7. SITE 207

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Location: South Lord Howe Rise Position: 36°57.75', 165°26.06'E Water Depth: 1389 meters Total Penetration: 513 meters

Summary: The upper fossiliferous section represents an incomplete late Pleistocene to latest Cretaceous sequence which includes the most continuous and complete temperate late Cenozoic section collected in the Southern Hemisphere. Below the acoustic basement (309 m) approximately 50 meters of unfossiliferous silty claystone is underlain by rhyolitic lapilli tuffs and vitrophyric rhyolite flows.

BACKGROUND AND OBJECTIVES

General

Site 207 is the southern one of two sites located on the Lord Howe Rise. The rise, with continental crustal thickness, is one of the major structural elements of the Tasman Sea Basin. Data available prior to site survey indicated the presence of a dominant deep acoustic reflector continuous over most of the area, but with localized indications of deeper structure. This site was proposed as a pair of holes—one near the top of the local structure where the dominant reflector was relatively close to the sea floor and another down the flank where an expanded section could be sampled above the dominant reflector. The site, as drilled, is on the southern portion of the Lord Howe Rise, northwest of Bellona Gap and just west of the crest of the rise.

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Contour interval 400 fathoms. (After Mammerickx, Chase, Smith, and Taylor, 1971. Bathymetry of the South Pacific, Charts 11 and 12; Scripps Institution of Oceanography, California.)

Review by the JOIDES Panel on Pollution Prevention and Safety indicated that the site and proposed drilling program were acceptable providing there was continuous coring with maximum drilling precautions. The site was to be abandoned and the hole filled with mud if any detectable evidence of hydrocarbons was observed in the cores.

Site Survey

Site survey for Site 207 was carried out by R/V Kana Keoki during September 1971. The site is located on the southern part of the Lord Howe Rise northwest of Bellona Gap and just west of the rise crest. In the region surrounding the site (Figure 1) the sea floor slopes gently and smoothly to the south.

The acoustic basement structure (Figure 2) is relatively rugged and is dominated by a deep east-west trending basin south of a roughly circular rise which has relief of over 1.0 sec (approximately 1 km). Sedimentation has smoothed and obscured the more rugged basement structure. Deep reflectors in the basin pinch out against the basement highs, while middepth and shallow reflectors converge and continue across the structure (Figure 3). An east-west crossing of the basement high (Figure 4) shows that the sedimentary structures are generally uniform about the high. At the site, it is difficult to distinguish between the major middepth reflector and acoustic basement as the two are close enough to converge. To the south of the site, the deepest reflector (acoustic basement) often disappears, possibly due to steep slopes on the structure.

The drilling site was selected just off the top of the structure to sample as complete a section as possible.

OPERATIONS

Site Approach

Site 207 was drilled near the top of the southern flank of the basement high which had been defined by the site survey. The approach was from the southeast on a track closely parallel to the R/V Kana Keoki survey approach track which permitted occupying the site after one crossing (Figure 5). The technique of dropping the beacon underway provides a very nice tie from the underway profile to the on-site sonobuoy record.



Figure 1. Bathymetry at Site 207 (uncorrected meters). R/V Kana Keoki site survey September 1971.



Figure 2. Acoustic basement structure at Site 207 in seconds of reflection time. R/V Kana Keoki site survey September 1971.

Sonobuoy

The on-site sonobuoy record (Figure 6) correlates well with the underway profiles. The reflector data are summarized in Table $1.^2$ Reflector 6 at 0.37 sec is the major intermediate reflector in the basin south of the site. Reflector 7 at 0.42 sec is underway acoustic basement. South of the site this reflector is intermittent due to slope problems.

Drilling Program

This was the first of the two sites planned on the Lord Howe Rise, and coring was to be continuous following the recommendation of the Safety and Pollution Prevention Panel. Five cores (Table 2) were obtained down to a depth of 47 meters at Hole 207. Following the fifth core an unsuccessful attempt was made to obtain a downhole temperature measurement using the special "stinger" extension. The extension was broken off downhole, and only the instrument package was recovered. To avoid possible trouble from the extension, the bit was raised just above the mud line and Hole 207A was spudded at 1045 on 12 December. Hole 207A (Table 3) produced 50 additional cores with penetration to 513 meters. Coring began at 47 meters below the sea floor and was continuous through the remainder of the section with the exception of four 3 to 5 meter wash intervals made to adjust the drill pipe breaks to the drilling deck level. Additional inhole temperature measurements were made at several levels before the sediment consolidation reached a point which precluded penetration by the probe. These measurements were made with the retractable unit and were incorporated routinely into the program of continuous coring.

Coring operations were terminated with Core 50A prior to reaching any clearly defined basement, but well after penetration of the deepest acoustic reflector observed at the site. The site was abandoned at 1500 on 15 December.

LITHOLOGY

General

Site 207 was continuously cored to a depth of 513 meters below the sea floor. Fifty-five cores were cut (5 in Hole 207 and 50 in Hole 207A). Five core barrels came up empty. A total of 212.3 meters (47%) of core was recovered.

²Final correlations of sonobuoy profiles using laboratorymeasured velocities, other physical properties, and lithologic boundaries are presented in Part II of this Initial Report.



Figure 3. Seismic profile across basement high (SE-NW) at Site 207. Note converging and truncated reflectors. R/V Kana Keoki site survey September 1971.

Summaries of each of the cores are given in Appendix F. The lithologic sequence can be divided into five units. These are, from top to bottom:

1) Unit 1 (0 to 142 m) – middle Miocene to Pleistocene: Foraminiferal nannofossil ooze and nannofossilforaminiferal ooze.

2) Unit 2 (142 to 309 m) Paleocene to middle Miocene. – Foraminiferal-nannofossil ooze to foraminiferal-bearing nannofossil ooze and clay nannofossil ooze (or chalk), with subordinate siliceous fossil-bearing (to rich) foraminiferal-nannofossil ooze.

3) Unit 3 (309 to 357 m) Maastrichtian. – Glauconitic silty claystone (sandstone at the very base).

4) Unit 4 (357 to 433 m). Upper Cretaceous. - Rhyolitic (pumiceous) lapilli tuffs, and vitrophyric rhyolite flows (fragmented in part).

5) (433 to 513 m). Upper Cretaceous or older. – Vitrophyric rhyolite flows, fragmented in part.

Unit I (Foraminiferal-Nannofossil Ooze and Nannofossil-Foraminiferal Ooze)

This unit was penetrated by Cores 1 to 5 and 1A to 10A. It consists of white to very light gray ooze, with streaks and patches of shades in green, blue, and gray. Drilling

disturbance is high. Percentages of foraminifera and nannofossils both vary around the 50% mark. Minor constituents of the sediment are volcanic glass shards and zeolites (up to 2%). At the base of this unit Eocene sediments are intercalated between Miocene sediments (Core 10A). Both have the same macroscopic appearance, and this disconformity could only be detected by paleontological age determinations. Overthrusting or slumping may have caused this disconformity.

Unit 2 (Foraminiferal-Nannofossil Ooze to Foraminiferal-Bearing Nannofossil Ooze and Clay Nannofossil Ooze (or Chalk), with Subordinate Siliceous Fossil-Bearing (to Rich) Foraminiferal-Nannofossil Ooze)

This unit was penetrated by Core 11A to the bottom of Core 28A. The boundary between Units 1 and 2 is the regional unconformity, also found at Sites 206, 208, 209, and 210. Though very similar in macroscopic appearance to Unit 1 (including color), this unit is distinguished by its siliceous content, consisting of siliceous fossils or chert. Siliceous fossils are found in Cores 12A to 22A, in amounts from a few to 35%. They are mainly radiolarians and sponge spicules. Chert fragments (broken up by drilling) are found in Cores 11A and 28A. Except for a small overlap in



Figure 4. East-west seismic profile showing lateral atent of the basement structure at Site 207. R/V Kana Keoki site survey.

Core 22A (in which siliceous fossils occur as a trace only), chert and siliceous fossils are mutually exclusive. This suggests a mobilization and reprecipitation of the biogenic silica. Further support for this is found in the occurrence of partially silicified nannofossil ooze (porcellanite) and ghosts of foraminifera in the chert.

There is a gradual change in lithology downwards. From Core 20A clay-sized, noncalcareous material becomes a component of the sediment, increasing irregularly downwards. X-ray diffractometry and infrared spectrophotometry analyses, however, do not reveal the presence of recognizable clay minerals. Their identity remains unsolved at present. Glauconite in small quantities (up to a few percent) is found in Cores 26A to 28A. Higher content of clay-sized material seems to be associated with greater induration. In Cores 27A and 28A the sediment becomes indurated enough to be called a chalk. Sedimentary structures, mainly burrows and mottling, are better preserved in the more indurated rock.

Unit 3 (Glauconitic Silty Claystone – Sandstone at the Very Base

This unit extends from Core 29A to the top of Core 34A. Transition from Unit 3 to 2 is gradational. The unit

consists of noncalcareous, dark olive-colored silty claystone. Silt-sized clastic particles, recognizable under the light microscope, include quartz, potash feldspar, plagioclase, and glauconite, with subordinate light and dark volcanic glass, mica, and opaque grains (mainly pyrite). Rare sponge spicules, silicoflagellates, and arenaceous foraminifera were observed also. X-ray diffraction indicates the presence of subordinate clay minerals (mainly montmorillonite). Infrared spectrophotometry, however, did not reveal the presence of clay minerals. Both techniques, on the other hand, showed cristobalite to be the main constituent of the sediment. This mineral must be present mainly in clay-sized foraminifera. Tridymite is present in relatively large amounts.

At the base of Unit 3 a glauconitic sandstone layer, a few centimeters thick, occurs directly overlying the volcanic rocks of Unit 4. The sand consists of glauconite grains and of particles derived from the underlying volcanics, such as quartz, potash feldspar, plagioclase, volcanic glass, and volcanic rock fragments. Apart from these, microcline and microperthitic orthoclase grains are also present. They are most likely derived from a plutonic (granitic) or high-grade metamorphic source. Accessory minerals include opaque grains and rare epidote. A large



Figure 5. Seismic profile along approach to Site 207. Glomar Challenger December 1971.

proportion of the sand grains are well rounded, indicating extensive reworking by currents (and/or wave action).

Unit 4 (Rhyolitic, Pumiceous Lapilli Tuffs and Vitrophyric Rhyolite Flows – (Fragmented in Part)

This unit was penetrated in Core 34A to the top of Core 42A. The bulk of the sequence consists of fragmental volcanics. Sizes of fragments vary greatly, from a few millemeters to several centimeters (maximum size observed: 8 cm). Pumice fragments have been observed in some parts (e.g., Cores 39A and 40A). They generally occur in finer grained (<1 cm) parts of the sequence, which are fragments of vitrophyric rhyolite, whose mode of origin is less certain. They most likely represent autobrecciated flows. The ground mass consists of perlitic glass, moderately altered and/or devitrified, and has varying colors (shades of brown, black, gray, and green). Phenocrysts consist of quartz, potassium feldspar, and plagioclase. Potash feldspar content was determined by staining. Universal stage analyses give an average plagioclase composition of An 30 (andesine). Fluidal banding in the fragments is common. Secondary hematite locally gives the rock a bright red color. Secondary pyrite locally lines perlitic cracks.

It is not always possible to clearly distinguish between flow breccias and flows proper. Locally, intervals without or with few fragments are observed (e.g., in Core 35A). These may represent lava flows. The lithology, however, is the same as for the fragments in the brecciated parts. Discontinuous core recovery makes it impossible to distinguish individual units (flows).

Lithoidal rhyolite fragments have been observed in various samples (e.g., from Cores 35A, 39A, and 40A). They are probably accessory fragments from rhyolites crystallized at deeper levels.

Unit 5 (Vitrophyric Rhyolite Flows, Fragmented in Part)

Rocks of this unit were penetrated in Cores 42A to 48A. Lithologically, these volcanics are very similar to those of Unit 4, the only difference being the apparent absence of potash feldspar, phenocrysts, and accessory components. The main differences from Unit 4 are the gross texture of the rocks and the amount of alteration. Texturally, the rocks show far less brecciation. Fluidal banding is common throughout the entire unit. These rocks therefore most likely represent lava flows. The localized fragmented horizons may represent auto-brecciated flows, although discontinuous core recovery makes it impossible to distinguish discrete flows. Vericular rhyolite was noted in Cores 44A and 47A.

The rocks have been devitrified and/or argillized to a far greater extent than in Unit 4. Locally, quartz grain "floating" in a clay matrix is all that remains. The clay is mainly montmorillonite. In Core 47A, however, the clay mineral is halloysite, with minor montmorillonite.

All stages of divitrification and argillization can be observed in Unit 5. Most typical are spherulites which start forming in the centers of the glass fragments, bounded by perlitic cracks. These spherulites consist of a cryptocrystalline intergrowth of cristobalite and anorthocalse. In hand specimens, these spherulites give a "salt and pepper" appearance to the rocks.

Other secondary minerals are a prismatic (blocky) zeolite (clinoptilolite), chalcedony, opal, and chlorite (?).

Phenocrysts consist of quartz and plagioclase and some rare pseudomorphs after ferromagnesian minerals. The composition of the plagioclase is similar to the one in Unit 4 (andesine). No potash feldspar was observed. The high degree of alteration in large parts of the unit, however, makes it difficult to determine if this absence is characteristic for the entire unit.

		Site 207 Sonobuoy I	Data			
Reflector	Depth (sec)	Nature	Estimated Velocity Structure (m/sec)	Estimated Depth (m)		
1	0.17	Weak	1600	136		
2	0.22	Weak	1600	176		
3	0.27	Moderate	1600	216		
4	0.31	Moderate/strong	1600	248		
5	0.34	Strong/moderate	1600	272		
6	0.37	Strong-basin, intermediate reflector	3000	317		
7	0.42	Strong underway acoustic basement	3000	392		
8	0.45	Strong/moderate	4400	493	-	

TABLE 1

Discussion

There are two main items which need further discussion, viz, changes in depth of deposition during the time span represented by the sequence cored and the mode of deposition of the rocks, more particularly the rhyolites.

The rhyolites of Unit 5 do not show clear signs of subaqueous extrusion. However, criteria for this are not clear-cut. The pumiceous lapilli tuffs in Unit 4 provide the best evidence for suggesting that at least part of the rocks in that unit are of subaerial or very shallow marine origin (for further discussion, see Chapter 15).

