37. CHEMICAL ANALYSES AND MINOR ELEMENT COMPOSITION OF LEG 29 BASALTS

A. T. Ovenshine, Gary R. Winkler, U. S. Geological Survey, Menlo Park, California Peter B. Andrews, New Zealand Geological Survey, Christchurch, New Zealand and

Victor A. Gostin, University of Adelaide, Adelaide, South Australia

ABSTRACT

Drilling on Leg 29 recovered basalts at five sites. Three of the basalts were from pillow flows, one from a massive lava flow, and one from an intrusion. No major differences between the rocks from the five sites are found in major oxide, normative mineralogical, and minor element composition. However, there are minor differences of unknown significance between the sites close to New Zealand and the sites close to Australia. Chemically, the Leg 29 basalts are intermediate in composition between oceanic tholeiites and oceanic alkalic basalts.

INTRODUCTION

Basalt basement was reached and penetrated at Sites 278, 279 (Hole 279A), 280 (Hole 280A), 282, and 283. The purpose of this paper is to provide abbreviated descriptions of the basalts based mainly on shipboard examination of cores and thin sections and to present major oxide chemical analyses, minor element compositions, and CIPW norms.

LITHOLOGY OF LEG 29 BASALTS

Tables 1, 2, and 3 contain analytical data and normative compositions of basalts recovered from the five sites. In the following general descriptions of the basalt cores, the focus is upon the major megascopic properties of the rocks as they would be viewed by field geologists. Detailed studies by specialists are included elsewhere in this volume.

Site 278

Drilling at Site 278 penetrated the upper portion of a pillow flow, although the precise upper contact was not recovered. Cores recovered from several complete or nearly complete pillows suggest that they are up to 20 cm thick with black glass rinds as much as 3 cm thick. The outer surface of the rind is frequently composed of orange palagonite. Inferred interpillow material is principally micritic limestone (Figure 1) containing fragments of glass, basalt, and palagonite. Tests of partially recrystallized foraminifera can be recognized in thin sections of the interpillow limestone which was probably developed from a foraminiferal-nannofossil ooze.

In hand specimen the fresh rock is olive-black (5Y2/1) porphyritic basalt, with phenocrysts of plagioclase comprising from 5% to 15% of the volume. Rare, white calcite-filled amygdules averaging 0.5 mm in diameter occur preferentially in, and just beneath, the glassy pillow rims. In thin section the basalts consist of labradorite, altered acicular pyroxene, iddingsite(?) developed after olivine, and abundant veinlets of sparry calcite, chalcedony, and micritic limestone. Unaltered olivine euhedra occur in the fresh glass of the pillow rims.

Site 279 (Hole 279A)

Basement recovery at Hole 279A consisted of three core sections of fine-grained dark gray basalts interpreted as the upper few meters of a massive flow. The basalts of Core 12, Section 1 are vesicular to amygdaloidal, subaphanitic to very fine-grained plagioclaseporphyritic basalts, that grade to nonvesicular and nonporphyritic fine-grained basalts toward the base of the core. The basalts of Core 13, Section 1, and Core 13, Section 2 are in lithic continuity with the vesicular basalts of Core 12, Section 1, but differ from them in their slightly coarser grain size and lack of vesicles and phenocrysts. The nonvesicular basalts are subfriable, and during drilling many of the short core segments were abraded to barrel shapes.

In the upper 30 cm of Core 12, Section 1, vesicles and amygdules comprise up to 40% of the rock volume. The vesicles range from 0.5 to 4 mm in diameter, averaging 1.5 mm. Approximately 50% of the vesicles are unfilled. The remaining 50% are partly to completely filled with white calcite, chlorite, or a gray-blue zeolite(?). Pyrite subhedra occur sparingly in some of the vesicles.

