7. SITE 296

The Shipboard Scientific Party1



Figure 1. Location of Site 296 and track of Glomar Challenger. From map: "Topography of North Pacific," T. E. Chase, H. W. Menard, and J. Mammerickx, Institute Marine Resources, Geol. Data Center, Scripps Institution of Oceanography, 1971. Contour depths in kilometers. Scale: 1:6,500,000.

SITE DATA

Position: 29°20.41'N; 133°31.52'E.

Water Depth (from sea level): 2920 corrected meters (echo sounding)

Bottom Felt At: 2958 meters (drill pipe)

Penetration: 1087 meters

Number of Holes: 1

Number of Cores: 65

Total Length of Cored Section: 612 meters

Total Core Recovered: 312.1 meters

Percentage of Core Recovery: 51%

Oldest Sediment Cored: Depth below sea floor: 1087 meters

Nature: Volcanic tuff and conglomerate Age: Early Oligocene(?) Measured velocity: 2.7 km/sec

Principal Results: Site 296 was drilled on a sediment-covered terrace high on the west flank of the Palau-Kyushu Ridge. The stratigraphic sequence consists of 453 meters of late Oligocene to Pleistocene ash-bearing, clav-rich, and clavey nannofossil oozes/chalks overlying more than 634 meters of early to late Oligocene volcaniclastics in which the hole bottomed (terminated). The clayey chalk-ooze interval provides an excellent biostratigraphic reference section and record of Neogene planktonic events beneath the Kuroshio Current. Displaced littoral foraminifera indicate that portions of the ridge were at or near sea level during the late Oligocene, whereas Neogene bathyal species document later subsidence of the ridge. The boundary between Oligocene volcaniclastics and younger chalks may coincide with rifting of the ridge after initial opening of the Parece Vela Basin in the late Oligocene.

BACKGROUND AND OBJECTIVES

Background

Site 296 is located on a northwest-west-trending structural bench or terrace near the northern terminus of the Palau-Kyushu Ridge adjacent to the Nankai Trough (Figures 1, 2). This major ridge has been most recently viewed as a remnant arc of Paleogene age (Karig, 1971; Uyeda and Ben-Avraham, 1972) which has undergone subsidence, perhaps beginning in the Oligocene in conjunction with rifting of the arc. Importantly, the Palau-Kyushu Ridge is rather broad at its northern terminus providing a topographically isolated platform for the accumulation of calcareous pelagic oozes above the calcium carbonate compensation depth (CCD), thus preserving a record of mid-Tertiary to Recent planktonic events in the overlying Kuroshio Current system.

Precruise study of LDGO Vema-21 reflection profiles revealed a suitable area displaying 500 to 700 meters (0.4-0.5 sec) of acoustically transparent sediment on the west side of the Palau-Kyushu Ridge which was confirmed by additional Glomar Challenger profiles (Figure

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Figure 2. Track of Glomar Challenger approaching and departing Site 296. Other well-navigated tracks, for which bathymetric and reflection data were available, are also shown.

3); this area was ultimately selected as the site for Hole 296 (Figure 4). Thus, drilling at this position near the ridge terminus offered an equally attractive setting for recovery of data bearing on planktonic biostratigraphy of this mid-latitude area, as well as ridge history. Moreover, it was also thought that possible variations in the position of the convergence zone between the northward-moving warm Kuroshio Current and subarctic water of the Oyashio Current might have occurred in this region in response to major climatic fluctuations during the Neogene. If so, these variations might be expressed as stratigraphic variations in planktonic biofacies at Site 296.

Objectives

One of the major overall goals of Leg 31 was to obtain a series of biostratigraphic reference sections at sites traversing tropical through subarctic zoogeographic provinces extant in the marginal northwestern Pacific. The outstandingly complete Eocene through Pleistocene sequence recovered at Site 292 on Benham Rise formed the initial and wholly tropical biostratigraphic section within this series. The primary objective at Site 296 was the recovery of an equally valuable reference secion deposited beneath the subtropical transitional water mass of the Kuroshio Current, the major western boundary current in the Pacific Ocean. It was hoped that such a mid-latitude site would allow recovery of planktonic assemblages containing tropical, subtropical, and perhaps temperate species. The assemblages recovered would provide assistance in correlation of well-established tropical planktonic zonations with cooler water biofacies as well as detailing paleooceanographic trends.

The second major objective at Site 296 was centered on recovery of paleontologic and lithologic data which might aid in deciphering the age, origin, and history of the controversial Palau-Kyushu Ridge. It was anticipated that variations in occurrence and character of volcanic debris might prove especially helpful in this respect.

OPERATIONS

Site 296 had been planned primarily as a biostratigraphic hole with the additional objectives of exploring the possible subsidence, basement character, and volcanic activity on the Palau-Kyushu Ridge. Locating such a biostratigraphic hole was difficult because the area chosen had to be substantially shallower than the CCD, far enough north to produce a mid-latitude flora and fauna, and beneath sufficiently productive waters. The northern end of the Palau-Kyushu Ridge presented the only suitable site area. A favorable area was located on LDGO Vema-21 reflection profile, which showed about 500 meters of transparent sediment on a bench along the west side of the Palau-Kyushu Ridge. The constant thickness of the sediment on this bench suggested that the section consisted of pelagics, with little reworking. The track, headed directly from Site 295 toward the point on the Vema track which crossed this plateau. However, it was decided to survey the area briefly to insure that we were not adjacent to local topographic highs.

Presite Survey

The short survey began by reducing speed to 8.5 knots along a 020° track between sites and continuing until reaching the area of the basement core of the ridge. A second survey leg, along 255° contained data, that together with the *Vema* profile and our 020° survey, indicated that the plateau trended north-northwest. The data also revealed an adequate sediment pile between the major basement high and the edge of the plateau. After continuing 4 miles beyond the chosen site, the course was reversed, and a 16-kHz beacon was dropped. Because of a faulty signal from the beacon while descending to the bottom, the profiling gear was quickly restreamed, and a third pass made over the site. A 13.5kHz beacon was dropped on this pass.

The water depth was 2920 meters, and the hole spudded in on 10 July with an initial punch core. Continuous coring, at a rate of about one core per hour (Figure 5), continued in nannofossil ooze to a depth of 472 meters (Table 1). At this depth the sediment changed gradually to coarse volcaniclastics, which formed acoustic basement. A decision was made to core and drill through the volcanics which were thought to be relatively thin because flows or very coarse material had not been encountered and the fossil content had remained high. However, after 700 meters of drilling with cores taken at an increasing spacing with depth (Table 1) and with the loss of fossil control, it was decided to abandon the site.



Figure 3. Glomar Challenger seismic reflection profiles across the Palau-Kyushu Ridge near Site 296.

The final, 65th, core was run in conjunction with an inclinometer test which showed the hole to be vertical. The pipe was pulled out of the hole at 1900 on 14 July.

Because there was insufficient current to run an onsite sonobuoy, a sonobuoy was dropped upon leaving the site. The ship left at low (2 knots) speed with one gun streamed. After completing a semicircle about 1 mile in diameter, the ship proceeded toward Site 297.

LITHOLOGY

Site 296 is located on a sediment-covered terrace on the western side of the Palau-Kyushu Ridge at a water depth of 2920 meters. The lithologic sequence is outlined in Table 2 and on Figure 5.

Unit 1

The unit is predominantly a 453-meter-thick clayey nannofossil ooze/nannofossil clay, but exhibits the following possible subunits: (A) The 0-44.5 meter interval (Cores 1 to 5) is a foraminifera-rich nannofossil clay, with occasional volcanic ash zones, ash-rich zones, or interbeds. The colors are grays and olive grays; (B) The 44.5-63.5 meter interval (Cores 6 and 7) is a foraminifera/clay-rich to clayey nannofossil ooze, with colors in grays, greenish-grays, and olive-grays; (C) This subunit

is defined on the basis of a decrease in the foraminifera content and an increase in micarb. The basic lithology is a clay/micarb-rich nannofossil clay. It is found between 63.5 and 92.0 meters (Cores 8 to 10), with colors of green-gray, olive-gray, and gray; (D) Cores in the 92.0-206.0 meter interval contain a subunit, which is basically a clayey (clay-rich) nannofossil ooze and chalk. The lithification to chalk first becomes apparent in Core 15, Section 4 (134.5 m) and occurs in zones or interbeds with ooze thereafter. Local volcanic ash, ash-rich zones occur. The colors are predominantly light gray, and light olive-gray; (E) A color change to yellow-gray, grayish-orange, and yellow-brown occurs at 206 meters. The basic lithology (clayey nannofossil ooze/chalk) remains the same; however, radiolarians appear as a lithologic component, from 5% to a high of 20% in radiolarian-rich zones. Volcanic ash, ash-rich zones are common and become increasingly prevalent in Cores 28, 29, and 30 (253.5-282.0 m); (F) At 282 meters, lithification is dominant in the sediments. Although the upper 6.5 meters (Core 31) is a nannofossil-rich claystone, the dominant lithology is a gravish-orange to yellow-brown clayey (locally clay-rich) nannofossil chalk. Volcanic ash interbeds occur in Cores 35 and 36. The subunit is defined as occurring between 282 and 343.5 meters (Cores



Figure 4. Bathymetry in the vicinity of Site 296. The discontinuous and en-echelon nature of the basement ridge segments shows well. Dashed lines are bathymetric control from Bracey (1966, p. 11).

31 to 37); (G) A noticeable color change from yellowbrown and pale orange to green-gray occurs at 343.5 meters in Core 37. The lithology in this core is a clayrich nannofossil chalk; however, beginning in Core 38 (348.5 m) very significant amounts of volcanic ash, ashy chalk, clayey ash interbeds occur through Core 48 (445 m). There is a slight tendency to a decreasing ash influence in Cores 45 and 46. The X-ray analyses show significant amounts of augite (4.8%-25.3%) in the bulk, 2-20 μ m, and <2 μ m fraction in the interval 370.9 (Core 40) to 435.5 meters (Core 47).

In general, for Unit 1, bioturbation is most noticeable from 254 meters to 453 meters. The age is late Oligocene to late Pleistocene/Holocene.

Unit 2

Unit 2 begins at 453 meters and is a series of subunits of tuffs, lapilli tuffs, volcanic sandstones and siltstones, and occasional ash-rich/bearing nannofossil chalks.

Tuffs

The tuffs and lapilli tuffs contain lithic and crystal fragments that are essentially textural variants from adjacent volcanic sources. A major lithic type is a porphyritic pyroxene andesite, with phenocrysts of oscillation zoned plagioclase, hypersthene, augite, opaques, and occasionally hornblende in a pilotaxitic to hyalophitic or glassy matrix. The matrix consists of plagioclase microlites in a microgranular to glassy groundmass of pyroxene, with scattered specks of opaques. The size of the phenocrysts is essentially constant, whereas the groundmass grain size is highly variable.

1) Phenocrysts and clasts: The plagioclase phenocrysts usually show oscillatory zoning in the range from Anas to Ans2 and are full of tiny black specks of pyroxenes and opaques, which are often remelted to a brown glass. Augite is the predominant pyroxene, but hypersthene is often present. The hypersthene like the plagioclase may be inclusion filled, with the inclusions remelted to a brownish glass. In some cases the augite is replaced by hornblende. Along with euhedral opaque phenocrysts, the plagioclase and pyroxene often form large glomeroporphyritic aggregates, which themselves occur as lithic fragments in the tuff. Other abundant lithic fragments are vesicular basalt, pumice, spherulitic andesite, and brown glass. The distribution of the lithic fragments is such that the finer-grained volcanics such as the pumice, glass, and hyalopolitic-textured andesites predominate in the lower portions of the unit. Fragments with coarser-grained matrices are most abundant near the top of the unit. Still other clasts are the phenocryst phases of the andesites. These are predominantly plagioclase crystals, with associated augite, hypersthene, opaques, and hornblende. In general, the crystal fragments are more abundant in the finer-grained tuffs.

The clasts are generally angular to subangular, and poorly to moderately well sorted. Many of the tuffs are found as well-defined 3-meter-thick units, grading upward from lapilli tuffs to tuffs. In any graded unit, the sorting improves as the grain size becomes finer. Deeper in the section flattened clasts show a slight imbricate structure.

2) Matrix: The matrix of the tuff (from 10% to 20%) is composed primarily of palagonitized glass, which is often spherulitic. Near the top of the unit, the tuff has textures reminiscent of a welded tuff, or graywackes, where grain boundaries of the clasts often merge into that of the matrix. The palagonite matrix becomes more clayey and serpentinized with depth, so that deeper than Core 60, the matrix is essentially a clay-serpentine mixture. In addition, smaller fragments of lithics and minerals become more prevalent in the groundmass of the deeper cores. Small vugs in the matrix are often filled with acicular radiating crystals of zeolites. Calcareous nannofossils, shallow-water benthonic foraminifera, and micarb (algae) are found in the groundmass of Cores 53 and 56.

