The Shipboard Scientific Party

SITE DATA

Data Occupied: 3 November 1972

Data Departed: 7 November 1972

Time on Site: 108 hours

Position:

lat 29°37.05'S long 112°41.78'E

Water Depth (from sea level): 4696 corrected meters (echo sounding)

Water Depth (from drill floor): 4706 corrected meters (echo sounding)

Bottom Felt At: 4712 meters (drill pipe)

Penetration: 346 meters

Number of Holes: 1

Number of Cores: 41

Total Length of Cored Section: 346 meters

Total Core Recovered: 248.75 meters

Percentage Core Recovery: 71.9

Oldest Sediment Cored:

Depth below sea floor: 304.3 meters Nature: Green-gray clay Age: Aptian Measured velocity: 1.6 km/sec

Basement:

Depth below sea floor: 0.35 sec DT (seismic profiler) 304.3 meters (drilled) Inferred velocity of basement: 1.75 km/sec Nature: Altered basalt

Principal Results: The principal result was to determine the basement age is aptian. Another important result is the discovery of the hiatus in the Upper Cretaceous and lower pirt of Paleocene.

BACKGROUND AND OBJECTIVES

Two different models for the origin of the southeastern Indian Ocean have recently been proposed:

1) by the spreading of the sea floor between Australia and India (Crawford, 1969; Heirtzler, 1971; Veevers et al., 1971) or between Australia and southeast Asia (Ridd, 1971; Audley-Charles et al., 1972). The essence of these reconstructions is that Gondwanaland was continuous westward from southwest Australia.

2) as an original oceanic embayment in Gondwanaland—the *sinus australis* of Dietz and Holden (1971), which is implicit in the classical reconstruction of Du Toit (1937), and which is supported by Smith and Hallam (1970) and McElhinny and Luck (1970). McKenzie and Sclater (1971) reassembled the Gondwanaland fragments back to the Upper Cretaceous (75 m.y. ago) from sea-floor magnetic anomalies, and found that:

"It is also difficult to understand how the fit proposed by Du Toit (1937) and refined by Smith and Hallem (1970) could evolve to the Upper Cretaceous arrangement proposed in this paper. Recent paleomagnetic observations (McElhinny and Luck, 1970) have strongly supported Du Toit's fit, and it is therefore unlikely that major errors are present in Du Toit's reconstruction. Perhaps the solution of these difficulties will come from the results of the Deep Sea Drilling programme."²

The Du Toit reconstruction conflicts with the geological evidence from Western Australia (Veevers et al., 1971) of a Permian to Jurassic facies change from non-marine in the south to marine in the north. Falvey (1972) attempted to reconcile this conflict by adopting an extended eastern margin of India (now beneath the Tibetan Plateau), which lay against the southwest Australian margin.

A common element of these models, as modified above, is that there was land west of Australia and that this land has been dispersed by sea-floor spreading. The next stage is to outline the salient details of the assembly of the continents and the timing of their dispersal.

The geology of Western Australia sets constraints on the timing: rupture occurred in the basal Cretaceous approximately 130 m.y. ago (Veevers et al., 1971), and the first phase of dispersal was completed by the Upper Cretaceous. The geology of Western Australia is not so pertinent in detailing the reassembly and subsequent spreading except for one point, the lineament of the southwest margin and the parallel Perth and Carnarvon

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²Page 504.

basins and the Darling Fault. Veevers et al., (1971) and Veevers (1971) relate the development of the Permianto-Jurassic Perth and Carnarvon basins to rifting along a lineament that became a spreading ridge in the Cretaceous, so that the lineament is normal to subsequent spreading. On the other hand, Falvey (1972) interprets the lineament as the trace of a transform fault, and thus parallel to spreading. From these interpretations, Veevers et al., (1971) imply west-east spreading, whereas Falvey (1972) implies northwest-southeast spreading. It was to test such ideas that Site 259 was drilled.

SITE SURVEY

Details of the site area (Figures 1, 2) are provided by reflection profiles of Lamont-Doherty Geological Observatory (Conrad-11), the Australian Bureau of Mineral Resources (BMR 18), and by a survey from Glomar Challenger, all navigated by satellite fixes. The site area is bounded by the 4900-meter-depth contour to form a southwest-sloping elevated area with a sharp ridge (4200 m) in the northeast part of the map area, and smaller scarps in the south and west. The sediment on the elevated area is no thicker than 0.4 sec, and is acoustically transparent. Outside the immediate area the transparent layer is thicker (≥ 1.0 sec) and is overlain by stratified sediments. It was the discovery (by Conrad-11) of Paleocene sediment (L. H. Burckle, personal communication) on the elevated area that made this an attractive site for drilling. The basement reflector has the character and the shape of oceanic basalt.

Site 259 is located in an elongate steep-walled depression in the acoustic basement which is filled with 0.35 sec of sediment.

OPERATIONS

The approach to Site 259 was essentially along Conrad-11's track out of Fremantle. As the proposed site was approached the track of Glomar Challenger was found to be about one mile to the east of Conard's track but parallel to it. Although there was a general similarity to Conrad's seismic record, significant differences were noted. Glomar Challenger turned on a southerly, then northeasterly course to cross both Conrad's track and to cross our own track and to observe the east-west extent of the sedimentary cover. It was soon realized that the bottom and subbottom topography were somewhat more complex than originally supposed. The ship turned and made two more runs parallel to Conrad's track before finally selecting a place that appeared to have 350 meters of sediment of the same type as that found on Conrad's track.

The beacon was dropped at 0430, 3 November 1972 while underway at 7.5 km/hr (4 knots). Survey gear was retrieved and the ship positioned over the beacon. The bottom hole assembly and drill pipe were run in and bottom tagged at 4712 meters. The hole was spudded at 1700, 3 November and continuously cored to a total depth of 5058 meters or 346 meters below sea floor. Details of the coring are included in the coring summary, Table 1.

Weather conditions were generally good and positioning was excellent, allowing rig-floor operations to proceed unhindered. Basement was reached at 307.5 meters below sea floor in Core 33 and 38.5 meters of basalt were cored.

A deviation survey was run on Core 39 and indicated a deviation of 2° at 5042 meters. Coring was completed at 0600, 7 November. The drill string was pulled and we were underway to Site 260 at 1630, 7 November.

LITHOLOGY

Site 259 was drilled to a depth of 346 meters. The hole penetrated a sedimentary sequence 304.3 meters thick ranging in age from Quaternary to Lower Cretaceous. The sediments rest on a basement of highly altered basalt breccia that grades downward into relatively fresh, moderatly brecciated basalt.

The sedimentary sequence can be divided into four distinct units based on color and composition (Table 2).

Unit 1 (0.0-60.0 m)

Unit 1 consists of approximately 60 meters of lightorange to light-brown, interbedded, nanno ooze, nanno clay, and zeolite clay. The sediments are stiff, well bedded, and usually highly contorted. Individual beds range from 1 to 5 cm in thickness.

Nanno ooze and zeolite clay are the dominant lithologies. The ooze typically is about 80% nannos, 15% zeolite, 5% clay and mica with minor amounts of dolomite rhombs. Coarse fractions (greater than 62μ in diameter) contain traces of heavy minerals, volcanic glass, and sponge spicules. The zeolite clay has an average composition of 75% clay and 25% zeolite with small quantities of quartz, feldspar, and nannos. Coarse fractions consist chiefly of quartz, forams, and micarb fragments with traces of sponge spicules, fish remains, silicoflagellates, and black micronodules.

Numerous small nodules ranging from 0.5 to 2 cm in diameter are present between 30- and 33-meters depth. Some nodules are dark gray to black, possibly manganese or iron oxide in composition. Others are light gray and consist of dolomite- or zeolite-cemented clay.

Fossils are abundant throughout Unit 1, consisting chiefly of coccoliths, discoasters, forams, and sponge spicules with lesser quantities of Radiolaria, diatoms, and silicoflagellates. Several thin sandy layers in the upper part of the unit consist largely of shallow-water forams and carbonate fragments in a clay matrix. These layers, which range from 2 to 5 cm in thickness, are poorly sorted but show no obvious grading. The presence of shallow-water forams suggests transport of material from the shelf into deeper water.

The upper part of Unit 1 is Quaternary in age whereas the lower part ranges from lower Eocene to upper Paleocene. A major disconformity probably exists in the unit somewhere between 9 and 27 meters in depth. The base of the unit at 60 meters is marked by a fairly distinct color change to dark or dusky yellowish brown and by a marked decrease in the abundance of nannofossils.

Unit 2 (60.0-103.0 m)

Unit 2 consists of about 43 meters of dark or dusky yellow-brown zeolite clay and zeolite-rich clay. The clay is strongly contorted with vertical striping and no



Figure 1. Location of Site 259 and generalized stratigraphic columns of Site 259 and adjacent sites.



Figure 2. Site area showing ships' tracks, bathymetry, and sediment isopachs.

bedding is preserved. Typical specimens are about 50%-65% clay, 20%-25% zeolite, 10%-20% cristobalite and microcrystalline quartz, and a few percent of heavy minerals, quartz, fledspar, and nannos. Fossils are scarce in the clays but forams are present in most coarse fractions. Olive-black nodules are locally present between 80 and 90 meters in depth. These consist of zeolite and opaque material, possibly iron oxide.

Unit 2 is Cretaceous. The base of the unit is marked

by a slight color_change to yellow brown and light brown and the reappearance of abundant nannofossils.

Unit 3 (103.0-154.0 m)

Approximately 50 meters of zeolite-rich nanno clay with lesser zeolite-rich clayey nanno ooze comprise Unit 3. The sediments are moderate yellowish brown and light brown in color with some greenish-gray zones. Most of the clay has been strongly disturbed and brec-