In hand specimen the basalts of Core 12, Section 1 have a color index of 70, and consist of 1% white plagioclase phenocrysts ranging to 4 mm in length, in a subaphanatic to very fine-grained groundmass that consists of equal parts of feldspar laths and pyroxene plus glass(?). In thin section the basalts of Core 13, Section 2 have an intersertal texture and consist of 0.2-0.4 mm labradorite laths (35%), anhedral pyroxene (30%), patches of partially altered crystallite-rich brown glass

	Lab. No.									
	M119 549W	M119 950W	M119 951W	M119 952W	M119 -953W	M119 954Wa	M119 955W	M119 956W	M119 957W	
	Sample 278- 35-3	Sample 278- 35, CC	Sample 279A- 12-1	Sample 279A- 13-2	Sample 280A- 23-4	Sample 282- 18-2	Sample 282- 18-3 ^a	Sample 283- 18-1	Sample 283- 19-1	Average
SiO2	48.1	48.3	48.1	48.9	46.7	46.6	35.6	46.3	45.7	47.3
A1203	16.8	17.2	15.1	15.0	17.3	17.8	13.5	16.1	15.8	16.4
Fe ₂ O ₃	3.8	3.9	2.7	3.3	3.2	5.6	5.1	7.5	5.6	4.4
FeO	3.8	3.6	5.9	6.3	6.2	2.7	2.8	3.7	3.4	4.4
MgO	6.4	6.4	6.0	7.0	6.3	2.7	3.8	6.8	7.8	6.2
CaO	14.1	14.4	13.2	10.6	10.6	12.1	19.9	6.3	5.7	10.9
Na ₂ O	2.0	2.2	2.6	2.3	2.6	3.1	3.2	3.5	3.5	2.7
K ₂ 0	0.29	0.25	0.22	0.20	0.07	0.63	0.63	0.68	0.52	0.36
$\tilde{H_2O+}$	1.3	1.2	0.60	1.5	1.7	1.3	1.6	2.8	3.1	1.7
H20-	1.7	1.6	1.2	1.8	1.9	2.0	2.4	5.1	5.9	2.7
TiO ₂	0.91	0.88	1.3	1.4	0.91	1.6	1.2	1.8	1.8	1.32
P205	0.10	0.10	0.19	0.18	0.07	0.24	0.19	0.22	0.26	0.17
MnO	0.11	0.11	0.16	0.12	0.17	0.17	0.23	0.18	0.14	0.14
^{CO} 2	0.59	0.52	1.6	0.03	1.6	2.5	11.2	0.08	0.09	0.88
Total	100	101	99	99	99	99	101	101	99	99.57

TABLE 1 Major Oxide Analyses of Leg 29 Basalts

Note: Analytical methods described in USGS Prof. Paper 575-B, p. 187-191. Analyst: Hirbert Kirshenbaum. ^aNot used in computing average.

	Sample								
	278-35-3	278-35, CC	279A-12-1	279A-13-2	280A-23-4	282-18-2	282-18-3	283-18-1	283-19-1
Quartz	2.8	1.6	2.2	3.8	1.8	6.0		1.3	0.6
Orthoclase	1.7	1.5	1.3	1.2	0.4	3.8	3.8	4.2	3.3
Albite	17.2	18.8	22.5	20.1	22.6	27.0	27.4	30.9	31.7
Anorthite	36.6	36.7	29.6	31.0	36.2	33.8	20.8	27.3	27.7
Wollastonite	12.6	13.1	10.8	9.1	2.9	4.2	2.6	1.3	0.1
Enstatite	16.2	16.1	15.3	18.0	16.1	6.9	5.2	17.6	20.8
Ferrosilite	2.6	2.2	6.9	7.0	7.8				
Forsterite							3.1		
Magnetite	5.6	5.7	4.0	4.9	4.8	4.8	6.4	7.6	6.6
Hematite						2.5	0.8	2.6	1.4
Ilmenite	1.8	1.7	2.5	2.7	1.8	3.1	2.3	3.6	3.7
Apatite	0.2	0.2	0.5	0.4	0.2	0.6	0.5	0.5	0.7
Calcite	1.4	1.2	3.7	0.1	3.7	5.9	25.7	0.2	0.2
Total	98.7	98.8	99.3	98.3	98.3	98.6	98.6	97.1	96.8
Salic	58.4	58.6	55.6	56.1	61.1	70.6	52.0	63.6	63.3
Femic	40.3	40.3	43.7	42.2	37.2	28.0	46.6	33.5	33.5

