NANNOFOSSILS	RADIOLARIA
	1 0	Diatomaceous foraminiferal nannofossil chalk ooze, light greenish gray with lesser		ENE		Emiliania huxleyi or G. oceanica	Collosphaera tuberosa
	2	olives and yellowish gray, slight to moderate mottling,		1SI OCENE		Gephyrocapsa	LUDETOSA
		pyritic, H <sub>2</sub> S odor.		YLE		oceanica	Amphirhopalum ypsilon

Figure 4. (Continued).



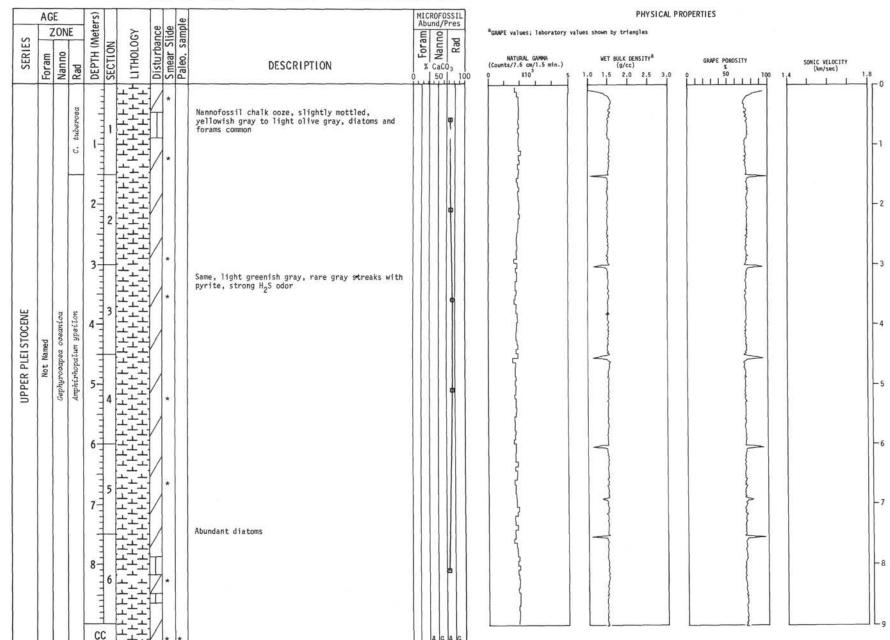
# PHYSICAL PROPERTIES



**SITE 157** 

					SITE	: 15	7 HOLE: CORE: 1 Cored Interv	1: 1	10-19	m	
S	AGE Z(	ONE	Aeters)	βGY	nce	ample		Abi	CROFOS	_	PHYSICAL PROPERTIES <sup>#</sup> GRAPE values; laboratory values shown by triangles
SERIES	Foram	Rad	DEPTH (Meters)	LITHOLOGY	Disturba	Paleo, sample	DESCRIPTION	E ** Foram	OUUEN CaCO	Long Rad	NATURAL GAMMA         WET BULK DENSITY <sup>®</sup> GRAPE PORDSITY         SONIC VELOCITY           (Counts/7.6 cm/1.5 min.)         (g/cc)         x         (km/sec)         1.8           0         x10 <sup>3</sup> 5         1.0         1.5         0         500         100         1.4         1         1         1.6
3		8		VOID 1 1 2 2 3		*	Nannofossil - foram chalk ooze, greenish gray, diatoms and rads common, spicules and silico- flagellates present	đ	A G	þ	
UPPER PLEISTOCENE		Gephyrocapea oceanica Collosphaera tuberosa	4 5 6	NOT OPENE 4	D						
			7		111111111	*	Nannofossil - foram chalk ooze, mottled, light olive gray to yellowish gray Same, light greenish gray			0	
			8111111	6 NOT OPENE	Đ		Same, massive, yellowish gray glauconite present				

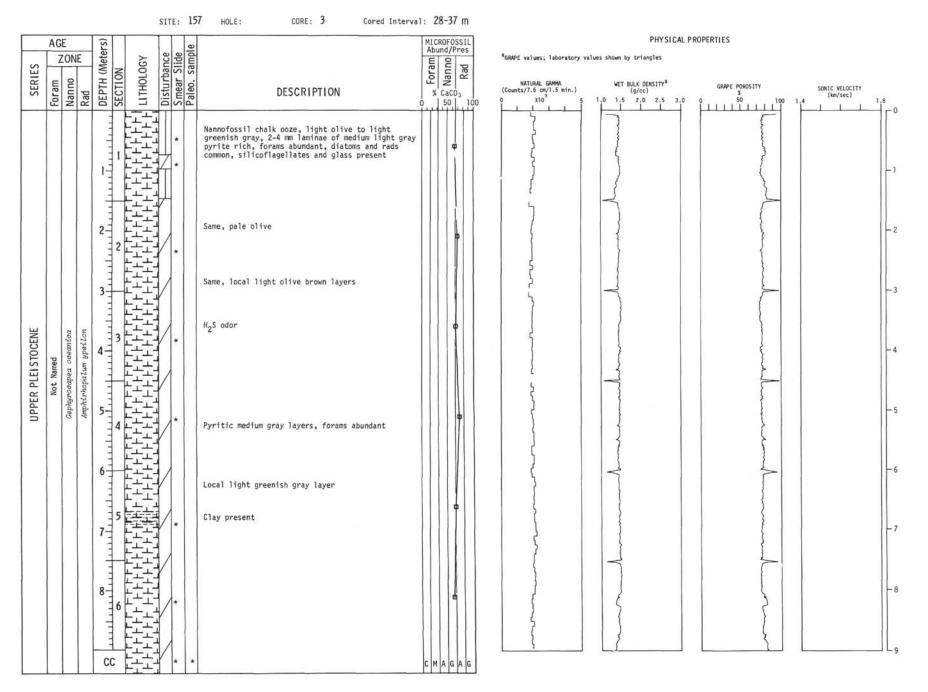




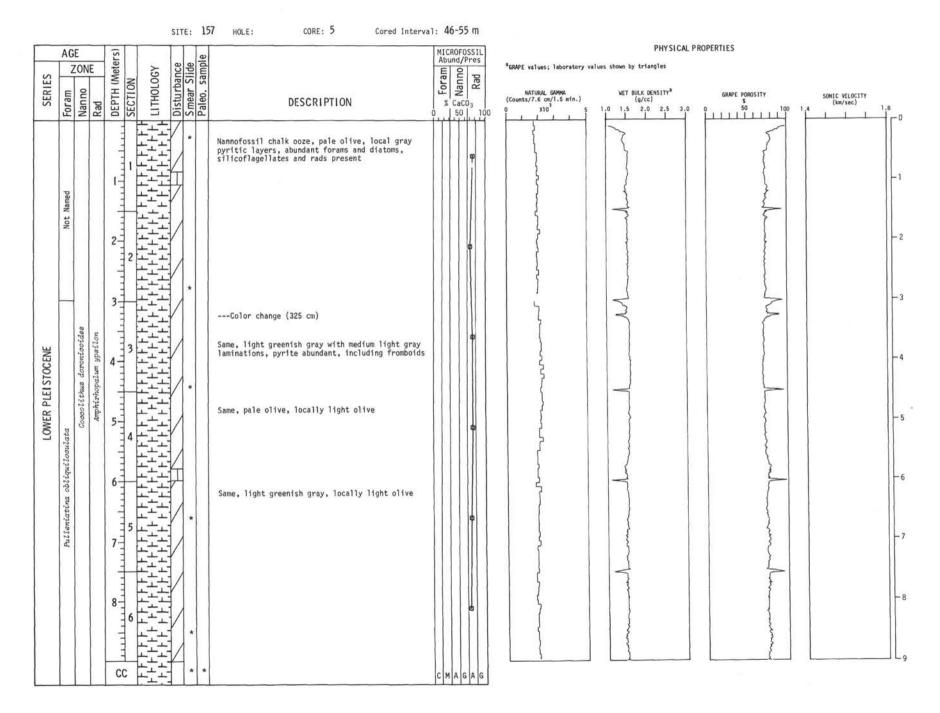
HOLE:

CORE: 2

Cored Interval: 19-28 m

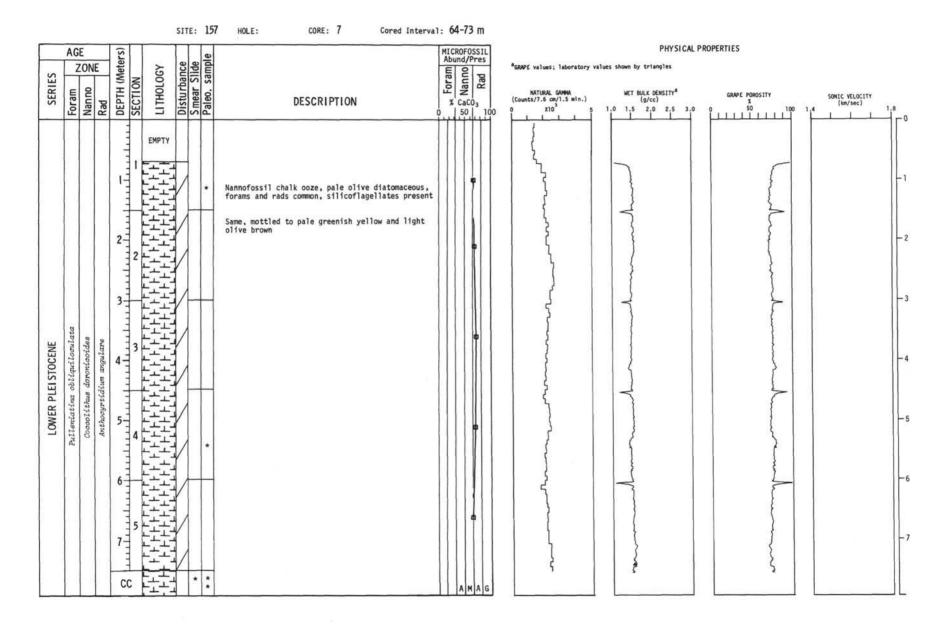


	AGE	_	-1	er			1		e		Ab	unc	FOSS J/Pre	s	5
DERIED	T	NE Ped	Nau	DEPTH (Meters)	JEGITON	LI THOLOGY	Disturbance	Smear Slide	Paleo. sample	DESCRIPTION	Foram	Nonan	ACO3	RdU	GRAPE values; laboratory values snown by triangles     NATURAL GAMMA WET BULK DENSITY <sup>a</sup> GRAPE PORDSITY SONIC VELOCITY     (Counts/7.6 cm/1.5 mfn.) (g/cc)      (km/sec)
								*		Nannofossil chalk ooze, pale olive, few gray pyritic streaks, diatomaceous, forams abundant, rads and silicoflagellates present			φ		
				2221						Same, few light olive brown streaks					
INVLINE	P	notran		3 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1									¢		
	Not Named	norinan negrocifingen		5111111				*		Same, few light greenish gray and dusky yellow layers			6		
				6				*					0		
				7	ניבביביב			*							
				8 Thurlin	121212								4		



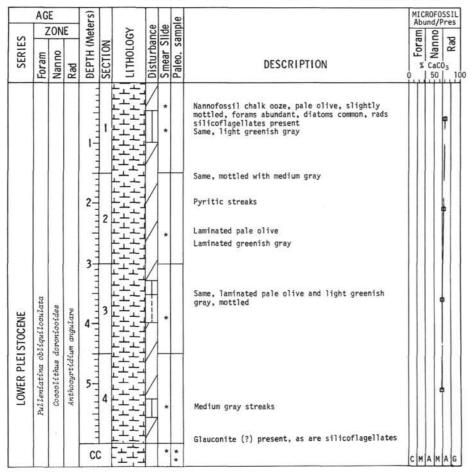
_				_	SITE: 1	57	HOLE: CORE: 6 Cored Interv					PHYSICAL F	DADEDTIES	
	AGE	ONE	Aeters)	GΥ	ide			Abu	ROFOSSIL nd/Pres		<sup>a</sup> GRAPE values; laboratory val			
SERIES	Foram	Rad	DEPTH (Meters)	LITHOLOGY	Disturbance Smear Slide Palen camule		DESCRIPTION	e Foram	CaCO <sup>3</sup> Rad	100	NATURAL GAMMA (Counts/7.6 cm/1.5 mfn.) 0 X10 <sup>3</sup> 5	WET BULK DENSITY <sup>a</sup> (g/cc) 1.0 1.5 2.0 2.5 3.0	GRAPE POROSITY SONIC S 100 1.4 () 111111111111111111111111111111111111	: VELOCITY m/sec} 1.8
						Nann	nofossil chalk ooze, very light olive , pale olive, diatomaceous, rads		φ					
			11111		*	\$111	coflagellates common							-1
			2						φ					-2
			3			Same	e, locally mottled light olive brown				5			-3
OCENE	uilooulata	doronicoides lum ypsilon	4				,		φ					-4
LOWER PLEI STOCENE	Pulleniatina obliquilooulata	Coccolithus doronicoide Amphirhopalum ypsilon	4											
LOW	Pulleni	Cocco	5-11-1	4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4					٥					- 5
			6		$\wedge$						<u> </u>			-6
			7	+_+ +_+ +_+ +_+ +_+ +_+ +_+ +_+ +_+ +_+					ø					-7
		_	11		*	-								
		angulare	811111	6 4 4 4		abund	e, pale olive, diffusely laminated, diatoms dant , very pale olive, gray pyritic laminations, odor		đ					-8
		A.	E CC		Т • ;	*		См	AMAG	G	Į		<u> </u>	

SITE 157





CORE: 8



	AG		-	ers)			a	e	ple			ROF0	SSIL
SERIES	Foram	Nanno Na		DEPTH (Meters)	SECTION	LITHOLOGY	Disturbance		Paleo. sample	DESCRIPTION	E Foram	CaC 50	Rad
	2			1	1			*		Nannofossil chalk ooze, yellowish olive brown mottled to moderate olive brown Color change Same, pale olive, rads, diatoms, forams common, silicoflagellates and discoasters present Pyrite framboids			
UPPER PLIOCENE	Fulleniatina obliquiloculata	Discoaster browneri	Pterocanium prismatium	3	3			*				8	
				5	4					Same, pale grayish olive			
				C	;				*		RP	A Po	AG

STTE: 157 HOLE .

CORE: 10

Cored Interval: 91-99 m

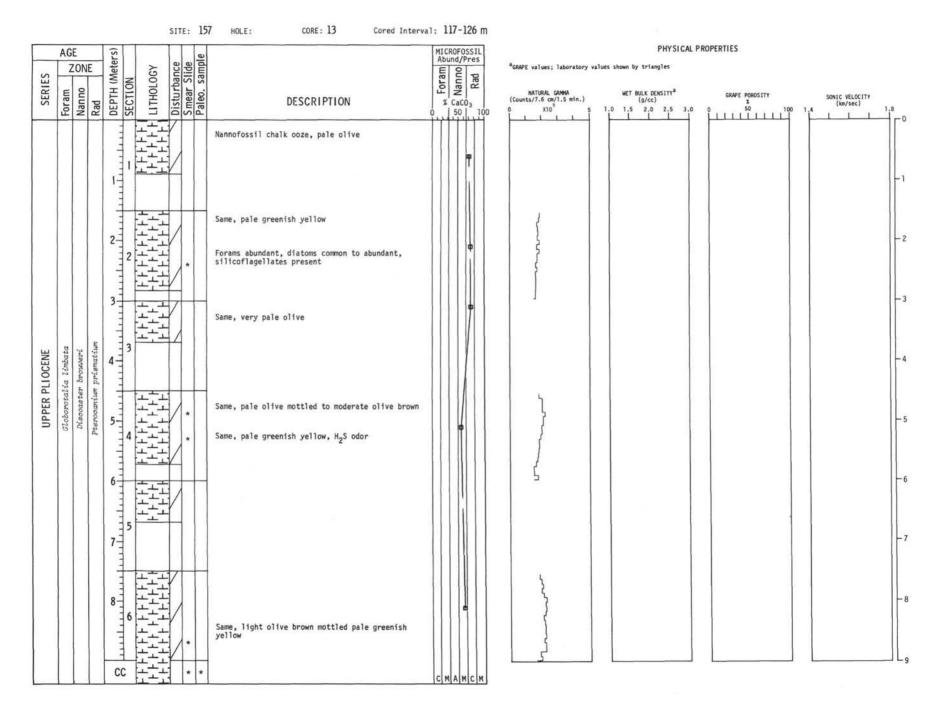
	AG			ers)					e		MI A	LCR	ROF nd/	OSS	IL
SERIES	Foram	Nanno No		DEPTH (Meters)	SECTION	К ОТОНАТИ	Disturbance	Smear Slide	Paleo. sample	DESCRIPTION		N LOI GIII	Can Sanno	:03	DEN 10
UPPER PLIOCENE	Pulleniatina obliquilooulata	Discoaster broweri	Pterooanium priematium	2-3-	2	┥ <u>╗</u> ┝┾┾┥ <u>╗</u> ┝┾┍┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝		*		Nannofossil chalk ooze, olive brown, alternating with mottled pale olive and light olive brown, strong H2S odor Forams and diatoms common, rads, glauconite, and silicoflagellates present Same, yellowish gray and light olive gray Same, pale olive and light olive brown Zone abundant in diatoms			All		
	G. Limbata			5-	4			*	*	Same, yellowish gray Same, pale olive to olive brown Same, light olive brown, glauconite present				0	

