# The Shipboard Scientific Party1

# SITE DATA

Date Occupied: 10 January 1973

Date Departed: 12 January 1973

Position: 63°56.99'S; 105°09.34'E

Water Depth: 3529 corrected meters (echo sounding)

Water Depth (adopted): 3544 meters (drill pipe from rig floor)

Total Penetration: 474.5 meters

Number of Cores: 20

Total Section Cored: 189.5 meters

Total Section Recovered: 65.6 meters

Percentage Core Recovery: 35%

**Oldest Sediment Cored:** 

Depth below sea floor: 474.5 meters Lithology: Clay and chert Age: Middle Oligocene

Principal Results: Clayey sediments totaling about 474 meters thick and ranging in age from Quaternary to late Oligocene were cored at Site 268. Ice-rafted pebbles and granules are common well down into the lower Miocene part of the section, and isolated granules occur in the Oligocene sequence. Sediments in which chert has formed were deposited prior to the interval of abundant ice rafting—in the Oligocene and earliest Miocene. Thin sand and silt laminae found in the upper 150 meters probably represent turbidite deposition. Silty clays in the lower two-thirds of the section show few structural characteristics of turbidites and the dominant sedimentation process is inferred to have been related to deep ocean currents.

### BACKGROUND

Site 268 is located on the lower continental rise just north of the Knox Coast of Antarctica (Figure 1) in a water depth of about 3500 meters. There have been no



previous geophysical surveys at this site, although magnetics data acquired aboard *Glomar Challenger* enroute indicate that the underlying crust is older than anomaly 20 (50 m.y.). A magnetic quiet zone may underlie the site. The total sediment cover at this site is expected to exceed 1.5 km, if conditions here are similar to surveyed areas along the lower continental rise to the east. The *Challenger* seismic profile in the vicinity of Site 268 is shown in Figure 2.

The objectives at this site were to extend the reconnaissance profile of holes nearer shore in the hope of recovering terrigenous material bearing on the geologic history of this sector of Antarctica, and to generally examine glacial, sedimentological, and volcanic history of the adjacent continent. Several hours of presite seismic surveying were required to ensure the absence of potentially hazardous geological structures.

## **OPERATIONS**

The area originally intended for Site 268 was briefly surveyed by *Glomar Challenger*. However, because of the presence of numerous icebergs, this area was abandoned. A new site some 30 miles to the southeast and relatively free of icebergs was surveyed. The site chosen was on the lower continental rise and on a gentle slope to the northwest. Several subbottom reflecting horizons were present.

Following completion of the survey (see Figure 3), the towed gear was retrieved and the ship reversed course and steamed back over the selected site using the PDR as a guide. The beacon was dropped at 1047 on 10 January in 3529 meters (PDR corrected) of water and positioning in the automatic mode was acquired by i200 hr.

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Figure 1. Location of Site 268 and bathymetry. Contours in fathoms (corrected). Solid line is track line for Glomar Challenger seismic profile shown in Figure 2.



Figure 2. Glomar Challenger acoustic reflection profile on approach to and departure from Site 268. Vertical scale is in seconds of two-way reflection time. The last portion of the approach profile (starting at point A) is the Challenger survey for Site 268, the detailed track of which is shown in Figure 3.



Figure 3. Detailed dead-reckoning survey by Glomar Challenger for Site 268. Point A corresponds to beginning of survey and to point A on the seismic reflection profile in Figure 2.

Hole 268 was spudded in at 2215 on 10 January and drilling with intermittent coring was carried out to a depth of 474.5 meters subbottom (Table 1).

No serious problems were encountered with either the drilling and coring operations or with the dynamic positioning system, although during the first several hours on site these operations were barely maintained because of rough seas and 35- to 40-knot winds.

The last core was brought onboard at 1120 on 12 January, the drill string and bottom-hole assembly were recovered and the ship got underway at 2155 12 January 1973.

## LITHOLOGY

The sequence cored at this site has been divided into three units, shown in Table 2. Several criteria could have been selected in order to break up the section into units. An arbitrary choice was to use the first and last occurrence of nanno ooze to define the limits of Unit 2. Several other lithological characteristics change at or near the unit boundaries defined in this way.

Ice-rafted pebbles are common in Unit 1; some icerafted granules are found in Unit 2; in Unit 3, here are very rare pebbles and granules, which are possibly of ice-rafted origin. The lowest granules found are in Core 19.

Semilithified sediment was found at a shallower depth at this site than in previous sites in the Southeast Indian Ocean. This sediment was also noticeably fissile, probably because of the abundance of lamination. The sedimentology of this site is discussed in greater detail (Piper and Brisco, this volume).

### Unit 1

Three main types of sediment are found in this unit. (1) Diatomaceous sediment, ranging from diatom ooze