TABLE 2 CIPW Weight Norms of Leg 29 Basalts

	Basa
	29
	Leg 29 Ba
	Elements,
	Minor
3	for
TABLE	Analyses
	e Spectrographic
	Quantitativ

alts

	M119549W	M119550W	M119551W	M119552W	M119553W	M119554W	M119555W	W119556W	M119557W	
	Sample 278-35-3, 9.2 cm	Sample 278-35, CC	Sample 279A-12-1, 10 cm	Sample 279A-13-2, 13 cm	Sample 280A-234, 143-150 cm	Sample 282-18-2, 21 cm	Sample 282-18-3, 16 cm	Sample 283-18-1, 121-134 cm	Sample 283-19-1, 6 cm	Average
Ті %	0.44	0.48	0.95	0.97	0.64	1.25	1.14	1.56	11.1	1.07
(mqq) nM	920	1000	1450	1000	1530	1600	2000	1560	1340	1378
Ba	N4	N4	51	55	20	83	58	37	47	40
Co	31	38	40	43	53	36	44	32	49	ŧ
S	400	400	200	150	400	370	280	255	260	302
Cu	91	86	90	100	147	107	52	77	115	96
iz	140	130	86	82	200	120	100	69	87	113
Sc	26	27	43	45	40	48	43	55	49	42
Sr	120	130	160	150	77	170	122	133	147	134
~	200	230	270	264	280	390	265	390	350	293
Y	14	13	22	20	26	40	38	32	38	27
Zr	56	54	85	82	57	130	98	150	110	16
Ga	12	11	13	10	16	20	15	21	19	15
Yb	1	1	2	2	4	\$	4	ŝ	4	ę



Figure 1. Upper part of a basalt pillow from Site 278. Interpillow light gray micritic limestone developed by recrystallization of foraminiferalnannofossil ooze adheres to the black chilled glass pillow rim (XI).

(25%), and magnetite (10%). Accessory minerals are calcite, chlorite, analcite(?), and an unidentified zeolite. Basalts from the vesicular part of Core 12, Section 1 are similar, but the intercrystalline glass has fewer crystallites and is less altered.

Site 280 (Hole 280A)

poulos. N = not detected at value shown.

Drilling at Hole 280A recovered approximately 5 meters of core consisting of basalts intercalated with claystone and siltstone. The basalt layers are interpreted as apophyses of a larger intrusion, because of silicification, pyritization, and lithification of the intercalated sediment, and because of the extensive brecciation and minor mineralization in the sediments overlying the basalts. The apophyses are 2.0, 0.5, and 1.1 meters thick.

The rock is greenish-black, holocrystalline, subophitic, fine-grained, pyroxene-labradorite basalts containing magnetite in amounts ranging from 2% to 10% by volume. Recognizable olivine phenocrysts are confined to the lowermost apophysis. The rock is extensively altered to serpentine group minerals, and locally contains minor amounts of pyrite. The most altered portions of the core expand and disintegrate on wetting.

Site 282

Drilling penetration of basement at Site 282 resulted in recovery of 7.6 meters of core interpreted as altered pillow basalt, or possibly altered broken pillow breccia and aquagene tuff. Veins of calcite, chlorite, limonite, and steatite(?) up to several centimeters in width, pervade the highly fractured core. Radially oriented calcite fractures are common in recognizable pillow segments, but are generally confined to the greenish-black, partially devitrified glass rims. In some instances, calcite, accompanied by minor pyrite, appears to have completely replaced basalts (Sample 29-282-18-2, 53-60 cm).

In hand specimen the basalts are greenish-gray (5 GY 7/2), ranging to bluish-gray and light bluish-gray. Locally, plagioclase and pyroxene crystals can be recognized under the hand lens, but generally individual crystals are indistinguishable. Thin sections of the rare, unaltered patches of basalts show subophitic intergrowths of zoned plagioclase and magnetite-rich pyroxene in a devitrified groundmass. Most of the rock is pervasively altered to calcite, brown iron oxide, and serpentine group minerals.

Minute grains of native copper were recognized on slabbed core surfaces, but quantitative spectrographic analysis for copper (Table 3) did not show unusually high values in comparison to the other Leg 29 basalts.

