The Shipboard Scientific Party1

SITE DATA

Date Occupied: 13 December 1973 (0757Z)

Date Departed: 16 December 1973 (1300Z)

Time on Site: 77.05 hours

Position:

Latitude: 14°49.63'S Longitude: 146°51.51'W

Water Depth (sea level: 2641.0 corrected meters, echo sounding

Water Depth (rig floor): 2656.8 corrected meters, echo sounding

Bottom Felt at: 2659.0 meters, drill pipe

Penetration: 745.0 meters

Number of Holes: 1

Number of Cores: 32

Total Length of Cored Section: 298.5 meters

Total Core Recovered: 147.1 meters

Percentage Core Recovery: 49.3%

Oldest Sediment Cores:

Depth below sea floor: 745.0 meters Nature: Foraminiferal volcanogenic sandstone Age: Early Eocene Measured velocity: 2.2 km/sec

Basement: Not reached

- **Principal Results:** The site was drilled in order to determine if the northwest part of the Tuamotu chain is coeval with the Hawaiian-Emperor bend and to study the history of reefal growth and erosion in this area. A simplified geological history at the site may be given as follows:
 - (1) Eruption of basaltic edifices on older oceanic sea floor at some time prior to 49-51 m.y.B.P.

(2) Deposition of volcaniclastic sandstones and siltstones of shallow water origin at rates averaging 65-70 m/m.y. as these edifices were eroded.

(3) Formation of reefs, at least as old as 49-50 m.y.B.P.
(4) Pelagic sedimentation from early Eocene time to the present, with at least three interruptions. Floods of reefal debris entered the basin as turbidite units during middle Eocene and early Miocene time.



Figure 1. Location of Site 318 plotted on a portion of the bathymetric chart of the South Pacific of Mammerickx et al. (1973).

Inasmuch as basaltic basement rocks were not drilled, we can give only minimum ages of edifice construction. Nevertheless, even minimum ages of 49-51 m.y.B.P. are slightly older than recent estimates of the age of the Hawaiian-Emperor bend. Comparison with other ages in the Tuamotu chain suggests that volcanism has been episodic, like that of the Hawaiian chain. Reflectors in the basin are irregular and discontinuous; this irregularity may be due to influx of sediments from ridges bounding it. Comparison of reflectors with drilling results at this site is complex but suggests that the 0.56 to 0.60 (underway) and 0.68 (sonobuoy) sec reflectors, represent an unconformity in the Eccene at a depth of approximately 530 meters and the top of very hard limestones at a depth of about 620 meters, respectively. We continued evaluation of the heave compensator and continued to conclude that its use results in no better core recovery, nor less drill-induced deformation in cores than conventional drilling. Further comparative testing in continuously hard rocks seems indicated.

BACKGROUND AND OBJECTIVES

The original objectives of Site 318 (Figure 1) were (1) to determine the date of cessation of volcanism of near the bend of the putatively continuous Tuamotu/Line Islands seamount chain; (2) to study the petrology of the volcanic rocks that have built the principal Tuamotu edifices; and (3) to attempt to decipher the geological history of the main group of major atolls

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on the Tuamotu Ridge crest by a study of the sedimentary, presumably turbidite, debris that lies above volcanic rocks. The age of volcanism in this area has direct bearing on Morgan's (1971, 1972) suggestion that this particular area marks a bend in a single Tuamotu/Line Islands chain; a bend that he suggested records a major change in the direction of motion of the Pacific plate. Morgan's bend would therefore be similar to, and coeval with, the bend in the Hawaiian-Emperor chain. Whether the nature of formation of Hawaiiantype chains in linear (Morgan, 1971, 1972) or episodic (Shaw and Jackson, 1973), the age of bends in such chains should be the same according to hypothesis. Clague and Dalrymple (1973) have estimated the Hawaiian-Emperor bend age as about 42 m.y.B.P., although more recent data from Larson, Moberly, et al., (1973), if correct, would lead to estimated bend ages for that chain as being somewhat older. Subsequent to our own drilling in the Line Islands, we believe serious doubt has been cast on the hypothesis that a Line Islands/Tuamotu connection exists; nevertheless, the Tuamotu chain itself has characteristics that resemble the Hawaiian chain (Dana, 1876; Jackson and Shaw, 1975), and it now appears equally important to determine whether or not the birth of the Tuamotu chain coincided with that of the Hawaiian chain. Existing age dates in this area are sparse. Deep Sea Drilling Site 76 (Haws et al., 1972) recovered Eocene reefal debris in Pliocene turbidites along the northern flank of the Tuamotu chain, about 175 km north of our proposed site, but only 27 meters of penetration was achieved. Dredge hauls to the northwest of our proposed site yielded fossil ages in the range 37.5-43 m.y. (Burckle and Saito, 1966) and 37.5-53.5 m.y. (Repelin, 1919), but in terms of basement ages, none of the above data is in any way definitive.

In addition to our interest in basement ages at this site, the Tuamotu Islands in this area form one of the most extensive atoll complexes on earth, and it was hoped to reconstruct a history of atoll formation in this area by drilling reefal debris that presumably lies above basement.

The site considered most likely to meet these objectives was chosen on the *Conrad*-11 refraction profile at 1800, 4 September 1967 (Figure 2) at a position of about 15.2°S; 146.8°N, where a ridge was observed between the major Tuamotu atolls in this area that appeared to contain ponded sedimentary rocks in its basined upper surface. Water depth at this site was measured as about 2600 meters; depth to acoustic basement appeared to be about 0.75 sec.

OPERATIONS

Predrilling Site Surveys

The original site was selected on the basis of a *Conrad*-11 seismic profile that crossed the Tuamotu Ridge on 14 September 1967 (Figure 2). The bathymetry indicated that the *Conrad* line crossed a basin that could have been receiving sediments from five atolls, including Rangiroa, one of the largest in the Tuamotu group.

Glomar Challenger approached Site 318 by attempting to intersect the *Conrad* line before it crossed the north scarp of the Tuamotu Ridge, northwest of Ahe Atoll (Figure 3). This approach was made inasmuch as we were not certain of the location of the Conrad line, except that it passed close by Ahe Atoll and passed over the south scarp of the ridge between the Rangiroa and Arutua atolls (Figure 3). In this manner we planned to cross the grain of the structure and be able to navigate on the submarine morphology. The assumed intersection with Conrad-11 was made on a course of 100° True and a course change to 215° True was made at 1540 (local), 12 December; shortly afterwards (0140Z, 13 December) Ahe Atoll was picked up on the radar at 16 n. mi. distance, bearing 140°. The north scarp of the Tuamotu Ridge was crossed as well as the perched basin shown on the Conrad-11 line (compare Figures 2 and 4). On crossing a high ridge, presumably the northern limit to the trough that we intended to drill, we changed course to 196° True and crossed to the major basin that lies between raised edifices of the Tuamotu Ridge. On the Conrad-11 line this basin is 2 hr wide. After steaming for ~ 1 hr and 15 min, the beginning of another ridge, presumably the one leading up to Rangiroa, appeared. We decided that we were actually east of the Conrad-11 line instead of west-the basin being narrower along our line-course was reversed (Figure 5) and we steamed back to a suitable section and dropped the beacon at 2157, 12 December (0757Z, 13 December).

The PDR depth of 1416 fathoms at Site 318 was corrected to 2641 meters (Matthew's tables, Area 41), giving a derrick floor to mudline depth of 2656.8 meters. The bottom proved to be solid in this area, and the drill pipe depth was 2659.0 meters; the latter depth was used for the site.

Reflectors on both the *Conrad*-11 and *Glomar Challenger* lines are irregular and show rapid lateral changes in thickness and strength (Figures 2 and 4). On the *Conrad*-11 line reflectors were noted at 0.08?, 0.20, 0.40, 0.56, and 0.75 sec. On the *Glomar Challenger* line reflectors at roughly 0.1, 0.20, 0.40, 0.60, and 0.75 were noted. The deeper reflectors are quite vague on both lines.

Sonobuoy Survey

Inasmuch as current directions in the general area between the Tuamotu atolls were known to be weak and irregular, a sonobuoy was released at 0535Z, 13 December prior to beacon drop, while underway to the site (see Figure 6); recording on EDO 2 (4-sec sweep). A prominent reflector at 0.69 sec corresponds to the deepest reflector, at approximately 0.71 sec on the profile (Figure 4). Details of the sonobuoy survey are given below.

Drilling Program

Because of the very limited time available at Site 318, the most reasonable drilling strategy appeared to be (1) to take a punch core at mudline to establish the age and erosional character of the surface; (2) to retrieve only every third core drilled, washing the two intermediate intervals until such time as this became impossible due to increasing rock coherence; and (3) to core continuously once such hard formations were reached. We followed this program through Core 21, but, in washing the first and second interval following Core 21,



Figure 2. LDGO Conrad-11 seismic reflection profile showing physiographic features used in approach to Site 318.



Figure 3. Location of Site 318, with Conrad-11 and Glomar Challenger Leg 33 tracks. Bathymetry from Mammerickx et al. (1973).

washing times became long. We were advised by the Drilling Superintendent that we might have a jammed core catcher. We therefore pulled Core 22 after only one washing interval, and pulled Core 23 immediately following, in response to the Cruise Operations Manager's request for a core at that time. Subsequent to Core 23 we washed as drilling time dictated (in general, washing two cores for one if drilling times were 45 min or less, one for one if drilling times were 1 to $1\frac{1}{2}$ hr, and taking every core that required as long as $1\frac{1}{2}$ hr to drill). Penetration rates became very slow subsequent to Core 21, and the drilling and coring were finally terminated in hole at 2045 (local), 15 December 1973, in order to meet our arrival schedule in Papeete. Bit wear at that time seemed nominal; the final cores were nearly full diameter.

Heave Compensator

Inasmuch as we had accepted the responsibility for a geologic evaluation of the heave compensation device, and since it had previously been operated only in Tertiary rocks, and then with somewhat ambiguous results (see Site 317 report, this volume), we agreed to leave the compensator in the drill string at all times at Site 318, but to take alternate cores with the device locked out and operating throughout the section cored. Thus, in the early portion of the hole 10 even-numbered cores were taken with heave compensation, 10 odd-numbered cores without. During the drilling of Core 22, the Cruise Operations Manager ordered that the heave compensator be left operating for the remainder of the drilling at this site; as a result we are unable to compare results beyond Core 20. Core recovery in holes drilled with the compensator locked out (odd-numbered Cores 1-19) was sensibly the same as recovery in the cores with the compensator operating (even-numbered Cores 2-20). In the former cores recovery averaged 52.9%; in the latter, 53.2% (see Table 1). Recovery in soupy oozes and rocks interbedded with cherts was poor in both modes; firm oozes and chalk gave good recovery in both modes. Recovery time was somewhat slower with the heave compensator operating than without it, no doubt because of its operational complexities (see Site 317 report); the time from delivery of an even core on deck to the arrival of an odd one totaled 11.9 hr versus the



Figure 4. Glomar Challenger seismic reflection profile on approach to Site 318, showing physiographic features used in locating the site (see Figure 2).

opposite situation, where the time of delivery of compensator-drilled holes totaled 13.8 h (see Table 1). Deformation was somewhat less severe in cores drilled without the heave compensator (about 24% were intensely deformed) than in cores drilled with it (about 27% were severely deformed).



Figure 5. Glomar Challenger track line on approach to and departure from Site 318.



Figure 6. Glomar Challenger sonobuoy record obtained while underway on approach to Site 318.

