# 9. SITE 178

# The Shipboard Scientific Party<sup>1</sup>

#### ABSTRACT

At Site 178, 777.5 meters of sediment and 17 meters of basalt were cored and drilled. The upper 270 meters is gray mud with ice-rafted erratics; glacial-marine sediments are as old as mid-Pliocene. From 270 meters to 742 meters there are interbedded muds, diatomaceous sediments, thin poorly sorted silts and sand turbidites. At 742 meters, a distinctive varicolored claystone and chalk separates the turbidites above from the underlying 29 meters of pelagic shale above basement. The section ranges in age from Pleistocene through earliest Miocene. The 20 m.y. discrepancy between the earliest Miocene age and the anomaly 20 basement age is not resolved; however, no microfossils were recovered in the 27 meters of shale above the basalt and the basement may occur below the basalt drilled. Rates of sedimentation based on diatom and radiolarian zonations in the section younger than 1 m.y. are are from 70 to 180 m/m.y. and between early Pleistocene and earliest Miocene they are about 30 m/m.y. (Sedimentation rates uncorrected for compaction).

# SITE SUMMARY

Date Occupied: 1-5 July 1971.

Position (Satellite): Latitude: 56°57.38'N; Longitude: 147°07.86'W.

Number of Holes: One

Water Depth: 4218 meters.

Penetration: 794.5 meters below sea floor.

Number of Cores: 59.

Total Core Recovered: 212.5 meters, 40.9%.

Age of Oldest Sediment: Early Miocene.

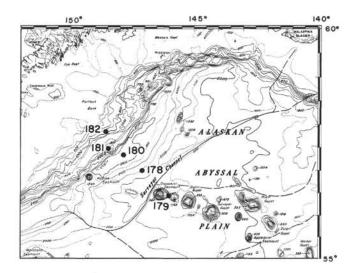
Acoustic Basement:

Depth: 777.5 meters. Nature: Basalt.

## BACKGROUND AND OBJECTIVES

#### Site Description

Site 178 is on the western Alaskan Abyssal Plain, 330 km from Kodiak Island at latitude 56°57.38'N and



longitude 147°07.86°W, in 4218 meters of water. The Aleutian Trench and Kodiak Seamount are west of the site, Surveyor Channel and Giacomini Seamount are on the east (Figure 1). Seismic reflection records indicate that a 800- to 1100-meter-thick section of turbidites and hemipelagic sediments occur at the site. This section extends from under the Aleutian Trench into the Gulf of Alaska and it is thought to be representative of the Alaskan Abyssal Plain stratigraphy.

The seismic records show single reflections and groups of reflections that can be traced more than 100 km. Lateral variation in layering is seen where moderate to small buried turbidity current channels disturb the otherwise parallel reflections. Vertically the record shows moderate variability from weak narrow reflectors to strong wide ones.

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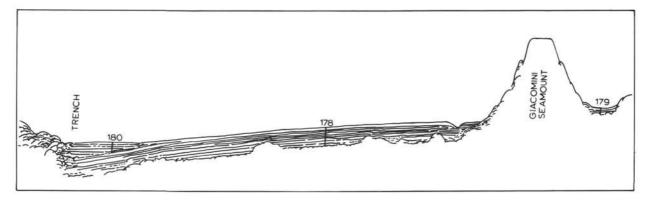


Figure 1. Diagrammatic section (based on reflection records) across the Aleutian Trench and adjacent abyssal plain showing setting of Sites 178 through 180.

# Site Objectives

The broad objectives in sampling at this site are to study tectonism of the continental margin of Alaska and investigate the geologic history of the Gulf of Alaska. The biostratigraphy of this section is needed to help establish Tertiary faunal and floral successions for high-latitude regions. The time span of glacially derived material at this site may indicate when strong glaciation began in Alaska. It is also possible that the ocean crust now in the Gulf of Alaska was once farther south and that it migrated north as new sea floor was generated (Atwater, 1970). This idea may at least be partially tested by studying the lithology and biostratigraphy at Site 178. It is also possible to date sediments filling the Aleutian Trench. The objectives of this site are closely tied to the similar objectives of the set of sites 178, 180 and 181.

## LITHOLOGIC SUMMARY

# **General Statement**

A 777.5 meter section of deep-sea sediments overlies olivine basalt at this site. The uppermost 280 meters consists mainly of thick bedded silty clay with sand- to pebble-sized metamorphic and sedimentary erratics.<sup>2</sup> Towards the base of this unit, the erratics decrease in number and silt laminae and beds increase in abundance. reaching a maximum of 6 per meter section. From 280 to 742 meters, erratics are absent and silty clay, diatomaceous sediments (diatom-bearing silty clays to diatom ooze) and thin bedded very poorly sorted silt and very fine sand turbidites are interbedded in varying proportions. At 742 meters, a readily identifiable marker unit exists; it consists of 6.5 meters of interbedded chalk and varicolored claystone (shades of brown and olive gray) with abundant, well-preserved burrows emphasized by color contrasts. Between this marker unit and the basalt, there are 29

meters of clay shales barren of microfossils. The contact between the basalt and the overlying clay shale is not baked, and the basalt is glassy and vesicular, indicating that the fine-grained sediments were laid down on submarine basalt flows. The following units are defined on the basis of different proportions of various sediment types.

#### **Lithologic Units**

#### Unit 1 (0-96 m; Cores 1-11)

This uppermost unit consists of very sticky, medium dark gray (N 4) silty clay, commonly with erratics (1 to 6 pebbles per meter exposed in the cut surface of the core). Erratics are especially abundant in two zones, from 24 to 33 meters and from 60 to 96 meters. They range in size up to 9-cm pebbles; the granules and pebbles are angular to subrounded with a few having recognizable striations of their surfaces. The lithologies include graywacke, sandstone, shale, and amphibolite. Laminae of silt occur very rarely. Diatoms are noticeably absent, although rare intervals of diatom-bearing silty clay and two layers of olive gray (5 Y 4/1) silty clay-rich diatom ooze (36 and 16 cm thick at 7 and 9 meters, respectively) are present. Several layers of light ash are visible (noted on summary core log); they are 1 to 14 cm thick, silt to fine sand sized, and occasionally graded.

#### Unit 2 (96-141 m; Cores 12-16)

This unit also consists of medium dark gray silty clay but the number of erratics decreases noticeably below 96 meters (less than 1 pebble per meter). Several diatombearing to diatom-rich silty clay intervals are present; they tend to be dark greenish gray (5 GY 4/1). Silty clay diatom oozes occur at 100 meters (8 cm thick), 106 meters (10 cm), 135 meters (5 cm) and 137 meters (215 cm). Silt laminae are rare.

## Unit 3 (141-280 m; Cores 17-29)

In this unit, the number of visible erratics is greater than in the second unit but low (1 pebble per meter). However, the abundance of silt layers increases in the silty clays which are olive gray (5 Y 4/1) and medium dark gray, and the diatoms decrease. The silt layers are generally 2 mm to 2 cm thick, although layers up to 12 cm thick occur rarely.

 $<sup>^{2}</sup>$ The relative abundances of erratics noted are based on visual inspection of cut core surfaces only. The abundances of erratics, based on X-radiography is explained in von Huene, Chapter 21. A more thorough discussion of silts is given in Chapter 23, this volume.

On the average, 1 to 6 layers are present per meter of core, with the abundance increasing with depth in the subunit. The silt layers frequently have sharp bases and grade upward into silty clays through a zone of interlaminated silt and silty clay. Concentrations of hornblende at the bases of silt beds are present in the interval between 175 and 180 meters. A few layers of light ash, 1 to 5 cm thick, occur, and two layers of very fine sand sized brown glass are visible at 152 meters (50 cm thick) and 224 meters (3 cm thick).

#### Unit 4 (280-357 m; Cores 30-37)

This unit consists of predominantly diatomaceous sediments (diatom-bearing silty clay to diatom ooze) with interbeds of silty clay and thin, poorly sorted silt to very fine sand turbidites. The silty clayey sediments are olive gray, dark greenish gray, and greenish gray (5 GY 5/1), with those sediments having higher concentrations of diatoms frequently tending towards greenish gray. Contacts between fine sediments of varying diatom concentration are not well defined. The sand or silt intervals of the turbidites generally are 2 to 15 cm thick but can be as thick as 55 cm. They comprise about 5 to 10 per cent of the upper part of the unit and generally increase down section to a maximum of 30 per cent. The poorly sorted nature of these turbidites is due more to admixed medium and fine silt than to clay. Sharp bases and grading are evident in many of the turbidites, and, occasionally, the diatom abundances increase upwards in the silty clays overlying sandy bases.

#### Unit 5 (357-410±10 m; Cores 37-39)

This unit consists predominantly of olive gray to greenish gray silty clay and diatom-bearing silty clay to diatom ooze; sandy or silty turbidites are generally absent. Occasional horizontal burrows, *Zoophycus*, are revealed by color contrasts in the silty clays. From 392 to 399 meters, there is an unusual occurrence of thin interbeds of olive gray fine silt and medium dark gray silty clay. The silt layers have a modal size of 6 to 8 microns, are composed dominantly of quartz with some carbonate, are well sorted and occur in laminae 2 mm to 2 cm thick, and are separated by 2- to 15-cm-thick silty clay intervals with no biogenous components.

#### Unit 6 (410±10 - 456 m; Cores 40-43)

This unit is characterized by thin, poorly sorted silt and very fine sand turbidites interbedded with olive gray and dark greenish gray silty clays. The sand-silt basal portions of the turbidites are up to 25 cm thick and may comprise around 35 per cent of the section. Well-preserved, pebblesized mud clasts are present above a basal sand in one turbidite at 447 meters. Some of the sand-silt beds may be cemented by calcium carbonate as indicated by one such bed recovered in the core catcher at 437 meters.

### Unit 7 (456-505±10 m; Cores 44-46)

This unit is characterized by olive gray to dark greenish gray silty clay with intervals of greenish gray clayey diatom ooze, clay-bearing diatom ooze, or diatom ooze, and rare beds of very fine sand.

## Unit 8 (505±10 - 560 m; Cores 47-48)

This unit consists predominantly of poorly sorted silt and very fine sand turbidites with the sand-silt intervals comprising a total of 30 to 50 per cent of the section and ranging in thickness from 1 to 40 cm. The silty clay intervals (pelagic and turbidite tails) are dark greenish gray to olive gray. Carbonate-cemented siltstone and sandstone were encountered at 553 meters.

#### Unit 9 (560-630 m; Core 49)

Turbidites are absent and olive gray and dark greenish gray silty clay and diatomaceous sediments characterize the interval between 615 and 645 meters. A 150-cm-thick greenish gray diatom ooze bed extends from 631 to 632.5 meters.

# Unit 10 (630-688 m; Cores 50-52)

This section consists of thin-bedded, poorly sorted silt and very fine sand turbidites and olive gray to dark greenish gray silty clays with rare diatomaceous sediments. Calcareous-cemented, poorly sorted fine sandstone (25 cm thick) is present at 658 meters.

#### Unit 11 (688-742 m; Cores 52-53)

Hard, olive gray and dark greenish gray silty clays with no or only trace biogenous components characterize this lowermost unit.

# Unit 12 (742-748.5 m; Core 54)

A readily identifiable unit occurs between 742 and 748.5 meters; it consists of interbedded chalk and varicolored claystone. Four chalk intervals (25 to 60 cm thick) are present, and they vary from light olive gray (5 Y 6/2) to light greenish gray (5 GY 7/1) to grayish orange pink (5 YR 7/2). They are separated by claystone intervals (25 to 300 cm thick) that range from dark yellowish brown (10 YR 4/2) to moderate brown (5 YR 4/4) to to light olive gray (5 Y 6/2) and pale yellowish brown (10 Y 6/2). Some zones of the claystones contain as much as 20 per cent diatoms and radiolarians, while others are barren. The entire unit is mottled by burrowing, with most of the burrowing being horizontal (including *Zoophycus*).

#### Unit 13 (748.5-777.5 m; Cores 55-57)

A 31-meter-thick olive gray, dusky yellowish brown (10 YR 2/2) and dark yellowish brown (10 YR 4/2) clay shale unit occurs just above the basalt basement. It is barren of biogenous materials.

#### Unit 14 (777.5-794.5 m; Cores 58-59)

Olivine basalt in contact with the overlying olive gray shale was recovered at 777.5 meters. The basalt continues downwards to at least 794.5 meters, at which point the drilling was terminated.

#### PETROGRAPHY OF THE BASALT

Hole 178 penetrated 17 meters of basalt of which 40 cm were recovered. In hand specimen the basalt core is dark gray and generally unaltered. In thin section, the rock ranges from glass with scattered phenocrysts of olivine to rather coarse-grained holocrystalline basalt. Between the two textural extremes, the rock shows progressive increase in crystallinity and, particularly interesting, shows a feathery texture of the olivine.

The glass phase of the rock is a translucent brown to nearly opaque. Olivine phenocrysts range from few to numerous in the thin sections, but have a very irregular distribution. Most are euhedral but occasional phenocrysts show the start of additional terminal growth.

The intermediate stage of crystallization consists of small plagioclase laths and a very fine grained ground mass containing feathery olivine. Occasional phenocrysts and crystals of olivine can be seen to form nuclei from which the feathery lattice structure has grown. A different aspect of the growth consists of a network of acicular crystals up to 2 mm in length that dominate the slide. By extrapolation through a series of slides, the olivine nature of all the acicular and feathery forms can be demonstrated.

Coarse-grained variations of the basalt have a diabasic sort of texture with laths of plagioclase, interstitial pyroxene and much chloritized groundmass. In a few coarsegrained sections, relict phenocrysts of olivine have altered to a mixture of serpentine, carbonate, and iddingsite. Several of the coarsest-grained rocks have no indication of olivine.

The core section is incomplete and rather broken up; samples from both of the cores lack a consistent pattern of crystallinity. In fact, the glassiest samples are from the lower part of the section. It is, therefore, probable that the cored section was through a series of pillows on the sea floor or multiple thin surface flows.

#### PALEONTOLOGIC SUMMARY

# Introduction

Calcareous microfossils are rare or absent in post-Lower Miocene sediments encountered from 0 to 720 meters (Cores 0 to 53) at Site 178 due to deposition below the calcium carbonate compensation level (CCL). In contrast, siliceous microfossils are present throughout the entire Pleistocene through Lower Miocene section with diatomrich muds and diatomites predominant within the 287- to 742-meter interval (Cores 30 to 53). A striking coccolithrich chalk<sup>3</sup> containing Lower Miocene foraminifera occurs in Core 54 (742 to 749 meters) and appears to represent the terminal phase of a regionally significant lithologic event produced by a lowered CCL and/or prolific production of calcareous nannoplankton.

The base of the lower Pleistocene *Eucyrtidium* matuyamai radiolarian zone could not be precisely located at Site 178 although it occurs at the base of or below Core 22 (195 m). Apparent Pliocene radiolarian faunas occur below Core 22. The Pliocene-Pleistocene boundary is arbitrarily located at the base of Core 22 (195 meters) on the basis of this tentative radiolarian zonation; diatom and planktonic foraminiferal zonation tend to support this placement of the boundary.<sup>4</sup> North Pacific Diatom (NPD) Zone V, thought to correlate in part with the base of the Olduvai geomagnetic event, spans Cores 18 through 28 (152 to 243 meters). In addition, *Globigerina pachyderma* exhibits a change from sinistral to dextral coiling between Cores 14 to 18 (123 to 159 meters) representing a major increase in surface temperature thought to correlate with the lower portion of Blow's (1969) basal Pleistocene zone N22.

Pliocene radiolarian species are present as low as Core 37 (360 meters). Radiolarians do not provide good stratigraphic control below Core 22. Pliocene planktonic foraminifera are present through Core 29 (255 meters).

Pliocene diatom floras encompassing NPD Zones V through X are recognized from Core 28 through Core 52 (690.5 meters) with an Upper Miocene flora (NPD Zone XIV) recognized in Core 53.

Well preserved planktonic foraminfera of Lower Miocene age are present in the multicolored chalk encountered in Core 54 (742 to 749 meters) and are correlated with Blow's (1969) zone N5 and the lower portion of zone N6. Beggren (1972) currently correlates the base of zone N5 with an estimated radiometric age of 21 m.y. and the top of this zone with an age of 19.5 m.y. A limited coccolith flora in Core 54 supports the Lower Miocene age of this unit.

#### Calcareous Nannofossils

Most of the sediment examined from Site 178 was deposited below the calcium carbonate compensation level; therefore, few calcareous nannofossils were found in Cores 1 through 50 (0 to 638.5 meters). A few nondiagnostic species such as Coccolithus pelagicus were noted in isolated samples and these are assumed to have been deposited in association with turbidites. The Core 51 core catcher (667 meters) contains reworked Oligocene-Miocene coccoliths and discoasters deposited within a gray turbidite silt. Core 54 (742 to 749 meters) recovered a multicolored chalk composed almost exclusively of broken or isolated rims derived from Coccolithus miopelagicus, Coccolithus sp. cf. pliopelagicus, and Cyclicargolithus bukryi Wise n. sp. Discoaster deflandrei, including forms heavily overgrown by secondary calcite, is also common in some samples of this chalk. The calcite overgrowths on the discoasters are similar to those described at Site 172 which indicates that the nannofossil ooze deposited at the Site 178 underwent some amount of early diagenesis and cementation prior to deep burial under 742 meters of clastic sediment. This would explain the surprising absence of compaction features in the chalk. Zoophycus burrows show little or no flattening and delicate features such as partially etched but intact coccosphaeres show no distortion at all. Apparently, the nannofossil ooze was sufficiently well lithified during early cementation to prevent significant compaction during the subsequent loading of hundreds of meters of clastic sediment.

The nannofossil assemblage from the chalk is highly restricted to a few cold-water species including forms not

<sup>&</sup>lt;sup>3</sup>See Wise, Chapter 15, this volume, for a detailed discussion of chalk ultrastructure and origin.

<sup>&</sup>lt;sup>4</sup>See Chapter 21 for 1.8 m.y. paleomagnetic boundary.

found in warmer waters. Coccolithus pliopelagicus closely resembles a cold-water member of the Coccolithus sp. cf. pelagicus lineage. Cyclicargolithus floridanus. The cosmopolitan Discoaster deflandrei is the only astrolith present. As a rule, Discoaster is a warm-water genus with well over a hundred species reported from low and middle latitudes. The relatively small numbers of Discoaster deflandrei present in Core 54 and those reworked into higher strata are the only Neogene discoasters reported to date from the Gulf of Alaska.

Planktonic foraminifera indicate that the chalk is Early Miocene in age. The calcareous nannofossils are considered to be Early Miocene as well, certainly not younger than Middle Middle Miocene. Diatoms extend down to Section 3 in Core 54 and these are classified Late Miocene in age; perhaps these have been reworked downward by burrowing organisms. Above the chalk, siliceous organisms are the dominant constituent of the planktonic fauna and flora. As evidenced by the records at Sites 177 and 178, the cold-water-restricted calcareous nannoplankton floras made only minor contributions to the post-Lower Miocene fossil record in northeast Pacific during the remainder of the Neogene. This has resulted in the breakdown of the "standard" Neogene calcareous nannofossil zonation based on their occurrence (summarized by Martini and Worsley, 1970).

The occurrence of nannofossil chalk at this high latitude (56°N) and at such great depth (4960 meters below sea level) is somewhat remarkable. Although tectonics could be invoked to explain this phenomenon, it is noteworthy that the chalk is underlain by pelagic clays rather than turbidites. Thus, it is most probable that the chalk represents a sharp downward excursion of the carbonate compensation level produced by an exceptionally high production of calcareous phytoplankton. These supplied large amounts of skeletal carbonate to the sea floor and increased the sedimentation rate sufficiently to depress the carbonate compensation level several hundreds of meters. It has previously been noted (Tracy et al., 1971) that in the equatorial Pacific, the Early Miocene was a time of widespread carbonate deposition. Perhaps favorable current conditions allowed the incursion of carbonates into the Gulf of Alaska during this time.

#### Diatoms

Varying abundances of diatoms in good to poor states of preservation were found throughout the sequence penetrated at Site 178. The interval from 0 to 740 meters (Cores 1 to 53) represents a continuous record of high diatom productivity from the Holocene through the Upper Miocene interrupted only by intermittent influxes of terrigenous sand, silt, and mud turbidites. The amount of older reworked microfossils is negligible.