Site 283

Drilling at Site 283 encountered basalts that were originally either pillow lava or broken pillow breccia. This interpretation is not straightforward, however, because a color change in the overlying sediments could indicate weak metamorphism by an intrusion. The interpretation as a pillow flow or broken pillow breccia is based on: (1) abundant devitrified glass in the upper part of the core, (2) occurrences of breccia containing both glass and basalts, (3) amygdules occurring throughout the core, and (4) one core segment (283-19-1, 6 cm) containing the upper half of a pillow with a glass patina and chilled rim, grading downwards to holocrystalline basalts.

The upper 20 cm of the core consists of light greenishgray devitrified volcanic glass with abundant veinlets of calcite and serpentine and one tabular mass of dark red palagonite that is probably altered interpillow glass. The rock consists of rare altered feldspar laths (2%) in an altered groundmass consisting of clay minerals, zeolite, and serpentine.

The remaining 1.3 meters of core consists of dark greenish-gray amygdaloidal basalts (80%), calcite veins (15%), and dark fragments of serpentinized glass (5%). In several places the rock exhibits autoclastic, or tectonic breccia textures. In thin section the original texture

of the fine-grained rock, which may have been intersertal and subophitic, is largely obscured by the radial habit of the alteration products. Altered plagioclase laths comprise 20%, black opaque minerals (10%), partially altered olivine and pyroxene (30%), serpentine(?) after glass (10%), calcite (5%), with the remainder consisting of clay minerals, serpentine, and zeolite.

CLASSIFICATION

Comparison of major oxide, normative mineralogical, and minor element compositions (Tables 1, 2, and 3) does not indicate major differences between the basalts from the five sites. Minor differences of unknown significance, however, do exist between the basalts from the sites near New Zealand (278 and 279) and the basalts from sites near Australia (280, 282, and 283). The "New Zealand" basalts show, in major oxide composition, lower Na₂O and K₂O values than the "Australian" basalts (Table 1). Similarly, contents of Zr, Ga, and Yb are lower in the New Zealand basalts than in the Australian ones.

The problems in classifying the Leg 29 basalts occur because their composition tends to be close to various proposed chemical boundaries. For example, on the AFM diagram commonly used to separate subalkaline from alkaline and calcalkaline basalts (Figure 2), the Leg 29 basalts straddle the boundary of Irvine and Baragar (1971, p. 528). Somewhat improved discrimination is offered by the plot of weight percent alkalies against silica (Figure 3). This plot has also been used to discriminate alkaline and tholeiitic rocks of Hawaii (MacDonald and Katsura, 1964, p. 87) and Japan, Korea, and Manchuria (Kuno, 1968, p. 627). Considered as a group, the Leg 29 basalts straddle the highalumina versus alkali-olivine field boundary. This was also true for many samples from DSDP Legs 15 and 17 (Bence et al., 1973, p. 997 and 1000).

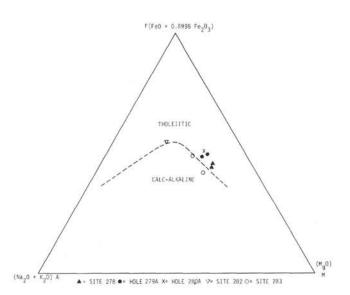


Figure 2. AFM diagram, Leg 29 basalts (after Irvine and Baragar, 1971, p. 528).

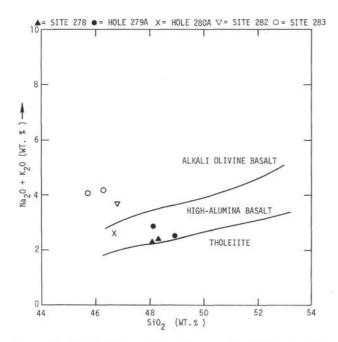


Figure 3. Alkali-silica variation diagram, Leg 29 basalts (after Kuno, 1968, p. 627).

Often considerable emphasis in classification is placed on differences in alumina content between calcalkaline and tholeiitic basalts. Generally, calcalkaline basalts 283Y. The "New Zealand" basalts show, in major oxide contain 17%-20% Al2O3, whereas tholeiitic basalts contain 12%-17%. The contrast should be especially prominent on a plot of Al2O3 against normative plagioclase composition (Figure 4). Considered as a group, the Leg 29 rocks straddle the field boundary, as do abyssal basalts from the Mid-Atlantic Ridge (Irvine and Baragar, 1971, p. 536).