Core	Date (Jan. 1973)	Time	Depth From Drill Floor (m)	Depth Below Sea Floor (m)	Length Cored (m)	Length Recovered (m)	Recovery (%)
1	10	2355	3554.0-3563.0	0.0-9.0	9.0	8.9	98
2	11	0129	3582.0-3591.5	28.0-37.5	9.5	8.6	91
3	11	0310	3610.5-3620.0	56.5-66.0	9.5	2.4	25
4	11	0420	3639.0-3648.5	85.0-94.5	9.5	2.8	29
5	11	0542	3667.5-3677.0	113.5-123.0	9.5	2.7	28
6	11	0740	3688.5-3698.0	134.5-144.0	9.5	0.6	6
7	11	0930	3698.0-3707.5	144.0-153.5	9.5	1.8	19
8	11	1050	3724.5-3734.0	170.5-180.0	9.5	3.7	39
9	11	1208	3753.0-3762.5	199.0-208.5	9.5	7.2	76
10	11	1329	3781.5-3791.0	227.5-237.0	9.5	9.5	100
11	11	1500	3810.0-3819.5	256.0-265.5	9.5	0.3	3
12	11	1635	3838.5-3848.0	284.5-294.0	9.5	1.5	16
13	11	1851	3857.5-3867.0	303.5-313.0	9.5	0.6	6
14	11	2044	3876.5-3886.0	322.5-332.0	9.5	0.3	3
15	11	2235	3895.5-3905.0	341.5-351.0	9.5	0.9	9
16	12	0125	3905.0-3914.5	351.0-360.5	9.5	1.2	13
17	12	0345	3933.5-3943.0	379.5-389.0	9.5	3.3	35
18	12	0635	3962.0-3971.5	408.0-417.5	9.5	4.3	46
19	12	0920	3990.5-4000.0	436.5-446.0	9.5	3.0	32
20	12	1120	4019.0-4028.5	465.0-474.5	9.5	2.0	
Total					189.5	65.6	35

TABLE 1 Coring Summary, Site 268

TABLE 2 Lithologic Units, Site 268

Unit	Lithology	Subbottom Depth (m)	Unit Thickness (m)	Age
1	Clay, silty clay, sand, and diatom ooze	0-~160	~160	Pliocene to Quaternary
2	Clay, silty clay, and clay nanno ooze	~160-228	~68	Early Miocene
3	Silty clay, laminated silty clay and clayey silt, and chert	228->474.5	~256.5	Mid-Oligocene or older to early Miocene

to diatom-bearing silty clay. Colors are olive- or yellowish-gray (5Y6/4-10Y4/2-5Y4/1). Pebbles, granules, and coarse sand, presumably ice rafted, are common. (2) Clay, silty clay, and clayey silt, with brownish color (10YR4/2-5YR3/4-5YR4/1). Diatoms are present in trace amounts only. Ice-rafted granules and coarse sand are in places concentrated in thin beds, but rarely scattered. Ice-rafted pebbles are rare. Laminae of medium silt are common (average one every 5-10 cm). (3) Beds of fine sand and coarse silt generally 3-10 cm thick. These generally have sharp bases, and grade finer upwards into clayey silt, and are interpreted as turbidites.

These three sediment types appear uniformly distributed throughout the unit. Cores 6 and 7 recovered 70 cm of granite, gneiss, and gabbro in lengths of up to 28 cm. This is presumed to be from ice-rafted boulders. The sequence of cored lithologies suggests at least seven boulders. There is no evidence for the cause of such a concentration of boulders, but it may have resulted from an iceberg melting sufficiently for the remaining sediment laden ice to sink.

#### Unit 2

Three main sediment types are found in this unit. (1) Olive-gray silty clay and clay, some diatom bearing, with infrequent coarse sand, granules and pebbles, presumably of ice-rafted origin. (2) Olive-gray nanno clay and clay nanno ooze, some diatom bearing. Coarse sand is very rare and granules and pebbles are lacking. Forams are rare and are generally fragmentary. (3) Grayish-red and brown clay and silty clay. This lithology is best developed in the lower part of Core 9, where thin laminae of medium silt are common (one every 1-2 cm). The silt includes about 15% size-sorted diatoms, Radiolaria, and sponge spicules; in the silty clay, biogenous material is less concentrated (2%-5%).

# Unit 3

This unit consists dominantly of silty clays and porcellaneous cherts, with minor amounts of partly chertified silty clay. The cherts appear to have formed by lithification of the silty clays, since all the variations in color and bedding structures in the silts can also be found in the cherts. The two main lithological types are: (1) Dark yellowish-brown (10GR4/2) laminated silty clay and (clayey) silt, similar to that in Unit 2. Silt laminae are very abundant, sometimes several per centimeter. Silt-sized diatoms and Radiolaria are concentrated in the laminae. In Core 14 and below there is a similar lithology of a grayish-olive color (10Y4/2, also 5GY4/1). In places, silt laminae are less abundant (one per 2 cm), and in parts of Cores 19 and 20 are completely absent. Similar chertified lithologies, both laminated and unlaminated, are common. Small amounts of diatoms and radiolarians are found in the silty clays throughout the unit. (2) Grayish-red and brownish-gray silty clay, with no laminae, and similar chertified lithologies.

Because of limited recovery, general statements about the abundance of chert are difficult. There is no evidence for marked variation in abundance in different parts of Unit 3.

Rare granules and fine pebbles are found in the cherts. A 4-cm granite pebble in Core 14 may be evidence of ice rafting, but could be downhole contamination similar to that in Unit 4 of Site 274. A granule-sized clast was found in a grayish-red chert in Core 11, Section 1. A dark greenish-gray chert at 19-2, 78-81 contains a thin bed of sand and granules.

In Core 10 and many of the lower cores, occasional silt laminae have a dark color which extends about 1 mm into the surrounding silty clay. This is apparently diagenetic, and associated with an unidentified opaque mineral seen in smear slides to coat or replace biogenic silica. At 10-2, 137, small dark concretions of granule size have formed in such a lamina.

Silt laminae are generally planar. No cross-lamination has been seen. In the lower part of the unit, many silt laminae are discontinuous and lenticular. A silt lamina at 18-3, 40 shows apparent small load deformation. The biogenic component of the silt laminae is generally size sorted, but varies in abundance with the abundance of biogenic material in the interlaminated silty clay.

Core 13-1 contains a trace of nanno ooze. A grayishred silty clay in 17-3 contains 1-mm flecks predominantly of carbonate, including some calcareous nannofossils. Occasional calcareous nannofossils are found in silty clays in Core 20.