The base of Pleistocene NPD Zone I is placed at 46 meters (178-6-3, 130-132cm), the base of NPD Zone II at 125 meters (178-15-2, 36-38cm), and the base of combined NPD zones II and IV at 152 meters (178-18-2, 45-47cm). The base of Pliocene NPD Zone V is placed at 243 meters (178-28-4, 5-7cm), the base of NPD Zone VII is tentatively placed at 360 meters (178-37-4, 133-135cm), and the base of NPD VIII is tentatively placed at 465 meters (178-46-6,

121-123cm). The base of Pliocene NPD Zone IX and X is tentatively placed at 690.5 meters (178-52, CC).

Upper Miocene NPD Zone XIV was recognized in Core 53 at Site 178 but the base of this zone could not be defined. The oldest diatom-bearing sediment recovered at Site 178 is placed within NPD Zone XIV and occurs at 746 meters (178-54-3, 126-127cm).

The preservation of diatom frustules is generally poor due to dissolution from 289 to 746 meters (178-30-2, 62-64cm to 178-54-3, 126-127cm) with only the more robust tests been preserved. Thus, *Coscinodiscus marginatus, C. marginatus* forma *fossilis*, and *C. symbolophorus* occur in abundance within this interval whereas weakly silicified diatoms including various species of *Denticula*, *Thalassiosira*, and *Nitzschia* have been dissolved.

#### **Planktonic Foraminifera**

Foraminifera are absent or rare in most of the sediments recovered at Site 178. Paucity of foraminifera and other calcareous microfossils apparently reflects continuous deposition below the calcium carbonate compensation level during most of the interval represented. Rare and occasionally common specimens of poorly preserved foraminifera were found in Cores 1 to 14 (0 to 123 meters) and in Cores 18 to 29 (150 to 270 meters) within a sequence characterized by muds, silts, and glacial erratics. Relatively rapid burial by terrigenous clastics apparently provided minimal protection from aggressive bottom water during this interval. Predominantly diatomaceous sediments recovered from 270 to 742 meters (cores 30 to 53) are barren of foraminifera with the exception of rare benthic specimens in Cores 31, 33, and 34 reflecting the slower rate of sedimentation and consequent exposure of tests during this interval. Common and well-preserved foraminifera are present in Core 54 (742 to 749 meters) within a multicolored coccolith chalk. A clay shale encountered from 749 to 778 meters (cores 54 to 58) proved to be barren of all microfossils.

Planktonic foraminiferal faunas encountered from 0 to 123 meters (Cores 1 to 14) are dominated by sinistral coiling populations of Globigerina pachyderma representing mean surface temperatures lower than 5°C and are correlated with Pleistocene zones N23 and N22. Dextral coiling populations of this same species dominate faunas from 150 to 225 meters (cores 18 to 29) and probably represent surface temperatures higher than 15°C within the Alaskan Current gyre. Correlations with Sites 173 and 174 suggest this latter warm event encompasses the later Pliocene and early Pleistocene equivalent to zones N21 and N22. Globigerina bulloides and G. quinqueloba constitute common accessory species within the Plio-Pleistocene interval with one unusual assemblage at 242 meters (178-28-3, 65-67cm) characterized by an abundance of Globorotalia scitula.

Planktonic foraminifera recovered from the chalk in Core 54 (742 to 749 meters) include Catapsydrax dissimilis, C. stainforthi, C. unacava s.l., Globorotalia obesa, G. opima nana, G. praescitula, and Globoquadrina cf. praedehiscens indicating this important lithologic horizon is Lower Miocene in age equivalent to Blow's (1969) zone N5 through perhaps the lower portion of Zone N6.

# **Benthonic Foraminifera**

Sparse calcareous and arenaceous benthonic foraminifera occur at scattered horizons within the Plio-Pleistocene clastic sediments from 0 to 270 meters (Cores 1 to 29). These faunas are composed of species common to modern high-latitude lower bathyal-abyssal faunas and include *Cibicides wuellerstorfi, Gyroidina soldani, Melonis pompilioides, Pulleria bulloides,* and *Uvigerina auberiana.* Admixtures of species displaced from shallower water include common specimens of *Eilohedra levicula, Epistominella exigua,* and *Globobulimina affinis* with graded intervals containing rare abraded specimens of neritic species such as *Elphidium clavatum.* 

Benthonic foraminifera are common and well preserved within the lower Miocene chalk of Core 54 (749 to 778 meters). These faunas are relatively diverse but retain a lower bathyal character with common specimens of *Cibicides kullenbergi*, *Nuttallides umbonifera*, *Pleurostomella elliptica*, *Pullenia bulloides*, and several species of *Stilostomella*.

## Radiolaria

Cores 1 through 22 (0 to 195 meters) contain reasonably rich radiolarian faunas. Cores 23 to 53 (195 to 720 meters) contain only sporadic radiolarian occurrences with few age-diagnostic species.

Cores 1 through 8 (0 to 69 meters) contain assemblages representing the upper Pleistocene-Holocene Artostrobium miralestense Zone of Hays (1970). The upper limit of Druppatractus acquilonius occurs at 51 to 60 meters (178-6, CC to 178-7, CC) within this zone and is correlative with an estimated radiometric age of 0.3 m.y. Cores 9 through 15 (69 to 125 meters) represent Axoprunum angelinum zone; the top of this zone (178-8, CC to 178-9-1, 136-138 cm) has an estimated radiometric age of 0.4 m.y. whereas the base of the zone has an age of 0.9 m.y. Cores 15 through 22 (125 to 196 meters) represent the lower Pleistocene Eucyrtidium mutuyamai zone, but its base cannot be recognized with certainty. Below Core 22 (195 meters) sporadic specimens of Lamprocyrtis heteroporos indicate Pliocene ages through at least Core 37 (363 meters). Cores 52 (686.5 to 690.5 meters) and 54 (742 to 749 meters) contain Miocene species that are not diagnostic of any specific zones.

#### Spores and Pollen

Samples examined from Site 178 generally contain few or no palynomorphs. Pollen occurring in amounts worthy of note (Cores 24 and 42) is so badly preserved and oxidized that specific identification is impossible. Moreover, identifiable grains (*Tsuga, Alnus, Salix, Pinus*, psilate and trilete spores) were present in such small quantities that no paleoclimatological determinations could be made.

Coaly fragments occur commonly throughout the cores. When samples were high in diatomaceous material, no palynomorphs or coal fragments were observed. This is certainly the result of dilution by terrigenous sediment and by rapid deposition of organic silica.

# PHYSICAL PROPERTIES

Physical properties were measured at Site 178 in the standard manner as on previous sites. Time was available to measure physical properties on more sections than previously because of long core retrieval time and a faster GRAPE carriage speed.

GRAPE porosity shows some good correlations with lithology. For instance, hard fissile muds show a marked increase in porosity as do ash beds and diatom-rich sediments. Clean sands also correlate with a well-defined porosity increase, but when dirty, the sands produce a good low-porosity peak in the GRAPE records. Larger ice-rafted erratics are easily spotted. However, to be meaningful, the GRAPE records must be interpreted with care because almost all of the above effects can be confused in the records with effects from drilling disturbance. Another effect that invalidates the absolute GRAPE porosity value is the small core diameter resulting when harder rock is cored. Porosity and density values for the shales, sandstones, and limev mudstones (all harder rocks) from this site are invalid in an absolute sense but are considered very good in a relative sense.

In the upper 100 meters of core, porosity values generally range between 45 and 60 percent. Fairly pure mud shows a GRAPE variation of only 5 percent and with ash and silt admixtures, porosities range as much as 20 percent. The range in porosity increases somewhat between 100 and 300 meters correlating with the increasing frequency and amount of the silts, ashes and diatomaceous sediments recognized visually. Below 300 meters, large ranges of GRAPE values indicate large variation in sediment type or a change from a more monolithologic sequence to significantly greater variation of beds with a large range in porosity (as great as 50 percent in one section). One of the greatest ranges in values results when a diatomaceous bed is found in the same section with a sand.

Sonic velocity measurements were made on the least disturbed and most consolidated material in one or more sections of all but two cores. The material was taken out of the core liner for measurement. Three measurements were made on each sample and, on some shales, measurements were made parallel and perpendicular to bedding planes.

#### CORRELATION BETWEEN REFLECTION RECORDS AND THE STRATIGRAPHIC COLUMN

Site 178 was selected from seismic records made in 1970 by the USGS (von Huene, 1972). Because of their good quality, these records were relied on most for a comparison with lithologies in the drill hole and to extrapolate lateral variations in specific lithologic intervals (Figure 2). However, all depths are calculated using travel times picked from the on-station record because the site may be located as much as 2 km from the ship's track along which the USGS seismic record was made.

Velocities used in calculating depths from travel times are a composite of those reported in the previous section and sonobuoy velocities (von Heune, 1972). Because of poor velocity control in the lower part of the hole due to interval coring, poor recovery, and wide ranges in velocity values, the seismic and lithologic sections are not correlated with as great a certainty as in the upper part of the hole. In some instances, such as Unit 4, changes in reflective character that approximate lithologic units were relied on most.

Table 1 summarizes the correlation of seismic information and lithology. Five records (two USGS, two *Challenger*, and one University of Washington) were available



Figure 2. Seismic record (U.S.G.S.) near Site 178 showing lithologic and acoustic character is best in the multiple for Units 1-4 and better in the primary for Units 5-13.

Unit	Cores	Interval (sec)	Velocity (km/sec)	Depths (m)	Lithologic Description	Acoustic Character	Thickness	Interpretation of Lithologic and Acoustic Information
1	1-11	0-0.13	(1.5)	0-96	Mud, dark gray, sticky, few to abundant erratics.	Sharp, fine, unevenly spaced, reflections with variable ampli- tude; stronger reflec- tion at 0.07 sec (53 m)	96 m	Massive continuous beds of gray mud, with variable amounts of erratics broken by thin diatom-rich muds and ash probabl with a very continuous 5 to 10-m-thick diatom-rich mud bed or sand in the middle of the unit. The unit was probably deposited during the most intense period of glaciation (0.7 t 0.9 m.y. long).
2	12-16	0.13-0.19	(1.5)	96-141	As Unit 1 with more diatom-rich intervals and fewer erratics.	A "translucent" unit with weak, fuzzy, non-persistent, low- frequency reflections.	45 m	Same as above with more frequent diatom-rich mud interval that may have gradational contacts and less distinct bedding (May indicate slightly less intense glacial environment.)
3	17-29	0.19-0.37	(1.5)	141-280	As Unit 1 with more silt and fewer diatom-rich intervals.	Generally weaker low- frequency but moder- ately continuous reflections that are stronger; unit at same depth as a channel.	139 m	Same as 96 m with more frequent continuous silt beds. Silt below 182 m probably derived from last overbank activity of a tributary of Surveyor Channel system about 1.5 to 2.0 m.y ago. The bottom of this unit probably received the first erratics from occasional times of heavy sea ice and a large contribution of finely ground mountain-valley glacial debris to give muds their gray color.
4	30-37	0.37-0.45	(2.0)	280-357	Mud interbedded with diatom-rich mud, diatom ooze, and fine sand to silt turbidites.	Strong continuous high-frequency reflec- tions and low-frequency reflections grading into Unit 5.	77 m	Irregular lensing and pinching beds of mud interbeds with diatom-rich intervals and silt and sand deposits from either a single large channel to the west or a series of smaller shift- ing channels.
5	37-39	0.45-0.51	(1.7)	357-410±10	As in 4 without silt and fine sand turbidites.	As in 4, with weaker reflectivity on <i>Slave</i> and and <i>Challenger</i> record, but stronger on primary.	53 m 1	Massive beds of mud. Probably a period of lower Pliocene abyssal conditions without effects of deep-sea turbidity current channels.
6 & 7	39-46	0.51-0.63	(1.7)	410±10-505±10	Mud with interbedded silt and fine sand tur- bidites grading to diatom-rich intervals in lower part.	Strong low-frequency reflections in upper part separated from weaker reflections in lower part by a single strong reflection on <i>Challenger</i> record, consistently strong low-frequency on U.S.G.S. record.	95 m	Mudstone beds broken by silts and some fine sand beds. Lower part broken by beds of diatom oozes. Abyssal condi- tion with some turbidity current action. First ashes indicate beginning of present volcanic period on the Alaska Peninsula
8	47-48	0.63-0.70	(1.6)	505±10-560	Mainly silt and sand turbidites with some mud.	Weak low-frequency reflections, now per- sistent, that pinch out toward trench.	55 m	A thick unit of sand and silt turbidites with breaks of mud. This unit is derived from a channel to the east pinching to one-half its thickness in 100 km.

 TABLE 1

 DSDP Site 178 – Acoustic and Lithologic Summary

Thick sand and silt interbedded with mud	Generally hard mud and silt with a possible silt to sand bed near the middle of the unit. Bedding appears very irregular at drill site but on the other side of a basement hump to the west the beds at the same depth are very well stratified. Site 178 may be in an unusual basement area, and the configuration of the basalt flow may cause unique lithology.	Continuous beds with a rough upper surface indicating pelagic deposition.
70 m	110 m	
Very strong continuous low-frequency reflec- tions becoming wider toward trench.	Weak to moderate dis- continuous low- frequency reflections showing very indis- tinct signs of stratification.	Very strong low- frequency reflections with a somewhat irregular upper surface.
Silt and very fine sand turbidites with mud intervals.	Variable sequence of sediments from mud to silt and sand.	Chalk underlain by claystone.
560-630	630-742	742
(2.0)	(2.7)	
0.70-0.76	10 & 11 50-53 0.76-0.84 (2.7)	
49	50-53	54-57
6	10 & 11	12 & 13 54-57

for interpretation and each record provided some information not available in the others. Figure 2 is a USGS record that shows most of the main acoustically determined features.

## SUMMARY AND CONCLUSIONS

Site 178 is on the Alaskan Abyssal Plain about midway between Surveyor Channel and the Aleutian Trench. The abyssal sediment section is often more than 800 meters thick. The age of basement at this site, based on magnetic anomaly 20 (Pitman and Hays, 1967), is about 50 m.y. In seismic records, the layered section can be followed under the Aleutian Trench and in some instances seen under the continental slope (von Huene, 1972). The configuration of reflections suggests that the section consists of turbidite layers interbedded with some hemipelagic and pelagic layers.

This site was one of three chosen to study tectonic mechanisms operating along the continental margin of Alaska and to investigate the geologic history of the Gulf of Alaska. The site was also drilled in an attempt to establish the biostratigraphy of siliceous microfossils at this latitude.

Core recovery was poor because of a malfunction in the bottom-hole assembly. Accordingly, the stratigraphic section is not well known, especially below 360 meters where continuous coring was terminated and interval coring was started. The hole was drilled to a total depth of 794.5 meters which includes about 17 meters of coring in basalt.

Sediments at Site 178 may be grouped into three main lithologic sequences with each sequence suggesting a pronounced change in environment. The upper 270 meters consist of dark and medium gray muds with ice-rafted erratics. The erratics are most abundant in the first 100 meters, and decrease in the lower part of the unit where the number of silt beds increase. From 270 to 742 meters, muds, diatomaceous sediments, and thin, very poorly sorted silt and very fine sand turbidites are interbedded in varying proportions. At 742 meters, a distinctive varicolored mottled claystone and chalk was recovered. Between this horizon and the basalt basement, there are 27 meters of shale composed of pelagic clay.

The lower sequence indicates a typical abyssal environment and a sudden short interception of the bottom by the calcium carbonate compensation level. This lower Miocene depression of the compensation level may have been synchronous with major uplift of the Chugach-St. Elias Mountains. This may also have been a time when the Kula Ridge was being subducted (Atwater 1970). The intermediate sequence marks a change from the previous pelagic and hemipelagic sedimentation to turbidite sedimentation. It is a time of abyssal turbidity current channels broken by short periods of quiescence as marked by zones rich in diatoms. Continued uplift of the mountains fringing the Gulf of Alaska and short periods of mountain-valley glaciation may have provided changing types of sediment. An increasing proximity to land by westward plate movement may also be reflected in the lithology. The first volcanic ash, possibly from the Alaska Peninsula, is in the upper Miocene section. The upper sequence reflects strong glaciation and the uppermost 100 meters indicate exceedingly rapid denudation on land.

The sedimentary section ranges in age from Pleistocene through early Miocene. Calcareous microfossils are absent through much of it because water depths are below the calcium carbonate compensation zone. However, siliceous microfossils are present in most of the cores with diatom floras being locally abundant. Microfossil assemblages date the upper gray muds containing erratics as Pleistocene and the glacial erratics and gray mud continue about halfway through the Pliocene section. The distinctive varicolored chalk and mudstone interval in Core 54 contains lowermost Miocene nannofossils, planktonic foraminifers, radiolarians, and diatoms. Unfortunately, the 29-meter interval between the varicolored horizon and basement proved barren of microfossils.

On the basis of the diatom ages, rates of sedimentation in the Pleistocene section vary between 170 and 180 m/m.y. assuming the beginning of Pleistocene time at the 1.8 to 2.0 m.y. boundary in Core 23. From Core 28 to Core 54, the rate is 35 m/m.y. assuming a 20 m.y. age at the base of Core 54.<sup>5</sup>

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## APPENDIX A. OPERATIONS

## Pre-drilling Site Survey

Site 178 was approached from the east on a course of 277° T navigated by Loran A. About 6 hours prior to arrival, it was established that *Challenger* was 5 miles north of the proposed track line. Rather than correcting for the

northerly set, *Challenger* remained on course to a point due north of the site and then came to  $180^{\circ}$  T and 6 knots to make a seismic record crossing the site selection record at right angles. The *Challenger* record showed that no turbidity current channels or faults hindered the objectives for Site 178 and upon reaching the plotted position the geophysical gear was brought aboard, the ship stopped, and the beacon dropped.

# **Drilling Program**

Site 178 is located in 4218 meters of water on the flat Alaskan Abyssal Plain. The water depth was calculated from Matthews Tables (Area 44 was used because Areas 45 and 24 were not available) and the "Hawaii factor" (i.e., 4198 + 14 = 4212 meters to the transducer). The vessel was positioned directly over the beacon.

Site 178 was continuously cored to a depth of 258.5 meters below the sea floor and cored at selected intervals between this depth to the total depth of 794.5 meters. Thin hard sandstone layers encountered at various levels below 427.5 meters did not affect the drilling and several good sandstone cores were recovered. Recovery was poor below the 7-meter-thick varicolored mudstone in Core 54 (749 meters) because of the extreme fissility of the shale encountered. Basalt occurred at 777.5 meters; a total of 3 meters was obtained in Cores 58 and 59.

Hole conditions at Site 178 were excellent largely due to the high percentage of mud in the sedimentary column. The weather was excellent, with calm seas and no currents.

The recovery at Site 178 was 40.9 percent. The unusually low recovery was probably caused by a malfunction in the bottom-hole assembly causing the core barrel to turn.

#### **Drilling Specifications**

This hole was cored and drilled from 4229 meters to 5023.5 meters for a total sub-mudline penetration of 794.5 meters. Of the 794.5 meters penetrated, 275 meters were drilled without coring and 519.5 meters were cored with 212.5 meters of core recovered (see Table 2 for the coring summary). The recovery was 40.9 percent. The drilling assembly was the same as used on Holes 177 and 177A; a Smith 4 cone button bit was run.

Hole conditions were very good, positioning was good, and the weather was excellent, but recovery of cores was a problem. The 40.9 percent recovery was marginal for the scientific objectives and disappointing to the operational people. After the drilling assembly was pulled out of the hole, it was discovered that the support housing was sanded up which apparently caused the inner barrel to rotate.

Hole 178 was abandoned and filled with 240 barrels of 10.2 p.p.g. mud.

# Post-drilling Site Survey

No post-drilling survey was run.

<sup>&</sup>lt;sup>5</sup> All rates are uncorrected for compaction.

