## 50. PRELIMINARY LIPID ANALYSES OF CORE SECTIONS 18, 24, AND 30 FROM HOLE 402A

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## INTRODUCTION

A preliminary analysis of the solvent-extractable lipids from three samples from the Aptian-Albian core sections of Hole 402A, which was drilled on the mid-continental slope of the Bay of Biscay (47°52.48'N, 08°50.44'W), has been carried out. The aliphatic and aromatic hydrocarbons, carboxylic acids and free hydroxy acids have been examined by gas chromatography and gas chromatography-mass spectrometry.

A list of the samples is given in Table 1, together with relevant background data taken from the Summary Log. Core 402A-18 was a composite sample of Sections 1-4, obtained via the Institut Français du Pétrole; Cores 402A-24 and 30 were collected from the Scripps Institute of Oceanography. An additional core, 401-5 (supplied in a frozen state by the shipboard geochemist via British Petroleum Ltd.), was extracted, but the low concentrations of extractable lipids at comparable concentrations to a blank analysis precluded further investigations.

A preliminary attempt has been made to relate the extractable lipids of these black shales to the extent of diagenesis/maturation and to the paleoenvironment of deposition.

## EXPERIMENTAL PROCEDURES

### General

Cores 402A-24 and 30 and Core 401-5 were kept frozen until extraction; Core 402A-18 was supplied in a freeze-dried state.

All operations prior to GC were carried out in a clean room equipped with a filter unit (Microflow Pathfinder Ltd.) to remove continuously all particles  $>0.5 \mu m$  from the incoming air stream. Nylon gowns, hats, overshoes, and gloves were worn and the room was frequently vacuum cleaned; a fume hood (Bigneat Ltd.) circulated air through molecular sieve to remove contaminants and solvent vapors. All solvents were Mallinckrodt Nanograde and examined for purity by concentration 100 ml to  $20\,\mu$ l) and analysis by GC as described below. Glassware was cleaned by boiling in detergent (Decon 90), rinsing ( $\times$ 6) in double-distilled water (distilled from KOH/KMnO<sub>4</sub>), heating (3 hr, 600°C) and finally, rinsing twice with methanol and twice with dichloromethane immediately before use. Pasteur pipettes and glass vials were heated at 600°C and rinsed as above; caps were lined with P.T.F.E. discs washed in dichloromethane. Double distilled water was extracted with dichloromethane before use. Concentrated HCl, mercury, aluminum foil, and glass wool were extracted with dichloromethane and the glass wool heated (600°C); fused

TABLE 1 Sediment Data for Samples Examined

Sample (Interval in cm)	Mean Depth (m)	Age	Lithology
402A-18-1, 4	298-308	Albian	Composite of carbonaceous marly limestone and carbonaceous calcareous mudstone
402A-24-1, 0-20	360	Lower Albian	Carbonaceous marly limestone
402A-30-1, 65-75	420	Upper Aptian	Carbonaceous marly chalk
401-5-1, 1-5 (all 0-5)	113-122	Mid Eocene	Nannofossil chalk

KOH was dissolved in water to yield a 60 per cent solution and was extracted with dichloromethane. The boron trifluoride/methanol complex was found to be sufficiently clean in the volumes used and the *bis*(trimethylsilyl)trifluoroacetamide contained only minor amounts of low molecular weight components that did not interfere with the analyses of alcohols. Silica-gel thin layers were purified by continuous elution in ethyl acetate (> 20hr) and removal of contaminated gel. The N<sub>2</sub> for solvent evaporation was "oxygen-free" grade and was passed through activated molecular sieve (5Å). Extracted fractions were stored in dichloromethane (ca. 50 µl).

## **Extraction** (Figure 1)

After thawing (ca. 1 hr) each sample was pulverized. The lipids were extracted with dichloromethane/methanol ( $\times 2$ , 2:1, 200 ml) by ultrasonication (10 min, Dawe Soniprobe 7533A, equipped with Microtip) with cooling (ice-bath).<sup>1</sup> After centrifugation (10 min; 2500 rpm), the yellow-brown supernatants were decanted; the extraction was repeated with 100 ml solvent and the combined extracts concentrated at 30°C under water pump vacuum. The residues from solvent extraction were stored at  $-20^{\circ}$ C after drying.