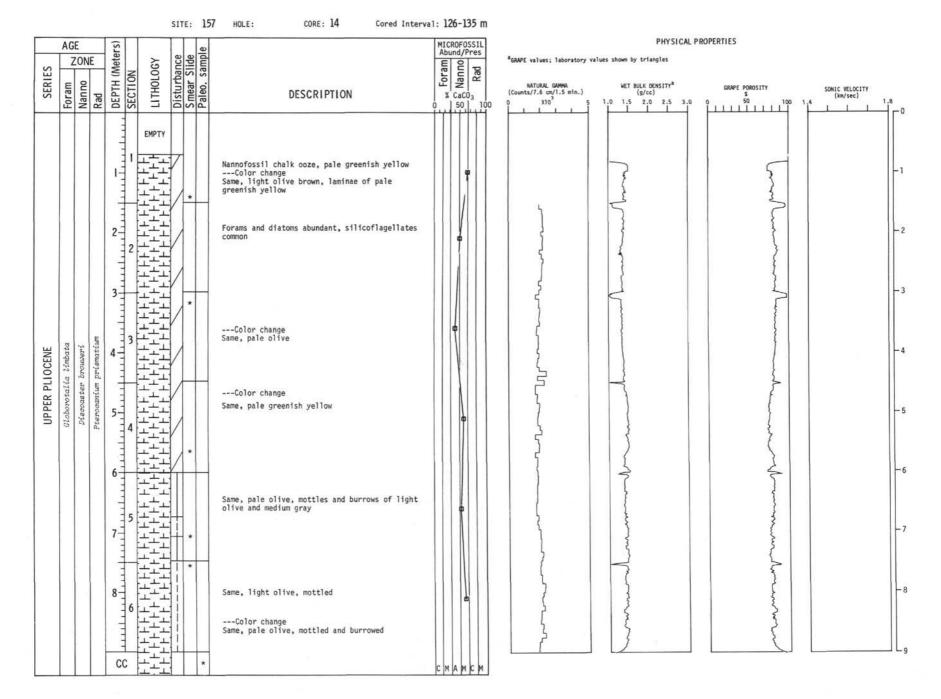
SITE: 157 HOLE:

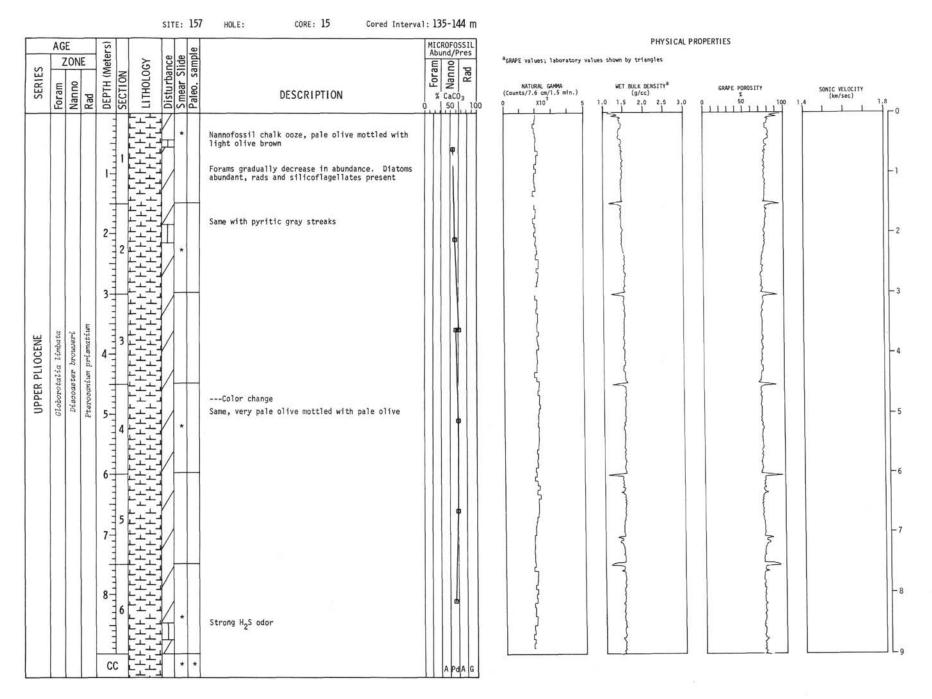
CORE: 11

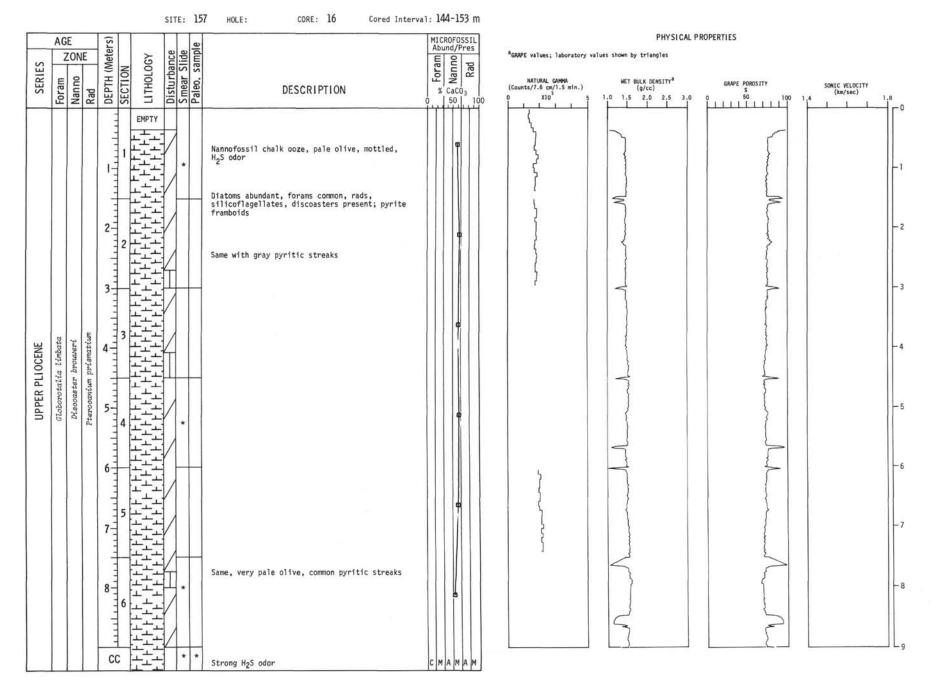
Cored Interval: 99-108 m

AGE 2		MIC	ROF0 und/f	SSIL
Sample Sance		oram	CaC 50	Rad
SERIE: Foram Nanno Rad DEPTH (A DEPTH (A LITHOLO LITHOLO LITHOLO Disturba Paleo, sa	DESCRIPTION	x o	CaCi 50	 0 <sub>3</sub>  1(
UPPER PLIOCENE Cloborotatia Limbata Discocarta Limbata Discocar				