	Cored Inter	val Below		Reco	overed
Core	Derrick Floor (m)	Sea Floor (m)	Cored (m)	(m)	(%)
1 2 3 4 5	4229.04235.0 4235.04244.0 4244.04253.0 4253.04262.0 4262.04271.0	0-6 6-15 15-24 24-33 33-42	6 9 9 9	6.0 4.0 5.0 6.0 5.5	100.0 44.4 55.6 66.7 61.1
6 7 8 9 10	4271.0-4280.0 4280.0-4289.0 4289.0-4298.0 4298.0-4307.0 4307.0-4316.0	42-51 51-60 60-69 69-78 78-87	9 9 9 9	4.5 CC 5.5 5.0 4.0	50.0  61.1 55.6 44.4
11 12 13 14 15	4316.0-4325.0 4325.0-4334.0 4334.0-4343.0 4343.0-4352.0 4352.0-4361.0	87-96 96-105 105-114 114-123 123-132	9 9 9 9	2.5 3.5 3.5 0.5 3.0	27.8 38.9 38.9 5.6 33.3
16 17 18 19 20	4316.0-4325.0 4370.0-4379.0 4379.0-4388.0 4388.0-4397.0 4397.0-4406.0	132-141 141-150 150-159 159-168 168-177	9 9 9 9	4.5 7.0 3.5 4.0 8.0	50.0 77.8 38.9 44.4 88.9
21 22 23 24 25	4406.0-4415.0 4415.0-4424.0 4424.0-4433.0 4433.0-4440.0 4440.0-4449.5	177-186 186-195 195-204 204-211 211.0-220.5	9 9 9 7 9.5	5.0 1.5 3.5 4.0 7.5	55.6 16.7 38.9 57.1 78.9
26 27 28 29 30	4449.5-4459.0 4459.0-4468.5 4468.5-4478.0 4478.0-4487.5 4516.0-4525.5	220.5-230.0 230.0-239.5 239.5-249.0 249.0-258.5 287.0-296.5	9.5 9.5 9.5 9.5 9.5	3.5 1.5 9.0 6.0 1.5	36.8 15.8 94.7 63.1 15.8
31 32 33 34 35	4525.5-4535.0 4535.0-4544.5 4544.5-4554.0 4554.0-4563.5 4563.5-4573.0	296.5-306.0 306.0-315.5 315.5-325.0 325.0-334.5 334.5-344.0	9.5 9.5 9.5 9.5 9.5	2.0 4.5 7.5 9.0 2.0	21.1 47.4 78.9 94.7 21.1
36 37 38 39 40	4573.0-4582.5 4582.5-4592.0 4611.0-4620.5 4620.5-4630.0 4655.5-4656.5	344.0-353.5 353.5-363.0 382.0-391.5 391.5-401.0 426.5-427.5	9.5 9.5 9.5 9.5 1.0	2.5 6.0 1.5 7.5 CC	26.3 63.2 15.8 78.9 -
41 42 43 44 45	4656.5-4666.0 4666.0-4675.5 4675.5-4685.0 4685.0-4694.5 4694.5-4704.0	427.5-437.0 437.0-446.5 446.5-456.0 456.0-465.5 465.5-475.0	9.5 9.5 9.5 9.5 9.5	CC CC 1.5 8.5 1.5	- 15.8 89.5 15.8
46 47 48 49 50	4725.0-4734.5 4734.5-4744.0 4763.0-4772.5 4791.5-4820.0 4848.5-4858.0	496.0-505.5 505.5-515.0 534.0-543.5 591.0-600.5 629.0-638.5	9.5 9.5 9.5 9.5 9.5	2.0 2.5 1.5 1.5 4.5	21.1 26.3 15.8 15.8 47.4
51 52 53 54 55	4886.5-4896.0 4915.0-4919.0 4945.0-4949.0 4971.0-4978.0 4978.0-4987.5	657.5-667.0 686.0-690.0 716.0-720.0 742.0-749.0 749.0-758.5	9.5 4.0 4.0 7.0 9.5	2.5 2.5 2.5 7.0	26.3 62.5 62.5 100.0
56 57 58 59	4987.5-4997.0 4997.0-5006.5 5006.5-5016.0 5016.0-5023.5	758.5-768.0 768.0-777.5 777.5-787.0 787.0-794.5	9.5 9.5 9.5 7.5	CC 0.5 1.5 1.5	- 5.3 15.8 20.0
		Total	519.5	212.5	

TABLE 2 DSDP Site 178 Coring Summary

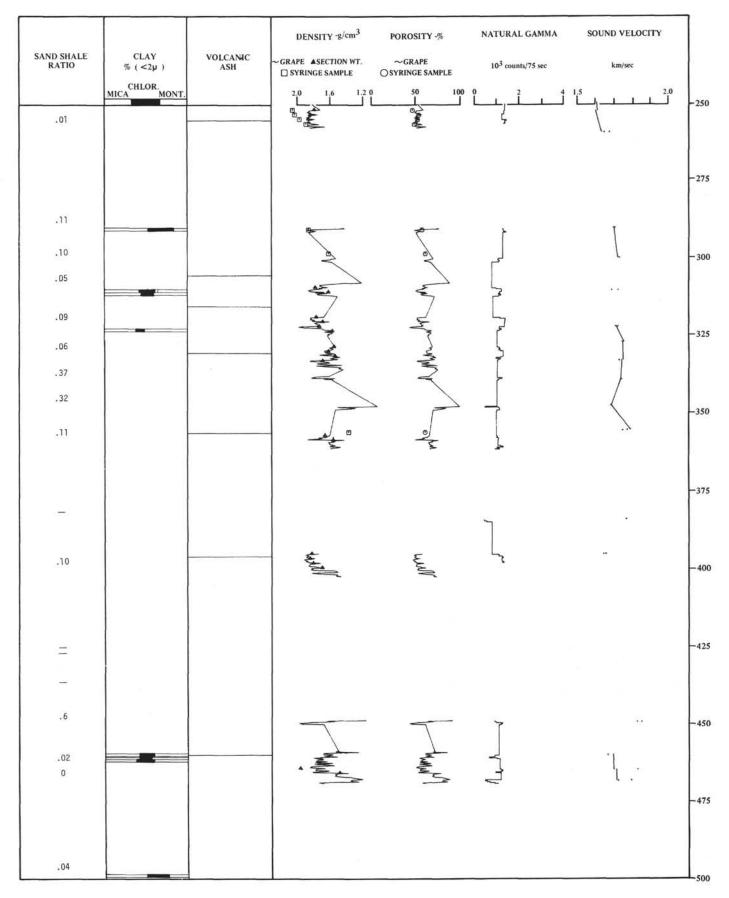
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METERS	DIA- TOMS	FORAM- INIFERA	NANNO– RADIO– FOSSILS LARIANS		CHRONO- STRATI- GRAPHY	GRAPHICAL LITHOLOGY	RECOVERY CORE NO.	LITHOLOGIC DESCRIPTION
0— 25—	I ZDZ	N22/237		ARTOST ROB IUM MI RAL ESTENSE			1 2 3 4	Gray SILTY CLAY, with occasional ASH laminae, beds of MUD RICH DIATOM OOZE, SILT and SAND. Pebbles and granules scattered throughout section.
50—	п			ARTOSI			5 6 7 8	
75—	ZQAN	N22		AXOPRUNUM ANGELINUM	PLESITOCENE		9 10 11 12	
125—	NPDZ III						13 14 15	
150—	VI ZOAN			EUCYRTIDIUM MATUYAMAI			16 17 18	
175—	IV-V ZŪAN	N21/22		EUC			19 20 21 22	
200-	1224				DCENE		23 24 25	
225—	IIV ZOAN	N21			LATE PLIOCENE		26 27 28	
250								

# **SITE 178**

			DENSITY -g/cm <sup>3</sup>	POROSITY -%	NATURAL GAMMA	SOUND VELOCITY	
SAND SHALE RATIO	CLAY % ( <2µ )	VOLCANIC ASH	~GRAPE ▲SECTION WT. ☐ SYRINGE SAMPLE	~GRAPE OSYRINGE SAMPLE	10 <sup>3</sup> counts/75 sec	km/sec	
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METERS	DIA- TOMS	FORAM- INIFERA	NANNO- FOSSILS	RADIO- LARIANS	CHRONO- STRATI- GRAPHY	GRAPHICAL LITHOLOGY	RECOVERY	CORE NO.	LITHOLOGIC DESCRIPTION
250-		N20-N21						29	
275 —									
300-			1	T			H	30	
300-	IIV ZQAN							31 32	
	NPD							33	
325 —		l.				<u>, , , , , , , , , , , , , , , , , , , </u>	-	34	
		?						35	
350 -					CENE			36	
					E PLIOCENE			37	
375 —					LATE				
								38	
400 -								39	
	NPDZ VIII								
425-								41	
			1			+ + + +	+	42	CARBONATE CEMENTED SANDSTONE SILTY CLAY CHALK
450—							H	43	
				2		<del>7.17.17.</del> 17.		44	
								45	
475—	XI ZOAN				PLIOCENE				
500								46	

SITE 178



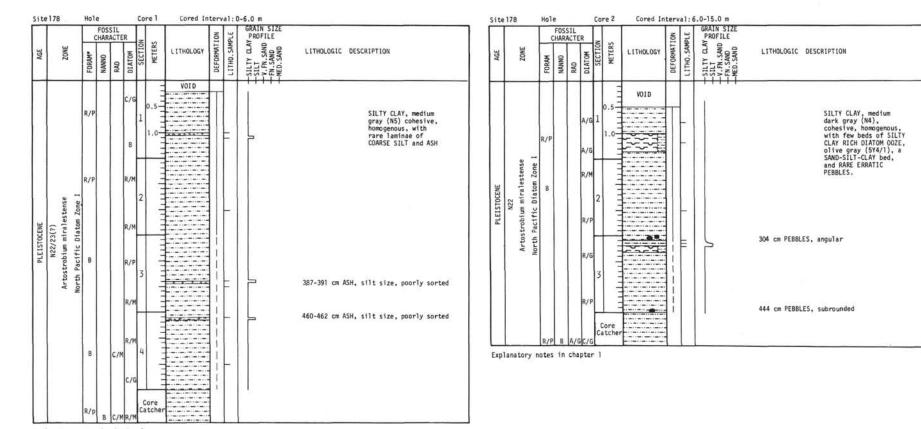
ts		BIOST	FRATIGRAPHY				ž		
METERS	DIA- TOMS	FORAM- INIFERA	NANNO- FOSSILS	RADIO- LARIANS	CHRONO- STRATI- GRAPHY	GRAPHICAL LITHOLOGY	RECOVERY	CORE NO.	LITHOLOGIC DESCRIPTION
500—								46 47	
525—									
								48	
550—	Z IX 2								
575—	ZQAN	?			PL IOCENE				
							4	19	
600 —									
625—								50	
650—	X ZOAN								
	N				PL I OCENE		-	51	
675—					EARLY			50	
700 —								52	
725 —	2				MID-PLIOCENE		5	53	
750	NPDZ   XIV?	- <u>-</u>	DISCOASTER DEFLANDREI		LOWER MIOCENE		5	54	VARICOLORED CLAYSTONE INTERBEDDED WITH CHALK

**SITE 178** 

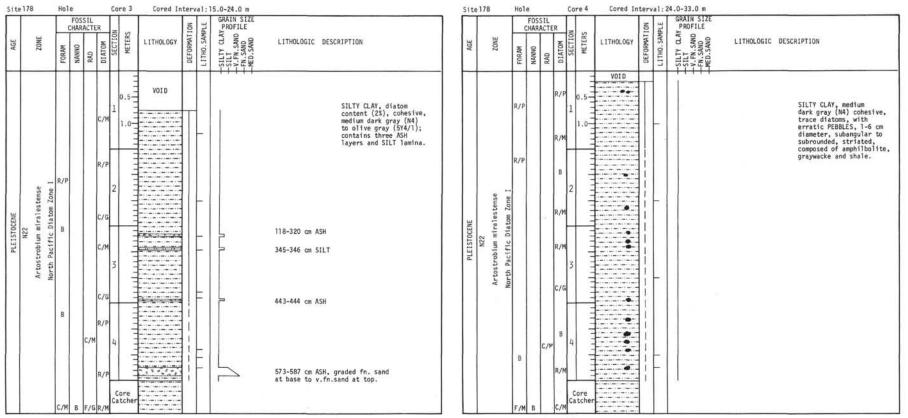
SAND SHALE RATIO .84	CLAY % (<2µ) MICA MONT.	VOLCANIC ASH	DENSITY -g/cm <sup>3</sup> ~ GRAPE ASECTION WT. □ SYRINGE SAMPLE 2.0 1.6 1.2 0 0	POROSITY -% ~GRAPE OSYRINGE SAMPLE	NATURAL GAMMA 10 <sup>3</sup> counts/75 sec	SOUND VELOCITY km/sec 1.5 2.0 	- 500
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750-	1	1			MIOCENE ?			55 56 57	Olive gray, fissil, CLAYEY SHALE
775—								58 59	Basalt
800-									
825—									
850—									
875 —									
900-									
925—									
950—									
975—									
1000									

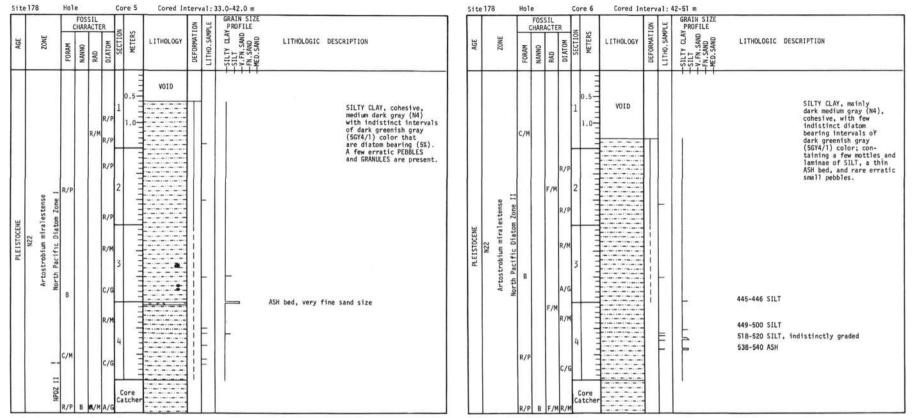
			DENSITY -g/cm <sup>3</sup> POROSITY -% NATURAL GAMMA SOUND VELOCITY	
SAND SHALE RATIO	CLAY % (<2µ)	VOLCANIC ASH	~GRAPE ▲SECTION WT. ~GRAPE 10 <sup>3</sup> counts/75 sec km/sec □ SYRINGE SAMPLE ○SYRINGE SAMPLE	
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				-800
				-825
				- 850
				-875
-				- 900
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Explanatory notes in chapter 1 \*PLANKTONIC FORAMINIFERA SITE 178

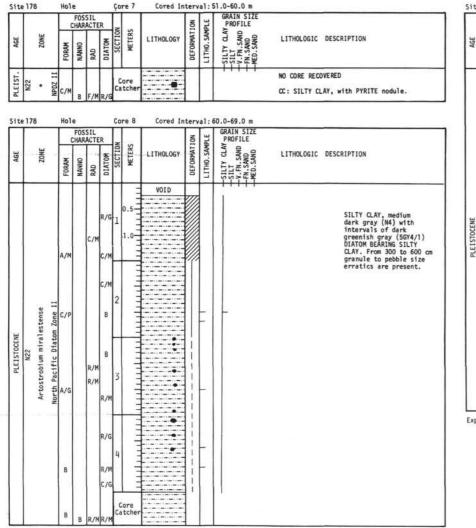


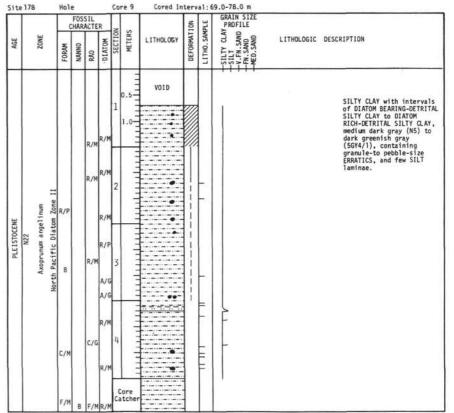
Explanatory notes in chapter 1



Explanatory notes in chapter 1

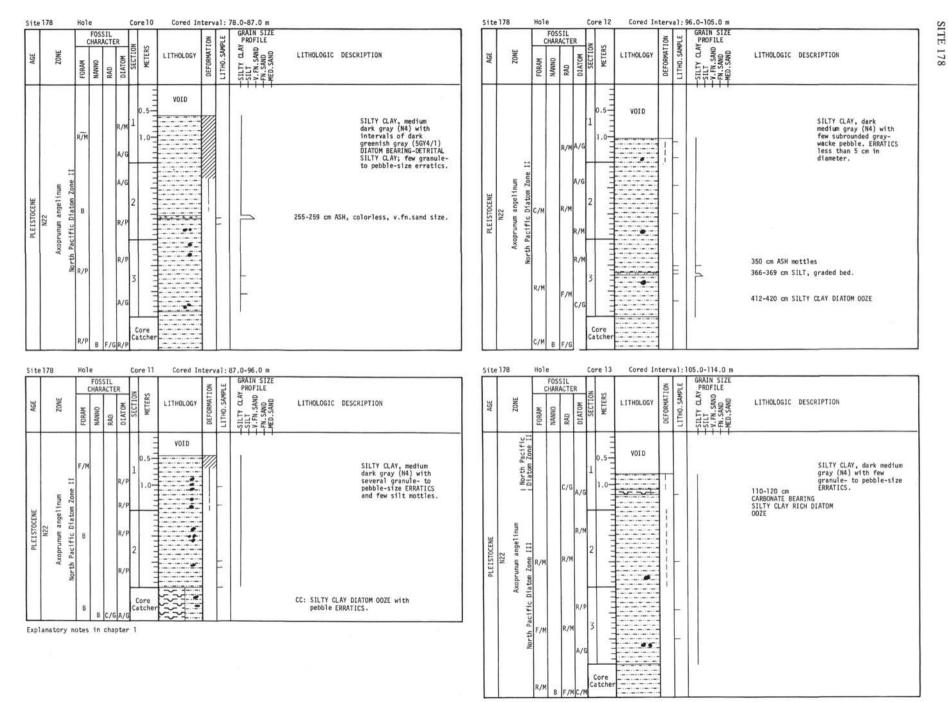
**SITE 178** 



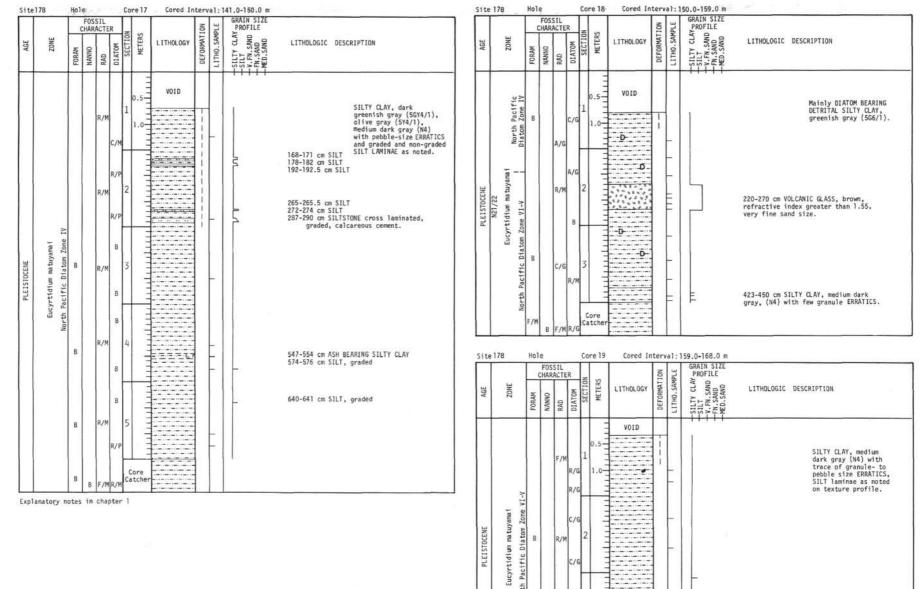


Explanatory notes in chapter 1

**SITE 178** 



Site 178	Ho1	e		Co	re 14	2 1	Cored	Inter	val:	114.0-123.0 m		Sit	e 17	8	Hole			ore 16	Cored In	terva	al:132.0-141.0 m	
AGE ZONE		FOSS CHARA ONNEN	CTER	SECTION	METERS		THOLOGY	DEFORMATION	LITHO.SAMPLE	GRAIN SIZE PROFILE DIAL ARD DIAL STILL CLAY BURSAN	LITHOLOGIC DESCRIPTION	AGE		ZONE	Cł	TOSSIL TARACT ONNON	ER	METERS	LITHOLOGY	DEFORMATION	CRAMPLE CRAMPLE CLITHO.SAMPLE CLITHO.SAMPLE CLITHO.SAMPLE CRAMPLE CLITHO.SAMPLE CRAMPLE CRAMPLE CLITHO.SAMPLE CRAMPLE CRAMPLE CLITHO.SAMPLE CRAMPLE CLITHO.SA	LITHOLOGIC DESCRIPTION
PLEIST. N22	NP 02 111	в	R/M I		ore tcher			-			NO CORE RECOVERED. CC: SILTY CLAY, medium dark gray (N4)			Pacific Zone III				0.5-	VOID		11111	Mainly SILTY CLAY medium dark gray (N4) with trace of
Site 178	0	FOSS	IL CTER DV4	NO	WETERS		Cored 1	TION		GRAIN SIZE PROFILE ONVS'NJ- LIIS LIIS	LITHOLOGIC DESCRIPTION			North Pac	в		A/G A/G	1.0	VOID			(N4) with trace of granule- and pebble-size ERRATICS, and DIATOM- DETRIAL SILTY CLAY, dark greenish gray (56Y4/1).
PLEISTOCENE tidium matuyamai	ic Diatom Zone III co co		RJ F/M Rj	1 /P	1.0		•				0-150 cm SILTY CLAY, medium dark gray (N4) with few granules. 130 cm PEBBLE 150-300 cm DIATOM BEARING and DIATOM RICH-DETRITAL SILTY CLAY, dark greenish gray (550/471).	PLEISTOCENE		Eucyrtidium matuyamai Diatom Zone IV	в	R/I	A/G A/G	3	**			305-310 cm SILTY CLAY BEARING DIATOM OOZE
BLEISIG BLEISIG	a North Pacif	BR	A/G A, A/M R, hapti	Ca Ca	ore	-6	•		- 2		CC: PEBBLE, subrounded, graywacke, 9-cm diameter.			North Pacific Di		c/I	C/G 4 R/P	4 Core		ių,	-	605-608 cm - graded SILT



