

## 8. X-RAY MINERALOGY OF DEEP SEA DRILLING PROJECT LEGS 51 THROUGH 53, WESTERN NORTH ATLANTIC

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### INTRODUCTION

A series of 140 samples from Holes 417A, 417B, 418A, and 418B were examined semiquantitatively for their X-ray mineralogy. We selected most samples from the dominant lithologies rather than by coring the same sedimentary sequence at Sites 417 and 418, only 8 km apart (Figure 1). Both sites are located at the foot of the Bermuda Rise (Figure 2) along the M0 magnetic anomaly on crust of lower Aptian age. Sediments are mainly abyssal pelagic clays, with various authigenic minerals low in carbonates.

### METHODS

#### Analysis Conditions

The mineralogical data in Tables 1 and 2 are for samples dried at 70°C, powdered and X-rayed under the following conditions: CuK $\alpha$  radiation, nickel filter, 36 kv/24mA, and 1° detection slit. Goniometer speed was 0.5° = 2θ/min, paper speed 10 mm/min. All samples were X-rayed from 3° to 40° = 2θ.

Clay minerals were identified on oriented glass smear mounts. To obtain the optimal texture, the <2 μm fraction was first disintegrated with a KLN 582 ultrasonic generator and then centrifuged to clearness. The residue was mixed to a homogeneous paste and finally placed on the glass slide. It was X-rayed in an untreated state and glycolated, when necessary, after heating.

#### Factor Analysis

For the bulk mineralogy, the factors used are those listed in Table 3. When multiplied by their peak heights, they add up to 100 per cent. Because of the large amount of amorphous material — i.e., volcanic glass, opal, and amorphous clay minerals — and without using the monochromator, diffuse scattering was relatively high. Thus, if these components vary, the factors have to be redetermined. These factors can only be used for sedimentary types that are very similar to those above; for example, they cannot be applied to nearshore sediments on account of their higher crystallinity.

For the clay mineralogy, we used the method discussed by Biscaye (1965). The factors in Table 4 were multiplied by the peak area, determined planimetrically; the sum of

these was 100 per cent. The kaolinite/chlorite peak at 12.3 Å = 2θ was divided by peak splitting at 24 to 25° = 2θ — a factor obtained from the ratio:

$$\frac{3.58 \text{ Å peak area}}{3.58 \text{ Å} + 3.54 \text{ Å peak area}}$$

The carbonate contents were determined gasometrically by carbonate bomb (Müller and Gastner, 1971). X-ray numerical data for all holes are presented in tabular form on the core barrel sheets following the site report chapters. Mineral identifications were made manually and peak areas measured with the aid of a template having graduated triangles. The amount of amorphous material was not determined quantitatively.

### RESULTS

X-ray results are listed in Tables 1 and 2. A graphic representation in Figures 3 through 5 places the samples at their proper stratigraphic depths.

Unit 1 and 2 pelagic clays are characterized mainly by a chlorite component in the dominant clay universal spectrum. Rhodochrosite and dolomite appear in the base of Unit 2 above the first clinoptilolite occurrences in Unit 1. Palygorskite occurs in the nonzeolitic interval of mid-Eocene radiolarian clays. Multicolored Cretaceous clays are rich in clinoptilolite and cristobalite from former radiolarian tests.

Palygorskite increases in abundance downhole in the Upper Cretaceous. In the Middle Cretaceous black clay facies, the clay-mineral assemblage changes to dominantly mixed-layer varieties; illite peaks are asymmetric; zeolites plus cristobalites disappear; and quartz (and, periodically, calcite) becomes an important constituent. Several authigenic phases were detected, including pyrite, gypsum, barite, and apatite (phosphates).

### REFERENCES

- Biscaye, P. E., 1965. Mineralogy and sedimentation of recent deep-sea clay in the Atlantic Ocean and adjacent seas and oceans, *Geol. Soc. Am. Bull.*, v. 76, no. 7, p. 803-832.  
Müller, G. and Gastner, M., 1971. The "Carbonate Bomb," a simple device for the determination of the carbonate content in sediments, soils, and other materials, *N. Jb. Mineral.*, v. 10, p. 466-469.