Volcanic Sandstones and Siltstones

A second sublithology present in Unit 2 are volcanic sandstones and siltstones. These are significant subunits starting at 750 meters. The sand/silt sized volcanic fragments are subrounded/subangular to angular, illustrate excellent graded units, poor to good sorting, and have excellent sedimentary structures including cut and fill, load casts, slumping, foreset bedding, and megaforeset bedding. The fragmental material is identical to that found in the tuffs.

Volcanic Ash-Rich/Bearing Nannofossil Chalks-Volcanic Ash

These occur as bearing (2%-10%), or rich (10%-25%) components in nannofossil chalks, or clay-rich nannofossil chalks. These sublithologies were present in Cores 52, 54, and 56. Characteristics of this sublithology



Figure 5. Hole summary diagram, Site 296.

TABLE 1 Coring Summary, Site 296

	Cored Interval	Ganad	Reco	vered			
Core	(m)	(m)	(m)	(%)	Remarks ^a		
1	0.0-6.5	6.5	5.9	91	Punch core		
2	6.5-16.0	9.5	7.7	81			
3	16.0-25.5	9.5	5.9	62			
4	25.5-35.0	9.5	5.0	53			
5	44 5-54 0	9.5	5.5	74			
7	54.0-63.5	9.5	4.3	47			
8	63.5-73.0	9.5	5.0	53			
9	73.0-82.5	9.5	9.5	100			
10	82.5-92.0	9.5	9.0	94			
11	92.0-101.5	9.5	5.1	54			
12	101.5-111.0	9.5	4.7	50			
13	120 5-130 0	9.5	0.5	100			
15	130.0-139.5	9.5	5.5	58			
16	139.5-149.0	9.5	8.0	84			
17	149.0-158.5	9.5	7.0	74			
18	158.5-168.0	9.5	2.9	31			
19	168.0-177.5	9.5	8.5	89			
20	177.5-187.0	9.5	4.9	52			
21	196 5-206 0	9.5	0.0	91			
23	206.0-215.5	9.5	3.4	36			
24	215.5-225.0	9.5	7.1	75			
25	225.0-234.5	9.5	5.5	58			
26	234.5-244.0	9.5	6.6	69			
27	244.0-253.5	9.5	3.1	34			
28	253.5-263.0	9.5	6.6	69			
29	203.0-272.5	9.5	5.0	94			
31	282 0-291 5	9.5	71	75			
32	291.5-301.0	9.5	6.1	64			
33	301.0-310.5	9.5	5.2	55			
34	310.5-320.0	9.5	4.5	48			
35	320.0-329.5	9.5	7.3	77			
36	329.5-339.0	9.5	8.1	85			
38	348 5-358 0	9.5	0.1	42			
39	358.0-367.5	9.5	5.2	55	0		
40	367.5-377.0	9.5	5.0	53			
41	377.0-386.5	9.5	3.1	33			
42	386.5-396.0	9.5	1.3	14			
43	396.0-405.5	9.5	1.1	12			
44	405.5-415.0	9.5	0.4	4			
45	424 5-434 0	9.5	5.4	55	-		
47	434.0-443.5	9.5	0.3	3	30 bbls mud.		
48	443.5-453.0	9.5	0.3	3	1020-0-071-0201200		
49	453.0-462.5	9.5	0.4	4			
50	462.5-472.0	9.5	0.2	2			
51	472.0-481.5	0.5	0.0	0.0			
Drill	491.0-548.0	2.5	0.0	0.0	50 bbls mud.		
52	548.0-557.5	9.5	2.7	30.0	e e e e e e e e e e e e e e e e e e e		
Drill	557.5-567.0		10000				
53	567.0-576.5	9.5	0.6	1.0	ost street to		
Wash	576.5-624.0			1.0.0	50 bbls mud.		
54	624.0-633.5	9.5	3.9	43.0			
SS	652 5-662 0	0.5	14	15.0			
Drill	662 0-690 5	9.5	1.4	15.0			
56	690.5-700.0	9.5	10.0b	100.0			
Drill	700.0-709.5			0.00.000	50 bbls mud.		
57	709.5-719.0	9.5	4.5	47.0	13		
Drill	719.0-747.5						
58	747.5-757.0	9.5	4.4	46.0			
59	785.5-795.0	9.5	1.2	13.0			
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TABLE 1 - Continued

	Cored Interval Below Bottom	Cored	Recov		
Core	(m)	(m)	(m)	(%)	Remarks ^a
Drill	795.0-823.5				50 bbls mud.
60	823.5-833.0	9.5	1.1	12.0	
Drill	833.0-861.5	101010	12.2.5.4.7.7.		
61	861.5-871.0	9.5	2.0	21.0	
Drill	871.0-899.5				
62	899.5-909.0	9.5	2.2	23.0	
Drill	909.0-966.0				50 bbls mud.
63	966.0-975.5	9.5	4.9	52.0	
Drill	975.5-1070.5		1.1165	00000000	50 bbls mud.
64	1070.5-1080.0	9.5	5.1	53.0	
65	1080.0-1087.0	7.0	3.0	43.0	
Total	1087.0	612.0	312.1	51.0	

^aSee Figure 5 for graph of drilling rates.

^bExtra 0.5 meter beyond core barrel.

U	nit and Descriptions	Depth (m)	Thickness (m)	Age
1A	Foraminifera-rich nannofossil clay	0-44.5	44.5	Early Pleistocene- Pleistocene/Holocene
1B	Foraminifera/clay- rich, (clayey) nanno- fossil ooze	44.5-63.5	19.0	Early Pleistocene
1C	Clay/micarb-rich nannofossil ooze	63.5-92.0	28.5	Late Pliocene- early Pleistocene
1D	Clayey nannofossil ooze/chalk	92.0-206.0	114.0	Late Miocene- Late Pliocene
1E	Clayey nannofossil ooze/chalk with radiolarians and radiolarian-rich zones	206.0-282.0	76.0	Early Miocene- late Miocene
1F	Clayey nannofossil chalk	282.0-343.5	61.5	Late Oligocene- early Miocene
1G	Clay-rich nannofossil chalk with extensive ash, ashy chalk and clayey ash interbeds	343.5-453.0	109.5	Late Oligocene
2	Volcanic tuffs, lapilli tuffs, volcanic sand- stones/siltstones	453.0-1087.0	≌634	Early (?) to late Oligocene

TA	BLE 2		
Unit Descriptions, Depths,	Thicknesses,	and Ages,	Site 296

included: lamination, thin bedding, sharp basal contacts, associated "fall" fragments, and bioturbation. Some of the ashy zones exhibited current traction features. The ash content varied from 7% to 30%.

Unit 2 is at least 65 meters thick, with an age of early(?) to late Oligocene. The early Oligocene date is based on marginal nannofossil floras which were recovered in Core 65 (1087 m).

Lithologic Interpretations

The geologic history of the Site 296 area is characterized by a dominant eruptive volcanic phase up through late Oligocene. This activity waned from late Oligocene to the Holocene, being replaced by pelagic nannofossil sedimentation.

The volcanogenic subunits of Unit 2 illustrate the combined effects of pyroclastic accumulation by settling through the water column, depositional characteristics associated with settling, and minor redistribution by gravity transfer and bottom current mechanisms.

1) The lapilli tuffs and tuffs, with their range of sorting from poor to moderate; the grading of the volcanic fragments (lapilli to tuff); the angular-subangular fragment shape; and in particular the glassy matrix are indicative of a near-direct accumulation from a volcanic eruptive process. Shallow-water foraminifera and algal carbonate masses serve to indicate a shallow source for the pyroclastic products during certain eruptive periods. Graded bedding, imbricate structures, and current structures may emphasize sedimentary processes occurring during the initial accumulation, as well as transport over the bottom to deeper water. In general, there is increasing evidence of sedimentary structures toward the base of the unit.

Diagenetic changes associated with time and depth have imparted characteristics to these lapilli tuffs and tuffs. These changes include: consolidation, devitrification, hydration to palagonite, or hydration leading to diffuse fragment boundaries, and alteration of fragments to a clay-serpentine material. These characteristics are especially noticeable deeper in the cored section.

2) The angularity, grading, and composition of the volcanic ash, and its association with nannofossil chalks in Cores 52, 54, 56, and 57 support the idea that they are ash-fall accumulations which settled through the water column.

3) Volcanic sandstones and siltstones contain subrounded to subangular clasts and show moderate to good sorting and current structures. This evidence can be used to infer postdepositional redistribution by bottom currents or gravity transfer processes as single events or a concurrent interplay of volcanic settling and redistribution.

The thickness and age of the volcanic sediments in Unit 2 provide support for a rapid build-up (317 m/m.y.) of a volcanic pile. Factors such as oversteepening, liquification, or earthquakes could trigger a gravity transfer mechanism, and subsequent transfer processes on the sea floor. The lithologic variations throughout Unit 2 and the large- and small-scale graded units support periodicity of the eruptive and depositional processes in this area.

4) A rather significant waning of eruptive volcanic activity, and a decrease in sediment redistribution by bottom transfer processes is noted in Unit 1. A decrease in volcanic activity is noted upwards from Core 28, and the decrease becomes more evident toward the top of the unit.

Pelagic nannofossil sedimentation is quite apparent in Unit 1, showing occasional volcanic eruptive cycles throughout its 453-meter thickness. Trace amounts of detrital minerals (quartz, feldspar, and heavies), plus clay minerals (up to 25%), attest to contamination of the nannofossil ooze via settling in the water column or settling from a bottom turbid current or drift. Radiolarians are present in the upper portions of the cored interval, becoming absent with increasing depth, and reappearing again in Cores 22 to 38. They are generally absent thereafter. Foraminifera are also variable in their content, but occur down to Core 56 of Unit 2.

PHYSICAL PROPERTIES

Bulk Density, Porosity, and Water Content

In general, a minor increase in bulk density can be observed in a downward direction in this hole, and no significant difference can be seen between the nannofossil ooze, nannofossil chalk, and carbonate containing volcanic ash (Figure 5). Water content reveals the same irregular pattern, with a slight overall decrease in downward direction.

Vane Shear

The shear-strength measurements of the nannofossil ooze show a well-defined trend to 50 meters increasing rapidly over this interval. At greater depths, the data show more scatter and the rate of strengthening decreases. The scatter of measurements between 50 to 155 meters is partly due to drilling deformation, but also may be a function of periodic variation in natural consolidation/lithification observed at depths greater than 100 meters.

A comparison with the brown mud at Sites 294/295 shows, that for a given depth, the nannofossil ooze has a higher shear strength. Early incipient lithification and greater internal friction probably account for higher shear strengths of the nannofossil ooze. Additional discussion will be found in Bouma and Moore (this volume).

Sonic Velocities and Thermal Conductivities

Sonic velocities and thermal conductivities were both measured on the same core section. Syringe samples for water content were also collected at the same locations where thermal conductivities were measured, as long as the cores were soft enough. During the continuous coring, cores were brought up at about 1-hr intervals. This situation did not always allow thermal uniformity in the core to be attained, and most conductivity measurements were made 1.5 to 2.5 hr after cores were brought on deck. In a few cases, it was recognizable that the results were affected by thermal nonuniformity.

The results of sonic-velocity and thermal-conductivity measurements are tabulated in Tables 3 and 4, and summarized in Figures 5 and 6. Sonic velocity is almost constant in the upper 250 meters, while thermal conductivities increases slightly. At depths of about 260 and 410 meters, discontinuities are present in the trend of sonic velocities, while thermal-conductivity measurements show little change in trend at 260 meters. At the depth where sonic velocity became 1.8 km/sec, the sediment became too hard for thermal-conductivity measurements. A slightly slower velocity in vertical direction than in horizontal direction was suggested by a few measurements.

Thermal conductivities estimated from water content coincide well with the values by needle-probe method with few exceptions, suggesting that Ratcliffe's formula was applicable for deeply buried sediments.