В

C/6

R/M

C/G

-

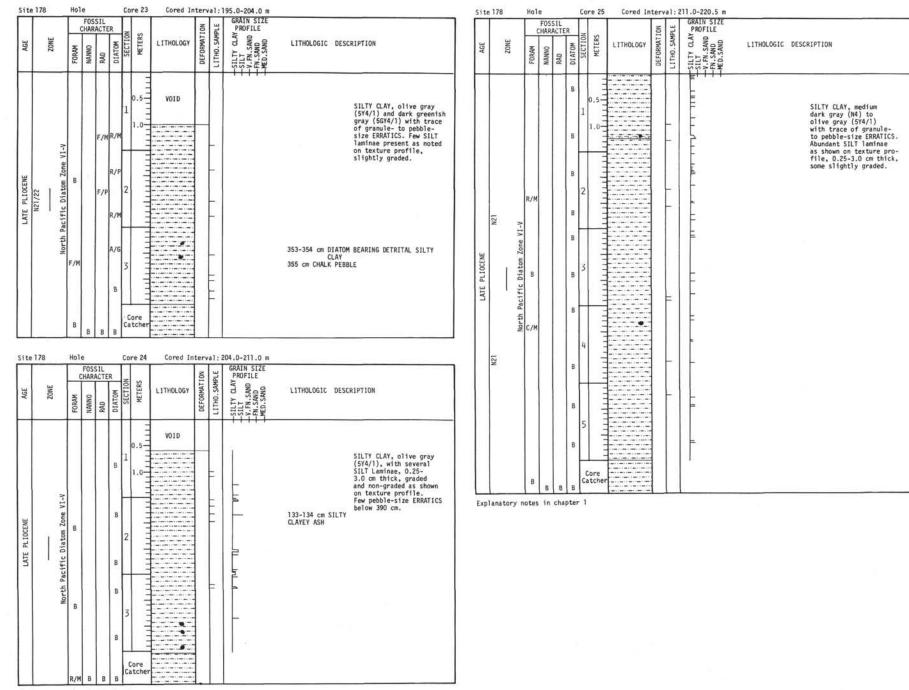
Core Catche .

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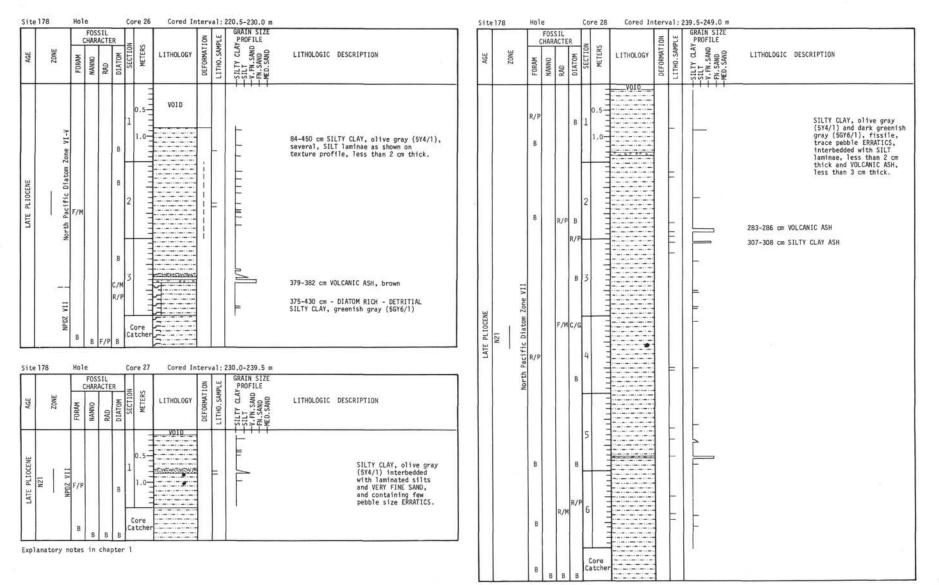
SITE : 178

Site 178	Hole Core 20 Cored Interval: 168.0-177.0	m	Site 178	Hole		Core	21 Cored 1	Interv	/a1:	177.0-186.0 m	
AGE ZONE	FOSSIL CHARACTER CHARACTER NOTINO CHARAC		AGE ZONE	FOS CHAR WANNO	ACTER	SECTION	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	-SILTY CLAV -SILT -V.ENU CLAV -V.ENU CLAV -V.ENU CLAV -V.ENU -SAND -FIL -FIL -FIL -FIL -FIL -FIL -FIL -FIL	LITHOLOGIC DESCRIPTION
PLEISTOCEHE Eucyrtidium matuyamai Moodo Buordsia Diceber 7000 Vr.0	PTN 9010 R/M 2 R/M 2 R	SILTY CLAY, olive gray (5Y4/1) and dark greenish gray (56Y4/1) with few granule- to pebble- size ERRATICS as noted. Numerous SILT laminae up to 6 cm thick as shown on texture profile.	PLEISTOCENE N21/22 ucyrtidium matuyamai		B R/M B	2					SILTY CLAY, olive gray (5Y4/1) to dark greenish gray (5GY4/1) with trace of pebble- to granule- size ERRATICS. SILT laminae, 0.5-12 cm thick are present as shown on texture prfoile; a few are slightly graded with hornblende concentrations at base.
	R/M 5		Site 178	Hole		Core	22 Cared 1	Intoni		186.0-195.0 m	
	B B B B B B B B B B B B B B B B B B B		AGE ZONE	FOS	SIL ACTER UV UV UV UV UV	NO	LITHOLOGY	- DEFORMATION	LITHO.SAMPLE	GRAIN SIZE PROFILE PROFILE NEWS SUD PROFILE PROFILE SWD NEWS NEWS NEWS NEWS NEWS NEWS NEWS NEWS	LITHOLOGIC DESCRIPTION
Explanatory	B R/P Core		PLEISTOCENE cyrtidium matuyamai	FACIFIC DIATOM 20Ne	B C/G B	0. 1 1.	5		-		SILTY CLAY, olive gray (SY4/1) to dark greenish gray (SGY4/1); some intervals are very firm.
			L L	B B	F/MR/F	Catc	her				

**SITE 178** 

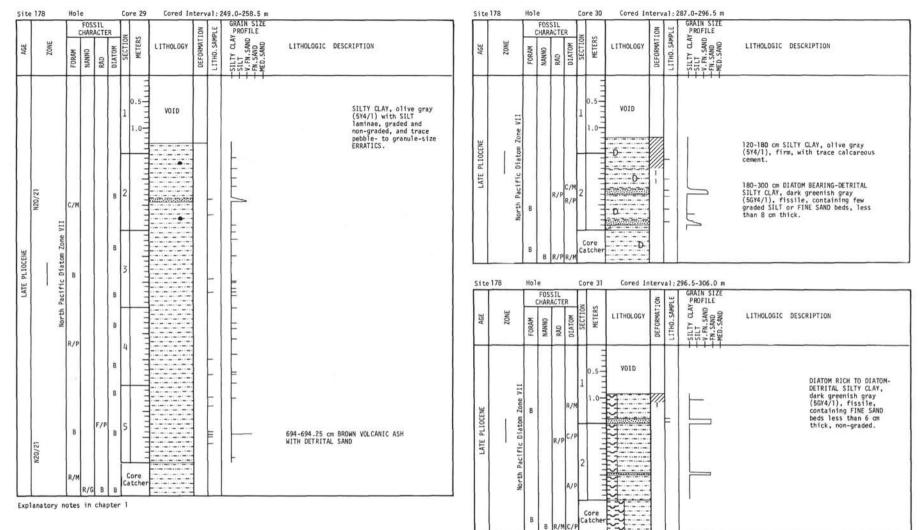


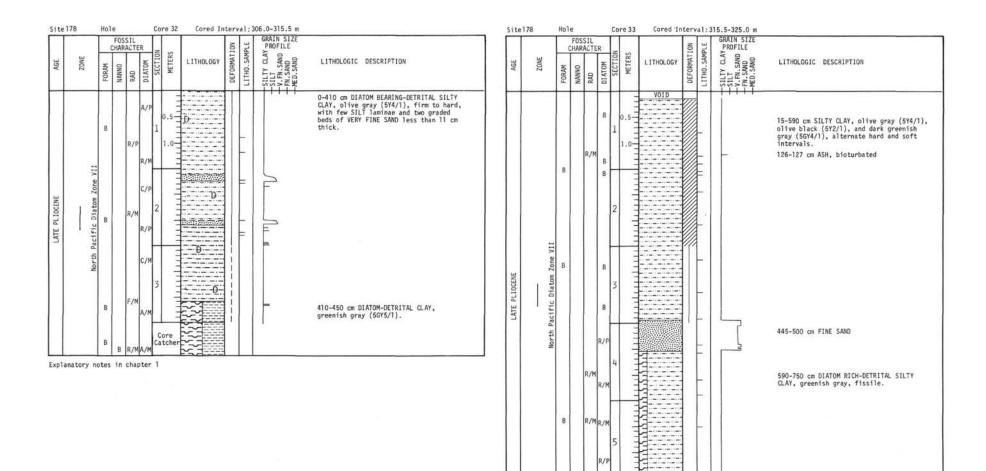
**SITE 178** 



Explanatory notes in chapter 1

**SITE 178** 

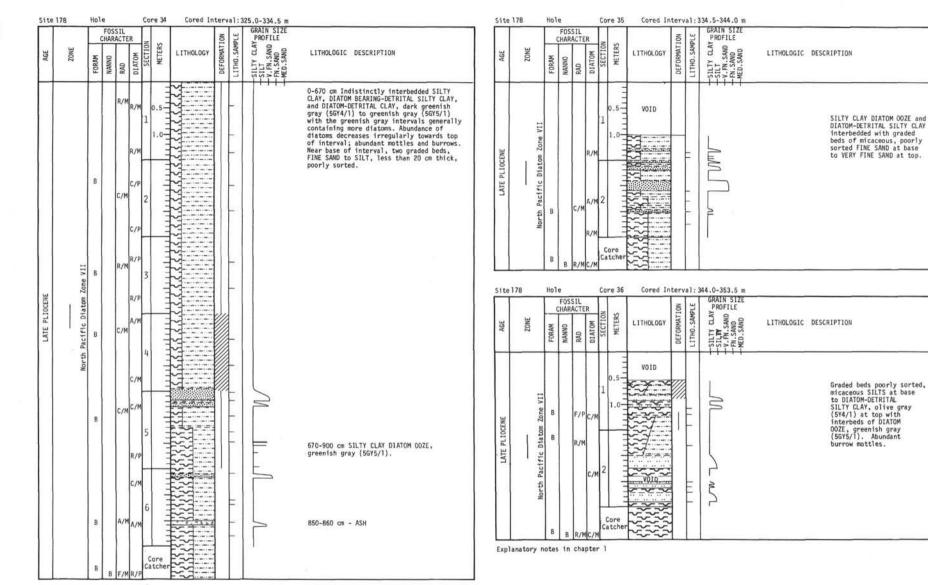




F/PR/F

Core Catcher

**SITE 178** 

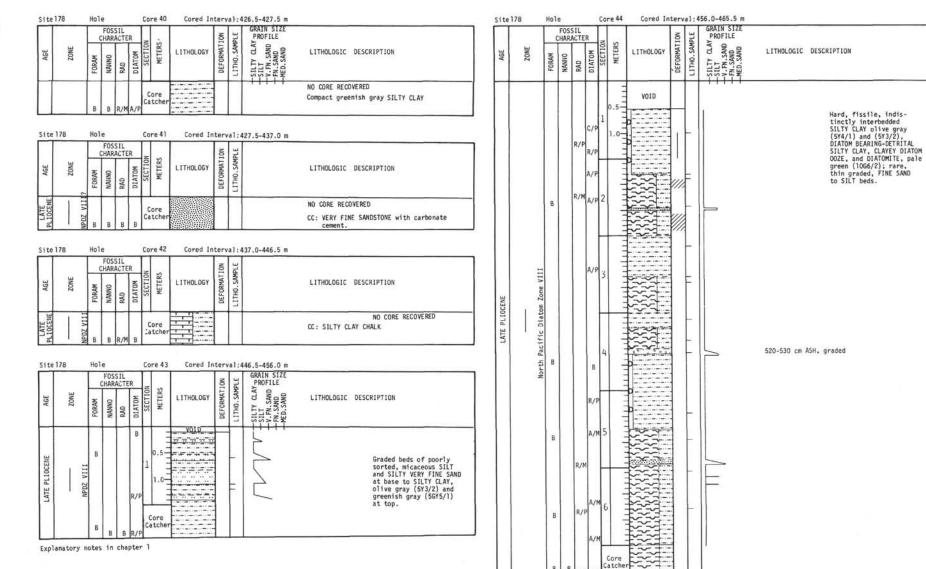


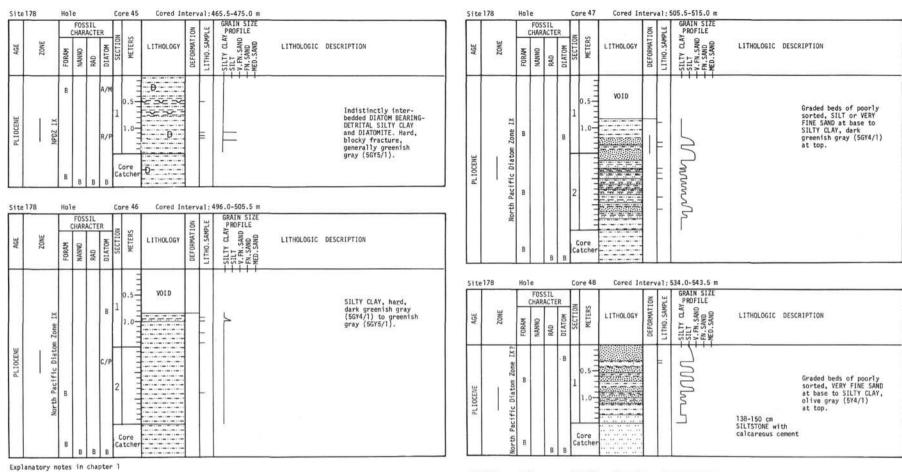
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Explanatory notes in chapter 1

SITE 178

Site 17	78 H	Hole		Cor	e 37	Cored I	nterv	a1:35	3.5-363.0 m				Site	178	Hole		C	ore 39	Cored Int	erval	:391.5-401.0 m	
AGE	ZONE	FORAM 23	SSIL RACTER OVU	SECTION	METERS	. I THOLOGY	DEFORMATION	LITHO.SAMPLE	-SILTY CLAVe BUE -SILT -V.FN.SAND AND FN.SAND FN.SAND -FN.SAND -MED.SAND	LITHOLOGI	DESCRIPTION		AGE	ZONE	CH	OSSIL ARACT UNNN	ER WO.	METERS	LITHOLOGY	DEFORMATION		LITHOLOGIC DESCRIPTION
LATE PLIOCENE	North Pacific Diatom Zone VII	B B B B B B B B	А/М А// R/P A/ C/ F/M A/ A/ A/ A/ A/ A/ A/	2 M3 M3 M Cot							CLAYEY DI DIATOM-DE CLAY, oli to greeni (5GY5/1) with grad poorly so	ttly bedded SILTY ATOM 002E and TRITAL SILTY ve gray (SY4/1) sh gray interbedded led beds of rted VERY FINE ILT at base to oop.	LATE PLIOCENE	North Pacific Diatom Zone VIII «	в		в в в в в в в в в в	0.5			n - - 	140-700 cm interbedded FINE SILT. 2 mm - 2 cm thick, well sorted, and SILTY CLAY, medium dark gray (N4), 2-4 cm thick.
Site 17	78 Н	Hole		Cor	e 38	Cored I	nterv	al:38	2.0-391.5 m							R/1	A/P		225572		1	700-720 cm VOLCANIC ASH
AGE	ZONE	FORAM NANNO 12	STL ACTER OVU	SECTION	METERS	ITHOLOGY	DEFORMATION	LITH0.SAMPLE	BLAIN SIZE - V.FN. SAND - V.FN. SAND - FN. SAND - FN. SAND - MED. SAND	LITHOLOGI	DESCRIPTION				В		- AV M	1.1.1		-		720-900 cm CLAYEY DIATOM 00ZE, greenish gray (505/1) grading upwards to DIATOM BEARING- DETRITAL SILTY CLAY, light
LATE PLIOCENE	h Pacific Diatom	8 8 B	A/M C/ R/M R/	P 1	<u>איזיקיקיקי</u>	•					00ZE, gre (5GY5/1) DETRITAL olive gra all with	tly interbedded YEY DIATOM enish gray and DIATOM- SILTY CLAY. y (5Y3/2), abundant 1 worm burrows.	Expl	anatory r		B R/I	MC/P	Core				olive gray (5Y5/1).