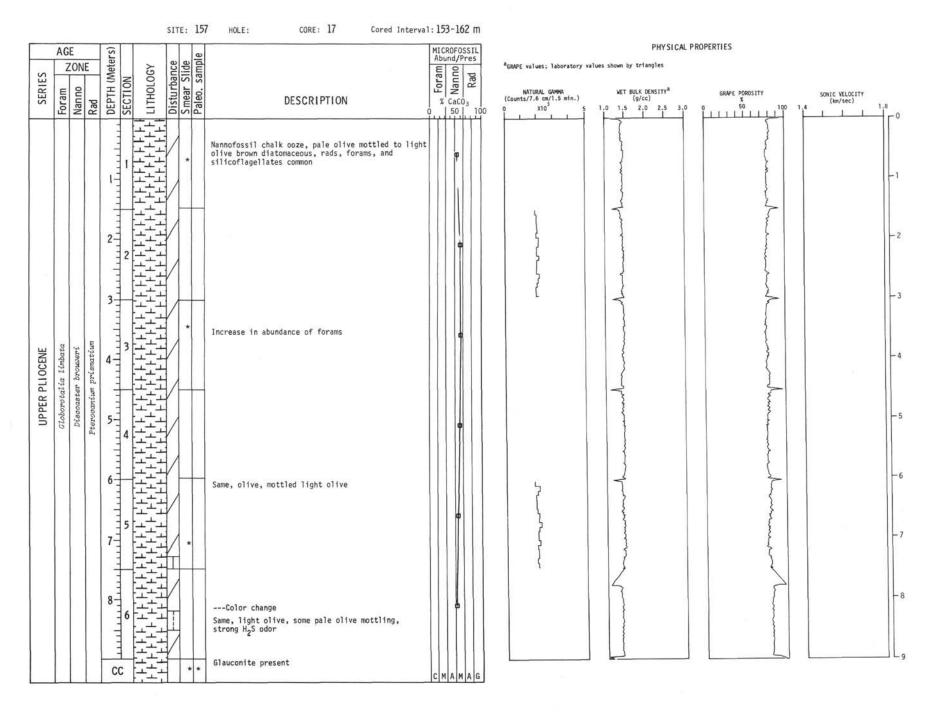


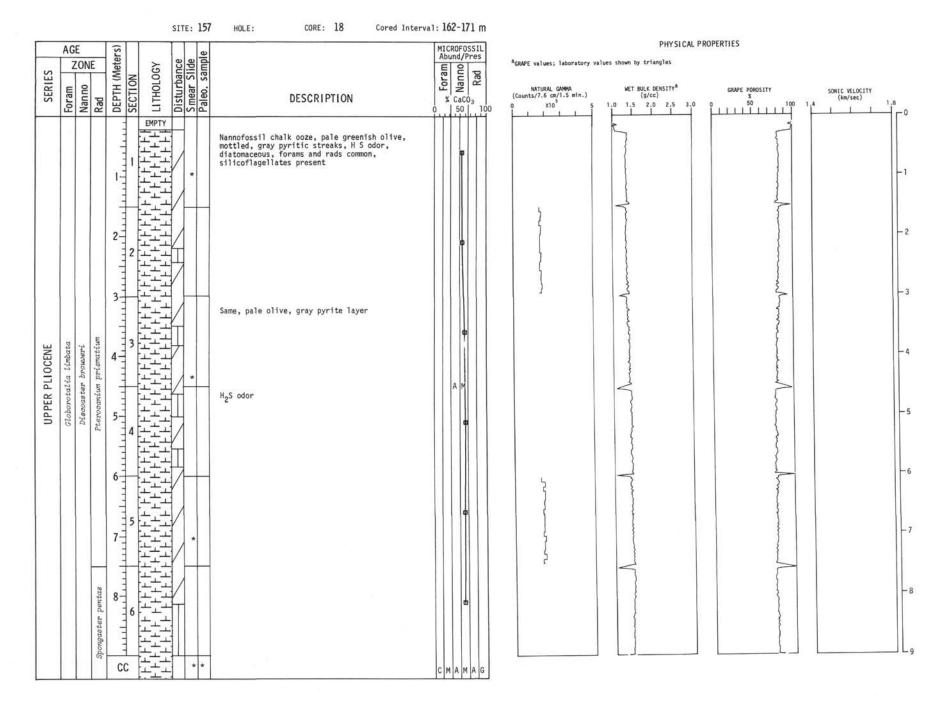




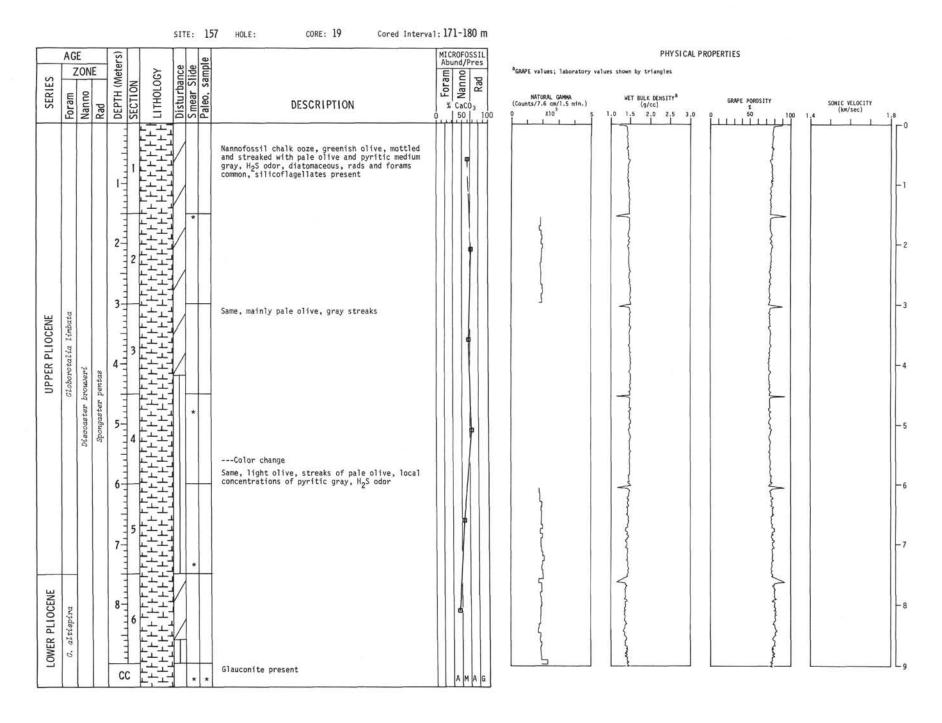


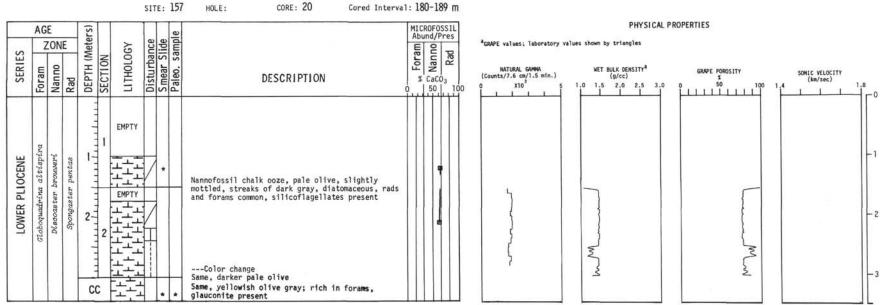
SITE 157





SITE 157





SITE: 157

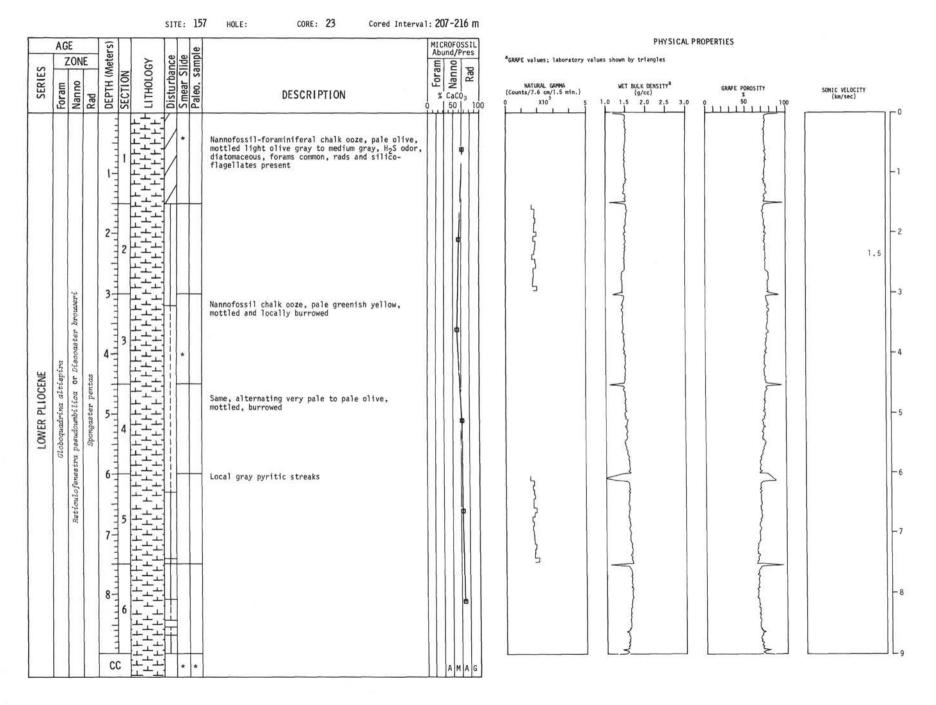
HOLE:

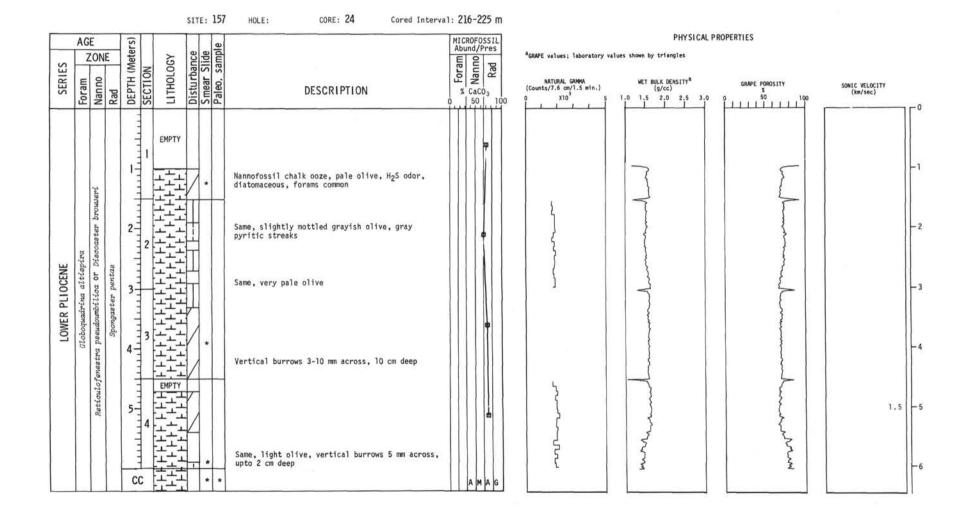
Cored Interval: 180-189 m

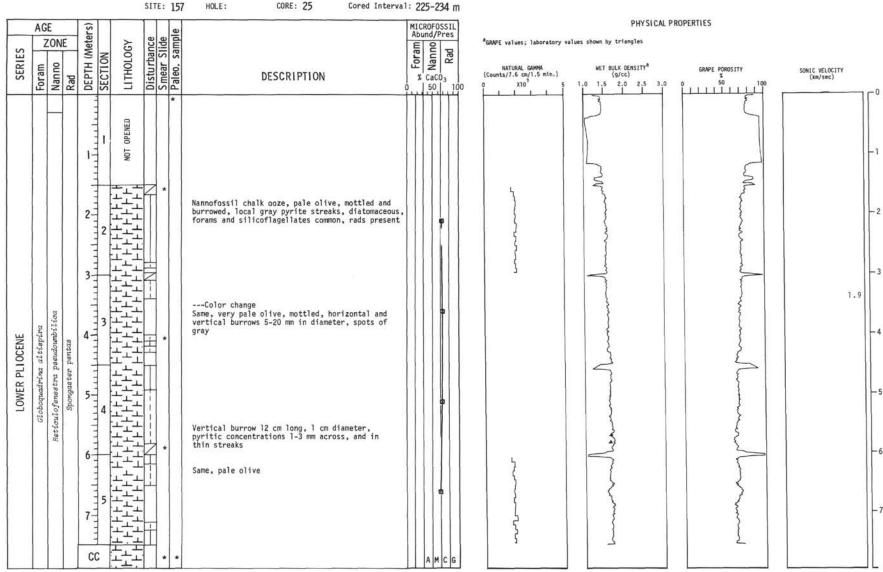


1				1	SITE:	157	HOLE: CORE: 21 Cored Interva				
ES		ONE	DEPTH (Meters)	SECTION	Disturbance Smear Slide	. sample		Foram Por	ROFOS und/P OuueN	Kad Rad	PHYSICAL PROPERTIES <sup>a</sup> GRAPE values; laboratory values shown by triangles NATURAL CANNA UPT OUT OUT OFFICE/
SEF	Foram	Rad	DEPT	SECTION	Distu	Paleo	DESCRIPTION	2	CaCO	3 100	NATURAL GAMMA         WET BULK DENSITY*         GRAPE POROSITY         SONIC VELOCITY           (Counts/7.6 org/1.5 min.)         (g/cc)         %         %         %         %           0         X10 <sup>3</sup> 5         1.0         1.5         2.0         2.5         3.0         50         100         1.4         1.4         1.4         1.4         1.4         1.4         1.4         1.4         1.4         1.4         1.4         1.4         1.6         1.8         1.6
		recoaster prowert	2				Nannofossil chalk ooze, pale olive mottled darker pale olive; H2S odor, diatomaceous, forams and silicoflagellates common		æ.		
LOWER PLIOCENE	Globoquadrin	uuaa or vuscoaster prouveru Spongaster pentae	3 4 11111111111111111111111111111111111						8		
		nerrowwowaetra pseudowworroa	6 7 7	44444 4444444444444 44444444444444			Same, mottled pale greenish yellow and olive gray; H <sub>2</sub> S odor		0		
			8 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1		*	*	H <sub>2</sub> S odor		A M	CG	

					SITE:	157	HOLE: CORE: 22 Cored Interv	a1:19	98-207 n	n				
	AGE ZC	ONE	Meters)	GY	lide	ample		Abi	ROFOSSII		<sup>a</sup> GRAPE values; laboratory v	PHYSICAL F alues shown by triangles	ROPERTIES	
SERIES	Foram	Rad	DEPTH (Meters)	LITHOLOGY	Disturbance Smear Slide	Paleo. s	DESCRIPTION	Foram	CaCO <sub>3</sub>	100	NATURAL GAMMA (Counts/7.6 cm/1.5 min.) 0 X10 <sup>3</sup> 5	WET BULK DENSITY <sup>8</sup> (g/cc) 1.0 1.5 2.0 2.5 3.0	GRAPE POROSITY	SONIC VELOCITY (km/sec)
							Nannofossil chalk ooze, pale olive, mottled, diatomaceous, rads and forams common, silico- flagellates present		φ 					-1
			21111				Same, streaks of dark to medium gray		8			- Announce	h	- 2
	lepira	odster. promete	3		4		Same, very pale olive		Ð		}	m	nation	- 3
	alti	pent	4		/.  -									- 4
MOT	Globoquadrina	odg Sbou	5				Same, pale olive, mottled light olive and medium gray		¢					- 5
		nafornorrau	6								7			-6
			7		*						}	4		1.5 -7
			8 1111111						đ					-8
			E S			*			AMC	G		<u>}</u>		