The extract was saponified (6% KOH in methanol; 5 ml) by leaving to stand overnight at ambient in the dark under N<sub>2</sub>. The insoluble residues after extraction were dried, weighed, and stored for further analysis. Water (6 ml) was added and the neutral lipids were extracted in dichloromethane ( $3 \times 3$  ml), washed with water (4 ml) and the water washings combined with the aqueous fraction. The neutral extract was evaporated to dryness at 30°C. The aqueous solution was acidified (*p*H1) with concentrated HCl and extracted with dichloromethane ( $3 \times 3$  ml).

<sup>&</sup>lt;sup>1</sup>Sections 402A-18-1 to 4, supplied as a freeze-dried residue by Dr. B. Tissot, was extracted with hexane/ propan-2-ol (1:4) and was saponified as above; the acid and non-saponifiable fractions were treated as above except that elemental sulfur was removed during the TLC separations.



\*Core 402A-18 freeze dried and extracted with hexane/propan-2-ol (see text).

Figure 1. Analytical scheme for the extractable lipids of core samples from Hole 402A.

Sulfur was removed from acid and non-saponifiable fractions by shaking with elemental mercury (lg) and hexane (1 ml), allowing the sulfide precipitate to settle (>3 hr), and removing the supernatant. The precipitate was washed with hexane (× 2; 2 ml) and the complete procedure was repeated with a further aliquot of mercury; the hexane layers were combined and evaporated under N<sub>2</sub>. Boron trifluoride/ methanol complex (14%, 0.2 ml) was added to the total acid fraction and the mixture heated (70°C; 5 min). Water was added to the methyl esters extracted into dichloromethane  $(3 \times 1 \text{ ml})$ , followed by solvent evaporation under N<sub>2</sub>.

## Thin-Layer Chromatography (TLC) (Figure 1)

The non-saponifiable and methylated total acid (FAME) fractions were further separated on silica gel G (0.25 mm) with dichloromethane, and with hexane/ethyl acetate (9:1) as developers, respectively. Separated fractions were located (UV light; 254  $\mu$ m, Rhodamine 6G in ethanol), and fractions were collected from the total non-saponifiables as follows: Rf = 0.75-1.00 (Rf = tetracosane standard), 0.20-0.75 (includes ketones), 0.05- 0.20 (Rf = cholesterol), 0.0-0.05. Fractions were collected from the FAME as follows: Rf = 0.65-1.00 (Rf = methylstearate standard), 0.05-0.65 (includes hydroxy acids, OH FAME), 0.0-0.05. The hydrocarbons were further separated into aliphatic and aromatic fractions with hexane as developer, by comparison with authentic standards (tetracosane and pyrene).

All TLC fractions were collected from the silica gel by elution with dichloromethane (10 ml). Separated fractions containing OH FAME and alcohols, respectively, were silylated with *bis*(trimethylsilyl)trifluoroacetamide ( $20 \mu$ l; 5 min, ambient).

## **Control Experiments**

For each sample, the entire analytical sequence (except for GC-MS) was carried out in parallel under identical conditions except that no core sample was present.

## Gas-Liquid Chromatography (GLC)

Analyses were carried out on a Carlo Erba FV2151 equipped with an SL490 splitless injector, under the following conditions:  $10 \text{ m} \times 0.25 \text{ mm}$  glass coated with OV-1, 60 to 260°C at 6°/min, after 1 min, at ambient, 0.4 kp/cm<sup>2</sup>N<sub>2</sub>, injector 275°C, detector 275°C.

# Combined Gas Chromatography-Mass Spectrometry (GC-MS)

Analyses were carried out on a Finnigan 4000 GC-MS coupled via a Finnigan interface to a DEC PDP-8e computer (M.J. Humberston, unpublished results) under the following conditions: column  $20 \text{ m} \times 0.32 \text{ mm}$  i.d. glass coated with OV-1 (Jaeggi), programmed from 120 to 260°C at 6° or 8°/min, flow rate ca. 2 ml/min He, source 250°C, filament current  $250 \mu$ A, electron energy 35 eV, scan rate 2.5 sec.