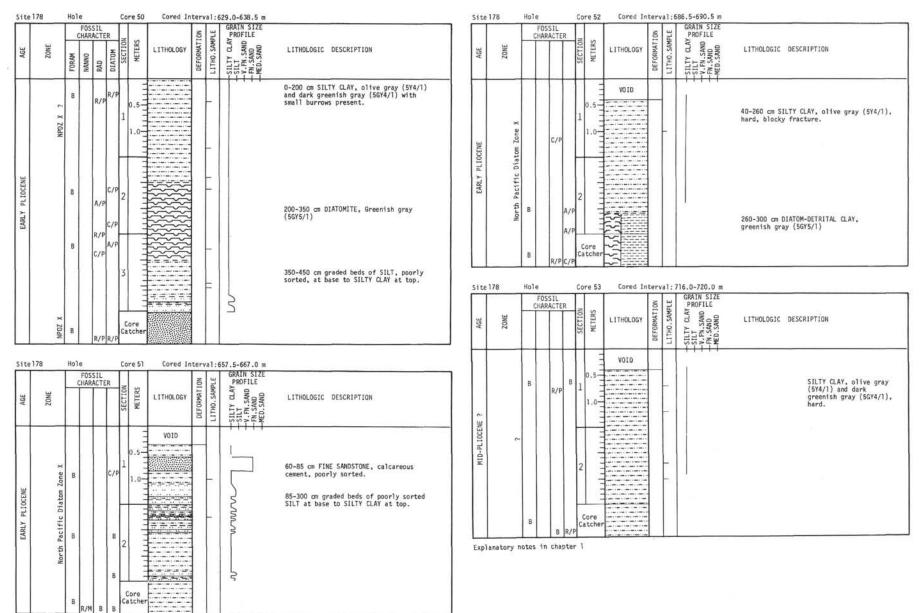




Site178 Hole. Core 49 Cored Interval: 591.0-600.5 m GRAIN SIZE FOSSIL PROFILE DEFORMAT ION SAMPLE CHARACTER SILTY CLAY METERS ZONE AGE LITHOLOGY LITHOLOGIC DESCRIPTION NANNO RAD DIATOM FORAM LITHO. VOID R Graded beds of poorly B FINE SAND at base to SILTY CLAY, olive gray (5Y4/1) and dark PL IOCENE IX 2 3 Tommer. ZOPN P 3H 3H THE C ş greenish gray (5GY4/1) at top. Core Catche

Explanatory notes in chapter 1

**SITE 178** 



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Explanatory notes in chapter 1

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**SITE 178** 

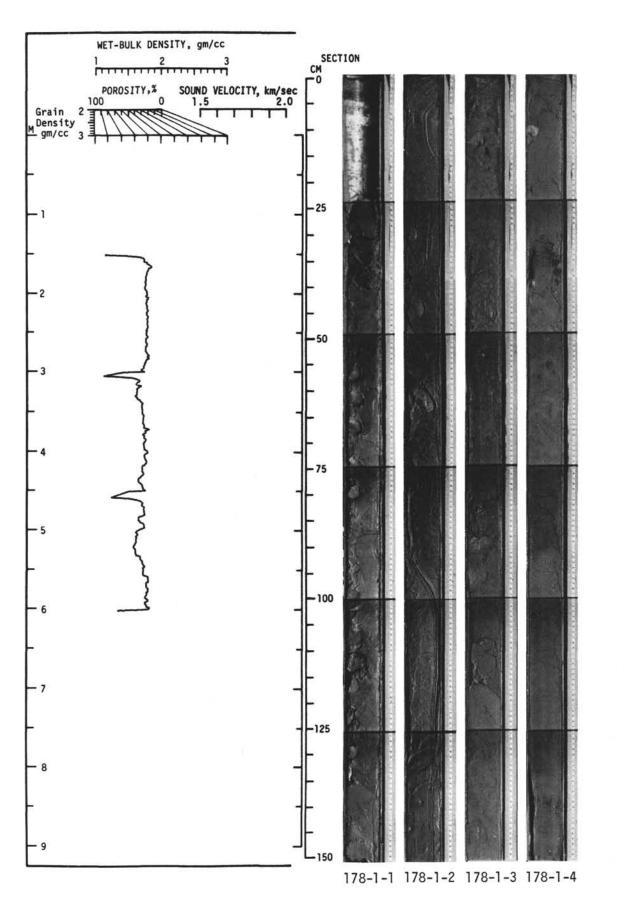
Site178	Hole Core 54			Site178	Hole	Core 55 Core	d Interval:74	9.0-758.5 m	
AGE ZONE	FOSSIL CHARACTER NOILUS	GRAIN ST. BOEFORWATTON LITTHO.SAMPLE SILTY CLAV SAMPLE ST. T. M. SAMPLE		AGE ZONE	FOSSIL CHARACTER	METERS	DEFORMATION LITHO.SAMPLE		LITHOLOGIC DESCRIPTION
	B C/P R/P 0.5	VOID	32-690 cm varicolored CLAYSTONE, dark	MIOCENE 2		Core Catcher			NO CORE RECOVERED
N77	1		brown (SYR4/4), light olive gray (5Y6/2), pale yellowish brown (10YR6/2), containing up to 20 percent diatoms and radiolarians in a few indistinct intervals,	Site 178	Hole	Core 56 Core	d Interval:75	8.5-768.0 m	
	F/M F/P R/P A/P F/M C/P A/P		radiolarians in a few indistinct intervals, interbedded with CHALK, light olive gray (5%C2), light greenish gray (5%Y7/1) and grayish orange pink (5%R7/2). Well preserved burrows over entire interval.	AGE ZONE	FOSSIL CHARACTER	METERS METERS	DEFORMATION LITHO.SAMPLE	-SILTY CLAY B -SILT CLAY B -V.FN. SAND B -V.FN. SAND B -FN. SAND B	LITHOLOGIC DESCRIPTION
				MIOCENE ?	B B I	Core Catcher			NO CORE RECOVERED. CC: CLAY SHALE, fissile, olive gray (5Y4/1)
1 Zone	в С/РС/М -			Site178	Hole	Core 57 Core	d Interval:76	8.0-777.5 m	
/ MIOCE N5/6 defland	7 400 10 10 10 10 10 10 10 10 10 10 10 10 1			AGE ZONE	FOSSIL CHARACTER	METERS METERS	DEFORMATION LITHO.SAMPLE	GRAIN SIZE PROFILE ONVS'OBW- ALTIS- ALTIS-	LITHOLOGIC DESCRIPTION
N5/6 D1sc	B B 4 F/M B 5			2 JNBCOTM Explanatory		B B D D D D D D D D D D D D D D D D D D			100–150 cm CLAY SHALE, fissile, possible 12° dips, dusky yellowish brown (10YR2/2), dark yellowish brown (10YR4/2).
N	B B B B		690-750 CLAY SHALE, olive gray (5Y4/1), fissile, possible 6 degree dips on bedding planes.						

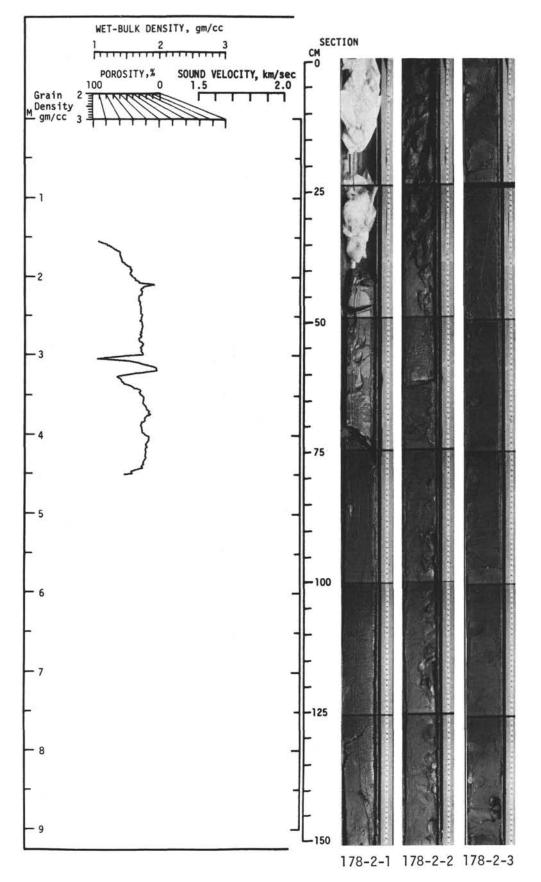
Explanatory notes in chapter 1

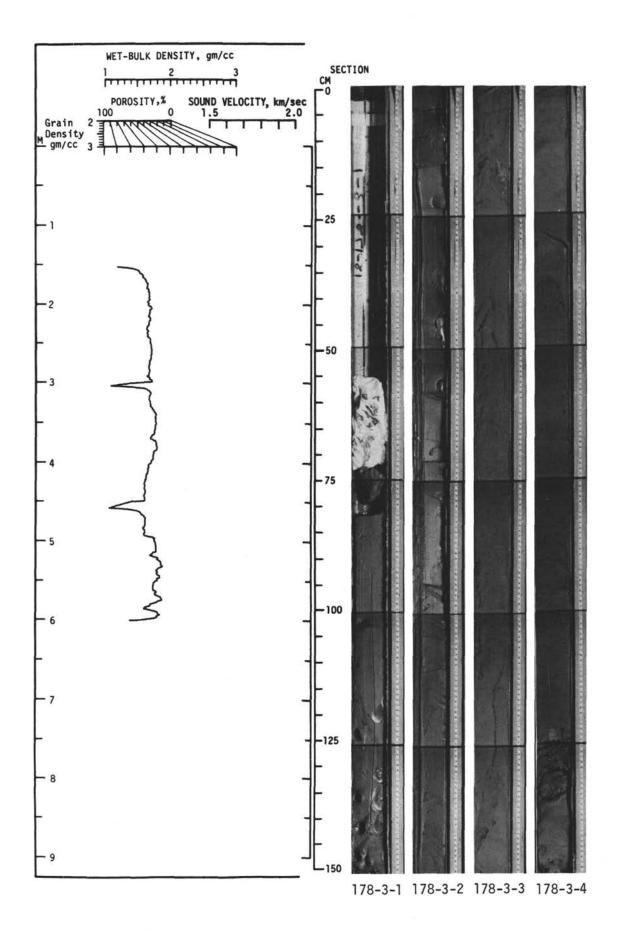
ite		T	e FOS			T	-		2		GRAIN SIZE	
AGE	ZONE		HAR	ACTE	R	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	-SILTY CLAY - -SILT CLAY - -SILT 33 -V.FN.SAND - -FN.SAND -	LITHOLOGIC DESCRIPTION
		В	В	В	В	1	2.5	VOID			1	20-30 cm CLAY SHALE, olive gray (5Y4/1), no baking evident, possibly recovered contact with underlying basalt. 30-150 cm BASALT.
							cher					
yde alle	178 JNOZ	Hol	FOS	SIL	R	SECTION	METERS 65	Cored Ir	DEFORMATION	LITHO.SAMPLE	787.0-794.5 m GRAIN SIZE PROFILE ONVS'NJ_ APROFILE ONVS'NJ_ APROFILE ONVS'NJ_ APROFILE	LITHOLOGIC DESCRIPTION

