8. SITE 396: 23 °N, MID-ATLANTIC RIDGE

Shipboard Scientific Party¹

SITE DATA

Date Occupied: 11 January 1976 (0200)

Date Departed: 14 January 1976 (2230)

Time on Hole: 3 days, 20 hours, 30 min

Position: 22°58.88 'N, 43°30.95 'W

Water Depth (sea level): 4450 corrected meters, echo sounding

Water Depth (rig floor): 4460 corrected meters, echo sounding

Bottom Felt at: 4460 meters, drill pipe

Penetration: 221.49 meters

Number of Holes: 1

Number of Cores: 25

Total Length of Cored Section: 221.49 meters

Total Core Recovered: 133.28 meters

Percentage Core Recovery: 60 per cent

Oldest Sediment Cored:

Depth sub-bottom: 125 meters Nature: Foraminifer-nannofossil ooze Chronostratigraphic unit: middle Miocene Measured velocity: 1.5-1.8 km/sec

Basement

Depth sub-bottom: 96 meters sub-basement Nature: Basalt

Principal Results: Site 396 was drilled at latitude 22°58.88 'N, longitude 40°30.95 'W, in a water depth of 4450 meters. The site is about 1.4 km from the western edge of a sediment pond, Atlantic survey Site AT-6, within normalpolarity magnetic anomaly 5 (8.7 to 10 m.y.). It is at the approximate conjugate position, on the opposite side of the ridge axis, to Site 395. The regional bathymetry and the magnetic anomalies are well lineated, and no evidence of fracture zones is apparent in the vicinity of the sediment pond. One hundred twenty-five meters of sediment were continuously cored; the basalt/sediment contact was recovered in undisturbed form. The sediments above the basalt contact are tentatively assigned to the *Discoaster exilis* Zone (middle Miocene; ~ 14 m.y.). The sediments can be divided into two units. The first, from 0 to 117 meters subbottom, is a nannofossil ooze interbedded with lesser foraminifer-nannofossil ooze and foraminiferal sands. The second, from 117 to 125 meters, is brown calcareous basal clay interbedded with marly nannofossil ooze. Recovery in the sediments was 80.7 per cent.

We drilled 96 meters beneath the sediments, at which depth the bit became worn. Basement recovery was 33 per cent. The cored rocks consist lithologically of fairly uniform phyric basalt pillows with a multitude of glass zones, lithified sediment veins, and contact zones. The lava is plagioclase (5 to 15%) and olivine (1 to 3%) phyric. Clinopyroxene is not a phenocryst phase. The groundmass of all the rocks is fine grained, often spherulitic, with feldspar laths around 0.1 mm long. There is no basis for discrimination of lithologic units from hand-specimen petrology. Two varieties of carbonate veins cross many of the samples. The first is a lithified (baked?) sedimetary carbonate, observed only in the upper three cores; the second is a recrystallized white vein carbonate. The presence of the sedimentary carbonate in the basalts suggests that the upper part of the sequence was intruded into or extruded onto carbonate ooze.

Despite the hand-specimen and petrographic uniformity of the basalts, compositional variation is significant, particularly in Mg/(Mg + Fe) values (0.58 to 0.66), Ni concentrations (111 to 189 ppm), and in Sr values. These variations do not simply reflect varying proportions of modal phenocryst content, since A1₂O₃ and CaO concentrations are approximately constant (16.7% and 12.0%, respectively). Most of this variation can be attributed to fractional crystallization involving olivine and subordinate plagioclase. Three comagmatic sequences are observed, differentiated primarily by abundance of Sr. One is the topmost basement core. The second is from the base of this unit (127 m sub-bottom) to 191 meters, and contains Sr in concentrations of 154 ppm (average). Variations in TiO₂, Ni, and Zr within this inverval suggest a single comagmatic sequence involving olivine control. The third sequence is from 201 meters to the bottom of the hole (221 m), and contains Sr in concentrations of 131 ppm (average).

The magnetic inclinations fall into three separate groups. The upper group, identified by only two samples, from near the top of basement, is reversed at a shallow angle of -20° . The middle group, which extends to about 70 meters sub-basement, is scattered around an average inclination of $+35^{\circ}$. The lower magnetic group has average inclinations of -5° , and coincides with the lower geochemical unit. The average intensity of the entire 96 meters of sub-basement is $1.53 \pm 0.15 \times 10^3$ emu/cm³. This value is less than half that obtained at Hole 395A, and suggests either that the basaltic pillows of Site 396 are more extensively oxidized than the units of Hole 395A, or perhaps that this section was extruded during a magnetic transition.

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In summary, three geochemical units and three magnetic units occur in Hole 396. The geochemical breaks coincide with the magnetic breaks.

OPERATIONS

Atlantic survey area AT-6 was surveyed by R/V Atlantis II (WHOI-M. Purdy, Chief Scientist) and R/V Akademik Kurchatov (USSR). The results of these excellent site surveys (Purdy et al., this volume) enabled us to locate the optimum drilling location in minimum time, aboard Glomar Challenger. The reasons for selecting the particular sediment pond we drilled are given in Chapter 16, this volume. The sediment pond is situated near the center of the well-developed magnetic anomaly 5 (10 m.y.) east of the Mid-Atlantic Ridge axis. Both the regional topography and the magnetic anomalies are characterized by long linear troughs and highs. There is no evidence for fracture zones either in the magnetics or the topography. Seismic refraction (sonobuoy) results over the sediment show seismic structure representative of "normal" ocean crust of this age (~400 m of crust with Layer 2A velocities, ~1.5 km of crust with Layer 2B velocities) although the results are not ideal (Purdy, this volume). The topography, together with the location of the sediment pond and extent of crustal blocks causing the magnetic anomalies, is given in Figure 1. Super-imposed is the track made by Glomar Challenger in approaching the site and dropping the beacon.

The site was selected as a prime location for deep crustal drilling. On Leg 45, sufficient time was not available to begin drilling a deep hole. Our primary purposes for going to survey area AT-6 were to drill a pilot hole that would provide Leg 46 scientists with additional information regarding the feasibility of deep drilling, and to compare results here with our results on Site 395 drilling. Since the ultimate aim at this site was to study the ocean crust beneath the sediments, we tried to select the site in a region of minimum sediment thickness. As at Site 395, the eventual plan was (if the pilot hole were successful) to set a re-entry cone and case to basement. Minimum sediment thickness was therefore highly desirable. It was an operational necessity, however, to have about 100 meters of sediment for spudding in and supporting the bottom-hole assembly. Since magnetic anomaly 5 encompasses the entire sediment pond, we were not confined in our drilling to any particular location within the pond, as we were at Site 395. Seismic reflection profiles (Figure 2) show a two-way reflection time of 0.3 sec near the center of the pond. We chose the western edge of the pond, where a small step in basement or basement high minimized the sediment overburden. We dropped the beacon on the first pass over the western edge of the sediment pond (previously collected site survey data enabled us to do this), where sediment thickness corresponds to 0.15 sec two-way travel time (135 m if we assume $\overline{v} = 180$ km/sec). We then steamed across the sediment pond, reversed course to obtain reflection records while the beacon was on the sea floor, and retrieved geophysical gear; we did this to confirm sediment thickness.

SEDIMENTS

Lithologic Description

We cored sediments continuously from 4450.00 meters below sea level to 4575.24 meters at Site 396, and recovered 101.19 meters (80.7%) in 14 cores. Core 4 had no recovery, because the plastic sack in the core catcher broke and the sediments washed out as the core barrel was pulled to the ship. Excluding Core 4, recovery was 87.4 per cent. The sediments range from Pleistocene to middle Miocene above basalt. Sediments in the core catcher of Core 21 within the basalt are upper Miocene, and apparently fell down the hole.