Scattering in values is much larger for the thermal conductivities than for the sonic velocities, and this may partly be explained by a technical problem of sampling: sonic-velocity measurements were made on lessdisturbed parts of sediments based on selection in a split core, while thermal-conductivity measurements were made at points apparently selected through the plastic liner before splitting, and it was normally impossible to

Sonic-Velo	city Measu Site 296	rements,
Sample (Interval in cm)	Depth in Hole (m)	Velocity (km/sec)
1-3, 77	3.77	1.528
2-6, 50	14.50	1.506
3-3, 35	19.35	1.521
4-2, 52	27.52	1.499
4-2, 54	27.54	1.499
5-3, 55	38.55	1.517
6-4,100	55.04	1.522
8-3 37	66.87	1.526
8-4, 37	68.37	1.513
8-4,90	68.90	1.513
9-2, 111	75.61	1.511
9-5, 33	79.33	1.501
10-3, 43	85.93	1.501
10-6, 38	90.38	1.511
10-6,42	90.42	1.530
12-5.74	108.24	1.506
12-5,63	108.13	1.510
13-3, 35	114.35	1.517
13-3, 36	114.36	1.520
16-3,63	143.13	1.517
16-3, 28	142.78	1.528
16-3, 36	142.86	1.523
17-5, 20	155.20	1.556
17-5, 78	155.31	1.518
18-2, 36	160.36	1.570
19-2, 62	170.12	1.509
20-2, 52	179.52	1.522
21-3, 26	190.26	1.506
22-5, 115	203.65	1.533
23-3, 71	209.71	1.510
24-4, 20	220.20	1.536
26-5.0	240.50	1.529
27-3, 30	247.30	1.521
28-4,60	258.60	1.608
28-4,89	258.89	1.621
28-4, 110	259.10	1.620
29-5,0	269.00	1.615
30-5, 64	279.14	1 724
31-3, 30	285.30	1.672
34-5,80	317.30	1.738
35-5, 147	327.47	1.762
35-5, 144	327.44	1.776
36-5, 148	336.98	1.778
36-5, 144	336.94	1.721
38-2 110	351 10	1.805
39-4, 132	363.82	1.749
40-3, 80	371.30	1.791
43-1,63	396.63	1.800
45-3, 101	419.01	2.120
54-3, 62	627.62	2.875
54-3,68	627.68	2.809
56-1 20	699 70	4.034
56-7, 28	699.78	2.625
56-1, 2	690.52	2.785
56-1, 2	690.55	2.816

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identify the degree of deformation exactly. Consequently, it is conceivable that the real values of the thermal conductivities may be always slightly higher than the measured ones.

GEOCHEMICAL MEASUREMENTS

Alkalinity, pH, and salinity measurements are summarized in Table 5.

Alkalinity

The average alkalinity for Hole 296 is 2.20 meq/kg. Six of the values are higher than the surface seawater reference value of 2.35 meq/kg, and they are found in Cores 1 to 25. Two cores (Core 40, Section 1, and Core 45, Section 2) show very significant low values of 0.59 and 0.78 meq/kg, respectively. Both cores are from Unit 1, Subunit 1G, a clay-rich nannofossil chalk, with extensive ash, ashy chalk, and clayey ash interbeds.

pН

The average pH values obtained by both punch-in and flow-through methods were all below that of the seawater reference at the site (8.17 to 8.19). The seven punch-in pH values averaged 7.24, while the 12 flowthrough values averaged 7.65. The most noticeable change in pH values is that shown by the flow-through values of Cores 40 and 45. These two values exceed 8.4. The two high pH values correspond to the two very low alkalinity values (0.59 and 0.78) reported. These cores of Unit 1, Subunit 1G contained substantial volcanic ash interbeds and volcanic components in the biogenous sediments.

Salinity

Twelve salinity measurements at Site 296 averaged $35.6^{\circ}/_{\circ\circ}$. A fairly definite increase in salinity with depth is noticed, except for some variability in Cores 40, 45, and 52. All 12 values and their average were above the overlying seawater reference value of $34.4^{\circ}/_{\circ\circ}$.

PALEONTOLOGIC SUMMARY

Introduction

The thick and unusually complete sedimentary section penetrated at Site 296 ranges in age from Pleistocene to late Oligocene and may include basal sediments as old as early Oligocene. Nannofossils and planktonic foraminifera occur quite abundantly throughout most of the sequence, whereas radiolarians, when present, occur rather intermittently. The state of preservation for all groups ranges from poor to good, with a general decrease in quality downward in the hole.

Age and zonal determinations at Site 296 were made with the use of calcareous nannofossils, planktonic foraminifera, radiolarians, silicoflagellates, and diatoms in Pleistocene sediments. Only planktonic foraminifera and calcareous nannofossils were utilized in Pliocene sediments. These two groups together with radiolarians were utilized in the Miocene and late Oligocene portion

Sample		Thermal	Conductivit	ty in 10 ⁻³ cal/cm sec°C					
(Interval	Hole Depth	Needle		From Water					
in cm)	(m)	Probe	Average	Content					
1-3.35	3	1.75							
1-3, 55	4	2.03	1.90	2 08 +0 11					
1-3, 105	4	1.93	1.50	2.00 10.11					
2-2,	· · · ·		2.12						
2-2, 110	9	2.12		2.01 ±0.11					
6-6,40	14	2.38	2.47	2.19 ±0.12					
6-6, 110	15	2.57	12000000						
3-3, 40	19	2.25	2.15						
3-3, 110	20	2.05	1.07	2.17 ± 0.12					
4-2,40	27	2.01	1.97	2.02 ± 0.11					
5-3 40	38	2.01	2.04	217 ± 012					
5-3, 105	39	2.00	2.01	2.17 20.12					
6-4,40	49	1.89	2.00	2.19 ±0.13					
6-4, 110	50	2.10		CUSS EAVER					
7-2, 40	56	2.07	2.22	2.13 ±0.12					
7-2, 110	57	2.47		S					
8-3, 40	67	2.17	2.13	2.19 ±0.13					
8-3, 110	68	2.09	2.00						
8-4,40	68	1.88	2.09	0.00.14					
8-4, 110	69 75	2.30	2.22	2.33 ±0.14					
9-2, 40	76	2.23	2.22	2.22 ± 0.13					
9-5, 40	79	1.95	2.04	2.22 ± 0.13 2 14 +0 12					
9-5, 110	80	2.12	2.01	2.11 20.12					
10-3,40	86	2.09	2.09	2.06 ± 0.11					
10-3, 110	87	2.08	17 CENTO						
10-6,40	90	2.32	2.29	2.19 ± 0.12					
10-6,110	91	2.26							
11-3, 40	95	2.02	2.07						
11-3, 110	96	2.12	2.22	2.11 ± 0.12					
12-5, 75	108	2.24	2.22	2.19 ±0.12					
14-5 40	127	2.20	2.25	1 86 +0 10					
14-5, 110	128	2.35	2.20	1.00 10.10					
16-3, 40	143	2.28	2.28	2.29 ±0.13					
16-3, 110	144	2.28							
16-5,40	146	2.08	2.23	2.36 ± 0.14					
16-5, 110	147	2.38							
18-2, 40	160	2.44	2.28	0.04 .0.10					
18-2, 110	101	2.11		2.34 ± 0.13					
19-2,40	170	2.15	2.10						
20-2, 40	179	2.58	2.40						
20-2, 120	180	2.22							
21-3,40	190	2.21	2.08						
21-3, 110	191	1.94		2.32 ± 0.13					
25-4, 40	230	2.52	2.32						
25-4, 120	231	2.12							
26-5,40	241	2.38	2.19						
20-5, 120	242	2.45	2.02	(Hard part)					
28-4 120	250	1.61	2.05	(fait part)					
30-5 60	279	217	218	(Soft part)					
30-5, 110	280	2.19	2.10						
31-2, 24	284	2.45	2.24						
31-2, 104	285	2.03	00.000						
32-3, 30	295	2.59	2.57						
32-3, 120	296	2.55							
33-4, 60	306	1.64	2.06						
33-4, 104	307	2.47	2.05						
35-5, 25	320	2.52	2.65						
36-5 34	336	2.01	2.85						
36-5, 115	337	2.79	2.05						
37-3, 32	342	2.50	2.49						
37-3, 114	343	2.47	1000 1000 1000 1000						
38-3, 33	352	2.85	2.52						
38-3, 121	353	2.18							

TABLE 4 Thermal Conductivities Measured at Site 296

Thermal Conductivity in 10^{-3} cal/cm sec °C



Figure 6. Thermal-conductivity results, Site 296.

of the sequence. Calcareous nannofossils of possible early Oligocene age are present in the basal portion of the section.

Both planktonic foraminiferal and calcareous nannofossil evidence places the Pliocene-Pleistocene boundary between Cores 7 and 8 (63.5 m), based upon the zonal schemes of Blow (1969) and Bukry (1973b). Virtually all of the Pliocene nannofossil and foraminiferal zones are recognized in Cores 8 through 17 (63.5-158.5 m). The Miocene/Pliocene boundary is placed between Cores 17 and 18 (158.5 m) on the basis of calcareous nannofossils, or alternately, between Cores 16 and 17 (149.0 m) on the basis of planktonic foraminiferal zones recognized by Ujiié and Oki (in press). If the zonal correlations of Berggren and Van Couvering (1973) are utilized in interpreting nannofossil and planktonic foraminiferal ranges at Site 296, then placement of the Miocene/Pliocene boundary utilizing these two groups would coincide at the base of Core 17 (158.5 m).

The Miocene sequence is marked by the apparent absence of several nannofossil, planktonic foraminiferal, and radiolarian zones. This suggests that significant intervals of erosion or nondeposition occurred on this portion of the Palau-Kyushu Ridge during this epoch. For example, two calcareous nannofossil subzones are missing between Cores 17 and 18 (158.5 m), with one subzone missing in Core 23, and another absent between Cores 27 and 28 (253.5 m). The absence of these various subzones represents gaps of 1 to 4 m.y., following the zonal scheme of Bukry (1973b). Furthermore, four planktonic foraminiferal zones (N10 through N13) are missing within Core 26 (234.5-244.0 m), and at least one radiolarian zone is absent within the Core 26 through 28 interval.

Despite the apparent loss of several microfossil zones within Miocene sediments at Site 296, calcareous nannofossil and planktonic foraminiferal evidence places the late Miocene/middle Miocene boundary within Core 23 (approximately 209 m), and the middle Miocene/early Miocene boundary within, or at the base of Core 28 (263 m).

Differences in definition cause the controversial Miocene/Oligocene boundary to be placed between Cores 33 and 34 (310.5 m) on the basis of calcareous nannofossil zonation, and within Core 36 (334 m) on the basis of planktonic foraminiferal zonation. A thick late

Sample	Depth	p	Н							
(Interval in cm)	Below Sea Floor (m)	v Sea Fl r (m) Punch-in thr		Alkalinity (meq/kg)	Salinity (°/00)	Lithologic Units				
Surface seawate	er reference	8.17	8.19	2.35	34.4					
1-3, 144-150	4.5	7.31	7.58	3.23	34.9	Unit 1				
2-5, 144-150	14.0	7.31	7.37	3.32	34.9	Cubunit 1 A				
5-3, 144-150	39.5	7.35	7.55	2.93	35.2	Subuint 1A				
10-5, 144-150	90.0	7.24	7.57	2.35	34.4	Unit 1, Subunit 1C				
15-3, 144-150	134.5	7.14	7.36	3.03	34.4	Unit 1,				
20-3, 140-150	182.0	7.32	7.35	2.74	34.6	Subunit 1D				
25-3, 144-150	229.5	7.02	7.24	2.83	35.2	Unit 1,				
30-4, 144-150	278.5	-	7.35	1.96	35.2	Subunit 1E				
35-5, 144-150	327.5		7.39	1.37	36.6	Unit 1, Subunit 1F				
40-1, 143-150	369.0		8.42	0.59	36.3	Unit 1,				
45-2, 144-150	418.0	-	8.44	0.78	38.2	Subunit 1G				
52-1, 144-150	549.5	-	8.16	1.27	37.4					
Average	real graduitation	7.24	7.65	2.20	35.6	Unit 2				

 TABLE 5

 Summary of Shipboard Geochemical Data, Site 296

Oligocene sequence is present in Cores 34 through 63 (310.5-975.5 m). Cores 64 and 65 (1070.5-1087 m) contain only a few rare, and relatively long-ranging nannofossils whose maximum age cannot be older than early Oligocene. This is based on the exceedingly high sedimentation rates represented by the thick volcaniclastic sequence at Site 296. However, the actual age range of the species identified is mid-early Eocene to early Oligocene. Possible reworking of some of these forms cannot be ruled out since the specimens are poorly preserved. In any event, the basal sediments at Site 296 are most probably late Oligocene in age and no older than early Oligocene.

Finally, it is important to note that a few specimens of larger benthonic foraminifera (*Lepidocyclina* sp.), displaced from a littoral environment, occur in Cores 55 and 56 (662.0-693.5 m) and clearly indicate that portions of the Palau-Kyushu Ridge were at or near sea level during the late Oligocene.