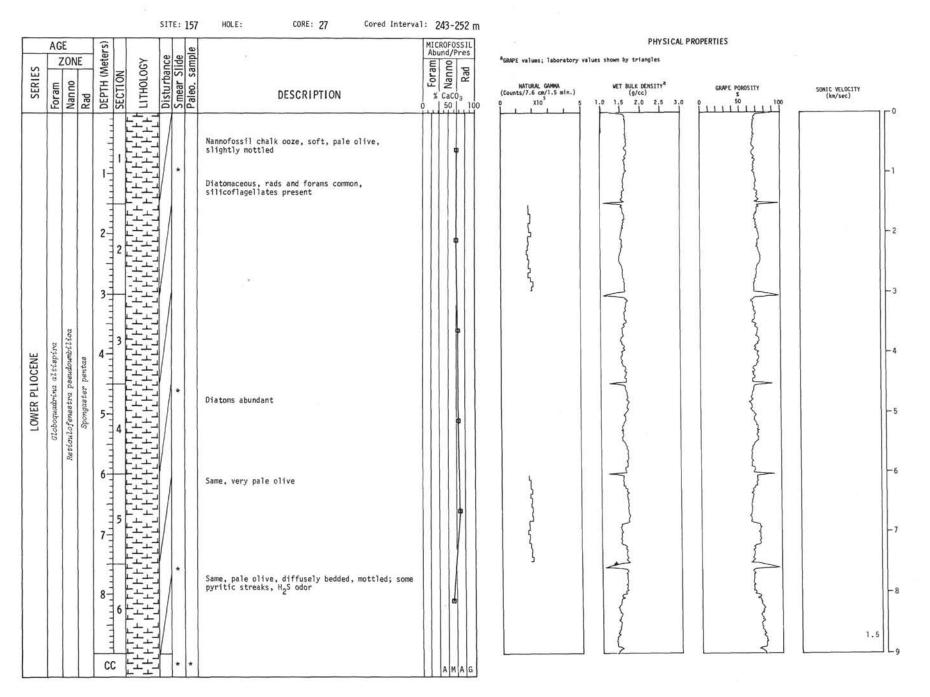


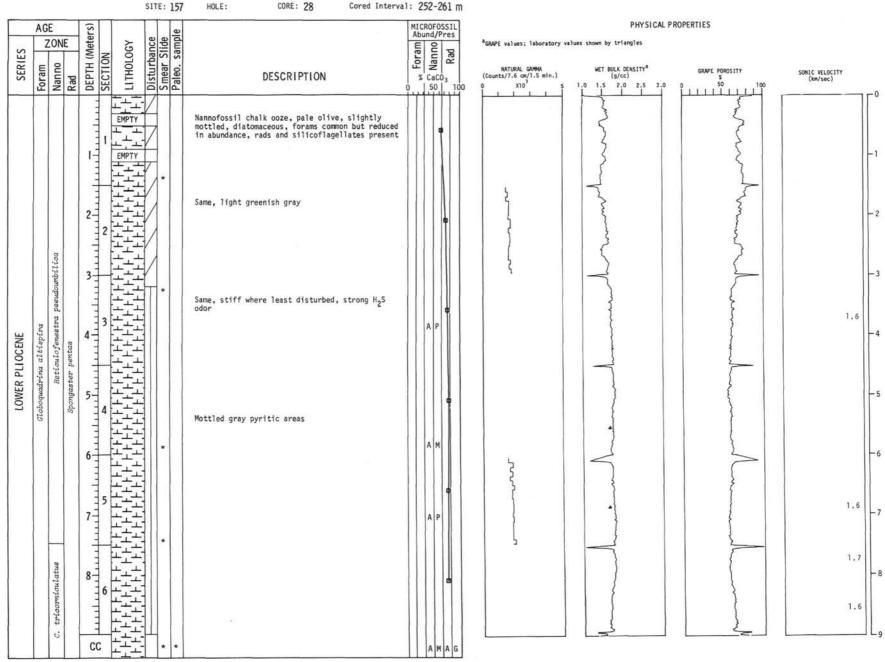


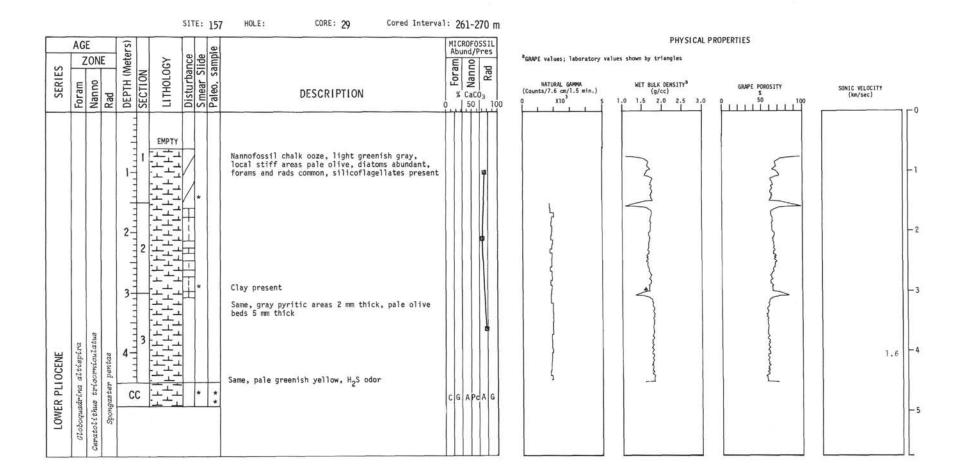


SITE: 157

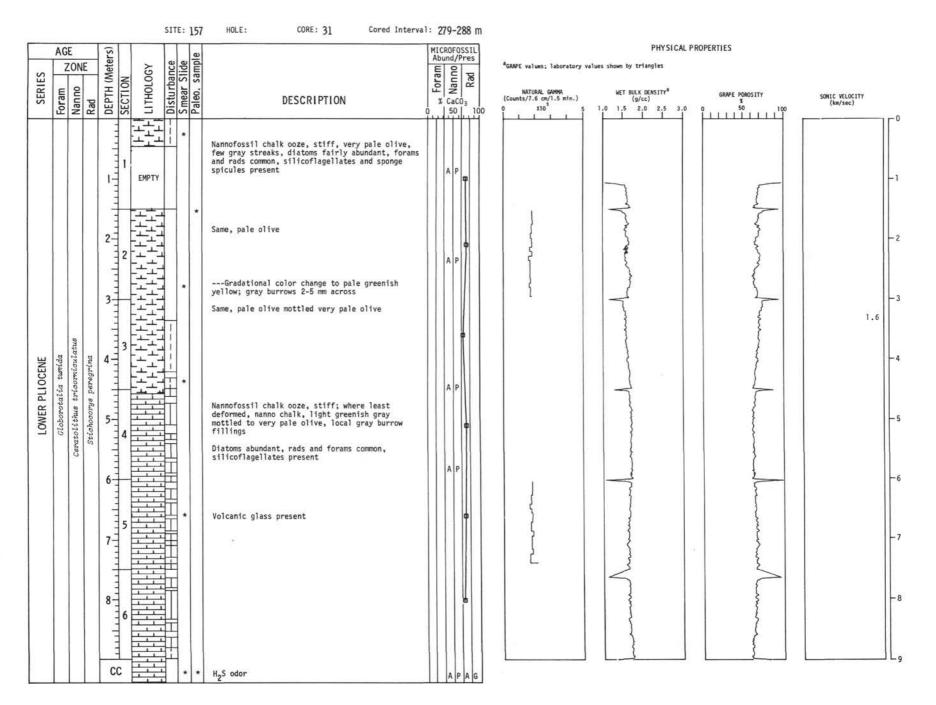
Т	AGE ZO	ONE	Meters)	N DGY	ance Slide		Foram PUL	ROFOSSIL and/Pres Paulon Roperati	<sup>a</sup> GRAPE values; laboratory v	PHYSICAL P alues shown by triangles	ROPERTIES	
	Foram	Rad	DEPTH (Meters)	LITHOLOGY	Disturbance Smear Slide Paleo. sample	DESCRIPTION	%	CaCO <sub>3</sub>	NATURAL GAMMA (Counts/7.6 cm/1.5 min.) D X10 <sup>3</sup> 5	WET BULK DENSITY <sup>a</sup> (g/cc) 1.0 1.5 2.0 2.5 3.0	GRAPE POROSITY 1 0 50 100	SONIC VELOCITY (km/sec)
			111111111			Nannofossil chalk ooze, stiff, pale olive, slightly mottled, few gray streaks, forams and diatoms common, rads and silicoflagellates present		F				
			2		*	Same, very pale olive		æ	}	-		1.4
			31111			Color change Same, pale olive mottled light olive, burrows up to 1 cm diameter, not stiff Same, stiff					maha	
1000	altispira emidnebilion	entas	4	3		Same, burrowed, one burrow 10 cm long		e				
	Globoquadrina altispira	sucofenesura peenaa Spongaster pentas	5			Same, very pale olive, burrows 5 mm diameter filled with grayish olive Color change Same, greenish gray		٥				
		Retro	6			Same, mottled very pale olive, stiff			5			
			7		*	Broadly layered. Diatoms abundant H <sub>2</sub> S odor		ø		and man	andren	
			1111									