Hole 396 is on the western edge of a small northerly trending sediment pond about 3×10 km in dimensions and 300 meters deep in its deepest part (Purdy, et al., this volume). The pond is somewhat larger than that cored at Site 395, and lies between two prominent northtrending ridges in a region of generally north-lineated topography. The bounding ridges rise to about 1000 meters above the sediment pond, but relief is generally lower and not as steep as at Site 395. The ridges, nevertheless, are in many places free of sediment. The sediments at Site 396 have therefore largely slumped or been swept into the pond from the adjacent ridges.

The sediments can be divided into two units, as follows.

UNIT I (Cores 1 through 13, Section 6, 145 cm, 0 to 1117.15 m sub-bottom, 4450.00 to 4567.15 m below sea level) is nannofossil ooze interbedded with lesser

TABLE 1 Coring Summary, Site 396

Core	Date (January 1976)	Time	Depth From Drill Floor (m)	Depth Below Sea Floor (m)	Length Cored (m)	Length Recovered (m)	Recovery (%)
1	11	1525	4460.00-4465.00	0.00-5.00	5.00	5.00	100
2	11	1712	4465.00-4474.28	5.00-14.28	9.28	7.30	79
3	11	1830	4474.28-4483.84	14.28-23.84	9.56	5.54	58
4	11	1950	4483.84-4493.10	23.84-33.10	9.26	0.00	0
5	11	2110	4493.10-4502.37	33.10-42.37	9.27	9.00	97
6	11	2210	4502.37-4511.92	42.37-51.92	9.55	8.00	84
7	11	2320	4511.92-4521.47	51.92-61.16	9.24	9.15	99
8	12	0102	4521.47-4530.72	61.16-70.41	9.25	8.10	88
9	12	0200	4530.72-4539.95	70.41-79.64	9.23	6.65	72
10	12	0310	4539.95-4549.18	79.64-89.18	9.54	8.27	87
11	12	0412	4549.18-4558.41	89.18-98.41	9.23	8.60	93
12	12	0525	4558.41-4567.96	98.41-107.96	9.55	8.20	86
13	12	0640	4567.96-4577.20	107.96-117.20	9.24	9.24	100
14	12	0825	4577.20-4586.74	117.20-126.74	9.54	9.54	100
15	12	1315	4586.74-4596.05	126.74-136.05	9.31	4.95	53
16	12	1800	4596.05-4605.30	136.05-145.30	9.25	5.83	63
17	12	2240	4605.30-4614.55	145.30-154.55	9.25	0.00	0
18	13	0530	4614.55-4623.56	154.55-163.56	9.01	4.78	53
19	13	1005	4623.56-4632.82	163.56-172.80	9.26	3.00	32
20	13	1215	4632.82-4642.06	172.80-182.06	9.24	0.00	0
21	13	1340	4642.06-4651.33	182.06-191.33	9.27	1.03	11
22	13	1730	4651.33-4660.59	191.33-200.59	9.26	5.70	62
23	13	2030	4660.59-4670.24	200.59-210.24	9.65	0.70	7
24	13	2310	4670.24-4679.49	210.24-219.49	9.25	3,60	39
25	14	0310	4679.49-4681.49	219.49-221.49	2.00	1.10	55
Total					221.49	133.28	60



Figure 1. Bathymetry and magnetics of survey area AT-6, with Glomar Challenger approach track superimposed. Sediment pond is shaded.



Figure 2. Approach profile of Glomar Challenger to Site 396.

foraminifer-nannofossil ooze and rare foraminiferal sands. Colors range from pale brown to pale yellowish brown, and are slightly darker in Cores 1 through 4 than in 5 through 13. The lower part of Core 13 darkens to yellow-brown toward the top of Unit II. The darker colors result from a larger proportion of brown clays and associated reddish hydrous iron oxides. A small deformed patch of brown clay occurs in Core 5, Section 3.

Deformation caused by drilling is intense throughout Unit I. In general, sedimentary structures are preserved only in the lowest sections of each core. Beds and traces of beds are evident in Cores 6, 7, and 10, but are not very prominent. Core 11, Section 4, and Core 13, Sections 5 and 6, have beds every 10 to 50 cm, characterized by faint color and slight variations in "grittiness" (proportion of foraminifers to nannofossils). Each bed is very sharply defined. Bioturbation apparently did not occur. Throughout Unit I, both foraminifers and nannofossils occur. Traces of hydrous, amorphous-looking iron oxides, sponge spicules, and fish debris are present in most smear slides, but heavy minerals and volcanic glass are notably absent. The only significant variation among the components of the sediments occurs in local foraminiferal sands, and in a noticeable general increase in the proportion of foraminifers in Cores 7 and 8; this proportion drops sporadically from there through Core 14.

The general characteristic of Unit I is its uniformity. Partly, this is because structures have been obliterated by drilling deformation. But mostly, the proportion of foraminifers to nannofossils varies little in the less deformed sections. The contrast to Site 395 is marked in this respect. The sequence at Site 395 grades downward from foraminifer-nannofossil ooze (5 to 20%foraminifers), with interbedded foraminiferal sands, to nannofossil oozes. In Hole 396, this sequence is reversed: nannofossil oozes with less than 1 per cent foraminifers abruptly grade into foraminifernannofossil ooze with 5 to 20 per cent foraminifers in Core 7.

At this time, we can only speculate about the reasons for these differences. Perhaps different histories of subsidence, relative to the foraminiferal lysocline at Sites 395 and 396, were involved. The average distance of sediment transport from adjacent ridges was also probably different at the two sites. This, coupled with differences in slopes of the ridges, would cause foraminifers and nannofossils in sediments slumped from the ridges or carried from them in turbidity currents to sort differently. The lack of well-preserved obvious graded beds in both the Site 395 and the Site 396 cores, however, makes this interpretation difficult to assess without detailed grain-size analysis. But the average finer grain size of Site 396 sediments, together with the lack of volcanic glass, heavy minerals, and basaltic sand grains (all present at Site 395), suggests that Site 396 is farther from any possible turbidite sources than Site 395. Site 396 is about 1.4 km from the closest edge of the sediment pond, compared with about 0.3 km at Site 395.

UNIT II (Core 13, Section 6, 145 cm to Core 14, Section 6, 10 cm: 117.15 to 125.34 m sub-bottom, 4567.15 to 4575.34 m below sea level) is dark yellowish brown to dark brown calcareous basal clay interbedded with pale yellowish brown marly nannofossil ooze. Foraminifers are scarce in the oozes and virtually absent in the clays. Staining of coccoliths with amorphous hydrous iron oxides is intense even in the layers less rich in clay. Bioturbation is not evident. The clays contain as little as 15 per cent CaCO₃ (carbonate bomb analysis), and are quite stiff. They are generally fragmented and brecciated by drilling. The marly nannofossil oozes are relatively soft and pliant. Traces of manganese oxide micronodules are present. A well-preserved fish tooth about 5 mm long was found in brown clay in Core 14, Section 1.

The clays are most prevalent in Core 14, Sections 1 through 3, where they have been drastically mixed with marly nannofossil ooze. Coherent clay pieces up to 10 cm long occur in this mix. Sections 4 and 5 are homogeneous marly nannofossil oozes. The sediment/basalt contact at the top of Section 6 is in this ooze. Sediments were pressed about 50 cm farther into Section 6 during drilling, however; rounded to angular basalt pieces are embedded in deformed ooze to that point.

The clays in Cores 13 and 14 are typical in appearance of basal Fe-Mn-rich clays at many DSDP sites, and are very similar to basal clays at Site 395. These clays are considered to be precipitates from hydrothermal metal-rich solutions emanating from the basement rocks (Timofeev et al., this volume), though they have rarely been described in the context of the Mid-Atlantic Ridge.