At 18-3, 131-140, there are blocks of silty clay dipping at about 35° set in a slightly softer matrix. The sediment is stiff to semilithified. The occurrence may be due to either drilling deformation, or to synsedimentary slumping. The former interpretation is preferred. However, one chert bed at 15-1, 109 has microfaulting which apparently developed prior to chertification.

## Petrology of Ice-rafted Debris

Numerous clasts of continental-type rocks occur to a subbottom depth of 440 meters in Hole 268. Most were doubtless deposited from melting icebergs calved from ice shelves and glacier tongues around Antarctica.

Visual inspection and probing of each split core section revealed a total of 105 clasts larger than about 1 cm. Most by far were in Unit 1. Much of the silt and sand fraction in the core is also likely to be of glacial origin. Size, shape, and roundness measurements of the clasts are summarized in Chapters 22 and 23, this volume. Although the largest clasts in most clast-bearing intervals are between 3 and 6 cm long, much larger ones, up to at least 28 cm long, are present in Sections 6-1 and 7-1. The clasts are mostly equant to tabular, and with sphericities that range from 0.55 to 0.9. Mean roundness for the four core sections with 10 or more pebbles ranges from angular to subangular. The range in roundness is also small within each sample and only one or two very angular or rounded pebbles were found. Several subrounded pebbles appear to have been broken in half, probably during glacial transportation, but possibly during drilling.

The effects of glacial abrasion, implying transport at the base of a wet-base glacier, were seen clearly on only two clasts, both of which are gabbro. One has a flat face with fine parallel striae and the other a well-developed "keel" formed by abrasion as the pebble, held firmly in the basal ice of the glacier, was dragged through unconsolidated subglacial sediment. Several other clasts have indistinct striae or facets.

Clast lithologies were determined in hand specimens, and 20 of these were examined in thin section. Half of the clasts are granitic, and are typically fine- to coarsegrained, equigranular biotite-microcline granite with no prominent fabric, and biotite or biotite-hornblende granodiorite. A few are altered, like clast 5 of Section 5-1, which consists of extensively sericitized plagioclase, quartz, and elongate patches of chlorite, probably replacing biotite. Another (1, CC, 1) shows clear evidence of cataclastic deformation and mineralization. The rock consists of about 20% quartz with strongly undulose extinction and orthoclase and sericitize plagioclase in equal proportions. Quartz occurs as large, noncomposite or complexly sutured crystals a millimeter across, and as aggregates of fine crystals 50  $\mu$ m across ("wreaths of mortar") forming a matrix for the larger crystals. A gray metallic mineral, possibly galena, occurs in several veins up to 0.2 mm wide in this pebble. Two other pebbles (1, CC, 5 and 18) have traces of finegrained brassy sulfides.

Mafic plutonic rocks, which make up 9% of the total, are mainly unaltered fine- to medium-grained clinopyroxene gabbros, but include a sericitized hornblende gabbro (5-2, 1) and two probable diorites. In most of the gabbros the crystals appear randomly oriented, but in 7-1, 4 and 7-2, 1 apparent long axes of plagioclase laths are closely aligned, suggesting derivation from a layeredcumulate mafic intrusion. Both samples are strongly magnetic and carry about 10% of disseminated magnetite blebs about 100  $\mu$ m across.

The metamorphic rock types comprise mainly regionally metamorphosed granitic gneiss, and quartzofeldspathic and mafic schists. The gneisses are composed mainly of quartz, oligoclase, and microcline with minor brown biotite and a little chlorite. The mafic schists consist almost entirely of green hornblende. A pebble from 5, CC (9) consists mainly of sericite (pinnite)-flecked cordierite crystals a millimeter across, with scattered quartz and patches of fine-grained chlorite and epidote. Most crystals show the pseudohexagonal twinning which is characteristic of cordierite. The rock probably resulted from thermal metamorphism of a pelitic sequence.

The sedimentary rocks include several lithologies. The most common clasts are very quartzose, well-sorted, fine- to medium-grained sandstones with well-rounded grains in a quartz cement. They are very much like the Devonian orthoguartzites of the Beacon Supergroup of the Transantarctic Mountains. A second lithology is hematite-cemented quartzose sandstone with grains ranging from angular to rounded. This sandstone is medium grained and moderately sorted and was probably derived from a "red-bed" sequence. The third lithology is represented by two pebbles (5-2, 3 and 4) of greenish-gray semilithified claystone with about 10% silty quartz. A smear slide of 4 shows scattered sponge spicules, diatom fragments, and well-preserved wood fragments. The lithology and degree of lithification are comparable with lower Miocene strata in Holes 268 and 269. There are also three pebbles of very light gray indurated siltstone, probably from the same, or a similar, sequence to that from which the quartz-cemented sandstones were derived.

Two volcanic pebbles were recovered. The basalt (5-22, 14) has a few scattered phenocrysts of plagioclase and clinopyroxene set in a groundmass of semiopaque devitrified glass with flow-aligned plagioclase microlites and microphenocrysts (An<sub>70</sub>). A few microphenocrysts of clinopyroxene are also present. Sample 5-2, 10 is a lava of intermediate composition with a strongly developed trachytic texture and without visible mafic minerals. Most of the plagioclase microlites have parallel extinction, indicating a composition of about An<sub>20</sub>. Phenocrysts are about An<sub>35</sub>. The thin section also shows scattered patches of chlorite and epidote and an area 1.55 mm across of finely crystalline quartz.

Other pebbles include three of vein quartz and several fine-grained pebbles that have not been identified.

## Provenance

The scattered distribution of the clasts in the sediments, the variety of lithologies, and the glacial features of two of the pebbles, in fine-grained deep-sea strata, provide compelling evidence for transport by ice rafting from a continental landmass.