Calcareous Nannofossils

Abundant, diverse, well-preserved nannofossil assemblages were recovered from throughout the Neogene part of the section. There is a marked decrease in the species diversity and state of preservation in the Oligocene assemblages. Fourteen of the 17 zones established by Bukry (1973b) were recognized in samples from the Holocene to Oligocene interval of the Cenozoic.

Floods of the very small coccolith, *Emiliania huxleyi*, were observed in samples from Cores 1 through 3, Section 1, which clearly places these samples in the Holocene-late Pleistocene *E. huxleyi* Zone.

A complete sequence of Pleistocene nannofossil zones can be recognized in Core 3, Section 1 through Core 7. The *Gephyrocapsa oceanica* Zone, and the *G. caribbeanica* and *Emiliania annula* subzones are readily identified. Core 8 marks the first occurrence of discoasters; consequently, this appearance is interpreted as indicating the top of the Pliocene.

A normal progression of Pliocene and late and middle Miocene zones can be recognized in Cores 8 through 28, except for the early Pliocene Ceratolithus acutus Subzone, the late Miocene Triquetrorhabdulus rugosus and Discoaster bellus Subzones, and the middle Miocene Coccolithus miopelagicus Subzone. These subzones should be present in the intervals between Cores 17 and 18, Core 23, Sections 2 and 3, and Cores 27 and 28, respectively. According to Bukry (1973a), 0.9 m.y. is represented by the Ceratolithus acutus and Triquetrorhabdulus rugosus subzones, nearly 4 m.y. is represented by the Discoaster bellus Subzone, and 0.6 m.y. is represented by the Coccolithus miopelagicus Subzone. The absence of these various subzones may reflect a hiatus in the sedimentary record, or it may indicate that zone-defining species were just not recognized.

There is a normal sequence of foraminiferal zones representing continuity through the interval, with the missing nannofossil zones described above. Similarly, a break in the continuity of foraminiferal and radiolarian zones within Core 26 occurs within the nannofossil Discoaster kugleri Subzone of the Discoaster exilis Zone, and no interruption in the nannofossil sequence is apparent. This clearly demonstrates that while the three fossil groups are in general agreement with regard to age determinations, some specific differences need to be resolved.

Cores 34 through 63 represent a thick, fine to coarse, clastic sediment interval. A normal progression of the three late Oligocene nannofossil zones (Cyclicargolithus abisectus Subzone of the Triquetrorhabdulus carinatus Zone, Sphenolithus ciperoensis Zone, and S. distentus Zone) was recognized in these cores. The sphenolith index species for the latter two zones are unusually small and occur sparsely in these samples, and their recognition requires rather extreme microscopic techniques. The boundary between the S. ciperoensis Zone and the S. distentus Zone cannot be clearly defined. If a specimen of S. ciperoensis in the sample from Core 56 is correctly identified, then the boundary must be between Cores 56 and 57. However, if the first occurrence of S. ciperoensis is in reality in Core 52, then the boundary lies between Cores 52 and 53. Consequently, until this problem can be resolved, the interval represented by Cores 53 through 56 is designated a transitional interval.

Cores 64 and 65 contain only a few specimens of the nannofossil species *Helicopontosphaera compacta, Dictyococcites bisectus,* and possibly *Cyclococcolithina formosa.* These species have reported occurrences ranging from the middle early Eocene to the early Oligocene. While they may not be very age definitive, they may provide some indication of the possible minimum and maximum ages for these samples. If the identification of *C. formosa* in Core 65 is correct, and if they are not reworked specimens, then the sample can be no younger than early Oligocene.

Foraminifera

A nearly continuous biostratigraphic succession was observed from the lower Pleistocene (lower N22 Zone) to the upper Oligocene (lower N21 Zone) at Site 296, except for the absence of the N13 to N10 zones in the lower middle Miocene, between Core 26, Section 5 and Sample 26, CC. It is interesting that the hiatus correlates roughly with a similar gap in the fossil record observed at Site 292. A hiatus at this same time has also been recognized in many of the sedimentary basins on the Pacific coast of the Japanese Islands. Therefore, this hiatus may be regional in nature and not due to local events.

Zone P21 of the late Oligocene is represented by faunas in samples of a layer of more than 731 meters (Cores 38 through 64) thick, providing evidence of a truly rapid rate of sedimentation.

Some larger benthonic foraminifera, including *Lepidocyclina (Eulepidina)* cf. formosa and probably others, are scattered throughout Core 56, Section 2 and Sample 56-1, 137 cm. The subgenus *Eulepidina* ranges from Aquitanian to middle Oligocene; consequently, this age agrees with that determined by the planktonic foraminifera. A displaced fauna, consisting of four species of small benthonic foraminifera that are characteristic of shallow water, and a fragment of coral were found in Sample 55, CC.

Radiolarians and Silicoflagellates

Radiolarians occur in the samples from Site 296 in the topmost (Cores 1 and 2) and in the middle part of the section (Cores 21 to 47), but they are virtually absent from Cores 45 to 65. It was also noticed that, although Site 296 is located beneath the track of the warm-water Kuroshio Current, several age-diagnostic index forms are absent, and ranges of others are different from those reported from lower latitude areas.

A well-preserved Pleistocene assemblage is present in Core 1 as would be expected in surface sediments of a modern warm-water region, but abundance is sharply decreased in Core 2. Radiolarians are completely absent from Cores 3 through 23, except for occasional rare specimens of collosphaerids, found in Cores 21 and 24.

The Ommatartus antepenultimus Zone is recognized in sediments of Core 25 and in part of Core 26. Radiolarians are rare in Core 27. Sections 2 and 4 of Core 28 belong to the Cannartus laticonus Zone of Moore (1971). Sediments equivalent to the Cannartus pettersoni Zone are apparently missing from this site despite continuous coring. Sample 28, CC to Core 30, Section 4 belong to the Calocycletta costala Zone; the interval from Sample 30, CC to at least Core 33, Section 2 is assigned to the Calocycletta virginis Zone. Sediments within Core 33, Section 4 to Core 44, Section 1 contain only marginal abundances of radiolarians as well as the absence of guide species making age assignment difficult. Radiolarians are absent from Core 45 to the base of Site 296 (Core 65).

A few specimens of the silicoflagellate *Dictyocha fibula* and its variety, *D*. f. var. *aculeata*, were observed in Core 1 representing the initial recognition of silico-flagellates on Leg 31.

Diatoms

Core 1 contains the Quaternary species Coscinodiscus excentricus, C. lineatus, C. nodulifer, C. radiatus, Nitzschia marina, Pseudoeunotia doliolus, Rhizosolenia bergonii, and Roperia tesselata. These specimens occur in frequent abundance and are moderately well preserved and represent tropical species.

Preservation of diatoms in Core 2 is poor, but a few specimens were recovered which belong to the species *Rhizosolenia bergonii*, *Thalassionema nitzschioides*, and *Thalassiothrix longissima*. None of these are age diagnostic.

The sample from Core 3 contains only a few fragments of *Thalassionema nitzschioides* and *Thalassiothrix longissima*. No additional diatoms were found in samples from Cores 4 through 65.

SUMMARY AND INTERPRETATIONS

Summary

Hole 296 was spudded in a water depth of 2920 meters on a structural bench about 30 km west of the crest of the Palau-Kyushu Ridge (Figures 2, 3, and 4) initially penetrating an acoustically transparent layer composed of pelagic oozes. Drilling ultimately penetrated a total of 1087 meters into an entirely sedimentary section representing two major depositional phases in the Holocene through late Oligocene history of the rise.

Continuous coring was maintained to a depth of 472 meters due to biostratigraphic objectives revealing that this portion of the Palau-Kyushu Ridge is capped by a 453-meter-thick Holocene/Pleistocene through late Oligocene sequence of nannofossil oozes and chalks containing variable amounts of volcanic debris. These sediments were placed in Unit 1 and further subdivided into seven secondary units based upon degree of lithification, variable percentage of foraminifera, and abundance of volcanic ejecta. Chalks initially appear at 134.5 meters (Core 15), with complete lithification to chalk at 282 meters (Core 31). Abundant planktonic foraminifera and calcareous nannofossils occur throughout Unit 1; however, radiolarians are found only sporadically with diatoms and silicoflagellates restricted to Pleistocene sediments. Unfortunately, the apparent absence of several foraminiferal, nannofossil, and radiolarian zones mars the continuity of the Miocene portion of this section with paleontologic gaps present in the Core 17 through 28 sequence representing an estimated stratigraphic loss of 1 to 4 m.y. Although these gaps interrupt the biostratigraphic continuum present in this section, they alternately provide evidence of an early mid Miocene period of active erosion or nondeposition on the rise.

Volcanic ash and evidence of bioturbation increase in the lower portion of Unit 1, with distinct ash interbeds beginning at 348.5 meters (Core 38) forming a transitional zone with underlying Unit 2 which is predominantly composed of volcaniclastic material. In fact, continuous coring was halted at 472 meters in this latter series of late Oligocene volcanic ashes, lapilli tuffs, and volcanic sandstones due to poor core recovery. The boundary between the ooze-chalk sequence of Unit 1 and underlying Unit 2 was placed at 453 meters. This lithologic transition apparently correlates with the reverberant acoustic "basement" noted in continuous reflection profiles and a sonobuoy record at Site 296 (Figures 3 and 5). These same records suggest that volcaniclastic Unit 2 fills troughs between buried highs of the ridge and onlaps older basement ridge flanks.

Interval coring within Unit 2 penetrated a strikingly thick (634 m) series of coarse volcaniclastics to a depth of 1087 meters at which point the hole was abandoned due to loss of fossil control and the prospect of many more meters of similar sediment before any underlying lithology might be reached. The major lithic type within the tuffs and lapilli tuffs of Unit 2 is porphyritic andesite with a matrix composed primarily of palagonitized glass. Volcanic sands, sandstones, and siltstones increase in abundance from 750 meters to 1087 meters and display well-preserved sedimentary structures including cross bedding, cut-and-fill, graded bedding, and load casts. The structures attest to vigorous redistribution of this material by bottom currents in association with slumping and mass movement on the Oligocene ridge slope (see Bouma, this volume); enclosing tuffs and lapilli tuffs exhibit sorting and structures indicative of near-direct accumulation from a volcanic eruptive process. Additional evidence of the displaced nature of these deposits is provided by the occurrence of littoral and neritic species of benthonic foraminifera (including the larger tropical foraminifer *Lepidocyclina*) along with algal debris in Cores 55 and 56. In fact, many of the individual beds within Unit 2 are marked by sharp basal contacts and in all likelihood represent single instantaneous events associated with volcanic eruptions, ash falls, and mass movement downslope with attendant erosion of underlying material resulting in a depositional rate exceeding 140 m/m.y.

Somewhat surprisingly, calcareous nannofossil and planktonic foraminifera are generally present in low abundance throughout Unit 2 and indicate that the bulk of the volcaniclastic sequence can be placed within a portion of the late Oligocene correlative with planktonic foraminiferal Zone P21 of Blow (1969) and the *Sphenolithus distentus* Zone of Bukry (1973b) emphasizing the rapid deposition of this unit. However, fossils are sparse in the basal part of the unit with a meager nannofossil flora indicative of an early Oligocene age.

Interpretation

The primary biostratigraphic objective was most certainly fulfilled by recovery of the continuously cored sequence of Pleistocene through late Oligocene nannofossil oozes and chalks assigned to Unit 1. These sediments contain an excellent calcareous paleontologic record of planktonic evolution and productivity within the mid-Kuroshio Current system for the past 26-27 m.y., complementing the equivalent equatorial record obtained at Site 292. Estimated rates of deposition within the latest Oligocene through late Miocene portion of the chalk-ooze sequence average about 9 m/m.y., whereas Plio-Pleistocene oozes accumulated at an estimated rate of 32 m/m.y. (Figure 7), signifying accelerated productivity and overturn within the Kuroshio surface water coincident with the onset of climatic deterioration in the late Neogene. Variations in species abundance (Ujiié, this volume) also appear to represent responses to latitudinal shifts of critical isotherms in the marginal northwestern Pacific with subtropical (transitional) and tropical species alternately dominant during portions of the Miocene, Pliocene, and Pleistocene. However, the present study does not suggest that the major Oyashio-Kuroshio convergence zone migrated as far south as Site 296 (latitude 29°N) even during the most severe periods of Quaternary refrigeration.