## RESULTS

For Cores 402A-24 and 402A-30, all of the blank fractions equivalent to the fractions examined (aliphatic hydrocarbons, aromatic hydrocarbons, FAME and OH FAME) showed virtually no measurable contaminants by GLC analysis (under identical conditions) of injection volumes equivalent to those used for the fractions themselves. The same situation applied to the aliphatic hydrocarbon and OH FAME blanks of Core 402A-18; the aromatic hydrocarbon blank showed two components which were not aromatic hydrocarbons. The FAME blank of 402A-18 contained n-C<sub>16:0</sub>, n-C<sub>16:1</sub>, n-C<sub>18:0</sub>, and n-C<sub>18:1</sub> at concentrations half those observed in the sample; the distributions for Core 18 FAME are corrected accordingly.

Quantities of individual fractions were too low for accurate weighing; Table 2 shows the concentrations of selected individual components in the fractions, expressed as ng/g extracted sediment (dry weight). Although only a few components have been quantitated, Table 2 shows that Core 402A-30 has the highest concentration of extractable lipids.

In this preliminary study, tentative structural assignments are based mainly on relative retention times and comparison of full mass spectra with authentic standards (where available), comparison with literature spectra or spectral interpretation. Distributions of compound classes for comparison between samples were obtained by examination of the appropriate mass fragmentograms.

#### Aliphatic Hydrocarbons

## Normal and Branched Alkanes

The distributions of *n*-alkanes in the three samples range from C<sub>13</sub> to C<sub>34</sub> (402A-18), C<sub>16</sub> to C<sub>34</sub> (402A-24), and C<sub>15</sub> to C<sub>34</sub> (402A-30) with CPI values of 1.07, 1.38, and 2.14, respectively (Figure 2). All show bimodal distributions maximizing at C<sub>16</sub> or C<sub>19</sub> and C<sub>25</sub> or C<sub>27</sub>, with a smooth distribution of lower molecular weight components. Core 18 shows the highest relative abundances of *n*-C<sub>13</sub> to C<sub>17</sub>,

 
 TABLE 2

 Concentrations<sup>a</sup> of Selected Individual Lipids in Core Samples From Hole 402A

	Sample		
	18-1 to 4	24-1	30-1
Aliphatic Hydrocarbons			
n-C17:0	17	4	11
n-C18:0	13	8	15
Hop-17(21)-ene	32	27	60
17βH-homohopane	86	66	160
Aromatic Hydrocarbons			
Perylene	0.5	b	b
FAME			
n-C16.0	11	3	12
n-C24:0	4	0.15	10

<sup>a</sup>Nannogram of component/g dry weight extracted sediment; based on GLC peak height relative to a standard.

<sup>b</sup>Concentration too low for quantitation by GLC.



Figure 2. Distribution of n-alkanes (unbroken lines), pristane and phytane (broken lines) in core samples: (A) 402A-18, (B) 402A-24, (C) 402A-30.

and of pristane and phytane (Figure 2) which are associated with a complex envelope of unresolved components not present in Cores 24 and 30. This suggests shipboard contamination of the composite Core 18. Low levels of pristane and phytane are present in Cores 24 and 30.

#### Steranes and Sterenes

Several rearranged sterenes  $(I,R=H,CH_3,C_2H_5)$  were present as major components in all three samples, with a similar distribution; Figure 3 shows these components in Cores 24 and 30 by way of m/e 257 which is the base peak in the spectra of these compounds (Rubinstein et al., 1975).



Figure 3. Partial mass fragmentograms of m/e 257 (A and B) and 271 (C and D) for aliphatic hydrocarbon fractions from Cores 402A-24 and 402A-30.

The six major components in increasing order of elution are two C<sub>27</sub> (I,R=H), three C<sub>28</sub> (I,R=CH<sub>3</sub>) and one C<sub>29</sub> (I,R=C<sub>2</sub>H<sub>5</sub>) components. Tentative assignments were based on comparison of the mass spectra with literature examples (Rubinstein et al., 1975) GC-MS analyses of Core 18 were carried out before and after storage (5°C, 8 weeks) in solvent; the fragmentogram of m/e 257 showed that the rearranged sterenes were virtually absent, presumably as a result of atmospheric oxidation. This emphasizes the necessity of rapid analysis and careful storage for these compounds.