The composition and texture of the thin sandstone layer at the base of Unit 3 suggest rather strong currents bringing foreign (plutonic or metamorphic) material to the site and locally reworking the underlying volcanics. The rest of Unit 3 is a silty claystone. Its main constituent is α -cristobolite, the origin of which is uncertain. It may be derived from either siliceous fossils or volcanic glass. The fauna in the silty claystone suggests a neritic environment with restricted circulation. Paleontological data suggest a rapid deepening of the site over the interval of Cores 28A to 25A. By middle early Eocene a depth was reached similar to that of the present day. No depth changes since then are suggested by the fauna.

In conclusion, one can say that the data suggest a subsidence at the site from terrestrial (or very shallow water) to present-day depth over a period from at least Upper Cretaceous to early Eocene, remaining constant since then.

Part of the volcanics in Unit 4 are clearly tuffs and are the products of explosive eruptions. The rest of this unit is interpreted as rhyolitic flows, part of which are autobreccias. Although some intervals superficially resemble ignimbrites, detailed study did not reveal positive diagnostic ash-flow characteristics. The volcanics in Unit 5 are more homogeneous than those in Unit 4. They are interpreted as rhyolitic flows, which are autobrecciated in part.



Figure 6. On-site sonobuoy profile at Site 207.

TABLE 2Coring Summary – Hole 207

Core	Date	Time	Depth from Drill Floor (m)	Depth Below Sea Floor (m)	Cored (m)	Recovered (m)	Recovery (%)
1 2 3 4 5	12/12 12/12 12/12 12/12 12/12 12/12	0635 0715 0745 0820 0855	1399-1405 1405-1414 1414-1423 1428-1437 1437-1446	0-6 6-15 15-24 29-38 38-47	6 9 9 9	5.5 7.2 9.0 7.5 8.8	92 80 100 83 98
Total					42	38.0	90

Note: Echo sounding depth (to drill floor) = 1393 meters;

drill pipe length to bottom = 1399 meters.

BIOSTRATIGRAPHY

General

Paleontological examination of cores from Holes 207 and 207A reveals an incomplete latest Cretaceous to late Pleistocene sequence containing abundant and moderately to well-preserved planktonic foraminifera and calcareous nannofossils throughout much of the sequence and abundant Radiolaria in sediments of middle Eocene age. Five distinct episodes of deposition can be distinguished as follows: late Miocene to Pleistocene, early middle Miocene, early to middle Eocene, Paleocene, and Late Cretaceous. A sequence of events is defined as follows:

1) Late Cretaceous deposition of noncalcareous mudstones containing a few poorly preserved late Maastrichtian foraminifera.

2) Paleocene deposition, apparently highly condensed, of about 20 meters of calcareous mudstones which appear to become increasingly more clastic towards the base. These mudstones contain common poorly preserved planktonic foraminifera and nannofloras that are easily correlatable with the zonations of Edwards (1971) and Jenkins (1966). The low numbers of poorly preserved Maastrichtian planktonic foraminifera, calcareous nannofossils, and Late Cretaceous Radiolaria present in most of Core 28A appear to be reworked since this interval contains Paleocene benthonic foraminifera.

3) Early to late middle Eocene deposition of 140 meters of foraminiferal-nannofossil ooze containing abundant, moderately well preserved planktonic foraminifera and nannofloras. The middle Eocene is marked by an abundance of Radiolaria and other siliceous microfossils. Near the base, the presence of several neritic calcareous nannofossil species suggests proximity of this site to land – perhaps the Karamea Peninsula of Fleming (1962, fig. 6). This part of the sequence is readily correlatable with the relevant parts of both the New Zealand Stage classification (Bortonian to Mangaorapan stages) and with the zonal scheme of Edwards (1971).

4) Erosion and/or nondeposition of latest Eocene to early Miocene sediments followed by a sequence of sediments assumed to reflect a complex history of slumping that possibly occurred in the late middle Miocene. The sequence in this interval, based primarily on the foraminiferal sequence is as follows: 207A-CC: Earliest Tongaporutuan to latest Waiauan (early late Miocene to latest middle Miocene).

207A-9-1, 75-77 cm to 207A-9-3, 75-77 cm: Lillburian (middle Miocene)

207A-9-4, 75-77 cm to 207A-9-5, 75-77 cm: Clifdenian (middle Miocene)

207A-9-5, 115-117 cm to 207A-9, CC: Kaiatan (late Eocene); also present is a Miocene foraminiferal assemblage

207A-10-1, 110-112 cm to 207A-10, CC: Late Lillburian (middle Miocene)

207A-11-2, 75-77 cm to 207A-13, CC: Bortonian (late middle Eocene)

An age inversion occurs in this sequence with Late Lillburian sediments immediately overlying the Bortonian sequence. The disturbed sequence from 207A-9-1, 75 cm to 207A-10, CC appears to be best explained by a complex sequence of slumping which took place in the Late Lillburian (late middle Miocene) during which time slump blocks of various ages were transported to the area in the vicinity of Site 207. This event may in part be related to an interval of intense andesitic volcanism and tectonism in Northland, New Zealand, represented by the Waitamata Group (including Manukau Breccia) of early Miocene age followed by massive and widespread slumpiing (Onerahi Chaos) involving redeposition of Late Cretaceous to early Miocene strata.

5) Late Miocene to Pleistocene continuous deposition of about 120 meters of foraminiferal nannofossil oozes containing foraminifera and calcareous nannofossils with temperate affinities. As in Site 206, this part of the sequence has an upward increasing sedimentation rate (see Table 4) and does not have a New Zealand on-land lithologic counterpart. Planktonic foraminiferal frequency changes reflect oscillations of subtropical and subantarctic water masses throughout the sequence and indicate that cold (glacial) cycles have occurred since the late Miocene. Correlation with the Pliocene-Pleistocene of New Zealand is difficult because of general lack of New Zealand microfossil biostratigraphy within this interval. Climatic cycles within the late Miocene enable correlation with the Kapitean and Tongaporutuan stages.

Paleodepth Interpretation

Benthonic foraminifera were examined in eight samples ranging in age from the Late Cretaceous to the late

SITE 207

TABLE 3Coring Summary – Hole 207A

Core	Date	Time	Depth from Drill Floor (m)	Depth Below Sea Floor (m)	Cored (m)	Recovered (m)	Recovery (%)	
1	12/12	1140	1446-1455	47-56	9	8.6	96	
2	12/12	1214	1455-1464	56-65	9	1.0	11	
3	12/12	1250	1464-1473	65-74	9	9.0	100	
4	12/12	1320	1473-1482	74-83	9	6.5	72	
5	12/12	1400	1482-1491	83-92	9	9.0	100	
6	12/12	1440	1491-1500	92-101	9	8.4	93	
7	12/12	1505	1500-1509	101-110	9	9.0	100	
8	12/12	1555	1509-1518	110-119	9	8.5	94	
9	12/12	1630	1518-1527	119-128	9	9.3	100	
10	12/12	1710	1532-1541	133-142	9	5.3	59	
11	12/12	1755	1541-1550	142-151	9	6.1	68	
12	12/12	1855	1550-1559	151-160	9	1.0	11	
13	12/12	1935	1559-1568	160-169	9	8.0	89	
14	12/12	2020	1568-1577	169-178	9	1.9	21	
15	12/12	2100	1577-1586	178-187	9	6.0	67	
16	12/12	2145	1586-1595	187-196	9	3.8	42	
17	12/12	2230	1595-1604	196-205	9	3.5	39	
18	12/12	2315	1604-1613	205-214	9	8.9	99	
19	12/12	2355	1613-1622	214-223	9	77	86	
20	12/13	0035	1622-1631	217-223	9	9.0	100	
21	12/13	0115	1636-1645	237-246	9	8.8	98	
22	$\frac{12}{13}$	0200	1645-1654	246-255	9	1.0	11	
23	$\frac{12}{13}$	0250	1654-1663	255-264	9	8.0	89	
24	$\frac{12}{13}$	0340	1663-1672	255-204	9	6.3	70	
25	$\frac{12}{13}$	0430	1672-1681	273,282	9	5.4	60	
26	$\frac{12}{13}$	0520	1681-1600	275-202	0	4.8	53	
27	$\frac{12}{13}$	0645	1690-1699	202-201	9	24	27	
28	$\frac{12}{13}$	0750	1600-1708	291-300	0	0.7	8	
20	$\frac{12}{13}$	0030	1708-1717	300-309	0	2.7	30	
30	$\frac{12}{13}$	1020	1717-1726	318-327	0	1.5	17	
31	$\frac{12}{13}$	1140	1720-1720	220.320	9	1.5	51	
32	$\frac{12}{13}$	1240	1729 1747	220 348	9	4.0	0	
32	$\frac{12}{13}$	1240	1747-1756	210 257	0	3.6	40	
33	$\frac{12}{13}$	1505	1756 1765	257 366	9	3.0	33	
35	$\frac{12}{13}$	1620	1765 1774	357-300	9	2.0	33	
36	$\frac{12}{13}$	1740	1703-1774	275 294	9	2.9	32	
37	$\frac{12}{13}$	1940	1792 1703	201 202	0	0.4	4	
38	$\frac{12}{13}$	1040	1702 1901	202 402	9	0.0	1	
30	$\frac{12}{13}$	2025	1901 1910	402 411	9	0.1	7	
39 40	12/13 12/12	2025	1001-1010	402-411	9	0.0	0	
40	12/13 12/14	2130	1010-1019	411-420	9	0.0	9	
41	12/14 12/14	0120	1019-1020	420-429	9	0.0	22	
42	12/14 12/14	0130	1831-1840	432-441	9	2.1	23	
43	12/14 12/14	0245	1840-1849	441-450	9	0.4	4	
44	12/14	0705	1049-1058	450-459	9	1.4	10	
43	12/14	0705	1838-186/	459-468	9	4.5	50	
40	$\frac{12}{14}$	1220	100/-10/0	408-4//	9	0.0	0/	
4/	12/14	1230	18/0-1885	4//-480	9	3.3	51	
48	12/14	1/05	1885-1894	486-495	9	0.5	12	
49	12/14	1840	1894-1903	495-504	9	0.0	0	
50	12/14	2130	1903-1912	504-513	9	0.0	0	
Total					450	212.3	47	

Note: Echo sounding depth (to drill floor) = 1393 meters; drill pipe length to bottom = 1399 meters.

Pleistocene. Cretaceous paleodepth determination is difficult because the benthonic foraminifera have little relationship with present-day and Cenozoic forms. The rarity of planktonic foraminifera and absence of calcareous nannofissils may suggest restriction of oceanic circulation which is generally associated with shelf deposition.

Paleocene benthonic foraminifera (207A-27, CC and 207A-26, CC) are distinctly shallower than all younger

sediments. based on relatively high frequencies of discorbids, polymorphinids, *Cibicides* and *Anomalinoides*, and higher frequencies of *Bolivina*. Planktonic foraminifera and calcareous nannofossils have much lower frequencies than in the younger sediments, suggesting shallower depths. Upper bathyal depths are suggested by the foraminifera.

Rapid increase in depth of deposition to middle bathyal depths occurred within the early Eocene (207A-25, CC).

TABLE 4 Sedimentation Rates in the Late Cenozoic

	Time Represented	Sediment Thickness (m)	Sedimentation Rate m/10 ⁶ yr.	
Pleistocene	1.8	27	15	
Pliocene	3.7	50	14	
Late Miocene	4.5	45	at least 10	

These depths have remained much the same until the present day with only minor fluctuations.

Thus, a major subsidence appears to have taken place at Site 207 during the Paleocene and earliest Eocene with little vertical movement since.

Foraminifera

The 33 fossiliferous cores obtained during continuous coring at Site 207 represent an incomplete latest Cretaceous to late Pleistocene sequence of abundant and well-preserved planktonic foraminifera. This sequence will be of special value in providing additional resolution for planktonic foraminiferal biostratigraphy, evolution, and paleoclimatology in temperate waters of the Southern Hemisphere. The planktonic foraminiferal sequence is readily discussed under six age groupings as follows:

Pleistocene, Pliocene, and Late Miocene:

An apparently continuous sequence of over 120 meters (Cores 1 to 5 and 1A to 8A) of foraminiferal-nannofossil oozes containing extremely abundant, well-preserved planktonic foraminifera. This sequence represents the most continuous, complete temperate late Cenozoic section in the Southern Hemisphere. Tropical elements are essentially absent at this site. The relative closeness of Site 207 to the Subtropical Convergence is of particular interest, as fluctuations in subtropical and subantarctic water masses since the late Miocene are clearly reflected by frequency changes in the planktonic foraminifera. Quantitative work is required paleoclimaticdetailed evaluation of the for paleooceanographic oscillations. However, one trend that is readily apparent is that a distinctly cool interval marks the latest Miocene in 207A-4-3, 75 cm to 207A-5-1, 75 cm. During this interval, subantarctic waters migrated northward over Site 207, reflecting a cooling equivalent to those of glacial stages of the Pleistocene. This cooling has been detected previously in New Zealand sections much further to the south during the Kapitean Stage (Kennett, 1967; 1968). Representation of this cooling event at the lower latitude of Site 207 is thus of additional value in the evaluation of the magnitude of these and other oceanographic changes in the late Cenozoic. Coiling direction switches in Globigerina pachyderma are valuable guides to paleooceanographic change.

The Pliocene-Pleistocene boundary (between 207-3, CC and 207-4-1, 50 cm is marked by the first appearance of *Globorotalia truncatulinoides* and the Miocene-Pliocene boundary (between 207A-4-2, 75 cm and 207A-4-3, 75 cm by the appearance of nonkeeled *Globorotalia* as in New Zealand on-land sections. Earlier late Miocene strata

equivalent to the Tongaporutuan Stage contain a warmerwater fauna than that of the latest Miocene (Kapitean Stage). The late Miocene cooling represented at Site 207 represents the first late Cenozoic cold interval and is synchronous with the first major late Cenozoic buildup of ice on Antarctica.

Early Late Eocene and Early Middle Miocene

A disturbed interval of the sequence is represented by Cores 9A and 10A with middle middle Miocene immediately overlain by the late Miocene and underlain by early late Eocene sediments, which are in turn underlain by late middle Miocene sediments. These in turn are underlain by undisturbed early to middle Eocene sediments. Within the disturbed sequence the *Orbulina* bioseries is well represented in the section from 207A-9-1, 75 cm to 207A-9-5, 75 cm.

Early to Middle Eocene (207A-11-2, 75 cm to 207A-25, CC)

A fine, apparently continuous sequence of abundant, well-preserved planktonic foraminifera can be correlated with the Bortonian (late middle Eocene) to Mangaorapan (middle early Eocene) stages of New Zealand.