Sediments in Basalt

The topmost basalts drilled at Site 396 are in intimate contact with recrystallized calcareous sediments in cracks and with largely palagonitized glass breccias spalled from pillows. The limestones are very pale yellow to very pale brown and occur in 0.2- to 5-cm cracks and cavities. Recrystallization is essentially complete, but ghosts of foraminifer tests are visible in some thin sections. There is no trace of nannofossils.

Basalt pillows apparently intruded the sediments while the sediments were extremely fluid; they were probably no more than a few meters thick at the time. Contraction cracks in the basalt apparently sucked up sediment as they formed. Delicate flow patterns in these cracks are preserved in the lithified sediments. The sediments do not appear to have contained large percentages of Fe and Mn at the time of basalt intrusion. Core 21 penetrated a small pocket of sediments. Only a small amount of this material was recovered. It is a pale yellow brown foraminifer-nannofossil ooze, and contains upper Miocene microfossils that apparently fell down the hole.

Biostratigraphy (by Ansis Kaneps, DSDP)

The physiographic setting of Site 396 is analogous to that of Site 395, but the site is situated on the opposite (east) side of the Mid-Atlantic Ridge crest. The section penetrated also consists of foraminifer-nannofossil oozes, but unexpectedly the oldest sediments cored are middle Miocene (13.6 \pm 1.6 m.y.) rather than upper Miocene, as was predicted by sea-floor spreading anomaly ages. There is no explanation for this discrepancy at this time. As at Site 395, reworking and drilling disturbance complicate the biostratigraphic sequence.

The oldest sediments cored (Core 14) belong to the Discoaster exilis Zone and the Globorotalia fohsi fohsi Zone. Overlying these sediments (with probable disconformity) is lower Pliocene (Cores 12 and 13), for which assignment to the Ceratolithus rugosus Zone is suggested (~4 to 4.4 m.y.). This is in turn disconformably overlain by upper Pliocene sediments (Cores 8 through 11), characterized by the foraminifer species Globorotalia miocenica, along with other upper Pliocene markers, such as G. tosaensis, G. patenis, G. exilis, and Globigerinoides fistulosus. The remainder of the section is lower Pleistocene, except for a thin upper Pleistocene and/or Recent sequence in Core 1. Reworked Pliocene species (Globoquadrina altispira, Globorotalia multicamerata, Sphaeroidinellopsis spp.) occur as high as Core 5 (some 25 m above the Pliocene/Pleistocene contact). The bulk of the section at Site 396 appears to have been deposited during late Pliocene through early Pleistocene time (between about 2.7 and 1.0 m.y.); the remainder of middle Miocene through Recent time is poorly represented. This abnormal thickening of the upper Pliocene and lower Pleistocene, together with the evidence for extensive reworking, strongly suggests concentration of sediments, by redeposition from surrounding hills, in the basin in which Site 396 is situated.

Magnetics

No sediments at Site 396 were sufficiently undisturbed to bother taking magnetic measurements.

Physical Properties and Chemistry

On Figure 3 are plotted variation with depth of porosity, grain density, water content, and CaCO₃ content (as determined using the carbonate bomb). GRAPE density and porosity were determined only on relatively undisturbed cores (from Core 5), but are not plotted here. In general, as at Site 395, GRAPE density increases and porosity decreases with sediment compaction. As at Site 395, separate cube porosity measurements were made on least-disturbed samples, which should more accurately reflect *in situ* conditions. As at Site 395, porosity and water content are fairly con-



Figure 3. Physical properties and CaCO3 contents of sediments versus depth, Hole 396.

Grain densities given are densities of the sediments determined dry, so that most of the effects of drilling disturbance and addition of water have been eliminated. The small "spike" in the plots of grain density and carbonate content (Core 5) correspond to the small patch of calcareous red clay described earlier.

The percentage of CaCO₃ is somewhat low in Core 1, increases to between 80 and 90 per cent from Core 5 through Core 13, and drops sharply in the basal clays. These clays are much lower in CaCO₃ than those of Site 395, but recovery at Site 396 was more complete. We attribute the low CaCO₃ in Core 1 to a decline in the rate of carbonate sedimentation, resulting from subsidence of the sea floor closer to the lysocline. The content of pelagic clays is higher as a result.

The water contents, porosities, and saturated bulk densities of the sediments from Hole 396 (27 samples) are given in Table 2.

IGNEOUS ROCKS

Lithologic Summary

Hole 396 was drilled for 96 meters sub-basement. The cored rocks consist of a fairly uniform sequence of phyric basalt pillows, with a multitude of glass zones and lithified sediment veins and contact zones. The lava is plagioclase and olivine phyric; clinopyroxene is not a phenocryst phase, and is generally spherulitic to

TABLE 2 Physical Properties of Sediments, Site 396

Sample (Interval in cm)	% Water Content (E/A)	Porosity (E/D)	Wet-Bulk Density (g/cm ³) (A/D)
1-2, 19-21	0.461	0.701	1.52
2-6, 65-67	0.362	0.610	1.69
3-4, 98-100	0.346	0.589	1.70
5-1, 104-106	0.358	0.603	1.69
5-3, 105-107	0.351	0.598	1.70
5-3, 125-127	3.55	0.604	1.70
5-5, 60-62	0.358	0.605	1.69
6-3, 24-26	0.386	0.632	1.64
6-6, 96-98	0.386	0.632	1.64
7-3, 69-71	0.368	0.611	1.66
7-6, 142-144	0.343	0.588	1.71
8-5, 46-48	0.380	0.626	1.65
8-6, 92-94	0.395	0.654	1.65
9-4, 90-92	3.52	0.599	1.70
10-2, 26-28	0.350	0.596	1.71
10-5, 23-25	0.361	0.613	1.70
11-4, 92-94	0.352	0.599	1.70
12-2, 40-42	0.348	0.592	1.70
12-4, 126-127	0.345	0.590	1.71
13-2, 47-49	0.338	0.581	1.72
13-4, 52-54	0.362	0.607	1.68
13-6, 121-123	0.338	0.582	1.72
14-1, 119-121	0.426	0.675	1.58
14-2, 47-49	0.416	0.662	1.59
14-3, 75-77	0.455	0.703	1.55
14-4, 98-100	0.402	0.649	1.61
14-5, 136-138	0.378	0.625	1.66

skeletal, even in the groundmass. The proportion of phenocrysts varies from about 7 to about 15 per cent; of this only about 1 to 3 per cent is olivine, and the rest is plagioclase. The feldspar phenocrysts vary from 0.5 to 3 mm in length, and are often zoned, but the core composition is about An₇₆. They are generally euhedral; coexisting with them are a few very rounded plagioclases, quite possibly xenocrysts, a very small proportion of which contain brown-red spinel. The same is true of the olivines: some rounded crystals include spinel, but generally the olivines are altered in large degree to a brown iddingsitic material. Both the rounded and subrounded olivines have a 2V~90° (~Foss), and are around 1 to 2 mm in diameter. The spinels present are usually included in feldspar phenocrysts, but do occur rarely in the groundmass. They are deep red-brown in the upper part of the core, but toward the bottom they become browner and lose the ruby-red tinge shown by the spinels in the upper cores. The groundmass of all these rocks is fine grained, often spherulitic-with feldspar laths around 0.1 mm long-and dendritic or spherulitic clinopyroxene. None of the basalts have ophitic texture. The clear, very pale brown glass zones contain phenocrysts of olivine and plagioclase.

Two varieties of carbonate veins cross many of the samples. The first is lithified (baked?) sedimentary carbonate. Veins of this material are pale brown, in parts are relatively uncrystallized, and may contain rare patches of virtually unaltered foraminifer-nannofossil ooze. This carbonate type occurs as a coating on glassy surfaces, as veins 0.2 to 1.0 cm wide, and as "blebs" in some of the basalt samples. It is restricted to the upper cores, 14, 15, 16. The second type is a characteristic recrystallized white vein carbonate. It occurs generally in much thinner veins, about 1 mm wide or less. The presence of included sedimentary carbonate in these basalts suggests that the upper part of the sequence was intruded into a carbonate ooze.