TABLE 1 Coring Summary, Site 259

	Date (Nov		Depth From Drill Floor	Depth Below Sea Floor	Cored	Recov- ered	Recov- ery
Core	1972)	Time	(m)	(m)	(m)	(m)	(%)
1	3	2100	4712.0-4720.0	0.0-8.0	8.0	7.7	96
2		2220	4720.0-4729.5	8.0-17.5	9.5	0.0	0
3	4	0000	4729.5-4739.0	17.5-27.0	9.5	2.8	29
4		0130	4739.0-4748.5	27.0-36.5	9.5	9.4	99
5		0255	4748.5-4758.0	36.5-46.0	9.5	6.7	71
6		0425	4758.0-4767.5	46.0-55.5	9.5	9.5	100
7		0550	4767.5-4777.0	55.5-65.0	9.5	7.5	79
8		0720	4777.0-4786.5	65.0-74.5	9.5	9.0	95
9		0845	4786.5-4796.0	74.5-84.0	9.5	9.5	100
10		1010	4796.0-4805.5	84.0-93.5	9.5	8.5	90
11		1145	4805.5-4815.0	93.5-103.0	9.5	6.5	70
12		1315	4815.0-4824.5	103.0-112.5	9.5	9.5	100
13		1450	4824.5-4834.0	112.5-122.0	9.5	9.5	100
14		1620	4834.0-4843.5	122.0-131.5	9.5	8.9	94
15		1755	4843.5-4853.0	131.5-141.0	9.5	9.4	99
16		1935	4853.0-4862.5	141.0-150.5	9.5	5.3	56
17		2120	4862.5-4872.0	150.5-160.0	9.5	5.3	56
18		2300	4872.0-4881.5	160.0-169.5	9.5	3.5	37
19	5	0045	4881.5-4891.0	169.5-179.0	9.5	7.7	81
20		0220	4891.0-4900.5	179.0-188.5	9.5	8.0	84
21		0400	4900.5-4910.0	188.5-198.0	9.5	6.0	63
22		0550	4910.0-4919.5	198.0-207.5	9.5	2.9	30
23		0715	4919.5-4929.0	207.5-217.0	9.5	9.5	100
24		0900	4929.0-4938.5	217.0-226.5	9.5	5.2	55
25		1040	4938.5-4948.0	226.5-236.0	9.5	8.0	84
26		1225	4948.0-4957.5	236.0-245.5	9.5	9.4	99
27		1410	4957.5-4967.0	245.5-255.0	9.5	6.8	72
28		1550	4967.0-4976.5	255.0-264.5	9.5	8.1	85
29		1735	4976.5-4986.0	264.5-274.0	9.5	4.5	47
30		1925	4986.0-4995.5	274.0-283.5	9.5	9.0	95
31		2120	4995.5-5005.0	283.5-293.0	9.5	8.0	84
32		2255	5005.0-5014.5	293.0-302.5	9.5	5.0	53
33	6	0105	5014.5-5024.0	302.5-312.0	9.5	2.0	21
34		0415	5024.0-5027.0	312.0-315.0	3.0	1.3	43
35		0630	5027.0-5033.5	315.0-321.5	6.5	2.4	37
36		0940	5033.5-5038.5	321.5,326.5	5.0	4.8	96
37		1245	5038.5-5043.0	326.5-331.0	4.5	2.3	51
38		1620	5043.0-5047.5	331.0-335.5	4.5	4.5	100
39		1920	5047.5-5049.0	335.5-337.0	1.5	1.5	100
40	7	0115	5049.0-5052.5	337.0-340.5	3.5	0.3	10
41		0600	5052.5-5058.0	340.5-346.0	5.5	3.1	55

Note: Echo-sounding depth (to drill floor) = 4706 meters; drill-pipe length to bottom = 4712 meters.

ciated by drilling and bedding is rarely apparent. Where present, beds range from 1 to 10 cm in thickness and consist of alternating color bands. Nodules are sparse to absent in Unit 3 but many of the clays contain white specks of carbonate, chiefly dolomite rhombs. An average composition for the ooze is 50%-60% nannos, 20% cristobalite and microcrystalline quartz, 10% clay, 12% zeolite, 2% forams, and 1% dolomite rhombs. Trace amounts of quartz, opaques, gypsum, and dolomite rhombs occur in the coarse fraction.

Unit 3 is Lower Cretaceous (Albian) in age. The base of the unit is indicated by a sharp color break and a virtually complete disappearance of nannofossils.

Unit 4 (154.0-304.3 m)

Unit 4 is a very uniform sequence of dark greenishgray claystone about 150 meters thick. The claystone is stiff to indurated and usually highly deformed. Rare preserved beds are regular repetitions of 1 cm-thick layers of dark greenish-gray and olive-black clay. Clay and microcrystalline cristobalite and quartz make up 95% of average specimens; detrital quartz, feldspar, and heavy minerals make up the remainder. Most specimens also contain trace amounts of feldspar and chlorite. The mineralogy suggest a volcanic origin for much of the clay, either by weathering of exposed volcanics or by alteration in situ of pyroclastic material.

Nodules are common throughout Unit 4 and range from 2 mm to 10 cm in diameter. Most are light gray in color and consist of dolomite- or zeolite-cemented clay. A few pyrite nodules, up to 4 cm across, are present and small crystals of pyrite are common in the coarse fraction. Euhedral gypsum crystals, 1-2 mm long, are also present in the clays. Fossils are rare in this unit although forams are usually present in the coarse fraction. Tiny opaque globules generally less than 1 mm in diameter are also present in the clays.

Unit 4 is Lower Cretaceous (Aptian) in age with the upper boundary corresponding to an unconformity between Albian and Aptian.

TABLE 2 Major Lithologic Units of Site 259

Interval (m)	Unit	Description	Age	Thickness (including gaps) (m)	Cores
0.0-60.0	1	Soft, grayish-orange clay-rich nanno ooze and zeolite clay	Quaternary to upper Paleocene	60.0	1-7
60.0-103.0	2	Dark-yellow-brown zeolite clay	Cretaceous	43.0	7-11
103.0-154.0	3	Light-brown zeolite- rich nanno clay and clayey nanno ooze	Lower Cretaceous (Albian)	51.0	12-17
154.0-304.3	4	Greenish-gray zeolite-bearing claystone	Lower Cretaceous (Aptian)	150.3	17-33
304.3-346.0		Basalt		41.7	33-41

Basalt Basement

Basalt basement was encountered at 304.3 meters and drilled to a depth of 346 meters. Approximately 20 meters of basalt were recovered from this interval. The sediments immediately overlying the basalt are somewhat indurated but show no signs of an intrusive contact.

The upper part of the basalt consists of light-gray to greenish-gray breccia. This grades downward into darkgray fine-grained basalt at about 25 meters below the sediment-basement contact. The breccia consists of angular to subangular fragments of altered basalt up to 50 cm across. The fragments grade from light gray in the center to light green near the margins and are typically separated by zones of black to dark-green altered glass(?) 1-2 cm thick. Veinlets and small masses of calcite and chlorite are common throughout the breccia and a few geodes of quartz are locally present. The solid basalt is dark gray to light gray in color and is generally fresher than the breccia.

The basalt is fine- to medium-grained and sparsely porphyritic with a variolitic, intergranular to intersertal groundmass. The average grain size ranges from 0.3 to 0.5 mm with a few phenocrysts up to 5 mm long. A few small vesicles up to 0.5 mm in diameter are usually present.

Typical specimens consist of small, ragged laths of labradorite set in a matrix of poorly crystallized clinopyroxene, iron oxide, and clay minerals. The clinopyroxene forms radiating, sheaf-like masses indicating incipient crystallization from glass. Rare microphenocrysts of clinopyroxene are augite $(2V, \simeq 50^{\circ})$ and the groundmass crystals are probably the same. The iron oxide is chiefly magnetite which occurs in small clusters and fern-like masses of octahedra. Olivine is present in some specimens but never makes up more than 2%-3% by volume. It occurs in small corroded crystals, 0.3-0.5 mm in diameter, which have been replaced by clay minerals and iron oxides.

All of the basalts are altered but the degree of alteration decreases markedly with depth. In the most altered specimens, green montmorillonite makes up 30%-40% by volume and replaces olivine, glass, clinopyroxene, and, occasionally, plagioclase. In the

fresher specimens, clay rarely exceeds 10% and occurs chiefly as vesicle fillings and as pseudomorphs after olivine.

The mineralogy of the basalt indicates a tholeiitic composition typical of oceanic basement. The grain size and texture indicate rapid cooling, probably in contact with water, suggesting extrusion onto the sea floor.

Preliminary Interpretation

Site 259 is at a water depth of 4712 meters. Based on the presence of abundant calcareous nannofossils and forams in the uppermost sediments, this depth is above the present carbonate compensation level which averages about 5000 meters in the world's oceans. Carbonate fossils are abundant throughout Unit 1 suggesting that similar conditions existed in the early Tertiary. However, a hiatus occurs in the sequence between the Quaternary and the lower Eocene so most of the Tertiary record is missing.

Terrigenous sand and silt are absent in the sediments at this site; however, a few thin layers of sand-size shallow-water forams occur in Unit 1. These suggest transport of material from a shallow-water-shelf environment into deeper water. The transporting mechanism was presumably turbidity currents and the thinness of the layers and their bimodal grain distribution (sand and clay) suggest deposition at the distal ends of the flows.

Unit 2, which is Cretaceous in age, consists chiefly of zeolite clay and contains only 1%-2% calcareous fossils. The existing fossils show strong dissolution effects suggesting that the scarcity of nannofossils in this unit is due to solution of carbonate below the carbonate compensation level rather than to changes in planktonic production. In contrast, the reappearance of relatively abundant nannofossils in Unit 3 probably reflects a return to shallow water conditions.

Unit 4, consisting of dark greenish-gray and black cristobalite claystone makes up nearly half of the sedimentary sequence. Zeolites generally make up 1%-2% of this unit and trace amounts of quartz, feldspar, heavy minerals, and chlorite are ubiquitous. Calcareous fossils are very rare and there is no evidence of bioturbation in the sediments. Finely divided pyrite is abundant and presumably is responsible for the dark colors. Small, euhedral gypsum crystals are also present throughout the unit.

The paucity of carbonate fossils and absence of terrigenous sand or silt suggest deposition in quiet, deep water. The mineralogy of the clay suggests a volcanic origin for much of the unit, either by weathering of exposed volcanic rocks or by alteration in situ of pyroclastic material. The quartz grains are typically rounded and frosted and probably are windblown in origin.

The contact between the basal sediments and the basaltic basement appears to be normal, with no evidence of baking, hydrothermal alteration, or metasomatism in the sediments. The basement consists of fine-grained tholeiitic basalt typical of oceanic spreading centers.

BIOSTRATIGRAPHY AND PALEONTOLOGY

General

The Cenozoic is represented by only fragmentary material. The top core (0.0-8.0 m) consists of Pleistocene foraminiferal nanno ooze. Cores 2 and 3 (8.0-27.0 m) have no diagnostic fossils. The zeolite nanno ooze of Core 4 (27.0-36.5 m) is lower Eocene, and the interval between Cores 4 and 8 (36.5-65.0 m) is upper Paleocene. The remaining sediments (Cores 8-33; 65-304.3 m), composed largely of claystone and some zeolite nanno ooze, are Cretaceous, mainly Albian and Aptian.

Cores 11-17 (93.5-154.0 m) ore the most interesting paleontologically. These sediments are very rich in calcareous faunas such as benthonic and planktonic foraminifera, nannoplankton, and Calcisphaerulidae. Also present are ostracods and, in one level, bivalvia. Radiolaria are abundant. Of considerable interest is the sudden appearance of this rich fauna within Core 17, apparently not connected with a major lithological change. No calcareous forms are present below this interval; only sparse arenaceous foraminifera and Radiolaria are present. Above this interval the sediment contains chiefly palynomorphs which abound in dinoflagellates, spores, and pollen. See Figure 3.

Biostratigraphy

Quaternary: Core 1 (0.0-8.0 m). Globorotalia truncatulinoides truncatulinoides Zone; Emiliana huxleyi to Pseudoemiliania lacunosa Zone. Foram nanno ooze, partially affected by CaCO₃ dissolution. Globorotalia inflata, a temperate-water form is predominant, G. truncatulinoides truncatulinoides is common. Faunal assemblages indicate that the sediments in the lower part of the core were deposited in waters slightly warmer than those of the upper part.

Age unknown: Cores 2, 3 (8.0-27.0 m). Core 2 had no recovery. Core 3 consists of zeolite clay with very rare, inconclusive, arenaceous foraminifera.