Late Paleocene to Early Eocene

Interval 207A-26-3, 80 cm to 207A-26-4, 77 cm contains a rich and varied assemblage of planktonic taxa. Taxa present include Globorotalia (Morozovella) dolabrata Jenkins, G. (M.) aequa marginodentata Subbotina, Globorotalia (Planorotalites) australiformis Jenkins, and Truncorotaloides pseudotopilensis Subbotina. This assemblage is Waipawain age. There is a considerable amount of Upper Teurian contamination in this interval.

Paleocene

Interval 207A-27-1, 126 cm to 207A-27-2, 72 cm contains much richer assemblages of foraminifera than those underlying this interval. Agglutinated taxa are uncommon, calcareous benthonic taxa are dominant, and planktonic taxa are quite common. Taxa present include Pseudoclavulina anglica anglica Cushman, Citharina rakauroana (Finlay), Neoflabellina semireticulata (Cushman and Jarvis), Gavelinella becarriiformis (White), Alabamina creta (Finlay), Buliminella creta (Finlay), Tappanina selmensis (Cushman), Globigerina (Subbotina) triloculinoides Plummer, Globorotalia (Turborotalia) pseudobulloides (Plummer), Globorotalia (Acarinina) mckannai (White), G. (A.) acarinata (Subbotina), and a single specimen of Globorotalia (Planorotalites) pseudomenardii Bolli. This assemblage is clearly Mid to Upper Teurian in age equivalent to the lower part of the G. triloculinoides Zone of Jenkins (1966).

Interval 207A-28-1, 137-139 cm contains a varied assemblage of agglutinated and calcareous benthonic Teurian taxa, including *Pseudoclavulina anglica* Cushman, *Gaudryina whangaia* Finlay, *Alabamina creta* (Finlay), and *Gavelinella beccariiformis* (White). No planktonic foraminifera were extracted from this interval, which is probably Lower Teurian in age and in part equivalent to Jenkins, *G. pauciloculata Zone*.

The major change of lithology and microfauna (foraminifera) which occurs immediately below this interval is taken at the Maastrichtian-Paleocene boundary.

Maastrichtian (207A-29-1, 20 cm to 207A-34-1, 8 cm)

The interval 207A-33-3, 40 cm to 207A-34-1, 8 cm contains no foraminifera. Interval 207A-29-1, 20 cm to 207A-33-2, 119 cm contains *Hyperammina elongata* Brady, *Ammodiscus pennyi* Cushman and Jarvis, *Cyclammina* cf. *elegans* Cushman and Jarvis, *Gaudryina healyi* Finlay, and *Dorothia elongata* Finlay. The latter two species indicate a Haumaurian age. One sample within this interval (207A-31-2, 116-118 cm) contains *Heterohelix globulosa* (Ehrenberg), *Globigerinelloides volutus* (Shite), and *Hedbergella monmouthensis* (Olsson), indicating that at least the 207A-29-1, 20 cm to 207A-31-2, 118 cm interval falls within the *G. (R.) circumnodifier* Zone of Webb (1966).

Reworked Maastrichtian planktonic foraminifera and molluscan prisms into Paleocene sediments are noted in Samples 207A-27-2, 70-72 cm and 207A-28-1, 137-139 cm. The latter sample includes specimens of G. (R.) circumnodifier (Finlay) and G. (A.) mayaroensis Bolli. Preservation of calcareous taxa is poor, and much of the fauna has obviously been destroyed since deposition.

Calcareous Nannofossils

Materials from Holes 207 and 207A yielded nannofloras indicating the presence at this site of an incomplete latest Cretaceous to sub-Recent sequence which, particularly in the lower part, closely resembles, both lithologically and paleontologically, the age-equivalent sequences of eastern New Zealand. The age groupings identified are, in downhole order, as follows:

Pleistocene (207-1-1, 60 cm to 207-4-1, 25 cm: Nannofossil-Foraminiferal Ooze)

The Pleistocene sediments contain abundant, wellpreserved calcareous nannofossils. Species commonly present are: Emiliania huxleyi, Gephyrocapsa oceanica, other Gephyrocapsa spp., Pseudoemiliania lacunosa, Coccolithus pelagicus, Cyclococcolithus leptoporus, Helicopontosphaera kamptneri, and Cyclococcolithina macintyrei. Other species present in small quantities, or as isolated specimens, are: Rhabdosphaera claviger, Syracosphaera pulchra, Pontosphaera alboranensis, Pontosphaera distincta, Pontosphaera japonica, Pontosphaera pacificus, Scapholithus ganerotus, Cyclolithella anulla, Oolithotus antillarum, Discoaster perplexus, Scyphosphaera cf. campanula, and a holococcolith species.

Considerable fluctuations in environmental conditions have occurred during the Pleistocene, as shown by the changes in the calcareous nannofossil assemblages. The environment of deposition was mid-subtropical in the lowermost Pleistocene sediments (Sections 3-4 and 3-5), corresponding to a present-day latitudinal position slightly north of the location of Hole 207. This changed through a warm subtropical (Section 3-4) to a marginal tropical/warm subtropical environment (Section 3-3). A warm-water environment persisted for a period (Section 3-3 to 2-2) but with several slightly cooler fluctuations (Section 3-2 and 3-5). A gradual cooling through a warm subtropical (Section 2-2) to a cool subtropical environment followed (Section 2-1). This cool subtropical environment persisted throughout the period of uppermost sediment deposition and corresponds to conditions found in the most recent sediments of this latitude (Burns, in press).

The Pleistocene sediments can be fitted to the zonal scheme of Martini (1971) as follows: 207-1-1, 60 cm, NN21; 207-1-2, 25 cm to 207-1-3, 25 cm; NN20; and 207-1-4, 25 cm to 207-4-1, 25 cm; NN19.

Late Pliocene (207-4-2, 25 cm to 207-5-6, 25 cm; Nannofossil-Foraminiferal Ooze)

The late Pliocene sediments contain abundant, moderately preserved calcareous nannofossils. Species commonly present are: Gephyrocapsa spp., Pseudoemiliania lacunosa, Coccolithus pelagicus, and Cyclococcolithina macintyrei. Other species present in small numbers are: Discoaster brouweri, Discoaster pentaradiatus, Discoaster surculus, Umbilicosphaera mirabilis, Helicopontosphaera kamptneri, Syrocosphaera pulchra, Pontosphaera spp., Rhabdosphaera claviger, Cyclococcolithus leptoporus, Scyphosphaera apstenii, and Scyphosphaera campanula.

The late Pliocene sediments of Hole 207 are characterized by an extremely low content of discoasters. Discoaster content of the sediments is often as low as 2% to 5% of the total assemblage and hence makes stratigraphical zonal definition difficult. The zonal scheme of Martini has been fitted to the late Pliocene sediments, for the present, by placing the boundaries at the first or last appearance of a few good, well-preserved zonal index calcareous nannofossils. It is, however, clear that the zonal scheme for such southern latitude sequences requires further consideration.

The low numbers of discoasters in the sediments indicate a late Pliocene environment of deposition in generally cool southern subtropical conditions.

The late Pliocene sediments have been conditionally fitted to the zonal scheme of Martini (1971) as follows: 207-4-2, 25 cm to 207-4-3, 25 cm, NN18; 207-4-4, 25 cm to 207-4-5, 25 cm, NN17; and 207-4-CC to 207-5-6, 25 cm, NN16.

Early Pliocene to Middle Miocene (207-5, CC to 207A-10, CC; Foraminiferal-Nannofossil Ooze)

This interval contains abundant, more or less moderately preserved diverse unwinnowed nannofloras having temperate affinities and therefore less easily correlated with the zonal scheme of Martini (1971). Correlation with the age-equivalent strata of New Zealand is at present also difficult to achieve due to lack of detailed knowledge of the on-land sequence, but present indications are that the latter part of this sequence will correlate much more easily with those of northern New Zealand than with those of central and southern New Zealand. For convenience, a slumped block of late Eocene age within this interval is dealt with separately below.

Late Eocene (207A-9-5, 120 cm to 207A-9, CC; Foraminiferal-Nannofossil Ooze)

Taxa present in this interval, represented by three richly nannofossiliferous samples. include: Chiasmolithus oamaruensis, Coccolithus eopelagicus, C. ovalis, Cyclicargolithus reticulatus, Discoaster tani nodifer, D. saipanensis, Reticulofenestra bisecta, R. hampdenensis, R. placomorpha, Spenolithus morifermis, Thoracosphaera prolata, Zygrhablithus bijugatus (abundant), and rhabdoliths. Clearly these assemblages conform to the late Eocene 2/5 (mid-Kaiatan) Chiasmolithus oamaruensis zones of both Edwards (1971) and Martini (NP18, 1971). Their disordered position is ascribed to mid Miocene slumping as a semiconsolidated block. Whatever, their fortuitous occurrence in this sequence is particularly interesting for, being significantly younger than any other Eocene assemblages present encountered, they help fill in part of a very large hiatus at this site (see below). If, as seems likely, the Eocene-Oligocene unconformity of Sites 206 and 208 to 210 (q.v.) was formerly present in this sequence, then the occurrence of these assemblages would imply that the event causing it occurred subsequent to the Late Eocene 2/5 interval. No obvious indication of stratigraphic direction could be determined from these assemblages.

Major Unconformity (Between 207A-10, CC and 207A-11-2, 30 cm)

An unconformity, involving all of the mid Eocene 7/7 to early Miocene 5/5, interval, occurs between Cores 10A and 11A as judged by a major floral turnover.

Mid Eocene (207A-11-2, 30 cm to 207A-24-4, 25 cm; Foraminiferal-Nannofossil Ooze to Nannofossil Ooze with Chert)

Taxa present in this interval, which contains moderately well preserved, abundant nannofloras throughout, include: Braarudosphaera bigelowi (top at 207A-21, CC); Chiasmolithus expansus; C. grandis; C. solitus; Chiphragmalithus cristatus (base at 207A-20-3, 25 cm and sporadic in Cores 12A and 13A); C. cristatus var. (cf. Nannotetrina fulgens, 207A-14, CC to 207A-16-1, 100 cm); Coccolithus eopelagicus; C. ovalis s.1.; Discoaster barbadiensis; D. distinctus, D. lodoensis (sporadic above 207A-21, CC); D. sublodoensis (207A-18-6, 100 cm to 207A-24-4, 25cm); Discoasteroides kuepperi (top at 207A-21-2, 100 cm); Reticulofenstra dictyoda; R. hampdenensis (base at 207A-12-1, 100 cm); R. placomorpha (base at 207A-16-2, 100 cm); Spenolithus moriformis; S. radians (top at 207A-22, CC); Zygolithus dubius; and Zygrhablithus bijugatus (abundant). The following biostratigraphic assignments are made, in downhole order, with the zonal scheme of Edwards (1971):

1) 207A-11-2, 30 cm to 207A-12-1, 100 cm: mid Eccene 6/7 (mid to early Bortonian) combined *Discoaster distinctus* and *Reticulofenestra hampdenensis* zones. Probably only the equivalent of the latter (early Bortonian) zone is present, but this cannot be confirmed due to the apparent unreliability of the top *Chiphragmalithus cristatus* datum level in this sequence.

2) 207A-12, CC to 207A-16-2, 100 cm: mid Eocene 5/7 (early Bortonian to late Porangan) part of the *Chiphragma-lithus cristatus* Zone.

3) 207A-16, CC to 207A-18-6, 25 cm: mid Eocene 4/7 (early Porangan) part of the *Chiphragmalithus cristatus* Zone.

4) 207A-18-6, 100 cm to 207A-20-3, 25 cm: mid Eocene 3/7 (late Heretaungan) part of the *Chiphragma-lithus cristalus* Zone.

5) 207A-20-3, 100 cm to 207A-21-1, 60 cm: mid Eocene 2/7 (mid Heretaungan) *Discoaster elegans* Zone.

6) 207A-21-1, 100 cm to 207A-24-4, 25 cm: mid Eocene 1/7 (early Heretaungan and possibly late Mangaorapan) part of the *Reticulofenestra dictyoda* Zone. The top *Spenolithus radians* datum level (207A-22, CC) may provide a useful further subdivision.

Judging from the above, a complete open-water mid Eocene 1/7 to 6/7 (early Heretaungan to early Bortonian) sequence is present. Despite this, accurate correlation with the zonal scheme of Martini (1971) is not possible; intervals 1 to 3 and 4 to 6 above appearing to be approximate correlatives of this *Chiphragmalithus alatus* (NP15) and *Discoaster sublodoensis* (NP14) zones, respectively. Even so these broad correlations are more accurate than those given by Edwards (1971) based on New Zealand material.

Early Eocene (207A-24-4, 100 cm to 207A-26-3, 25 cm; Nannofossil Ooze with Chert)

Taxa present in this interval, which contains abundant but variably preserved nannofloras throughout, include: Braarudosphaera bigelowi (base at 207A-25, CC); Chiasmolithus grandis, C. solitus; Coccolithus cavus s.l.; C. eopelagicus; Discoaster barbadiensis (base about 207A-24, CC); D. diastypus (top at 207A-26-2, 145 cm); D. distinctus (base about 207A-24, CC); D. lodoensis (base at 207A-25-4, 100 cm; Discoasteroides kuepperi (base at 207A-26-3, 66 cm), Marthasterites tribrachiatus (top at 207A-24, CC); Reticulofenestra dictyoda (base at 207A-25-1,100 cm); Spenolithus moriformis; S. radians (base at 207A-26-3, 25 cm); and Zygrhablithus bijugatus (dominant). The following biostratigraphic assignments are made, in downhole order:

1) 207A-24-4, 100 cm to 207A-24-5, 135 cm: early Eocene 5/5 (late Mangaorapan) part of the *Reticulofenestra* dictyoda Zone of Edwards (1971) plus the whole of the Discoaster lodoensis (NP13) Zone of Martini (1971).

2) 207A-24, CC to 207A-25-1, 100 cm: early Eocene 5/5 (late Mangaorapan) part of the *Reticulofenestra* dictyoda Zone of Edwards (1971) plus part of the Marthasterites tribrachiatus (NP12) Zone of Martini (1971). Note the different (improved) correlation between these two zonal schemes and that given by Edwards (1971).

3) 207A-25-2, 100 cm to 207A-25-4, 100 cm: early Eocene 4/5 (early Mangaorapan) *Discoaster lodoensis* Zone of Edwards (1971) and part of *Marthasterites tribrachiatus* (NP12) Zone of Martini (1971).

4) 207A-25, CC to 207A-26-3, 66 cm: early Eocene 3/5 (early Mangaorapan to late Waipawan) part of the *Chiasmolithus grandis* Zone of Edwards (1971) and, based on sequential position, the *Discoaster binodosus* (NP11) Zone of Martini (1970).