In addition, a series of nuclear methylated rearranged sterenes (II) was observed in Cores 24 and 30 and were characterized (Rubenstein et al., 1975) by m/e 271 (Figure 3). The corresponding fragmentogram for Core 18 was not obtained prior to storage; after storage these components were virtually absent. In Cores 24 and 30, the distributions were similar to each other. However, they occur as minor components, being present at ca. 1/15 of the concentration of the rearranged sterenes. The full mass spectra showed that the four major components are a  $C_{28}$  (II,R=H) and three  $C_{29}$  (II,R=CH<sub>3</sub>) compounds (Figure 3).

Steranes (m/e 217) were present in only small quantities in all three samples; no sterenes (m/e 215 and m/e 213) were observed.

#### **Triterpenes and Triterpanes**

The distributions are again very similar in all three samples, as indicated by the fragmentograms for m/e 191 (see Figure 4 for Cores 24 and 30) and m/e 189 (not shown). The structural assignments of the major components are listed in Table 3 and the chemical structures<sup>2</sup> appear at the end of the text in this chapter. The triterpanes were assigned by comparison of their mass spectra with those of standards, except for compound R, which was assigned by spectral interpretation.

A series of C<sub>27</sub> (III,R=H), C<sub>29</sub> to C<sub>31</sub> (III,R = C<sub>2</sub>H<sub>5</sub>, *iso*-C<sub>3</sub>H<sub>7</sub>, *sec*-C<sub>4</sub>H<sub>9</sub>) triterpenes, characterized by M<sup>+</sup>, M<sup>+</sup>-15, m/e 367, m/e 231, and m/e 191, were observed in all three samples. Only a spectrum of hop-17(21)-ene (Component G) was available for comparison, but Ensminger (1977), using the same criteria, also tentatively assigned components of this series in the Toarcian shales of the Paris Basin. The two C<sub>31</sub> components (L and M) are presumably C-22 epimers (IIIa). 22,90,30-Trisnorhop-17(21)-ene has only been identified to date in a surface marine sediment of the Norwegian Sea (Dastillung et al., 1977) and the C<sub>29</sub> component (F) has not been reported previously.

Three other triterpenes (C<sub>27</sub>, C<sub>29</sub>, C<sub>30</sub>) are also present in all of the samples. Compound J has a mass spectrum identical to that of neohop-13(18)-ene (IV,R=*iso*-C<sub>3</sub>H<sub>7</sub>). The spectra of the C<sub>27</sub> and C<sub>29</sub> components are similar to that of IV (R=*iso*-C<sub>3</sub>H<sub>7</sub>), except that m/e 218 is shifted to m/e 204 and m/e 176, respectively. It is, therefore, possible that both compounds have  $\Delta$  <sup>13(18)</sup> neohopene structures. However, the retention time of neohop-13(18)-ene is longer than that of hop-17(21)-ene, whereas the retention times for the proposed  $C_{27}$  and  $C_{29}$  compounds are reversed relative to their hop-17(21)-ene counterparts. Only the  $C_{30}$ compound has been identified to date in sediments, viz. the Toarcian shales of the Paris Basin (Ensminger, 1977).

# Aromatic Hydrocarbons

A feature of these fractions is the gross similarity in the overall distributions of Cores 24 and 30 (Figure 5), although there are differences in the trace components (see below). In contrast, Core 18, containing ca. 1/20 of total aromatic hydrocarbons shows significant differences in the distribution, for example, in the fragmentogram of m/e 252 (=e.g. M  $\ddagger$  of perylene) (Figure 6). The major component in Core 18 was shown to be perylene, from the mass spectrum and co-injection with a standard, which is either absent or is a minor component of the other two. The major component in Cores 24 and 30 is an unknown aromatic hydrocarbon of molecular weight 218.

Among the minor components a number of triterpenoid-derived compounds were tentatively assigned from comparison of their mass spectra with literature examples (Spyckerelle, 1975), although the distributions varied from sample to sample. These include structures VI (Core 24 only), VII (Cores 24 and 30), VIII (Cores 18 and 30), IX (Cpre 18 only). In addition, Cores 24 and 30 contain components whose spectra are consistent with nuclear-methylated aromatic hopanes as follows: structure X (two components in sample 24; M  $\ddagger$  360, M-15, M-29). XI (two components in Sample 30; M  $\ddagger$  342, M-15, M-29). Core 18 shows two components with M  $\ddagger$  = 292, whose spectra compare favorably with literature spectra of XII (Spyckerelle, 1975).