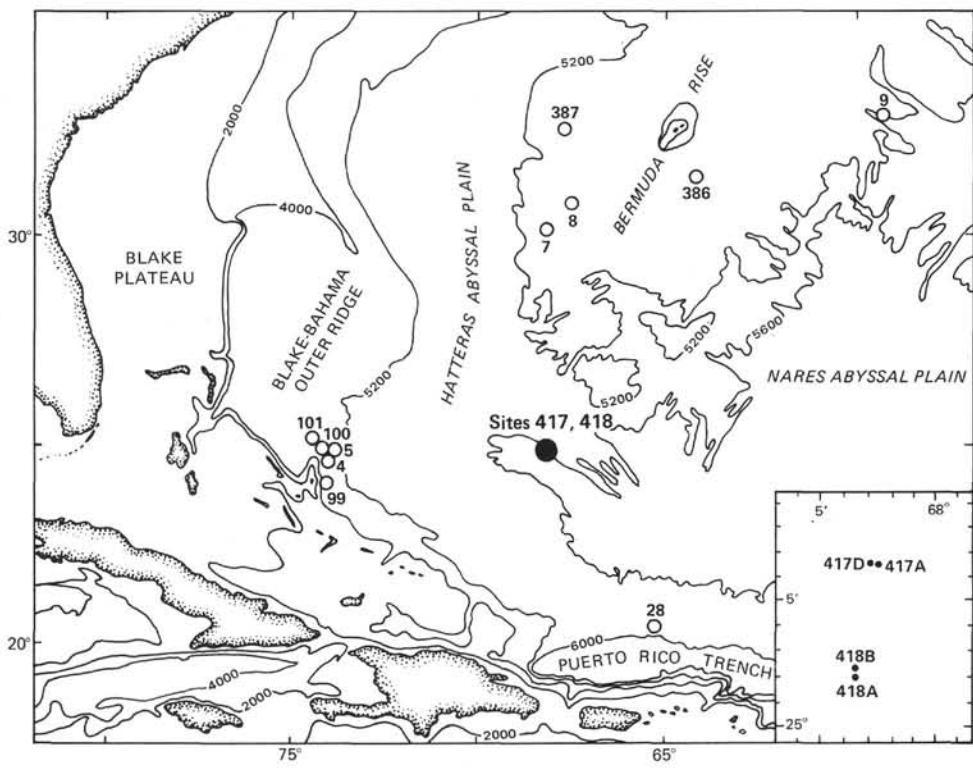


Figure 1. Location of Sites 417 and 418.