The relative abundance of clasts more than a centimeter across at this site and their absence about the same latitude at Site 269, 1200 km to the west, points to a strong local influence in the quantity of glacial debris deposited in this sector of the Subantarctic oceans. However, the presence of gabbros, possibly from a layered intrusion, of "red beds," and of orthoquartzites suggests that the rocks were derived from an area larger than the adjacent 30° sector of the continent, where such rocks are not known. Charnockitic rocks, which are common along the Wilkes Land coast, are inexplicably absent in the ice-rafted debris in this hole.

#### Interpretation

In contrast to the sites to the north, calcareous sediments are restricted to the early Miocene, and the lower part of the core has cherts and diatom-bearing sediments. The site is on the continental rise, and has received much sediment from the continent. Many of the laminae in Unit 3 appear the result of contour currents; in contrast, Unit 1 is dominated by turbidites.

### PHYSICAL PROPERTIES

Wet-bulk density measurements using the GRAPE technique were obtained on one or more sections from each of Cores 1-10, 17, and 19 at this site. Several additional determinations of wet-bulk density and porosity were obtained from the syringe samples. Sonic-velocity measurements were made on nearly all cores. Representative data are plotted in Figure 6. The sonic-velocity determinations for Cores 1-10 were made on the unsplit, lined cores while those for 11-20 were made on chunks of sediment.

With a couple of notable exceptions, sonic velocity increases regularly with depth at a rate of about 50 m/sec per 100 meters of depth. Wet-bulk densities are extremely variable within single cores, sections of cores, and from core to core. Based on the preliminary lithologic descriptions, most of these variations can be accounted for by changes in the relative proportions of sand, silt, clay, chert, and biogenous components. A good example of this is found in Core 10, Section 1, which contains a clay nanno ooze ( $\rho B = 1.60$  to 1.70 g/cc), silty clay ( $\rho B$ = 1.55 to 1.70 g/cc), and clayey silt ( $\rho B = 1.70$  to 1.80 g/cc).

Several very sharp changes occur in the physical properties of the sediments cored. Figure 4 shows the variation of sonic velocity, wet-bulk density, and acoustic impedance with depth. The peak in acoustic impedance at about 30 meters subbottom is caused by a sand layer (ca 30 cm thick) and correlates well with an acoustic reflector at about 0.04-0.05 sec two-way travel time which appears both on the profiler and sonobuoy records. Also, the very large change in physical properties at 115 to 140 meters subbottom is correlatable with a reflector at 0.13 to 0.14 sec.

Sonic-velocity measurements on chunks of semilithified and laminated clay and siltstone in several of the cores below Core 11 showed anisotropy, with velocities measured parallel to the bedding generally being 5% to 10% greater than velocities measured normal to the bedding.

Shipboard determinations of alkalinity, pH, and salinity are shown in Figure 5. Alkalinity values are quite high (3.91-6.16 meq/kg), except for the determination from deepest in the hole (2.1 meq/kg in Core 17). Salinity and pH show little variation, except for a low value of salinity corresponding to the low alkalinity in Core 17.

## BIOSTRATIGRAPHIC SUMMARY

A largely detrital section of undetermined age at the base to Recent at the top was recovered at Site 268. The lowest dated core is Core 17 for which nannofossils and foraminifera indicate a probable late Oligocene age. The section is, in general, sparsely fossiliferous and preservation of calcareous as well as siliceous fossils is poor throughout. Foraminifera are present only in Cores 1, 8, 12, and 17 and as previous Antarctic sites, diversity is limited to one or two species in any assemblage. Nannofossils occur only in Cores 8-11 and 17-20 and are,



Figure 4. Shipboard measurements of sonic velocity and wet bulk density at Site 268. Calculated acoustic impedance is shown.

likewise, of very limited diversity and poor preservation. Datable radiolarian assemblages occur in Cores 1, 4, and 5; in the remainder of the section radiolarians are absent or so poorly preserved that age determinations are impossible. Diatoms are scarce and generally occur as fragments; they are absent below Core 12. Unequivocal age determinations were possible only for Cores 1, 2, and 5. Poor assemblages of Miocene diatoms occur in Cores 4 through 12.

Based on combined data from the various fossil groups the probable distribution of ages for cores at Site 268 is as follows. Cores 1, 2: late Pleistocene to Recent; Core 3: early Pleistocene(?); Core 4: Pliocene; Core 5: late Miocene or early Pliocene; Cores 6, 7: late-mid Miocene?; Cores 8-16: early Miocene to late Oligocene; Core 17: late Oligocene; Cores 18-20: unknown.

#### Foraminifera

Except for occurrences in four cores, the section at Site 268 is barren of foraminifera. Core 1 contains a fauna consisting entirely of left-coiling *Globorotalia pachyderma*, and belongs in the *G. pachyderma* assemblage zone of late Miocene to Recent age (see report for Site 266). The core catcher of this core contains frequent specimens of this species along with abundant ice-rafted sand- to pebble-sized material. A sample from Section 6, 82 cm is virtually a *G. pachyderma* ooze, containing very little other material.

The other occurrences of foraminifera at Site 268 are in Cores 8, 12, and 17; these belong in the *Catapsydrax unicavus* assemblage zone of late Oligocene/early Miocene age. In Core 8 *Catapsydrax dissimilis* is also present.



Figure 5. Shipboard measurements of pH, alkalinity, and salinity in sediment pore waters at Site 268.

As at previous Antarctic sites, the lack of diversity of the foraminiferal assemblages indicates deposition in a high-latitude environment. Moreover, despite its relatively shallow depth, Site 268 has been below the carbonate compensation depth for practically all of late Oligocene/early Miocene to Recent time.