Lower Eocene: Core 4 (27.0-36.5 m). Globorotalia subbotinae/Globorotalia formosa formosa Zone; Discoaster lodoensis to Tribrachiatus contortus Zone. The complete sequence of lower Eocene nannoplankton zones is represented in the zeolite clay and nanno ooze of this interval, which is apparently strongly condensed due to CaCO₃ dissolution. Benthonic and planktonic foraminifera such as *Globorotalia aequa*, *G. formosa*, and some *Acarinina* species, are present. Radiolaria are very rare and poorly preserved.

Upper Paleocene: Core 5-Core 7, Section 3 (36.5-60.0 m). Globorotalia velascoensis Zone; Discoaster multiradiatus Zone. Only the upper part of the upper Paleocene, consisting of zeolite nanno ooze is present. However, it is considerably thicker than the lower Eocene above. Rare, probably transported, planktonic foraminifera are represented mainly by some Acarinina species. Calcareous nannoplankton are common, Radiolaria are rare and poorly preserved.

Cretaceous s.I.: Core 7, Section 4-Core 11, Section 4 (60.0-99.5 m). A lithologic break occurs in Core 7 between Sections 3 and 4. Radiolaria become frequent in Core 7, Section 4, suggesting a Cretaceous age. Paleocene foraminifera continue to the base of Core 7 but are probably due to contamination as a result of drilling disturbance. Below the much reduced and incomplete Tertiary section is an interval that contains almost exclusively Radiolaria, indicating a Cretaceous s.l. age. The Cretaceous-Tertiary boundary is placed within Core 8 based on the Radiolaria.

Lower Cretaceous, Albian: Core 11, Section 5-Core 17, Section 3 (99.5-155.0 m). This interval of zeolite-rich clay contains abundant microfaunas with benthonic and planktonic foraminifera, including different species of *Hedbergella*, nannoplankton, Radiolaria, and Calcisphaerulides. Present in lesser numbers are ostracods, and, in one level, bivalves. The nannoplankton indicate middle Albian (*Prediscosphaera cretacea* Zone), the planktonic foraminifera indicate Albian, and the benthonic foraminifera indicate upper Albian. The composition of the very rich *Hedbergella* fauna, poor in species, indicates comparatively deep open-sea deposition, still above the lysocline, and a temperate climate.

Lower Cretaceous, Aptian: Core 17, Section 4 to Core 33 Section 1; (155.0-304.3 m). A distinct change occurs between Sections 3 and 4 Core 17. Nannoplankton drops from an estimated 55% by volume in the upper interval to zero in the lower unit. The richcalcareous and arenaceous foraminifera are replaced by a monotonous and very poor arenaceous fauna. Arenaceous foraminifera are frequent only in the basal part of the lower unit. Calcisphaerulidae and ostracods are completely absent in the lower unit, but palynomorphs, with dinoflagellates are abundant, indicating a lower Aptian age for Cores 25-33 and an upper Aptian age for Cores 18-24. The absence of calcareous fossils suggests deposition below the lysocline, or possibly somewhat shallower if deposition was in a cold-water environment. The poor arenaceous foraminiferal fauna is of deepwater character, but may also have lived at somewhat shallower depth in colder, higher latitudes. The absence of planktonic organisms and the presence of bottomliving arenaceous foraminifera exclude euxinic conditions.

Paleontology

For more information on the individual fossil groups briefly discussed below, refer to the special reports in this volume.

FORAMINIFERA	NANNOPLANKTON	RADIOLARIANS	AGE	CORES DEPTH NO. (m)	LIT	THOLOGIC DESCRIPTION
Globorotalia truncatulinoides	Emiliani huzleyi to Peeudoemiliania lacunosa	Radiolaria common to rare-moderate well preserved. Quaternary assem- blage of moderate	QUAT	1 8	± ± ± 	
G. subbotinae-G. formosa	Discoaster lodoensis, Marthasterites tri- brachiatus, Discoaster binodosus, Marthaster-	diversity includ- ing Euchitonia elegans; Carpo- canistrum Spp., Ommatartus tetra- thalamus and sev- eral species of	L. EOCENE	2 17.5 3 27 4	- <u>7-7</u> - -7-0	Unit 1 Pale orange, grayish-orange and dark yellowish orange clay rich nanno ooze and interbedded zeolite clay.
Globorotalia velascoensis	ites contortus.	the Pyloniidae family. Radiolaria range from common to rare; preservation very poor in all	UPPER PALEOCENE	5 36.5 46 6 55.5 7 65		
	Barren of calcareous nannoplankton.	cases due to re- crystallization and corrosion. Radiolaria gener-	CRETACEOUS s. 1.	8 9 84 10 93.5	z = -2 z = -2 z = -2 z = -2 z = -2 z = -2 z = -2	Unit 2 Dark yellowish-brown to dusky yellowish brown zeolite clay and zeolite- rich clay.
Rich association of benthonic foraminifera, Hedbergella infraore- tacea, H. ploinispira, Hedbergella Sp.	Prediscosphaera cretacea	ally abundant but very poorly preserved; spe- cies diversity low.	LOWER CRETACEOUS ALBIAN	103 12 12.5 13 122 14 131.5 15 141 16	<u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u> <u>z</u>	Unit 3 Moderate yellowish brown and light brown zeolite- rich nanno clay and clayey nanno ooze some greenish- gray zeolite rich clay.
	Barren of calcareous nannoplankton.	Few to rare Radiolaria; very poorly preserved.	Eous	17 160 18 19 20 189.5 21 198 22 207.5 23		Unit 4 Dark greenish-gray and olive black bedded zeolite bearing clay.
Association of primitive anenaceous foraminifera.			LOWER CRETACI APTIAN	24 217 226.5 25 236 26 245.5 27 255 28 264.5 29 264.5 29 274 30		
		Barren		283.5 31 293 32 302.5		
			**	34 312 35 321.5 36 326.5 37 331 38 335.5 37 337	× × × × × × × × × × × × × × × × × × ×	Basalt basement greenish-gray to light gray basalt breccia and tholeiitic basalt.

Figure 3. Biostratigraphic zones, Site 259.

Foraminifera

All samples of Core 1 (0.0-8.0 m) contain very rich asseblages of predominantly planktonic foraminifera. *Globorotalia inflata* is strongly predominant in most samples. Present in decreasing order of frequency are:

Abundant: Globorotalia inflata, G. dutertrei, Globigerina bulloides.

Common: Globorotalia truncatulinoides truncatulinoides, G. menardii, G. crassaformis, Globigerinoides trilobus s.l., G. conglobatus.

Rare: Pulleniatina obliquiloculata, Globorotalia cultrata, Hastigerina siphonifera, Globigerinita glutinata, Globigerinoides ruber.

Present in only some of the examined samples are: Orbulina universa, Globorotalia tumida tumida, Candeina nitida, Sphaeroidinella dehiscens.

Benthonic forms are not numerous, but are rich in species.

Core 3 (17.5-27.0 m) only contains very rare specimens of *Glomospira* and *Haplophragmoides* which are insufficient for an age determination.

Core 4 (27.0-36.5 m) contains a very poor planktonic foraminiferal fauna. In Section 3, casts of the *Globorotalia aequa-G. subbotinae* and the *Globorotalia subbotinae-G. formosa* groups are present along with complete specimens of Acarinina primitiva, A. soldadoensis, and Chiloguembelina aff. wilcoxensis. The planktonic foraminifera are accompanied by numerous benthonic forms.

Core 4, CC, Core 5-Core 7, Section 3 and the top of Core 8 (36.5-60.0 m) contain an assemblage of planktonic and benthonic foraminifera including *Acarinina acarinata, A. mckannai, A.* aff. primitiva, *Globorotalia imitata, Chiloguembelina* aff. wilcoxensis. Typical benthonic forms are *Tappanina*, and *Aragonia*.

A rich assemblage of planktonic and benthonic foraminifera is present in Sample 14, CC, including several species of *Hedbergella*. Similar but less abundant foraminifera are also present in Cores 11 (lower part)-13, 15, 16, and upper part of Core 17(103.0-155.0 m). The fauna indicates a temperate water, open-sea environment of deposition, with a water depth equal to shelf slope or deeper, but above the lysocline.

Core 17 Section 4 to Core 30 (155.0-283.5 m) contain poor arenaceous foraminifera of mainly dwarfed *Glomospira* sp., *Ammodiscus* sp., *Bathysiphon* sp., *Trochammina* sp., *Reophax* sp., and *Textularia* sp. The arenaceous foraminifera of Cores 31-33 (283.5-304.5 m) are somewhat richer and more diversified.

Nannoplankton

Rich well-preserved temperate- to warm-water Pleistocene assemblages are present in Core 1. A complete sequence of lower Eocene to upper Paleocene nannoplankton zones (*Discoaster lodoensis*, *Tribrachiatus orthostylus*, *Discoaster binodosus*, *Tribrachiatus contortus*, *Discoaster multiradiatus*) is present from Core 4-Core 7, Section 3 (27.0-60.0 m). The assemblages are composed of solution-resistant taxa, *Discoaster*, *Fasciculithus*, and some *Coccolithus*. Based on a typical Albian *Prediscosphaera cretacea* Zone assemblage, present in Cores 13-14, the interval Core 11, Section 5-Core 17, Section 3 (99.5-155.0 m) is referred to the middle Albian. The lower part of the sedimentary section, Core 17, Section 4-Core 33 (155.0 m to the basalt contact at 304.3 m), is barren of calcareous nannofossils.

Radiolaria

Quaternary Radiolaria range from common to absent in Core 1 (0-8.0 m). Preservation is moderately good but decreases rapidly with depth in the core.

Cores 3-30 (17.5-283.5 m) contain Radiolaria in fluctuating abundances. Preservation is so poor that few specimens can be identified and none can be used for age determination. Some specimens from Core 8 and lower, however, can be classified on the family or even the generic level indicating the material is Cretaceous, i.e., *Dicolocapsa* spp., *Dictyomitra* spp., and several Spongodiscidae.

Palynomorphs

Only the Cretaceous sediments, Cores 10-33, were investigated for palynomorphs. In the upper part, Cores 10-17 (84.0-160.0 m) are barren. Cores 18-33 (160-304.3 m) contain rich dinoflagellate assemblages typical of the Aptian. *Hystrichosphaera ramosa* var. *reticulata* in Cores 19-23 is suggestive of upper Aptian; ?Oligosphaeridium asterigerum in Cores 25-27, and ?Dingodinium albertii in Core 29 to Core 33, Section 1 are indicative of lower Aptian.

Spores and pollen occur in Core 18 to Core 33, Section 1. The presence of *Crybelospirites stylosus* in Core 31 suggests a *C. stylosus* Zone age (Neocomian) for the lower part of the interval. Higher in the section, at least up to Core 20, the flora is correlated with the Microcachryidites-Assemblage of Neocomian to Aptian age.

Calcisphaerulidae

Seven *Pithonella* species which occur exclusively at Site 259 are distinguished in the Albian Cores 12-17 (103.0-160.0 m). These species were not recorded at Site 260, Cores 9-15 (234.0-300.5 m). Based on other fossil evidence, however, this interval at Site 260, which contains a *Pithonella* association of nine different species, is approximately equivalent in age to Cores 12-17, Site 259.