Unfortunately, the age of the Palau-Kyushu Ridge and its probable pre-Oligocene history were not clarified by the results from Site 296. The basement core of the ridge, which apparently underlies the volcaniclastics could consist of earlier volcanics, deformed sediments, or even a metamorphic-plutonic complex, and of an age as possibly as old as Cretaceous. The late Eocene volcaniclastic apron west of the Palau-Kyushu Ridge, drilled at Site 290, and mixed Eocene volcanics and sediments on Palau afford a firm minimum age further south, but there is no well-defined apron west of Site 296, and the assumption that the Palau-Kyushu Ridge has a similar history along its entire length is hazardous. A Cretaceous origin would now hang on the very



Figure 7. Sedimentation rates-Site 296. Based on correlations of planktonic foraminiferal zonation with time scale of Berggren (1972).

speculative assumption that the several reworked late Cretaceous planktonic foraminifera in Hole 290 were derived from the Palau-Kyushu Ridge.

If, for the moment, the Palau-Kyushu Ridge is considered to have a similar history along its length, then it witnessed two Tertiary volcanic pulses which were related to island arc tectonism. The first peaked in the late middle Eocene to late Eocene time and formed the major apron west of the ridge. The second, of late (?) early Oligocene to late Oligocene age, probably did not greatly enlarge the apron. An eastward thinning of the post Eocene (Site 290) pelagic brown clay which covers the distal part of the apron from 120 meters to less than 60 meters adjacent to the ridge (unpublished *Antipode* 4, 14 reflection and 3.5-kHz profiles), indicates that the influx of volcanic material to the head of the apron ceased only 20-30 m.y. ago.

Although volcanism along the Palau-Kyushu Ridge ceased by the end of the Oligocene as marked by the boundary between Unit 1 and Unit 2 in Hole 296, it continued in the Shikoku and Parece Vela basins to the east, and constitutes a prominent volcanic sequence on Guam, Yap, and other segments of the eastern arc systems. The cutoff of volcanism along the Palau-Kyushu Ridge thus corresponds well with the beginning of volcanism on the next ridge east and with the probable extensional development of both the Shikoku and Parece Vela basins. The most coherent integration of present information would support the idea that until the late Oligocene, the Palau-Kyushu Ridge was part of a frontal arc and contained the volcanic chain with attendant rapid deposition of volcaniclastics at rates as high as 140 m/m.y. on the ridge and ridge flanks. Near the Oligocene-Miocene boundary, the Parece Vela and Shikoku basins were formed by crustal extension within the part of the frontal arc near the volcanic chain. Within a short time, the volcanic chain, which stays with the active frontal arc, had left the Palau-Kyushu Ridge and became reestablished along the new frontal arc. As the Shikoku Basin opened to the east, the steep rifted eastern flank of the Palau-Kyushu Ridge developed (Figures 3 and 4) and with increasing width of the basin, the source moved further from the ridge. The coarse volcaniclastics were succeeded by wind and waterborne ash deposition which declined to a low level by the end of the Oligocene (Donnelly, this volume).

These abandoned ridges, or remnant arcs (Karig, 1972) seem to subside after rifting together with the cooling basins to either side. The displaced larger late Oligocene benthonic foraminifera present in Unit 2 provide clear evidence that portions of the ridge were at or near sea level and harbored a tropical reef biota at that time. Moreover, initial accumulation of nannofossil oozes (Unit 2) suggests major subsidence of the ridge occurred during this same interval allowing deposition of dominantly biogenic debris beneath the then through the flowing Kuroshio Current. Prior to this event the Kuroshio Current may well have been deflected eastward by the arc-ridge complex or perhaps split into an eastern and western fork. Further study is needed to sort out these various possibilities of reorganization of surface flow.

The presence of lower bathyal benthonic foraminifera throughout the late Oligocene-Pleistocene chalk-ooze unit at Site 296 also points to major subsidence during the post-rifting phase of ridge history. However, submergence of all portions of the ridge to depths in excess of 150 meters may not have occurred until the late Pleistocene judging from the continual presence of displaced shallow-water benthonic foraminiferal species throughout most of this sequence.

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APPENDIX A	
Summary of X-Ray, Grain Size, and Carbon-Carbonate Results, Site	296

	Sample Depth				Bulk Sam	ple	2-2	0µm Fra	ction	<2	μm Fra	tion	(rain Si	e		Ca	rbon Carbo	nate	
Section	Below Sea Floor (m)	Lithology	Age	1 1	2	3	Maj 1	2	3	1	2	3	(%)	(%)	(%)	Classification	(%)	(%)	(%)	Comments
296-1-2 296-1-3 296-2-6 296-3-1 296-3-3 296-3-4 296-5-4 296-5-4	2.5 4.2 14.2 16.4 19,0-19.2 20.5 39.8 40.3	Unit 1 Gray, olivegray foraminifera-rich 1 A Nannofossil clay with volcanic ash zones	E. Pleist. -late Pleist- Holocene	Cale. Cale. Cale.	Quar. Quar. Quar.	Mica Mica Mica	Quar. Quar. Quar.	Plag. Plag. Mica	Mica Mica Plag.	Mica Mica Mont.	Quar. Quar. Quar.	Plag. Mont. Mica	32.9 10.7	53.5 54.3	13.6 35.0	Sandy Silt Clayey Silt	3.4 5.4 5.7 3.7 6.0 3.3	0.4 0.3 0.1 0.3 0.1 0.3	25 43 47 28 49 25	
296-6-4 296-6-5	50.0 51.3	B Gray, foraminifera/ clay-rich-clayey nannofossil ooze	Early Pleistocene														6.9 4.4	0.1 0.3	57 35	
296-8-4 296-9-4 296-10-1 296-10-5 296-12-4 296-13-2 296-14-4 296-15-3 296-15-3 296-16-4	68.5 69.1 78.9 83.9 89.2 107.0 113.5 125.8 133.3 134.0 144.8	IC Green, olive-gray clay-rich nannofossil ooze	Late Pliocene to early Pleistocene	Cale.	Quar.	Mica	Quar.	Mica	Plag.	Mica	Quar.	Mont.	0.6 0.6 58.5	33.6 21.7 25.5	65.7 77.7 15.9	Silty Clay Clay Silty Sand	3.4 6.8 5.9 6.4 7.1 7.5 6.6 5.7 9.5	0.4 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1	25 56 48 52 58 61 54 47 79	
296-17-4 296-19-2 296-20-3 296-21-2 296-22-2 296-22-2 296-22-4 296-28-5 296-28-5 296-28-5 296-30-4 296-31-5 296-36-1 296-35-1 296-36-2	154.5 170.3 181.9 189.4 198.8 208.7 220.5 228.8 239.7 260.1 277.4-277.7 288.1-288.2 311.5 321.0 331.9	1D Clay-rich nannofossil ocze and chalk with radiolarian/ clay-rich nannofossil chalks	Late miocene -late Pliocene	Cale, Cale, Cale, Cale, Cale, Cale, Cale, Cale,	Mica. Quar. Quar. Quar. Mica Plag. Quar. Plag.	Quar. Mica Mica Mica Quar. Mica Mica Mica Mont.	Quar. Quar. Quar. Quar. Quar. Plag. Quar. Plag.	Mica Mica Plag. Plag. Quar. Plag. Quar.	Plag. Plag. Mica Mica Mica Mica Mica	Mica Mica Quar. Quar. Mont. Mont. Mont.	Mont. Quar, Mont. Mica. Quar, Quar, Quar, Quar,	Quar. Mont. Mica Mont. Mica Mica Mica Plag.	2.0 4.3	50.6 49.3	47.4 46.4	Clayey Silt Clayey Silt	8.2 6.8 8.5 6.8 8.0 6.4 7.9 6.0 5.2 7.2 7.2 7.4 3.5 7.5 7.9 9.1	$\begin{array}{c} 0.1 \\ \hline 0.1 \\ 0.1 \\ 0.1 \\ 0.1 \\ 0.1 \\ 0.1 \\ 0.1 \\ 0.0 \\ 0.0 \\ 0.0 \\ 0.0 \\ 0.0 \\ 0.0 \\ 0.0 \\ 0.0 \\ 0.1 \\ 0.1 \\ 0.0 \end{array}$	68 56 70 56 66 53 65 50 43 60 61 28 62 66 76	Kaol. in <2μm (5.4%) Kaol. in <2μm (2.2%) Kaol. in <2μm (3.1%) Kaol. in <2μm (2.1%) Augi. in 2-2μm (7.5%) Kaol. in <2μm (2.6%)
296-38-2 296-39-2 296-40-3 296-40-3 296-40-4 296-40-4 296-40-4 296-41-2 296-41-2 296-41-2 296-41-2 296-41-2 296-42-1 296-45-1 296-46-4 296-47-1	350.8 360.2 370.9 371.0* 372.2-372.7 373.5 377.1 378.7 379.3-379.4 380.3 387.0-387.2 397.2 415.7 429.6 433.5	1E Greenish-gray clayey, clay-rich nannofossil chalks with volcanic ash interbeds	Early Miocene to late Miocene	Plag. Plag. Plag. Plag. Plag. Plag. Plag. Plag. Plag. Plag. Plag. Calc.	Mont. Augi. Augi. Plag. Augi. Mont. Augi. Mont. Augi. Calc, Plag.	Augi, Calc, Calc, Augi, Mont, Calc, Mont, Augi, Mont, Augi,	Plag, Plag, Plag, Plag, Plag, Plag, Plag, Plag, Plag, Plag, Plag,	Augi. Augi. Augi. Augi. Augi. Augi. Augi. Augi. Augi. Augi. Quar.	Quar. Chlo. Quar. Quar. Quar. Quar. Quar. Quar. Quar.	Mont. Plag, Mont. Plag, Plag, Mont. Plag, Plag, Plag, Mont.	Piag. Mont. Plag. Mont. Plag. Mont. Augi Plag. Plag.	Augi. Augi. Chlo. Augi. Augi. Augi. Quar. Quar. Augi. Quar.	0.2 86.6 19.4 72.4 0.4 53.7 28.0 73.4	65.3 7.9 64.8 17.5 74.9 36.7 55.1 20.0 73.5	34.4 5.4 15.8 10.1 24.7 9.6 16.9 6.5	Clayey Silt Sand Sandy Silt Silty Sand Clayey Silt Silty Sand Sandy Silt Sandy Silt	3.0 8.5 7.5 0.3 7.5 6.5 3.4 7.7	0.0 0.0 0.0 0.1 0.0 0.0 0.0	25 71 62 2 62 54 28 64	Clin. in <2μm (2.6%) Augi. in <2μm (8.8%) Augi. in bulk (6.6%) Clin. in <2μm (1.8%)
296-52-1 296-52-1 296-54-1 296-54-3 296-55-1	548.7 549.3 624.8 628.2 653.5	Unit 2 Volcanic tuffs sand- stones, siltstones with ash-rich nannofossil chalks	Early to late Oligocene (?)	Plag.	Calc.	Augi.	Plag.	Augi.	Quar.	Plag.	Augi.	Mont.	1.0 52.9 57.6	77.6 26.8 20.3	21.4 20.3 22.0	Silt Sand-silt-clay Sand-silt-clay	6.2 0.8	0.0 0.0	51 6	

Note: Complete results of X-ray Site 296 will be found in Part V, Appendix I. X-ray mineralogical legend on Appendix A, Chapter 2.



Ag-Aq23

čc.