### Nannofossils

A 474-meter section of siliceous oozes and cherts (Cores 1 to 20) was sampled at Site 268. Within the siliceous oozes there were a few calcareous horizons which contained nannofossil assemblages (Sample 10-1, 10 cm; Core 13-1; Core 17-1-top; Samples 17-2, 129 cm to 17-3, 9 cm). These assemblages were generally poorly preserved. Occasional good nannofossil specimens allowed species identification to be made, but single shields and separated centers of placoliths were common. Species present include: Coccolithus miopelagicus, small Reticulofenestra sp., Discoaster deflandrei, Cyclicargolithus floridanus, Coccolithus eopelagicus, Reticulofenestra umbilica, Dictyococcites abisecta, D. bisecta, Chiastomolithus altus, Dictyococcites scrippsae, Reticulofenestra cf. gartneri, Sphenolithus moriformis, Coccolithus pelagicus, Reticulofenestra hillae.

The low abundance of nannofossils in the assemblages and their generally poor preservation made it impossible to give accurate age designations for the calcareous horizons. However, some biostratigraphic approximations can be given as follows. Mid to early Miocene: Core 10-1-top to Sample 10-1, 10 cm; early Miocene: Core 13-1; mid Oligocene: Core 17-1-top; mid Oligocene: Sample 17-2, 130 cm and Sample 17-2, CC; lower Oligocene: Core 17-3, 9 cm.

#### Radiolaria

Radiolaria are rare and moderately to poorly preserved in most of the Neogene sediments (Cores 1 to 12), and trace to absent in older sediments.

The only radiolarian zone recognized is the *Helotholus vema* Zone (Core 3, Section 2 to Core 7). The overall assemblages in Cores 8 to 13 indicate Miocene age, and further subdivision of this interval into zones is impossible. From Core 14 downward, sediments are barren of Radiolaria except at Core 17, Section 1 where only *Stylosphaera coronata coronata* is observed.

#### Diatoms

Diatoms are typically in very poor abundance and condition. When diatom remains are found they are fragmented. Very few entire frustules were found from Cores 1 through 12. Below Core 12 diatoms are absent.

Core 1 through Sample 2, CC contain the *Coscinodiscus lentigenosus* Zone along with a few reworked Oligocene diatoms. Only at Sample 2-1, 60 cm were there enough diatoms to give assurance to this assignment. In Cores 3 and 4 there were too few identifiable diatoms to make an assignment. Cores 5, 6, and 7 contain a portion of the *Denticula hustedtii* Zone. Cores 8 through 12 and below show such paucity of diatoms that they cannot be zoned.

## Palynology

Eight samples, taken from all three lithological units, were macerated to recover palynomorphs from this site. Sampling was concentrated in the basal part of Unit 3, in the hope of establishing the age of the oldest part of the sequence penetrated. However, none of the samples yielded diagnostic fossils.

Core 20 yielded abundant acritarchs, chiefly leiospheres, which are not diagnostic as to age. A few pollen grains referable to the Late Cretaceous-Tertiary genus *Nothofagidites* are present, and recycled Permian pollens were notably common. *Nothofagidites* spp., and *Proteacidites* spp., another Cretaceous-Tertiary from genus, occur in Cores 16-19. Cores from the upper units, i.e., Cores 7, 8, and 9 were barren of identifiable palynomorphs, probably due to extreme dilution with terrigenous material.

## SUMMARY AND CONCLUSIONS

Site 268 is the southernmost comprising the transect between the ridge crest and the Antarctic continent and along 105°-110°E longitude. The site is located on the lower continental rise off the Knox Coast in a water depth of 3544 meters. Sediments here are probably more than 1.5 km thick, but only 474.5 meters were penetrated during drilling. Sediments throughout the hole are chiefly silty clays, diatom-rich near the top and nannofossil-rich at lower levels, while chert is common in the lower 200 meters of the drilled section. Ice-rafted detritus is common down to the early Miocene portion of the section and isolated granules are found in the Oligocene sediments.

A few strata identifiable as turbidites occur in the form of thin sand units and silt laminae near the top of the sequence but below Core 7 silt laminae more closely resemble deposits from contour currents (Piper and Brisco, this volume). Thus, turbidity currents apparently did not reach this site before the Pliocene and then not in substantial quantities; contour currents apparently deposited a considerable amount of the silty sediment before the Pliocene. As in the pelagic sites on Leg 28, nannofossil sediments underlie diatomaceous ones at Site 268, but these possibly represent the upper part of a thick transition zone since nannofossils are present in only sparse amounts in the cherty sequence at the bottom of the hole (below Core 11). The highest level at which calcareous nannofossils are abundant is in the early Miocene (Core 8), within the time span representing the similar lithologic change at Site 267 (late Oligocene-early Miocene).

A prominent unconformity occurs between Pliocene (Gilbert) sediments in Core 7 and early Miocene ones in Core 8; throughout this interval the sedimentation rate averages to  $\sim 2$  m/m.y. Diatomaceous clays above this break accumulated at  $\sim 31$  m/m.y. and nannofossilbearing clays below it at  $\sim 41$  m/m.y. Extrapolation of the sedimentation rate observed between Cores 14 and 17 yields an age of early Oligocene for the oldest sediments penetrated at the site.

Magnetic lineation studies suggest that the underlying crust at Site 268 is about 55 m.y. old. If correct, this implies a very high average sedimentation rate of 60-70 m/m.y. for the lower 1000 meters of unsampled sedi-

ments at this site. This area has presumably undergone major vertical tectonism associated with the early rifting of Antarctica and accompanied by large-scale terrigenous sedimentation.