Although there are three chemical units, coinciding with three magnetic units, the lithology—to the extent examined on shipboard—is uniform and not divisible into more than one unit.

Geochemistry

Basalts were recovered to a total sub-basement depth of 96 meters. All are extrusive, petrographically uniform, plagioclase-olivine phyric basalts. Total phenocryst abundance varies along the core from 5 to 15 per cent; plagioclase and olivine phenocryst contents average about 10 and 2 per cent, respectively. All the samples are altered, and show yellow-brown discoloration developing from cracks. Many of these cracks contain carbonate. Samples selected for chemical analysis (Appendix I, this volume; Bougault et al., this volume) were chosen to avoid veins and yellowish brown alteration zones. Even so, almost all are moderately altered, and undergo relatively high loss on ignition, yielding values from -2.0 to 3.1. These values are similar to the most altered samples analyzed from Hole 395A.

Despite the petrographic uniformity of these basalts, three different units have been defined: P_a , P_b , and P_c . The subscripts a, b, and c (rather than 1, 2, and 3, conventional for Site 395) were chosen because the causes of geochemical variations in Hole 396 are less certain than at Site 395; there is more overlap, and combinations of partial melting, fractionation, and mixing could explain them. The groups are more distinct magnetically than geochemically (see Paleomagnetism, this chapter). The averages of the analyses of these units and their standard deviation (σ) values are reported in Table 3. The boundary between P_a and P_b is situated between Samples 14-6, 97-99 cm and 14, CC, 17-19 cm, and the boundary between P_b and P_c is between Samples 22-3, 108-110 cm and 22-4, 135-137 cm.

Composition varies significantly among these basalts, particularly in Mg/(Mg + Fe) values (0.58 to 0.66) and Ni concentrations (111 to 189 ppm). These variations do not simply reflect varying proportions of modal phenocryst content, since Al₂O₃ and CaO concentrations are approximately constant (e.g., 16.7%) and 12.0%, respectively). Most of the variation can be attributed to fractional crystallization involving olivine and subordinate plagioclase. Thus, the sequence of increasing TiO₂ and Zr with decreasing MgO, Mg/(Mg +Fe), and Ni concentrations, in samples from Core 21 through Core 15 to Cores 16, 18, and 19 (Bougault et al., this volume), suggests a single comagmatic sequence involving olivine control. The sample from Section 14-6 is probably a more evolved member of this sequence (Mg/[Mg + Fe] = 0.58), derived by plagioclase and olivine (and pyroxene?) fractionation, with plagioclase the dominant phase (e.g., plagioclase: olivine = 80:20).

TABLE 3 Averages of Chemical Analyses (%) of the Three Basaltic Units Defined

	for Site		
	Pa	Pb	Pc
Unit	1	4	4
SiO ₂	49.3	49.68	49.45
2		0.13	0.35
Al203	16.25	16.78	16.73
2 0		0.10	0.15
$Fe_2O_3(t)$	9.96	9.33	9.48
2 0		0.24	0.16
FeO	5.36		
MnO	0.16	0.16	0.16
MgO	7.0	8.17	7.55
		0.31	0.26
CaO	11.57	11.98	12.03
		0.09	0.07
Na ₂ O	2.40	2.41	2.46
-		0.07	0.03
K20	0.31	0.24	0.21
<i></i>		0.04	0.05
TiO ₂	1.50	1.27	1.29
2		0.02	
P205	0.15	0.12	0.11
2 0		0.01	0.01
Total	98.60	100.14	99.47
LoI	-2.0	-2.8	2.25
		0.3	0.21

As already mentioned for Site 395 (Chapter 7, this volume), a more complex process (or processes) is required to explain Sr behavior (average of 131 ppm in Unit P_c and 156 ppm in Unit P_b). This process could be alteration or mobilization of Sr, non-equilibrium of Sr between liquids and fractionating plagioclases, etc. On the other hand, the constancy of Zr/Ti implies a homogeneous source with respect to these two elements.

Paleomagnetism in Basalts, Hole 396

The sampling and measurement techniques for paleomagnetism at this hole were identical to those described for Hole 395A. All samples appeared to be plagioclase phyric pillow basalts, with no obvious grouping into individual units by lithology. The relatively low values of intensity and the high median destructive field shown in Table 4 are consistent with this observation. The integrated average magnetization intensity of the entire 100 meters of sub-basement basalt is 1.53 $\pm 0.15 \times 10^{-3}$ emu/cm³. This value is less than half the value of 3.78×10^{-3} emu/cm³ obtained for Hole 395A. This may be because the basaltic pillows of Hole 396 were more extensively oxidized than the units of Hole 395A, or, perhaps because this section of igneous crust was extruded during a magnetic transition. There is no obvious trend in intensity with depth. Sample 396-18-1, 109-111 cm (Table 4) was specifically sampled in an oxidized and veined zone. Clearly, the effect of alteration (presumably low-temperature oxidation) is to reduce the intensity of magnetization. This result is consistent with previous studies of low-temperature alteration of submarine basalts.

The stable magnetic inclinations occur in three separate groups defined by depth. The upper group, consisting of two samples from Core 14, Section 6, are reversed at a shallow angle of -20° . The middle group, which consists of most of the samples in the hole, is scattered around an average inclination of $+35^{\circ}$, and corresponds to chemical unit P_b. The lower magnetic unit, with inclinations close to -5° , coincides with the bottom geochemical unit in this hole.

The causes of the low inclinations may be (1) secular variation of the earth's magnetic field, with three separate eruptive events; (2) fault-block rotation; or (3) combinations of the above (Johnson, Chapter 14, this volume). The low intensities of magnetization, compared with Site 395, appear to reflect (1) increased oxidation of the magnetic minerals in the older Site 396 basalts, and (2) reduced availability of the magnetic mineral (Johnson, Chapter 15, this volume).

Physical Properties of Basement Rocks

The saturated bulk densities of 28 basalt samples from Site 396 are given in Table 5. In general, the densities are uniform and comparable to those of the aphyric basalt units of Site 395.

TABLE 4 Paleomagnetism Data, Site 396

Sample (Interval in cm)	Stable Inclination (°)	Intensity of Remanence (× 10 ⁻³ emu/cm ³)	Median Demagnetizing Field (Oe)	Quality of Orientation ^a	Comments
14-6, 14-16	-21.8	1.52	700	М	Phyric pillow
14-6, 97-99	-19.8	1.21	550	Μ	Phyric pillow
14, CC, 17-19	+41.5	1.55	490	MG	Phyric pillow
15-2, 91-93	+38.2	1.64	700	M	Phyric pillow
15-3, 138-140	+36.1	1.51	425	G	Phyric pillow
15-4, 50-52	+37.0	.788	550	Μ	Phyric pillow
16-1, 90-92	+35.0	1.51	350	MG	Phyric pillow
16-2, 57-59	+36.4	1.43	500	MG	Phyric pillow
16-3, 68-70	+33.1	.926	550	Μ	Phyric pillow
16-4, 91-93	+35.2	1.27	300	MG	Phyric pillow
18-1, 68-70	+33.8	.944	550	Μ	Phyric pillow
18-1, 109-111	+37.7	.251	450	MG	Oxidized phyric pillow
18-2, 42-44	+32.8	1.26	350	Μ	Phyric pillow
18-3, 124-126	+27.9	1.53	225	G	Phyric pillow
18, CC, 10-12	+29.4	1.16	450	Μ	Phyric pillow
19-1, 22-24	+29.4	1.55	250	G	Phyric pillow
19-2, 48-50	+32.0	1.64	510	M	Phyric pillow
21-1, 70-72	+34.3	.911	550	G	Phyric pillow
22-1, 139-141	+33.3	.811	700	MG	Phyric pillow
22-2, 131-133	+34.5	1.42	550	MG	Phyric pillow
22-3, 108-110	+36.6	.750	700	MG	Phyric pillow
22-4, 135-137	-4.8	4.53	450	G	Phyric pillow
22, CC, 138-140	-2.5	1.61	450	Μ	Phyric pillow
23-1, 67-69	-4.2	2.05	450	М	Phyric pillow
23-1, 102-104	-4.4	2.16	450	М	Phyric pillow
24-1, 107-109	-8.3	1.11	900	MG	Phyric pillow
24-2, 80-82	-6.0	1.88	450	G	Phyric pillow
24-3, 138-140	-6.9	2.83	450	G	Phyric pillow
25-1, 111-113	-4.2	2.39	450	G	Phyric pillow

 ^{a}M = moderate, MG = moderately good, G = good.