Indeed, we are unable to amend our previous evaluation—it may be that the heave compensator installed on *Glomar Challenger* keeps the bit more firmly in the bottom of the hole than does the conventional drill. If so, returned cores drilled with the compensator show no improved effects, better recovery, time saving, or less severe deformation than those cut with the conventional drill. It is a pity that we were unable to compare the performance of the unit directly in basalt, inasmuch as it was planned for Leg 34 to use it largely in that rock type. On other rock types equally difficult to core, it

SITE 318

TABLE 1 Coring Summary

1 2 3 4 5	13 Dec 13 Dec 13 Dec 13 Dec 13 Dec	0555 0740	2659.0-2666.5				1101
2 3 4 5	13 Dec 13 Dec 13 Dec 13 Dec	0740		0-7.5	7.5	7.5	100
3 4 5	13 Dec 13 Dec 13 Dec	0045	2695.0-2704.5	26.5-36.0	9.5	2.0	21
4 5	13 Dec 13 Dec	0845	2723.4-2733.0	64.5-74.0	9.5	1.5	16
5	13 Dec	0945	2752.0-2761.5	93.0-102.5	9.5	9.3	98
	10 2000	1055	2780.5-2790.0	121.5-131.0	9.5	9.2	97
6	13 Dec	1205	2809.0-2818.5	150.0-159.5	9.5	9.5	100
7	13 Dec	1310	2837.5-2847.0	178.5-188.0	9.5	9.5	100
8	13 Dec	1420	2866.0-2875.5	207.0-216.5	9.5	4.5	47
9	13 Dec	1530	2894.5-2904.0	235.5-245.0	9.5	2.4	25
10	13 Dec	1650	2923.0-2932.5	264.0-273.5	9.5	5.0	53
11	13 Dec	1810	2951.5-2961.0	292.5-302.0	9.5	9.5	100
12	13 Dec	1930	2980.0-2989.5	321.0-330.5	9.5	2.9	30
13	13 Dec	2105	3008.5-3018.0	349.5-359.0	9.5	2.1	22
14	13 Dec	2230	3037.0-3046.5	378.0-387.5	9.5	4.4	46
15	13 Dec	2355	3065.5-3075.0	406.5-416.0	9.5	2.7	28
16	14 Dec	0120	3094.0-3103.5	435.0-444.5	9.5	3.6	37
17	14 Dec	0250	3122.5-3132.0	463.5-473.0	9.5	2.0	21
18	14 Dec	0445	3151.0-3160.5	492.0-501.5	9.5	2.5	26
19	14 Dec	0625	3179.5-3189.0	520.5-530.0	9.5	1.9	20
20	14 Dec	0750	3208.0-3217.5	549.0-558.5	9.5	7.0	74
21	14 Dec	0955	3236.5-3246.0	577.5-587.0	9.5	3.5	37
22	14 Dec	1405	3255.5-3265.0	596.5-606.0	9.5	5.0	53
23	14 Dec	1625	3265.0-3274.5	606.0-615.5	9.5	3.2	34
24	14 Dec	2025	3274.5-3284.0	615.5-625.0	9.5	6.0	68
25	14 Dec	2245	3284.0-3293.5	625.0-634.5	9.5	4.7	49
26	15 Dec	0100	3293.5-3303.0	634.5-644.0	9.5	6.2	64
27	15 Dec	0310	3303.0-3312.5	644.0-653.5	9.5	2.4	25
28	15 Dec	0645	3322.0-3331.5	663.0-672.5	9.5	3.4	36
29	15 Dec	1020	3341.0-3350.5	682.0-691.5	9.5	2.0	21
30	15 Dec	1430	3360.0-3369.5	701.0-710.5	9.5	4.5	47
31	15 Dec	1800	3379.0-3388.5	720.0-729.5	9.5	2.1	22
32	15 Dec	2045	3398.0-3404.0	739.0-745.0	6.0	5.1	93
Total					298.5	147.1	49.3

appears to have little other effect than the conventional bumper sub that is removed to allow its operation. In washing cycles it continues to be detrimental because of time loss.

Departure

The Captain of Glomar Challenger estimated 26 hr steaming time for the 230 n. mi. from Site 318 to an 0600 rendezvous 17 December with the Papeete pilot boat; the Drilling Superintendent estimated 9 hr were required to pull the string and stow the heave compensator (the bottom-hole assembly had been magnafluxed at Site 317). We therefore planned to have our last core from this site on deck at 1900, 15 December. The final core (Core 32) was on deck at 2040, 15 December. The string was pulled and Glomar Challenger got underway at 0300 (local), 16 December on a course of 020°, turned back, passed over the beacon and headed for the pass between Rangiroa and Arutua atolls on a course of 205°. The profiler record (Figure 7) shows the southern edge of the drilled basin, confirming that we drilled near the Conrad-11 line.

LITHOLOGIC SUMMARY

The stratigraphic column at Site 318 can be subdivided into five lithologic units (Figure 8):

Unit 1 (0 to between 35.5-64.5 m)—Nannofossil-foraminifer firm ooze that contains some graded beds with shallow water skeletal debris of foraminifers, bryozoans, echinoids, corals, mollusks, and altered volcanogenic material.

Unit 2 (between 35.5-64.5 and between 245 and 264 m)—Foraminifer and nannofossil firm ooze, changing downward to soft chalk, with a few graded beds with sand-sized grains of volcanogenic and shallow water biogenous origin, and a bed of breccia with basalt, chert, and bioclastic debris.

Unit 3 (between 245-264 m and between 416-435 m)—Foraminifer nannofossil chalk, becoming harder with depth, and with chert nodules in the lower part, below 379.5 meters. Thin laminae of sand with rare volcanogenic grains.

Unit 4 (between 416-435 m and between 530-549 m) —Yellowish and greenish nannofossil and foraminifer limestone with chert nodules, clayey laminae and graded beds with volcanogenic and shallow water skeletal debris.

Unit 5 (between 530 to 549-745 m)—Green and greenish-gray clayey limestone, and volcanogenic siltstone and sandstone, mainly as graded beds. Shallow water skeletal debris is abundant in some sandstone layers.

Unit 1—Ooze (0-35.5 m)

Pale orange nannofossil-foraminifer ooze and foraminifer ooze. Interbedded with these calcareous oozes are graded layers 20 to 60 cm in thickness containing sand- to granule-sized grains of volcanogenic and carbonate debris. This debris includes palagonite, greenish



Figure 7. Glomar Challenger seismic reflection profile on departure from Site 318. Annotated reflectors are correlative with those on the Conrad-11 profile shown in Figure 13.

glass, rock fragments, and, in some layers, shallow water skeletal fragments of bryozoans, mollusks, and large and small foraminifers including *Amphistegina* and *Heterostegina*. The volcanogenic and bioclastic material is mainly concentrated at the base of graded beds; higher parts of the beds consist chiefly of planktonic foraminifers and nannofossils. In thin section, some layers have a packstone texture, with mainly globigerine foraminifers, as whole tests, and a few globorotaliids in a matrix of nannofossils. Other layers have wackestone textures, with about 50% nannofossils and 50% planktonic foraminifers. The foraminifers are either empty or filled with micrite matrix. (The terms wackestone and packstone are used according to the classification of Dunham, 1962.)

The ooze is vaguely mottled in color shades both slightly lighter and darker than the pale orange host sediment.

As shown by X-ray data (Cook and Zemmels, this volume), some of the transported fossils probably con-



Figure 8. Graphic log showing lithologic units, carbonate content, and drilling rates at Site 318.

tain aragonite in their shells. About 1% dolomite and a trace of barite were also detected by X-ray diffraction.

Much of the ooze has the firmness of packed sand, and in Core 2, Section 2, near the base of the unit, the sediments have the induration of chalk.

Unit 2-Ooze and Chalk (64.5-265.5 m)

Foraminifer-nannofossil firm ooze, changing downward to soft chalk. Some layers are foram-rich and contain rare sand-sized grains of palagonite. A few of these layers show graded bedding. Colors range from white at the top of the interval to bluish-white and greenishwhite in Cores 6, 7, and 8, and white again near the base of the unit in Core 9. Thin wisps and streaks of gray ooze occur throughout. The foram-rich layers tend to be shades of pale yellowish-gray. At the base of the unit is a thin bed of well-indurated breccia, consisting of pebbles of basalt, chert, and shallow water skeletal debris (Figure 9).

The foram-rich layers in this section have a packstone texture with sand-sized planktonic (99%) and benthonic (1%) foraminifers. Also present in these packstones are very rare fragments of bryozoans, tiny calcareous sponge spicules and grains of palagonite, and feldspar, all set in a matrix of nannofossil micrite. The finer grained sediments, in thin section, have wackestone textures, some with a wide range of sizes of planktonic foraminifers (40-600 μ m) set in a matrix of nannofossil micrite. The ratio of matrix to foraminifers ranges from about 1 to 2. Most foraminifers are filled with micrite; some are completely empty. In the lower part of the unit, some of the thinner walled globorotaliid foraminifers have slightly crushed tests.



0.5 cm

Figure 9. Lower Miocene, well indurated, brownish polygenetic breccia from the base of lithologic Unit 2. Diverse fragments include rounded aphyric vesicular basalt clasts with ferruginous halos, gray angular chert clasts, and light yellow skeletal and reef debris (Core 10, Section 1, 137-140 cm; 268 m below the sea floor).

X-ray data (Cook and Zemmels, this volume) reveal the presence of small amounts of dolomite in gray ooze in Cores 4 and 5, and barite in Cores 4 and 7.

The breccia bed at the base of the unit consists of angular fragments, up to 1 cm in diameter, of a wide variety of clasts. The proportions of clast types are roughly 50% volcanogenic, 40% sedimentary, and 10% biogenic clasts. Volcanogenic clasts consist of basalt, with plagioclase of An₄₀₋₄₅ composition and augite phenocrysts (15%) in a sideromelane or palagonite groundmass; quartz syenite, containing quartz, orthoclase, minor plagioclase, and aegerine augite; trachyte, composed almost entirely of fine-grained sanidine; oligoclase andesite; holocrystalline basalt, with subcalcic augite microphenocrysts and groundmass pigeonite; palagonite, generally of the gel variety, but also including fibrous types with abundant spherulites; bioclastic debris is composed of mollusks, benthonic and planktonic foraminifers, bryozoans, and sedimentary rock fragments of silicified foraminifer-nannofossil limestone.

Unit 3—Varicolored Chalk, Chert, and Sandstone (265.5-416.0 m)

Foraminifer-nannofossil chalk in shades of yellowishgray, yellowish-white, pale orange, grayish-orange, pinkish-white, and white with occasional wisps, streaks, laminae, and burrow mottles of light olive-gray and light gray. The induration ranges from soft chalk at the top of the unit to hard chalk at the base. Much of the chalk has been fractured and crushed in the coring process, reducing it to a paste in some intervals. Chert, in the form of olive-gray and yellowish-brown nodules occur very sporadically in the lower part of the unit, beginning at 379.5 meters in Core 14.

Within the chalk are occasional thin laminae of foram-rich sand, with about 1% volcanogenic grains of pale green palagonite and minor feldspar. In thin section, the sands have a packstone texture, with grains of both planktonic (~90%) and benthonic (~10%) foraminifers in a matrix of nannofossil micrite. These tests are commonly fragmental and partly dissolved and recrystallized.

X-ray data show barite and analcite in addition to volcanogenic components in the acid-insoluble fraction.

The bulk of the chalk is wackestone, with planktonic foraminifers of a wide range of sizes (40-500 μ m) in a nannofossil micrite matrix. The micrite is commonly charged with tiny calcareous sponge (?) spicules, lying in the bedding plane.

Unit 4-Limestone and Chert (416.0-530.0 m)

Yellowish-gray and pale greenish nannofossil and foraminifer-nannofossil limestone, with common chert nodules. The unit is characterized by darker, more greenish colors than Unit 3 and by the prevalance of laminations, burrows, and mottles. Near the top of the unit these burrows are dark and are set in a lighter host, but lower down the rule is lighter burrows in a darker host. Burrows are generally somewhat flattened parallel to bedding.

Evidence of soft sediment deformation is common, especially in Core 16. Recumbent folding and tiny brittlefracture step faults occur in Section 2 of Core 16.

Streaks and laminae of foram-rich packstone with greenish palagonite grains occur scattered throughout the unit. These are more numerous, thicker, and coarse grained toward the base of the unit, where they generally form the lower parts of graded beds. The upper parts of these beds consist of nannofossil-foraminifer wackestone with decreasing proportion of foraminifers (mainly planktonic) upward. The limestone is layered on a scale of about 250 μ m, marked by changes in foraminifer concentrations in the matrix. Sponge spicules and radiolarians are rare in the upper part of the unit, but become more common near the base. In Core 18, radiolarians are almost as abundant as foraminifers in one thin section.

Chert nodules are scattered throughout the unit, at a frequency of 1 or 2 per meter.

The base of the unit is chosen at the bottom of Core 19, in which an artificial sand or drilling breccia was recovered. This breccia consists of volcanogenic grains (mainly palagonite), chert, and bioclastic debris (planktonic and large benthonic foraminifers including Nummulites, Heterostegina, Cycloclepeae?, and Asterocyclina, bryozoans, calcareous algae, corals, and mollusks).