A summary of the results obtained by classifying the Leg 29 basalts according to the schemes shown in Figures 2, 3, and 4 is presented in Table 4. Table 5 may give the best appreciation of the intermediate character of the rocks. Clearly, the Leg 29 basalts have chemical aspects of both oceanic tholeiites and oceanic alkaline basalts, reported by Engel et al. (1965, table 5).

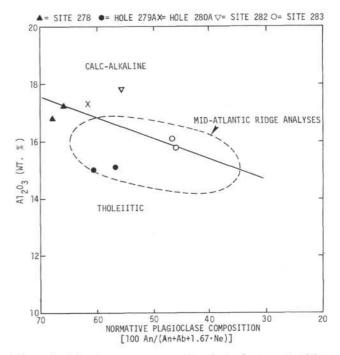


Figure 4. Alumina versus normative plagioclase composition, Leg 29 basalts (after Irvine and Baragar, 1971, p. 536).

REFERENCES

- Bence, A. E., Papike, J. J., Chandrasekharam, D., and Cameron, M., 1973. Petrology of basalts from Leg 15 of the Deep Sea Drilling Project; the central Caribbean, and Petrology of basalts from Leg 17 of the Deep Sea Drilling Project; central Pacific Basin: Am. Geophys. Union Trans., v. 54, no. 11, p. 995-1001.
- Engel, A. E. J., Engel, G. G., and Havens, R. C., 1965. Chemical characteristics of oceanic basalts and the upper mantle: Geol. Soc. Am. Bull., v. 76, p. 719-734.
- Irvine, T. N. and Baragar, W. R. A., 1971. A guide to the chemical classification of the common volcanic rocks: Canadian J. Earth Sci., v. 8, p. 523-548.
- Kuno, H., 1968. Differentiation of basalt magmas. In Hess, H. H. and Poldervaart, A. (Eds.), Basalts: New York (Interscience), p. 623-688.
- MacDonald, G. A. and Katsura, T., 1964. Chemical composition of Hawaiian lavas: J. Petrol., v. 5, p. 82-133.

	Classification of Leg 29 Basalts						
		Classification Method					
Hole	AFM (Figure 2)	Alkali-Silica Varia- tion Diagram (Figure 3)	A1 ₂ O ₃ vs. Normative Plagio- clase Composition (Figure 4)				
278	Tholeiitic basalt	High-alumina basalt	Tholeiitic basalt ^a				
279A	Tholeiitic basalt	High-alumina basalt	Tholeiitic basalt				
280A	Tholeiitic basalt	High-alumina basalt	Calcalkaline basalt				
282	Tholeiitic basalt	Alkali olivine basalt	Calcalkaline basalt				
283	Calcalkaline basalt	Alkali olivine basalt	Calcalkaline basalt ^a				

TARLE A

^aOne of the two analyses plots on the field boundary.

Weight (%)	Oceanic ^a Tholeiite	Leg 29 ^b Basalt	Oceanic ^c Alkalic Basalt
SiO ₂	49.94	47.3	47.41
A1203	16.69	16.4	18.02
Fe ₂ O ₃	22.01	4.4	4.17
FeO	6.90	4.4	5.80
MgO	7.28	6.2	4.79
CaO	11.86	10.9	8.65
Na ₂ O	2.76	2.7	3.99
K20	0.16	0.36	1.66
H ₂ O+		1.7	
н20-	-	2.7	2.45
TiO ₂	1.51	1.32	2.87
P205	0.16	0.17	0.92
MnO	_	0.14	-
CO_2	-	0.88	
Sum	99.17	99.57	100.73
ppm			
Mn	-	1378	-
Ba	14	40	498
Co	32	41	25
Cr	297	302	67
Cu	77	96	36
Ni	97	113	51
Sc	130	42	815
Sr	130	134	815
V	292	293	252
Y	43	27	54
Zr	95	91	333
Ga	-	15	-
Yo	—	3	1
Li	9	1211	11

 TABLE 5

 Comparative Chemical Compositions of Oceanic Basalts

^a42 analyses, Engel et al., 1965, p. 731. ^b8 analyses, this report. ^c10 analyses, Engel et al., 1965, p. 725.