TABI	E 5	
Bulk Densities of	Basalt,	Site 396

Core	Section	Interval (cm)	Wet-Bulk Density g/cm ³ (A/D)
14	6	14-16	2.77
14	6	97-99	2.82
14	CC	17-19	2.81
15	2	91-93	2.82
15	3	138-140	2.86
15	4	50-52	2.84
16	1	90-92	2.84
16	2	57-59	2.84
16	3	68-70	2.82
16	4	91-93	2.84
18	1	68-70	2.78
18	1	109-111	2.71
18	2	42-44	2.79
18	3	124-126	2.86
18	CC	10-12	2.82
19	1	22-24	2.80
19	2	48-50	2.85
21	1	70-72	2.77
22	1	139-141	2.81
22	2	131-133	2.82
22	3	108-110	2.81
22	4	135-137	2.86
22	CC	138-140	2.82
23	1	102-104	2.77
24	1	107-109	2.76
24	2	80-82	2.85
24	3	138-140	2.85
25	1	111-113	2.85

Sit	e 396	Hole	6		Cor	e 1	Cored In	nterv	a1:	1460-4465 m (0-5 m subbottom)	Site	396	96 Hole		C	ore 2	Cored Int	erval	: 4465-4475.28 m	(5.00-14.28 m subbottom)
AGE	ZONE	FORAMS	FOSS HARAC SONNYN	IL TER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FORAMS	FOSSIL CHARACTE SONNYN	R	SECT TON METERS	L1THOLOGY	DEFORMATION	1111.509712	LITHOLOGIC DESCRIPTION
PLEISTOCINE	ti E. ovata Mixed G. oceanica-E. huxleyi	6, calida calida 6, calida calida est	s in (Chapt	0 1 2 3 4 5 Co Ca		VOID		· cc ·	<pre>DVRAMINIFER-NANNOPOSSIL 002E Light yellowish brown (10VR 6/4) to very pale brown (10VR 7/4) foraminifer-nannofossil ope, highly disturbed, or soupy. Small bed of foram sand in Section 5. Smoor Slides: 1-63, 2-46, 3-75, 4-140, 5-50, 5-77, CC Porams 4-40% Fe-0x Tr Nannos 55-80% Clay 5-10% Diatoms Tr Sponge Spic. Tr Smoor Slides: (all but 5-77) 10VR 5/4 (all but 5-77) 5-77 10VR 6/4 Silt 0% Sand 20% Silt 0% Silt 10% Clay 93-100% Clay 70% Carbon Carbonate 2-100 9.7 0.1 80 Grain Size 10VR 7/4 4-144 0.8 12.4 86.8</pre>	PLETSTOCENE	Nixed G. caribbeanica G. crassaformis viola G. crassaformis hesi				0 0.5 1 1.0 2 2 2 2 3 3 3 5 5 6	VOID VOID VOID VOID VOID VOID VOID	× 0000000 × × × × × × × × × × × × × × ×	10YR 6/4 10YR 7/6 10YR 6/4 10YR 6/4 10YR 6/4	<pre>FORMINITER-MANNOFOSSIL OOZE Soupy to highly deformed light yellowish brown (10TB 6/4) to brown (10TG 7/6) formaminfer-nannofossil ooze</pre>

	715	FOSSIL CHARACTER	200	MPLE	
AGE	ZONE	FORAMS NANNOS	TITHOLOGY	LITH0.5A	LITHOLOGIC DESCRIPTION
PLEISTOCENE	Nixed G. caribbeanica G. crassaformis viola		0 VOID 0.5 1 1.0 2 2 3 4	* 20 8. *	FORAMINIFER-NANNOFOSSIL 002E 10YR 6/4 Soupy to highly deformed light yellowish brown (10YR 6/4) or very pale brown foramin ifer-nannofossil coze. 10YR 6/4 Smear Slides: 1-90, 2-80, 4-100, CC Forams Tr - 1% Nannos 99% Fe-0% Fe-0% Tr Sponge Spic. Tr Fish Debris 10YR 6/4 Sand 0% Silt Tr - 1% Clay 99-100% Carbon Carbonate 2-100 10.3 0.0 86 Grain Size 4-50 0.1 15.7 84.2 10YR 6/4 Site 396, Core 4, 4483.84-4493.10 = (23.85-33.10 m) 10
			4	GZ	10YR 7/4 Site 396, Core 4, 4483.84-4493.10 = (23.85 subbottom): NO RECOVERY

		c	FOSSIL	N			NOI	IPLE	
in the second se	ZONE	FORAMS	NANNOS	SECTIO	METERS	LITHOLOGY	DEFORMAT.	LITHO.SAM	LITHOLOGIC DESCRIPTION
				0		VOID	٧		NANNOFOSSIL 002E
				1	0.5		0000-		10YR 7/3 Light yellowish brown (10YR 6/4) to very pale brown (10YR 7/3 or 10YR 7/4), soupy to highly deformed namofossil ooze. Darker patches approach marly oozes (pelag clay plus ooze).
					1.0				10YR 7/4 10YR 8/4 Smear Slides: 1-81, 2-90, 3-117, 4-67, 5-7 6-75
				F	1111				10YR 7/4 Forams Tr-7% Sand 0% Nannos 95-100% Silt 1-5% Fe-Ox Tr Clay 95-99%
				2	and and a			•	3-108 10YR 7/4 Forams Tr Sand 0% Nannos 45% Silt Tr Carb. Unspec. 1% Clay 99% Fish Remains 1% Clay 50% Fe-0x 2-3%
	brouweri			3	in the				10YR 6/4 Carbon Carbonate 2-100 10.6 0.0 88 Grain Size
C SUCCESS	mbilica-D			2	duct			•	10YR 7/4 10YR 4/4 10YR 4/4 10YR 8/3
22	Mixed R. pseudo			4	mhadan			•	10YR 6/4 CHEM, SAMPLE
	G. tosaensis			5	munum			* cc	10YR 6/4
				6	ruluuluu				
	G. tos			6 Co Ca	retcher				