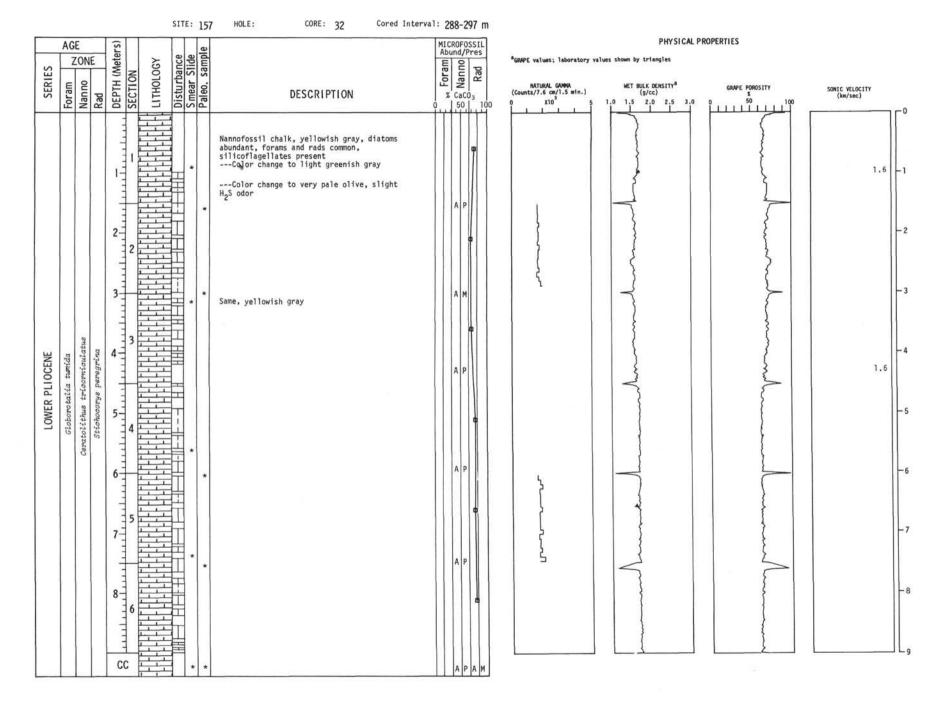




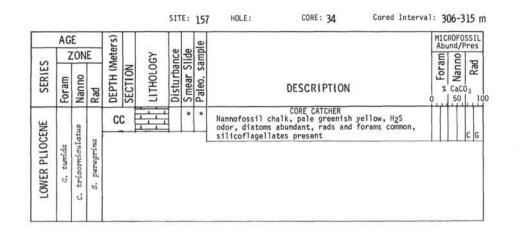


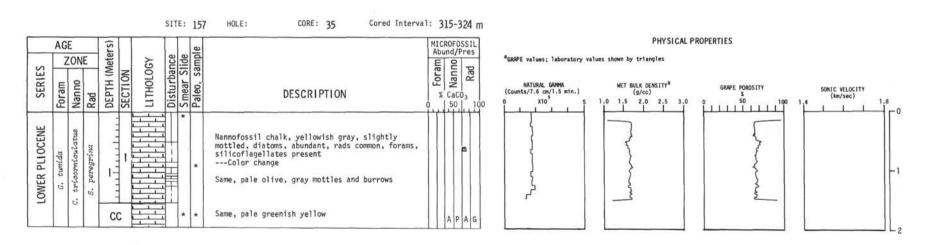
	ZO	NE	eters		7	ce	mple		Ab	ROFOSSI und/Pres	-	<sup>a</sup> GRAPE values; laboratory va	PHYSICAL F lues shown by triangles	KOPERIES	
JUNICO	Nanno	T	DEPTH (Meters)	SECTION	ГІ ТНОГОСУ	Disturban	Paleo, sample	DESCRIPTION	Foram	CaCO <sub>3</sub>	000	NATURAL GAMMA (Counts/7.6 cm/1.5 min.) 0 X10 <sup>3</sup> 5	WET BULK DENSITY <sup>®</sup> (g/cc) 1.0 1.5 2.0 2.5 3.0	GRAPE POROSITY \$ 0 50 100	SONIC VELOCITY (km/sec)
	G. attispira			115			*	Nannofossil chalk ooze, light greenish gray to pale olive, diatomaceous, forams and rads common, silicoflagellates present		Ø				harrow	
		ae	2-				e:	Same, light greenish gray		•		}	T	when	
		Spongaster pentas	3-					Same, very pale olive					m	mand	
	tricomiculatus		4							ø					
	otalia tumida Ceratolithus tridorr									8			- marine	harris	. 1
	Globorotalia tumida Ceratolithus		5-	11				Light olive burrows				L.			
		puing	7-	5				Same, light greenish gray, H <sub>2</sub> S odor, stiff in undisturbed areas, local mottles to light olive		æ					
		Stichocorys peregrina					*					Ş			
		St	8	6						6					

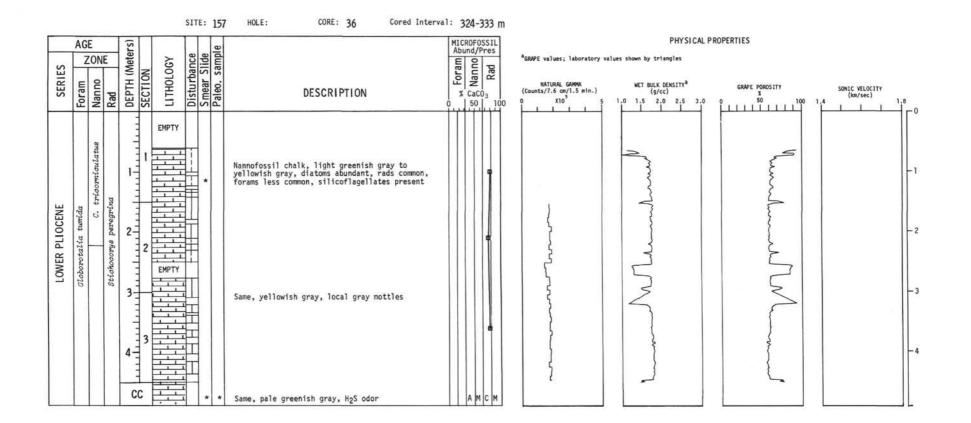




10	AGE	ONE	Aeters)	ζ	nce Iide ample		At	CROFOS		<sup>8</sup> GRAPE values; laboratory v	values shown by tri	PHYSICAL PI angles		
SERIES	Foram	Rad	DEPTH (Meters)	SECTION LITHOLOGY	<ul> <li>Disturbance</li> <li>Smear Slide</li> <li>Paleo. sample</li> </ul>	DESCRIPTION	Foram	% CaCO	<sup>2</sup> Rad	NATURAL GANMA (Counts/7.6 cm/1.5 min.) 0 X10 <sup>3</sup> 5	WET BULK (g/c	DENSITY <sup>a</sup> c) D 2.5 3.0	GRAPE POROSITY	SONIC VELOCITY (km/sec)
			1-		*	Nannofossil chalk, pale olive grading down to light pale olive, H2S odor, diatomaceous, forams and rads common, silicoflagellates present		A P	7					
			2-			Same, burrows are dusky yellow		A						
LOWER PLIOCENE	Globorotalia tumida	Ceratolitmus tricomiculatus Stichocorys peregrina	4-			Same, pale olive mixed with medium light gray in 5-20 mm beds and in burrows. Zone rich in pyrite Same, pale olive Same, yellowish gray		Ð			munu		mundum	
10/	Glob	Ceratolii Stich	6-			Same, light greenish gray		A P	2		m			
			7-			Same, yellowish gray		ΑP	5		- Marine		man have	
			c		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	H <sub>2</sub> S odor								1,.







AGE

SERIES

UPPER MIOCENE plesiotumida quinqueramus S. peregrina

6

à

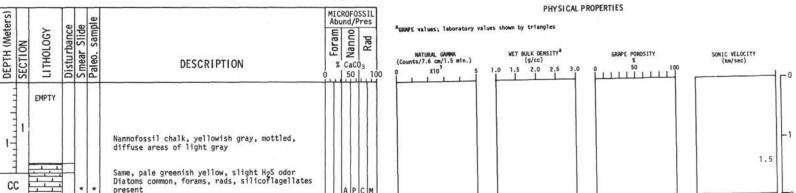
ZONE

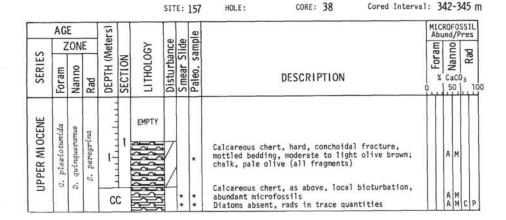
Foram Nanno Rad

DEPTH (Meters)

CC

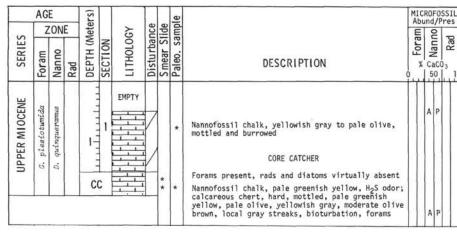
Cored Interval: 333-342 m





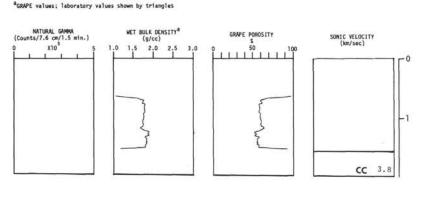
HOLE: CORE: 39 Cored Interval: 345-350 m

100



SITE: 157

PHYSICAL PROPERTIES

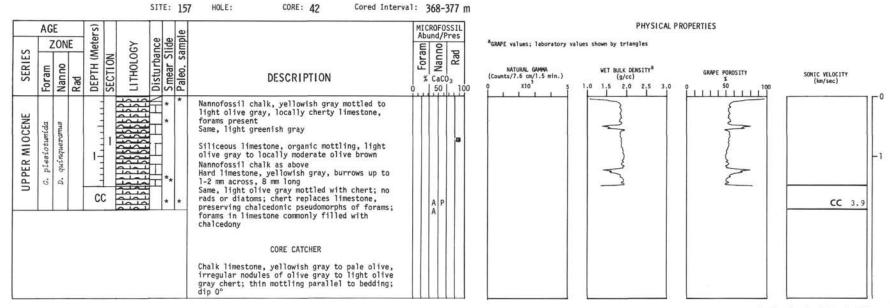


**SITE 157** 

	AG			ers)				e			ROFO	
SERIES	Foram	Nanno	Rad	DEPTH (Meters)	SECTION	КЭОТОНТІ	isturb	Paleo, sampli	DESCRIPTION	Foram	CaCO 50 Nanno	Rad
UPPER MIOCENE	G. plesiotumida	D. quinqueranus				EMPTY			CORE CATCHER Chalk limestone, yellowish gray; and calcareous chert, yellowish gray to light olive gray, mottled, laminated Forams in limestone, but no rads or diatoms			

CORE: 41 Cored Interval: 359-368 m SITE: 157 HOLE: Nanno Rad Rad Meters) SECTION MICROFOSSIL Abund/Pres wervoy K CacO3 0 50 10 AGE Disturbance Smear Slide Paleo, sample ZONE LI THOLOGY SERIES Foram DESCRIPTION 100 0 EMPTY UPPER MIOCENE G. plesiotumida D. quinqueramue Chalk limestone, hard, fine grained, cherty, yellowish gray mottled moderate olive brown and yellowish gray; chert mottles moderate olive brown Chalk, yellowish gray and light olive gray Calcareous chert, pale olive Chalk, yellowish gray Tur CC CORE CATCHER Calcareous chert, light olive gray, interbedded with chalk limestone, yellowish gray

+



Sonic Velocity (km/sec) 4.2

	AG			ers)			-		ole			ROFO	
SERIES	Foram	Nanno	Rad	DEPTH (Meters)	SECTION	LITHOLOGY	Disturbance	Smear Slide	Paleo. sample	DESCRIPTION	Foram	Ca So Nanno	
UPPER MIOCENE	G. plesiotumida	D. quinqueramus		1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-			Ζ		*	Nannofossil chalk limestone, laminated to nodular, yellowish gray to very light gray, bioturbation and mottles, stringers of nodular chert, mottled light olive gray to olive gray, 1-5 mm thick		A	