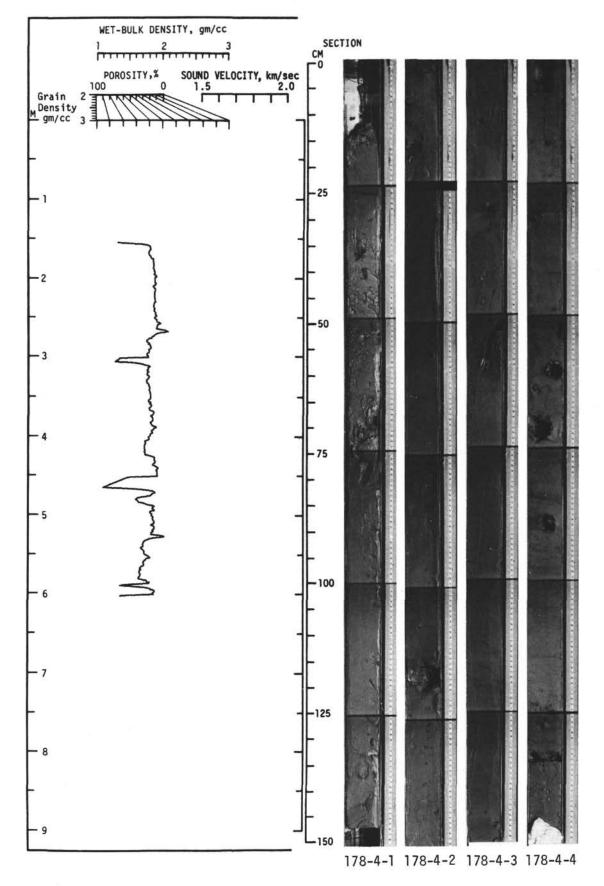
Explanatory notes in chapter 1

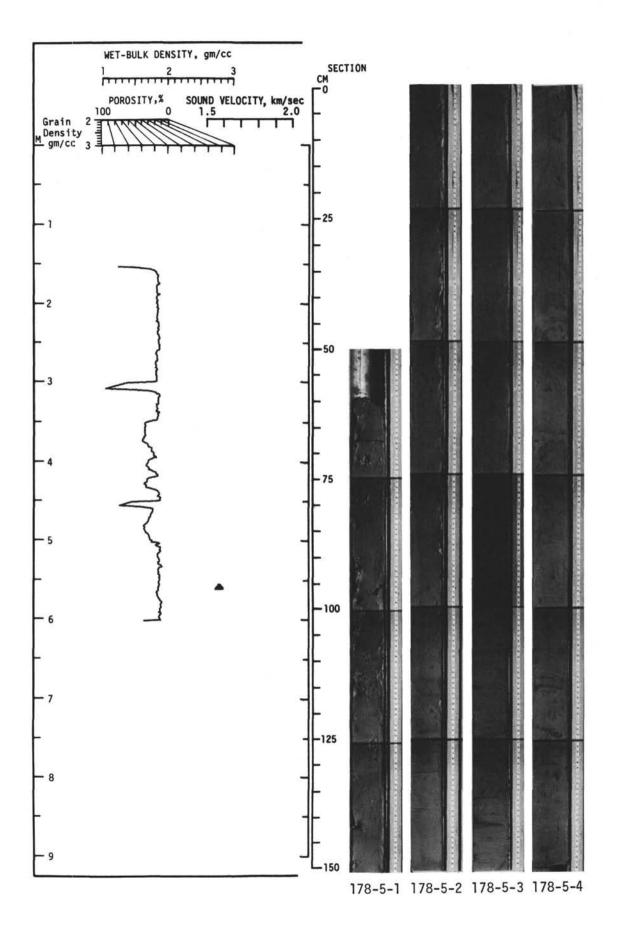
Core Catcher

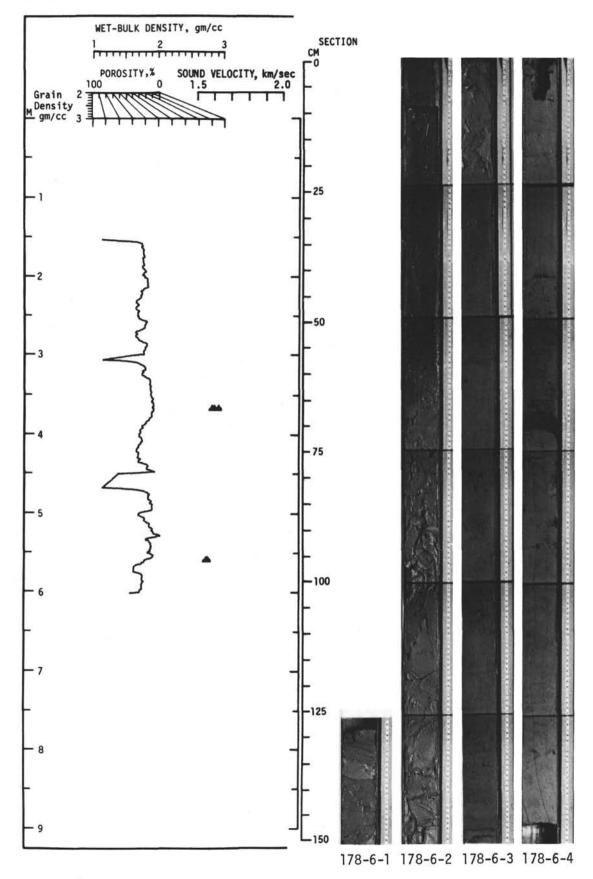


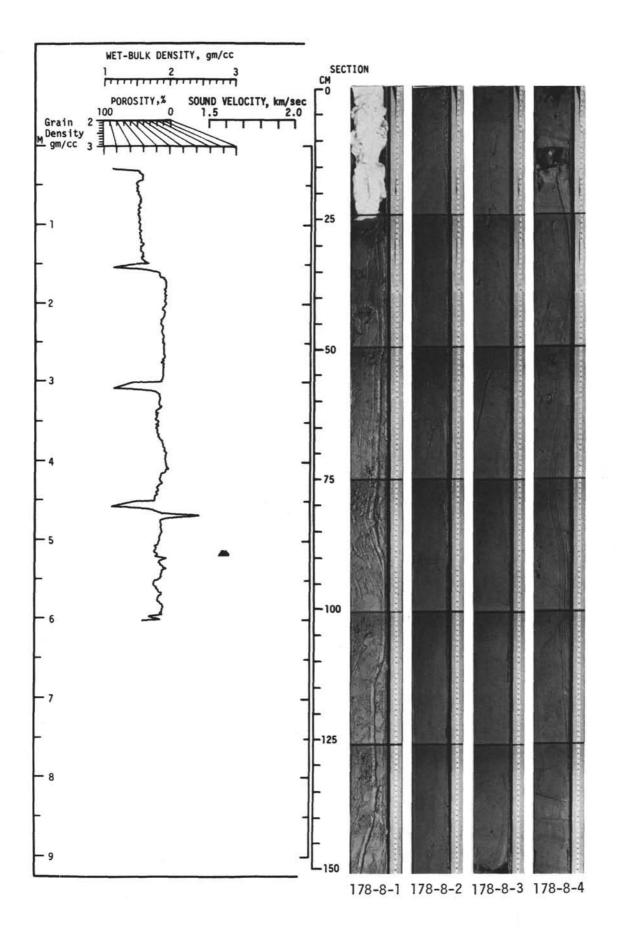


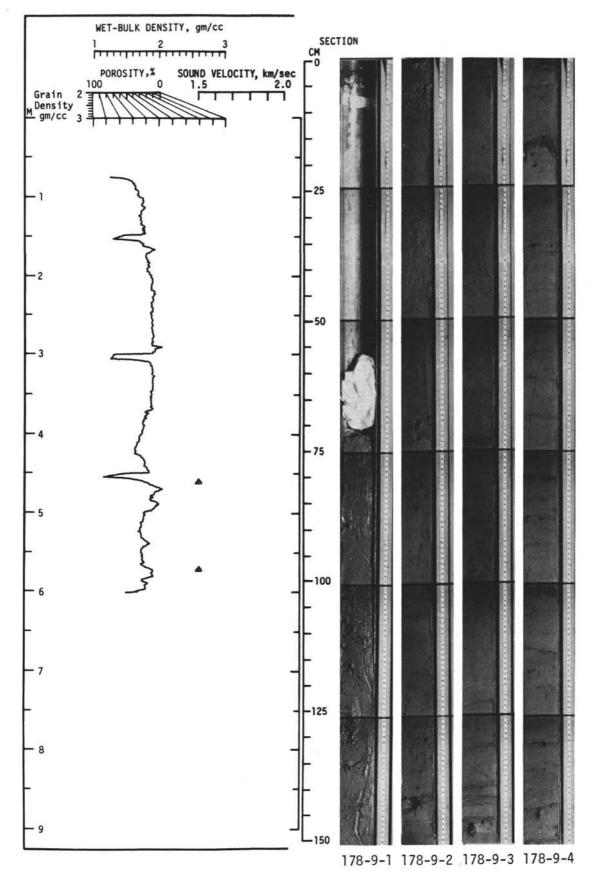


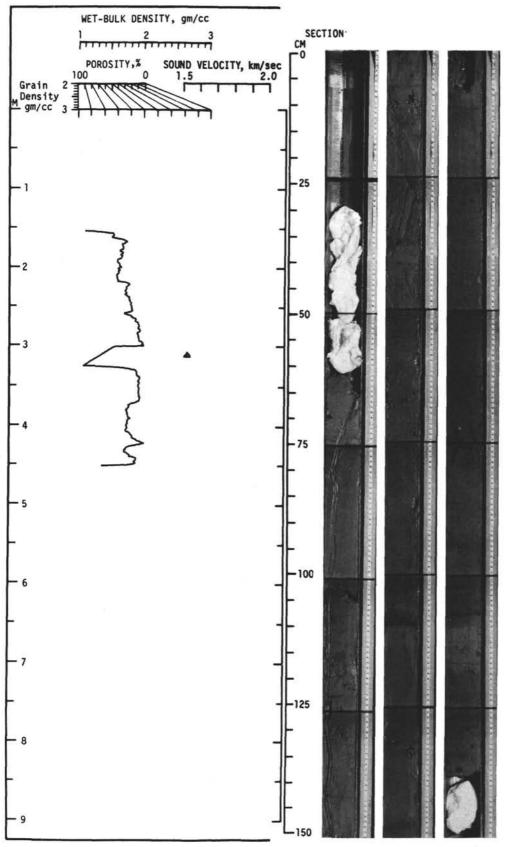






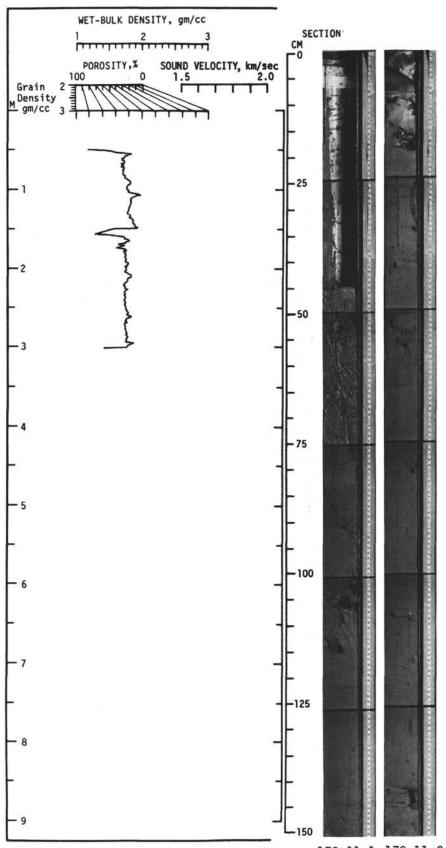




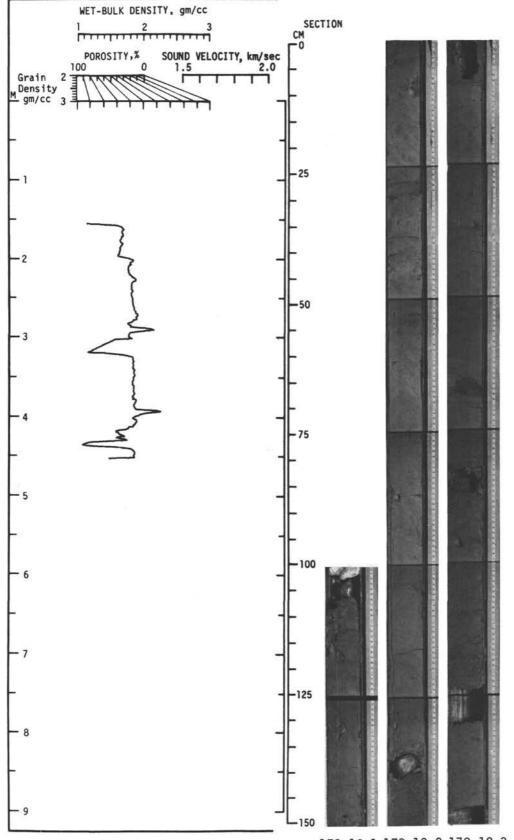


**SITE 178** 

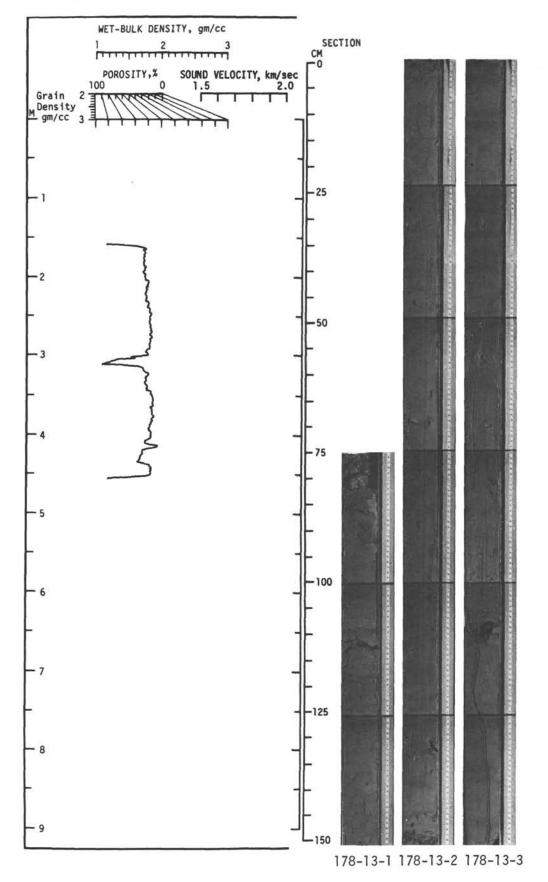
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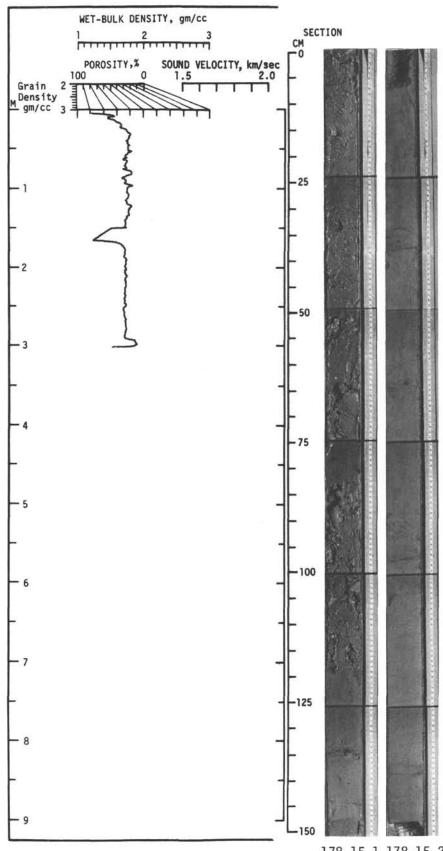


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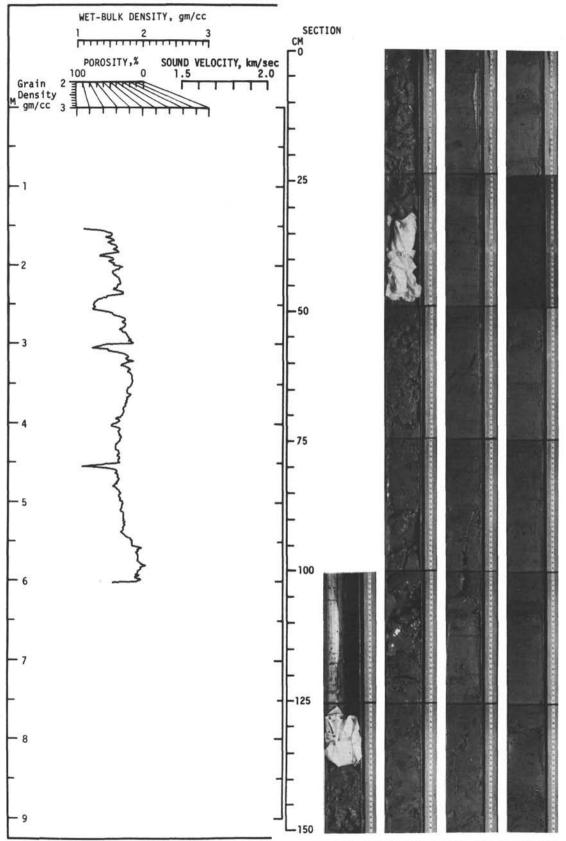


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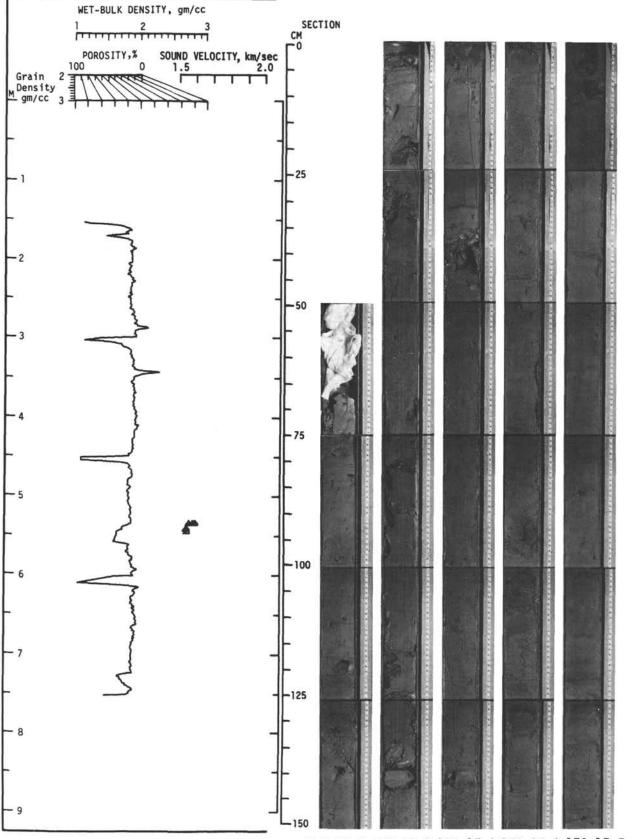




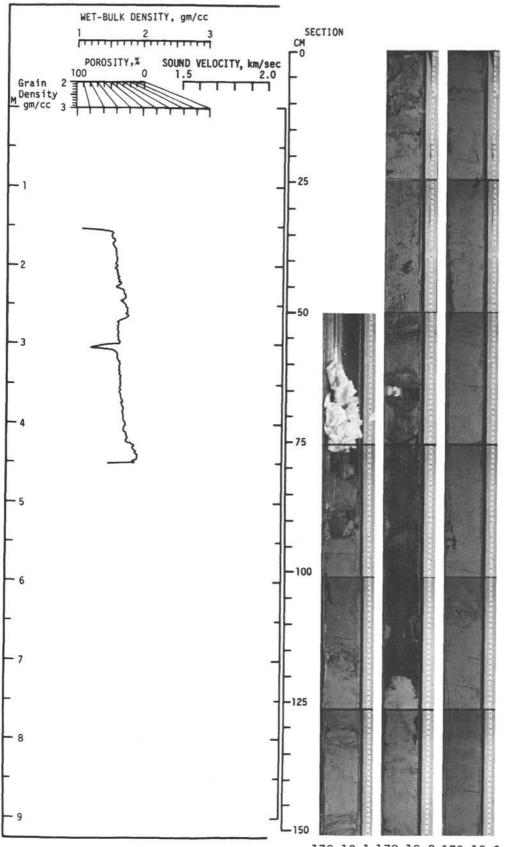
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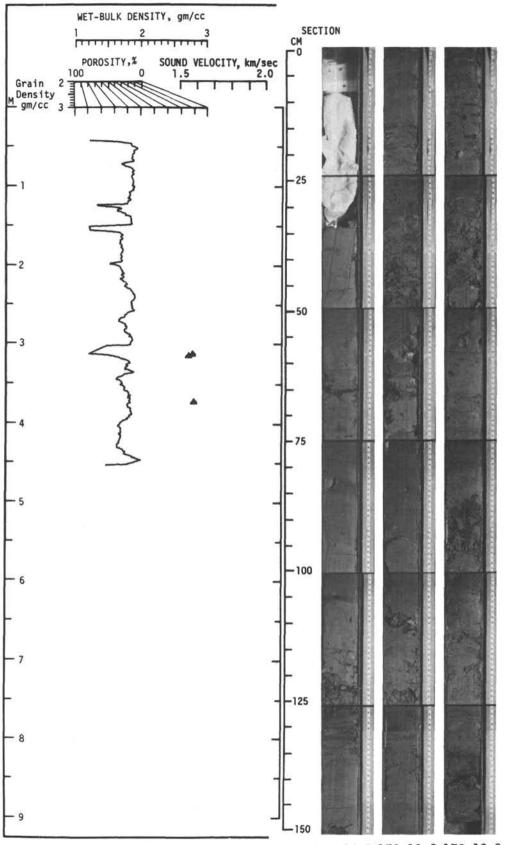
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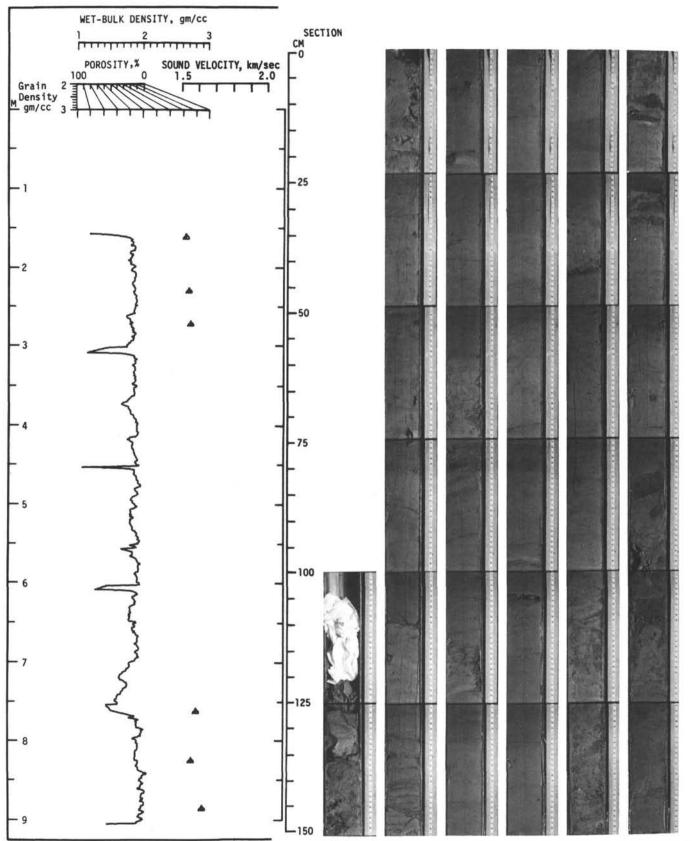
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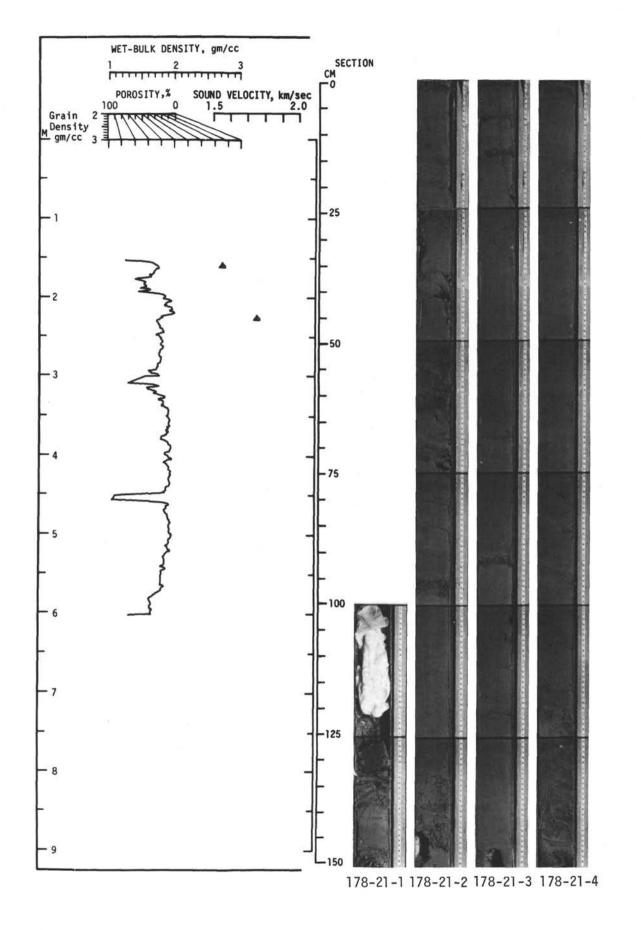
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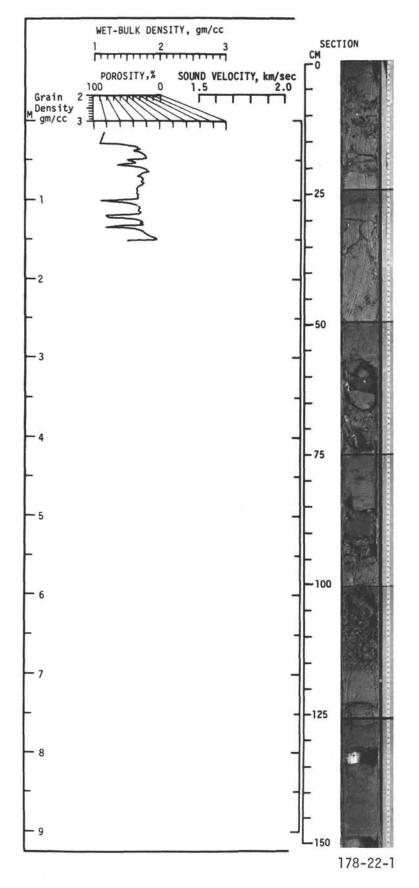


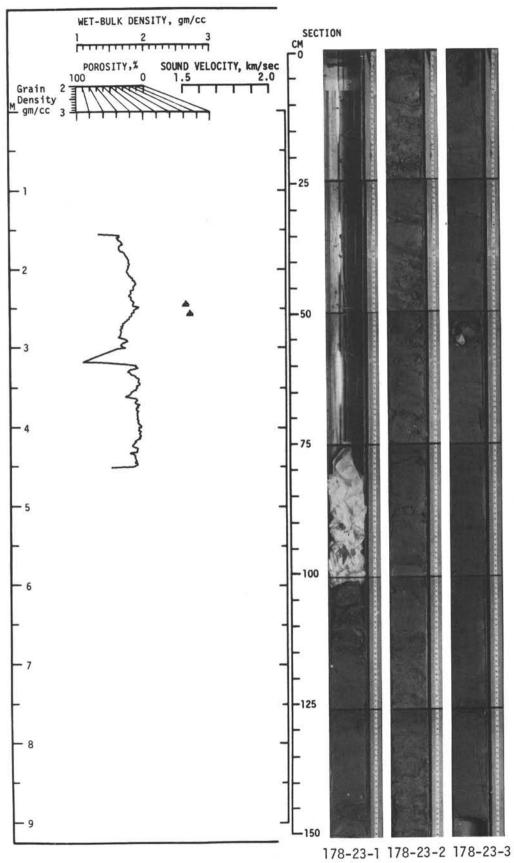
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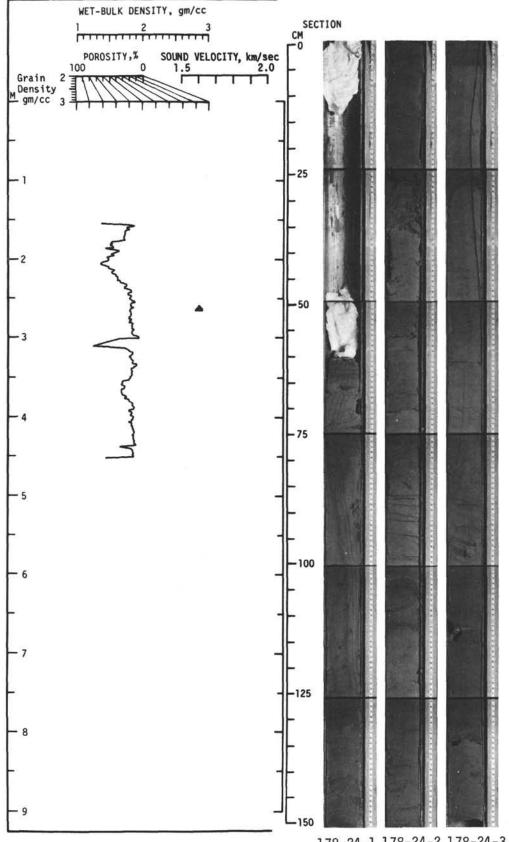


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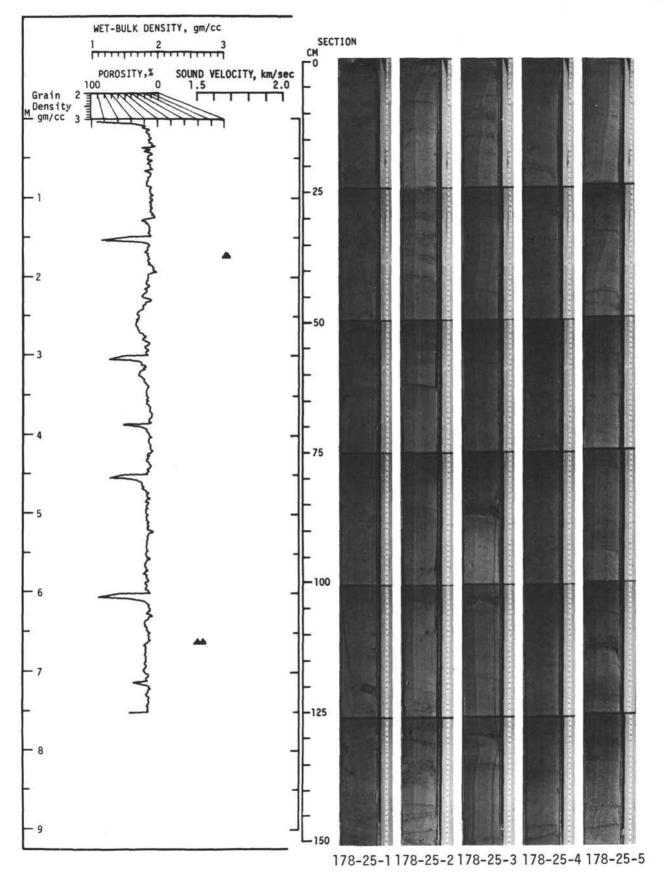