This difference in *Pithonella* species—a similar difference is apparent for the ostracods—indicates a difference in the environment, or a slight age difference within the Albian of the two intervals.

Pithonella is common to abundant in most samples examined in Site 259, Cores 12-17; more frequent than in any other Leg 27 site, except for Cores 31-33 of Site 261, where they occur in floods.

Ostracods

Six species in open nomenclature occur in Cores 12-17 (103.0-160.0 m). *Arculicythere*? sp. A, by far the most numerous, occurs in most of the examined samples. All six species are present in Sample 14, CC, which contains the richest ostracod fauna of the interval. The dominance of smooth forms, mainly Bairdiids and

Cytherella, and the virtual absence of the Cythereis group (only one specimen present), which is indicative of Lower Cretaceous near-shore and shelf deposits, indicates a depth of deposition deeper than the shelf.

Bivalves

Numerous exfoliated fragments of Aucellina sp. A (Family Buchiidae) were recovered from Sample 14, CC (131.5 m). A definite species identification is prevented by the small size of the fragments. However, the fragments closely resemble the juvenile phase of Aucellina cf. gryphaeoides from the Great Artesian Basin of Australia. The fragments are most closely comparable to Aptian-Albian species, particularly with Albian forms. This age is in agreement with the age obtained from foraminifera and nannoplankton.

Fish debris

Fish remains are fairly frequent in the Paleogene, Cores 4-8 (27.0-74.5 m); in Cores 9-33 (74.5-304.3 m) they are rare, and often absent altogether. They are particularly rare in the basal part of the Cretaceous interval.

GEOCHEMICAL MEASUREMENTS

Alkalinity, pH, and salinity data are summarized in Table 3 and in Figure 4. Analytical methods used for determining these values are discussed later in this section.

Alkalinity

Alkalinity values at Site 259 are in general low, showing a mean of 2.64 which is close to the mean of 2.49 for seawater at this site. A low alkalinity value is commonly ascribed to a slow rate of deposition (see earlier DSDP reports) and this would also seem to be compatible with the probable slow rate of sedimentation at Site 259. However, the consistent decrease of alkalinity with depth suggests that postdepositional modification of interstitial solutions by such factors as precipitation of minerals from solution and the increase in overburden pressure may also be significant.

pH

Only three "punch-in" values were obtained; consequently, this discussion will be limited to the "flow-through" results. Throughout most of the column, pH values are below that of seawater, ranging from 7.24 to 8.09. At most DSDP sites pH values are below those of seawater, consequently Site 259 fits into the normal pattern. There is also a consistent increase in pH with depth (Figure 4) which is probably the result of diagenetic changes although the nature of the processes involved is unknown.

Salinity

There is a consistent decrease in salinity with depth at Site 259, (Figure 4) with values ranging from 35.50/00 at the top of the hole, to $34.1^{\circ}/_{00}$ at a depth of 282 meters. This decrease may result from the removal of ions from solution, particularly Ca++ and SO4--, as gypsum is present in Units 3 and 4. The apparent co-existence of gypsum with a comparatively dilute solution (maximum salinity 35.5% (00) is somewhat problematical, particularly as the gypsum may have crystallized out in localized pockets isolated from the bulk of the interstitial fluids.

PHYSICAL PROPERTIES

Bulk-density, sound-velocity, porosity, vane shearstrength, and residual pore-pressure measurements were made on materials recovered at Site 259. Density, porosity, and sonic velocity are plotted alongside the site summary sheets. Continuous GRAPE density (and porosity) are plotted alongside the core photographs. A description of the testing procedures and discussion of wet bulk-density determinations and vane shear results are included in later chapters in Part IV.

Site 259 produced the largest quantity of sediment suitable for physical properties testing of the five sites drilled. Coring was continuous through the unconsolidated sediment column and in most cases each core had some portion which displayed only mild layer distortion. These areas of relatively low disturbance are toward the bottom of the core, usually in the lowest section. Due to layer color variation, photographs from Cores 1, 8, 14, and 16 show well the wide range of drilling disturbance. Many of these color changes are due to alternating inorganic clay and nanno ooze layers. Since both materials have undergone the same history after deposition, the variation of their properties with depth may be compared.

Density

Wet bulk density was determined on fresh unsplit core sections by continuous GRAPE measurements and by

		pI	H			
Sample (Interval in cm)	Depth Below Sea Floor (m)	Punch-in	Flow through	Alkalinity (meq/kg)	Salinity (°/00)	Remarks
		8.00	8.21	2.44	36.0	Reference seawater
1-6, 0-6	6.5-6.56	7.46	7.24	3.13	35.5	
5-5, 0-6	44.5-44.56	7.30	7.35	2.93	35.2	
10-5, 0-6	90.5-90.56	7.13	7.40	2.83	34.9	
15-6, 0-6	139.5-139.56	-	7.15	2.54	35.2	
20-5, 0-6	185.5-185.56	0-0	7.53	2.83	34.4	
25-3, 0-6	230.0-230.06	1 1 - 1	7.60	2.35	34.1	
30-6, 0-6	282.0-282.06	-	8.09	1.86	34.1	
		8.34	8.29	2.54	35.8	Reference seawater

TABLE 3



Figure 4. Geochemical analysis of interstitial water at Site 259.

weighing each section. A third value was determined from syringe samples taken routinely for water contents or immediately adjacent to each vane shear test, at which point two syringe samples were taken. In eight other locations samples were taken for shore-laboratory determination of density. The profile drawn on the site summary corresponds to GRAPE data from relatively undisturbed sediments, considered the most accurate indicator of in situ density at Site 259. The sectionweight determinations are considerably lower than these GRAPE data because although both methods are density averages of a specific sediment volume, the section-weight method includes more low-density disturbed sediment. (In most cases, core sections containing relatively undisturbed sediments also contain large areas of lower density disturbed material and often a slurry of even lower density. The section weight averages all materials within the section. However, even in unusual core sections, which may contain 150 cm of relatively undisturbed sediment, the section weight method should measure a lower density, because of the inevitable higher disturbance of the sediment immediately adjacent to the core liner. While section weight is a total core cross-sectional average, the GRAPE is not and averages a lower percentage of the outer material within its pencil-thin gamma ray beam.) The syringe values, while more accurate than the section weight data, are slightly lower and less accurate than GRAPE determinations.

Major density variations occur at Site 259 within the top 55 meters of the sedimentary column. However, these variations appear more dependent on sediment type than on depth. Adjacent layers of nanno ooze and zeolitic clay do not have the same density. Within the surface core ooze density is 1.50-1.70 g/cc while clay density is only 1.25-1.35 g/cc, even though the clay is in a deeper layer. Nanno clay, found between these sediments, shows an intermediate density of about 1.45 g/cc. Clay density rises to about 1.60 g/cc at 55 meters with most of the increase occurring in a zone of low recovery between 20 and 35 meters. Fewer measurements were made on nanno ooze. Ooze density is approximately 1.70 g/cc at 45 meters and 1.80 g/cc at the 55-meter level. Behavior is difficult to trace because, although layers of each sediment type were found within most cores, much of the material was too highly disturbed for testing.

Below 55 meters the sediment is predominantly inorganic clay. Density variation is less dramatic, even where nanno ooze is abundant in layers between 105 and 154 meters. Density rises steadily to 1.80 g/cc between 55 and 154 meters, then drops to 1.65 g/cc. (Nanno ooze is not found in the clay below 154 meters.) Density then remains relatively constant until 255 meters and again rises gradually to 1.80 g/cc above the basalt.

The density fluctuations correspond closely to the four units identified on the basis of color and composition. These units include the intervals of 0.0-60.0 meters, 60.0-103.0 meters, 103.0-154.0 meters, and 154.0-304.3 meters.

Porosity

Porosity plotted on the site summary sheets was determined from continuous GRAPE readings and from syringe samples. Only those syringe samples which were taken in relatively undisturbed material are shown. Porosity is directly related to wet bulk density and mineral grain density. Variation in porosity follows variation of bulk density very closely. The effect of grain density is as follows. In two sediments with the same bulk density, the sediment with a higher grain density also has the higher porosity. Grain density of mostsediments, however, does not vary enough to overcome the close dependence of porosity on bulk density. A continuous readout is available for GRAPE measurements by working with the GRAPE density trace shown beside the core photographs. A variable porosity scale has been included above the wet bulkdensity scale for this purpose. The correct scale to use depends on grain density of the sediment. Although a porosity plot may be computer-produced in a manner similar to GRAPE density plots, it is necessary to specify the grain density for each sediment layer. In some cores mineralogical variations occur often enough to make this procedure approach hand reduction.

Sonic Velocity

Sonic-velocity measurements shown on the site summary sheets were made on relatively undisturbed sediment and on basalt specimens prepared by trimming to produce parallel flat surfaces. Sonic velocity increases gradually from 1.50 km/sec at the sediment surface to 1.60 km/sec at 295 meters. Velocity in basalt ranges from 3.4 to 5.0 km/sec in the deepest basalt recovered. On some specimens, measurements were made in the vertical direction, perpendicular to bedding. (Sonic velocity on Glomar Challenger is normally measured in the horizontal direction-parallel to bedding planes.) In another study, velocity was measured on the same basalt specimens at 5°C and at 21°C. Comparison of the measurements in both studies is inconclusive due to data scatter. Poor Hamilton frame transducer contacts and fissures in the basalt are suspected to have caused the scatter.

CORRELATION OF SEISMIC REFLECTION PROFILE WITH DRILLING RESULTS

At this site, drilling records show that 304.3 meters of sediment and 42 meters of basalt were penetrated below the sea floor. The seismic profiles show a single strong reflector at a depth of 0.35 sec. It was assumed that this reflector is basaltic basement. (Figure 5.)

The 0.35-sec DT through 304.3 meters of sediment gives an average velocity of 1.73 km/sec. The velocity measured in the ship laboratory is approximately 1.5-1.6 km/sec.



Figure 5. Seismic profile approaching Site 259.

A sonobuoy was run while the ship was on-station drilling. Side reflections prevented this record from showing other possible reflectors in the sediment column. The steep slope of the basement is evident from the isopachs shown in the site survey map (Figure 2). Additional evidence for the slope of the basement was provided when the ship left this site. As the seismic gear was placed in the water, the basement depth was found at 0.3 sec, i.e., 0.05 sec less than indicated on the record when the ship came on this station. The incoming profile is regarded as the more reliable so the accepted value is 0.35 sec.

Compressional wave velocities measured for basalt from this site range from 3.4 km/sec at the top of the recovered basalt to about 5.0 km/sec for the lower pieces recovered about 38.5 meters below the sedimentbasalt interface. These values are somewhat lower than average basalt and probably result from the fact that the basalt is altered.

SUMMARY AND CONCLUSIONS

Summary

Site 259 is located in 4706 meters of water near the southeastern edge of the Wharton Basin. Near the site,

seismic profiles show 0.4 sec or less of transparent sediment over basement, and a piston core taken by Lamont-Doherty Geological Observatory in 1967 shows that the sediment at or near the surface is Paleocene (L. H. Burckle, personal communication). One hole was drilled at Site 259 through 304.3 meters of sediment and 41.7 meters of basalt.