## Carboxylic Acids (FAME)

#### Normal and Branched Alkanoic and Alkenoic Acids

The distributions of the three samples are dominated by components ranging from n-C10:0 to n-C32:0 maximizing at  $n-C_{16:0}$ . In the deepest sample (402A-30) a greater relative abundance of long-chain acids is present (Figure 7) and  $n-C_{24:0}$  is only slightly less abundant than  $n-C_{16:0}$ . Saturated branched acids, dominated by iso- and anteiso-C15:0, are very minor components of all of the samples; acyclic isoprenoid acids are also present in very low concentrations. Two mono-unsaturated acids, n-C16:1 and n-C18:1, were detected at all three levels, but their concentrations were considerably less than those of the corresponding saturated acids; the low concentrations precluded assignment of the position and geometry of the double bond, and the concentration of the two compounds decreased with increasing depth. No polyunsaturated fatty acids were detected.

#### **Triterpanoic and Triterpenoic Acids**

Triterpanoic acids belonging to three series of the hopane family were detected as major components in the FAME from all three depths (Table 4). There is a similar distribution in each case, as shown by the fragmentograms of m/e 191 (Figure 8) although the greatest similarity is in Cores 18 and 24; the  $C_{32}$  components are of greater abundance relative to their  $C_{31}$  and  $C_{33}$  counterparts in Core 30. Structural assignments were based on the characteristic

<sup>&</sup>lt;sup>2</sup>Each chemical structure is designated by a roman numeral.



Figure 4. Mass fragmentograms of m/e 191 for aliphatic hydrocarbons in Cores 402A-24 and 30.

ions M<sup>+</sup>, M<sup>+</sup> -15, M<sup>+</sup>-side chain, m/e 191, m/e 148 + R (where R = the substituent on Ring E). The three series (17\alphaH,21\betaH; 17\betaH,21\alphaH; 17\betaH,21\betaH) were distinguished by the relative abundances of the ions m/e 191 and m/e 148 + R (Van Dorsselaer, 1975). In all three samples,  $17\beta$ H,  $21\beta$ H-bis-homohopanoic acid (XIIIb) is the major triterpenoid acid (Figure 8). Table 4 lists the compounds assigned in Core 402A-30 from their full mass spectra, and they appear to be present in Cores 18 and 24 from comparison of full mass spectra (where obtained), retention data and the distributions in the m/e 191 fragmentograms (Figure 8): the C31, C32, C33 members of the  $17\alpha H, 21\beta H$  series (components A, E, H; XIIIa,b,c), the C<sub>31</sub>, C<sub>32</sub>, C<sub>33</sub> members of the  $17\beta$ H,21 $\alpha$ H series (components C, G., J; XIIIa,b,c), and the C<sub>31</sub>, C<sub>33</sub>, C<sub>34</sub> members of the  $17\beta$ H,  $21\beta$ H series (components F, K, L; XIIIa,c,d). In Cores 18 and 30 a component showing the characteristic fragment ions (m/e 191 and m/e 235) for a C30 triterpanoic acid (methyl ester) was detected but no M <sup>±</sup> was obtained.

At all three depths, two mono-unsaturated  $C_{32}$  triterpenoid acids were present in virtually the same relative

proportions to each other (Table 4). The spectra (e.g., Figure 9) show M  $\ddagger$  = m/e 482, M<sup>+</sup>-15, 367 (M<sup>+</sup>-side chain), 231 (100%), and 191 and are consistent with the compounds being *bis*-homohop-17(21)-enoic acids (XIV). Compounds of this type have not been reported in organisms or sediments although Ensminger (1977) reported a series (C<sub>30</sub>-C<sub>35</sub>) of  $\Delta^{17(21)}$  alkenes in the Toarcian shales of the Paris Basin. The two compounds appear to be stereoisomers but the position at which they are isomeric is not known. In Core 30, two other C<sub>32</sub> mono-unsaturated triterpenoic acids were observed.