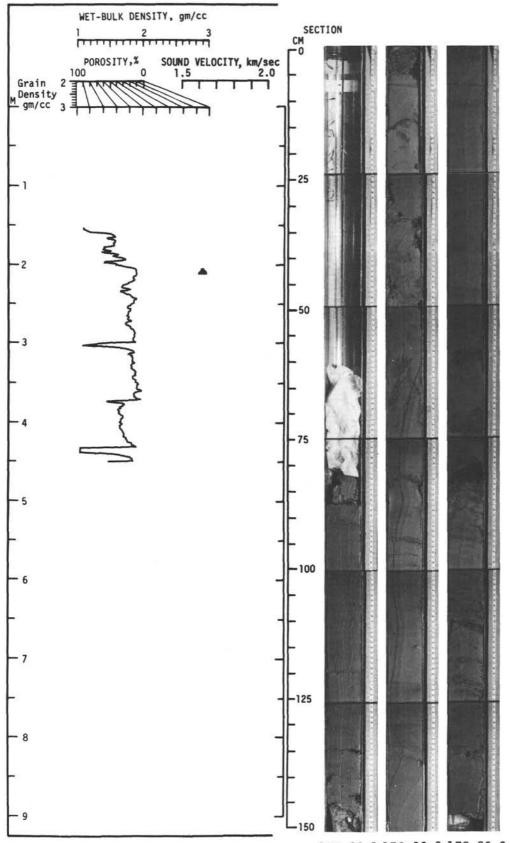




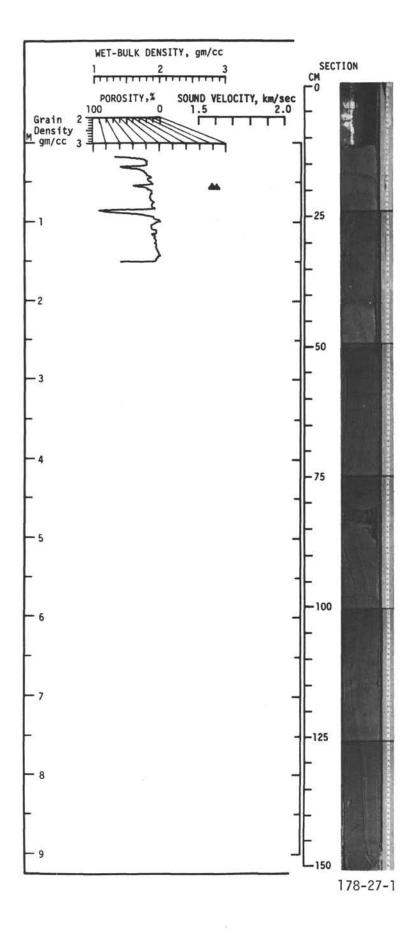


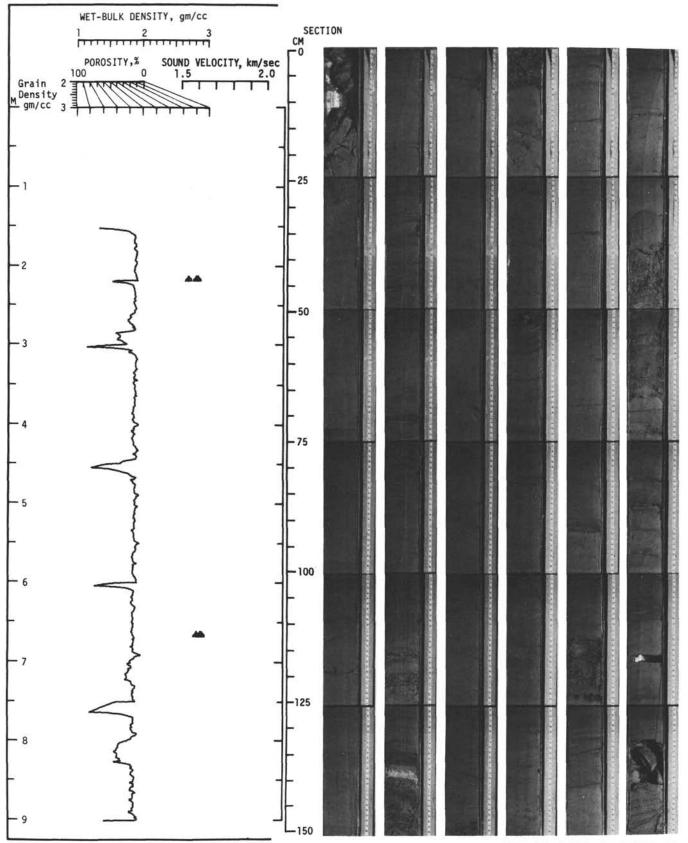
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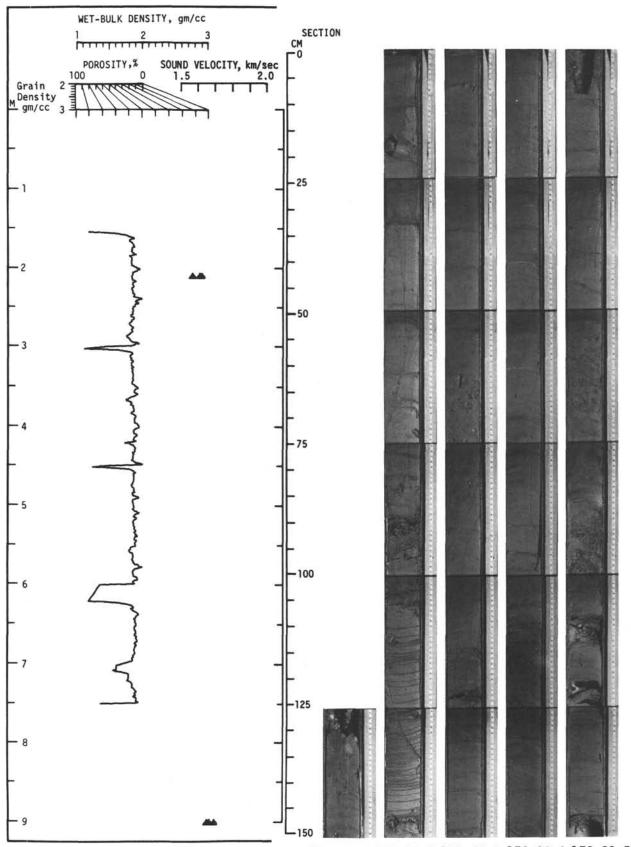


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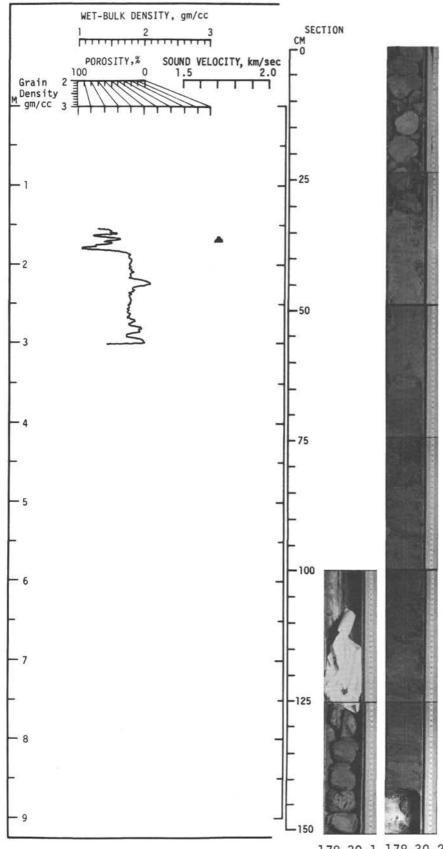




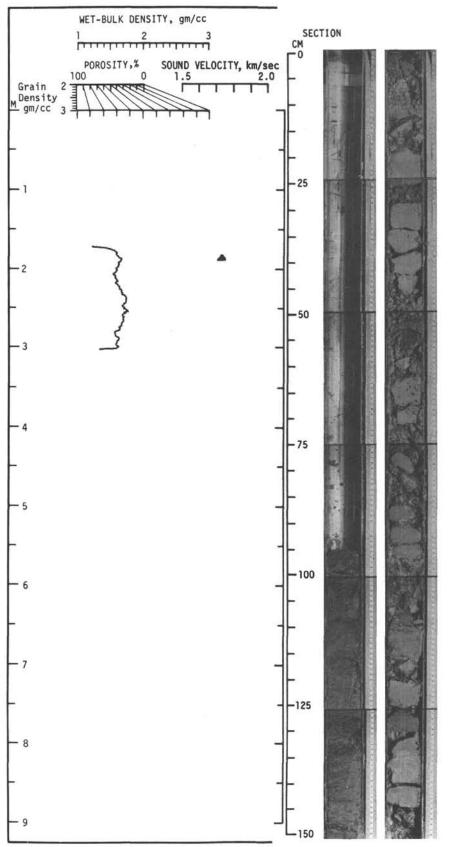
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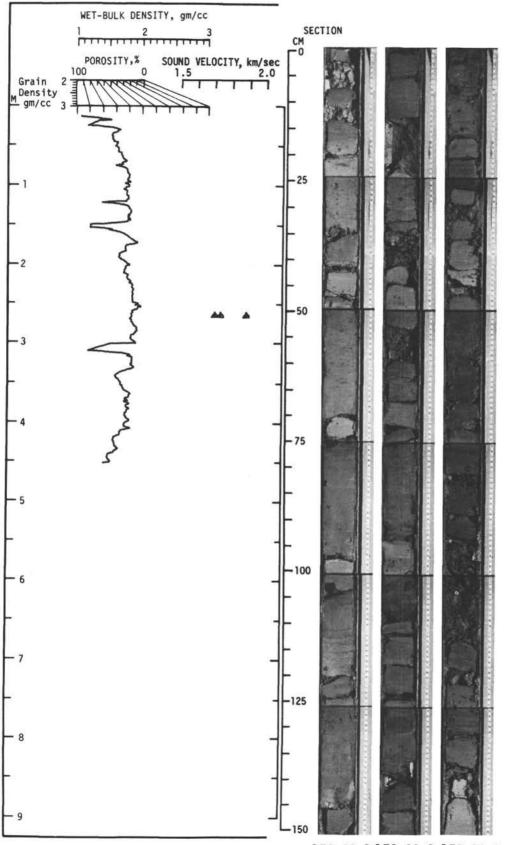
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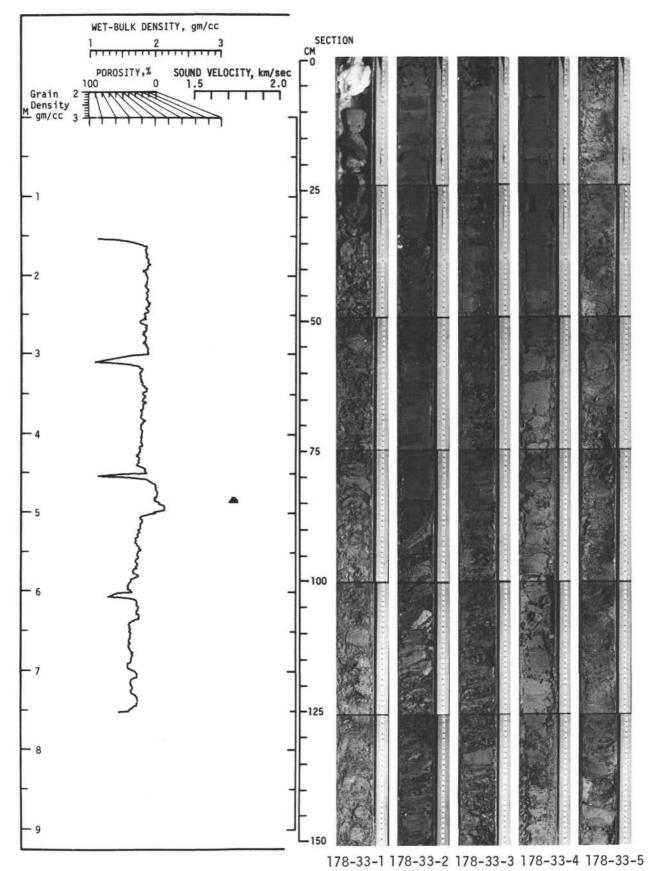
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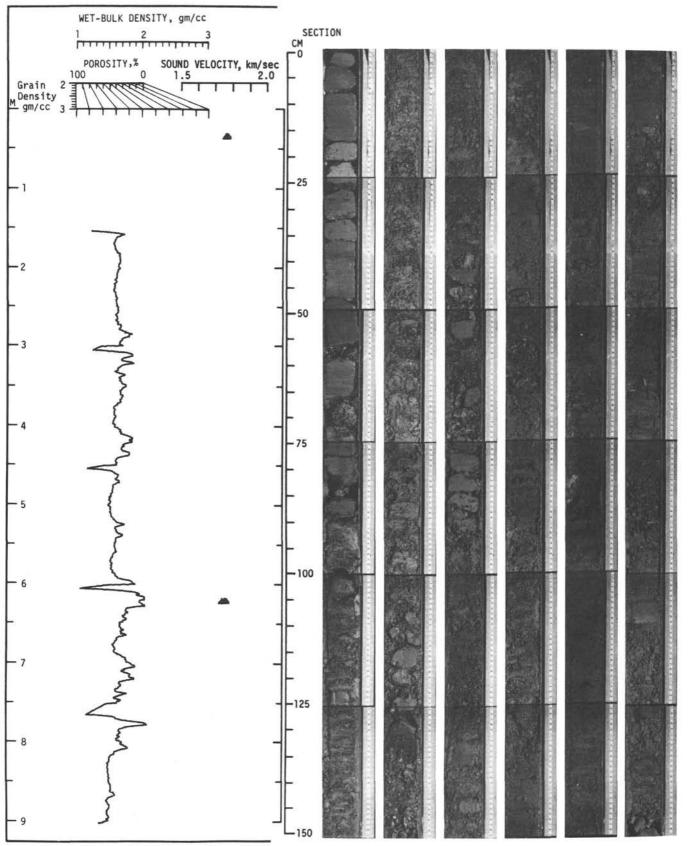


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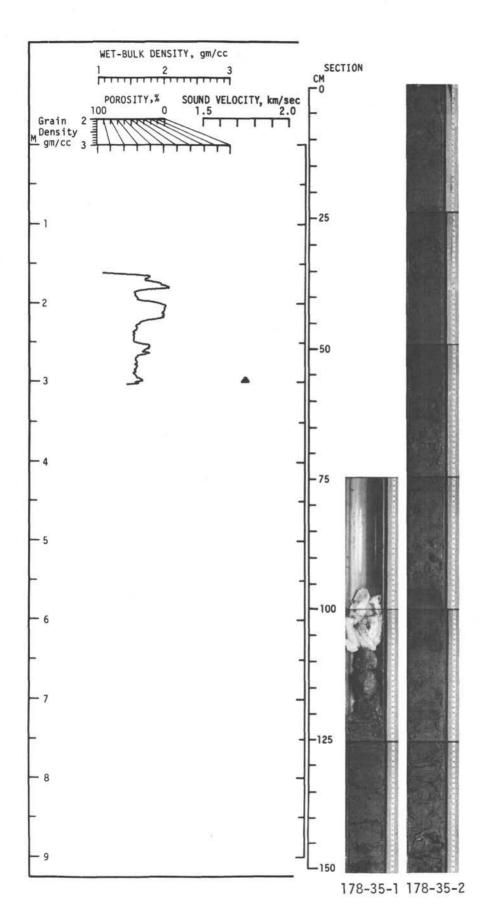


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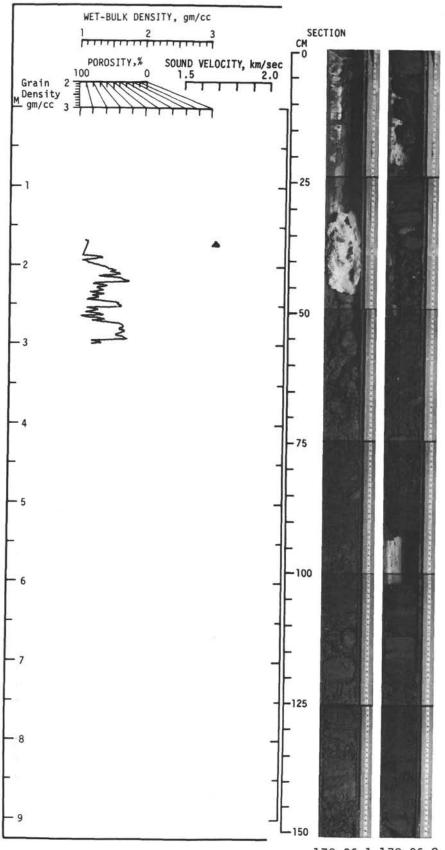




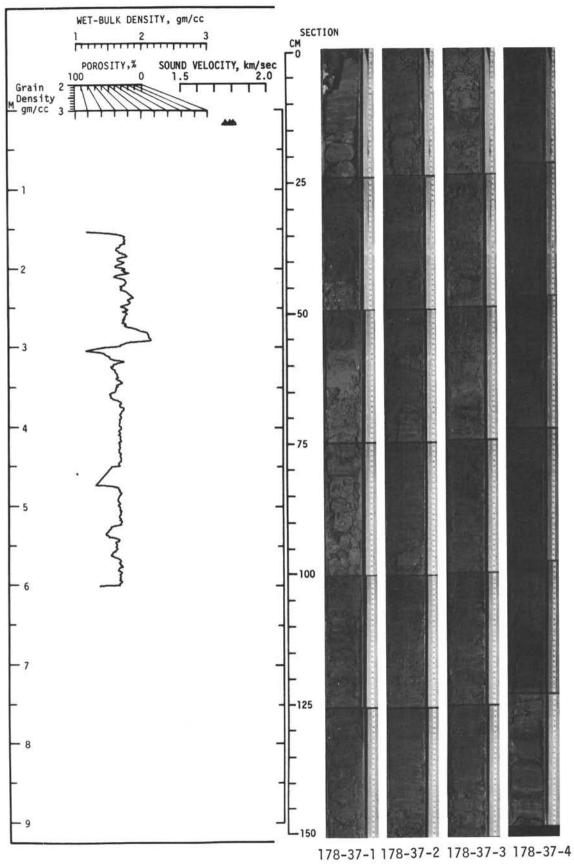
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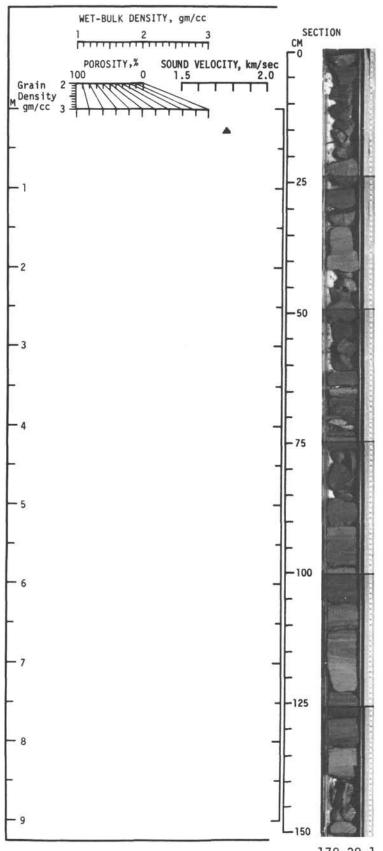


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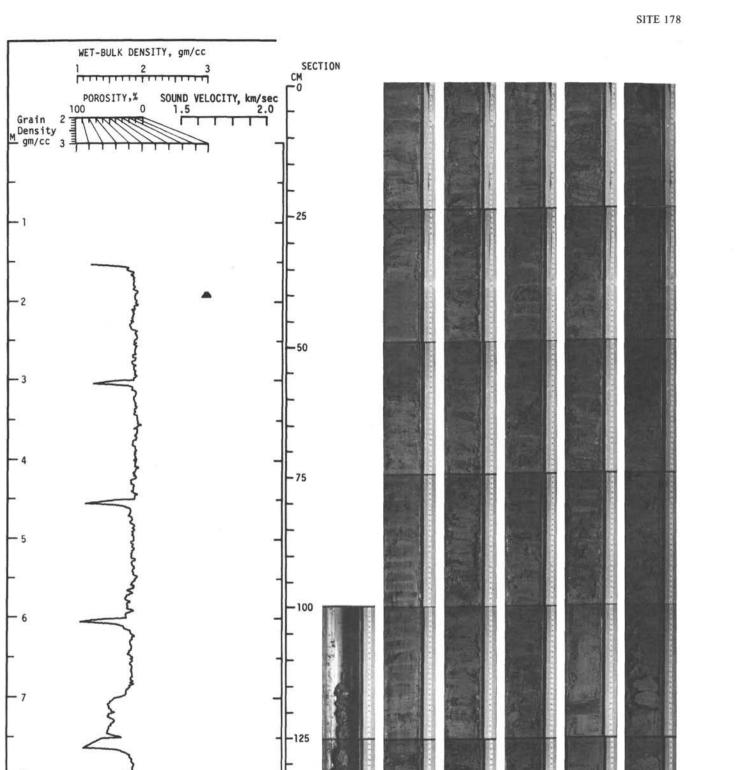