The above indicates that a complete, but rather condensed (relative to the overlying siliceous microfossil-rich sediments), open-water early Eocene 3/5 to 5/5 (late Waipawan to late Mangaorapan) sequence is present. Particularly notable in these floras is the presence of poorly preserved (overgrown) braarudosphaerid pentaliths and segments. These warm-water taxa have a well-known general restriction to sublittoral and hemipelagic sediments, although the reasons for this distribution are disputed. In this particular case, for reasons given in the report by

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Edwards (Chapter 18), their occurrence is attributed to the presence of large shoal areas around the southernmost part of the New Caledonia Basin (cf. Fleming, 1962, figs. 4 and 6; Stevens and Suggate, in press, figs. 11.19 and 11.20) Perhaps significantly, the lowest sample in which this group was observed, (Sample 207A-25, CC), also contains very rare, badly corroded *Discolithina pulchra*—a "neritic" species. The following species were observed in 207A-25-2, 25 cm, the sample in which braarudosphaerids are present in their greatest numbers (common): *Braarudosphaera bigelowi, B. discula, Micrantholithus* cf. *flos*, and *M. inaequalis*. Also observed was a single, poorly preserved (overgrown) ascidian sclerite cluster (see report by Edwards, Chapter 18).

Paleocene-Eocene Boundary (207A-26-3, 66 cm to 207A-26-4, 25 cm; Nannofossil Ooze)

Taxa present in this interval, represented by four richly nannofossiliferous samples, indicate the repetitive occurrence of two distinctly different open-water assemblage types having different age implications (see below). The reason for this situation is not known, but could be due to slumping following a disconformity involving part or all of the late Paleocene 3/3 to early Eocene 2/5 interval.

In Samples 207A-26-3, 66 cm and 207A-26-4, 25 cm taxa present include Chiasmolithus grandis, C. solitus, Coccolithus cavus, "C." pertusus, Discoaster diatypus, Discoasteroids kuepperi, Marthasterites tribrachiatus (common), and Zygrhabdithus bijugatus (dominant, plus rarer Chiasmolithus consuetus, Markalius astroporus, Spenolithus moriformis, and Zygolithus concinnus. Single specimens, attributed to contamination or reworking, of the late Paleocene index species Discoaster multiradiatus were found in both samples. Apart from this, these assemblages conform to the early Eocene 3/5 (late Waipawan) part of the Chiasmolithus grandis Zone of Edwards (1971) and are probably correlative with the Discoaster binodosus (NP11) Zone of Martini (1971).

In Samples 207A-26-3, 71 cm and 207A-26-3, 125 cm taxa present include Chiasmolithus bidens?, C. sp. cf. grandis, Coccolithus cavus, "C." pertusus, Discoaster diastypus, D. mediosus?, D. multiradiatus, Semihololithus kerabyi, (lower sample only), Toweius eminens, and Zygrhablithus bijugatus (very abundant) plus rarer Biantholithus sparsus, Chiasmolithus consuetus, Discoaster lenticularis, Ellipsolithus macellus, Fasciculithus tympaniformis, Hornibrookina n.sp. (minute), Markalius astroporus, Spenolithus moriformis, and Toweius tovae. These assemblages appear to be correlative of the late Paleocene 2/3(latest Teurian to earliest Waipawan) Discoaster mediosus Zone of Edwards and of the combined Discoaster multiradiatus (NP9) and Marthasterites contortus (NP10) zones of Martini (1971). A single specimen of the early Eocene index species Marthasterites tribrachiatus found in the lower sample is tentatively attributed to contamination. Much coccolith debris is present.

Late Paleocene (207A-26-4, 100 cm to 207A-26, CC; Nannofossil Ooze)

Taxa present in this interval, represented by two richly nannofossiliferous samples, include: *Chiasmolithus bidens*,

"Coccolithus" pertusus, Discoaster multiradiatus, Fasciculithus tympaniformis, Prinsius martinii martinii, Semihololithus kerabyi (very abundant), and Toweius eminens plus rarer Chiasmolithus consuetus, Coccolithus cavus, Discoaster lenticularis, Hornibrookina n.sp. (minute), Markalius astroporus, Prinsius bisulcus, Spenolithus moriformis, Thoracosphaera operculata, and Zygolithus concinnus. These open-water assemblages conform to the late Paleocene 1/3 (late Teurian) Discoaster multiradiatus Zone of Edwards (1971) and appear to be correlatives of the combined Discoaster multiradiatus (NP9) and Marthasterites contortus (NP10) zones of Martini (1971).

Mid Paleocene

(207A-27-1, 126 cm to 207A-207, CC; Nannofossil Chalk)

Taxa present in this interval, represented by four richly nannofossiliferous samples, include: Chiasmolithus bidens, "Coccolithus" pertusus, Fasciculithus tympaniformis, Prinsius martinii, and Toweius tavae plus rarer Coccolithus cavus, Heliolithus kleinpelli, Markalius astroporus, Thoracosphaera operculata, Zygodiscus sigmoides, Zygolithus concinnus, and Z. junctus. These diverse open-water assemblages conform to the mid Paleocene 2/3 (mid Teurian) Heliolithus kleinpelli zones of both Edwards (1971) and Martini (1971; NP6). The core-catcher sample contains clay-grade rounded isotropic clastic debris similar to that characteristic of the underlying intervals (q.v.).

Early Paleocene (207A-28-1, 78 cm; Clay Chalk)

This sample, which immediately overlies a prominent bedding plane (disconformity?), contains a small, low diversity flora which includes rare Chiasmolithus danicus, Coccolithus cavus, and Prinsius martinii plus occasional Conococcolithus panis, Hornibrookina teuriensis (an Austral realm element), Thoracosphaera operculata (fragments only), and Zygodiscus sigmoides. This assemblage conforms to the early Paleocene 4/5 (early Teurian) Prinsius martinii Zone of Edwards (1971) and is correlatable with the combined Chiasmolithus danicus (NP3) and Ellipsolithus macellus (NP4) zones of Martini (1971). No reworking from older strata is evident. The associated debris not demonstrably of biogenic origin includes a very abundant clay-grade rounded isotropic clastic material and abundant fine silt and clay-grade calcite. By analogy with similar Austral regime Paleocene nannoflora/sediment associations in nearby New Zealand, Site 207 was shallow sublittoral (Sensu lato) during the deposition of this sample. The nearest land may have been in the northern South Island area (Fleming, 1962, figs. 4 and 6), some hundreds of kilometers to the southeast.

?Late Maastrichtian

(207A-28-1, 79 cm to 207A-28, CC; Clay and Clay Chalk)

Taxa present in this interval, represented by four samples, include: Arkhangelskiella cymbiformis, Deflandrius spinosus, Kamptnerius magnificus, and Nephrolithus frequens (a well-known Austral and Boreal element) plus rarer Cretarhabdus surirella, Deflandrius cretaceus, Eiffelithus turriseiffeli, Glaukolithus amphipons, Markalius astroporus, and Microrhabdulus decoratus. These assemblages conform to the Maastrichtian 3/3 (late Haumurian) part of the *Nephrolithus frequens* zones of both Edwards (1971) and Cepek and Hay (1969). They are more diverse than the equivalent Austral realm assemblages of nearby New Zealand and probably reflect more open-water conditions.

However, all four samples contain unusually high proportions of coccolith "shard" fragments. Since the associated clastic debris is of clay grade and thus likely to have been deposited under low energy conditions, it seems probable that these assemblages represent the redeposited "fines" resulting from the winnowing of an older deposit. Further support for such an origin is provided by the distribution of the calcareous nannofossils in this interval-they are abundant in the lower part (207A-28, CC and 207A-28-1, 137 cm) but very much rarer in the upper part (Samples 207A-28-1, 85 cm and 207A-28-1, 79 cm). Since the relative topographic position, but not necessarily the depth, of this site is unlikely to have changed since the deposition of this interval, the hypothetical source of the nannofossils need not have been shallower. If this theory is correct, the age of this interval is, on sequential grounds, within the late Maastrichtian (Late Haumurian) to early Paleocene 4/5 (early Teurian). No nannofossils characteristic of the inferred age of this interval were found apart from single, very robust, specimens of the early Paleocene (including Danian) species Conococcolithus panis in the lower two samples. The higher of these samples also contains obvious, but very minor, contamination, such as: Cyclicargolithus sp., Discoaster sp. (6 armed), Fasciculithus tympaniformis (Mid Paleocene), Reticulofenestra sp?. and Spenolithus moriformis, from uphole Cenozoic.

Unfossiliferous Interval (Cores 29A to 33A; Silty Claystone)

This interval, which is noncalcareous, was sampled at 207A-29-1, 80 cm, 207A-29, CC, and 207A-33, CC, but did not yield nannofossils apart from very occasional specimens of Paleogene taxa such as *Reticulofenestra* sp.?, *R. dictyoda, R. placomorpha*, and *Zygrhablithus bijugatus*. Clearly these specimens, which are abundant uphole, represent contamination—a conclusion consistent with the presence in this interval of Maastrichtian foraminifera (Dr. P. N. Webb, New Zealand Geological Survey, Lower Hutt, personal communication, samples from 207A-30-1 and 207A-31-2). In this and the overlying intervals (q.v.), the major clastic component present is rounded isotropic material of clay and fine silt grade.

Radiolaria

The radiolarians are practically absent in the five cores recovered in the first hole of Site 207. Isolated corroded individuals of *Lamprocyclas maritalis*, *Lithomelissa* sp., *Druppatractus* sp., *Spongodiscus* sp., and a few other species have, however, been recognized in 207-1, CC and 207-2, CC.

A rich radiolarian fauna was encountered only in Cores 207A-11, CC through 207A-21-5 (late middle Eocene to early middle Eocene). Below this level in 207A-21, CC and 207A-22, CC they become very rare and poorly preserved. Rare and poorly preserved Upper Cretaceous radiolarians (*Amphipyndax* sp. and *Dictyomitra andersoni*) were found in 207A-28, CC.

In spite of their abundance in the middle Eocene assemblages, no short-ranged species known for the tropical Pacific were found to make possible a zonation of this rather long sequence. Several events observed by the investigation of the core-catcher samples might be useful for such a purpose.

The following species are among the most common throughout most of the sequence: Lychnocanium aff. bellum, Velicucullus magnificus, Petalospyris eupetala, Lophocyrtis biaurita. Other species such as Eusyringium lagena, Phormocyrtis striata, Amphicraspedum prolixum, Dictyomitra sp., Anthocyrtoma sp., and Thyrsocyrtis sp. have been recorded only at some levels. They might be taken into consideration for a future zonation of the middle Eocene middle latitude radiolarians.

PHYSICAL PROPERTIES

Bulk Density

Fore Cores 2 through 27, values of bulk density were calculated from the GRAPE curves. Below Core 27 (300 m) the cores consisted of individual rock fragments. In this situation the GRAPE technique gives unreliable values, and densities were determined by weighing rock samples of known volume. Because of the motion of the ship and the limited equipment available, these determinations are probably not better than $\pm 5\%$. The density values are plotted against subbottom depth in the hole summary illustration.

The sequence cored can be divided, on the basis of density, into four parts, roughly corresponding to major lithologic changes:

1) In the top 50 meters of the sequence the density is low, about 1.6 gm/cc. This represents the essentially unconsolidated part of the sequence.

2) Below 50 meters compaction increases, and the density rises into the range 1.59 to 1.77 gm/cc, although it is quite variable due to disturbance of the sediments by drilling. Density continues in this range down through Core 20A (230 m).

3) From Core 20A to Core 23A the density rises as the sediments become increasingly lithified and cherty to reach a maximum of about 2.09 gm/cc in Core 23A (255 m). Individual chert fragments have densities as high as 2.24 gm/cc. The pronounced change in lithology between Cores 28A and 29A corresponds to a drop in density.

4) In Core 34A the density rises abruptly again accompanied by a similar change in sonic velocity. This corresponds to the top of the rhyolites. From Core 34A to the bottom of the hole (357 to 500 m) the sequence sampled consisted of glassy rhyolites having a mean density about 2.10 gm/cc. The upper part of the rhyolite section (Cores 34A to 36A), with densities ranging from 2.25 to 2.84 gm/cc, is markedly denser than the lower part of the sequence.

Sonic Velocity

Sonic velocity measurements were made on one or two samples from each core. These values are also plotted in the hole summary. The essentially uniform lithology of the upper 250 meters of the sequence cored is reflected in the velocity measurements which give virtually uniform values in the range 1.58 to 1.66 km/sec, although this does rise slightly through the last 50 meters and reaches a value of 1.91 km/sec in Core 25A (275 m). Velocities from Core 25A through Core 33A remain in the range 1.74 to 2.16 km/sec, although the velocity of individual chert fragments is as high as 4.59 km/sec. Velocities in the volcanic sequence are very variable since the rhyolites are quite strongly altered in parts. The mean value is about 3.35 km/sec, but individual values range from 1.94 to 5.32 km/sec. The upper part of the rhyolitic sequence (Cores 34A to 36A) has generally higher velocities (4.45-5.32 km/sec), reflecting the increased density of these cores.

Thermal Conductivity and Heat Flow

Thermal conductivity values on sediment cores recovered from this site range between 2.8 and 3.7 m cal/°C cm sec (TCU), uncorrected for ambient pressure and temperature conditions at the sea floor. The values show an overall increase with depth down to about 290 meters beneath the sea floor, below which the sediment cores are too indurated for the standard needle probe technique. The increase in thermal conductivity with depth corresponds to a similar increase in density, although individual values do not generally show such a correspondence. Probably the variability in both parameters within each core is sufficiently large to obscure any detailed correlation. The relatively high values of thermal conductivity, in comparison with other sites of this leg, is probably the result of higher calcium carbonate content of the sediments compared with other sites.

Due to various operational difficulties, no reliable in situ sediment temperatures were measured downhole at this site. A measurement at 110 meters (Core 8A) gave 7.4 $\pm 0.2^{\circ}$ C, but it seems likely that this was the temperature inside the core barrel before coring, rather than that of the sediment.

SUMMARY AND CONCLUSIONS

Drilling at Site 207 was planned as one of the two holes to investigate the nature and history of the Lord Howe Rise, as well as providing part of the biostratigraphic tie between New Zealand temperate faunas and floras and those of tropical and subtropical latitudes. The site is located on the southern Lord Howe Rise northwest of the Bellona Gap. The lithostratigraphic section provides some indications of the development of the rise.

At the base of the cored sequence are the Upper Cretaceous rhyolites (flows, breccias, and pumiceous lapilli tuffs) of Units 4 and 5. A marked basement high (Figure 4) shown by the site survey may represent the eruptive center.