Unit 5—Limestone and Volcaniclastic Sediment (549.0-745.0 m)

Green and greenish-gray clayey limestone, siltstone, and sandstone, mainly as graded beds. Hard and cherty

in the upper 100 meters (Cores 20 through 24), slightly softer below. The graded beds, which average between 5 and 10 cm thick, generally have siltstone or sandstone at the base, and pass upward into increasingly burrowmottled limestone and clayey limestone. The colors become increasingly darker towards the base of the unit, except for yellowish-gray silty and sandy layers rich in ash or palagonite. The proportion of CaCO₃ decreases regularly toward the base of Unit 5 (see Figure 8). Soft sediment deformation features, such as small step-faults, deformed burrows, inclined beds, and overfolds, are common at many levels within Unit 5.

In thin section, the sand layers have a packstone texture and are seen to consist of mixtures of biogenous and volcanogenic grains with a nannofossil-clay matrix. Most of the sand layers in the upper part of the unit consist chiefly of planktonic foraminifers, with lesser amounts of shallow water bioclastic debris (tests of large foraminifera including Asterocyclina, Nummulites, and Heterostegina, bryozoans, echinoid spines, calcareous algae) and volcanogenic materials (green and brown palagonite, basalt fragments, and feldspars, including both plagioclase and anorthoclase; see Figure 10). In the lower part of the unit, in Cores 30 through 32, the proportions are reversed; these sediments are volcanogenic sandstones, with occasional laminae or beds rich in planktonic foraminifers or shallow water bioclastic debris (especially discocyclinids). Some layers (e.g., Core 30, Section 3, 58-60 cm) consist almost entirely of slightly altered splintery volcanic glass shards. The



0.5 cm

Figure 10. Middle Eocene, light greenish-gray limestone from lithologic Unit 5. Basal part of a graded bed is composed of masses of discocyclinid benthonic foraminifers and a few scattered green palagonite grains (Core 26, Section 1, 137-140 cm; 636 m below the sea floor). micrite matrix of the sandstones and the upper, finer grained parts of graded layers are nannofossil rich in the upper part of Unit 5, changing gradually to montmorillonite clay rich in the lower part. X-ray data reveal montmorillonite, clinoptilolite, and volcanogenic minerals at all levels sampled in Unit 5.

GEOCHEMICAL MEASUREMENTS

The results of pH, alkalinity, and salinity analyses of interstitial waters are summarized in Table 2 and Figure 11. Nine samples were analyzed. The shipboard CaCO₃ content values for 21 selected sediment samples are shown in Figure 8. Procedures for analysis were those routinely performed aboard ship of *Glomar Challenger*.

pH Values

From 0 to 553 meters, pH values remain consistently below those for surface seawater. They range from 7.20 to 7.39 from 50 to 553 meters. The three methods used to measure pH show corresponding results. A slight increasing trend in Unit 5 may be noted in the lowest (553 m) measurement.

Alkalinity

From 0 to 438 meters, alkalinity values remain uniformly above those for surface seawater which is 8.21 at this site. The alkalinity values show little deviation in this zone, ranging from 2.93 to 3.42 meq/kg. There is a decreasing trend reflected in the lowermost measurement (0.59 meq/kg) at 553.0 meters, which corresponds to the change to lithologic Unit 5.

Salinity

From 0 to 438 meters, salinity values range from 35.5 to $36 \,^{\circ}/_{00}$ remaining near to, but lower than, the value of 36.3 for surface seawater at this site. Only the lowest measurement ($36.6 \,^{\circ}/_{00}$) at 553 meters in Unit 5 is slightly above that of the surface seawater value.

CaCO₃

All of the sediments in Units 1 through 4 have calcium carbonate values between 90% and 100%. Near 600 meters, in Unit 5, as the quantity of volcanogenic clay, siltstone, and sandstone increases, the calcium carbonate content concomitantly decreases (see Figure 8).

PHYSICAL PROPERTIES

The physical properties methods, presentation in hole and core plots, presentation in tables, and definitions are discussed briefly in the Physical Properties section of the Site 315 report and in detail in Appendix I, this volume, and therefore will not be discussed in detail here.

The GRAPE analog data are displayed in the core scale graphs only. Where the sediment was soft and the core liner completely filled, the analog GRAPE shorebased computer program requires no diameter corrections, and the data are plotted as a single solid line. Where the analog GRAPE scans cores of hard rocks with varying diameter, the data are presented as two lines. The solid line is the routine analog data without correction for core diameter and assuming a 6.61-cm

TABLE 2 Summary of Shipboard Geochemical Data

Sample	Depth Below	p	н	Colorimetric		Combination	Potentiometer
(Interval in cm)	Sea Floor (m)	Punch- in	Flow- through	Alkalinity (meq/kg)	Salinity (°/00)	Electrode pH	Alkalinity (meq/kg)
Surface Seawate	er	8.25	8.18	2.44	36.3	8.23	2.39
1-4, 144-150	6	7.85	7.53	3.03	35.8	7.80	3.06
4-4, 144-150	96	7.28	7.20	3.23	35.5	7.38	3.24
6-4, 144-150	156	7.28	6.99	3.42	35.5	7.28	3.38
8-1, 144-150	207.5	7.34	7.20	3.03	35.5	7.40	2.95
10-3, 144-150	268.5	7.30	7.16	3.03	35.5	7.40	3.02
12-1, 142-150	322.5		7.32	2.64	36.0	7.45	2.60
14-2, 144-150	381	7.20	7.19	2.93	36.0	7.35	2.79
16-2, 144-150	438		7.15	2.05	36.0	-	-
20-3, 144-150	553.5	—	7.39	0.59	36.6	—	-



Figure 11. Graphic log of shipboard pH, alkalinity, and salinity measurements at Site 318.

diameter. This line is presented in case it is felt necessary to consult photographs, measure diameters, and apply other diameter corrections for a discrete interval, or simply for the manipulation of data. The dotted line represents data to which diameter corrections have been applied as described in Appendix I, this volume. Where rock segments were very short, the data appear as a series of peaks; only the maximum density of the peak in these cases represents good density values, and the density values of the shoulders of the analog peaks should be ignored.

The GRAPE Special wet-bulk density, gravimetric wet-water content, porosity, sound velocity (perpendicular and parallel to bedding), absolute velocity anisotropy, percentage velocity anisotropy, acoustic impedance, and reflection coefficients are presented in Table 3, with most of these parameters, graphically displayed in the site plots and core plots.

The physical properties data are primarily presented in tables and graphs and only briefly discussed. Detailed discussion and interpretation about the interrelationships of the laboratory physical property data are treated more thoroughly in Chapter 26, this volume.

Results

Sound velocity, wet-bulk density, wet-water content, porosity, acoustic impedance, and reflection coefficients were measured or calculated from ooze, chalk, limestone, chert, and volcaniclastics, from 0 to 745 meters below the sea floor. The physical properties of the sedimentary rocks recovered at Site 318 are summarized into six stratigraphic groups based primarily on sound velocity, wet-bulk density, wet-water content, and impedance data as shown in Figure 12. The boundaries are indicated by the dashed horizontal lines, and listed in Table 3. This grouping is summarized in Table 4, which lists the lithology, depth interval, and typical values of the characteristic physical properties.

The characteristic physical property intervals at Site 318 are as follows: (1) Quaternary to Miocene foramnanno ooze and chalk from 0 to 265.5 meters below the sea floor; (2) Miocene to Oligocene foram-nanno chalk from 265.5 to 408 meters; (3) Oligocene to Eocene foram-nanno, nanno-chalk and limestone from ~408 to ~494 meters; (4) Eocene nanno limestone from ~494 to ~540 meters; (5) Eocene nanno limestone with minor interbedded clastics from ~540 to 581.5 meters; and (6) Eocene nanno limestone, clayey limestone, silicified limestone, and volcaniclastics from 581.5 to ~745 meters. The last three units are rather arbitrary and they could be possibly lumped. The last unit could also be arbitrarily subdivided below 700 or 715 meters.

Interval 1, from 0 to 265.5 meters, has variably low velocity wet-bulk density, and impedance with high water content and porosity. The lower boundary is marked by conglomerate and breccia with high velocity clasts.

Interval 2, from 265.5 to \sim 408 meters, has uniformly low velocity, wet-bulk density, and impedance, and the wet-water content and porosity are high. The lower boundary is marked by a small but significant increase in velocity, wet-bulk density, and impedance, and a corresponding decrease in wet-water content and porosity.

The third interval (3) from ~408 to ~494 meters has fairly uniform parameters, with moderate wet-bulk density and velocity. The lower boundary at 494 meters is marked by an abrupt change to relatively high velocity and relatively high wet-bulk density limestone from ~494 to ~540 meters. However, this fourth (4) interval (~494 to ~540 m) is based only on a few data points and may not exist in situ as represented here. The lower boundary of ~540 meters is distinctly marked by a significant and abrupt decrease in velocity, wet-bulk density, and impedance, with a decrease in porosity and wet-water content.

The fifth (5) interval from \sim 540 to 581.5 meters has moderate velocity, wet-bulk density, and impedance, probably the result of abundant clastics. Its lower boundary is marked by an abrupt increase in velocity, wet-bulk density, and impedance, and decreases in wetwater content and porosity.

The sixth and last group from 581.5 to ~ 745 meters is characterized by variably high to low parameters, which tend to decrease with increasing depth (porosity and wet-water content increase). The contrasting parameters in this interval are the result of interbedding of limestone, silicified limestone, clayey limestone, and volcaniclastics.

Discussion

The reflection coefficients calculated are only typical of what may be present at the arbitrary boundaries of major groups of physical properties. The few that are calculated are not any indication of the location or number of such impedance mismatches as are present.

Sound velocity has a tendency to be faster parallel rather than perpendicular to bedding. A few negative anisotropy values were calculated, but it cannot be precluded that these result from error. In general, anisotropy in the Quaternary to Eocene sediments and sedimentary rock above ~540 meters is very slight (0 to 2%), with an occasional high (6%) anisotropy in a few hard layers. Below ~540 meters the Eocene limestone and clastics have significantly higher anisotropy of 0 to 20% with 0 to 12% being typical. See Chapter 26, this volume, for more information.

CORRELATION OF REFLECTION PROFILE WITH DRILLING RESULTS

The reflectors in the basin drilled are not regular; they display lateral changes in strength and "thickness" across the basin (Figures 2 and 4). On the *Conrad*-11 (Figure 13) line, reflectors lie roughly at 0.08, 0.20, 0.40, 0.56, and 0.75 sec; on the *Glomar Challenger* line (Figure 7) reflectors could be discerned, after a fashion, at 0.20, 0.40, 0.60, and 0.75 sec. The deep reflectors are quite vague and are characterized by wavy and indistinct tops. An underway sonobuoy was run (Figure 6) and the only really distinct reflector seen lies at 0.68 sec (Figure 14).

Upon drilling, it was found that the upper reflectors (0.08, 0.20, and 0.40 sec) could not be recognized or correlated with any obvious lithologic breaks. The 0.56-sec reflector of the Conrad-11 line and the 0.60-sec reflector of the Glomar Challenger line may represent an unconformity in the section at approximately 530 meters. Using a mean velocity of 1820 m/sec from the mudline to 530 meters, a reasonable velocity based on the shipboard data (see Physical Properties section), the twoway travel time to the unconformity should be 0.58 sec, a value that falls between the Conrad-11 and Glomar Challenger 0.56-sec and 0.60-sec reflectors. The 0.68-sec reflector, or the 0.75-sec reflector of the Conrad profile (Figure 13), probably represents the shallowest occurrence of the very hard limestones encountered at approximately 625 to 670 meters depth (Figure 8). The velocity of these limestones is quite high; many are greater than 3 km/sec. If the 0.68-sec reflector lies at a depth of 625 meters, the interval velocity to that depth would be 1.84 km/sec. Figure 14 is an analysis of the sonobuoy record (see Figure 6 for a photograph of the record).