Four sedimentary sequences were recognized. Lightbrown nanno clay and ooze and zeolite clay extend from the sea floor to a depth of 60 meters. Then follow 43 meters of dark-yellow-brown zeolite clay, 51 meters of zeolite-rich nanno clay, and finally 150.3 meters of darkgreenish-gray cristobalite claystone overlying basalt breccia with a depositional contact. The greenish-gray basaltic breccia grades 70wnward into dark-gray finegrained basalt at about 25 meters below the sedimentbasalt contact.

Fossils show that all but the top 60 meters of sediment is Cretaceous. The Cenozoic section is subdivided as follows: 8 meters is Quaternary, 8-27 meters is of unknown age, 27.0-36.5 meters is lower Eocene (indicated by a complete sequence of four nannoplankton zones), and 36.5-60.0 meters is upper Paleocene, confirming L. H. Burckle's determination (personal communication) from a piston core. The boundary between the upper Paleocene and the Lower Cretaceous is not visible as a physical break, but this may be due to mixing during coring or to original reworking. The upper, calcareous part of the Cretaceous sequence is Albian, as indicated by nannoplankton and foraminifera. The lower part is Aptian with arenaceous foraminifera.

Accumulation rates calculated from these data are only approximate due to the uncertain location of precise biostratigraphic boundaries in the drilled section. Figure 6 was compiled on the arbitrary assumption that each biostratigraphic division identified was complete. This is demonstrable at the site for the lower Eocene only. The compiled curve therefore should show minimum rates only. The time-scale adopted here is shown in Figure 9.

In the terms of these approximations, the curve shows that the accumulation rate was 1.6 m/m.y. in the lower Eocene, 12 m/m.y. in the upper Paleocene, 14 m/m.y. in the Aptian. The only comment warranted here is that the upper Paleocene and Lower Cretaceous rates are of the same order of magnitude, and the lower Eocene rate is an order of magnitude smaller.

Conclusions

Deposition at Site 259 began in an unknown part of the Aptian with extrusion of basalt and the accumulation of green-gray clay. The basalt was altered before deposition of the overlying sediment, probably by rapid cooling in contact with seawater during extrusion onto the sea floor. A thickness of 150 meters of green-gray clay accumulated during the Aptian at a rate of approximately 25 m/m.y. Stagnation of the bottom water is suggested by the uniformly dark tone of the sediments; uniform, though crude, lamination; fairly abundant pyrite; the impoverished biota of primitive arenaceous foraminifera and rare Radiolaria; and the reduction of the other microfossils. The arenaceous foraminifera indicate cold or deep water, or both. The impoverished biota is probably due to selective solution. The stagnation at the very bottom must have been milde to have allowed the existence of the bottom-dwelling foraminifera but stronger anaerobic conditions must have prevailed a short interval beneath the sediment surface. The composition of the clay suggests a volcanic origin either by alteration in situ of volcanic ash or by inflow of the weathering products of exposed volcanic rocks.

A preliminary interpretation of the Aptian sequence is that it was largely derived from a volcanic source associated with the early stages of continental rupture and drift, represented ashore by the basal Cretaceous Bunbury Basalt (Veevers, 1971); and that it was deposited during the initial opening of an oceanic rift in which oceanic circulation was somewhat inhibited.

Sedimentation in the Albian, led to the accumulation, presumably above the lysocline, of 51 meters of lightcolored zeolite-rich nanno clay and nanno ooze similar to part of the approximately equivalent sequence at Site 257 (Leg 26) 500 km to the west-southwest. The sediment of this age at Site 257 rests above 15 meters of barren clay, on basalt, so that the entire unit of greengray claystone at Site 259 wedges out between these sites, presumably as a result of sea-floor spreading to the west.

The rest of the Albian is yellow-brown zeolitic clay with poorly preserved nannofossils and Radiolaria. The lack of foraminifera probably indicates deposition below the lysocline.

A long hiatus representing most of the Upper Cretaceous and lower part of the Paleocene has no obvious lithological expression, probably due to original reworking or to disturbance during coring. Reworked Upper Cretaceous Radiolaria from near the base of the upper Paleocene caly probably indicate a source on the Australian continental margin. The succession of upper Paleocene yellow-brown clay overlain by pale orange nanno ooze is interpreted as indicating a relative downward movement of the lysocline. Partial dissolution of the nannoplankton in the thin lower Eocene interval may explain its low accumulation rate. After the deposition of clay of unknown age, Quaternary nanno and foraminiferal ooze and zeolitic clay were deposited, and the carbonates were partially affected by dissolution. Ten km southward, at Station 150 of Conrad-11 (Figure 2), Paleocene ooze is exposed at the sea floor, showing that the 36.5 meters of Eocene and younger sediments of Site 259 are not preserved everywhere on the elevated area around Site 259.

The elevated position of Site 259 may have caused other gaps in the record. The Upper Cretaceous clay found at Site 257 is lacking at Site 259, and seismic profiles show that a fuller Cenozoic section is possibly present in the deeper part of the Perth Abyssal Plain.

PROGRAM TO STUDY AIR TRANSPORTED DUST AND SEDIMENTS SUSPENDED IN SURFACE WATERS

A limited program to study air-transported dust and suspended surface sediments was undertaken on this



Figure 6. Sediment accumulation rates, Site 259.

cruise. The Western Australian coast is one of several areas of the world where a major desert is near the sea coast and off-shore winds are likely to carry considerable dust out to sea. In addition, these winds could blow surface waters off shore, causing an upwelling of nutrient-rich water and stimulating a biogenic component to sediments deposited on the sea floor. A third factor which may influence the nature of the sediments is the discharge of rivers from the Australian drainage basins into the eastern Indian Ocean.

Informal estimates indicate that river discharge of sediments is small on the average, but flash floods may carry a large sediment load near the estuaries. Firm figures should be established for these rates.

It is the opinion of oceanographers that upwelling is not strong off Western Australia but may occur in a more significant manner off northwest Australia or in the Timor Sea area. It may also occur in local areas that may not be significant to the marine biologist or oceanographer but may be significant geologically.

To provide data to study these effects, the following efforts were carried out:

1) The number of dust particles/cc of air was measured 3-4 times each day by the weatherman.

2) Surface observations were made 3-4 times each day by the weatherman. These observations included wind, clouds, pressure, air and sea-surface temperatures, dew point, sea, and swell.

3) Sea and cloud observations were made each 2 hours by the ship's officers.

4) Satellite cloud maps were obtained directly from the satellite by the weatherman approximately once each day.

5) Weather maps for the western Australian region were received by radio approximately twice each day.

6) A sample of surface water was filtered through a preweighed millipore filter each 6 hours. Surface temperature and salinity were also measured at that time. Occasionally, silver filters were used so samples could be analyzed by X-ray spectroscopy.

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Image: Second	ite 259	н	Hole		_	Core	1	Cored I	inter	val:	: O-8 m	Cito S	50	Hol	0		Core	e 2	Cored	nter	val - 8	.0-17.5 m	
AC AC<	AGE ZONE	Longie	FORAMS 23	ARACT VINON	OTHERS 3	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FORAMS	FOS	ADIOLARIA DI CLER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION
S & AG	JuriERMary AG catulinoides truncatulinoides udoemiliania lacunosa (mixed assoc.) 201		AG AG A	Grunnen G G G FM	OTHERS	100. 11 1. 2				132 140 - 61 - 92 -130 XR - 488 - 599 - 66 - 110 - 147 - 147 - 10 - 110 - 110	Core is undisturbed to weakly deformed near base. Color chiefly very pale orange (10YR8/2) with alternations and layers of pinkish gray (SYR8/1), light brown (SYR8/6) and moderate orange (SYR8/4). FORAM BEARING NANNO 007E Smear slides 1-140, 2-61, 3-110, 4-26, 5-93 Composition Mannos 80% Forams 9% Sponge spicules 5% Rads 2% Carb. frags. 2% Diatoms 1% Silicoflag. 1% Second State State State Sand 3% Micarb 5% Sand 3% Micarb 5% Sand 3% Micarb 5% Carb. 5% Sand 3% Micarb 5% Sand 3% Micarb 5% Carb. 5% Sand 3% Micarb 5% Carb. 5% Sand 3% Micarb 5% Carb. 5% Sand 3% Micarb 5% Carb. 5% Carb. 5% Sand 3% Micarb 5% Carb. 5% Carb. 5% Sand 3% Micarb 5% Carb. 5% Carb. 5% Carb. 5% Carb. 5% Carb. 5% Carb. 5% Sand 3% Micarb 5% Carb. 5% Carb. 5% Carb. 5% Carb. 5% Carb. 5% Carb. 5% Sand 3% Micarb 5% Carb.	39Y	20NE 655	KE NONE FORAMS Z FORAMS	NE NONE NANNOS 252 0 NANNOS	CVVP RADIOLARIA TTA CVP RADIOLARIA	Cot Cot MOILDES 0 1 1	Balaw reeker 2 Salaw 2 1 1 2 1 2 1 2 1 2 1 2 1 2 1 2 1 2 1	LITHOLOGY V01D Cored LITHOLOGY		-75 -75	7.5-27.0 m 10YR5/4 74	LITHOLOGIC DESCRIPTION LITHOLOGIC DESCRIPTION Core is weakly deformed, stiff clay. Chiefly moderate yellowish brown (10YR5/4 with streaks and patches of dark yellowis orange (10YR6/6). <u>ZEOLITE SILTY CLAY</u> Smear slides 1-75, 2-75, CC <u>Texture Composition</u> Clay 69% Clay 7 Silt 31% Zeolite 2 Quartz T Feldspar T
	que (F) Globorotalla trunca (M) Emiliania huxityi to Pseud	(M) EMNITANIA NUXIEVI TO FSEUD	A A A A A A A A A A A A A A A A A A A	G RM RM dAAN 3NON dA	Fd	4 5 6				- 26 - 52 - 75 - 105 - 119 - 64 62 - 93	Uuartz 2% Feldspar 1% Nannos Tr. 0paques Tr. 5 BULK X-RAY (3.30 m) Calcite 96% Quartz 3% Mica 1% 10YR5/4 10YR5/4	Expla	natory	RP note	NONE NONE S	dhhh	Cot Cat					10YR5/4	

SITE 259

75% 25% Tr. Tr.