Core Catcher 3GY 6/1

5GY 6/1

N6

Emiliania

Pseud N22

207



208



Ag- Ag- B

5GY 6/1 with N5

5Y 4/1 to 5GY 6/1

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Core -

Catche

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SITE 296

209

te	296	HOI	E066	TI	-	Cor	'e /	cored in	terv	a1:5	4,0-63.5 m						
	-	0	HAR	ACTE	R	NO	S		TION	MPLE							
AGE	ZONE	FORAMS	NANNOS	RADS	STLICO.	SECTI	METER	LITHOLOGY	DEFORMA	LITHO.SA	LITHOLOGIC DESCRIPTION						
						1	1.0	۷010 ۲	000			Intense deformation to near-breccia; extreme mottling and mixing of colors; greenish gray (SGY 6/1) and olive gray (SY 4/1), with gray (N6, N4). FORAMINIFERA/CLAY-RICH NANNOFOSSIL 00ZE CLAY-RICH NANNOFOSSIL 00ZE (Minor Lith) Smear: CC Texture Composition BSS Clay 699 Mannofossils					
caribbeanica Subzone	uribbeanica Subzone		Am-	В		2	of contrastor		0000000		5GY 6/1 with N6	7% Foraminifera 2% Micarb 1% Quartz 1% Heavy minerals Tr% Volcanic glass					
ALY PLEISTOCENE	N22 bzone/Gephyrocapsa ca					3		VOID	000								
EAR	Emílianía annula Su		Am-			4			0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0		5GY 6/1 and 5Y 4/4	with N4					
		Aq-	Aq-	в		Ca				čć.	5GY 6/1						

Sit	e 296	Ho	e		0	ore l	8 Cored I	nterv	al:6	63.5-73.0 m
Γ			FOS	SIL	ER	-		ION	PLE	
AGE	ZONE	FORAMS	NANNOS	RADS	SILICO. DIATOMS	METER	LITHOLOGY	DEFORMAT	LITH0.SAM	LITHOLOGIC DESCRIPTION
LATE PLIDCENE	N22 Cyclococcolithina macintyrei Subzone		Ag-	в	-	0.5 1.0 2 2				Very intense drilling deformation leading to extreme vertical streaking of greenish gray (SGY 6/1) with medium light gray (N6) = light and medium to medium dark gray (N5-N4) = dark; deformation less intense, lower in core. MICARB/CLAY-RICH NANOFOSSIL 00ZE Smears: 4-50, CC Texture 000000000000000000000000000000000000
	21 entaradiatus Subzone		Ag- Am-	В		4			* 50	SGY 6/1 with N6 SGY 6/1 with N5 and N4 - <u>Carbon Carbonate 4-50</u> SGY 6/1 6.8, 0.1, 56 with N6 <u>Carbon Carbonate 4-111</u> 3.4, 0.4, 25
	0. pc	Ag	Ag-	В		Core Catch			οc.	5GY 6/1

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SITE 296

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Ar Explanatory notes in chapter 1 VOID

L_1

Core

Catche

0 1-0

0 0 ----

5GY 6/1

5GY 6/1



Hole Site 296 Core 14 Cored Interval: 120.5-130.0 m FOSSIL DEFORMATION SAMPLE METERS ZONE LITHOLOGIC DESCRIPTION LITHOLOGY AGE CO. LITHO. FORAMS NANNOS NDS Intense deformation to breccia; color dominant greenish gray (56Y 6/1) to light gray (N7) at base. 0 0 0 0 CLAYEY NANNOFOSSIL 00ZE Smears: 4-75, CC Texture 100% Clay 0 5GY 6/1 Composition 50% Nannofossils 44-47% Clay minerals 2- 5% Foraminifera 1% Micarb 0 0 0 0 0 Carbon Carbonate 4-75 6.6, 0.1, 54 111 VOID +--HILLI I Subz EARLY PLIOCENE asymmetricus ò 0 0 Q Discoaster 1111 11 1 5GY 6/1 Cm 75 ò 0 11111111 00000 9 - 1-1 5GY 6/1 with N7 Ag Ag 5GY 6/1 * cc Core Catche

Explanatory notes in chapter 1



Explanatory notes in chapter 1

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Site 296 Hole Core 18 Cored Interval: 158,5-168.0 m FOSSIL NO SAMPL METERS STLICO. DEFORMATI ZONE LITHOLOGIC DESCRIPTION LITHOLOGY AGE NANNOS RADS ITHO. MWS Deformation slight to moderate; dominant color very light gray (N8), with areas of light olive gray (57 6/1); (N2) grayish black in burrow areas burrowing noticeable in all sections especially in chalky areas; or bud is conded VOID ÷. Subzone Cg- R NB MIOCENE primus ! ash bed is graded. -80 1..... 5Y 6/1 CLAYEY NANNOFOSSIL OOZE 1 Smear: 1-80 N17/N18 Ceratolithus p LATE Texture 100% Clay Composition 73% Nannofossils - . 25% Clay minerals 27 N2 2% Foraminifera 1 PYRITITE (from worm burrow) (minor) Smear: 2-27 -5Y 6/1 Composition 80% Pyrite 20% Nannofossils 1 Ag- Ag-+---Core N7 Catche 1 and 5Y 6/1

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Explanatory notes in chapter 1

Site 296		Hole	_	_	Co	re 2	3 (Cored In	terv	al:20	06.0-215.5 m					_	Site	296	Hole			Cor	e 24	Cored Ir	ter	val:2	15.5-225.0 m		
AGE	ZONE	FORAMS	SONNA	RADS	SECTION	METERS	LIT	THOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOG	IC DES	SCRIPTION			AGE	ZONE	FORAMS	FOSSHARA	RADS RADS BILICO. BILATOMS	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOG	SIC DESCRIPTION
MIDDLE MIDGENE LATE MIDGENE NIS	Discoaster hamatus Discoaster meorectus Subzone?	Ag	Ag- Ag-	B B Rm	1 2 3	0.5-				* 123 14 * cc	5¥ 7/2	Intense dri yellowish g deformation *Section 1 MICARB-RICH Smear: 2-12 Texture 95% Clay 5% Silt 0- 5% Sand CLAYEY NANN Smears: 3-1 Texture 80% Clay 15% Silt 5% Sand Carbon Carb 5.4, 0.1, 5 X-ray 2-123 80.7% Calc 7.9% Quar 7.1% Mica 4.3% Plag	111ing : ray (5) with c CLAYEN 3 (1 00FOSSI1 4, CC <u>5</u> 3 1 00FOSSI1 4, CC <u>5</u> 3 1 00FOSSI1 4, CC	de formation; c Y 7/2); modera chalk layers d tclan 2. Y NANNOFOSSIL Composition 25% Clay miner 25% Clay miner 5% Quartz 2% Feldspar 1% Foraminife 60% Nannofo 60% Nannofo 60% Nannofo 30-35% Clay mi 3- 5% Foramin 3- 5% Foramin 3% Micarb Tr% Volcani 2-123)	olor dominant te-slight n lower half 00ZE (Chalk) als ra rals ssils nerals ifera r c glass						Rm	0 1 1 2 3	1.0		000000000000000000000000000000000000000	*75	5Y 7/2	Color yell in Sections slight in Sand 4. CLAYEY NAU- Smears: J Texture 90% Clay 10% Silt Carbon Cart 7.9, 0.1, 6	ow gray (5Y 7/2); drilling breccia s l, 2, 6 to moderate deformation Section 4; chalk zones in Sections NOFOSSIL 00ZE (Chalk) 75, 4-50, CC <u>Composition</u> 63-71% Minorals 3-7% Micarb 1% Foraminifera Tr% Feldspar bonate 4-50 65
Explanat	tory n	otes	in c	hapt	er 1																	4	11			50			

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SITE 296

Explanatory notes in chapter 1

Ag- Ag- Rm

Fm

MIDDLE MIDCENE

N15 Discoaster hamatus 5

Core Catcher VOID

1

 5Y 7/2

	2.50		FOS	SIL ACTE	R		6 25	Cored II	NOI	FE	23.0-234.5 11			_
AGE	ZONE	FORAMS	NANNOS	RADS	STLICO.	SECTION	METERS	LITHOLOGY	DEFORMAT	LITHO.SAM		LITHOLOGIC	DESCRIPTION	
			Am-			1	0.5		00000			Drilling brecc: color dominant to medium yell zones in Section CLAYEY NANNOFO: Smears: 3-75, 4 Texture 85% Clay 10% Silt E% Sand	la to moderately deformed; grayish orange (10YR 7/4) wo brown (10YR 5/4); chalk ons 3, 4. SSLL 002E (Chalk) 495, CC <u>Composition</u> 40-45% Clay minerals	
NE	litus timus (R)		Ag-	Cg		2	ant and an		0000000		10YR 7/4	CLAY-RICH NANNO Smear: 4-115 Texture	5% Foraminitera 2-3% Micarb 1-2% Radiolarians 1-2% Heavy minerals 1% Quartz Tr-1% Volcanic glass DFOSSIL ASH (Minor Lith) Composition	
MIDDLE MIDCE	N14 Catinaster coa atartus antepenal		Ag-			3	drandram		0 	* 75		40% Sand 40% Silt 20% Clay	38% Volcanic glass 30% Nannofossils 15% Clay minerals 7% Feldspar 5% Micarb 3% Quartz 2% Heavy minerals	
	Ome			Ra			111 111					RADIOLARIAN/CL/ (Minor Lith) Smear: 4-145 Texture 80% Clay 15% Silt	AY-RICH NANNOFOSSIL CHALK <u>Composition</u> 55% Nannofossils 20% Clay minerals	
	f Subzone		Am-			4	undana		1	* 95 * 115	-	5% Sand	15% Radiolarians 3% Micarb 2% Feldspar 2% Heavy minerals 1% Quartz 1% Foraminifera	
	D. kugler	Ag-	Ag-	Cg		Ca	ore tcher			* CC	10YR 5/4 5Y 6/4	<u>Carbon Carbona</u> 6.0, 0.0, 50	t <u>e 3-75</u>	

FOSSIL SAMPLE DEFORMATION METERS ZONE LITHOLOGIC DESCRIPTION LITHOLOGY RADS SILICO. DIATOMS SECTIO AGE LITH0.5 NANNOS ORAMS Drilling breccia - moderate-slight drilling deformation down-core; colors grayish orange (10YR 7/4), with medium dark gray (N4), and moderate yellow brown (10YR 5/4); colors streaked in many cores. (8) 0.5 SI VOID 0 CLAYEY NANNOFOSSIL 00ZE (Chalk) Smears: 3-75, CC Texture Composition 100% Clay 40-45% Nannofos 0-0 Composition 40-45% Nannofossils 0 matartus 25-30% Clay minerals 3-10% Micarb 0 5% Quartz 2- 5% Radiolarians 2% Foraminifera 0 Cg Cm-2% Heavy minerals 1% Feldspar 1-+-10YR 5/4 1% Volcanic glass RADIOLARIAN/ASH-RICH NANNOFOSSIL CHALK (Minor Lith) Smear: 5-103 Texture 65% Clay 20% Silt Composition 56% Nannofossils 20% Volcanic glass 15% Radiolarians kugleri Subzo 10YR 7/4 * MIDDLE MIOCENE 15% Sand 3% Heavy minerals 2% Micarb 1% Quartz 1% Feldspar N14 1% Foraminifera Discoaster Tr% Glauconite Carbon Carbonate 4-73 5.2, 0.0, 43 Fm 10YR 5/4 <u>X-ray 4-74</u> (Bulk) 73.7% Calc 10.1% Quar 7.6% Mica 7.0% Plag 1.5% Mont + Rn 103 Ag- Ag- Cg cc. Core 5Y 7/2 Catche

Core 26 Cored Interval: 234.5-244.0 m

Explanatory notes in chapter 1

Site 296

Hole

te 296 Hole	Core 27 Cored Interval: 244.0-253.5 m	Site 296	Hole	Core 28 Cored Interval	: 253.5-263.0 m
FOSSIL CHARACTEI SONE RADIS RADIS	NOTITION STATE NOTITION STATE NOTITION NOTITION NOTITION NOTITION NOTITION NOTITION NOTITION NOTITION	AGE ZONE	FOSSIL CHARACTER SOUNNN SOUNNN SOUNNNN	DIATONS SECTION METERS METERS ASOTOHLIT DEFORMATION LITHOL SAMPLE	LITHOLOGIC DESCRIPTION
evoce source fin-	0.5- VOID 1 VOID 1.0- VOID 1.0- VOID 1.0- CLAYEY MANNOFOSSIL 002E (Chalk) Smears: 2-90, 3-90, CC 2	MIDDLE MIDCENE 1 Sphenolithus hetermorphus	Cannartus laticonus (R) -6V by by	0.5 1 VOID 1.0 2 2 4 4 4 4 4 4 4 4 4 4 4 4 4	Stiff chalk-coze; moderate drilling deformation; color grayish orange (1078 7/4), moderate yellow brown (1078 5/4); ash layers, burrows; dark colors are common. CLAY-RICH NANNOFOSSIL OOZE (Chalk) Smears: 1-125, 5-55 1078 7/4 95% Clay 50-55% Nannofossils 5% Silt 30-40% Clay minerals 5-7% Radiolarians 2-5% Micarb 1-2% Foraminifera 1% Volcanic glass 1078 5/4 NANNOFOSSIL-RICH VOLCANIC ASH (Minor Lith) 3% Silt 20% Nannofossils 1078 7/4 15% Clay 5% Feldspar 3% Quartz 4 1078 5/4 Ash at 2-82 has clay minerals 1078 5/4 Ash at 2-82 has clay mineral content of 15% 15% 15% 15% 15% 15% 15% 15% 15% 15%