PALEONTOLOGY

Biostratigraphic Summary

Site 318 was cored intermittently between the mudline and 745 meters, except for a portion of the middle Eocene section (Cores 22 to 27; at depths of 596.5 to 653.5 m) which was cored continuously. Nannofossils are present throughout the cored intervals, providing reliable age control for virtually all of the material recovered. Foraminifers are abundant and well preserved in the middle Miocene through Quaternary, but are generally poorly preserved in the Eocene through lower Miocene. Radiolarians of sufficient preservation

Sample	Denth in	II Bade	Compression	nal Sound Ve Aniso	elocity tropy	Temn	"Spe Wet- Den: 2-mi cot (g/	cial" Bulk sity ^a nute int cc)	Wet- Water Content Salt	Porosityb	Acoustic Impedance g 10 ⁵	
(Interval in cm)	Hole (m)	(km/sec)	(km/sec)	(km/sec)	(%)	(°C)	Beds	Beds	(%)	(%)	cm ² sec	Lithology
2-2. 61-63 3-1. 18-21 6-6. 85-87 7-3. 14-18 7-4. 119-121 7-5. 106-108 8-1. 34-36 9-1. 106-108 9-2. 50-52 10-1. 105-107 10-1. 122-124 10-3. 50-52 11-1. 126-128 11-3. 45-47 12-1. 175-77 12-2. 11-13 13-1. 146-148 14-1. 41-43 15-1. 124-126	28.61 64.68 158.85 182.14 184.69 186.06 207.34 236.56 237.50 265.05 265.15 265.22 265.40 267.50 293.76 295.95 321.75 322.61 350.96 378.41 407.74	1.968 1.612 1.620 1.636 1.639 1.665 1.698 1.827 1.776 3.222 2.763 3.956 4.496 1.763 1.763 1.786 1.729 1.709 1.683 1.694	1.654 1.669 1.839 1.788 1.656 1.658 1.835 1.700 1.713 1.681 1.684 1.034	-0.015 -0.004 -0.012 -0.012 -0.012 -0.012 -0.049 +0.029 -0.004 +0.002 +0.002	$\begin{array}{c} - & 0.91 \\ - & 0.24 \\ - & 0.65 \\ - & 0.67 \\ \end{array}$ $\begin{array}{c} - & 2.67 \\ + & 1.71 \\ - & 0.23 \\ + & 0.12 \\ + & 0.59 \\ - & 0.25 \end{array}$	23.0 23.0 22.0 22.0 22.0 22.0 22.0 22.0	1.829 1.668 1.631 1.646 1.818 2.156 2.273 1.830	1.674 1.726 1.760 1.726 1.801 1.805 1.805 1.825 1.825 1.916 1.831 1.828 2.920	33.14 42.69 36.23 36.02 39.56 32.11 31.62 31.14 13.97 25.10 12.31 29.75 30.64 31.48 32.16 27.13 30.38 27.65 28.86 29.03	60.61 71.21 59.09 60.30 68.28 58.38 55.65 53.75 30.12 57.05 56.07 56.70 58.05 51.14 55.44 52.98 52.84 53.06	3.60° 2.69° 2.69° 2.69° 2.77° 2.88 3.09° 3.24 3.09 6.95° 6.28° 3.23† 2.98 2.99 3.46 3.10 3.28 3.08 3.08 3.08	Nanno foram chalk, disturbed lump Nanno foram chalk, disturbed lump Foram nanno firm ooze, disturbed lump Nanno foram firm ooze Nanno foram firm ooze Foram nanno chalk Foram nanno chalk Foram nanno chalk Nanno foram limestone Volcanic breccia Nanno foram limestone Basalt pebble Foram nanno chalk Foram nanno chalk Spicule-rich foram nanno chalk Spicule-rich foram nanno chalk
15-2, 106-108 16-1, 95-97 16-2, 77-79 16-3, 120-122 16, CC 17-2, 5-10 17-2, 26-28 18-1, 92-94 18-2, 46-48 18-2, 46-48	409.06 435.95 437.27 439.20 439.51 465.05 465.26 492.92 493.96 494.15	1.983 1.754 1.861 1.844 5.067 1.977 1.680 2.214 3.617	1.976 1.730 1.752 1.735 4.997 1.973 1.669 2.278 3.681	+0.007 +0.024 +0.109 +0.109 +0.004 +0.011 -0.064 -0.064	+ 0.35 + 1.39 + 6.22 + 6.28 + 0.20 + 0.66 - 2.81 - 1.74	23.0 22.0 23.0 23.0 22.0 22.0 22.0 22.0	2.502 2.009	2.039 1.942 1.961 1.991 2.070 1.838 2.301	20.21 26.86 25.82 25.91 0.35 1.67 20.58 29.00 22.80 11.14	41.21 52.16 50.82 51.59 4.18 42.60 53.30 45.80 25.63	4.03 3.36 3.44 3.45 12.77 ^c 4.08 3.07 4.58 8.47	Foram nanno limestone Foram nanno limestone Foram-rich nanno limestone Chert Chert Foram-rich nanno limestone Foram-rich nanno limestone Foram-rich nanno limestone Nanno foram-rich volcanic silty sandstone Nanno foram limestone
19-2, 147-149 20-1, 92-94 20-2, 74-76 20-3, 77-79 20-4, 92-94 20-5, 87-89 21-2, 34-36 21-3, 84-86 21-3, 122-124	523.47 549.92 551.24 552.77 554.42 555.87 579.34 581.34 581.72	3.161 1.827 1.902 2.116 2.040 1.812 1.775 1.990 3.507	3.294 2.034 1.878 1.702 1.994 1.796 1.753 3.233	-0.133 -0.207 +0.024 +0.414 +0.046 +0.016 +0.022 +0.274	- 4.04 -10.18 + 1.28 +24.32 + 2.31 + 0.89 + 1.25 + 8.48	22.0 22.0 22.0 22.0 22.0 22.0 22.0 22.0	1.613	2.224 2.140 1.961 2.022 2.014 1.925 1.904 2.454	15.50 17.46 23.51 21.36 22.61 24.68 25.88 40.18 8.20	34.47 37.36 46.10 43.19 45.53 47.51 49.28 64.81 20.12	7.33 4.35 3.68 3.44 4.02 3.45 3.34 3.21 ^c 7.93	Foram nanno limestone Foram-rich nanno limestone Foram-rich nanno limestone Foram-rich nanno limestone Foram-rich nanno limestone Foram-rich nanno limestone Foram nanno limestone Nanno- foram-rich volcanic siltstone Foram-rich nanno micritic limestone
22-1, 98-100 22-2, 89-91 22-3, 82-84 22-4, 111-113 23-2, 37-39 23-3, 48-50 23-3, 52-54 24-1, 39-41 24-1, 143-145 24-2, 124-126 24-3, 69-71 24-4, 78 - 9	597.48 598.89 600.32 602.11 607.87 609.48 609.52 615.89 616.93 618.24 619.19 620.78	2.700 2.698 2.742 4.114 3.128 2.859 3.291 2.640 3.265 4.285 4.285 4.262	2.757 2.783 2.696 3.835 3.037 3.093 3.058 3.963 3.861 4.194	-0.057 -0.085 +0.046 +0.279 +0.091 +0.198 +0.207 +0.322 +0.401 +0.196	$\begin{array}{r} -2.07 \\ -3.05 \\ +1.71 \\ +7.28 \\ +3.00 \\ +6.40 \\ +6.77 \\ +8.13 \\ +10.39 \\ +4.42 \end{array}$	21.0 21.0 21.0 22.0 22.0 22.0 21.0 21.0	2.209 2.035	2.304 2.268 2.291 2.370 2.329 2.245 2.389 2.441 2.395 2.424	11.36 12.34 5.08 12.03 15.58 10.85 14.88 8.73 4.44 4.50	25.76 28.27 12.04 28.02 34.42 24.36 30.28 20.86 10.83 10.78	6.35 6.31 6.18 9.09 7.07 6.32c 6.94 5.37c 7.31 9.67 9.25	Nanno limestone, laminated Foram-rich nanno limestone, laminated Foram-rich nanno limestone, laminated Micritic limestone, laminated Foram-rich nanno micritic limestone Nanno foram silty sandstone Nanno micritic limestone Nanno vicleanic clayey siltstone Nanno micritic limestone Micritic limestone Micritic limestone

aminated	
Foram-rich nanno micritic limestone Foram-rich nanno micritic limestone Foram nanno-rich clayey micritic limestone Foram nanno-rich clayey micritic limestone Foram-rich nanno micritic limestone Calcareous clayey siltstone Foram-rich nanno micritic limestone Foram-rich nanno micritic limestone Foram-rich nanno micritic limestone Foram-rich nanno micritic limestone Nanno foram-rich clayey micritic limestone Nanno foram-rich micritic limestone Nanno foram-rich micritic limestone Nanno-rich micritic limestone Poram-rich nano micritic limestone Nanno-rich micritic limestone Nanno foram-rich clayey limestone, laminated Volcanic and carbonate, foram-rich sily sandstone, Nanno foram-rich clayey limestone, laminated Volcanic clayey siltstone Volcanic clayey siltstone Volcanic silty claystone, laminated Volcanic silty sandstone	
6.22 5.84 5.09 5.93 5.93 5.93 5.93 5.93 5.93 5.93 5.15 5.06 5.15 5.15 5.15 5.15 5.15 5.15 5.15 5.1	
21.38 27.60 27.60 27.61 27.69 38.04 56.80 27.65 27.55	
9.06 12.10 7.62 12.47 7.62 13.47 7.00 12.23 9.73 8.70 12.25 9.73 8.70 12.25 11.79 14.99 14.99 11.79 11.79 11.76 23.20 33.220 33.220 28.20	
2.360 2.281 1.557 2.281 2.285 2.268 2.285 2.285 2.285 2.285 2.285 2.285 2.285 2.285 2.285 2.285 2.285 2.285 2.264 2.265 2.266 2.2765 2.266 2.2765 2.2765 2.2766 2.2765 2.2777 2.2765 2.2777 2.2765 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.27777 2.277777 2.277777 2.277777777	
2.334	
21.0 21.0 21.0 21.0 21.0 21.0 21.0 21.0	
+ 8,43 +15,14 +17,91 +17,91 +17,91 +15,04 +10,95 +10,95 +14,55 +14,55 +14,55 +14,55 +10,40 +10,40 +10,40 +10,13 +1	
$\begin{array}{c} +0.222\\ +0.417\\ +0.417\\ +0.472\\ +0.472\\ +0.468\\ +0.308\\ +0.308\\ +0.308\\ +0.306\\ +0.306\\ +0.306\\ +0.306\\ +0.306\\ +0.398\\ +0.306\\ +0.306\\ +0.398\\ +0.306\\ +0.306\\ +0.308\\ +0.134\\ +0.231\\ +0.135\end{array}$	86 for basalt bulk density
2.635 2.562 2.562 2.563 2.635 3.111 2.936 2.942 2.942 2.942 2.942 2.942 2.942 2.942 2.942 2.942 2.942 2.724 2.942 2.724 2.942 2.724 2.933 2.724 2.518 2.724 2.518 2.724 2.518 2.724 2.723 2.723 2.723 2.733 2.23333 2.2333 2.2333 2.2333 2.233333 2.23333 2.23333 2.233333 2.233333 2.233333 2.23333 2.233333 2.23333 2.2333333 2.23333 2.23333333 2.233333 2.2333333 2.233333333	herts, and 2.5 tent) X (wet-
2.857 2.979 2.979 3.570 3.579 3.579 3.574 2.531 2.531 2.531 3.574 2.531 3.574 2.531 3.574 2.531 3.548 3.523 4.427 2.451 2.976 2.976 2.976 2.976 2.976 2.978 2.523 2.523 2.523 2.523 2.528 2.523 2.528 2.521 2.5323 2.5323 2.5323 2.5323 2.5323 2.5323 2.5323 2.5323 2.5323 2.5323 2.5323 2.5323 2.5323 2.5328 2.5323 2.5328 2.532	cs. 265 for cl st-water cont
627,03 629,34 630,66 637,16 637,16 640,06 641,20 644,20 644,27 664,23 664,23 664,23 664,23 664,23 664,23 703,22 703,22 703,22 703,25 703,25 703,25 713,996 739,96 742,61 741,46 744,05) for sed. rock corrected we
25-2, 53-55 25-3, 134-136 25-4, 116-118 26-5, 116-118 26-5, 70-72 26-4, 106-118 26-5, 70-72 27-1, 122-124 27-1, 122-124 27-1, 122-124 27-1, 122-124 27-1, 122-124 29-1, 139-141 30-2, 72-74 30-1, 139-141 30-2, 72-74 31-1, 105-107 31-2, 424 32-1, 96-98 32-1, 96-98 32-1, 96-98 32-4, 55-57 32-4, 55-57	$^{a}\rho_{g} \& \rho_{gc} = 2.70$ bPorosity = (salt

for age determination are present only in the upper Oligocene to upper Miocene, and in a part of the middle Eocene section. Sponge spicules dominate the siliceous fraction of the sedimentary rocks. Diatoms and silicoflagellates appear to be absent.