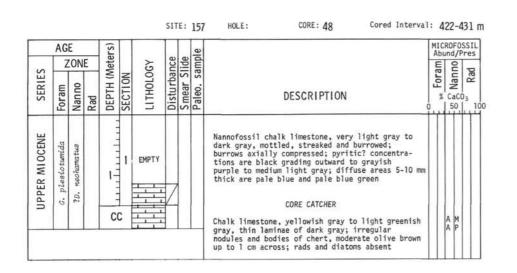
A	GE		 SLS)			ele			ROF0	
SERIES	T	Nanno	DEPTH (Meters) SECTION	<b>VIDUOUSY</b>	Disturbance	Paleo, sample	DESCRIPTION	Foram	CaC Solution	
UPPER MIOCENE	~	7D. quinqueramue	CC Sanic	Velocit	y (kn		Chalk limestone, yellowish gray, thin streaks light olive gray; irregular lenses of olive gray chert; silic. veins to 2 mm cross bedding 3.9; CC		A P	

	AG		_	ers)				-	e		MIC	ROF	OSSI Pre
SERIES	Foram	Nanno No	Rad	DEPTH (Meters)	SECTION	ГІ ТНОГОСУ	Disturbance	Smear Slide	Paleo. sample	DESCRIPTION	E . Foram	Cal 50	03
UPPER MIOCENE	G. plesiotumida	ND. nechamatus		1				*	* *	Chalk limestone, yellowish gray, massive, with nodules of olive gray chert		P	2

	AG			ers)					e			ROF0	
S	Z	ON	E	Mete	-	λĴ	ance	Slide	sample		Foram	ou	Rad
SERIES	Foram	Nanno	Rad	DEPTH (Meters)	SECTION	LITHOLOGY	Disturb	Smear S	Paleo. sample	DESCRIPTION	For *	1. 1. C. C. C. L.	0.000
UPPER MIOCENE	G. plesiotumida	7D. neohamatus		1- C(	1	EMPTY	/			Chalk limestone, hard, light greenish gray, mottled to greenish gray, pale olive; irregular concretionary bodies of chert, olive black, abundant bioturbation CORE CATCHER		A P	
										Thin section: silica replacing calcite matrix and fossil fragments Chalk limestone, yellowish gray, 1 mm concentra- tions of chert; irregular contact with massive chert, olive gray mottled to light olive gray			

	ROFO: nd/P						le					ers)			AG	
03	Nanno	%	Foram		SCRIPTION		aleo. sample	Smear Slide	sturbance	гі тногосу	SECTION	DEPTH (Meters		Nanno No	Foram	SERIES
4	50	Н	÷.				Å	SI	ō	=	SE	B	Rad	Na	요	-
				nd olive gray,	limestone, light vellowish gray a g, burrows compro	with stringe			7	EMPTY	1	1		PD. neohamatus	G. plesiotumida	UPPER MIOCENE
	A P			discontinuous	CORE CATCHER ale olive, mottle ay lenses, local tic concentratio	olive to med	*		1	<u></u>		C				

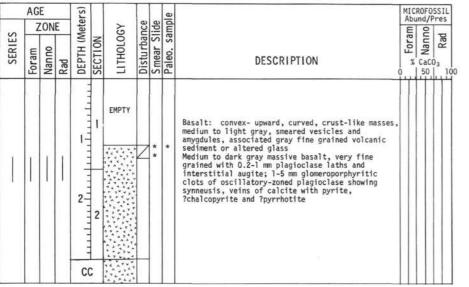
SITE 157

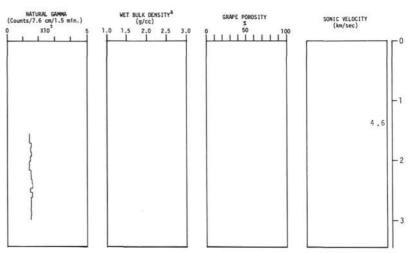


SITE: 157

HOLE:

CORE: 49 Cored Interval: 431-437 m





## PHYSICAL PROPERTIES

<sup>a</sup>GRAPE values; laboratory values shown by triangles

	AG	EZON	E	eters)		7	ce	de	nple				0SSI /Pre	s
SERIES	Foram	Nanno	Rad	DEPTH (Meters)	SECTION	LITHOLOGY	Disturbance	Smear Slide	Paleo. sampl	DESCRIPTION	Foram	Cal Solution	CÔ3	1 INdu
UPPER PLEI STOCENE OR HOLOCENE	Not Named	Cephyrocapea coecuica Dr Bmiliania huzleyi	Collospiaera tuberosa	2 3 4 5 7 7	2 3 4 5 6			* * * * * * * * * * *	*	Nannofossil-foraminiferal chalk ooze, light olive diatoms abundant, rads common Same, mottled pale olive to moderate olive brown, burrowed; silicoflagellates present Same, moderate greenish yellow; burrow 5 mm diameter moderate olive brown Same, light olive Same, pale olive, disturbed mottles and 10 mm burrows of light olive S mm burrows filled with light olive Same, mottled light olive and pale olive Burrows Massive Slightly mottled Same, light greenish gray mottled pale olive Light olive mottled greenish gray Same, mottled pale olive to very pale olive; light gray burrows and 3-8 mm streaks Same, light greenish gray; light gray 1-2 mm burrows and 3-8 mm discontinuous laminaes Same, pale olive, light olive interbedded 5 cm light olive interbed; otherwise pale olive Same, very pale olive, darker interbeds, medium light gray streaks Same, pale olive, streaks light gray, burrows up to 5 mm Same, with light greenish gray interbeds Large burrow 2 mm across		A		

SITE: 157 HOLE: A

CORE: 2

DESCRIPTION

Nannofossil-foraminiferal chalk ooze, H<sub>2</sub>S odor, pale olive grading down to very pale olive,

Light greenish gray to light olive burrows Medium light gray streaks

Same, pale olive mottled very pale olive and light gray

Forams reduced in abundance, diatoms abundant, rads common, silicoflagellates present

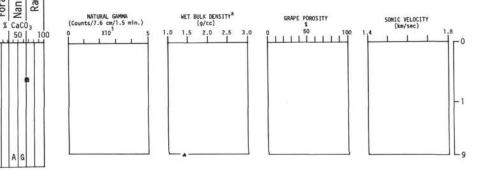
Cored Interval: 9-18 m MICROFOSSIL Abund/Pres Wand Our Bar % CaCO3

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AIG

## PHYSICAL PROPERTIES

<sup>a</sup>GRAPE values; laboratory values shown by triangles



<sup>1</sup>Gephyrocapea oceanica or Emiliania huxleyi <sup>2</sup>Gephyrocapea oceanica

Disturbance Smear Slide Paleo. sample

LI THOLOGY

EMPTY

1-1-1-1

حيث 1

-<u>-</u>---

Т. 1 1 Ê.

1

AGE

SERIES

UPPER PLEI STOCENE

Not Named

2

ZONE

Foram Nanno Rad

tuberosa

Collasphaera

DEPTH (Meters)

The Trees

CC

SECTION

A	GE ZON	IF	ters)		> 00	ple		Abu	ROFOSSI und/Pre	s	PHYSICAL PROPERTIES <sup>a</sup> GRAPE values: laboratory values shown by triangles
Foram	Γ		DEPTH (Meters)	SECTION	LITHOLOGY Disturbance	Paleo, sample	DESCRIPTION	Foram	CaCO3		NATURAL GAMMA WET BULK DENSITY <sup>a</sup> GRAPE PORDSJTY SONIC VELOCITY (Counts/7.6 cm/1.5 min.) (g/cc) 5 (m/sec)
			The function of the second			-	Nannofossil-foraminiferal chalk ooze, H2S odor, light greenish gray, local mottling to light olive Diatoms abundant, rads common, silicoflagellates present		φ		
			2-1-1-1-1-3	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2		e.	Same, pale olive with light olive mottles Same, with light gray di-fuse laminaes 5 mm dusky yellow burrow		æ		
Named	r oceanica	um ypsilon	4	3 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2		E	Same (pale olive) with laminae of light olive and light gray Same, with 1-3 mm diffuse laminae of light gray				
Not Na	Gephyrocapsa	Amphirhopalum	5-1-1	4444			Same with mottles and burrows of pale greenish yellow Same, very light olive		ð		
			61111			e)	Gradational change to pale olive Layer of light olive Same, very pale olive, mottles and streaks of light gray Same, very light olive Same, very pale olive Same, very light olive		ø		я.
			7 8 8 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1				Same, very pale olive, streaks of light gray Same, very light olive, slightly mottled		C		