Figure 6. Age vs. depth at Site 268.

BIOSTRA	ATIGRAPHY		105	(m)	HOLE	COLUMN	LITHOLOGIC	ACOUST. VEL.(kms-1)     BULK GRAPE     DENSITY D SYRINGE
		DIATOMS	AGE	DEPTH	268	COLUMN	DESCRIPTION	POROSITYA 1.4 1.6 1.8 2.0
	?	1	PLEISTOCENE	0	1			60 70 80 90
		?		50	3	125.25.25	CLAY, SILTY CLAY, SAND and DIATOM OOZE.	- <del>8</del> 0
	5		PLIOCENE	100	4 <b>–</b>			•~
		8				•• ••••		+
	terval.	tic.		150	- 7 <u>-</u>	مرد می بدر می	? ? ?	
-	ene aqe in this in	re and nondiagnosi		200	- 9_		CLAY, SILTY CLAY and CLAY NANNO OOZE.	- and the second
	ate general Mioce	ms present but ra	MIOCENE	250	10_		À	
	ssemblages indic	Diato		200	11	<b>AA</b>	SILTY CLAY, laminated SILTY CLAY and CLAYEY SILT, CLAYEY SILT, and	
		ARREN		300	<sup>12</sup> <u>–</u>		CHERT.	co o
	BARREN	8			14			

Figure 7. Graphic hole summary, Site 268.

BIOSTRA	TIGRAPHY		ACE	(m) H	HOLE	COLUMN	LITHOLOGIC	ACOUST.     BULK     DENSITY	VEL.(kms GRAPE SYRING	; -1) E
			AUL	DEPTI	268	COLONIA	DESCRIPTION	POROSITYA	1,8 2,0	
	BARREN			350	15 - 16	<b>.</b>		60 70 Φ	80 90 o	0
	Stylosphaero coronata coronata			400	17	<b>.</b>			ළ ლ <sub>დ</sub> დ	
	ten	BARREN			18			, ic	000	Turb
	BARR			450	- 20	A A		- 0 0		
								00	00	

Figure 7. (Continued).

		and the second se									
30V 30V 30V 30V 30V 30V 30V 30V	IC DESCRIPTION	AGE	ZONE	FOS CHAR TISSOJ	ACTER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION
R R G 1 0.5 CG G G G G G G G G G G G G G G G G G G	<ul> <li>SY 6/4) DIATOM 00ZE; rare erratic ially in lower part.</li> <li>Sec. 1 (76 cm): 97 diatoms 28 silt TR radiolarians, calcareous nannos</li> <li>(10YR 4/2) CLAYEY SILT.</li> <li>(3/4) SILTY CLAY, 2 coarse-sand-bearing</li> <li>b) washed out fine sand.</li> <li>(10YR 4/2), SILT RICH, DIATOM 00ZE; b) washed out fine sand.</li> <li>(5/2), SILT RICH, DIATOM 00ZE; b) sand granules.</li> <li>Sec. 5 (135 cm): 80% diatoms 15 silt 38 clay TR radiolarians, sponge spicules</li> <li>SILTY CLAY passing down to yellowish LAYEY SILT.</li> <li>D passing down into very fine NNO.</li> <li>clayey silt with many pebbles and % of sediment), and pieces of clayey highly disturbed alternation of beds, f clayey silt may be clasts.</li> </ul>	LATE PLEISTOCENE	PLEISTOCENE BRUNNES	D	A G	1 2 3 4 5 6 cat	0.5 1.0 1.0 1.1 1.1 1.1 1.1 1.1 1.1 1.1 1.1	V01DS		*91 CCC *108 *25 *108 *25 *108 *25 *108	Core soft throughout. Alternating beds of brownish gray (5YR 4/1) SiLTY CLAY. Scattered erratic pebbles, granules and sard. commoner in diatom-bearing Silth JCLAY. Scattered erratic pebbles, granules and sard. (Sec. 1, 143 cm): 68% clay 30% silt TR diatoms, radiolarians, sand two graded fine sand. Coarse silt beds Alternating thin beds of 5YR 4/1 silty clay and 5Y 4/1 diatom-bearing silty clay. Alternating thin beds of 5YR 4/1 silty clay and 5Y 4/1 diatom-bearing silty clay. Alternating thin beds of 5YR 4/1 silty clay and 5Y 4/1 diatom-bearing silty clay. <u>Bulk X-ray (33,1 m):</u> <u>Amorph. 42,1 %</u> Ident. 52,93 Chio. 2,13 <u>Mica</u> 21,33 <u>Mica</u> 21,33 <u>Mica</u> 21,33 <u>Mica</u> 21,33 <u>Mica</u> 21,33 <u>Mica</u> 21,33 <u>Mica</u> 22,33 <u>Chio. 2,178</u> <u>Mica</u> 22,33 <u>Chio. 2,178</u> <u>Mica</u> 22,33 <u>Chio. 2,178</u> <u>Mica</u> 22,33 <u>Mica</u> 22,01



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**SITE 268** 



te	268	Ho1	e		Co	ore 9	Cored In	terv	/al:	199-208.5 m
		F CH/	OSS	TER	2			ION	PLE	
MOL	ZONE	FOSSIL	ABUND.	PRES.	SECTIO	METERS	LITHOLOGY	DEFORMAT	LITH0.SAM	LITHOLOGIC DESCRIPTION
					1	0.5	VOID	•	*98 CC *145	Dark greenish gray (56Y 4/1) and olive gray (5Y 4/1) CLAY; rare granules and coarse sand. Color 5Y 4/1 Sec. 1 (98 cm); 90% clay 5% silt 3% diatoms TR sand
					2	and a data		1 1 1 1	*74 60 92* *93	Color 5GY 4/1 5YR 3/2 silty clay Color 5GY 4/1
TIVULIT					3	nutrutur	VOID		*64 *82	Grayish brown (5YR 3/2) SILTY CLAY; some medium silt thin laminae. Sec. 3 (82 cm): 40% clay 50% silt 5% diatoms
					4	mportur	VOID			2% sponge spicules TR radiolarians
					5	uluuluu	VOID	1		
		NFDR	p.g	PP	Ca	Core tcher		2	•	