The overlying silty claystone, also of Upper Cretaceous age (Unit 3), was probably deposited in a shallow marine environment with restricted (nonoceanic) circulation. This interpretation is based on a rarity of planktonic fossils and the fine-grained nature of the sediments. The latter suggests that the source area must have been either of rather low relief or at some distance. Alternatively, the environment of deposition was close to land, but restricted sufficiently to limit the grain size of the sediments. At the base of Unit 3 there is a larger sand-sized fraction in the form of material derived from the underlying rhyolite as well as from a granitic or metamorphic source.

Oceanic conditions began, as seen at the base of Unit 2, in middle Paleocene. Although the clay-sized terrigenous component is still relatively high in the lower portion of Unit 2, Units 1 and 2 are both clearly oceanic in nature. These two units are basically composed of carbonate oozes of Paleocene to Recent age which were deposited well above carbonate compensation depth. Unit 2 contains frequent to abundant siliceous microfossils in its upper part. The regional unconformity at this site is between Units 1 and 2 and separates late middle Eocene from early middle Miocene sediments. No reworked Oligocene microfossils have been found in the early middle Miocene beds. However, early late Eocene microfossils occur above the early middle Miocene fauna, introduced apparently by slumping. This site and Site 208 exhibit the largest time span in the unconformity. A second unconformity is present between Paleocene and Eocene sediments deeper in the section and represents the smallest time loss for this unconformity (latest Paleocene and earliest Eocene) which also occurs at Sites 206 and 208.

Benthonic foraminifera indicate rapid increase in depth of sedimentation from the relatively shallow depths in the Maastrichtian to depths approximately equivalent to those of the present day (1400 m) during the early Eocene.

A general reconstruction of the site history fits well in the regional picture. The rhyolites were erupted at or near sea level 94 m.y. B.P. K/Ar Date (McDougall, personal communication; see, also, Chapter 15, this volume). This is just prior to the earliest development of oceanic crust in the Tasman Sea at 80 m.y. B.P. (Hays and Ringis, 1972), and the activity may be related to early separation of the Lord Howe Rise from Australia. The silty claystone was laid down in the ensuing shallow restricted sea, with at least some of the material being derived from the rise. Oceanic conditions entered the area in latest Cretaceous time, and the rise underwent a slow subsidence from near sea level to the present upper bathyal depth by early Eocene. Relatively stable conditions have persisted since that time.

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NOTE CONCERNING THE APPENDICES

The appendices consist of tables of shore laboratory determinations of grain size, carbon content, and mineralogical composition, summary visual descriptions of the cores recovered from the site, photographs of the cores and, finally, an overall summary of the results of drilling at the site. The symbols used to represent lithology in the core summary forms are explained in Chapter 2 of this volume. The lithologic description of each core contains typical results of shipboard examination of smear slides of each lithology. In order to make the lithologic descriptions more complete we have also included many of the shore laboratory results. These are identified by being placed in square brackets.

APPENDIX A Grain Size Determinations, Holes 207 and 207A

Core, Section, Top of Interval (cm)	Depth (cm)	Sand (%)	Silt (%)	Clay (%)	Classification
Hole 207					
1-4,84.0 2-3,80.0 3-4,40.0 4-2,30.0 5-5,30.0	5.3 9.8 19.9 30.8 44.3	37.0 13.0 34.7 23.2 24.0	20.5 20.6 24.4 26.3 26.8	42.5 66.5 40.9 50.5 49.2	Sand-silt-clay Silty clay Sand-silt-clay Sand-silt-clay Sand-silt-clay
Hole 207A					
1-4,20.0 3-3,40.0 4-4,60.0 5-2,60.0 6-5,30.0	51.7 68.4 79.1 85.1 98.3	11.5 17.5 42.0 23.7 25.3	28.1 24.6 21.8 29.1 33.2	60.3 57.9 36.2 47.2 41.5	Silty clay Silty clay Sand-silt-clay Sand-silt-clay Sand-silt-clay
7-6,75.0 8-6,140.0 9-4,80.0 10-3,60.0 11-4,60.0	109.3 118.9 124.3 136.6 147.1	29.3 27.1 29.6 32.7 18.4	39.8 42.2 43.2 37.4 45.1	31.0 30.6 27.2 29.8 36.5	Sand-silt-clay Sand-silt-clay Sand-silt-clay Sand-silt-clay Clayey silt

APPENDIX A – Continued

Core, Section, Top of Interval (cm)	Depth (cm)	Sand (%)	Silt (%)		Classification
Hole 207A –	Continued				
13-6.60.0	168.1	4.9	45.4	49.6	Silty clay
14-2,30.0	170.8	4.1	49.6	46.3	Clayey silt
15-2,60.0	180.1	1.4	38.0	60.6	Silty clay
16-2,90.0	189.4	2.6	52.2	45.2	Clayey silt
17-2,70.0	198.2	5.5	40.1	54.3	Silty clay
18-3,35.0	208.4	4.4	46.4	49.2	Silty clay
20-3,48.0	226.5	4.4	51.2	44.4	Clayey silt
21-3,60.0	240.6	2.7	46.6	50.7	Silty clay
25-4,50.0	278.0	5.4	59.6	35.0	Clayey silt
26-3,96.0	286.0	5.9	46.1	47.9	Silty clay

APPENDIX B Carbon-Carbonate Determinations, Holes 207 and 207A

Core, Section,				
Top of	Depth	Carbon	Organic	
Interval	in Hole	Total	Carbon	CaCO ₂
(cm)	(m)	(%)	(%)	(%)
Hole 207				
1-2,80.0	2.3	11.5	0.0	95
2-3,123.0	10.2	10.2	0.0	84
3-4,50.0	20.0	11.1	0.1	92
4-2,40.0	30.9	11.3	0.1	93
5-5,40.0	44.4	11.2	0.1	93
Hole 207A				
1-4.30.0	51.8	11.4	0.1	94
3-3,60.0	68.6	11.5	0.1	95
4-4,70.0	79.2	11.6	0.1	96
5-2,70.0	85.2	11.6	0.1	96
6-5,40.0	98.4	11.6	0.0	96
7-6,80.0	109.3	10.7	0.0	89
8-6,145.0	118.9	11.4	0.0	94
9-4,90.0	124.4	11.3	0.0	94
10-3,70.0	136.7	11.5	0.0	95
11-4,65.0	147.1	11.3	0.0	94
13-6,70.0	168.2	10.9	0.0	90
14-2,40.0	170.9	10.9	0.0	90
15-2,70.0	180,2	11.1	0.0	92
16-2,75.0	189.3	11.2	0.0	93
17-2,80.0	198.3	11.0	0.0	92
18-3,43.0	208.4	10.9	0.0	91
20-3,60.0	226.6	10.8	0.0	89
21-3,70.0	240.7	11.0	0.0	92
25-4,54.0	278.0	10.6	0.0	88
26-3,96.0	286.0	9.2	0.1	76
27-2,18.0	292.7	10.5	0.1	87
28-1,135.0	301.4	5.4	0.5	41
31-2,30.0	331.8	1.2	1.2	0
33-2,2.0	349.5	1.4	1.4	0

Core	Cored Interval Below Sea Floor (m)	Sample Depth Below Sea Floor (m)	Diff.	Amor.	Calc.	Quar.	Plag.	Kaol.	Mica.	Chlo.	Mont.	Cris.	K-Fe.	Trid.	Clin.	Pyri.
Hole 2	07 Bulk Sample	8														
3	15-24	19.5	53.5	27.4	98.9	1.1	_	_	_	_		_	_	_	-	-
Hole 2	07 <2µ Fraction	1														
3	15-24	19.5	89.5	83.5	_	24.7	11.7	9.8	37.1	6.0	10.7	-	_	_	_	-
Hole 2	07A Bulk Sampl	les														
33	348-357	349.5 351.1	86.4 87.2	78.8 80.1	_	8.4 5.9	3.8 2.3	_	1.6 0.9	_	3.7 3.0	69.2 69.2	1.5 1.4	8.5 14.4	1.4 0.9	1.8 1.9
42	432-441	434.5	73.0	57.9	-	0.6	4.6	_	-	· _	89.3	5.5	_	-	_	-
46	468-477	471.7	82.9	73.2	-	8.1	20.3	-	-	—	32.9	38.7	-	_	-	-
Hole 2	07A 2-20µ Frac	tions														
33	348-357	349.5 351.1	81.1 81.2	70.5 70.6	_	14.9 21.6	6.8 10.6	_	3.0 2.1	_	_	60.5 42.6	2.8 5.1	7.4 6.5	2.3 6.0	2.3 5.5
42	432-441	434.5	80.2	69.1	-	1.2	21.8	-	-	-	51.3	23.9	-	1.9	-	-
46	468-477	471.7	88.2	81.6	_	3.1	10.3	-	-	_	46.4	38.3	-	1.3	0.6	-
Hole 2	07A <2µ Fracti	on														
33	348-357	349.5 351.1	87.9 88.9	81.1 82.6	_	0.7 0.9	0.7 0.7	_	_	_	10.0 5.0	79.0 79.1	 1.0	7.6 11.3	0.5	1.9 1.5
42	432-441	434.5	74.8	60.7	-	-	-	—	-	-	93.6	6.4	-	-	-	-
46	468-477	471.7	69.3	52.1	_	-	-	-	-	_	95.8	4.2	-	-	-	-

APPENDIX C X-ray Mineralogy Determination, Site 207

Core, Section, Interval Below Top (cm)	Thermal Conductivity (mcal/°C cm sec)	Standard Deviation	Ambient Core Tempera- ture (°C)	Remarks
2-4,100	0.002772	0.009022	18.92	
3-3,85	0.002881	0.009589	19.11	
4-3,45	0.002960	0.006095	20.82	
5-4,42	0.002959	0.006806	20.71	
1A-5,115	0.002946	0.006011	20.01	
2A-1,115	0.002946	0.009757	19.70	
4A-3,84	0.003002	0.005387	18.07	
5A-3,84	0.003093	0.007216	19.70	
6A-3,77	0.003242	0.006424	19.46	
7A-3,91	0.002982	0.006394	20.57	
8A-3,81	0.003076	0.010349	20.30	Stiff sed
9A-3,75	0.002983	0.005508	18.58	
10A-3,85	0.003438	0.006814	21.45	
11A-4,84	0.003159	0.005523	21.29	
13A-5,90	0.003149	0.008342	18.92	
14A-2,79	0.002930	0.004147	18.54	at 84 cm
15A-3,80	0.003221	0.006062	17.86	
16A-3,56	0.002959	0.010481	20.69	
17A-3,56	0.003209	0.009900	20.30	
24A-3,58	0.003646	0.007383	21.98	
25A-3,58 26A-3,58	0.003561 0.003605	$0.008681 \\ 0.008674$	22.07 22.54	

APPENDIX D Thermal Conductivity Measurements, Holes 207 and 207A

Si	te 207	н	ole		Co	re 1	Cored 1	nter	val:	0-6 m	Site	207	Ho1	le	C	ore 2	Cored Interva	al:	6-15 m
ACE	ZONF	CO COSCELL	FOSS HARA	SIL CTER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FOSSIL 2	ARAC ONNgy	PRES. 31	METERS	LITHOLOGY DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION
	LCNN	N	I A	G	1	0.5-				5Y8/1 <u>NANNO FORAM OOZE</u> 5GY6/1 Yellowish gray to greenish patches gray; creamy. of: N8 Smear Slide at Section 1, 75 cm:			Ņ	A	^M 1	0.5			N8 FORAM NANNO OOZE patches Very light gray with greenish of: gray to very light gray SG16/1+ patches; creamy.
TETOCENE		1	A	G	2			4 4 1	сс	Soft TorAms Soft TorAms Alta nannos N8 Soft TorAms Alta nannos Soft TorAms So		,	N	A	M 2				NB Smear Slide at Section 1, 110 cm: 53% nannos 45% forams 1% glass 5GY6/1 FORAM RICH NANNO 002E patches Greenish gray with of: N8 patches of very light gray; creamy.
MID TO I ATE BIE			A	G	3					Light Dive gray to 5Y6/1 white; creamy. Smear Slide at Section 3, 70 cm: 59% mannos 40% forams 5Y6/1- →5GY6/1	PLEISTOCENE	6 L NN	N	A	м 3			GZ	Smear Silde at Section 2, 70 cm: 78% nannos 20% forams 1% glass 1% mica 000 1% mica 1%
	61NN	M	A	G	4			4	GZ	589/1.↓ →5GY6/1 [37% sand, 21% silt, 43% clay]			N	A	м 4				grading into: 3% glass 5G(7) [13% sand, 21% silt, 67% clay] [84% CaCD ₃] <u>FORAM NANNO OOZE</u> Very light gray with greenish gray streaks; soupy. N8 Streaks of: 59% nannos 56(76/1 40% forams 1% glass
		FN	A A +	G G M	C	ore tcher		-		FORAM RICH NANNO OOZE 75% nannos 25% forams			N	A	м 5				NANNO FORAM OOZE Very light gray; soupy to creamy. Smear Slide at Section 5, 81 cm: 50% foram 48% nannos 2% glass

Core Catcher

F A N A FORAM RICH NANNO OOZE

78% nannos 20% forams 1% glass 1% zeolites

207	Hole		Co	ore 3	Cored In	nterva	al: 1	5-24 m		Site	207	Hole		Co	ore 4	Cored In	terva	a]:	29-38 m
ZONE	FO CHAR	ABUND.	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOG	SIC DESCRIPTION	AGE	ZONE	FOSSIL R	ABUND.	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO' SAMPLE	LITHOLOGIC DESCRIPTION
	N	A M	1	0.5		4		5B9/1 FORAM RICH P Bluish white streaks of: N5 FORAM NANNO Bluish white Soupy to cre	ANNO OOZE with some medium gray py to creamy. Smear Slide at Section 1, 80 cm 78% nannos 20% forams 1% glass 1% zeolites 00ZE with medium gray patches. amy. Smear Slide at Section 2, 75 cm 54% nannos 45% forams 1% glass	PLEISTOCENE	6LNN 8LNN	N	AM	1 1 2	1.0	┥┥┥┥╎┝┝┝┝┝┝╎┥┥┥┥┥┥┥┥	4	GZ CC	NANNO FORAM 002E 509/1 to Bluish white to greenish gray. Creamy. with streams patches of greenish gray. Creamy. with Smear Slide at Section 1, 75 cm: patches of 54% forams pure: 64% forams
	N	A M	3	111111111111		4		NANNO FORAM N8 with Very light (streaks yellow gray; of: 5Y8/1	002E ray with streaks of creamy. Smear Slide at Section 3, 70 cm. 55% forams 43% nannos 1% quartz 1% feldspar	IOCENE	_	N	A	3			4		NB to N9 with patches and streaks of N5 and 5Y6/1
6 L NN	N	AM	4			4	хм gz cc	5Y6/1 [X-ray - 27% 5B9/1 99% with [35% sand, 2 patches of: 5GY6/1 5GY6/1 5GY6/1 Bluish white creamy. 55B9/1	amorphous, 73% crystalline: calc., 1% qtz. 4% silt, 41% clay] [92% CaCO ₃] <u>002E</u> with greenish gray patches; Smear Slide at Section 4, 50 cm. 59% nannos	LATE PL	2 LNN	N	AM	4			4		
	N	AM		-				5GY6/1	40% forams 1% zeolites			N	A	м					

5

Catcher

FNR

9LNN

A A G Core <u>NANNO FORAM OOZE</u> Greenish gray and very light gray (with medium gray streaks). Creamy.