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SITE 296

Explanatory notes in chapter 1

Ag-Ag- Cm

8

E N6

* 55

* CC

Core Catcher 10YR 7/4

10YR 7/4

Site 296	Hole Core 29 Cored Inter	rval: 263.0-272.5 m	Site 296 Hole Core 30 Cored Interval: 272.5-282.0 m	
AGE ZONE	F05STL CHARACTER NOTICA SUNDAL	LITHOLOGIC DESCRIPTION	30V 30V 30V 30V 30V 30V 30V 30V	
EARLY MIDCENE NB Heritcoontoschaera amDilaperta	Counter 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	 Chalk breccia; moderate-intense deformation; color mainly (10TR 7/4) grayish orange; burrows common; scattered pumice fragments. CLAYEY NANNOFOSSIL 002E (Chalk) Smears: 4-90, 6-75 Toyr 7/4 BSS Clay 55-602 Mannofossils 5-72 Foraminifera 3-72 Radiolarians 3-72 Radiolarians 3-75 Composition 3-75 Composi	Set Set <td>/4); ts.</td>	/4); ts.
	by b	* N6	Ag- Am- Cm Core Core Catcher CC 10YR 7/4	
	AG AG- Cm Core Catcher	toyr 7/4		

Site 296	Hole Core 3	1 Cored Interval: 282	2.0-291.5 m	Sit	e 296	Hole Cor	e 32 Cored I	nterval	:291.5-301.0 m	
AGE ZONE	POSSIL CHARACTER NOLL SUBJECT SUBJECT		AGE	ZONE	FOSSIL CHARACTER SOULIDIG SUDVIDIG SWDLVIG SUDVIDIG SWDVDJ	LITHOLOGY	DEFORMATION	110,000	LITHOLOGIC DESCRIPTION	
EARLY MIOCENE NG 1 Sphenolithus belemnos/Helicopontosphaera ampliaperta Calococlasta viroins (R)	Am. 1 Cm 2 Am- 3 Am- 4 Am- 5 Ag- Am- Fg- Core Core Catch		Colors grayish orange (10YR 7/4); intense deformation to a chalk breccia with slight deformation; pumice fragments, burrows common. NANNOFOSSIL RICH CLAYSTONE Smears: 1-80, 4-80, 5-16, CC Texture Commonition 505 Silt 58-655 Clay minerals 45% Clay 50-60X Nannofossils 5% Radiolarians 5% Quartz 3-5% Wicarb 2-5% Volcanic glass 3% Sponge spicules Tr- 1% Feldspar N4 Radiolarian, foraminifera-bearing; locally volcanic ash-bearing. <u>Grain Size 5-15</u> 3.5, 0.0, 28 <u>X-ray 5-15</u> (Bulk) 75.87 CalC 9.03 Plag 6.3% Mica 6.2% Quar 2.8% Mont 10YR 7/4	EARLY MIDCENE	N6 Discoaster druggii Subzone/Sphenolithus belemnos Calocyclatta virginis (R)	Am- Fm 2 Am- Am- 3 Am- Cm- 4 Am- 5 Ag- Am- Rp Car	.5		10YR 7/4	Color gravish orange (10TR 7/4); chalk breccia and slight deformation; burrow mottling common. CLAYEY NANNOFOSSIL CHALK Smears: 2-120, CC Texture Composition 70-75% Mannofossils 100% Clay To-75% Mannofossils 20% Clay minerals 1-2% Micarb 1% Voicanic glass 1% Foraminifera Tr-2% Radiolarians Tr% Sponge spicules

Explanatory notes in chapter 1

			FOSS	SIL	R	N	S		LION	MPLE	
AGE	ZONE	FORAMS	NANNOS	RADS	SILICO.	SECTIO	METER	LITHOLOGY	DEFORMA'	LITHO.SA	LITHOLOGIC DESCRIPTION
EARLY MIDCENE	N6 Discoaster deflandrei Subzone / Discoaster druggi Subzone Dorcadospyris atenchus - Lychnocanoma elongata (R)		Cm_	Rp		1 2 3				80	Colors grayish orange (10YR 7/4) to moderate yellow brown (10YR 5/4); slight deformation; moderate to intense burrow mottling throughout CLAYEY NANNOFOSSIL CHALK Smears: 3-80, CC Texture 100% Clay 10% Clay minerals 1% Micarb 1% Foraminifera Tr% Sponge spicules Tr% Volcanic glass 10YR 7/4 \$ 10YR 5/4
		Ag-	Am-	Rp+		C Ca	ore tcher			* cc	10YR 7/4

Site	296		Ho1	e		Co	re 34	Cored In	terv	al:3	310.5-320.0 m
AGE	ZONE		FORAMS	FOS: CHAR SONNAN	ACTER SOUTHIS	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION
LATE OLIGOCENE	N5/6 Cyclicargolithus abisectus Subzone	Dorcadospyris atenchus - Lychnocanoma eiongata (R)		Cm-	Fm	1 2 3 4 5	1.0			100	Color grayish orange (10YR 7/4): chert breccia with slight deformation; intense burrows; some dark yellow brown (10YR 4/2) and yellow brown (10YR 5/2) colors. CLAY-RICH NANNOFOSSIL CHALK Smear: 1-100, CC Texture Composition 1003 Clay Co-708 Nannofossils 15-25% Clay minerals 4-5% Volcanic glass 38 Quartz Tr- 38 Radiolarians 2% Feldspar Tr% Sponge spicules Tr% Heavy minerals Carbon Carbonate 1-100 7.5, 0.1, 62 X-ray 1-100 (Bulk) 90.5% Calc 3.4% Quar 2.4% Mica 2.2% Plag 1.5% Mont
			Ag-	Am-	Rp-	c	Core atcher			* cc	10YR 7/4



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ite	296	Hol	e	_		Cor	ne 37	Cored In	terv	al:	339.0-348.5 m	
			FOS	ACTI	ER	N	s		NOI	APLE		
AGE	ZONE	FORAMS	NANNOS	RADS	SILICO.	SECTIC	METER	LITHOLOGY	DEFORMAT	LITHO.SA		LITHOLOGIC DESCRIPTION
						1	0.5	VOID	1	*		Color: pale yellow brown (10YR 6/2) - very pale orange (10YR 8/2); slight deformation; extensive burrows; natural slickensides in Section 2; color green gray (5GY 6/1) in Sections 4 and 5. CLAY-RICH NANNOFOSSIL CHALK Smears: 1-140, 2-80 Texture <u>Composition</u> 100% Clay <u>76%</u> Nannofossils
	e Tonoata (R)			Rp		2	contraction.			* 80	10YR 6/2	20% Clay minerals 1- 2% Volcantc glass 1% Opaques Tr% Radiolarians Tr% Sponge spicules
E ULIGULENE	22 hus abisectus Subzon hus - Lvchnocanoma e					3	unterdation				10YR 8/2	
LAT	P; Cyclicargolit Dorcadospuris atenc		Am	- Cm		4	trial contraint				5GY 6/1	
						5	and and and					
		Ag	- Am	Rn		Ca	ore				5GY 6/1	

	200		FOS	SIL	ER	N	s		NOI	4PLE		
AGE	ZONE	FORAMS	NANNOS	RADS	SILICO.	SECTIC	METER	LITHOLOGY	DEFORMAT	LITHO.SA		LITHOLOGIC DESCRIPTION
LATE OLIGOCENE	P22 Sphenolithus ciperoensis Dorcadospyris atenchus - Lychnocanoma elongata (R)	Ag-	Cm-	Fm		1 2 3	0.5			* 70	N8 5Y 6/1 N7 5Y 6/1 N7 5Y 6/1 5Y 6/1	Generally very light gray (N8) to light gra (N7); slightly deformed; very bioturbated and slickensides noted; color darkens in Section 3 to light olive gray (5Y 6/1). CLAV-RICH NANNOFOSSIL CHALK Smear: 2-70 <u>Texture</u> <u>Composition</u> 95% Clay 72% Nannofossils 5% Silt 20% Clay minerals 4% Foraminifera 2% Heavy minerals 1% Volcanic glass 1% Volcanic glass 1% Volcanic glass 1% Volcanic glass 5% Silt 74% Volcanic glass 25% Silt 74% Volcanic glass 25% Silt 74% Volcanic glass 25% Clay 10% Heavy minerals 1% Nannofossils <u>Carbon Carbonate 2-80</u> 3.0, 0.0, 25

Explanatory notes in chapter 1

ite 296	Hol	e		Core	e 39	Cored In	ter	va1:3	58.0-367.5 m	Si	te 296	н	ole		c	ore 40	Cored In	nterv	al:3	67.5-377.0		
AGE ZONE	FORAMS	POSSIL CHARACTER NOTICINE SET LITHOLOGY WORK OF LITHOLOGIC DESCRIPTION				AGE	-	ZONE FORAMS		SSIL	SILICO. W	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE		LITHOLOGIC DE:	SCRIPTION				
LATE OLIGOCENE P22 Sphenolithus Caperoensis	Lychnocanoma bipes (R) Dorcadospyris atenchus - Lychnocanoma elongata (N)	Am-	m Rp-	0.1 1 2 3 4	.0 			140 * 65	Colors - greenish gray (SGY 6/1) to grayish burowing; sil git deformation; ash beds have saverage sol or nither deformation; ash beds have sol or nither deformation; ash beds have sol or not or not deformation; ash beds have sol or not deformation; ash beds sol or not deformation; ash beds deformation; ash beds have sol or not deformation; ash beds deformation; ash beds have sol or not deformation; ash beds deformation; ash beds deformati	LATE 0LTEOCEN	P22 Cohener118buse of successive	Lychnocanoma bipes (R) Dorcadospyris atenchus - Lychnocanoma elongata (R)	g- An	Rm i- Rm-	1 2 3 4 4	0.5	VOID GEOCHÉM SAR		* 65	5GY 6/1 with N3	Interbedded ash d greenish gray (56 bases; graded, bu deformation for c CLAYEY NANNOFOSSI Smear: 4-65 Texture 1002 Clay CLAYEY VOLCANIC A Smear: 4-16 Texture 64.83 Silt 19.44 Cain Size 3-44 0.2, 65.3, 34.4 Grain Size 4-16 19.4, 64.8, 15.8 Carbon Carbonate 7.5, 0.0, 62 X-ray 3-44 (Bulk) 50.22 Plag 21.93 Mont 11.33 Augi 9.22 Caic 4.13 Mica 3.33 Quar X-ray 4-16 (Bulk) 56.33 Plag 13.43 Mont	ark gray (N3) and chalk Y 6/1); ash beds have sh rrows at top; slight ore. L CHALK Composition 50% Nannofossils 30% Clay minerals 33% Micarb 33% Foraminifera 33% Volcanic glass 53% Fordspar 15% Clay minerals 30% Foldspar 15% Clay minerals 10% Heavy minerals 10% Lay Minerals 10% Heavy Minerals 10%

Explanatory notes in chapter 1

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ray (N3) and chalk 1; ash beds have sharp at top; slight


5.3% Mont 3.8% Quar

Explanatory notes in chapter 1



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Explanatory notes in chapter 1

			FOS	SIL	ER	N	S		NOT	4PLE	
AGE	ZONE	FORAMS	NANNOS	RADS	SILICO.	SECTIO	METER	LITHOLOGY	DEFORMAT	LITH0.SAM	LITHOLOGIC DESCRIPTION
LATE OLIGOCENE	P22 Sphenolithus cipercensis		Fm-	В		1 Ca	0.5 1.0 1.0 itcher	VOID	1	130	Medium light gray (N5) with black (N1); sandy unit (ashy) at base of Section 1 = boundary to Unit 2. CLAY/ASH-RICH NANNOFOSSIL CHALK Smear: 1-130 Texture Composition 55% Clay 53% Nannofossils 30% Silt 20% Clay minerals 15% Sand 15% Glass, 11th 3% Meavy minerals 2% Feldspar

			FOS	SIL	R	N	~		NOI	MPLE	
AGE	ZONE	FORAMS	NANNOS	RADS	SILICO. DIATOMS	SECTIO	METER	LITHOLOGY	DEFORMAT	LITHO.SA	LITHOLOGIC DESCRIPTION
LATE OLIGOCENE	P22 inolithus ciperoensis					I 1	5	VOID			Unit 2 - (Lapilli) Tuff, black (N2), frial clasts consist of pumice fragments, glass mineral fragments, red pumice; subrounded angular outlines; sandy matrix of glass; Lapilli sizes for fragments with some ash- sizes. N2
	Sphe		Cm			Cot	re cher	VOID			

Explanatory notes in chapter 1

			FOS	SIL	R	N	~		NOL	4PLE	
AGE	ZONE	FORAMS	NANNOS	RADS	brarows	SECTIO	METER	LITHOLOGY	DEFORMAT	LITHO.SAM	LITHOLOGIC DESCRIPTION
						1	0.5	VOID			Black (N2), (Lapilli) Tuff, friable, fragment consist of pumice, glass, and minerals; sub- angular to angular, with sizes from very- fine sand to granules; some associated ASH-RICH NANNOFOSSIL CHALK.
				в	Ì	c	ore	VOID			