Sit	e 259	Но	le		C	ore 4		Cored 1	Inter	val:	27-36.5 m		Sit	e 2!	59	Ho1	e		Co	re 5	Cored In	iter	al:3	36.5-46 m	
AGE	ZONE	FORAMS	FOS CHAR SONNAN	RADIOLARIA	OTHERS 2	METERS	LI	THOLOGY 0.4 m zero section	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION	AGE		ZONE	FORAMS	FOSS	RADIOLARIA 21	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION	
		NONE	NONE	FVVP	Fd]	0.5- L 1.0-				- 40 75 -120	10YR5/4 10YR4/2 10YR4/2	Core mostly strongly deformed stiff clay or ooze. Colors chiefly moderate yellowish brown (10785/4) with streaks and patches of dark yellowish orange (10786/6 and a few patches of dark yellowish brown (10784/2). Chiefly pale yellowish orange in lower 2 sections.				?	CM		1	0.5	VOID	5	-110	Highly contorted stiff clay and cl ooze. Chiefly grayish orange (10TR streaks of dark yellowish orange (Grades to very pale orange (10YR8/ base.	ty nanno 7/4) with 10YR6/6). 2) near
	141	111 0190090	СМ	RVVP	Fd 2	2 -	Z			-675	74 10YR5/4	Class Class Call <						RVVP	^{id} 2	111111111		ß	67 - 75	ZEQLITE RICH CLAY Smear slide 1-110 Composition Clay Zeolite Nannos ZEQLITE BEARING 002E	84% 14% 2%
EOCENE	(F) G. subbotfnae - G. formosa (N) Tribrachiatus	M orthostylus 48	АМ	RVVP	Fd 3	5	Z 2 2	0Mn - 0Mn		XR - 75	26-28 10YR5/4 10YR7/4	MICARE RICH NAME OOZE Smear slides 5-75, 6-75, CC Texture Composition Silt 53% Nannos 70% Clay 46% Micarb 20% Sand 1% Dolomite rhombs Tr. ZEOLITE RICH MICARB AND NANNO 00ZE Smear slide 2-75 Texture Composition Clay 40% Nannos 55%	UPPER PALEOCENE	Clobander for affectantal	(N) D. binodosus	FM	СМ	NONE	-d 3			$\left\{ \right\}$	XR - 75	Smear slides 2-75, 3-75, 4-75, 5-7 <u>Texture Composition</u> '26-28 Silt 44% Zeolite Dolomite rhombs 10YR7/4 Minor dark nodules (N1) at top of BULK X-RAY (39.8 m) Calcite 76% Clinoptilolite 16% Quartz 4%	5, CC 90% 10% Tr. core,
LOWER		RM	RP	RVVP	Fd 4	-		07 07 07 07 07 07 07 07		65 - 80		Silt 46% Micarb 25% Sand 5% Zeolite 15% Clay 5% Iron Oxide Tr. Feldspar Tr. bolomite rhombs Tr. Core contains numerous dark (N1) nodules of zeolite cemented clay. BULK X-RAY (30.3 m)		(e)	s (N) T. contortus	FM	см	NONE	d 4	111111111111		$\left\{ \right\}$	- 75	Mica 25 Montmorillonite 25	
	() Discoaster himdos	PM	СМ	RVVP	Fd 5	-				-95	74 10yr8/6	Calcite 22% Quartz 17% Philipsite 4% Mica 3% Montmorillonite 3%			(N) D. multiradiatus	см	СМ	RVVP	'd 5	and the state of t			- 67 - 75	10yr8/2 74 10yr7/4	
		FM	АМ	RVVP	Fd (5				-75	10YR8/6		Exp	lar	natory n	CM	CM in c	thapt	Ca er 1	Core Ltcher			- cc	10YR7/4	
		FM	AM	RVVP	Fd	Core	in in		4	- cc															

SITE 259

Site 2	59	Hol	e		Con	'e 6	Cored	Inter	val:4	46-55.5 m	Site	259	1	lole			Core 7	Cored	Inter	val:	5.5~65 m	
AGE	ZONE	FORAMS	FOSS	RAD I OLARIA LI T	SECTION	METERS	LITHOLOG 0.5 m zero section	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE		FORAMS	ARACTE ARDIOLARIA	OTHERS 20	SECTION METERS	LITHOLOG	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION	
	D. multiradiatus			RVVP F	1	0.5-		1111111	- 75	Highly contorted stiff Clay and ooze. Chiefly grayish orange (10YR7/4) in upper and lower parts and pale yellowish orange 10YR7/4 (10YR8/6) in middle.					RVVP	Fd	0.5		ननननन	- 90	10YR8/2 10YR7/4 Strongly disturbed stiff clay Colors chiefly grayish orange with black streaks (MI) in upp and dark yellowish orange (107 smear with dark yellowish orange 10YR7/4 10YR7/4 Streaks in midd part chiefly moderate yellowis	nd ooze. 10YR7/4) r part 6/6) (10YR4/2) e. Lower brown.
	(N) T. contortus (N)	CM	CM CM	KVVP	2			1111111111	-975	ZEOLITE BEARING NANNO OOZE Smear slides 1-75, 2-75, 3-75, 4-75, 5-75, 6-60, 6-135, CC Texture Composition 95% 74 Clay S8% Nannos 95% 74 Silt 42% Zeolite 5% 10YR8/6 Forams Tr. BULK X-RAY (49.3 m) 2% 93%		sis tus		AM C	RVVP	Fd	2			-93	Chiefly zeolite bearing nanno zeolite rich clay with minor i <u>ZEOLITE BEARING NANNO 002E</u> Smear slides 1-90, 2-75, 3-75, 74 <u>Texture Composition</u> 10YR7/4 <u>Clay</u> 56% <u>Nannos</u> Silt 433 Zeolite Sand 1% Heavy mineral Feldspar Quartz	oze and on oxide. 5-80 90% 10% Tr. Tr. Tr. Tr.
	Densis		69	F	13	et al cons			XR - 75	Clinoptiolite 5% 26-28 Quartz 2%	PALEOCENE	ilia velascoens cer multiradia	(C M	RWP	Fd	3			XR - 75	26-28 ZEOLITE RICH CLAY Smear slides 4-75, 5-103, CC Composition Clay 10YR7/4 Zeolite Mica	80X 19X 1X
UPPER PALEOCENE	(F) GIODFOTAIIA VElasco er multiradiatus	СМ	CP	RVVP RVVP	14	transferration of the second		नननननननननननननननननननननननननननननननननननननन	- 75	10YR7/4 with streaks of 10YR6/6	Radiolaria) UPPER	ascoensis (F) Globrota om above) (F) Globrota tiradiatus (N) Discoast	om above)	CM C HONE	d RWP	Fd	4			- 95 -75 -XR -95	N1 Feldspar Opaques 10YR6/6 Smear slide 4-95 Composition Tron 0xide Quartz N1 BULK X-RAY (58.8 m) 10YR6/6 Calcite 87% Quartz 3% Overtz 3%	Tr. Tr. 99% 1% Tr. Tr.
	(M) Discoaste	см	см	F	15	and contractions)_6 <u>7</u> _75	10YR7/4 74 10YR8/2	CRETACEOUS (based on	(F) Globrotalia vela (contamination fro (N) Discoaster mult	(contamination fro	CM. C	P RVVP	Fd	5			6Z - 80 -103	IOTR3/4 Micla 2.5 IOVR7/4 BULK X-RAY (61.0 m) 44% OVR5/4 Clinoptiloite 32% IOVR7/4 Mica 9% IOVR2/2 Montmorillonite 8% IOVR2/2 Montmorillonite 1%	
			СМ	HVVA FF	4 6			1.1.1.1.1)_ 60	10YR7/4	Exp	lanator	y n	CM (in ch	apte	Core Catch r 1			-cc	10YR5/4	
		СМ	СМ	RVVP	Cat	ore			-135 - CC	0786/6 10787/4												

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Site 259		Hole	ł		Cor	e 8	Cored I	nter	val:	65-74.5 m			Site	259	Hole	8		Cor	re 9	Cored In	terv	a1:7	4.5-84		
AGE	ZONE	FORAMS	HARACT VILLE SOUNAN	OTHERS 33	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	C.	LITHOLOGIC DESCRIPTION		AGE	ZONE	FORAMS	FOSS	RADIOLARIA 12 FI	SECTION	METERS	LITHOLOGY 0.5 m zero section	DEFORMATION	LITHO.SAMPLE	LIT	HOLOGIC DESCRIPTION	
		СМ	RP RV	P	1	1.0111111	Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z		- 62	10YR4/2	Strongly disturbed stiff clay. Ch dark yellowish brown with distorted and streaks of grayish orange (10Yk dusky yellowish brown (10YR2/1). ZEOLITE CLAY Smear Slides 2-25, 3-75, 4-75, 5-75	lefly d layers R7/4) and 5, 6-125					FVVP	d 1	0.5	2 2 2 7 2 7 2 7 2 7 2 7 2 7 2 7 2 7 2 7	$\left\{ \right\}$	75	10YR2/2 H 10YR4/2 d 10YR7/4 g 2 5	ighly disturbed stiff clay. Color chi usky yellowish brown with smeared lay nd patches of dark yellowish brown an rayish orange. <u>EOLITE CLAY</u> mear Sildes 1-75, 2-75, 3-75, 4-75, 5	efly ers d
		RP	WE	Fd	2		2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z		- 25	10YR4/2	Texture Composition Clay 76% Clay Silt 24% Zeolite Heavy minerals Nannos Quartz & feldspar BULK X-RAY (68.3 m)	70% 28% 1% 1% Tr.					WP	2	11 IIIIIII	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	$\left\langle \right\rangle$	GZ -	74 10YR2/2 10YR4/2	exture Composition Tay 85% Clay 11 14% Zeolite and 1% Nannos Heavy minerals	72% 25% 3% Tr.
		NONE	NONE NO		-	11111111	2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z	$\left(\right)$	XR	26-28	Palygorskite 33% Clinoptiolite 21% Mica 13% Cristobalite 12% Quartz 11% Montmorillonite 9% Gypsum 1%						R	-	muni	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	$\left\{ \right\}$	XR	10YR//4 B P C M Q Q M 26-28 G	ULK X-RAY (77.8 m) alygorskit 28% borbaorillonite 28% borbaorillonite 16% uuartz 11% ita 8% ypsum 2%	
laria) multiradiatus n from above)	n trom above/	NONE	RWP		3		2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2		- 75	10YR4/2			d on Radiolaria				AVVP	3	man	Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z	$\left(\right)$	75	10YR2/2	lagioclase Z%	
ACEDUS (based on Radio) (N) Discoaster (contaminatio)	(CON UMILI NA C10)	NONE	NONE		4			$\left(\right)$	- 75	10YR4/2			CRETACEOUS (base		NONE	NONE	AVVP	4	contraction.	Z Z Z Z <tb< td=""><td>$\left\{ \right\}$</td><td>-75</td><td>10YR2/2</td><td></td><td></td></tb<>	$\left\{ \right\}$	-75	10YR2/2		
CRET			4M	Fd	5	- 11 ft	2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z 2 Z		- 28 - 75	nodule 5YR7/2 10YR4/2 &							d/\\)	a 5	and and the	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2		GZ 75	74 10yr2/2		
		RM	CP a	Fd	6	hundring		$\left(\right)$		10YR4/2 & 10YR5/4 10YR4/2							FVVP	6	ultra ban b		$\left(\right)$	—75	10YR2/2 10YR5/4		
		NONE	S RVVB		Co Cat	re cher 2		(-125 - CC	10YR4/2 & 10YR7/4 10YR4/2							FVVP	C Ca	ore	2 2 Z Z Z Z Z Z Z Z Z Z Z Z Z	$\left<\right>$	cc	10YR5/4		