SITE 268

Site	268	Ho1	e		Core	10	Cored In	ter	val: 22	27.5-237 m	Site	268	Н	ole		C	ore 11	Ce	ored In	terv	al: 2	256-265.5 m
AGE	ZONE	FOSSIL 2	VICE ABUND.	PRES. 3	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	Encen D	FOS HARA UNITED	SIL ACTER	SECTION	NETERS	LITH	HOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION
		N	F	P	2	1.0			*104 *141 GZ *84 *129 *137	01ive gray (5Y 5/1) CLAY NANNOFOSSIL DOZE. 5Y 3/1 CLAYEY SILT 10YR 4/2 SILTY CLAY 56 4/1 <u>Bulk X-ray (229.4 m)</u> : Amorph 48.3% Ident 51.7% 04.4% K-Fe 16.7% Plag 15.1% Mica - 22.4% Chlo 2.6% Hema 1.7% Amph 1.0% Dark yellowish brown (10YR 4/2) LAMINATED SILTY CLAY and CLAYEY SILT. Smear slide silty lamina, Sec. 2 (84 cm):	MIOCENE			N F F F	p p	1	0.5 1.0		701D		130	Semilithified moderate brown (SYR 3/4) SILTY CLAY. CHERT, grayish red (10R 4/2) black with granule clast dark greenish gray (56 4/1) blocks. Sec. 1 (130 cm): 65% silt 30% clay 2% doatoms 2% spong spicules TR radiolarians, silico- flagellates SYR 4/1 CLAYEY SILT 10Y 4/2 SILTY CLAY CHERT
MIDCENE				-	3	true true true				65% silt 15% clay 12% diatoms 5% sponge spicules Smear slide silty clay, Sec. 2 (129 cm): 55% clay 40% silt	Site	268 300Z	H	FOS CHAR/	SIL ACTE	SECTION O	WELEKS	LIT	Cored In HOLOGY	DEFORMATION AT	LITHO. SAMPLE	284.5-294 m LITHOLOGIC DESCRIPTION
					5	and			€]04	5% diatoms 1% sponge spicules TR radiolarians, silico- flagellates Core stiff to semilithified throughout; trace of very coarse sand grains.	MIOCENE	Catapsydrax unicavus	and the second se	NFDR	RRPR	1	0.5- 1.0- Core atcher				GZ *90	<pre>10R 4/2 CHERT Dark yellowish brown (10YR 4/2) laminated SILTY CLAY and CLAYEY SILT. Core semilithified or lithified (chert). 10YR 4/2 diatom bearing SILTY CLAY 10Y 4/2 diatom bearing CLAY 5 3/2 CHEFT</pre>
		NF	11	-	6	re			*17 *58 CC	Olive gray (5Y 4/1) laminated, diatom-bearing SILTY CLAY Brownish gray (5YR 4/1) SILTY CLAY Grayish red (10R 4/2) CLAY Brown gray (5YR 4/1) SILTY CLAY Dark yellowish brown (10YR 4/2) LAMINATED SILTY CLAY and CLAYEY SILT.	Exp	lanato	ry n	notes	: in	Char	ter 1	-				DT 3/2 UNERI
		D	-	- 0	Cato	her		1														