## Hydroxy Carboxylic Acids (OH FAME)

Complex mixtures of more polar acids were found in each of the three samples. Silylation of the mixtures indicated the presence of a number of hydroxy acids which was confirmed by fragmentograms for m/e 73, 75, and 89. From the full spectra, a series of  $\alpha$ -hydroxy acids ranging from n-C<sub>11:0</sub> to n-C<sub>22:0</sub>, maximizing at n-C<sub>14:0</sub>, was identified in Core 18, but was not observed in the deeper samples. The distribution is unusual in that even chain length compounds are only slightly more abundant than the odd chain length

TABLE 3 Aliphatic Triterpenoid Hydrocarbons of Cores 402A-18, 402A-24, and 402A-30

Peak <sup>a</sup>	Diagnostic Ions (%)	Tentative Assignment <sup>b</sup>
A	M <sup>+</sup> 368,353,191(100),187(50),176(90),163(70)	22,29,30-Trisnorneohop-13(18)-ene (IV,R=H) <sup>C</sup>
в	M+368,353,231(75),191(100)	22,29,30-Trisnorhop-17(21)-ene (III,R=H) <sup>C</sup>
С	M+370,355,191(100),149(40)	17αH-22,29,30-Trisnorhopane (V,R=H)
D	M <sup>+</sup> 370,355,191(70),149(100)	22,29,30-Trisnorhopane (V,R=H)
Е	M+396,381,367(2),204(45),191(100),189(40)	30-Norneohop-13(18)-ene (IV,R=C <sub>2</sub> H <sub>5</sub> ) <sup>c</sup>
F	M <sup>+</sup> 396,381,367(22),231(65),191(100),189(30)	30-Norhop-17(21)-ene (III,R=C <sub>2</sub> H <sub>5</sub> ) <sup>c</sup>
G	M+410,395,367(100),231(95),191(70),189(50)	Hop-17(21)-ene (III,R=iso- $C_3H_7$ )
н	M+398,383,191(90),177(100)	30-Normoretane (V,R=C2H5,21aH)
I	M+412,397,191(100)	C <sub>30</sub> triterpane
J	M <sup>+</sup> ,410,395,367(5),218(45),205(47),191(100)	Neohop-13(18)-ene (IV,R= $iso-C_3H_7$ )
K	M <sup>+</sup> 398,383,191(50),177(100)	30-Norhopane (V,R=C <sub>2</sub> H <sub>5</sub> )
L	M <sup>+</sup> 424,409,367(60),231(100),191(95),189(50)	Homohop-17(21)-ene (III,R=sec-C <sub>4</sub> H <sub>9</sub> ) <sup>c</sup>
М	M <sup>+</sup> 424,409,367(85),231(100),191(80),189(40)	Homohop-17(21)-ene (III,R=sec-C <sub>4</sub> H <sub>9</sub> ) <sup>c</sup>
N	M+426,411,369(2),205(28),191(100)	17αH-homohopane (V,R=sec-C <sub>4</sub> H <sub>9</sub> )
0	M <sup>+</sup> 412,397,369(4),191(100)	Hopane (V,R=iso-C <sub>3</sub> H <sub>7</sub> )
Р	M <sup>+</sup> 426,411,369(5),205(80),191(100)	Homomoretane (V,R=sec-C <sub>4</sub> H <sub>9</sub> ,21aH)
Q	M+426,411,369(5),205(100),191(50)	Homohopane (V,R=sec-C <sub>4</sub> H <sub>9</sub> )
R	M <sup>+</sup> 440,425,219(100),191(55)	Bishomohopane (V,R=sec-C <sub>5</sub> H <sub>11</sub> ) <sup>c</sup>

<sup>a</sup>See Figure 4.



Figure 5. Total ion current traces of aromatic hydrocarbon fractions from Cores 402A-24 and 30.

components.  $\beta$ -Hydroxy acids are only minor components of the OH FAME, and only in the case of *n*-C<sub>140</sub> could positive evidence be obtained from the mass spectral data. No evidence for significant quantities of  $\omega$ - or polyhydroxy acids was obtained.