510e 390	Hole	Core 6 Cored Inter	<pre>/al: 4502.37-4511.92 m (42.37-51.92 m subbottom)</pre>	Site	396	Hole		C	ore	7 Cored In	terva	a1:	4511.92-4521.47 m (51.92-61.16 m subbottom)
AGE ZONE	FOSSIL CHARACTER SNVNOJ	METERS METERS	THOLOGIC DESCRIPTION	ÅGE	ZONE	FORAMS	FOSSIL ARACTE SONNYN	R	SECTION	2 ≝ LITHOLOGY	DEFORMATION	LI THO. SAMPLE	LITHOLOGIC DESCRIPTION
PLIOCENE C. macintyrei	0. Cosaetists	$\begin{array}{c c c c c c c c c c c c c c c c c c c $	NANNOFUSSIL OOZE Soupy to intensely deformed, pale yellow brown (10YR 8/3) namnofossil ooze with almost white (10YR 8/6) foraminifer layers, the largest in Section 3. Smear Slides: 1-140, 2-70, 3-30, 5-66, 6-50, CC 10YR 8/3 Forams 0- 2% Sand -4 Namnos 98-100% Silt 1 - 2% Fe-0x Tr 3-3 10YR 8/3 Forams 50% Sand 3% Fe-0x 1% Clay 98-100% GZ 3-3 10YR 8/3 Forams 50% Sand 3% Fe-0x 1% Clay 47% Carbon Carbonate 10YR 8/6 10YR 8/6 10YR 8/6 10YR 8/6 10YR 8/7 Grain Size 10YR 8/3 CC 10YR 8/2 10YR 8/2 10YR 8/2 10YR 8/2 10YR 8/4 10YR 8/3	PLIOCENE	Mixed R. pseudoumbilica-D. broweri Globoretalia tosaensis				0 0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1		>		NANNOFOSSIL OOZE 10YR 8/3 Intensely deformed pails yellowish brown (10YR 7/4 and 10YR 8/3) nannofossil ooze. Smear Slides: 1-80, 2-80, 3-115, 4-80, 5-110, 5-150, 6-8, 6-40, 6-53, 6-60, 6-122, CC Porams 0-38 Nannos 99-100% Silt 0- 2% Pe-0x Tr Clay 98-100% Carbon Carbonate 2-100 10.6 0.2 86 10YR 8/3 Grain Size 4-50 0.1 26.3 75.7 10YR 8/3 Intense 1000 Intense 1000 Intense 1000 Intense 1000 10YR 7/4 Intense 1000 Intense 1000 Intense 1000 Intense 1000 Intense 1000 10YR 7/3 Intense 1000 Intense 1000 Intense 1000 Intense 1000 Intense 1000

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Site 396	Hole	Core 8 Cored Interv	1: 4521, -47-4530, 72	2 m (61.16-70.41 m subbottom)	Site	396	Hole	Cor	e 9 Cored Inte	erval:	4530.72-4539.95 m (70.41-79.64 m subbottom)
AGE ZONE	FOSSIL CHARACTER SOUNVN	METERS METERS ADOTOHLIT DEFORMATION	LLTHO.SAMPLE	LITHOLOGIC DESCRIPTION	AGE	ZONE	FOSSIL CHARACTE SWN04 SWN04	SECTION	LITHOLOGY	DEFORMATION LITHO.SAMPLE	LITHOLOGIC DESCRIPTION
D. tamaiis D. surcuius	G. miccenica	$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	10YR 8/4 10YR 8/3 10YR 8/3 10YR 8/3 10YR 8/3 10YR 8/3 10YR 8/3 10YR 8/3 10YR 8/3 10YR 8/3 10YR 8/3	NANNOFOSSIL 00ZE AND FORAMINIFER- NANNOFOSSIL 00ZE Highly deformed toward top, less so toward bottom, pale yellowish brown (10YR 8/3) to very pale brown (10YR 7/3 or 10YR 7/4) or yellow (10YR 8/2) octo. May once have been well bedded, alternating foram sand and nanno coze, now well mixed. Smear Slides: 1-130, 2-20, 2-80 Forams 10-25% Sand 0-35 Nannos 75-90% Silt 10-20% Fo-0X Tr Clay 77-90% Sponge Spic. Tr 2-125, 2-147, 3-40, 4-70, 5-110 Forams 0-5% Sand 0-Tr Nannos 95-100% Silt 0-44 Forams 55% Sand 3% Nannos 44% Silt 50% Fe-0X Tr Clay 47% Carbon Carbonate 2-100 10.8 0.8 83 Grain Size 4-50 8.4 24.8 66.8	PLIOCENE	Mixed R. pseudoumbilica-D. brouveri G. miocenia		0 1 2 3 4 5 <u>Co</u> cat		V V V V	 NANNOFOSSIL OUZE Highly deformed yellow (10YR 8/2) or pale yellow brown (10YR 7/3) nannofossil oose. Only slight color variations indicate some form of bedding. Otherwise, very uniform. Smear Slides: 1-130, 1-150, 2-80, 3-100, 4-100, 5-110, CC 10YR 8/2 Forams 0-24 Sand 0% Silt Tr 10YR 8/2 Carbon Carbonate 2-200 Silt Tr 10YR 7/3 Grain Size 4-50 0.0 35.8 64.1 10YR 7/4 10YR 7/4

Si	te 396		Hole Core 10 Cored Interval: 4559.95-4549.18 m (79.64-89.18 m subbottom)						4539.95-4549.18 m (79.64-89.18 m subbottom)	Site	396	5 1	lole		Co	re 11	1 Cored Int	erval	: 4549.18-45	58.41 m (89.18-9	98,41 m subbo	strom)			
300	101	20NE	FORAMS	COSSIL IRACTEI	R	SECTION	METERS	LITHOLO	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC DESCRIPTION	AGE		ZONE	FORMIS	FOSSIL ARACTEI SONNYK	SCOTTON	METERS	LITHOLOGY	DEFORMATION	LL1110, 2001 LL	LITHOL	OGIC DESCRIP	TION	
u toutout	Nixed R. pseudounbilica-D. brouweri	G. miocenica				0 0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	e e			• • • • • • • • • • • •	NANNOPOSSIL COZE Moderately to intensely deformed pale yellow brown (10YR 7/3) to very pale brown (10YR 7/4) nannofossil ocie. No structures. Very hemogeneous. Shades of color differences only. IOYR 7/4 Smart Sides: 1-120, 2-1007, 3-51 Framas Tr Sand 0% Namnos 99% Silt Tr Fe-Ox Tr Clay 99-100% FORMUNIFER-NANNOFOSSIL COZE IOYR 7/4 Moderately deformed pale yellow brown (10YR 8/3) foraminifer-nannofossil ocie. IOYR 7/4 Moderately deformed pale yellow brown (10YR 8/3) foraminifer-nannofossil ocie. IOYR 7/4 Moderately deformed pale yellow brown (10YR 8/3) foraminifer-nannofossil ocie. IOYR 7/4 Moderately deformed pale yellow brown (10YR 8/3 IOYR 7/4 Moderately deformed pale yellow brown (10YR 8/3 IOYR 8/3 Carbon Carbonate 2-100 10.7 0.9 82 Grain Size Carbon Carbonate IOYR 8/3 IOYR 8/3 IOYR 8/3 IOYR 8/3 IOYR 8/3 IOYR 8/3 IOYR 8/3	PLIOCENE	Mixed R, pseudoumbilica-D, brouweri	G. m. evoluta miocenica			0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	0.5- 1.0-		v	10YR 8/3 10YR 7/3 spots of 10YR 8/3 10YR 7/3 10YR 7/3 10YR 8/3 10YR 7/3 10YR 7/3 10YR 7/3 10YR 7/3 10YR 7/3 10YR 7/3 10YR 7/3	NANK Intensely to brown JOYR 24 unar structures at bedding marly layers Smear slides Forams Nannos Fe-Ox Carbon Carbo 2-100 Grain Size 4-52	DFOSSIL 002E D moderately (3 or very junofossil oo: scent abrupt surfaces. Mi - - - - - - - - - - - - -	deformed pu ale brown i.e. No sed: minor chains or browne: 40 3-140 cc Sand Silt Clay 87 74	ale yollow 1078 7/3 to imentary nges of color = 1078 6/4 4-60 5-98 0% Tr-2% 98-1008 0% 5% 95%

