

39. Sr-ISOTOPE AND RARE-EARTH ELEMENT GEOCHEMISTRY OF DSDP LEG 37 BASALTS

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ANALYTICAL METHODS

Strontium Isotopes

Only the freshest samples were selected for study. Each sample leached with 6M HCl before total dissolution in order to remove contaminant Sr derived from seawater. Samples were dissolved in a HF-HNO₃ mixture and Sr was extracted by conventional cation exchange techniques (Pankhurst and O'Nions, 1973). Isotopic analyses were performed on a VG Micromass 30 using techniques similar to those described previously (O'Nions and Pankhurst, 1973).

Rare-Earth Elements

Rare-earth element abundances were determined by mass spectrometric isotope dilution. A mixed-solvent anion exchange procedure (Hooker et al., in press) was employed for the separation of the REE and isotopic analyses were performed on the Oxford 12" mass spectrometer.

RESULTS

Strontium Isotopes

⁸⁷Sr/⁸⁶Sr ratios are presented in Table 1. The samples analyzed cover a range in major element chemistry. For example, TiO₂ contents range from 0.62% to 1.14%. The measured ⁸⁷Sr/⁸⁶Sr ratios for the samples from Hole 332A, however, are indistinguishable and average 0.70296. The five samples analyzed from Hole 332B represent most of the major element variation seen in Leg 37 basalts and have ⁸⁷Sr/⁸⁶Sr ratios ranging from 0.70287 ± 4 to 0.70308 ± 4. This range is a little outside analytical error, but the average value for the five samples is 0.70298. The sample analyzed from Hole 333A has an ⁸⁷Sr/⁸⁶Sr ratio of 0.70304 ± 4 and is indistinguishable from the mean values determined for Holes 332A and 332B. The one sample analyzed from Site 335 yielded an ⁸⁷Sr/⁸⁶Sr of 0.70316 ± 4 and is the highest value determined for a Leg 37 basalt in this study. It is not clear at present whether or not the slightly higher ⁸⁷Sr/⁸⁶Sr ratios observed in this sample and some of those from Hole 332B are simply due to alteration not completely removed by the leaching technique.

Rare-Earth Elements

The abundances of nine rare-earth elements (Table 2) have been determined in four Leg 37 samples. Three of these samples are from Hole 332B and one is from Site 335. The chondrite normalized REE distribution

patterns range from slightly light REE enriched (Site 355 sample) to light REE depleted (Figure 1). Ce_N/Yb_N ratios range from 1.30 to 0.58 and are compared with TiO₂ and Sr contents and ⁸⁷Sr/⁸⁶Sr ratios in Table 3.

DISCUSSION

The presumed magmatic ⁸⁷Sr/⁸⁶Sr ratios of these rocks of 0.70295-70300 is comparable to those obtained for Mid-Atlantic Ridge basalts from the southern section of the Reykjanes Ridge (O'Nions and Pankhurst, 1974) and from part of the Kolbeinsey Ridge north of Iceland (Pankhurst and O'Nions, unpublished). It appears to indicate a uniformity in the Rb-Sr evolution of ocean floor basalts (and the mantle source region from which they are derived) over a large portion of the northern Atlantic Ocean. It is significant, however, that other sections of the Mid-Atlantic Ridge show considerably lower ⁸⁷Sr/⁸⁶Sr ratios, down to 0.7023 (Hart, 1976).

The REE abundances (with the possible exception of Sample 332B-10-3, #2A) are similar to those reported for Mid-Ocean Ridge basalts (e.g., Kay et al., 1970; O'Nions and Pankhurst, in press; Schilling 1973, 1975). Unlike the Reykjanes Ridge and Iceland (O'Nions and Pankhurst, in press) the Leg 37 basalts do not show a correlation between Ce_N/Yb_N and ⁸⁷Sr/⁸⁶Sr. Further the spread of Ce_N/Yb_N values within Hole 332B alone from ~0.6 to ~1.3 is not consistent with what might be expected from Schilling's (1975) plume model for the Azores, since this range corresponds to a considerable portion of the variation which Schilling ascribes to the mixing of two end members.

In conclusion the Sr-isotope geochemistry carried out to date is compatible with all samples being derived from a mantle which is essentially homogeneous isotopically. It is unlikely that the range of Ce_N/Yb_N ratios could be solely due to differential partial melting of a source which is homogeneous with respect to these elements. Some minor heterogeneity in REE distribution or low-pressure crystal fractionation may be responsible.