## DISCUSSION

## Straight and Branched Chain Components

The n-alkanes and n-alkanoic acids (Figures 2 and 7) show bimodal distributions. The odd/even predominance in

TIC

750

TIC

750



Figure 6. Mass fragmentograms of m/e 252 for the aromatic hydrocarbon fractions from Cores 402A-18 and 30.

the higher *n*-alkanes (>C<sub>21</sub>) is indicative of a higher plant contribution, as is the even/odd predominance in the longer chain *n*-alkanoic acids (>C<sub>20</sub>). The major difference between the three samples in this carbon number range is the greater proportion of higher plant-derived components in Core 402A-30, presumably reflecting a slightly greater contribution of such organisms.

The absence of a unimodal distribution of *n*-alkanes with CPI  $\simeq 1$  reflects the known (Deroo et al., this volume) immaturity of the samples. Similarly, the absence of a particularly abundant *n*-C<sub>17</sub> alkane is in agreement with either a low algal input to the forming sediments or extensive bacterial degradation at that time. It is difficult to distinguish these two possibilities: the former is in agreement with the low abundance of morphologically recognizable algal debris (Doran et al., this volume). The latter situation has been observed in a depth study of contemporary algal mats (Cardoso et al., 1976).

The presence of high relative abundances of *iso*- and *anteiso*-acids is typically taken as evidence of a bacterial contribution to sediments (e.g., Leo and Parker, 1976). The occurrence of low relative concentrations in the samples from Hole 402A is, however, unlikely to be the result of low bacterial input because many gram-negative species (particularly abundant in the marine environment) often show low abundances of these acids relative to  $C_{16:0}$ ,  $C_{18:0}$ ,  $C_{18:0}$ ,  $C_{18:1}$  (Oliver and Colwell, 1973).

Although the major input to the samples is thought to derive from higher plants (Deroo et al., this volume), no evidence for significant quantities of long chain  $\omega$ -hydroxy or polyhydroxy acids could be obtained; such components are typical of higher plant cutin and suberin (Hunneman and Eglinton, 1972). This suggests that little or no hydrolysis of these plant biopolymers has occurred during diagenesis. Examination of the bound hydroxy acids is in progress.

## **Cyclic Components**

The most striking feature of the sterenes, triterpenoid alkanes and alkenes, and triterpanoic and triterpenoic acids is the similarity in the distributions of each of these structurally specific classes in all three samples. These similarities suggest that the environments of deposition of the black shale layers were similar although the periodic deposition was separated by a time span of up to ca.  $7 \times 10^6$  years.

The high relative abundance of unsaturated components to saturated components is again in agreement with the immaturity of the sediments.

The occurrence of high relative abundances of  $\Delta$  <sup>13(17)</sup> rearranged sterenes (I) is noteworthy. The presence of these diagenetic products of sterols is thought to indicate the presence of acidic conditions within immature sediments (Dastillung and Albrecht, 1977).

The distribution of triterpanes is as expected for an immature sediment, with the thermodynamically less stable  $17\beta$ H components predominating over their  $17\alpha$ H counterparts (e.g., Ensminger et al., 1977).

The distribution of triterpene hydrocarbons in the samples is unusual in that a number of the components have not been reported previously to occur in sediments: the  $\Delta$  <sup>13(18)</sup> alkenes; 22,29,30-trisnorneohop-13(18)-ene (IV, R=H) and 30-norneohop-13(18)-ene (IV, R=C<sub>2</sub>H<sub>5</sub>); and 30-norhop-17(21)-ene. If the proposed assignments of the  $\Delta$  <sup>13(18)</sup> components are correct, their occurrence may relate to the reported conversion of hop-17(21)-ene (III, R = *iso*-C<sub>3</sub>H<sub>7</sub>) to the corresponding  $\Delta$  <sup>13(18)</sup> rearranged hopene skeleton (IV, R=*iso*-C<sub>3</sub>H<sub>7</sub>) under acidic conditions (Berti and Bottari, 1968). This is borne out by the co-occurrence of the C<sub>27</sub>, C<sub>29</sub>, and C<sub>30</sub> members of each series.

In an analogous manner to the triterpenoid hydrocarbons, the predominance of the  $17\beta$ H components and the presence





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Figure 7. Distributions of n-alkanoic (unbroken lines), n-C<sub>16:1</sub>, and n-C<sub>18:1</sub> (dotted lines) acids in core samples: (A) 402A-18, (B) 402A-24, (C) 402A-30.

of triterpenoic acids reflects the immaturity of all three samples (Ensminger, 1977).