Site 396	5	Hole		Cor	e 11	Z Cored In	nterva	a1:	4558.41-4567.96 m (98.41-107.96 m subbottom)	Sit	e 39	96	Hole		Core	13 Cored Inte	rval:	4567.96-4577.20 m (107.96-117.20 m subbottom)
AGE	ZONE	FORAMS	RACTER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC DESCRIPTION	ÅGE		ZONE	FOSSI CHARACT SONNEN SWENDE	IL TER	SECTION	LITHOLOGY	LI THO . SAMPLE	LITHOLOGIC DESCRIPTION
PLIOCENF Mixed R. pseudoumbliten-D. brotwert	G. margaritae evoluta			0 1 2 3 4 5 6				• • •	NANNOFOSSIL 00ZE AND FORAMINIFER-NANNOFOSSIL 00ZE Very pale brown (107K 7/3 or 107K 7/4) nanno- fossil ooce, adderately to intensely deformed, plus one layer in Section 5 of gritty foraminifer- nannofossil ooce, pale yellow brown (107K 8/3). Nannofossil ooce is homogeneous. No obvious bedding planes except the foram-rich layer in Section 5. Smear Slides: 1-140 2-140 3-140 4-140 6-130 and cc Forams 77-38 Sand 0% Nannos 97-100% Silt Tr-2% Fe-0x Tr Clay 98-100% S-110 Forams 30 30% Sand 5% Nannos 70% Silt 25% Fe-0x Clay 70% Carbon Carbonate 2-101 10.8 0.0 89% Grain Size 4-50 0.0 32.9 67.4% 10YR 7/3	PLIOCENE		Ceratolithus rugosus G. m. margatitae G. margaritae evoluta			0 1 1 1.(2 3 4 5 6 Core Careford		·	NUNNDFOSSIL 002E Note: upon retrieving this core, the liner was damaged at the bottom of what is here shown as section 2. If was not possible to determine how much sediment was lost or what the precise interval of the sediment is above the damaged liner was both borken and compressed (twisted). It was necessary to have a shortened Section 2 and a lengthened Section 0 because of the damaged interval. The sediment is moderately to intensely deformed vory pale brown (10YR 7/4 or 10YR 7/3) nanno-fossil ooze. There are no sedimentary structures and the sediment is homogeneous except for suble color changes. 10YR 7/3 METALIFERROUS BROWN CLAY AND MARLY NANNOFOSSIL OOZE The base of the core contains brown (10YR 5/4 or 10YR 4/4) clays and marly clays. Smear slides contain more hydrous (?) iron and manganes (?) oxides than the nanofossil occes. This and their proximity to basement suggests that the clays are metaliferrous, similar to clays found above basalt at many places in the world oceans. Smear Slides: 0-125 1-130 2-90 3-100 4-100 5-100 for 4-200 5-200 for 5-200 fo

Т		c	FOSSIL				NO	PLE					
AGE	ZONE	FORAMS	NANNOS	SECTION	METERS	LITHOLOGY	DEFORMATI	LITH0.SAM		LITHOL	DGIC DESCR	IPTION	
PPLIOCENE -	G. margaritae margaritae			0	0.5	Dettertion of		* cc	10YR 7/4 10YR 3/3 10YR 5/6	BROWN CLAYS, OUZE AND BASJ The top of ti clays. These by drilling r nanofossil to broken line 3-80 in the are generally but they are soupy nannofo basalt. The were complet result of cor core as the l of Section 1 on the first tion forms.	MARLY NAN ALT he core coi- e are fair and mixed i soze. Thi- lithology i brown (1 smeared w. sosil ooze four topm ely encase apression - hasalt was 4-6 is des of the fo	NOFOSSIL 007 ntains maini yy firm, but yith soupy is is indicat rval from 1 column. Th MyR5/4 or 11 ith lighter. . Little du (107K 5/6) ost pieces 6 d in this mo of the ooze drilled. 7 cribed in mc llowing base	LE, NANNOFOSSIL iy brown t fractured prown ced by a -80 to a clays NR 4/4) -colored overlies of basalt into the the basalt ire detail alt descrip-
MI DDLE MIOCENE	Discoaster exilis Gioborotalia fohsi fohsi			3 4 5 6					10YR 5/6 10YR 5/6	Remark: The bi metallferrous the world occ. Smear Slides: Forams Nannos Fe-OX *Note: the pi to estimate : given below : given below : Smear Slides. Forams Nannos Fe-OX q Carbon Carbon 2-100 Grain Size 4-50	rown clays s clays cor cans. It: the most i not immed : 1-40 2- Tr-1% 97-99% 1-3% ercentage (in these si indicates i s in sedim by these si : 5-75 6 Tr-3% 97-100% Tr nate 5.2 0.1 1.7 49.	are similar mmon above b is important clay-rich i liately over 100 3-100 4 Sand Silt Clay of clays war mear slides. 57% clay and ent similar mear slides. -10 Sand Silt Clay 43% 4 48.8%	r to basal asalt in to note, ayers in lie basalt. 1-75* 0% 1% 99% s difficult The %CaCO ₃ i iron-mangañes* 0% Tr-3% 97-100%
	Discoaster exili			6 Co Ca)re atcher			•					

Explanatory notes in Chapter 1







VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

u	EG		SIT	TE	HOLE	с	OF	E	SE	ст.
4	5	3	9	6	Π		1	5		2

MAJOR ROCK TYPE - SPARSELY TO MODERATELY PLAGIOCLASE-OLIVINE PHYRIC BASALT

Macroscopic Description

0.

MD

A continuation of fine-grained to glassy sparsely phyric pillow basalts with abundant lithified sediments in cracks and attached to basalt, described in Section 1. Glass present as indicated on many pieces, and on both sides of cracks containing lithified sediments. Zones of rather great alteration next to such cracks. No particular variation in proportions of phenocrysts, which are similar in size and abundance to those of Section 1. Pervasively moderately to intensely altered.

	Inc.	NRM	MDF	Vp (1)	Vp (1)	D	P
91-93 cm:	+38.2°	1.64	700			2.82	

6

SITE

LEG

4 5 3 9

VISUAL CORE DESCRIPTION

FOR IGNEOUS ROCKS

palagonitized glass. Piece 11 is essentially hyaloclastite.

This sequence, like basalts in Core 14, extruded into ooze.

FINE-GRAINED BASALT

Macroscopic Description







SITE 396

SECT

4



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

u	G		SIT	E	HOLE	C	OF	E	SE	ст.
4	5	3	9	6			1	8		1

MAJOR ROCK TYPE - SPARSELY TO MODERATELY PLAGIOCLASE-OLVINE PHYRIC

Macroscopic Description Sections 1 through core-catcher

Fine-grained, very fine-grained, and glassy plagioclase-olivine phyric basalt. Plagioclase 2-7% (1-5 mm), olivine Tr-2% (1-3 mm), Glass present in many places as shown. Has cracks lined with carbonate, but no recrystallized foram-nanno ooze is present. Alteration is moderate to intense throughout, especially adjacent to cracks and glass. Vesicles are Tr-3%, less than 1 mm, and partially filled with secondary clays. Glass is partially palagonitized. There are minor fluctuations in the abundance of plagioclase phenocrysts; some pieces are virtually aphyric, many have no more than 3% plagioclase phenocrysts, others contain as much as 5-7%. Olivine is more constant 1-2%, but is generally partially altered to reddish secondary minerals. Grain size ranges from glassy, through subvariolitic zones of incipient crystallization, to fine-grained but essentially holocrystalline rocks, representing the transition from pillow selvage to pillow interior. It is doubtful whether a single pillow thicker than about 1 meter was cored in this interval.