The sediments recovered range in age from Quaternary to lower Eocene (Zone NP13; approximately 49-51 m.y.B.P.). One major unconformity is present between Cores 9 and 10 (245 to 264 m), and is indicated by the absence of part of the lower and middle Miocene section. Another major discontinuity may be present in the base of Core 15, between mid-Oligocene and upper Eocene; and a hiatus, similar or identical to one at Site 317, was found between the middle and upper Eocene of Cores 20 and 19, respectively.

Mixed faunal assemblages, which contain species whose ranges do not overlap, are characteristic of many of the samples examined. Middle Eocene radiolarians are commonly reworked into Neogene assemblages. Among the Eocene foraminifers, the presence of *Asterocyclina* sp. and a pelagic foraminiferal assemblage showing very little dissolution indicates reworking from a relatively shallow source. Thus, reworked debris from a shallower source area (presumably one or more of the nearby Tuamotu edifices) is present within much of the sedimentary material examined at this site.

Foraminifers

^cHorizontal.

A discontinuous and condensed, richly foraminiferal sequence spanning middle Eocene through Quaternary was cored at Site 318. Foraminifers are abundant and well preserved in middle Miocene to Quaternary rocks; with a few notable exceptions, they are poorly preserved in Eocene to lower Miocene rocks. In the upper Eocene through lower Miocene, poor preservation is a result of dissolution, while in the middle Eocene preservation deteriorates downward with increasing lithification (though foraminifers remain common). The faunas throughout are very similar to those in rocks of similar age at Site 317.

Quaternary

Core 1 is of Quaternary age and contains a typical low-latitude planktonic fauna, dominated by the *Globigerinoides sacculifer* group. *Streptochilus tokelauae* is common and extends into the underlying Pliocene and upper Miocene sediments.

Pliocene

Cores 2, 3, and 4 are of Pliocene age and contain a typical Pliocene faunal sequence, including *Pulleniatina spectabilis*; the occurrence of this species at Site 318 is the southernmost yet reported for this areally restricted species.

Miocene

The upper Miocene (Cores 5 and 6) is probably in depositional continuity with the overlying Pliocene, but a hiatus probably separates the *Globorotalia acostaensis* Zone sediments of Core 6 from the *G. menardii* Zone sediments of Core 7, which in turn probably disconformably overlies middle Miocene sediments of Cores 8



Figure 12. Summary of physical properties at Site 318. GRAPE analog data, both corrected and uncorrected for actual core diameter, are presented at the end of this chapter.

TABLE 4 Site 318 Summary of Stratigraphic Grouping on the Basis of Rock Physical Properties (Lab Temperature and Pressure)

Lithology	Depth (m)	Cores	Typical Wet-Bulk Density (g/cc)	Typical Wet-Water Content (wt %)	Typical Porosity (vol %)	Typical Sound Velocity (km/sec)	Typical Acoustic Impedance $\frac{g \ 10^5}{cm^2 \ sec}$	Typical Reflection Coefficient at Boundary (no chert)
Foram nanno ooze and chalk; Breccia and conglomerate at base: Quaternary to Miocene	0 to 265.5	1-10	1.63-1.83 (cgl. 2.2- 2.3)	31-43 (cgl. 12- 30)	53-71	1.61-1.97 (cgl. 2.7- 4.5)	2.8 (2.6-3.6)	
Foram nanno chalk: Miocene to Oligocene	265.5 to ~408	10-15	1.83	28	54	1.7	3.1	0.13
Foram nanno, nanno chalk and limestone: Oligocene to Eocene	~408 to ~494	15-18	2.0 (1.83-2.07) (chert 2.5)	24 (20-27) (chert 1.7)	47 (41-53) (chert 4.0)	1.8 (1.6-2.3) (chert 5.1)	3.7 (3.0-4.1) (chert 12)	0.30
Nanno limestone: Eocene	~494 to ~540	18-19	2.25	13	31	3.4	7.9	0.26
Nanno limestone; Minor interbedded clastics: Eocene	~540 to 581.5	20-21	2.0 (1.6-2.1)	21 (17-40)	46 (37-64)	1.8-2.1	3.2-4.4	
Nanno limestone, siliceous 581.5 to - limestone, clayey lime- stone, and volcaniclastic interbeds: Eocene	581.5 to ~745	21-32	1.8 and 2.3 (1.55-2.65) irregularly decreasing downward	32 and 10 (4-42)	(10-65) irregularly increasing downward	1.9-4.4 generally decreasing downward	3.5-10.0	0.42



Figure 13. Conrad-11 seismic reflection profile, showing reflector stratigraphy.

and 9, which contain *G. fohsi s.l.* Another (more pronounced) hiatus separates the latter from lower Miocene, *Globorotalia kugleri* Zone sediments of Cores 10 through 12, which are in depositional continuity with the underlying upper and middle Oligocene sedimentary rocks of Cores 13-15.

The upper Miocene faunas are monotonous; *Globi*gerinoides groups are dominant and *Globorotalia acosta*ensis is very scare. *Globorotalia kugleri* is dominant in its zone, along with common specimens of the early *Globi*gerinoides lineage.

Oligocene

The Oligocene faunas of Cores 13-15 are poorly preserved and apparently only the globorotalia opima and "Globigerina ciperoensis" zones are represented. As at Site 317, Globorotalia opima is scarce, and G. ciperoensis is not present (the zone is here interpreted as the interval between the last G. opima and first G. kugleri).

Eocene

The upper Eocene (Globorotalia cerroazulensis and Globigerapsis mexicana zones) are contained in rocks of



Figure 14. Analysis of sonobuoy record at Site 318 (preliminary).

Cores 16-19. Preservation shows a pattern similar to that at Site 317; that is, poor preservation (dissolution) in the uppermost Eocene, with improvement downward into the *G. mexicana* Zone. Also, as at Site 317, the upper Eocene disconformably overlies the middle Eocene *Globorotalia lehneri* Zone.

A remarkably well preserved fauna of the *Globigerapsis mexicana* Zone was found in a graded bioclastic bed in Sections 1 and 2 of Core 19. The upper portion of this bed contains commonly occurring specimens of, among others, *Hantkenina* cf. *H. alabamensis*, and specimens transitional between *Globorotalia centralis* and *G. cerroazulensis*. Test walls are often glassy and without the chalky surface of specimens from typically abyssal sequences. The good preservation, along with the presence of *Asterocyclina* in the coarser parts of the bed, indicates a relatively shallow source for the graded bed. Also, a penecontemporaneous origin is suggested by the lack of elements foreign to a *G. mexicana* Zone assemblage.

The uppermost *G. lehneri* Zone fauna of Core 20 is moderately well preserved, but below this level induration increasingly obliterates the foraminiferal assemblages, which are apparently recrystallized, partially infilled, and deformed by compaction. The lowest, questionably identified assemblages are those of Core 25, which probably belong in the *Hantkenina aragonensis* Zone of the lower middle Eocene, and Cores 30 and 32, which probably are of upper lower Eocene age (Acarinina densa Zone).

Large, shore to shelf, tropic to subtropic forms occur erratically but frequently from Cores 2 to 32 (total depth). Amphisteginids, heterosteginids, cycloclypeids (? Spiroclypeus sp., Core 2), discocyclinids (Asterocyclina sp., Core 19), and others (Core 32) occur often with reefal forms (corals, bryozoans, etc) in turbiditic sands (see Beckmann, this volume).

The age difference between large foraminifers and enclosing sediment is pronounced at the top of the hole (Core 2, Eocene-Oligocene large foraminifers in Pliocene sediment) but decreases to approximate equivalence toward the bottom.

Calcareous Nannoplankton

Cores recovered and nannoplankton zones present seem to indicate that rocks of the Quaternary, Pliocene, and upper Miocene are present without interruption from the sea floor to a depth of approximately 250 meters (Core 9, standard nannoplankton Zone NN9). Reworked middle Eocene nannoplankton species occur in Core 2 (26 to 35.5 m) in layers that contain reefal debris. Between Core 9 (235.5 to 245 m) and Core 10 (264 to 273.5 m) a major unconformity is present, indicated by the absence of nannoplankton Zones NN3 through NN8; the age gap is equivalent to the middle Miocene and part of the lower Miocene. Also, corecatcher samples of Core 9 as well as rocks in the upper part of Core 10 contain a remarkable amount of redeposited nannoplankton from the NP25 and lower NN1 interval, nearly masking their true age. It seems probable that part of Zone NN1 is missing between Cores 10 and 11, inasmuch as upper NN1 nannofossils were not found in Core 11, which should have been present if continuous sedimentation took place. Below Core 11 (at ~300 m) an obviously complete sequence was cored down to the lower Eocene nannofossil Zone NP13. At approximately 690 meters the younger, predominantly calcareous sediments become increasingly rich in green siltstones and limestones, with an increase in accumulation rate. With few exceptions, almost all of these lower rocks are barren of calcareous nannoplankton. From the lower Oligocene downwards (Core 6, at ~440 m) nannoplankton assemblages indicate "near shore environments," as Zyrghablithus bijugatus and representatives of the genera Braarudosphaera, Micrantholithus, and Discolithina occur quite frequently.

In the Quaternary and Pliocene part of the sequence calcareous nannoplankton are well preserved, but become moderate to poorly preserved in the Miocene, and solution effects and heavy overgrowth on discoasters can be seen throughout the Miocene section. In the Oligocene and upper Eocene solution effects are less obvious, but overgrowth is present on all nannoplankton species. With the change toward more silty to sandy sediments in the middle Eocene, preservation of nannoplankton becomes very poor, and species identification at certain levels is extremely difficult or impossible.

Radiolarians

Radiolarians of sufficient abundance and preservation for age determination are present in only two intervals of the sedimentary column cored at Site 318. The upper Miocene through upper Oligocene (Cores 5 through 13; which range in depth from 131 through 359 m) have a siliceous assemblage dominated by sponge spicules, but contain a sufficient number of diagnostic radiolarians for reliable age control. The middle Eocene *Thyrsocyrtis triacantha* Zone is represented in a moderately to poorly preserved radiolarian assemblage in Cores 22 through 25 (depth range 606 to 635 m). The remainder of the cored intervals are barren of radiolarians.

Within the radiolarian-bearing intervals, the following standard radiolarian zones can be recognized: Core 5 (121.5 to 131.0 m) is in the *Stichocorys peregrina* Zone. Core 6 (150.0 to 157.5 m) is in the *Ommatartus penultimus* Zone. Core 7 (178.5 to 188.0 m) is in the *Ommatartus antepenultimus* Zone. Cores 8 and 9 (207.0 to 245.0 m) are in the *Calocycletta costata* Zone. Cores 10 and 11 (264.0 to 302.0 m) are in the *Calocycletta virginis* Zone. Core 12 (321.0 to 330.5 m) is in the *Lychnocanoma elongata* Zone. Core 13 (349.5 to 359.0 m) is in the upper *Dorcadospyris ateuchus* Zone. Cores 14 through 21 are barren of radiolarians, or contain only unidentifiable fragments. Cores 22 through 25 (596.5 to 634.5 m) are in the *Thyrsocyrtis triacantha* Zone. Radiolarians are absent below Core 25.

Reworked radiolarians of middle Eocene age are rare to common in virtually the entire Neogene section (Cores 5 through 13). The physiographic setting of this site is favorable for the accumulation of erosional products from nearby Tuamotu edifices; thus, the extensive reworking of radiolarian debris is not surprising.

ACCUMULATION RATES

As shown in Figure 15, sedimentation rates at Site 318 have been high during the past 50 m.y. Situated in a trough atop a volcanic ridge and surrounded by atolls, the site has been receiving carbonate sediments of both pelagic and reefal origin as well as abundant volcaniclastic debris of clay to pebble size. Thus the high rates are not necessarily indicative of high productivity of plankton in the water column; turbidites are a major component of the section. From early to middle Eocene time the accumulation rate was approximately 65 to 70 m/m.y.; abundant reefal and volcanic material marks this part of the section. For the remainder of Eocene time the rate slowed to approximately 25 m/m.y. Through Oligocene time the rate further slowed to 12 m/m.y. as pelagic sedimentation dominated. After formation of the early-middle Miocene unconformity, the rate increased from about 13 to 25 m/m.y., marked by a heavy influx of reefal and volcanic material.