FORAM NANNO OOZE Creamy texture.

FORAM NANNO OOZE

Smear Slide at Section 5, 75 cm: 55% forams 43% nannos 1% glass 1% quartz

Smear Slide at Section 6, 40 cm: 54% nannos 45% forams 1% glass

30% forams 1% feldspar

68% nannos 1% glass

N8 with streaks of: 5GY6/1 and N6

216

Site 207

GE

EARLY PLEISTOCENE

5

6

Core

Catcher

G M A

T----

Ν

F

NR

Α Iм 69% nannos 30% forams 1% glass

FORAM NANNO OOZE

Sit	e 207	Ho	ole		Cor	re 5	Co	ored In	terv	al: 3	8-47 m				Si	te 207	Ho	ole	A	Cor	re 1	Cored Int	erval	47-56 m
AGE	ZONE	FOSSIL 2	FOSSI HARAC	DRES.	SECTION	METERS	LITH	IOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION			AGE	ZONE	Enceti O	FOSS HARAI	BRES.	SECTION	METERS	LITHOLOGY	DEFORMATION LITHO.SAMPLE	LITHOLOGIC DESCRIPTION
		N	A	M	1	1.0	┍┍┍┍┍┍┍┍┍┍┍┍┍┍		4		N8	EORAM NANNO 00ZE Very light gray. Creamy, homog Smear Slide 50% nar 46% for 2% gla 1% fel 1% pyr	teneous. at Section 4, inos ams iss dspar oxene?	75 cm:		NN14-NN15	N	I A	м	1	0.5 1.0		4.	N8 FORAM NANNO OOZE with streaks Very light gray with streaks and patches of: of other colors. Creamy texture. N5, 5Y8/1, N4 Smear Slide at Section 1, 100 cm: N3 patches and 50% nannos Streaks 2% glass
IOCENE	9	N	А	м	3										TOCENE	413	N	A	м	3			4	NANNO FORAM OOZE Creamy texture. Smear Slide at Section 3, 75 cm: 54% forams 45% nannos 1% glass
MIDDLE PL	LNN	N	A	M	4			+ + + + + + + + + + + + + + + + + + +	4		5B9/1 to N8	FORAM NANNO OOZE Bluish white to very light gra	у.		MIDDIF DI	Nu	N	A	м	4			4 GZ	[11% sand, 28% silt, 60% clay] [94% CaCO ₃] F <u>ORAM NANNO 00ZE</u> Creamy texture. Smear Slide at Section 5, 75 cm: 55% nannos 45% forams
		N	А	м	5				4	GZ CC	N8 with swirls of: 5GY6/1 and N5	[24% sand, 27% silt, 49% clay] [93% CaCO ₃] <u>NANNO FORAM OOZE</u> Very light gray with swirls of greenish gray and medium gr	ay.			-	N	A	м	5			4	
	2	N	A	M	6				4		5B9/1 with swirls of: N5	NANNO FORAM OCZE Bluish gray with rare swirls o medium gray. Creamy texture.	f			≤ I NN-I I NN≈	N	A	м	6			4	
	NN15	F N R	A A -	G M -	Co Cat	ore cher						FORAM RICH NANNO OOZE	83% nannos 15% forams 1% glass				F N R	A A -	G M -	Co Cate	ore cher			FORAM RICH NANNO OOZE 73% nannos 25% forams 1% glass 1% zeolites

	Site	e 207	Hole	A	C	Core	3 Cored Int	terval	: 65-74 m
LITHOLOGIC DESCRIPTION	AGE	ZONE	FOSSIL FOSSIL	RACTE	PRES. 2		LITHOLOGY	DEFORMATION	LITHOLOGIC DESCRIPTION
h <u>NANNO FORAM OOZE</u> White, with swirls of yellowish gray and rare dark gray streaks. Creamy texture. Smear Slide at Section 1, 100 cm: 5% forams 45% nannos 1% zeolites <u>FORAM RICH NANNO OOZE</u> 59% nannos 40% forams 1% glass			N	A	P 1	0.		4	N8 with streaks and irregular patches of: 5Y8/1, N4, N5 <u>NANNO FORAM 00ZE</u> Very light gray with streaks of yellowish gray and medium dark gray and medium gray. Creamy texture. Smear Slide at Section 2, 80 cm: 60% forams 39% nannos 1% plass
	EARLY PLIOCENE (OPOITIAN STAGE)	ZTNN-TTNN≈	N	A	3 P	5	┍┙┥┙┙┙┙┙┙┙┙┙┙┙┙┥┥┥┥	4	IX 97833 Z [18% sand, 25% silt, 58% clay] C [95% CaCO ₃]
						5	┿┥┾┾┿┿┿┿┝┝┝┝┝┝┝┝┝┝┝ ╞┝┝╞╞┾╞╞╞╞╞╞┝┝┝┝┝┝┝┝┝ ┨┨┨┥┨┨┨┨┨┨┨┨	4	
			FNR	A A -	G P	Con			NANNO FORAM OOZE 55% forams 45% nannos

Cored Interval: 56-65 m Site 207 Hole A Core 2 FOSSIL CHARACTER DEFORMATION LITHO.SAMPLE FOSSIL ABUND. PRES. SECTION METERS AGE ZONE LITHOLOGY MIDDLE PLIOCENE 1111 EMPTY ≈NN11-NN12 N9 wit swirls of: 5Y8/1 some streak of: N3 NAP 1 1 0 A G A P - -F N R Core Catcher

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Site 207	Hol	e A		C	ore	4		Cored	Inte	erva	1:	74-83 m	Sit	207	ł	Hole	A	C	ore 5	Cored In	terv	al:8	12-92 m
AGE ZONE	FOSSIL 2	ARAC	BRES.	SECTION		METERS	LI	THOLOG	Y	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE		FOSSIL PAR	ABUND.	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION
LATE MIDGENE (KAPITEAN STAGE) EARLY PLIDGENE (OPOITIAN STAGE) =MN11-NN12	N F N F	A A A A A A A	P G G P P	1 2 3 4 5	0. 1.		┙┥┥┓┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥	┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥┥		4	GZ CC	N8 with irregular streaks and patches of: 5Y8/1, N5 and N3 FORAM NANNO ODZE Very light gray. Streaks and patches of yellowish gray, medium gray and dark gray. Soupy to creamy texture. Smear Slide at Section 2, 80 cm: 55% nannos 13 forams 13 forams 13 forams 13 forams 14 fish debris [42% sand, 22% silt, 36% clay] [96% CaCO ₃] NANNO FORAM OOZE S4% forams 45% nannos 18 glass	(TONGAPORUTUAN STAGE) LATE MIDCENE (KAPITEAN STAGE)	sinni-linnis		F N N	A P A P		0.5-			GZ CC	N8 with irregular streaks and patches of; 5Y8/1 and N5 FORAM MANNO OOZE Very light gray with streaks and patches of yellowish gray and medium gray. Creamy texture. [24% sand, 29% silt, 47% clay] [96% caCO_3] Smear Slide at Section 3, 80 cm: 60% nannos 40% forams
																FNR	A G A P	Ca	Core				FORAM NANNO QOZE 55% nannos 45% forams



Explanatory notes in Chapter 1

Sit	e 207	Но	le A	0	Core	8	Co	red In	terv	al:	110-119 m				Site	e 207	Ho	ole	A	Cor	re 9	Cored Ir	iterv	al:	119-128 m			
AGE	ZONE	FOSSIL 2_	ARACTI ONNBY	PRES. 2	2001 1014	METERS	LITH	OLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLO	GIC DESCRIPTION		AGE	ZONE	Enseti O	HARA	SIL CTER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DE	SCRIPTION	
	6NN	N	A	M 1	0	.5.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.			4		N9	FORAM NANNO White. Soup	00ZE y texture.		AN STAGE)	≈NN6	F		G	1	1.0		4		2.5Y8/2	<u>FORAM NANNO OOZE</u> White (yellow tir	it). Creamy texture. Smear Slide at Section 1, 75 cm 60% nannos 40% forams	1:
N STAGE)				2	2				4				Smear Slide 60% nar 40% for	at Section 4, 75 cm: nos ams	INING (LILLBURN		F	Γ Α	G	2			4					
EARLY TONGAPORUTUA				3					4						MIOCENE		F	A	G	3	111111111111		4					
MIOCENE (WAIAUAN OR	enn~8nn∞	N	A	4		*****		++++++++++++++++++++++++++++++++++++	4						N STAGE) MIDDLE	NN₽	F	ι A : Α	G	4	1111111111111		4	GZ		[30% sand, 43% si <u>NANNO FORAM OOZE</u> White to pale yel [94% CaCO ₃]	lt, 27% clay] low. Greamy texture. Smear Slide at Section 4, 75 cm: 55% forams 45% nannos	:
LATE				5					4						(CLIFDENIA		N F	I A	G	5			4		gradation 2.5Y8/2 ↓ to	nal change from:		
					+	1,) MIOCENE	uensis	FN	A	GM		11111111				N9 with streaks 2.5Y8/4	(pale yellow) of:		
Evo		FNR	A A -	6 1 Ci	Corre	e ner			4	CC GZ		[94% CaCO ₃] [27% sand, FORAM NANNO	42% silt, 31% clay] <u>007E</u>	60% nannos 40% forams	MIXED LATE EOCENE AND	Chiasmolithus oamar	FNR	A A	G M 	6 Cat	ore ccher		4			FORAM NANNO OOZE	70% nannos 30% forams	

S	ite a	07	Ho1	e A		Cor	e 10	Core	d In	terval:	:133-142 m	Site	207	Hole	A	Co	re 11	Cored In	terva	1:1	: 142-151 m
	AGE	ZONE	F0SSIL ₹ -	OSSI RACT	PRES. B	SECTION	METERS	LITHOLO	ιGΥ	DEFORMATION LITHO.SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FOSSIL B. A	RACTER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION
	EARLY MIDDLE MICCENE (LILLEDURNIAN STAGE)	#NN5	F N N F N R	A A A -	G M	1 2 3 4	0.5			4 4 4 4 4	2.5Y8/4 (pale yellow) with streaks of: N9 (white) Pale yellow with streaks of white. Creamy texture. N9 2.5Y8/4 with streaks of: N9 ECRAM NANNO 00ZE As above. [33% sand, 37% silt, 30% clay] [95% CaCO ₃] Smear Slide at Section 4, 50 cm: 55% nannos 45% forams ZEOLITE BEARING NANNO FORAM 00ZE 1% glass	MIDDLE EOCENE (BORTOMIAN STAGE)	*Reticuiofenestra hampdenensis + Discoaster distinctus	N	A M A C	1 2 3 4 5	0.5	┙┥┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙┙	4	GZ CC	N9 FORAM NANNO ODZE White. Creamy texture. Contains broken (drilling) chert fragments (light gray - 2.5V7/2). Smear Slide at Section 4, 75 cm: 60% nannos 40% forams 2 [18% sand, 45% silt, 37% clay] [94% CaCO ₃]

F N R

A G Core A M Catche SITE 207

59% nannos 40% forams 1% zeolites

FORAM NANNO OOZE

	Core 12	Cored I	nterv	al:	151-160 m							Si	ite 2	07	Hole	Α	Cor	e 13	Cored In	terval:	160-169 m	
PRES. N	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE		LII	HOLOGIC	DESCRIPTI	ON	e.		405	AGE	ZONE	FOS CHARA	SIL ACTER	SECTION	METERS	LITHOLOGY	DEFORMATION LITHO.SAMPLE		LITHOLOGIC DESCRIPTION
	l l l.o Core Catcher				N9	FORAM N White.	ANNO OOZI Creamy to SPICULE I H NANNO (Stiff. Smear 5 4 <u>BEARING F</u>	Slide at 0% nannos 6% forams 2% rads 2% sponge ORAM AND	Section 1, 130 spicules 56% nannos 20% forams 20% forams 20% rads 3% sponge spicul 1% zeolite	0 cm: les es	ατησις εστές (δούτουτω, ετώς)	WINDLE EUCENE (DOWN UNITWE STARE)	Chiphragmalithus cristatus (upper)			1 2 3 4 5 6		╵┝╴┡└┝┟┝┟┝┠╹┠╹┍╹┍╹┍╹┍╹┝╹┝╹┝╘┝╘┝╘┝╘┝╘┝╘┝╘┝╘┝╘┝╘┝╘┝╘┝╘┝	4 4 4 4 6 7 C C	N9 with very rar of: N4	FORAM AND RAD RICH NANNO 002E White. Creamy texture. Smear Slide at Section 2, 75 cm: 89% nannos 5% forams 3% rads 2% sponge spicules 1% plant debris FORAM AND RAD RICH NANNO 002E White. Creamy texture. Smear Slide at Section 5, 75 cm: 69% nanos 15% forams 15% forams 15% rods 1% feldspar
															F A N A R C	G G	Co Cat	ore cher	1			SPONGE SPICULE BEARING 43% nannos 30% rads FORAM RICH RAD NANNO 20% forams 5% sponge QUZE 1% glass spicules 1% zolites 1% zolites 1% zolites

Hole A FOSSIL CHARACTER ZONE AGE FOSSIL ABUND. MIDDLE EOCENE (BORTONIAN STAGE) Chiphragmalithus cristatus (upper) F A N A R C