70-78 cm (Piece 1G): Sparsely phyric plagioclase-olivine basalt, subvariolitic to subophitic groundmass. Phenocrysts: plagioclase 5-7% (.2-4 mm) eu-subhedral; olivine 1% (.2-1.5 mm) eu-subhedral. Both olivine and plagioclase phenocrysts contain rounded inclusions of glass. Picotite is rare. Matrix (90%) contains skeletal olivine (1%), skeletal plagioclase (45%), titanomagnetite (1-3%, dust sized), and clinopyroxene (50%). It is partly altered to clay minerals, and is partly subvariolitic "sworls", and partly interlocking crystal laths.

130-136 cm (Piece 1N): Sparsely phyric plagioclase-olivine basalt, subvariolitic groundmass. Phenocrysts: plagioclase 10% (.2-3 mm) eu-subhedral; olivine 1% (.2-.8 mm) eu-subhedral. The groundmass is largely dark, microcrystalline subvariolitic "sworls" with about 5% tiny skeletal olivine crystals. Needle-like plagioclase and clinopyroxene in about equal proportions make up the sworls, which are dusted with fine opaques. Chrome spinel is accessory. The matrix is partly altered to clays, and has minor secondary calcite. Plagioclase phenocrysts are rounded, and contain

2	AI203	Fe ₂ O ₃	MgO	CaO	к20	TiO2	IL	Mg/(Mg-	+Fe)	Cr	Ni	Sr	Zr
-13	6 cm:												
7	16.93	9.68	7.9	12.09	0.29	1.27	-2.6	0.62		383	142	146	87
			Inc.	NRM	MDF	Vp (11)	Vp (1)	D	Ρ				
3-70	cm:		+33.8°	.944	440			2.78	-				
9-11	1 cm:		+37.7°	.251	450			2.71	100				











VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



MAJOR ROCK TYPE - SPARSELY TO MODERATELY PLAGIOCLASE-OLIVINE PHYRIC BASALT

Macroscopic Description

General description given in Section 1. Glass present in Sections 1B, 1E, 1K, 1L, 1N, 1O, and 1Q. Pieces 1E-1K are probably a section through a pillow. Plagioclase and olivine phenocrysts increase in abundance and grian size from Pieces 1E (plag. about 3%, ol. less than 1%) to 1 H and 11 (plag. 7%, ol. 3%, partly fresh) and then decrease again to Piece 1K. A similar sequence probably representing a pillow was recognized between Pieces 1L and 1Q. At glassy zones, reddish brown alteration (palagonite?) is present. Olivine most altered in intensely altered pieces; partly fresh in moderately altered pieces.

Thin Section Description

29-31 cm (Piece 1E): Moderately phyric plagioclase-olivine basalt, vitrophyric groundmass. Phenocrysts: plagioclase 13% up to 7 nm; olivine 2%, up to 2 mm; both subrounded, skeletal. Matrix ranges from glassy (pale brown) to dark brown and black, where spherulites coalesce. Within this dark matrix, tiny needles of plagioclase and crystals of olivine can be seen here and there. Portions of the glass are altered to palagonite. Plag. is An₇₀. Olivine phenocrysts contain blebs of clear glass.

	Inc.	NRM	MDF	Vp (1)	Vp (1)	D	Ρ
08-110 cm:	+36.6°	.750	700			2.81	-



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

LI	EG	100	SIT	re	HOLE	0	OF	RE	SE	ст.
4	5	3	9	6	П		2	2		4

MAJOR ROCK TYPE -- SPARSELY TO MODERATELY PLAGIOCLASE-OLIVINE PHYRIC BASALT

Macroscopic Description

General description given in Section 1. Glass present in Pieces 11, 1M, and 1N. Plagloclase abundance slightly lower from Pieces 1A to 1H (5-7%) than from Pieces 11 to 1Q (7-10%). Carbonate veinlets in most pieces, prominent in Pieces 1E, 1F, 1G, 1K, and 1L. Pieces 1D, 1E, and the top of Piece 1F are more vesicular (2%, less than 1 mm, empty) than other pieces in the section. Some reddish secondary minerals are in the carbonate vein in Piece 1E.

Thin Section Description

144-150 cm (Piece 1Q): Sparsely phyric plagioclase-olivine basalt, holocrystalline basaltic groundmass. Phenocrysts: plagioclase 10% (up to 5 mm) eu-subhedral, skeletal, contain glass inclusions; olivine 2% (2.5 mm) eu-subhedral, mostly altered to clays and Fe-hydroxides ("iddingsite"). Groundmass: plagioclase 40%; clinopyroxene 35%; olivine 3%; titanomagnetite 3%. Clay minerals, Fe-hydroxides, and vesicles present in trace quantiles. Plag. and olivine phenocrysts can form glomerophyric aggregates. Both are typically rounded.

SiO2	Al ₂ O ₃) cm:	Fe203	MgO	CaO	K20	TiO2	IL.	Mg/(Mg+Fe)	Cr	Ni	Sr	Zr
19.8	16.93	9.47	7.6	12.11	0.25	1.29	-2.3	0.61	309	123	135	86
			Inc.	NRM	MDF	Vρ (1)	Vρ (1)	D P				
35-13	cm:		-4.8°	4.53	450			2.86 -				









cm

0

50

100

150

MD

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MAJOR ROCK TYPE - MODERATELY PLAGIOCLASE-OLIVINE PHYRIC BASALT

FOR IGNEOUS ROCKS

Macroscopic Description

General description given in Section 1. From Section 2, Pieces 1P to Section 3, Piece 11 is an apparent cooling unit. No significant variation of plagioclase phenocrysts is observed in this cooling unit. Olivine is only about 1% in the glass of Piece 11, however, and 3% in the cooling unit interior. From Pieces 1J to 1P, plagioclase phenocrysts increase from 5% to 10%, and olivine phenocrysts from less than 1% to about 3%. Black spots of Mn are seen on carbonates of Pieces 1D, 1J, and 1M.

Thin Section Description

73-78 cm (Piece 11): Moderately phyric plagioclase olivine basalt, subvariolitic to intersertal groundmass. Phenocrysts: plagioclase 20% (.2-4 mm) subeuhedral; olivine 1% (.1-,4 mm) eu-subhedral. Plag. is An₇₀₋₇₅. Groundmass: plagioclase 10% acicular, skeletal; olivine 1% skeletal. Matrix mostly finely crystalline cpx.:plag. = 2:1, with a fine dust of titanomagnetite. Some phenocrysts of plagioclase and olivine are rounded. Some vesicles (1-2%) contain microcrystalline volcanic glass, partly altered.

SiO2	Al203	Fe203	MgO	CaO	K20	TiO2	IL	Mg/(Mg+Fe)	Cr	Ni	Sr	Zr
49.1	16.58	9.68	7.9	12.07	0.23	1,29	-2.0	0.62	226	124	128	86
			Inc.	NRM	MDF	Vρ (∎)	Vp (1)	D P				
138-14	10 cm:		-6.9°	2.83	450			2.85 -				



crysts: plagioclase 7% (1-2 mm) sub-euhedral; olivine 3% (.3-1 mm) skeletal, sub-euhedral, Groundmass: plagioclase 10% (.01-.4 mm) skeletal, acicular; olivine 1% (.1-.2 mm) skeletal, not elongate. Microcrystalline mesostasis composed of titanomagnetite "dust", and highly skeletal clinopyroxene and plagioclase. Largely replaced by clay minerals. Form of mesostasis is subvariolitic "sworls". Note: interval on thin section should be 144-149 cm.

LEG

SITE

4 5 3 9 6

	SiO2 96-100	Al ₂ O ₃ cm:	Fe ₂ O ₃	MgO	CaO	к20	TiO2	IL	Mg/(Mg+Fe)	Cr	Ni	Sr	Zr
1	49.7	16.71	9.49	7.4	11.99	0.13	1.29	-2.2	0.61	310	109	130	81
	111-11	3 cm:		Inc.	NRM 2.39	MDF 450	Vρ (II)	Vp (1)	D P				

CORE

2 5











Site 396





Site 396





Site 396