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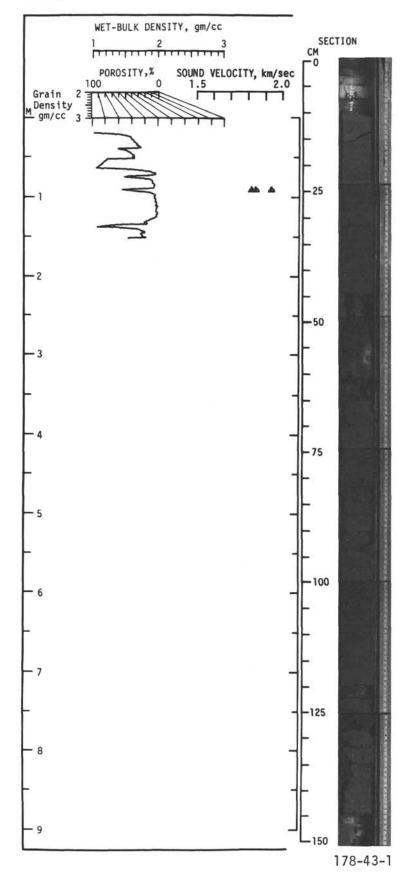


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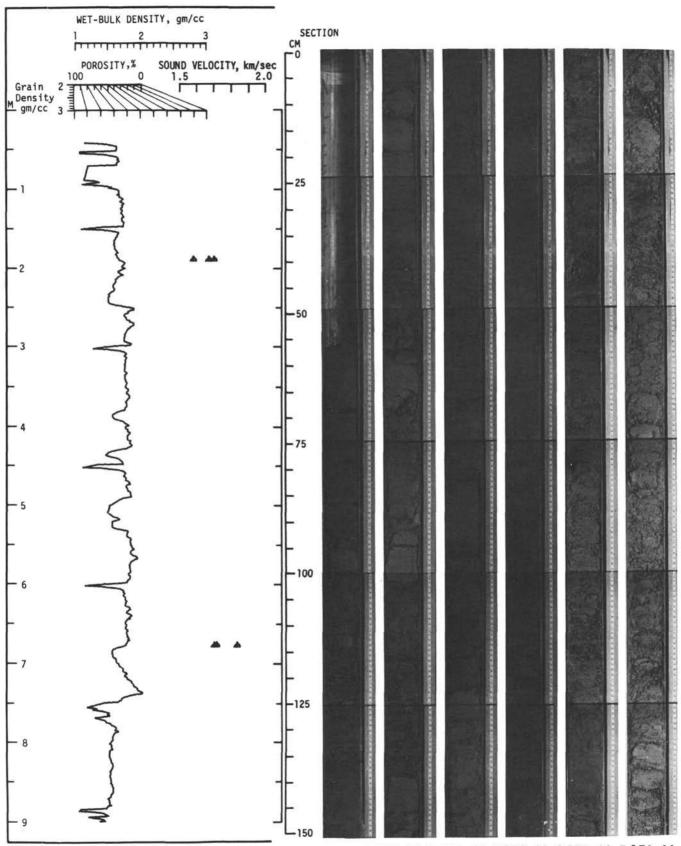


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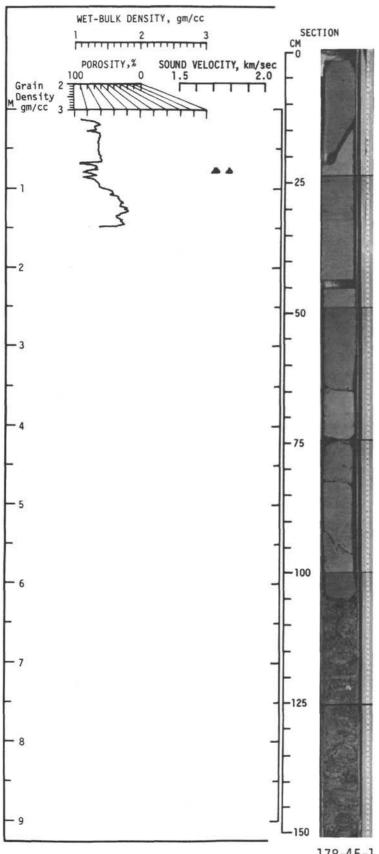
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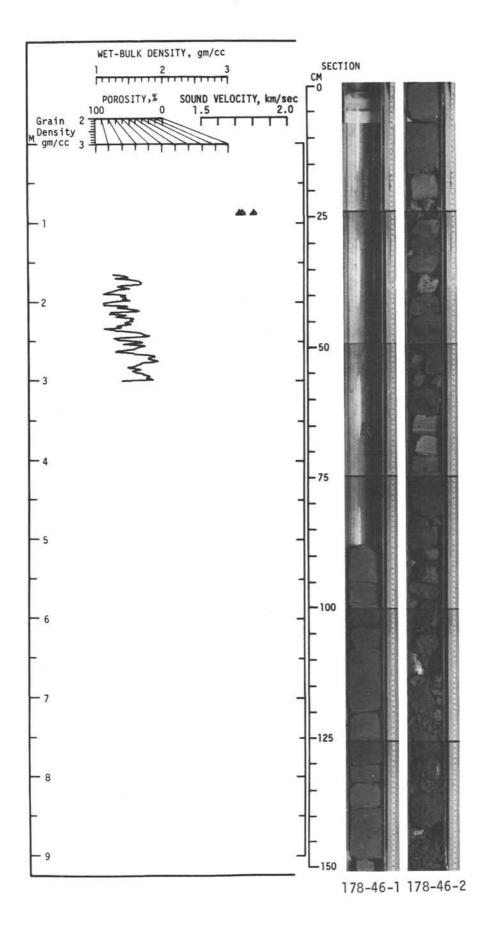
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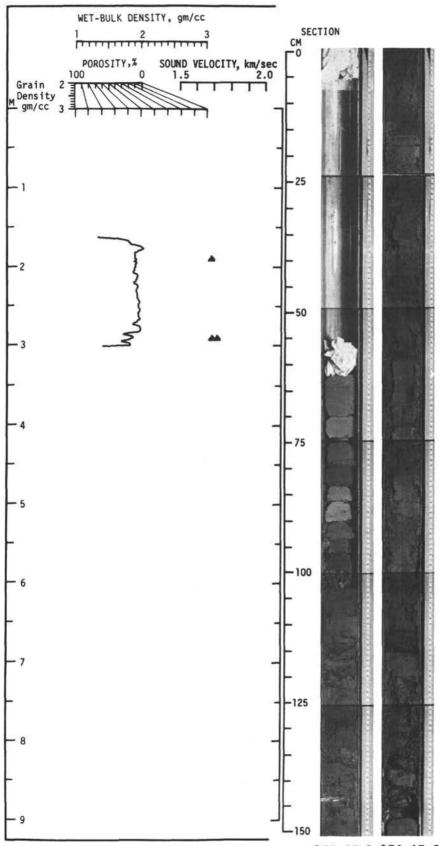


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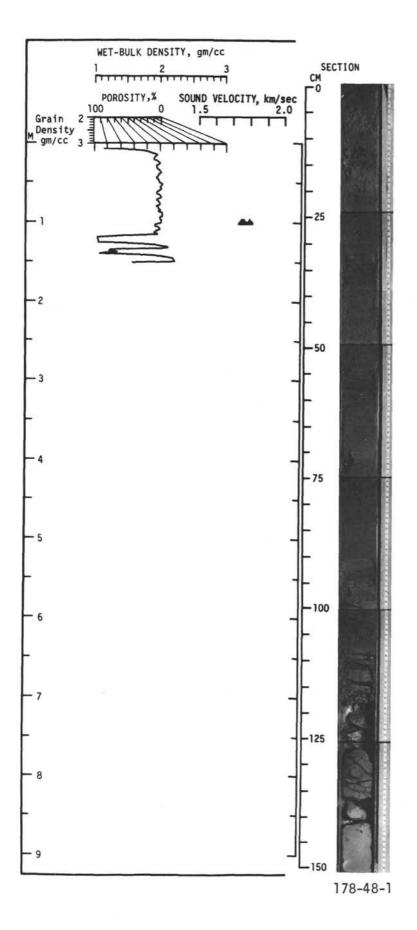


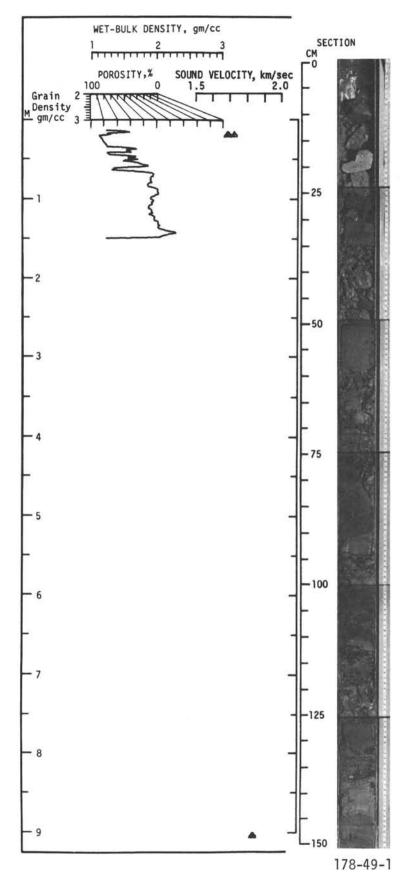
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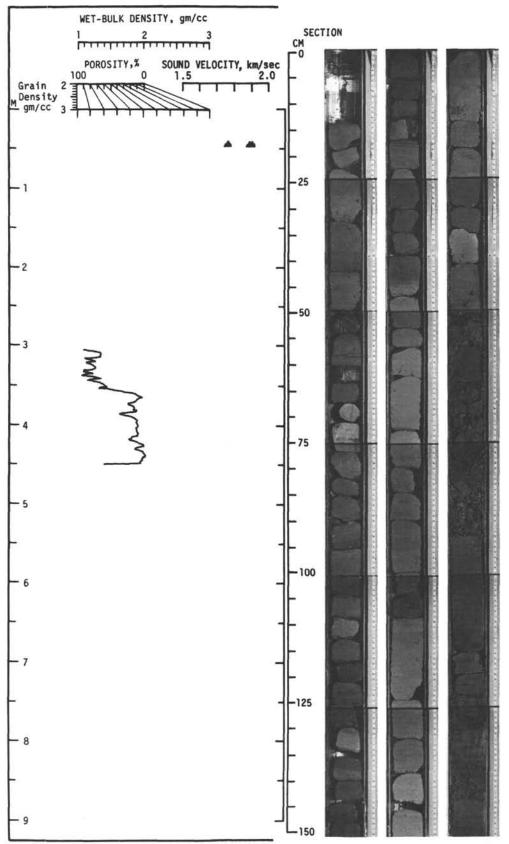




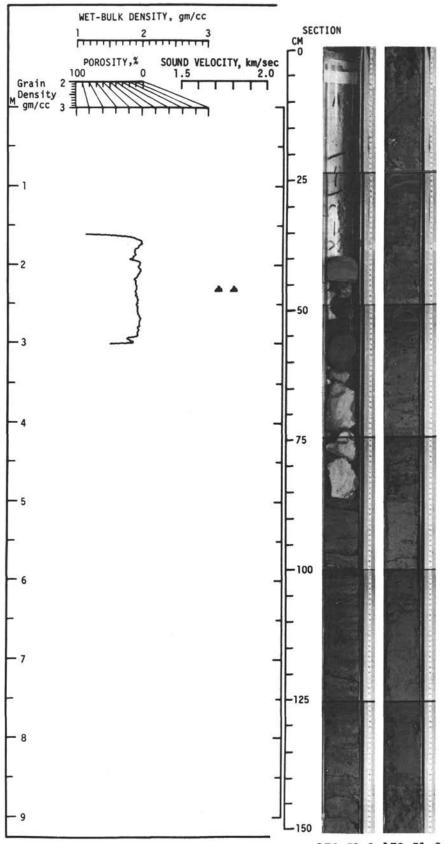
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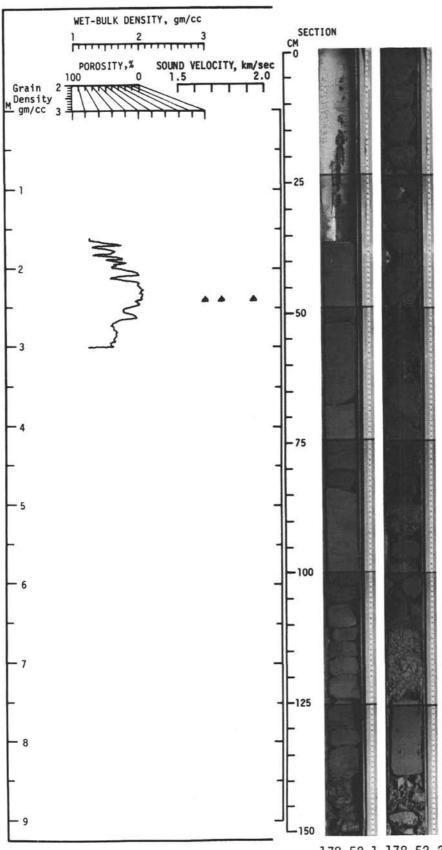




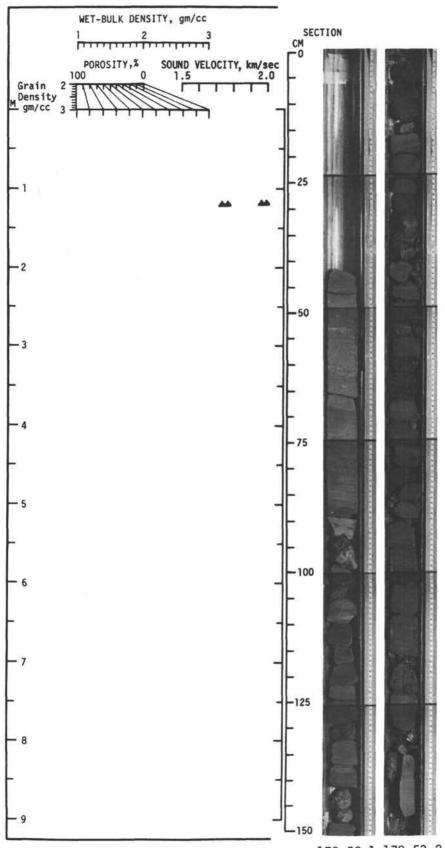
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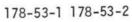


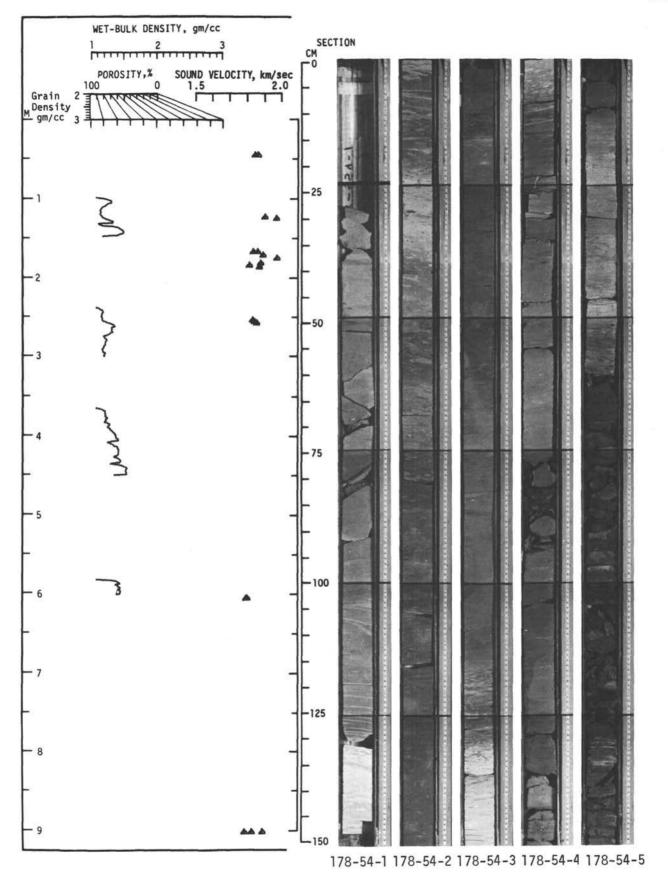
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178-52-1 178-52-2

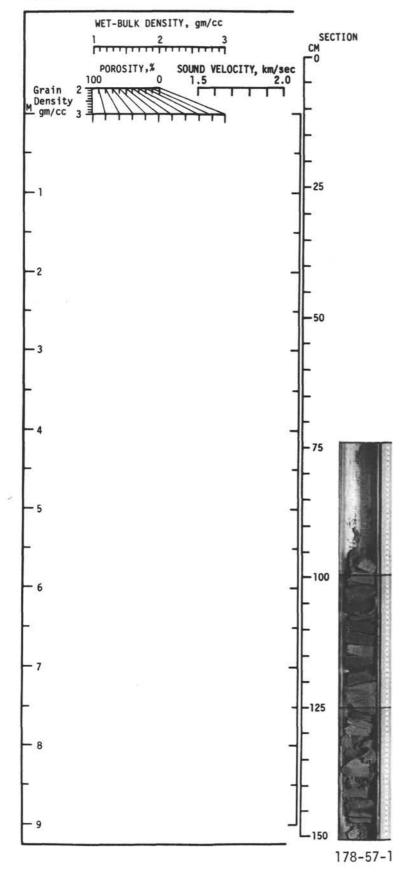


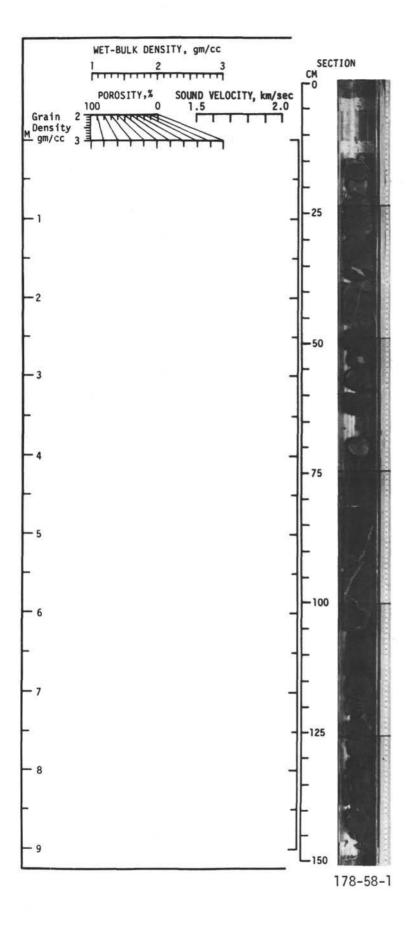


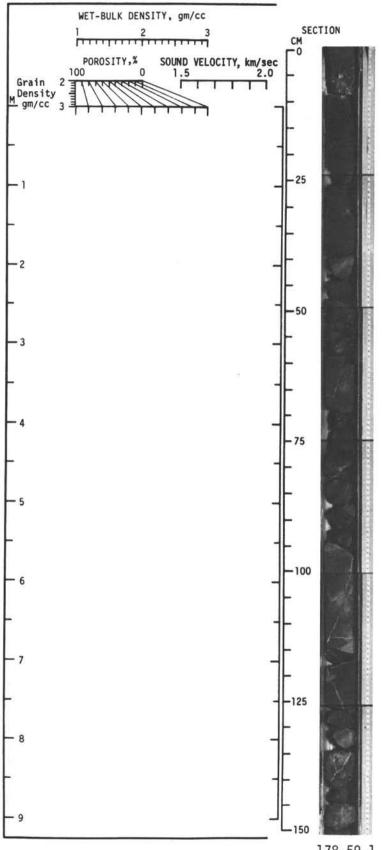


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SITE 178







178-59-1