Site 2	59	Hole	ŧ		Cor	re 12	Cored	Inter	val:	103.0-112.5 m		Sit	e 259	Ho	le		0	Core 13	Cored In	terva	1:11	2.5-122 m	
AGE	ZONE	FORAMS	FOSSI HARAC SONNAN	OTHERS OTHERS	SECTION	METERS	LITHOLOG	DEFORMATION	LITHO. SAMPLE		LITHOLOGIC DESCRIPTION	AGE	ZONE	FORAMS	FO CHA SOUNAN	RACTE	OTHERS 2	METERS	L1THOLOGY 0.5 m zero section	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	
			AP	ANN	1	0.5			-75	10YR5/4	Highly deformed stiff clay. Chiefly moderate yellowish brown with streaks of grayish orange and light brown. Greenish gray at base. Grades from zeolite clay at top to zeolite and nanno rich clay at bottom. ZEOLITE AND NANNO RICH CRISTORALITE CLAY				AF	AVVP		0.5-		V V V V V V V V V	-80 -89	5GY5/1 Extremely deformed stiff c of greenish gray clay smear brown clay. <u>ZEOLITE AND CLAY RICH NANNN</u> Smear slides 1-80, 1-89, CL Texture Compositi Clay 266 Nances	ay. Fragments ed with light <u>00ZE</u> <u>n 50%</u>
	(u	RM	AP	CS O	5 2	11111111111)	74 10YR5/4	Smear slides 1-75, 2-75, 4-75, 5-75, 6-73, 6-73, 6-73, 6-73, 6-73, 6-73, 6-73, 6-73, 6-73, 6-75, 6-75, 6-75, 6-75, 6-75, 6-75, 6-75, 6-73, 6-73, 6-73, 6-73, 6-75,		("	RM	AP	AWP	Cs 2	2		A A A A A A A A A	GZ	Silt 23% Clay and Sand 1% Zeolite Forams 56Y6/1 BULK X-RAY (115.8 m) 74 Calcite 35% 5YR5/6 Cristobalite 21% SYR5/6 Clinoptiolite 16%	ristobalite 30% 20% Tr.
LE-UPPER ALBIAN	n etacea (Middle Albid	RM	AP	VAL	3	munnin			a XR —75	26-28 10YR5/4	DUCK AFARI (100.3 m) Cristobalite 26% Clinoptilolite 22% Calcite 15% Montmorillonite 12% Quartz 9% Palygorskite 9% Mica 3% Tridymite 2% K-feldspar 1%	-UPPER ALBIAN	itacea (Middle Albia	RM				3		A A A A A A A A A A	XR	Monumorilionite 8% Quartz 7% Palygorskite 5% Mica 3% K-feldspar 2% 26-28 Tridymite 1% S6Y6/1 Plagioclase 1% 5YR5/6	
ER CRETACEOUS, MIDD	(F) Upper Albia Prediscosphaera cr	RM	A	P	4	Innihinh			-75	10YR5/4	Plagioclase 1%	R CRETACEOUS, MIDDLE	(F) Upper Albian Prediscosphaera cre					and and a sector		~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~		50Y6/1	
FOM	(N)	RM	AP			unninnin	$\begin{array}{c} 1 \\ 2 \\ 2 \\ 2 \\ 2 \\ 2 \\ 2 \\ 2 \\ 2 \\ 2 \\$			10YR5/4		LOWER	(N)							A A A A A A A A A A		50Y6/1	·
		NONE	AP		5	ununu			75	74 10YR5/4				RM	AM	AVP	5	5 -		A A A A A A	GZ	5775/6 74	
		RM	CP		6	1111111111			-75	5Y7/2 5YR5/6 5Y5/2				RM	AM	CWP	Cs	6			-50	10YR5/4 5GY6/1 5GY6/1	
		RM	CM		Cat	ore tcher		11111	-cc	5GY6/1				RM	CM	CVVP		Core Catche		\langle	-cc		

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Site 259	н	lole		. (Core	14	Cored 1	Interv	a1:1	22-131.5 m	Site	259	Ho1	e		C	ore 1	5 Cored In	terv	rval:131.5-141 m
AGE ZONE	autre -	FOS CHAR CHAR SOUNNEN	RADIOLARIA 22 S	OTHERS ²⁰	SECTION	METERS	LITHOLOGY	DEFORMATION	LITRO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FORAMS	FOSSCHARA	SADIOLARIA DE LA	OTHERS	METERS	LITHOLOGY 0.4 m zero section	DEFORMATION	LITHOLOGIC DESCRIPTION
WER CRETACEDUS, MIDOLE-UPPER ALBIAN (F) Upper Albian) Prediscosobhaera cretacea (Middle Albian)		см м	AVVP AVVP	Cs 3	0. 1 1 2 2 3	.5			75 GZ 75 XR 75	10YR5/4Greatly disturbed stiff clay. Chiefly moderate yellowish brown to 1.5 meters then chiefly greenish gray with some admixed moderate yellowish brown. ZEOLITE AND CLAY RICH NANNO 002E Smear Slides 1-75, 2-75, 3-75, 4-75, 5-75, CC74 Tay 73% Silt 26% Some 1% Dolomite rhombs Heavy minerals Carb. frags.74Silt 26% Sand 1% Carb. frags.756/1Sand 1% Calcite Palygorskite Montorillonite26-28BULK X-RAY (125,3 m) Calcite Palygorskite Montorillonite26-28BULK X-RAY (125,3 m) Calcite Palygorskite Mica Sig Sig Sig26-28BULK X-RAY (125,3 m) Calcite Palygorskite Mica Sig Sig Sig Sig Palygorskite566/1Bit Sig Palygorskite Mica Sig Sig Sig Palygorskite26-28Bit Sig Palygorskite Mica Sig Sig Sig Palygorskite26-28Bit Sig Palygorskite Mica Sig Sig Sig 	ER CRETAGEOUS, MIDDLE-UPPER ALBIAN	(F) Upper Albian Prediscosphaera cretacea (Middle Albian)	CM AM ?		LVVP CVVP FVVP	0 1	0.5			 Highly deformed and brecciated stiff nanno ooze. Chiefly light brown in upper half and olive gray and greenish gray in lower half. -75 <u>ZEOLITE RICH CLAY BEARING NANNO OOZE</u> Smear Sildes 1-75, 2-75, 3-75, 4-75, 6-75, 6-75, 6-75, 0-75 Wannos 72% GZ 74 GZ 74 GZ 74 BULK X-RAY (134.8 m) Calcite 51% Cliay and cristobalite 10% Forams 1% BULK X-RAY (134.8 m) Calcite 51% Cristobalite 15% Cristobalite 15% Cristobalite 15% Cristobalite 15% Cristobalite 14% Wantorrillonite 5% Quartz 5% Plagioclase 2% K-feldspar 1% Tridymite 11%
1046	RI RI CI	м G G G G G A G	CVVP AVVP CVVP CVVP	e Ms O C	5 5 Cor	e.her			-75 -75 -75	5G6/1 N1 streaks 5G6/1 74 5G6/1 10YR5/4	ГОМЕ	(w)	AM RM FM	СМ	TAVP AVP AVP	ο 4 Cs 5 ο 6	5 5 Core			67 67 74 67 74 -75 -119 -CC

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Site 259	Hole			Co	re 16	Cor	red Int	terva	1:14	41-150.5 m		Site	259	H	ole		(Core 1	7	Cored In	terva	1:1	50.5-160 m			
AGE ZONE	FORMIS	ARAC SONNAN	TER SASHLU	SECTION	METERS	LITHO	LOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION		AGE	ZONE	CODANC	ECHANNO CHANNA	RADIOLARIA STOR	OTHERS a	METERS	LI	THOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESC	RIPTION	
LOWER CRETACEOUS, MIDOLE-UPPER ALBIAM (F) Upper Albian (N) Prediscoshaera cretacea (Middle Albian)	? RM FM	Even Even Aven Aven Aven	0 00 00	1 2 4	0.5		╴┟┞╞┠┟┟┝┝┝┟┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝┝		102 -74 105 132 -75	Deformed and brecciated stiff on Some layering preserved in Secti Chiefly moderate yellowish brown mixed greenish gray in smears ar 10YR5/4 5GY6/1 CLAY AND ZEOLITE BEARING NANNO (Smear slides 1-102, 2-74, 3-105, Texture Composition Clay 77% Mannos Silt 23% Zeolite Clay and cristo Forams Heavy minerals 74 BULK X-RAY (144.3 m) 10YR5/4 26-28 Mica 3% 10YR5/4 5YS/6 5GS/1 layered 5YR5/6	:e. n 4. with jpatches. <u>IZE</u> 4-75 10% walite 10% 1% Tr.	LOWER CRETACEDUS, UPPER APTIAN (?) LOWER CRETACEDUS, MIDDLE-UPPER ALBIAN	(F) Upper Albian	(N) Freutscoppinger (N) Freutscoppinger (N) Automiger 20 DDOF arenaceous hanthonin and mane allanttonic	FI JRON	NOME AVVP AVVP	Cs Cs	0.5 1 1.0 2 3 3			<pre></pre>	-110 <u>67</u> 5 XR -75 -141 -83 -1111 -CC	5YR5/6 10YR5/4 74 5YR4/4 10YR5/4 26-28 5YR4/4 5Y5/6 5GY5/2	Weakly deform Chiefly light in upper part green in lowe clay and zeol lower part cl ZEOLITE CLAY Smear slides Texture Clay 733 Silt 273 BULK X-RAY (1 Calcite Montmorilloni Clinoptilolit Palygorskite Quartz Mica Plagioclase	ed stiff ooze and clay. brown and yellowish bro and olive brown and yel r part. Upper part chief ite rich nanno ooze and ay and zeolite rich clay 3-75, 3-141, 4-83, CC <u>Composition Clay</u> Zeolite Dolomite rhombs Nannos Quartz, feldspar, heavy minerals ITE <u>BEARING NANNO 00ZE</u> 1-110, 2-75 <u>Composition</u> Nannos Clay Zeolites Heavy minerals Forams Dolomite rhombs 53.8 m) 67% te 8% e 7% 5% 3% 2%	wn low ly 68% 30% 1% 1% 1% 1% 1% 1% 1% 1% 1%