Explanatory notes in chapter 1

ite	290			FOS	SIL	ER		0	corec II	NOI	PLE		
AGE		ZONE	FORAMS	NANNOS	RADS	SILICO. DIATOMS	SECTIO	METER	LITHOLOGY	DEFORMAT	LITHO. SAM		LITHOLOGIC DESCRIPTION
IGOCENE	P22	ciperoensis		Rm	в		C Ca	ore tcher				N2	LAPILLI TUFF recovered in core catcher only; similar to Cores 49-50.
LATE OL		Sphenol1thus											

Explanatory notes in chapter 1



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Explanatory notes in chapter 1

R Rm

Core Catcher

Sph

30% Clay 25% Sand

bioturbated, laminations Grain Size 3-115 52.9, 26.8, 20.3

Carbon Carbonate 1-82 0.8, 0.0, 6

25% Clay minerals

5% Nannofossils 5% Feldspar 3% Heavy minerals



Cored Interval: 690.5-700.0 m Site 296 Core 56 FOSSIL SAMPLE CHARACTER DEFORMATION METERS ZONE LITHOLOGY LITHOLOGIC DESCRIPTION AGE Sel Sel LITHO. FORAMS NANNOS (Lapilli) Tuffs - (Ash) Tuffs Colors: greenish black (56 2/1) and gray (N2); Colors greenish olack (So 277) and gray (AC) angular to subangular fragments with poor sorting in lapilli zones, good in ash zones; graded beds in all sections; sizes 0.5-2 mm for ash, 0.1 to 1.5 cm for lapill; sharp contacts at fine/coarse graded beds. TS N2 TS * TS Units Section 1: 0-74, graded, poor sorting, angular fragments, lapilli tuff; (74-144), angular, poor sorting, ash tuff. Section 2: 0-150, as in lower of Section 1 -Tapilli tuff. Section 3: 0-22, graded ash-lapilli coarse base of Section 2; 22-73, new unit, well sorted ash tuff; 73-150, ash-lapilli, poor sorted, angular. TS ر الد ال Sorted, angular. Section 4: 0-88, as 73-150 in Section 3: 88-124, Tapilli tuff, poor sorting, angular; 124-150, ash tuff well sorted, top of unit in Section 5 - 0-68. TS Section 5: 0-68, ash tuff; 68-73, nanno chalk, 73-150, fair sorted ash tuff. Section 6: well sorted ash tuff. TS distentus \$ + , + 1 Thin Sections: 1-67, 1-79, 1-141, 2-46, 2-136, 3-20, 3-29, 4-97, 6-10 OLIGOCENE Composition (General) Fragments: Volcanogenic (15%) including: glass ŝ P. 21 pensis porphyritic basalt vesicular basalt LATE glassy basalt N2 pumice cipe Mineral fragments (30%) including: plagioclase thus TS pyroxene pheno11 oxides amphiboles 5G 2/1 Foraminifera and carbonate Matrix: brown glass - devitrified and palagonitized (55%). Rp-NG Many fragments with diffuse boundaries with matrix. ASH CLAY-RICH NANNOFOSSIL CHALK N2 Smear: CC Texture 65% Silt Composition 36% Nannofossils 35% Clay 25% Volcanic glass 25% Clay minerals 10% Feldspar 3% Micarb 5G 2/1 1% Foraminifera Cp-5G 2/1 Core CC Catcher N5

Explanatory notes in chapter 1

Hole.

ite	296	Hol	e		C	ore 5	7	Cored In	nter	val:	709.5-719.0 m		Site	296	- 1	ole		Co	ore 58	Cored	inter	val:	47.5-757.0 m	
AGE	ZONE	FORAMS	FOSS	IL CTER SUP	DIATONS	METERS	L	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION	AGE	TONE		PUKAMS PUCAMS	SSIL RACTER SOLUS	DIATOMS	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE		LITHOLOGIC DESCRIPTION
LATE OLIGOCENE	Sphenolithus distentus		Cp-	В	A Contract of the second se	0.5- 1.0-		volo-		* 55 *CC	N2 - N3 - N2 -	(Lapilli) Tuff - (Ash) Tuff <u>Section 1</u> : 0-77 cm, grayish black (N2) subround-subangular clasts; clasts > matrix with grain to grain contact: clasts are volcanic plus some carbonate and foraminifera; 1 m - 1.5 mm average sizes up to 10 mm = lapilli. <u>Sections 1-2</u> : 77-150 cm to 60 cm of Section 2, dark gray (N3) volcanic fragments; lighter color due to less clast % and lighter color matrix; clast size increase to maximum = 10 mm; 1 mm - 10 mm = range of sizes. <u>Sections 2-3</u> : 60 cm Section 2 to 25 cm <u>Sections 3</u> , similar to Section 1, 0-77 cm. <u>VOLCANIC SILTSTONES/SANDSTONES</u> Sharp contact at 25 cm with volcanic sand- stone - siltstones. Unit at + (25 cm to 150 cm); medium dark gray (N4) and exhibits grading <u>coarse</u> + <u>fine</u> sharp with calcite stringer + <u>fine</u> sharp coarse (fine-very fine sand) Smear: 3-55, CC	LATE OLIGOCENE	Sphenolithus distentus		Ар	в	1 2 3	0.5			* TS TS	N3 + N5 N3-N5	VOLCANIC SANDSTONE/SILTSTONE Dark gray (N3) to medium gray (N5); bedding is of dark and medium colors: bedding thick- ness 0.5-1.0 cm to ±9 cm in Sections 2-3; some irregular thicknesses. Volcanogenic fragments subangular to subrounded; graded units apparent throughout. <u>coarse (N3)</u> sharp t 0.2 mm+1 mm <u>coarse (N3)</u> sharp increases t coarse (N3) <u>Thin Sections: 2-32, 3-16, sand of plagioclas</u> (f) palagont 6 20), provene (5), amphiboles (5), and opaques (5) in silt-clay size matrix (20) of palagont te clay, silt-size volcanic materials and clay minerals.
												Composition 65-74% Volcanic glass 15% Feldspar 10-15% Clay minerals	Exp [*] Site	anato 296	ry not H	es ir ole	chapt	er 1 Co	re 59	Cored 1	nterv	/al: 7	85.5-795.0 m	
Expla	natory	notes	in	chap	ter	1			-			1% Nannofossils	AGE	ZONE		FO: CHAI SON	SSIL RACTER	SECTION	METERS	LITHOLOGY	EFORMATION	TH0.SAMPLE		LITHOLOGIC DESCRIPTION

(Lapilli) Tuff 35-55 cm: salt-pepper volcanic tuff; medium light gray (N6) to dark gray (N3); subround -white (medium light gray) pumice fragments in dark ash matrix; sizes - 3-5 mm average -to 20 mm - lapilli. N6-N3 5GY 6/1, N7, N6 N4-N3 Volcanic Siltstone 55-78 cm: greenish gray (56 6/1), light gray (N7); medium light gray (N6); extremely fine aphanitic ash with 0.2-0.5 mm white pumice fragment layers.

Lapilli Tuff 78-150 cm: dark gray (N4) - medium dark gray colors; subangular to angular - close packed; darker colors show size increase to 2-3 mm average to 15 mm maximum.

Volcanic Silstone Thin Sections: 1-60, 1-67, silt size frag-ments of glass (fresh), amphiboles; plagio-clase, lithic frags; angular shapes, well sorted; few large lithic frags. Clay minerals (155).

Explanatory notes in chapter 1

Rp-

Sphenolithus distentus

LATE OLIGOCENE

VOID Ø

Ø Core Catcher

0

* TS-*TS

			FOS	SIL	Rz			ION	BLE		
AGE	ZONE	FORAMS	NANNOS	RADS	STLICO. DIATOMS SECTIO	METER	LITHOLOGY	DEFORMAT	LITHO.SAM		LITHOLOGIC DESCRIPTION
				в	1 Ca	0.5 1.0	VOID VOID	N3		N2 N3 N4 \$ 5G 6/1	<pre>(Lapilli) Tuff - (Ash) Tuff Section 1: 0-106 cm, continuation of Core 59, Section 1: 78-150 cm, dark gray (N3) to grayish black; angular, to subangular and poorly sorted; sizes - 0.2 - 15 mm (average 2-3 mm); grain to grain contacts with clasts > matrix. 106-150: medium dark gray (N4); better sorted; large % greenish gray (56 6/1) fragments; subround; clasts > matrix with size. Average = 1 mm to 10 mm - ash to lapilli. Coarse at 150 → finer at 106 cm (Section 1).</pre>

			FOS	SIL	ER	2			ION	PLE		
AGE	ZONE	FORAMS	NANNOS	RADS	SILICO. DIATOMS	SECT10	METERS	LITHOLOGY	DEFORMAT	LITHO.SAM		LITHOLOGIC DESCRIPTION
		4				1	0.5	VOID			N7 N2	(Ash) Tuff <u>Sections 1-2</u> : 62 cm, light gray (N7) and grayish black (N2) which continues into Section 2 to 62 cm; pumice fragments, glass, lithic fragments, scoria, mineral fragments in an altered glass matrix; poor sorting; size (1-2 mm average); subangular-angular.
						2		Ŵ			N7 + N2 - N6 - N7	Section 2: 62-120 cm, VOLCANIC SILTSTOME, modium light gray (NG) with some dark gray (N3) areas; bedding, flow structures including load costs; cut and fill, slump, pebble drop structures. 120-150 cm: same as 0-62 cm and continues to 100 cm Section 3: <u>0-62 cm</u> : same as Tuff (Lapilli) (N7-N2) 100-120 cm: Volcanic Siltstone - ripple marks 120-132 cm: Iuff (Lapilli) (N7-N2).
						3	the state of the s				N2 N6 N3 N4	134-140 cm: Tuff (Lapilli) dark (N3) 140-150 cm: Volcanic Siltstone
				B		c	Core atche				- N7+N2	

Explanatory notes in chapter 1

Site 29	6	Ho1	e	CTI	_	Cor	e 62	Cored In	ter	al: 1	899.5-909.0 m
AGE	ZONE	ORAMS	SONNAN	ACTE	STLTCO.	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION
				В		1 2 Cot	1.0 0.5 1.0 0.5 1.1 1.0 0.5 1.0 0.5 1.0 0.5 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0	VOID			Lapilli Tuff - (Ash) Tuff - Volcanic Silt/ Sandstones, grayish black (N2) - medium gray (N5); poor sorting with altered glass matrix or a light color; contacts of grains poor; a few large fragments are pumices, scorias. N2 Graded Units: Section 1 - 74-150 cm Section 2 - 0-100 cm 100-120 cm = Vol. silt/sandstone 120-134 cm 134-139 cm N5 139-150 cm = Vol. siltstone

Explanatory notes in chapter 1

Τ			FOS	SIL	ER				ION	FLE	
AGE	ZONE	FORAMS	NANNOS	RADS	STLTCO.	SECTIO	METERS	LITHOLOGY	DEFORMAT	LITHO.SAM	LITHOLOGIC DESCRIPTION
						1	0.5	V01D		* TS	(Lapilli) Tuff - (Ash) Tuff <u>Section 1</u> (130 cm) to <u>Section 4</u> (105 cm): medium gray (N5) to dark gray (N3); subrou subangular volcanic clasts in devitrified altered glass; clasts > matrix; poorly sor (0.5-30 mm), Average =2 mm. Volcanic clasts: pumice, lithics, mineral, glass fragments; some grading and imbricat <u>Section 4</u> : 105-133 cm, greenish gray (56 6 and medium dark gray (N4); fine grained
E OLIGOCENE	thus distentus					2					<pre>version of above (Average 1-2 mm); better sorting. <u>Section 4: 133-150 cm, sharp contact with above; medium dark gray (N4) and dark gray N5 (N3). * <u>Thin Sections: 1-1]8: Volcanic Siltstone, fragments: pumi</u></u></pre>
LAT	Spheno 11					3	tituli in tituli			040	 (40), fresh glass (10), porphyritic basalt (5), piagioclase (10), pyroxene (5), opaqu (2). Matrix: glass and small pumice fragme (28) and fine clay minerals. 4-130: (Lapilli) Tuff, fragments of pumice (25), porphyritic basalt (5), variolitic basalt (10), serpentinzed fragments (2), plagioclase (25), pyroxenes and oxides (3) in devitrified to palagonitized matrix (30)
						4	11111111111111	- 0		* TS	N4-N3

N5-N3

Explanatory notes in chapter 1

Core



SITE

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Explanatory notes in chapter 1















SITE 296












































SITE 296





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