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Site 207

SITE 207



N R

Catcher G

SITE 207

5% forams

3% sponge spicules 2% rads 1% zeolites



	PRES. 3	PRES. 18 SECTION METERS	Core 18 Cored In R Core 18 Cored In LITHOLOGY W LITHOLOGY	Core 18 Cored Interva	Core 18 Cored Interval: 2 R U01133 LITHOLOGY W01133 LITHOLOGY W0133 LITHOLOGY W0133 LI	Core 18 Cored Interval: 205-214 m	Core 18 Cored Interval: 205-214 m Site	Core 18 Cored Interval: 205-214 m Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207 Image: Site 207	Core 18 Cored Interval: 205-214 m Site 207 Hole	Core 18 Cored Interval: 205-214 m Site 207 Hole A	Core 18 Cored Interval: 205-214 m Site 207 Hole A Co R UILING V RULE V	Core 18 Cored Interval: 205-214 m Site 207 Hole A Core 19 Image: Rel to the state s	Core 18 Cored Interval: 205-214 m Site 207 Hole A Core 19 Cored Interval: 205-214 m Image: Rel of the state stat	Core 18 Cored Interval: 205-214 m Site 207 Hole A Core 19 Cored Interval: 205-214 m IR VI VI VI VI VI VI VI VI IR VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI VI
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	1	1		0.5-4-4-4-4 1 4-4-4-4 1.0-4-4-4 1.0-4-4-4 -4-4-4-4	1	0.5 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	0.5 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	0.5-1 1 1 1	0.5 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	0.5-1.4-1.4-1.4 RAD_SPONGE SPICULE AND 1 1.4-1.4-1.4 1.0-1.4-1.4-1.4 FORAM BEARING NANNO 002E 1.0-1.4-1.4-1.4 Light bluish white. Stiff 1.1.4-1.4-1.4 Smear Slide at Section 1, 75 cm: 84% nannos 84% nannos 94% nannos 10% forams	0.5-1.4-1.4-1.4 RAD_SPONGE SPICULE AND 1 1.4-1.4-1.4 1.0-1.4-1.4-1.4 FORAM BEARING NAME OOZE 1.0-1.4-1.4-1.4 Light bluish white. Stiff 1.0-1.4-1.4-1.4 Smear Slide at Section 1, 75 cm: 1.1.4-1.4-1.4 SMEAR Slide at Section 1, 75 cm: 1.1.4-1.4-1.4 Smear Slide at Section 1, 75 cm:	0.5-1.4-1.4.1.1.2 RAD. SPONGE SPICULE AND 0.5-1.1.2.1.2.1.2.1.2.1.2.1.2.1.2.1.2.1.2.1	0.5-1 1 0.5-1 0.5-1 0.5-1 1 1 1 1 1 1 1 1 </td <td>0.5-1 1 0.5-1 0.5-1 0.5-1 1 1 1 1 0.5-1 1 1 1 1 1 1 <td< td=""></td<></td>	0.5-1 1 0.5-1 0.5-1 0.5-1 1 1 1 1 0.5-1 1 1 1 1 1 1 <td< td=""></td<>
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M						Smear Slide at Section 5, 75 cm: 74 74 74 74 74 74 74 74 74 74 74 74 74 74 74 75 74	L-3	Smear Slide at Section 5, 75 cm: 74% nannos 20% forams 20% forams 20% forams 20% forams 20% forams 20% forams 3% sponge spicules 3% rads	L-3	Smear Slide at Section 5, 75 cm: 74% nannos L	Smear Slide at Section 5, 75 cm: 74% nannos 20% forans	Smear Slide at Section 5, 75 cm: 74% nanos L 74% nanos 20% forams 20% forams L 3% sponge spicules 3% rads 3% rads	L-3	L-3
M G G Ca	Core					SPONSE SPICULE BEARING 51% nannos 25% forams ST S RAD AND FORAM RICH 3% sponge 20% rads AND FORAM RICH 3% sponge 20% rads	Image: Spin start Spin start Spin start Spin start Spin start Spin start Spin start Spin start			Image: Spin start Spin start	Image: Spice	Image: spin start Image: spin st	$\begin{array}{c c c c c c c c c c c c c c c c c c c $	$\begin{array}{c} - 1 \\$

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Site 207	7	Hole	Α	С	ore 20	Cored	Inte	erval:	223-232 m			Site	207	Ho1	e A	Co	ore 21	Cored Ir	terval	237-246 m	
AGE	ZUNE	FO SSIL PAR	ABUND.	SECTION	METERS	LITHOLOG	Y	DEFORMATION LITHO.SAMPLE		LITHOLOGIC DESCRIPTION		AGE	ZONE	FOSSIL P	NOR CONTRACTE	SECTION	METERS	LITHOLOGY	DEFORMATION LITHO.SAMPLE		LITHOLOGIC DESCRIPTION
atus (lower)		N	A	1	0.5-			1	589/1				Discoaster elegans	N	A	1 1	0.5		1	N9	
Chioraomalithus crist				2	-			1		FORAM BEARING NANNO 002E Bluish white. Stiff to semilithified. Smear Slide at Section 2, 75 93% nannos 5% forams 1% rads 1% sponge spicules	5 cm:			N	AM	2			1		RAD, SPONCE SPICULE AND FORAM BEARING HANNO 002E Bluish white and white. Semilithified. Smear Slide at Section 2, 75 cm: 91% nannos 5% Fororams 2% rads 2% sponge spicules
E (HERETAUNGAN STAGE)	-	N	A M	3			-	1 GZ CC		[4% sand, 51% silt, 44% clay] [89% CaCO ₃]		HERETAUNGAN STAGE)	(upper)			3			GZ 1 CC	589/1	[3% sand, 47% silt, 51% clay] [92% CaCO ₃]
MIDDLE EOCEN	0.000			4	-		-	1				MIDDLE EOCENE (culofenestra dictyoda	N	A	4			1		
Discoaster				5						faint stratification at right angles to core length RAD, SPONGE SPICULE AND FORAM BEARING NANNO 0022E Bluish white. Semilithified. Smear Slide at Section 5, 75 91% nannos 5% forams 2% rads 2% rads	; cm:		Reti			5			٦	N9	
		F	A	6						RAD, CLAY(?), SPONGE 80% nannos 5% 1	forams			F	A	6			1		FORAM RICH NANNO QOZE 85% nannos 10% forams
		N R	AN	Ca	tcher					NANNO OOZE spicules 5% of	clay(?)			N R	A M R P	Ca	tcher				(clay bearing?) 5% clay?

Core	22	Cored In	terva	al: 2	246-255 m			S	ite	207	Hole	А	C	ore 23	Cored In	terval	255-264 m	
PRES. 31	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION			AGE	ZONE	FO CHAI	ABUND.	SECTION	METERS	LITHOLOGY	DEFORMATION		LITHOLOGIC DESCRIPTION
G Cor P Catcl	.5]		N9 117 → N8 589/1	FORAM RICH NANNO QOZE White to bluish white. Semilith Chert fragments. FORAM RICH NANNO OOZE	ified. 89% nannos 10% forams 1% glass						1	0.5		4	N9	<u>FORAM BEARING NANNO OOZE</u> White. Soupy to stiff.
									RETAUNGAN STAGE)	dictyoda (upper)			3			4		Smear S11de at Section 2, 75 cm: 89% nannos 10% forams 1% glass
									MIDDLE EOCENE (HEI	Reticulofenestra			4			1+ + 4		← Chert fragment (light olive gray, 5Y6/1)
													5			4 J2 4	N9	FORAM BEARING NANNO 00ZE White and bluish white. Soupy to semilithified. Chert horizons in Section 6. Chert replaces nanno ooze (irregular boundaries and semi-silicified ooze) Smear Slide at Section 5, 75 cm: 97% nannos 3% forams
											FN	A G A F	Ca	Core			589/1	<u>FORAM RICH NANNO 00ZE</u> 79% nannos 20% forams 1% zeolites

MIDDLE EDCENE (HERETAUNGAN STAGE) Reticulofenestra dictyoda (upper) F A N A R R

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Site 207	Н	ole	A	C	ore	24	Cored Ir	nterva	al: 2	54-273		Site	207	H	ole	А	Co	re 25	Cored	Inte	rval:	273-282 m
AGE ZONE		FOS CHARA TISSO4	ACTE	SECTION		MELEKS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION		AGE	ZONE		FOS CHAR	ACTER	SECTION	METERS	LITHOLOGY	DECODMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION
STAGE) upper)		N	A	1	0.	51111111111111				N9 FORAM BEARING NANNO OOZE White. Soupy, with medium to light gray angular (drilling) fragments of chert and porcellar (?, partly altered nanno ooze). Ghosts of microfossils can be se in the chert fragments. Smear Slide al	nite 2en t Section 3, 75 cm:		Reticulofenestra	<pre>dictyda (lower)</pre>		A M	1	0.5			-3	589/1 → 56Y8/1
EOCENE (HERETAUNGAN ofenestra dictyoda (2	2	hititititi				995 nann 5% foran	75 115	(MANGAORAPAN STAGE)		N		A P	2				-3	FORAM RICH NANNO OOZE Bluish white and light greenish gray. Semilithified to creamy. Smear Slide at Section 2, 54 cm: 90% nannos 10% forams
MIDDLE				3	5	hititit						EARLY EOCENE	Discosster lodons		N	A P	3					In semilithified part of Section 4 moderate
STAGE)		N	A	1	ł			1/2		589/1				N	N .	A M	4				GZ CC	mottling is visible. Burrows up to 15 mm, some subparallel, others vertical to bedding. ← Chert layer: 56Y6/1 Chert consists of =90% opal and 10% chalcedony. [5% sand, 60% silt, 35% clay] [88% CaCO ₃]
EARLY EOCENE (MANGAORAPAN : Retizulofenestra dictvoda		N	A	M	Core			+ + + + + + 2	- <u>1</u> 4 -1	N9 (Nanno ooze) As above. FORAM RICH NANNO OOZE	90% nannos 10% forams		Chiasmo-	grandis	F	A G M	Car	ore tcher				FORAM RICH NANNO OOZE 88% nannos 10% forams 1% glass 1% zeolites

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Site 207	Hole A Core 29 Cored Interval:	309-318 m	Site 207 Hole A Core 31 Cored Interval: 330-339 m
AGE ZONE	FOSSIL CHARACTER CHARACTER NOUTLYNO SECLION NOTHIN SECLION SECLION NOTHIN SECLION	LITHOLOGIC DESCRIPTION	Bit Markov FOSSIL CHARACTER I SUB I
	N 1 0.5 $\frac{PPPY}{2}$ 2 $\frac{P}{2}$	5Y4/1 <u>SLIGHTLY CALCAREOUS SILTY CLAYSTONE</u> Olive gray. 5Y2/1 (to 5Y4/1 + dry) <u>SILTY CLAYSTONE</u> Olive black (wet); non-calcareous. Some faint mottling at base of Section 2. <u>Smear Slide at Section 1, 30 cm:</u> grain size distribution: 30% silt 70% clay some recognizable minerals: 6% opaques (pyrite?) 1% iron oxide grains 5% feldspar(?) 10% quartz(?) 2% chlorite 2% glass 2% glauconite Some sporadic sponge spicules and arenaceous forams (in Section 309 to 357 meters). Silt and clay-size fraction in this unit (309-357 meters) consists mainly of - cristobalite. Montmorillonite is a minor constituent. <u>PYRITE AND ZEOLITE BEARING CLAYSTONE (noncalcareous). Grain size: 10% silt 90% clay minerals: 5% pyrite 5% zeolites(?)</u>	1 0.5 EMPTY 1 1.0 5Y2/1 (wet) 2 5Y4/1 (dry) [0% CaCOa] 2 5Y4/1 (dry) [0% CaCOa] 3 SILTY CLAYSTONE Olive black (wet). Moderately mottled (burrows sub-parallel to bedding). White spherical sand size grains (<<1%): Sponge Spicules. 3 Smear Slide at Section 2, 100 cm: grain size: 70% Clay 30% silt 3 Silty CLAYSTONE 1% for oxide 1% plant fossils(?) 5 Silty CLAYSTONE As above. Glauconite grains scattered throughout (<1%), From 112 to 150 cm parts with up to 5% glauconite. Moderately mottled.
Site 207	Hole A Core 30 Cored Interval:	318-327 m	Core SILTY CLAYSTONE
AGE ZONE	ABUNDI FERS ABUND METERS	LITHOLOGIC DESCRIPTION	Catcher As Core 30.
	1 0.5 1 1.0 Core Catcher	5Y2/1 SILTY CLAYSTONE Olive black. Homogeneous, with sporadic microfossils. Smear Slide at Section 1, 75 cm: grain size: 70% clay 30% silt some constituents: 8% opaque (pyrite?) 10% feldspar 1% glauconite ZEOLITE BEARING CLAYSTONE grain size: 90% clay 10% silt some constituents: 2% zeolites 1% feldspar	

Hole	А	Co	ore 33	Cored Ir	terv	al:	348-357 m		Site	207	Hole	А	Cor	re 34	Cored Inter	rval	: 357-366 m	
FO: CHAR TISSOJ	ACTEF	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE		LITHOLOGIC DESCRIPTION	AGE	ZONE	FO CHAI CHAI TI SSOJ	ABUND.	SECTION	METERS	LITHOLOGY	I TTUO CAMPLE	LI I HU. SAMPLE	LITHOLOGIC DESCRIPTION
		1 2 3	0.5			CC XM	5¥2/1 5¥2/1	SILTY CLAYSTONE Olive black. As Core 3, Section 4. [0% CaC0_3] At Section 2, 10 cm of 3 mm thick seams of pyrite crystals. X-ray - 79% amorphous, 21% crystalline: 8% qtz., 69% cris., 1% K-feld., 4% plag., 2% mica, 4% mont., 9% tridymite, 1% clin., 2% pyrite SILTY CLAYSTONE Olive black. Moderately mottled. Glauconite grains scattered throughout (<<1%).					1 2 Cat	0.5 1.0			SY6/1 S6Y2/1 variou colors variou colors	<pre>to SILTY CLAYSTONE AND SANDSTONE Light olive gray to greenish black. Irregular stratification. Some layers of pyrite. Grains consist of quartz, feldspar, glauconite, etc. VITROPHYRIC RHYOLITE Dark greenish gray (50'4/1) to greenish black (5Y2/1) with light greenish gray (56'8/1) to greenish gray (50'6/1) fluidal banding. Some light brownish gray (5Y6/1) fluidal banding. Some light brownish gray (5Y6/1) fluidal banding. Some light sorwnish gray (5Y6/1) fluidal banding. Some stare lapilli size (maximum 1 cm). Matrix contains up to 5% quartz phenocrysts. Fluidal banding dips up to 60°. <u>FRAGMENTED VITROPHYRIC RHYOLITE</u> Colors vary: olive black (5Y2/1); greenish gray (56'6/1); brownish gray (5Y84/1); light greenish gray (56'8/1), brownish black (5Y2/1); dre. Frag- ments are lapilli size and make up 50% of the rock. Fragments and groundmass not always clearly distinguishable. Glassy groundmass contains up to 5% quartz phenocrysts. Perlitic texture and con- choidal fracturing are common. Fluidal banding visible at intervals 1 (93-150) and 2 (77-82). Dips vary from 60 to 90°. Fragments are subangular. <u>RHYOLITE LAPILLI TUFF</u> Colorless to brownish black (5YR2/1) perlitic fragments with quartz phenocrysts. Fragments are ash and lapilli size, and are subrounded to sub- angular. NOTE: Most color given in the core summaries of these volcaries (Come 24 to 50) are vet colors</pre>
N		Ca	tcher	IX/N				l% glass l% glauconite	L									
								X-ray - 80% amorphous, 20% crystalline: 6% gtz., 69% cris., 1% K-feld., 2% plag., 1% mica. 3% mont., 14% tridymite, 1% clin., 2% pyrite] <u>SILTY CLAYSTONE</u> Same as Core 32.	AGE	ZONE	FO CHAI LISSOJ	A SSIL SSIL ACTER	SECTION	METERS			11110.0484	LITHOLOGIC DESCRIPTION
													1 2 3	0.5	EMPTY	1	5Y2/1 5YR2/1 vario colors	FRAGMENTED VITROPHYRIC RHYOLITE (possibly (APILLI TUFF) Oflive black and brownish black fragments in a greenish glassy groundmass (with perlitic texture and conchoidal fracturing). Some tiny green veinlets. Some scattered pyrite. Same as Section 1. Colors: Olive black (5Y2/1); brownish black (SYR2/1) when wet. Greenish fragments (5SY6/1) when dry. Matrix darker than fragments (5SY6/1) to 5SY4/1). IN0-117 cm: nearly all grayish black perlitic glass, without fragments (PITCHSTONE). Refr. index of glass = 1.500 (very acid). FRAGMENTED VITROPHYRIC RHYOLITE (Similar to Section 2) Glassy fragments (10-20% of rock) in a glassy groundmass. Groundmass is borwnish gray, black, and greenish gray; shows fluidal banding, perlitic texture, and conchoidal fracturing. From 0-47 cm abundant (10%) sanidine and quartz phenocrysts. From 47-150 cm, as Section 2; fluidal banding dipping 40-60°; no visible feldspar phenocrysts.
													Cat	ore tcher				FRAGMENTED VITROPHYRIC RHYOLITE Fragments are brownish gray, greenish gray, and black. Perlitic glassy groundmass with quartz phenocrysts. Possibly a tuff.