SUMMARY AND CONCLUSIONS

Site 318 was drilled very close to its preplanned location along the *Conrad*-11 track, in 2659 meters of water. One hole was drilled to a depth of 745 meters; nearly 300 meters were cored, with an average core recovery of more than 49%. It was hoped that the lowest reflector



Figure 15. Graphic log showing age and rate of sediment accumulation at Site 318.

observed on both the *Conrad*-11 and our own acoustic profiles, which lay at predrilling estimated depths between 650 and 700 meters, was basalt basement. We reached depths exceeding our original estimates, but did not succeed in penetrating or sampling basalt basement rocks at the site, although we used all of our remaining leg time in attempting to do so.

The geologic column drilled at the Tuamotu Ridge site was divided into five lithologic units:

1) Nannofossil-foraminifer ooze that contains some graded beds of shallow water skeletal debris made up of foraminifera, bryozoans, corals, echinoids, and mollusks, and altered volcanogenic material. This unit extends from 0 to between 35.5 and 64.5 meters, and is of Quaternary to upper Pliocene age; large foraminifers of Eocene to Oligocene(?) age are present as redeposited faunal elements.

2) Foraminifer and nannofossil firm ooze, changing downward to soft chalk that contains a few graded beds of sand-sized material of volcanogenic origin. The unit extends from depths of between 35.5 and 64.5 to between 245 and 264 meters, and is late Pliocene to middle Miocene in age. It is separated from Unit 3 by a middlelower Miocene hiatus between Cores 9 and 10.

3) Foraminifer-nannofossil, yellowish-gray chalk, that contains a breccia bed with pebbles of basalt, chert, and shallow water skeletal debris of reef and near-reef origin. This unit extends from depths of between 245 and 264 meters to between 416 and 435 meters. The unit is of lower Miocene to lower Oligocene age.

4) Nannofossil and foraminifer limestone, generally yellowish-gray to pale green in color, with abundant chert nodules and lenses that include some volcanogenic debris. The base of the unit is a coarse, pebbly sand containing abundant shallow water skeletal debris. The unit occurs between 416 and 435 meters to between 530 and

549 meters, is lower Oligocene through upper Eocene in age, and its base coincides with a minor hiatus and with the 530-meter acoustic reflector discussed below.

5) Green and grayish-green clayey limestone, siltstone, and sandstone, in graded beds as thick as 2.5 meters. The upper 100 meters of the unit contain some chert, and the amount of volcanogenic materials increases downward. The unit occurs between 530 and 549 meters depth to the bottom of the hole at 745 meters, and is of lower to middle Eocene age. At the base of the unit abundant coarse-grained reef debris characterized by large benthonic foraminifers was cored. The unit contains the acoustic reflector at 625 meters, discussed below.

A number of reflectors can be distinguished on the profiler records near Site 318, but all except two are irregular and discontinuous. This irregularity may be due to periodic but not necessarily chronologically equivalent influxes of sedimentary debris from the several edifices that surround the basin. Preliminary analysis of reflector stratigraphy suggests that the 0.56 to 0.60 sec reflector observed in the underway profiles corresponds to an unconformity between middle and upper Eocene strata, and that the 0.68-sec reflector observed in the sonobuoy records corresponds to the appearance of the very hard limestones at a depth of about 625 meters, some 90 meters beneath the unconformity.

The geological history of the general area drilled can be partially derived from bathymetric, profiler, and drill cores retrieved, but the volcanic history must remain largely speculative as a result of our inability to reach basement. The site was chosen near the northwest terminus of the Tuamotu chain, in an open sedimentary trough atop a ridge-like feature that we consider to be a portion of one of the Tuamotu constructional edifices. The basin proved to be filled with more than 745 meters of sedimentary rocks, beginning with an exceedingly heavy influx (as much as 70 m/m.y.) of coarse volcanogenic debris and shallow water foraminifers. Ages derived from several paleontological groups near the bottom of the hole are in close agreement, and yield an age at that depth of 49-51 m.y.B.P. If the basement rocks beneath are a portion of the Tuamotu Ridge, and if the volcanogenic sediments are derived by erosion of higher portions of the ridge or the nearby edifices marked by the coralline islands of Ahe, Rangiroa, and Arutua, then all of the basaltic edifices in the immediate area are older than that date. How much older, we cannot say, except that, considering the sedimentation rate and character and abundance of the volcanogenic debris at the bottom of the hole, it seems unlikely that the cessation of volcanism was prior to early Eocene time. Deep Sea Drilling Site 76 (Hays et al., 1972) also recovered reefal debris of probable lower Eocene age that had been redeposited in Pliocene turbidites on the northern flank of the Tuamotu chain about 175 km north of Site 318, which tends to indicate that most of the edifices in this general area were built; eroded, and capped by reefs by 50 m.y.B.P.

At that time the basin beneath Site 318 was receiving turbidity currents consisting of volcanogenic sandstones

and siltstones, intermixed with shallow water foraminiferal debris. This debris arrived at an average rate of 70 m/m.y. through much of lower and middle Eocene time, gradually decreasing in middle Eocene time. Nearly continuous pelagic sediments were deposited in the basin from that time to the present, although depositional hiatuses mark middle and late Eocene time, mid-Oligocene and upper Eocene time, and early through middle Miocene time. During the entire period from middle Eocene through Quaternary, the basin continued to receive small amounts of reefal and volcanogenic debris. Such small pebbles of basaltic rock as are present appear to be of alkaline character (the rocks appear to be of alkalic basalt-hawaiite types). Floods of reefal debris entered the basin in turbidite flows in middle Eocene and early Miocene time; turbidite deposition of reefal material has evidently continued to the present as shown by skeletal debris in Core l.

Early Eocene edifice ages in this area do not "extrapolate" to Line Islands edifice ages; indeed, at the conclusion of our Line Islands sites we were unable to document a southward younging of the Line Islands. The morphology and available age data in the Tuamotu chain, however, still suggest that it is of Hawaiian aspect and that it does young to the southeast (Dana, 1876; Clague and Jarrard, 1973; Jackson and Shaw, 1975). The age of basement basalt below Muroroa Atoll, approximately 1300 km to the southeast of Site 318 is reported to be 7 ± 1 m.y.B.P. (Chevalier, 1974), and the age of Gambier Island, approximately 1800 km from Site 318, if indeed it is a part of the Tuamotu chain, is given between 4.7 and 6.8 m.y. (Brousse, et al., 1972). These rather incomplete data suggest that the Tuamotu chain, like the Hawaiian chain (Shaw, 1973; Shaw and Jackson, 1973) becomes younger to the southeast in an episodic rather than a linear fashion.

Inasmuch as we failed to reach basement at Site 318 we can give only minimum ages for constructional volcanism in this part of the Tuamotu chain. Nevertheless, the minimum age of 49-51 m.y.B.P. is older than Clague and Dalrymple's (1973) estimate of 42 m.y. for the Hawaiian-Emperor bend. Their estimates of this age was in part based on a date of 46 m.y. at Koko Seamount. More recently Leg 32 of Deep Sea Drilling Project (Larson, Moberly, et al., 1973), reported fossil ages at Koko as much as 6 m.y. older than Clague and Dalrymple's radiometric dates on volcanic rocks there. Still more recently, Clague et al. (1975) dated Yuryoku Seamount at the Hawaiian-Emperor bend at 42.3 ± 1.6 m.y.B.P. Based on data at hand, it would appear that Tuamotu volcanism preceded Hawaiian-Emperor bend volcanism, at least by a few millions of years.

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Explanatory notes in Chapter 1

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AGE	ZONE	FOSSIL 5	ABUND.	PRES.	SECTION	METERS	LITHOLOGY	DEFORMATIC	LITHO.SAM		LITHOLOGIC DESCRIPTION
					0						
		N	A	M	1	0.5			68	58 9/1	White (N9) and bluish white (58 9/1) FORAM NANNO FIRM OUZE to CHALK. A few mm laminae of light gray (N7) color. Smear Slides: Forams C to A Nannos D
MIOCENE	otalia menardif NN9 tus petterssoni				2	minute		 		N9 to 5G 9/1	Rads R Sponges R
UPPER	7G1 obor Cannar					111111		1 1 1	148	N9	
					3	milin			88		
		F N R	A F	G M G	Cat	ore tcher			cc	- 255	
te	318	Hole	00001	_	c	ore 9	9 Cored I	nter	val:	235.5-245.0 m	
	w	CHA	RACT	TER	NO	ø		NOL	MPLE		
AGE	NOZ	FOSSIL	ABUND.	PRES.	SECTI	METER	LITHOLOGY	DEFORMAT	LITH0.5/		LITHOLOGIC DESCRIPTION
-	5				0			-	_		
E MIOCENE	talia menardii N9 etta costata or youngu	N	A	м	1	1.0			96 111	N9	White (N9) FORAM NANNO CHALK with a few mm-scale laminations. Rare (1%) sand-sized grains of dark volcanogenic material throughout. At 2.30 m: a pebble of very dusky green (55 Z/2) palagonitic clay, about 1 cm in diameter. Smear Slides:
MIDDL	761oborot N Calocycle	N	A	H	2	and and and			80	→ 5G 2/2	Nannos D to A Micrite R to A Rads R Sponges R to C
	1						1 1 1 1			N9	

AGE ZONE	FOSSTI D	HARAI UNDER	BRES.	SECTION	METERS	100000 00 00 A	LITHOLOGY	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FOSSIL P	ACTE	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC DESCRIPTION
LOMER MIOJENE Globorotalia kugleri Calocycletka kirginis	N N N N N F		м н м м	0 1 2 3 4	0.5 1.0				1100 120 125 87 56	Yellowish gray (5Y 7/2) FORAM NANNO CHALK, preserved as "biscuits" separated by intervals of drilling paste. The chalk "biscuits" are shattered by fractures induced by the coring process. 5Y 9/1 5Y 2/4 5Y 7/2 From 1.1 to 1.5 m are pieces of: A. BRECIA, consisting of angular fragments of various VOLCANIC ROCKS to 1 cm diam., mainly in shades of very dusky red (5R 2/4, 10R 2/2, etc.), vesicular, altered; SHELLY FORAM LIMESTONE, in shades of yellowish gray (5Y 7/2) and dusky yellow (SY 6/4); and SKELETAL DEBRIS, up to 5 mm. About 50% vol- canic, 40% sedimentary, and 10% shell debris. In T.S., scoria fragments, with plagicalase X1s, palagonite, large benthonic forams, gastropod (7). 5Y 7/2 B. BASALT, as rounded cm-size pebbles. Thin sections of 2 of these as follows: 1. Plagicalase augite phyric, fine- grained, sparely vesicular basalt with fine, ilmenite-rich, glassy groundmass; intersertal texture. Est. mode: vesicles (partly filled). 5%; plagicalase phenocrysts, 15%; augite phen., 5%; replaced augite phen., c5%; groundmass upite, 20%, groundmass ilmenite, 5%. 2. Vesicular vitrophyric basalt. Est. mode: vesicles (~2 mm), partly filled, 30%; groundmass plagicalase, 25%; groundmass plagical	LOWER MIDCENE	Globorotalia kugleri NNI Calocvcletta vireinis	N	A 1	0 1 2 3 4 5	0.5			25 81 30 -	5Y 8/1 N8 5Y 8/1 - N9 5Y 8/1 - 5Y 9/1 5Y 8/1	Pale yellowish gray (5Y 8/1) FORAM NANNO CHALK, mottled sparsely with pale gray (M8) and light blive gray (5Y 5/1) small (1-5 mm) burrows. At 2.8 m, a 5-cm layer of white (NY FORAM MANNO CHALK, with light gray (N6) mottles. Darker material concu- trated at upper and lower contacts At 3.3 m, a 5-cm layer of yellowisi white (5Y 9/1) NANNO FORAM CHALK, with dark-colored (volcanic) grain making up ~1% of the rock. At 8.7 to 8.8 m, dark-colored volcanic sand grains constitute ~1% of chalk. Chalk present as "biscuits", internally fractured by the coring process, and separated by a drillin paste. Smear Slides: (1-25 cm): Nannos D Forams R to C Rads R Sp. Spic. R Silicofl. R (core catcher): Forams A to D Nannos C to A Rads R Sp. Spic. R