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Site 259	Hole	÷		Co	re 20	Cored	Inte	rval	:179-188.5 m			Site	259	Hol	e		Core	e 21	Cored In	iterv	al:1	188.5-198 m			
AGE ZONE	FORAMS	FOSSI HARAC SONNEN	L TER SGENTU	SECTION	METERS	LITHOLO	DECODMATION	NOL INVISION		LITHOLOGIC DESCRIPT	ION	AGE	ZONE	FORAMS	FOSSI CHARAC	OTHERS OTHERS	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE		LITHOLOGIC DESCRI	IPT ION	
LOWER CRETACEOUS, UPPER APTIAM	dd poor arenaceous	R F R R R	P P P P	1 2 3 5 Ca	0.5			e 1	56Y4/1 & 5Y2/1 5Y6/4 clot 5Y2/1 5GY4/1 5Y2/1 8 26-28 5 5 5 5 5 5 7 4 5 5 7 4 5 7 4 5 7 4 5 7 4 5 7 4 5 7 4 5 7 4 5 7 4 7 5 7 4 5 7 4 7 5 7 4 7 5 7 4 7 5 7 4 7 5 7 4 7 5 7 4 7 5 7 7 4 5 7 7 4 5 7 7 4 5 7 7 4 5 7 7 4 5 5 7 4 7 1 5 7 7 4 5 5 7 4 7 1 5 7 4 7 1 5 7 7 4 5 5 7 4 7 1 5 7 7 4 5 5 7 4 7 1 5 7 7 4 5 5 7 7 4 5 5 7 7 1 5 7 7 7 1 5 7 7 7 7	Weakly to moderat Alternating layer and olive black. CRISIOBALITE CLAY Smear slides 3-75 Texture Clay 88% Silt 12% BULK X-RAY (182.3 Montmorillonite Cristobalite Quartz Mica K-feldspar Plagioclase Kaolinite Pyrite Chlorite	<pre>wely deformed stiff clay. s of dark greenish gray , 4-75, 5-75 Composition Clay and cristobalite 97% Heavy minerals 2% Zaolite 1% Quartz, feldspar, Tr. m) 38% 23% 23% 23% 23% 23% 23% 23% 33% 23% 33% 3</pre>	LOWER CRETACEOUS, UPPER APTIAN		poor arenaceous		dAM3 FP DETER	0 1 1 2 4 5 6		VOID		-75 -32 -15 XR -75 GZ -75 GZ -75	56Y4/1 5Y2/1 5Y7/2 nodule 74 5Y7/2 specks 26-28 74 5GY4/1 5Y2/1	Weakly to stroi in part. Chiefl olive black clu CRISTOBALITE CI Smear slides 1: <u>Texture</u> Clay 86% Silt 14% BULK X-RAY (19' Montmorillonit Cristobalite Quartz Mica K-feldspar Plagioclase Kaolinite Chlorite	ngly disturbed clay - sou y interlayered greenish W. AY -75, 2-75, 3-75, CC <u>Composition</u> Clay and cristobalite Zeolite Quartz Feldspar Heavy minerals Chlorite, mica 1.8 m) 2 43% 2 28% 12% 7% 5% 3% 3% 12% 1%	95% 2% 1% 1% 1% 1% 1%
													1			+	+	-	TVID	1	-00	C			

FP NO FP

Core Catcher







Site2	59	Hole			Cor	e 28	Cored	Inter	val:	255-264.5 m	Site	259	Но	le		Co	ore 29	Cored I	nterv	al:2	64.5-274 m	
AGE	ZONE	FORAMS	RADIOLARIA 201	OTHERS 3	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FORAMS	FOS CHAR SONNEN	SIL CACTER	OTHERS	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	L	ITHOLOGIC DESCRIPTION
LOWER CRETACEOUS, LOWER AFTIAN		poor arenaceous	FF RVI	P D Pm	1 2 3 4 5 6		2/bo1 2/bo1		-55 GZ -75 -70	Meakly to moderately deformed stiff clay. Chiefly dark greenish gray and olive black in color with streaks and nodules of very light gray and yellowish gray. N8 streak CRISIDBALITE CLAY Texture Sitreak Composition Texture 577/2 Clay and cristobalite 58 Sitreak Specks Sitreak 74 BULK X-RAY (258.3 m) Montmorilionite 402 Cristobalite 94 BULK X-RAY (258.3 m) Montmorilionite 74 BULK X-RAY (258.3 m) Montmorilionite 75 Az 74 BULK X-RAY (258.3 m) Montmorilionite 75 Az 74 BULK X-RAY (258.3 m) Montmorilionite 75 Az 76 Az 77 Bulk X-RAY (258.3 m) Montmorilionite 76 Az 77 Bulk X-RAY (258.3 m) 78 Streak 79 Streak 70 Streak 71 Bulk X-RAY (258.3 m) 72 N8 73 Streak 74 Streak 75 Streak 76 Streak 77 Streak 78 Streak 79 Streak 8 Specks	LOWER CRETACEOUS, LOWER APTIAN		poor arenacedus		FVP	1 	0.5-	@001 @001 @001		-45 -52	5Y7/2 module 74 5Y7/2 module N8 26-28 module 10YR5/4 module	Weakly to strongly deformed stiff clay with dolomite-cemented nodules. Chiefly dark greenish gray and very light gray nodules. QUARTZ AND CRISTOBALITE RICH CLAY Smear Sildes 2-75, 3-45, CC Texture Composition Clay 70% Clay and cristobalite 97% Silt 22% Quartz 1% Feldspar 1% Heavy minerals 1% Chiorits, dolomite Tr. BULK X-RAY (267.8 m) Montmorillonite 52% Quartz 20% Cristobalite 14% Mica 11% K-feldspar 3%
		CM	RP		Cat	cher		11111					CM	NONE	RVP	c	Core atche	r	$\left \right\rangle$	-cc		



2106 523	Hole			Co	'e 32	Cored	Inter	rval:	293-302.5 m	510	259	HOI	e -	_	LOI	re 33	cored in	nuerv	01:3	U2.5-312 m	
AGE ZONE	FORAMS	ARACT VIAN OSSIL	OTHERS OTHERS	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FORAMS	FOSS HARA SONNAN	RADIOLARIA CAL	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE		LITHOLOGIC DESCRIPTION
LOWER APTIAN	ns	NOWE	D Prr	1 2 3	1.0	001 001 001		 	Moderately deformed stiff to very stiff clay. Chiefly dark greenish gray and olive black colors with yellowish gray and light gray nodules. SV7/2 Smear slides 1-75, 2-75 Smear slides 1-75, 2-75 Clay and micro- Silt 10% SV7/2 Smear slides 1-75, 2-75 Composition Clay and micro- Silt 10% Torystalline quartz 98% Feldspar 10YR5/4 Quartz, clorite, dolamite rhombs, zeolite N7 nodules	LOWER CRETACEOUS, LOWER APTIAN		poor arenaceous	NONE	P	2 C(Ca	0.5	VOID D C P D C P D C P D C P D C P D C P		+7 XR XR	5Y2/1 92-94 92-94	Brecciated olive black stiff clay to 1-110 cm. Drilling change to basalt at 304.3 meters. Basalt breccia. Fragments highly altered with cores of gray N6 surrounded by altered glass 106/5/2 grayish green and grayish black (N2). Veinlets and vugs of carbonate. BULK X-RAY (303.4 m) (nodule) Dolomite 52% Quartz 28% Montmorillonite 10% K-feldspar 4% Siderite 3% Mica 2% Pyrite 1% BULK X-RAY (304.9 m) (veinlet) Dolomite 84% Montmorillonite 8% Siderite 5% Plagioclase 2% Quartz 1%
0S,	aceo			11	1					Sit	259	Hol	e		Co	re 34	Cored I	nterv	al:	312-315 m	
LOWER CRETACED	poor aren			4	tin filling	VOID				AGE	ZONE	FORAMS	FOSS	RADIOLARIA CITER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION
				5			2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	1 2 2 1		Ext	lanatory	notes	; in	chapt	1 Ca	0.5 1.0 tcher				5GY8/1 geode	Basalt breccia with fragments having gray core (N6) surrounded by altered glass of graytsh green color (10&Y5/2). Some zones of dark reddish brown (108.74). Veinlets of calcite. A few quartz geodes and opal crusts.
	CM	NONE	-	Ca	ore	VOID															

SITE 259

Site	259	Hole	Core 35	5 Cored In	nterv	a1:3	15-321.5 m	Sit	e 25	59	Hole	i	C	core 37	Cored In	terva	:326.5-331 m	
AGE	ZONE	FOSSIL CHARACTER SUNNAN SMD10LAR SMD10LAR SMD10LAR	SECTION	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE		ZONE	FORAMS	FOSSIL HARACT VINNON SOUNDIDE	OTHERS B	METERS	LITHOLOGY	DEFORMATION	L11HU.3MMPLE	LITHOLOGIC DESCRIPTION
			0.5- 1 1.0- 2			e 2	Basalt breccia. Fragments have gray (N7) cores and rims of greenish gray (SGY6/1). Freshest glass is black (N1). Some zones of dark reddish brown (10R2/4). Veinlets of calcite.							2	VOID		28	Gray basalt (N3) with some patches of grayish red (10R4/2). Numerous veinlets of calcite.
			Core	$\mathcal{O}_{\mathcal{O}}$				Sit	e 2	59	Hole	2	0	Core 38	Cored In	terva	:331-335.5 m	
Site	259	Hole	Core 36	; Cored Ir	iterv	a1:3	21.5-326.5 m	AGE		ZONE	SMNS	FOSSIL HARACT SON	ER HEKS	METERS	LITHOLOGY	FORMATION	HO. SAMPLE	LITHOLOGIC DESCRIPTION
AGE	ZONE	FORAMS INANNOS RADIOLARIA OTHERS	SECTION	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION				FO	NAN	10	0.5		DE	5	Basalt - dark greenish gray with veinlets of calcite. Non-vesicular, slightly porphyritic, fine-grained. Slightly
			0.5-1 1.0- 2 3	VOID			Basalt breccia. Chiefly dark gray (N3) with rims of grayish green (106Y5/2). Some zones of dark reddish brown (10R3/4). Veinlets and patches of calcite and opal. Alteration decreases with depth. 562/1 N3	Exi	pla	natory	notes	inch		2			5R4/2 patches	

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Site 259	H	ole			Core	39	Core	d Int	terv	/al:	335.5-337 m	Site	e 25	59	Hole	2		Co	re 41	C	ored In	terv	al:	340.5-346 m
AGE ZONE		NANNOS 23	RADIOLARIA 255	OTHERS ^B	SECTION	METERS	LITHOLO	GY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION	AGE		ZONE	FORAMS	FOSS HARA SONNYN	IL TER	SECTION	METERS	LIT	HOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION
					0. 1 1.	5 0	©			385	Basalt - dark greenish gray (56Y4/1) with veinlets of calcite. Fine-grained, non-porphyritic. Rare pyrite crystals in cavities. A few reddish zones (5R4/2).						C.	1	0.5	V	OID			Basalt. Medium light gray (N6) medium grained with a few plagioclase phenocrysts. Microvesicular. Calcite and chlorite(?) veinlets. A few red zones (5R4/2).
Site 259	H	ole		_	Core	40	Core	d Int	terv	al;	337-340.5 m								-					
		F0 CHA	RACTI	R	N	s			NOIL	APLE								2	1 3					
AGE ZONE		NANNOS	RADIOLARI	OTHERS	SECTI	METER	LITHOLO	.OGY	DEFORMAT	LITHO. SA	LITHOLOGIC DESCRIPTION							-						
					0. 1 1.	111111111111	VOID	200		P140	Basalt. Dark greenish gray (56Y4/1) with veinlets of carbonate. Non-vesicular, fine-grained, non-porphyritic.							3	the second s					
						1111												Ca	ore tcher	NOT	CORED			
					2	1111111111111	VOID					Exp	lar	natory	notes	in d	chapt	ter 1		1				
					Con	e her	NOT COR	ED																


























































259-30-1 259-30-2 259-30-3 259-30-4 259-30-5 259-30-6











SITE 259












SITE 259