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Site 207

AGE ZONE
Site 207	Hole A Core 36 Cored Interval:	375-384 m	Site 207 Hole A Core 40 Cored Interval: 411-420 m
AGE ZONF	F02S1IL CHARTES. SECTION METERS SECTION METERS ABOTOHLIT DEFORMATION LLTHO. SAMPLE	LITHOLOGIC DESCRIPTION	Yes Fossile CHARACTER IN ISOUTO Processile ISOUTO Processile ISOUTO Processile ISOUTO ISOUTO ISOUTO <thisouto< th=""> <thisouto< th=""> <thitholog< td=""></thitholog<></thisouto<></thisouto<>
	L Core Catcher	various FRAGMENTED VITROPHYRIC RHYOLITE colors No fluidal banding. Has several percent feldspar phenocrysts. <u>VITROPHYRIC RHYOLITE</u> Brownish gray (57R4/1). Glass has perlitic texture. Pyrite lines perlitic cracks.	I 0.5 EMPTY 1 0.5 Figure 1 1.0 1 * 1.0 1 * 5Y2/1 Figure 1.5 6 5Y2/1 7 0.1% gray to olive black. Predominantly pumice 6 Fragments, but 5-10% glassy (some black) fragments. SY4/1 Functe has cellular texture. Groundmass contains Core Core Catcher RHOLITIC LAPILLI TUFF Fragments (up to 4 cm) of black glass with fluidal banding. Some secondary hematite.
Site 207	FOSSIL	384-393 m	Site 207 Hole A Core 41 Cored Interval: 420-429 m
AGE ZONF	CHARACTER S BRES. FOSSIL ABUNO. WETERS SECTION METERS ABUNO. CAMPLE LITHO CAMPLE	LITHOLOGIC DESCRIPTION	POSSIL CHARACTER 11500 Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z
	Core Catcher	FRAGMENTED VITROPHYRIC RHYOLITE Fragments show fluidal banding.	Core Catcher Catcher
Site 207	Hole A Core 38 Cored Interval:	393-402 m	Site 207 Hole & Corred Laterweit 422 441 m
AGE ZONF	FOSSIL CHARACTER BEES: B	LITHOLOGIC DESCRIPTION	BIG LO FINE RECEIVED INCLASS AND
	Core Catcher	FRAGMENTED VITROPHYRIC RHYOLITE Black, brownish gray, and greenish gray fragments.	
Site 20	7 Hole A Core 39 Cored Interval:	402-411 m	
AGE	FOSSIL CHARACTER BEES BEELON HITHOLOGY HITHOLO	LITHOLOGIC DESCRIPTION	1.0 1 Image: Constraint of the constraint
	1 0.5 1.0 1.0 Core Catcher	RHYOLITE PUMICE (TUFF?) Dark yellowish brown. Contains quartz phenocrysts (bipyramidal) and pockets (up to 3 cm) filled with black glass. Some rare fragments of glassy rhyolite. 5GY4/1 PUMICEOUS LAFILLI TUFF Dark greenish gray. Fragments are mainly pumice (5-7 mm; up to 80% of rock) and some black glass and olive black laminated glass, with occasional quartz bipyramids and feldspar phenocrysts (~15%). No fluidal banding and little matrix. 115-120 cm: band of olive black fine-grained, non- pumiceous volcanics. RHYOLITIC LAPILLI TUFF Fragments of pumice and of vitrophyric rhyolite. Some secondary hematite. Fragments have sizes up to 8 mm.	2 iiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiii

SITE 207

Site	207	Hole	А		Cor	re 43	Cored In	terv	al:	441-450 m		Sit	e
Γ		FOS CHAR	SIL	R	N	\$		NOI	IPLE			Γ	1
AGE	ZONE	FOSSIL	ADUND.	PRES.	SECTIO	METER	LITHOLOGY	DEFORMAT	LITHO.SAN		LITHOLOGIC DESCRIPTION	AGE	
					1 Cat	0.5 1.0 core	EMPTY	1		56Y6/1	ARGILLIZED RHYOLITE Greenish gray. Faint fluidal banding, dipping 60°. Glassy groundmass altered to clay (probably montmorillonite and halloysite). ARGILLIZED RHYOLITE Greenish gray. Alteration product mainly halloysite, with some montmorillonite.		
Site	207	Hole	A		Cor	re 44	Cored In	terv	al:	450-459 m			
AGE	ZONE	FOS CHAR 11SSOJ	ACTE	PRES. 3	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION		
					1 Cat	0.5 1.0				5YR3/4	SPHERULITIC VITROPHYRIC RHYQLITE Grayish red (SR4/2). Quartz and feldspar phenocrysts in a Fe-oxide rich glass (devirified?). Abundant vesicles, filled with chalcedony. SPHERNLITIC VITROPHYRIC RHYQLITE Moderate brown. Vesicles filled with chalcedony. Fluidal banding, dipping 60°. SPHERNLITIC VITROPHYRIC RHYQLITE Dark reddish brown (10R3/4). Boundary at 62 cm is gradational. Same as between 11 and 62 cm. SPHERULITIC VITROPHYRIC RHYQLITE Light greenish grav (56Y8/1) and pale greenish yellow (10V3/2), and pale red (10K6/2). Some frag- mentation, especially in lower part. Fragments up to 1 cm. Note: May mark contact of discrete flows. SPHERULITIC VITROPHYRIC RHYQLITE Fluidal banding in various colors: dark greenish gray (56Y4/1), dark reddish brown (10R3/4), and light greenish grav (56Y8/1). Banding dips 60°. SPHERULTIC VITROPHYRIC RHYQLITE Fluidal banding in various colors: dark greenish gray (56Y4/1), dark reddish brown (10R3/4), and light greenish gray (56Y8/1). Banding dips 60°. SPHERULTIC WITROPHYRIC ARVOLITE		

e 207	7	Ho1	e A		Co	re 45	Cored In	terv	al:	459-468 m	
ZONE	ZUNE	F0SSIL ₽	RACT . ONUBA	PRES. BI	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION
					1	0.5	× / \ /	1	N	5G4/1 and 5YR4/3 various	SPHERULITIC VITROPHYRIC RHYOLITE Altered (argillized) pale green glass with wavy fluidal banding. A black unaltered glass lamina dips 45°. SPHERULITIC VITROPHYRIC RHYOLITE Dark greenish gray and brownish gray fluidal banding. Quartz and feldspar phenocrysts. Groundmass partly altered (zeolitized) and devitrified (chalcedony? and spherulites). Spherulites give rock a 'salt-and-pepper' appearance. Spherulites are in centers of perlitic fragments. FRAGHENTED VITROPHYRIC RHYOLITE Fragments of arcillized rhyolite (up to 3 cm).
					2 3 Ca	ore		1		various colors	Pragments of argillized rhyolite (up to 3 cm). Fluidal Danding still visible in fragments. Mosaic of colors, from light greenish gray (56Y8/1) to moderate reddish brown (1084/6). Fragmentation may be due to autobrecciation. Alteration products: montmorillonite and halloysite. <u>ARGILLIZED VITROPHYRIC RHYOLITE</u> (looks like top of core) Some parts are brecciated, containing fragments of altered rhyolite. Fluidal banding dips from 0° to 45°. <u>SPHERULITIC VITROPHYRIC RHYOLITE</u> Partly argillized. Colors vary, from dark greenish gray (56Y4/1) to light greenish gray (56Y8/1) to shades of grayish red (5R4/2). Phenocrysts of quartz and feldspar. Groundmass exists of altered perlitic glass with spherulites. Fluidal banding dips 60°.

Sit	e 207	Hole	A	Co	re 46	Cored	Inter	val:	468-477 m		S	ite	207	Hole	e A	Co	ore 47	Cored In	terval:	477-486 m	
AGE	ZONE	FOSSIL FOSSIL	BRES.	SECTION	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE		LITHOLOGIC DESCRIPTION		AGE	ZONE	FOSSIL PA	OSSIL RACTE	PRES. 2	METERS	LITHOLOGY	DEFORMATION LITHO.SAMPLE	Ка	LITHOLOGIC DESCRIPTION
				1	0.5	\/} /}// \/}			5GY4/1	SPHERULITIC VITROPHYRIC RHYOLITE Dark greenish gray. Altered parts and fragmented parts. The latter especially at intervals 0-60 and 80-110 cm. Fluidal banding is especially well developed in lower half of section (dipping 30°- 60°). Secondary Fe-oxide colors local areas red (10R4/6).						1	0.5	EMPTY	1		<u>VITROPHYRIC RHYOLITE</u> Black, vesicular. With parts of altered fragmented rhyolite.
				2						fragmentation. Veinlets of quartz, chalcedony, and pale green to white soft minerals. Trace amounts of pyrite and Fe-oxide. Most glass devitrified. Fluidal banding dips 10°-50°.						2			1	various colors	SPHERULITIC VITROPHYRIC RHYOLITE Vesicular. Altered and fragmented from 65 cm downwards. Groundmass of altered glass contains phenocrysts of quartz and feldspar (<5%). At 45 cm and at 135 cm some relatively unaltered glass that is vesicular (amygdaloidal). Vesicules filled with opal. Fluidal banding dips 25°-35°. Banding dis- rupted by Fe-oxide and chalcedony veinlets.
				3			*	ХМ	-	<pre>Feldspar phenocrysts. Black glassy bands at 93-95 and 120 cm. Fluidal banding dips 0°-15°. [X-ray - 73% amorphous, 27% crystalline: 8 % qtz., 39% cris., 20% plag. 33% mont.</pre>						3			1	various colors	ARGILLIZED SPHERULITIC VITROPHYRIC RHYOLITE Mottled layers of light greenish gray (56Y8/1), black (N2), and moderate brown (5YR3/4). Glassy groundmass is argillized. Quartz and feldspar phenocrysts. From 45 to 150 cm: light greenish gray (56Y8/1) and grayish red (5R4/2). At 75 cm a veinlet, containing quartz. Fluidal banding in whole section (dipping 15°
				4		/ \ / \				Same as above. Some fragmentation at 30-70 cm and 105-115 cm. Fluidal banding dips 10°-45°.						Ca	Core atcher				to 90°). Fragmentation from 3/140 to 3/150. <u>ARGILLIZED SPHERULITIC VITROPHYRIC RHYOLITE</u> Some veinlets with secondary hematite.
				Cc Cat	cher	$\sum_{i=1}^{N}$	/			FRAGMENTED VITROPHYRIC RHYOLITE Large fragment of black perlitic rhyolite (>7 cm).											

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AGE	ZONE	FOSSIL Z	ARAC . ONDAR	PRES.	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLI		LITHOLOGIC DESCRIPTION
					1	0.5	EMPTY			various colors	ARGILLIZED SPHERULITIC VITROPHYRIC RHYOLITE Mottled coloring: moderate yellow green (5677/4) and dark yellowish orange (10YR6/6). Groundmass altered. No apparent fragmentation. Quartz and Fieldspar phenocrysts (<2%). From 115 to 150 cm different coloring: dusky red (5R3/4) and pale olive (10Y6/2). At base of section some fluidal
					2					various colors	banding (dipping 50°). SIMILAR TO SECTION 1. Groundmass is altered to montmorillonite. Colors are: dusky red (SR3/4), moderate red (SR5/4), lig greenish gray (SB8/1), and greenish gray (differe shades). Remnants of perlitic glass gives rock a 'salt-and-pepper' texture. Fluidal banding dips 60° to 70°.
					3			1		various colors	SIMILAR TO SECTION 2. Colors are: 0-40 cm: bate greenish gray (564/1) 40-90 cm: pale yellowish green (10GY7/1) 90-150 cm: dark greenish gray. Some veinlets of reddish Fe-oxide. Fluidal banding dips 60°-90°.
					4					various colors	A <u>S SECTION 3</u> . Colors are: 0-55; 75-100; 120-133; 143-150: dark greenish gra (564/1). 55-75; 100-120; 133-143: light greenish gray (56Y8/1). Fluidal banding dips 45°-70°.
					5					various colors	As SECTION 4. Colors are: From 0-110 cm: greenish gray (5GY6/1) to dark greenish gray (5GY4/1), and some dusky red (5R3/4 mottles and veins (Fe-oxide). 110-150 cm: light greenish gray (5GY8/1) with mottles of grayish orange pink (1DR8/2) and moderate real (5R4/6). Fluidal banding dips 50°-6




































































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