M

6

Core Catcher

+ ++

CC

SITE 318

		F	055	IL.			1	12	5		
AGE	ZONE	FOSSIL E	ABUND.	PRES. 3	SECTION	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMP		LITHOLOGIC DESCRIPTION
					0						
	eri gata				1	0.5	VOID + + + + + + + + + + + + + + + + + + +		35	N9 	White (N9) and pale yellowish gray (SY 8/1) FORAM NANNO CHALK, with a few sandy foram-rich laminae, with ∿1% dark-colored volcanic grains. Smear Slides: (1-35 to 40 cm): Nannos C to A Forams R to C
ENE	kugl elon					1			10	5Y 8/1	Micrite C Volc. Glass R
LOWER MIDCI	Globorotalia NN1 Lychnocanoma				2				108	N9	(2-108 cm): Nannos D Forams R Rads R Sp. Spic. R
		FN	CA	PM	Co	ore					
					La	CLIEF	11111	1	L		
ite	318	Hole	055	1	-	ore	3 Cored I	nter	val:	349.5-359.0 m	
w	NE	CHA	RAC	ER	NO1	8		VLION	SAMPL		TTUN ALL DECONDENSION
AG	20	FOSSIL	ABUND.	PRES.	SECT	METE	LITHOLOGY	DEFORM	L1TH0.3		LITHOLOGIC DESCRIPTION
_					0						
					1	0.5	VOID				Very pale orange (10YR 8/1), pinkish white (SY 9/1), and white (N9) FORAM NANNO CHALK, with a few mm-scale burrow mottles of medium to light gray (N5 to N7).
		N	A	M		1				5YR 9/1	At 1.3 m, a 3-cm layer of pale yellowish gray (5Y 8/1) NANNO FORAM CHALK with alf
_	1				-	-	1 1 1			N9	dark-colored volcanic sand
						-	1 + + + + + + + + + + + + + + + + + + +				
OLIGOCENE	ia ciperoensis ris ateuchus				2	ofcodan				10YR 8/1	

		CHA	RAC	TER	N	~		ION	MPLE		
AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTI	METER	LITHOLOGY	DEFORMAT	LITH0.SA	LIT	HOLOGIC DESCRIPTION
					0						
	NP25	N	A	м	1	0.5	V010		20	N9	0.1-1.5 m: White (N9) and very light gray (N8) FORAM NANNO FIRM CHALK, with occasional laminae of light gray (N7), and faint white (N9) mottles. Rare dark-colored volcanic sand grains in foram-rich layers. 1.5 m:
DLIGOCENE	fa opfma 5	N	A	м	2	undun		1		N9 to N8	Ye llowish gray (51 G/1) and olive gray (5Y 5/1) CHERT, in a 5-cm chunk. 1.5-2.5 m: White (N9) to very light gray (N8) FORAM NANNO CHALK, mainly fractured
UPPER	61 oboro ta NP2								123 23	TOR 9/1 — GEOCHEM	2.5-4.5 m: A single graded unit, consisting of pinkish white (10R 9/1) NANNO CHALK grading down into white (N9) FORAM NANNO CHALK, orading
					3	dunda				N9 	by gradual increase in proportion of forams into very pale grayish orange (10YR 8/4) to grayish orange (10YR 7/4) to yellowish orange (10YR 7/4) NANNO FORAM CHALK, with dark-colored volcanic grains increas-
		FN	C A	р-М М	Co Cat	ore Lcher		i	140		Smear Slides: (1-20 cm; 2-123 cm; 3-23 cm): Nannos D Micrite C
											rorams x Rads R Sp. Spic. R (3-140 cm):
											nannos A Forams A Micrite A Fe-Specks R to C

		F CH/	OSS: RACI	TER	NO	s		NOI	MPLE		
AGE	ZOWE	FOSSIL	ABUND.	PRES.	SECTI	METER	LITHOLOGY	DEFORMAT	LITH0.S/	LI	THOLOGIC DESCRIPTION
					0						
		N	A	M	1	0.5	VOID		55	5Y 8/1 and 5Y 6/2	Pale yellowish gray (5Y 8/1, 5Y 8/2) FORAM NANNO HARD CHALK, with mm-scale mottles of light olive gray (5Y 6/1, 5Y 5/2) and laminae of light olive gray (5Y 5/2).
E OLIGOCENE	lia opima NP24	N	A	м		1.0				5Y 8/2	0.4-0.6 m: White (N9) NANNO FORAM CHALK, with laminae of pale orange (10YR 8/3) FORAM CHALK, with rare volcanic sand grains.
MIDOL	Globorota	N	A	н	2					10YR 5/4	$\frac{At 2.2 \text{ m}}{A 3-\text{cm}}$ chunk of moderate yellowish brown (10YR 5/4) CHERT.
	NP23	N	A	м		1			137	5Y 7/2	Smear Slides: (1-55 cm): Nannos A Forams A
		FN	C A	P M	Ca	ore tcher				51 8/2	Micrite A Foram perforations filled with red fe-specks.
											(2-137 cm): Nannos D Fe-Specks R to C Formes R

		F CH/	OSS	TER	NO	8		NOL	MPLE				
AGE	ZONI	FOSSIL	ABUND.	PRES.	SECTI	METER	LITHOLOGY	DEFORMAT	LITH0.5/	LIT	HOLOGIC DESCRIPTION		
UPPER EDCENE LOWER OLIGOCENE	Gioborotalia cerroazulensis NP19/20 1 NP21 NP21 2	N N N LOSS	A A A A	PRES	8 0 1 2 3	0.5		DEFOR	93 107 83	5Y 8/1 N9 and N8, N7, N6 SY 8/1 and N6 5Y 8/1 and N8 5G 9/1 10GY 8/1	0.9-2.1 m: Pale yellowish gray (5 FOSSIL and FORAM NANNO extensively burrowed a by light olive gray (5 2.1-3.2 m: Thinly laminated white gray (M8, M7, M6, M5) STOME. Tiny normal fau laminae, and at 2.6 to recumbent fold, with m folds on the limbs. 3.2-4.5 m: Yellowish gray (5Y 8/1 N8), white (M9), green (56 9/1) and pale yell (106Y 8/1) NANNO LIMES darker gray (N5, N6) a green (106Y 5/2) thin At 2.3 m: Two small pebbles of d volcanic material. In core catchers. Smear Slides: (2-83 cm):	Y 8/1) NA LIMESTON Ad mottle Y 5/2). (N9) and NANNO LIM Its displ 2.7 m is inor sheat), gray (1 ish white cowish gree owish gree tonk with d grayisl laminae. eeply alt ellowish l	NNO- E, d Sce a r r N7, en h t t t t t t t t t t t t t t t t t t
		N	A	M	Ca	ore tcher				10YR 4/2	Nannos Forams Fe-Specks (1-93 cm) white burrow Micrite	A to D R to C C ed area:	black concre tion
											Micrite Nannos Forams (1-107 cm) brownish no Nannos Forams	D C R to C nburrowed D C	area:

326



6 337	CHARACTER	N S	MPLE
ZONE	FOSSIL ABUND. PRES.	LITHOLOGY	LITHOLOGIC DESCRIPTION
		0	
Globigerapsis mexicana NP17/18	F C G F C G F C G N C G F C G	2 Core Catcher	An artificially graded sequence granule- to fine-grained sand, consisting of volcanogenic (pai- agonite) grains, cherty grains, and biogenous grains, including planktonic and large benthonic forams, bryozoans, algae, and molluscan fragments. Core catcher: Light olive gray (SY 5/1) CHERT, and pale green (106 8/1) NANNO LIMESTONE, with thin purplish gray (SP 5/2) laminae. Smear Slides: (2-140 cm) coarse fraction >1 mm abraded and rounded allocapics on Halimeda R Bryozoans A Gastropods R Large Forams B to C

300 FOSSIL CHARACTER 115501 NO S2 S2 S2 S2 S2 S2 LITHOLOGY S2 S2 NO S2 S2 S2 NO S2 S2 S2 NO S2 S2 S2 S2 S2 S2 LITHOLOGY S2 S2 NO S2 S2 S2 NO S2 S2 NO S2 S2S2 NO S2 S2 </th <th>Site</th> <th>318</th> <th>Hole</th> <th>£</th> <th></th> <th>C</th> <th>ore 20</th> <th>Cored I</th> <th>nter</th> <th>val:</th> <th>549.0-558.5 m</th>	Site	318	Hole	£		C	ore 20	Cored I	nter	val:	549.0-558.5 m
N F P/M VIII VIII VIII N F P/M N <th< td=""><td></td><td></td><td>F CHA</td><td>OSS</td><td>IL TER</td><td>NC</td><td></td><td></td><td>ION</td><td>MPLE</td><td></td></th<>			F CHA	OSS	IL TER	NC			ION	MPLE	
N F P/M 0.5- 1 0.5- 0	AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTIO	METER	LITHOLOGY	DEFORMAT	LITH0.SA	
N F P/M 0.5 1 0.5 0.5 0.5 0.5 0.5 0.5 0.5 0.5			Τ			0					
			N	F	Р/М	1	0.5			80	N5, 5Y 4/1 10GY 8/2

	ŭ	A	a.				B	3			
				0						100 M - 100000 M - 5	
	N	F	P/M	1	0.5			80		N5, 5Y 4/1 10GY 8/2	Pale yellowish green (10GY 8/2), pale green (10G 6/2), greenish white (10G 9/1), and 1ight greenish gray (5G 8/1) NANNO LIMESTONE with abun- dant burrows (of light material in darker). From 2.7-5.5 m, thinly laminated, on mm-scale, with may low-anole
					-	V010			=		faults and deformed phacoids (burrows), suggesting small-scale
			1.7		1					and	slumping.
	N	c	M	2	1					106 7/2	A few graded layers, most notably at 6.25 to 6.45 m.
				-						106 4/2	CHERT at 1.1 and 1.5 m (moderate yellowish brown, 10YR 5/4), and at 2.1-2.2 m (pale yellowish brown (10YR 6/2). Chert has ~0.5
					=						cm spots of partly silicified lime- stone within.
				3	- produce			130		5GY 2/1 GEOCHEM	Smear Slides: (1-80 cm, 3-30 cm, 4-148 cm): Nannos A Micrite A Forams R Fe-Specks R Volcamic Fragments R to C
9					-	VOID				100 510 100 011	
LdN				4					-22	10G 6/2, 10G 9/1, 5G 8/1 10YR 6/2	
					1			148			
					-	VOID		-			
				5	111111111					106 8/2, 5GY 4/1, 5GY 2/1 106 8/2 to 5G 7/2	
	F	c	M	Ca	ore tcher	VOID					

LITHOLOGIC DESCRIPTION

		F CH/	OSS RAC	IL TER	N	~		NOI	MPLE	
AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTIO	METER	LITHOLOGY	DEFORMAT	LITH0.54	LITHOLOGIC DESCRIPTION
					0					
		N	c	м	1	0.5	VOID		126	Interbedded and interlaminated pale green (106 8/2) and greenish gray (56 7/2) NANNO LIMESTOWE, and CLAYEY LIMESTOWE, with occassional beds and laminae of yellowish gray (5Y 7/2) and olive gray (5Y 5/2, 5Y 4/2) VOLCANIC SILTSTOWE and SANDSTONE, rich in forams, and laminae of grayish purple (5P 4/2).
MIDDLE EOCENE	91dN				2	malandan			30	Grayish black (N2) CHERT at 2.6 and 2.9 m, and dark brownish gray (5YR 3/1) CHERT at 3.4, 3.55, 3.8, 4.0, 4.35 m. Many graded layers, silty or sandy at base, and extensively burrowed near the top. Smear Slides: (1-126 cm):
		N	с	M	3	100 million			53 83 89	Micrite D Nannos C Forams R (2-30 cm) dark grainy bed: Nannos C Micrite D Forams R Green Clay Fragments R
		FN	C F	P P	Ca	ore tcher	VOID			