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TABLE 1
Sr-Isotope Data^a

Hole 332A					
Sample	8-2, #23	12-1, #9	17-1, #6	29-1, #9	
$^{87}\text{Sr}/^{86}\text{Sr}$	0.70302 ±4	0.70296 ±4	0.70300 ±4	0.70295 ±6 0.70299 ±4	
Hole 332B					
Sample	10-3, #2A	16-2, #8	25-1, #16	35-3, #4	37-3, #6
$^{87}\text{Sr}/^{86}\text{Sr}$	0.70296 ±7 0.70303 ±8	0.70292 ±4	0.70287 ±4	0.70308 ±4	0.70295 ±4
Hole 333A					
Sample	11-1, #12				
$^{87}\text{Sr}/^{86}\text{Sr}$	0.70304 ±4				
Site 335					
Sample	14-4, #10				
$^{87}\text{Sr}/^{86}\text{Sr}$	0.70316 ±4				

^aAll data relative to 0.70800 for Eimer and Amend SrCO₃ standard.

TABLE 2
Rare-Earth Element Concentrations

	La	Ce	Nd	Sm	Eu	Gd	Dy	Er	Yb
Hole 332B									
10-3, #2A	—	12.00	8.33	2.45	0.91	3.12	3.88	2.36	2.36
25-1, #16	—	6.41	5.00	1.71	0.70	2.52	3.19	2.10	2.04
35-3, #4	—	4.59	3.44	1.18	0.50	1.81	2.34	1.59	1.61
Site 335									
14-4, #10	—	6.52	6.22	2.37	0.91	3.57	4.38	2.94	2.87

Note: — = not determined.

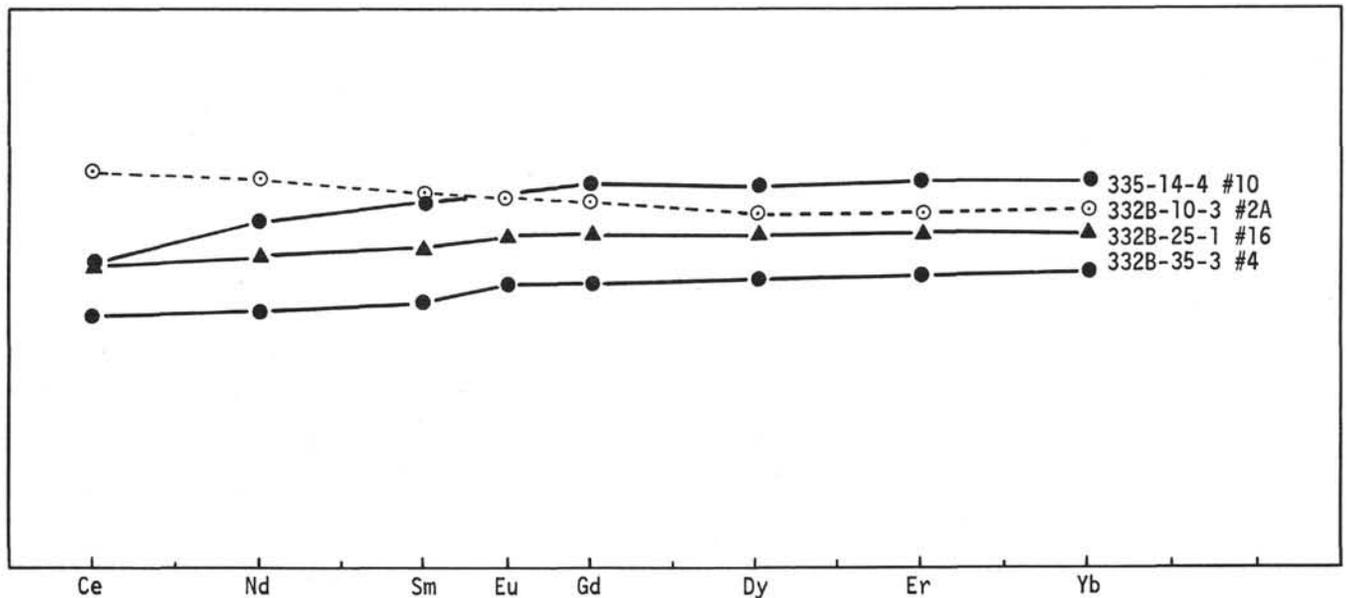


Figure 1. Chondrite normalized REE distribution patterns for Leg 37 basalts.

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TABLE 3
 Ce_N/Yb_N Ratios Compared with $^{87}Sr/^{86}Sr$, TiO_2 , and Sr

Sample	Ce_N/Yb_N	$^{87}Sr/^{86}Sr$	TiO_2 (wt %)	Sr (ppm)
Hole 332B				
10-3, #2A	1.30	0.70300 ± 7	1.05	112
25-1, #16	0.80	0.70287 ± 4	0.82	87
35-3, #4	0.72	0.70308 ± 4	1.12	100
Site 335				
14-4, #10	0.58	0.70316 ± 4	1.04	—

Note: = not determined.

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