Site	268	Hole		Cor	re 13	Cored	Inte	rval	1:3	03.5-313 m	Sit	e 268	Н	ole		Cor	re 15	Cored In	terv	al: 3	341.5-351 m
AGE	ZONE	FOS CHAR TISSOJ	ACTER	SECTION	METERS	LITHOLOG	DEFORMATION	UEFUKMALIUN	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	EDecti D	FOSS HARA	CTER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION
LATE OLIGOCENE/EARLY MIDCENE	Catapsydrax unicavus	NFDR	C P P	1 Cat	0.5 1.0	VOID		•11	06	Core semilithified on lithified (chert). Core spilled on deck; following lithologies present (in order of abundance): 10YR 4/2 SILTY CLAY 10Y 4/2 CHERT 10R 4/2 CHERT 10Y 4/2 Iaminated CHERT 10Y 4/2 Iaminated CHERT 10Y 4/2 Iaminated CHERT 10Y 4/1 Iaminated CHERT 10Y 4/2 Iaminated SILTY CLAY and CLAYEY SILT NANNO 002E Dark greenish gray (5GY 4/1) and dark yellowish brown	LATE OLIGOCENE/EARLY MIOCENE			N FDR		1 Cat	0.5 1.0	VOID		100 136	10Y 4/2 and 10R 4/2 CHERTS 10Y 4/2 SILTY CLAY with SILT laminae Core semilithified or lithified (chert). Smear slide of silty clay, Sec. 1 (136 cm): 60% clay 40% silt
										(10YR 4/2) SILTY CLAY.	Sit	te 268	В	lole		Co	re 16	Cored In	terv	al:	351-360.5 m
									11/14	Bulk X-ray (304.3 m):         Bulk X-ray (304.3 m):         Bulk X-ray (304.5 m):           Amorph.         -48.97         Amorph.         -50.27         Amorph.         -22.87           Ident.         -51.1%         Ident.         -49.85         Ident.         -77.22           Quar.         -21.1%         Calc.         -6.1%         Quar.         -48.6%           Cris.         -48.7%         Quar.         -26.6%         Cris.         -36.2%           K-Fe.         -6.8%         K-Fe.         -10.1%         K-Fe.         -15.6%           Plag.         -9.6%         Plag.         -77.3%         Plag.         -13.6%	AGE	ZONE		FOS CHARA TISSOJ	SIL	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC DESCRIPTION
Site	268	Hole		Con	re 14	Cored	Inte	rval	1: 3	Chlo 1.6% Chlo 5.0% Chlo 1.8% Hema 1.6% Mont 4.2% Amph 1.2%	CARLY MIDCENE					1	0.5			8 8	Alternating 10R 4/2 and 10Y 4/2 CHERT and 5Y 4/2 SILTY CLAY with numerous SILT laminae. Core semilithfied or lithified (chert). Smear slide of silty clay.
AGE	ZDINE	FOS CHAR	ACTER	CTION	ETERS	LITHOLOG	Y	NULLINN .	0. SAMPLE	LITHOLOGIC DESCRIPTION	LIGOCENE/						1.0	VOID	_	_	Sec. 1 (8 cm): 55% clay 45% silt
IE /		FOSS	PRES	SE	Ψ.		ou ou	UEPU	LITHO		LATE 0			N F DR		C Cat	ore tcher	<b>^^^^</b>			Sec. 1 (8 cm): 11 tamina, Sec. 1 (8 cm): 90% silt 10% clay
OCEN					-					Lore semillunitied or liunitied (chert).	Si	te 268	ł	lole		Co	ore 17	Cored In	iter	al:	379.5~389 m
SOCENE/EARLY MI				1	1.0	VOID	•	11	37	Grayish olive (10Y 4/2) laminated CHERT.	AGE	ZONE		FOS CHAR/ TISSOJ	ACTE	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION
LATE OLIG		N F D R		Cat	ccher				99	Grayish olive (10Y 4/2) laminated SILTY CLAY and CLAYEY SILT. Sec. 1 (144 cm): 70% silt 25% clay 25% clay 25% heavy minerals TR volcanic glass, sponge spicules 10Y 4/2 laminated CLAYEY SILT and SILTY CLAY. 10Y 4/2 laminated CHERT.	LIDDED OL TRANETIE	Chiastomolithus altus	Catapsydrax unicavus	N F R D	FR	1	0.5-	VOID		CC XM 1117 118	Core semilithified or lithified (chert). Mostly grayish olive (10Y 4/2) SILTY CLAY with many SILT laminae. Interbedded 5GY 4/1 CHERT as shown in lithology column. Smear slide of silty clay, Sec. 2 (74 cm); GGX clay 35% silt 1% diatoms no silt laminae no silt laminae NOR 4/2 and 5GY 4/1 chert Only a few broad diffuse silt laminae. 10R 4/2 silty clay, no silt laminae Bulk X-ray (380.9 m); Amorph 31.65 Amorph 41.25 Ident 66.45 Duar 50.55 Duar 45.85
											_										K-Fe.         - 15.4%         K-Fe.         - 14.9%           Plag.         - 12.4%         Plag.         - 10.8%           Mica         - 18.6%         Mica         - 17.4%           Chlo.         - 2.0%         Chlo.         - 2.4%           Amph.         - 1.1%         Pyri.         - 8.7%

No core catcher sample.

SITE 268

Site 268	Hold	2	Co	re 18	Cored	Inter	val: 4	408-417.5 m	Sit	e 268	Ho1	e		Core	19	Cored In	terv	val: 436.5-446 m
AGE ZONE	FOSSIL P	ACTER ONUBA	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO, SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FOSSIL F	OSSIL RACTI	PRES. Ja	SECTION	METERS	LITHOLOGY	DEFORMATION	UITHOLOGIC DESCRIPTION
٤			1	0.5			*54 1106	Core semilithified or lithified (chert). 5YR 4/1 silty clay, no silt laminae Dark greenish gray (5GY 4/1) SILTY CLAY; in most places, with many SILT laminae; 5GY 4/1 CHERT as shown in lithology column. Smear slide of silty clay, Sec. 1 (106 cm): 70% clay 30% silt	6		NFDR	R	p	0. 1 1. 2 Corre Catch	5-1-1-1-1-1-1-1-1-1-1-1			5Y 2/1 CHERT Core semilithified or lithified (chert). 5Y 2/1 CHERT Smear slide of silty clay, Sec. 1 (134 cm): 60% clay 35% silt 34 135 Dark greenish gray (5GY 4/1) SILTY CLAY; a few silt iaminae in Sec. 2. Chert in Sec. 2 (80 cm) has bed of sand and granules. 5GY 4/1 CHERT 5GY 4/1 CHERT
	NFDR		Car	ore		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1		SGY 2/1 CHERT About 15 horizons in core of diagenetic (?MnO <sub>2</sub> ) blackening of silt laminae, commonest in lower part of core.	Sit	a 268 SONE	FOSSIL R. 10H	e OSSIL RACTI	PRES. 33	SECTION	METERS 02	Cored In	DEFORMATION	LITHOLOGIC DESCRIPTION
								Grayish olive (10Y 4/2) silty clay.	2					1		VOID		Core semilithified or lithified (chert). Dark greenish gray (56Y 4/1) SILTY CLAY; rare silt laminae; some 56Y 4/1 CHERT, as shown in lithology column. Smear slide of silty clay, Sec. 1 (133 cm): 90% clay 10% silt

NFDR

2

Core Catcher

111

VOID

XM 86 93

