Explanatory notes in Chapter 1

MIDDLE EOCENE

Site 318		Hole			C	ore	22	Con	d In	ter	val:	596.5-606.0 m		Site	318	в н	ole			Cor	e 23	Co	red I	nter	val:	606.0-615.5 m
AGE	CUNE	FOSSIL PH	ABUND.	PRES. 33 7	SECTION	METERS		LITHOLC	GY	DEFORMATION	LITH0.SAMPLE		LITHOLOGIC DESCRIPTION	AGE		ZONE	E A TISSOL	RACTEL RACTEL	CELTINN	SECTION	METERS	LITHO	.0GY	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC DESCRIPTION
MIDLE EQCENE Globigerapsis kugleri	CL AN	FN	c c	P	0 1 2 3 4	0.5 1.0					77	alternating 10GF 7/2, 10GF 5/2, 5GY 3/2, and 10G 4/2	Alternating pale yellowish green (DGY 7/2), pale green (DG 6/2), grayish green (DGY 3/2) and grayish green (DG 4/2) NANNO LIMESTONE and LATEY NANNO LIMESTONE. Arg green (SG 2/2) NANNO LIMESTONE and SANOSTONE and SILTSTONE. Arg green ish black (SG 2/1) VOLCANIC SANOSTONE and SILTSTONE. Trevalent burrowing, of lighter mine. Burrowing, of	MIDDLE EDCENE	Globigerapsis kugleri(?)	STON	NF	FF		0 1 1 1 2 2 Core Catch	.5				1399 147 90 147 46 116	Graded beds, generally consisting of light greenish gray (56 8/1) NANNO LIMESTONE or CLAYEY NANNO LIMESTONE, commonly with darken-colored olive, gray (57 4/1, 57 3/1) SILTSTONE or dark greenish gray (56 4/1) fine- grained SANDSTONE, consisting mainly of volcanic (palagonite) grains and forams, at the base. Sand commonly laminated and x-strat. Burrowing of lighter into darker material becomes more intense towards the top of each graded unit. Occassional purplish gray (5P 4/2) laminae and streaks near the tops of graded units. Occassional internally deformed burrows. Graded units 1-15 cmthick, averaging about 5 cm. Smear Slides: (1-139 cm): Volcanic Glass A Micrite C Nannos R (1-147 cm): Nanos R (1-147 cm): Micrite A Sp. Spic. R (2-90 cm): (2-90 cm): Micrite C Feldspar R Green Clay R (2-147 cm): Micrite C Feldspar R Green Clay R (2-147 cm): Micrite C Feldspar R Volcanic Glass R to C (3-116 cm): Micrite A Volcanic Glass R to C (3-116 cm): Micrite A Volcanic Glass R to C Nannos C Na

ite 318	H	ole		0	ore	24	Cored	Inte	rva	:615.5-625.0 m	ite	318	Ho1	8		Co	re 25	Cored I	nter	val:	:625.0-634.5 m
AGE ZONE	-	FOS CHARA TISSOJ	SIL CTER	SECTION	METERS		LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FOSSIL R	ARACT . GNUBA	PRES, BIT	SECTION	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC DESCRIPTION
MIDDLE EGCENE NP15		NR	p	0 1 2 3 4	0.5				11	Graded beds, from 1-40 cm thick, averaging about 10 cm, generally consisting of 1ight greenish gray (5G 7/1) MANNO LIMESTONE and CLAFEY LIMESTONE, commonly with darker-colored olive gray (5Y 4/1, SY 3/1) SILTSTONE or fine-grained FORAM VOLCANIC SANUSTONE at the base, generally laminated or even cross-stratified, with an abrupt contact with the underlying unit. Dccassional beds with coarse-grained or granule-sandstone. Middle and upper parts of beds burrowed, with light material in darker host. Bur- rowing increase in intensity towards top of beds. Grayish purple (5P 4/2) laminae and wisps near tops of some beds. Grayish brown (SYR 3/2) and dark greenish gray (5G 4/1) CHERT blabs at 3.75 and 3.80 m, and a very dusky purple (5P 2/2) CHERT splotch at 4.20 m. Limestone in this core, especially the lighter-colored intervals and the purplish bands, are very hard and siliceous. Smear Slides: (2-62 cm): Micrite D Chert R to C Nannos R (3-50 cm HCL insolubles): (1) spheroidal blebs ~100µm diam. coarse chert replacing forans and/or rads. (2) green clay (1) + (2) = about 10-20% of rock.	MIDDLE EOCENE	Hantkenina aragonensis NP15	NF	cc	MP	0 1 2 3 4	e.e.			108 121	Graded beds, from 1-40 cm thick, averaging about 10 cm, each bed generally consisting of gravish green (56 Gr1, 56 Gr2, 567 Gr2, 10GY 5/2) MANNO LIMESTONE, MICRITE LIMESTONE, and CLAYE LIMESTONE, commonly with darker-colored olive grav (57 S/1, 57 Gr1) SLITSTONE or dark greenish black (56 Z/1) VOLCANIC SANOSTONE, generally fine-grained, but occassionally up to granule- size. Sandy layers commonly lami- nated and cross-stratified, with abrupt, erosional contacts with underlying unit. Middle and upper parts of beds burrowed, with bur- rowing commonly intense at the top of a bed and decreasing in inten- sity below; with light-colored material in a darker host. Some mottles internally deformed. Cherty blebs at 1.74 m. Smear Slides: (2-108 cm; 2-121 cm): Micrite A Nannos A Forams R 100µm chert blebs R to C (4-130 cm): Sp. Spic. A Diatoms C to A Micrite C Rads C

Site 318	Hole		C	ore 26	5	Cored I	nter	rval:	534.5-644.0 m	Site	318	Ho1	е		C	ore 27	Cored	Inter	val	:644.0-653.5 r	
AGE 7045	FOSSIL CHARAI VIII VIIIII VIIIII VIIIII VIIIII VIIIIII	SIL CTER .S384	SECTION	METERS	LI	THOLOGY	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FOSSIL 2	FOSS ARAC	SIL	SECTION	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE		LITHOLOGIC DESCRIPTION
ITODLE EOCENE NOTE			0 1 2 3	0.5				140	Graded beds, from 1-70 cm thick, averaging about 10 cm, each bed generally consisting of grayish green (SG 6/2, SG 5/1) LiMESTONE or CLAYEY LIMESTONE, commonly with a laminated or cross-stratified basal unit of greenish black (SG 2/1) FORAM-RICH VOLCANLC SAND- STOME (sait and pepper colors), or olive gray (SY 4/1 to SY 2/1) SILTSTOME. Contact with underlying bed abrupt, commonly erosional. Sandstome generally fine-grained, but coarse-grained or even granule- size grains occur at a few levels. Burrowing intense in upper parts of graded beds, decreasing down- ward; rarely do burrows cut base of a bed. Minor low-angle faults and shear folding, and deformed burrows, especially at -3.0 m. At 6.3-7.0 m, a single graded bed, with medium-grained foram-woclanic sandstome at base, laminated cross- stratified sandstome in middle and burrowed at top. Greenish black	MIDDLE EOCENE	NP 14	NF	F ?	P ?	0 1 2 Coc	0.5	VOID		142	HCL.	Graded beds, from 1-30 cm thick, averaging about 10 cm, each bed generally consisting of greenish gray (56 5/1) LIMESTONE or CLAYEY LIMESTONE, commonly with a laminated or cross-stratified basal part of olive gray (54 4/1, 55 2/1) SILT- STONE or dark greenish gray (56 4/1) to greenish black (56 2/1) "salt-and- pepper" FORAM-VOLCANIC SANDSTONE, generally fine-grained, but with a few coarse-grained to granule-bearing units. Burrowing intense near tops of graded beds, decreasing downwards. Smear Slides: (2-142 cm): Nicrite D Clay C Nannos R (2-145 cm HCL insoluble residue): 1) clay D 2) chert blebs ~100um C 1) + 2) = 10% of rock
ž			-		臣				(5G 2/1) at top, dark greenish gray (5GY 4/1) in middle, greenish gray	Site	318	Ho1	e		0	ore 28	Cored	Inter	val	663.0-672.5 m	
			4	Treedien				70	(Saf 6/2/ at Uase- A pyritized "chevron" burrow at 6.65 m - within graded bed. Smear Slides: (1-140 cm):	AGE	ZONE	FOSSIL 2	ARAC . UND	BRES - Sand	SECTION	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE		LITHOLOGIC DESCRIPTION
	N F F ?	FP?	5 Cat	re				95	Nannos R to C (5-95 cm): Volcanic Glass D Hicrite C	MIDOLE EDCENE	8 EGN				0 1 2 3	0.5	V01D		110		Graded bes, from 0.5-15 cm thick, averaging about 10 cm, each bed generally consisting of greenish gray (56 G/1 to 56 4/1) LIMESTONE or CLAYEY LIMESTONE, commonly with a basal part of laminated or cross- stratified olive gray (5Y 4/1 to 5Y 2/1) SiLTSTONE or greenish black (56 4/1 to 56 2/1) "salt-and-pepper" FORAM-BEARING VOLCANIC SANDSTONE, generally fine-grained, but coarse- grained or even granule-bearing in a few beds. Burrowing intense near tops of graded units, decreasing below. Wackestone textures in some graded units, with volcanic grains and forams suspended in clayey limestone matrix. From 2.1-2.15 m - laminated grayish blue green (56G 5/2) CLAYEY LIME- STONE. Smear Slides: (1-110 cm): Micrite A Forams C
												N F	F ?	P ?	Ca Ca	ore tcher					(3-15 cm): Micrite A Nannos C Volcanic Glass A to D Green Pyroboles R Fe-Specks C

SITE 318



ite	318	Hole	80		0	ore 31	Cored I	nter	rval	729.5 m
		F CH/	OSS	IL TER	NO	2		NOI	MPLE	
AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTI	METER	LITHOLOGY	DEFORMAT	LITHO.SA	LITHOLOGIC DESCRIPTION
					0				T	
		NF	C ?		1	0.5			36 138	Graded beds, from 0.5-15 cm thick, averaging about 5 cm, each bed consisting of greenish gray (56 6/1 through 56 4/1) and grayish blue green (586 5/2) to dusky blue green (586 3/2) CLAYEY LIMESTONE, commonly with a basal part of darker-colored to greenish black (56 2/1) SILTSTONE to SANDSTONE, with forams and vol- canic grains. Sandy parts generally
LOWER EDCENE	ELGN				2	multim				laminaied, with an abrupt, erosional or load-casted base. Burrowing common in upper parts of graded layers. Smear Slides: (1-138 cm) base of coarse-grained graded bed: Feldspar A Volcanic Glass A
				P ?	Cat	ore tcher	VOID			Clay C Micrite A (1-136 cm) top of graded bed: Micrite A Nannos A

		Hole						739.0-745.0 m
w	FOSS	TER	NO	8	LITHOLOGY	DEFORMATION	LITH0.SAMPLE	
ZON	FOSSIL ABUND.	PRES.	SECTI	METE				LITHOLOGIC DESCRIPTION
Acarinina densa(?) NP13?	N F	Р	0 1 2 3 4	0.5	V01D			Graded beds, from 1-250 cm thick, typically 10-15 cm, each bed con- sisting of greenish paray (56 6/1 to 56 4/1) CLAYEY LIMESTONE, with a basal part of olive gray (57 4/1) "SILTSTONE or greenish black (56 2/1) "salt-and-pepper" SANDSTONE, commonly laminated or cross-stratified. Burrowing commonly intense in upper parts of graded units; some burrows deformed by shearing. At 0.8-0.95 m: A granule-bearing bed with reverse (?] grading. At 1.9-2.4 m: One graded bed, with coarse-grained sandstone at base. At 3.45-6.0 m: One graded bed, with granules (including large benthonic forans) in basal 50 cm, s.to med, grained next 50 cm, laystone top, with bur- rows in upper 10 cm only. Base not seen. Smear Slides: (1-100 cm): Volcanic Glass A Feldspar C Clay A Micrite R Nannos R



NOTE: The upper scale is GRAPE Wet-Bulk Density (1.0 to 3.0 g/cc): solid lines (_____) are GRAPE analog data assuming a 6.61 cm core diameter; dotted lines (....) are GRAPE analog data adjusted for actual core diameter; circled (\mathfrak{O}) dots are the wet-bulk density calculated from two-minute counts on a stationary sample; the porosity nomogram allows a porosity scale to be determined by selecting the proper grain density (r_g) and extrapolating horizontally.













































SITE 318