Triterpenoid-derived aromatic hydrocarbons have not been reported to occur in the biosphere; in the sediments they are, therefore, likely to be diagenetic products formed at low temperatures. In contrast to the aliphatic triterpenoids, the distributions show differences among the samples. These differences appear to relate either to depositional environment or to diagenetic alteration, or both: (1) Core 18 contrasts with 24 and 30 in that no nuclear-methylated compounds are present. This must reflect a difference in input by way of nuclear-methylated

TABLE 4 Triterpenoid Acids of Core 402A-30

Tentative Assignment (ACID) <sup>b</sup>	Diagnostic Ions	
17αH,21βH-homohopanoic (XIIIa)	M <sup>+</sup> 470,455,369(3),249(64),191(100)	A
Bishomohop-17(21)-enoic (XIV)	M <sup>+</sup> 482,467,367(41),231(100),191(84)	в
17βH,21αH-homohopanoic (XIIIa)	M <sup>+</sup> 470,455,369(4),249(84),191(100)	C
Bishomohop-17(21)-enoic (XIV)	M <sup>+</sup> 482,467,367(57),231(100),191(83)	D
17aH,21BH-bishomohopanoic (XIIIb)	M*484,469,369(4),263(29),191(100)	E
176H,216H-homohopanoic (XIIIa)	M <sup>+</sup> 470,455,369(5),249(100),191(73)	F
17βH,21αH-bishomohopanoic (XIIIb)	M <sup>+</sup> 484,469,369(9),263(85),191(100)	G
17aH.21BH-trishomohopanoic (XIIIc	M*498,483,369(3),277(30),191(100)	н
176H,216H-bishomohopanoic (XIIIb)	M*484,469,369(9),263(100),191(67)	I
17BH,21aH-trishomohopanoic (XIIIc	M*498,483,369(13),277(100),191(94)	J
176H,216H-trishomohopanoic (XIIIc)	M*498,483,369(7),277(100),191(65)	K
176H.216H-C24 (XIIId)	M*512,497,369(4),291(100),191(69)	L

See Figure 8.

<sup>b</sup>By spectral interpretation, except for compound I for which a standard was available.

precursor triterpenoids, (2) taken with the overall similarity in total aromatic hydrocarbon distributions in Cores 24 and 30 (Figure 5), the presence in 30 of components VII, VIII and XI with 2 or 3 aromatic rings, rather than the analogs VI, VII, and X (1 or 2 aromatic rings) in 24, suggests a progressive aromatization with depth, (3) Core 18 alone shows two minor components (tentatively assigned as XII) whose skeletons are compatible with an origin in higher plant terpenoid precursors, and a high relative abundance of perylene.

Unfortunately, insufficient paleontological and sedimentary data are available for Cores 402A-18, 24, and 30 in order to relate the differences in the distributions of the aromatic hydrocarbons to these data. The triterpenoids of the hopane family found in sediments are thought to arise from tetra- and pentahydroxy bacteriohopanes (e.g., XV) (e.g., Van Dorsselaer et al., 1974), which have been shown to be major constituents of bacteria and blue-green algae (e.g., Rohmer and Ourisson, 1976); in the marine situation, where blue-green algae are not dominant in the phytoplankton, the likely source of hopane triterpenoids is bacteria. This is borne out by the low concentrations of visibly distinguishable algal debris in samples from Hole 402A (Doran et al., this volume). The only compounds which could be assigned to a higher plant origin are the high molecular weight n-alkanes and n-alkanoic acids in all three cores, and the minor components XII in Core 18. It is perhaps surprising that only traces of cyclic diterpenoid components were observed (e.g., from m/e 191 fragmentograms in Figures 4 and 8). The majority of the compounds assigned in the solvent-soluble lipid fractions are compatible with a bacterial origin. This contrasts with the terrigenous origin proposed for the bulk of the organic matter (Deroo et al., 1977) and therefore suggests extensive bacterial contribution and possible reworking at the time of deposition.

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IIIa



















 $\label{eq:XIIIa, R = CO_2H; b, R = CH_2CO_2H; c, R = (CH_2)_2CO_2H; d, R = (CH_2)_3CO_2H}$ 