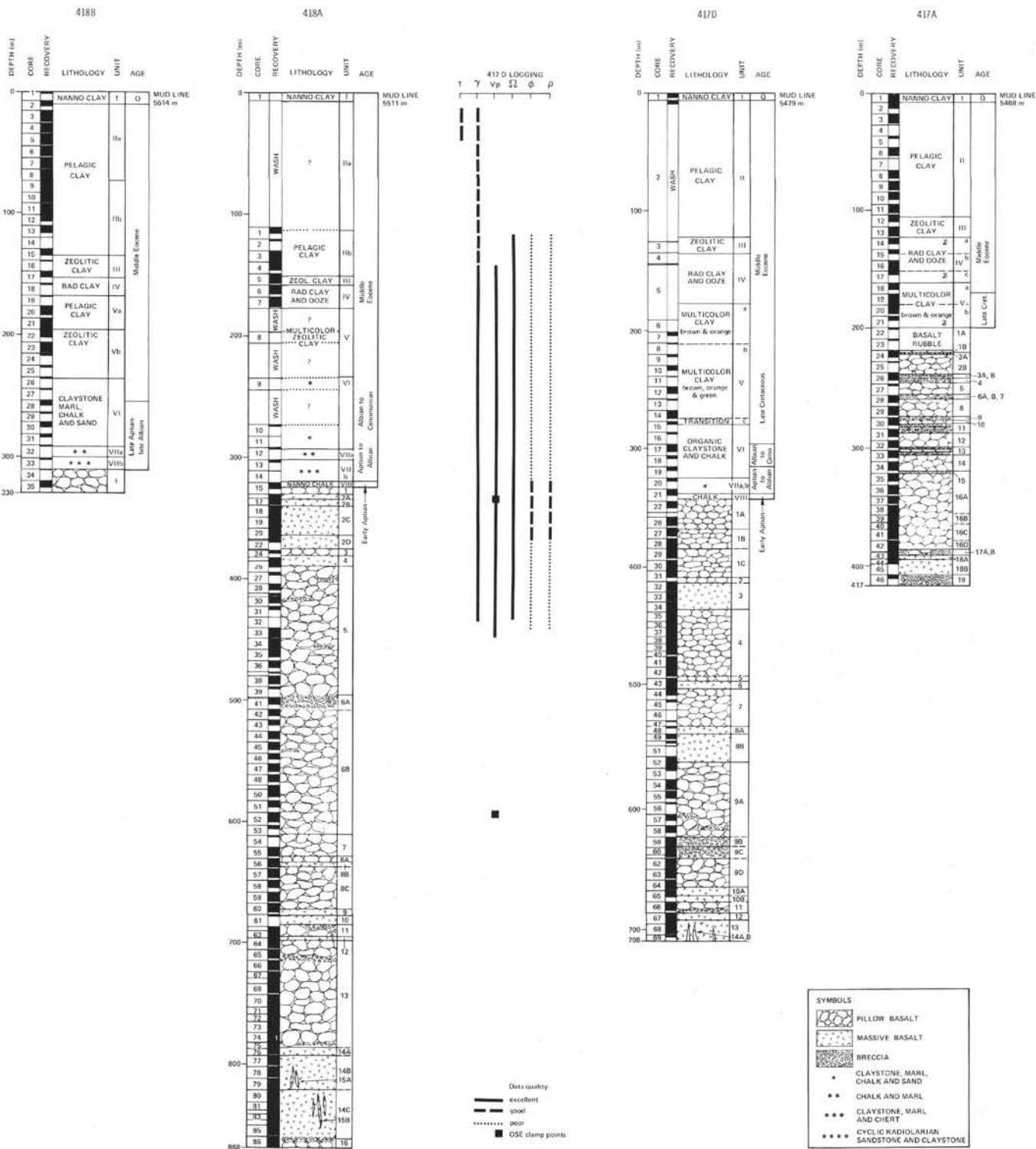


Figure 2. Stratigraphy of sediment and basement sections of Sites 417 and 418.





**TABLE 3**  
**Factor for Minerals Investigated, Bulk Mineralogy**

Mineral	Peak 2 $\theta$ (in degrees)	d (in Å)	Factor
Quartz	26.7	3.34	1.3
Feldspar	27.4-28.0	3.25-3.18	2.0
Pyroxene	29.9	2.99	2.0
Hornblende	10.5	8.42	2.0
Cristobalite	21.5-21.9	4.15-4.05	1.5
Clinoptilolite	9.8	8.99	2.0
Pyrite	33.0	2.71	2.0
Gypsum	11.7	7.56	2.0
Rhodochrosite	31.4	2.85	2.0
Barrite	25.9	3.44	2.0
Apatite	32.2	2.78	2.0

**TABLE 4**  
**Factors for Minerals Investigated, Clay Mineralogy**

Mineral	Peak 2 $\theta$ (in degrees)	d (in Å)	Factor
Smectite	5.2	17.0	1
Chlorite	12.3	7.2	2
Illite	8.8	10.0	4
Kaolinite	12.3	7.2	2
Palygorskite	8.4	10.5	1
Talc	9.4	9.3	1
Sepiolite	7.4	12.0	1
Mixed layer	Same as its parts		

## LEG 51 HOLE 417A

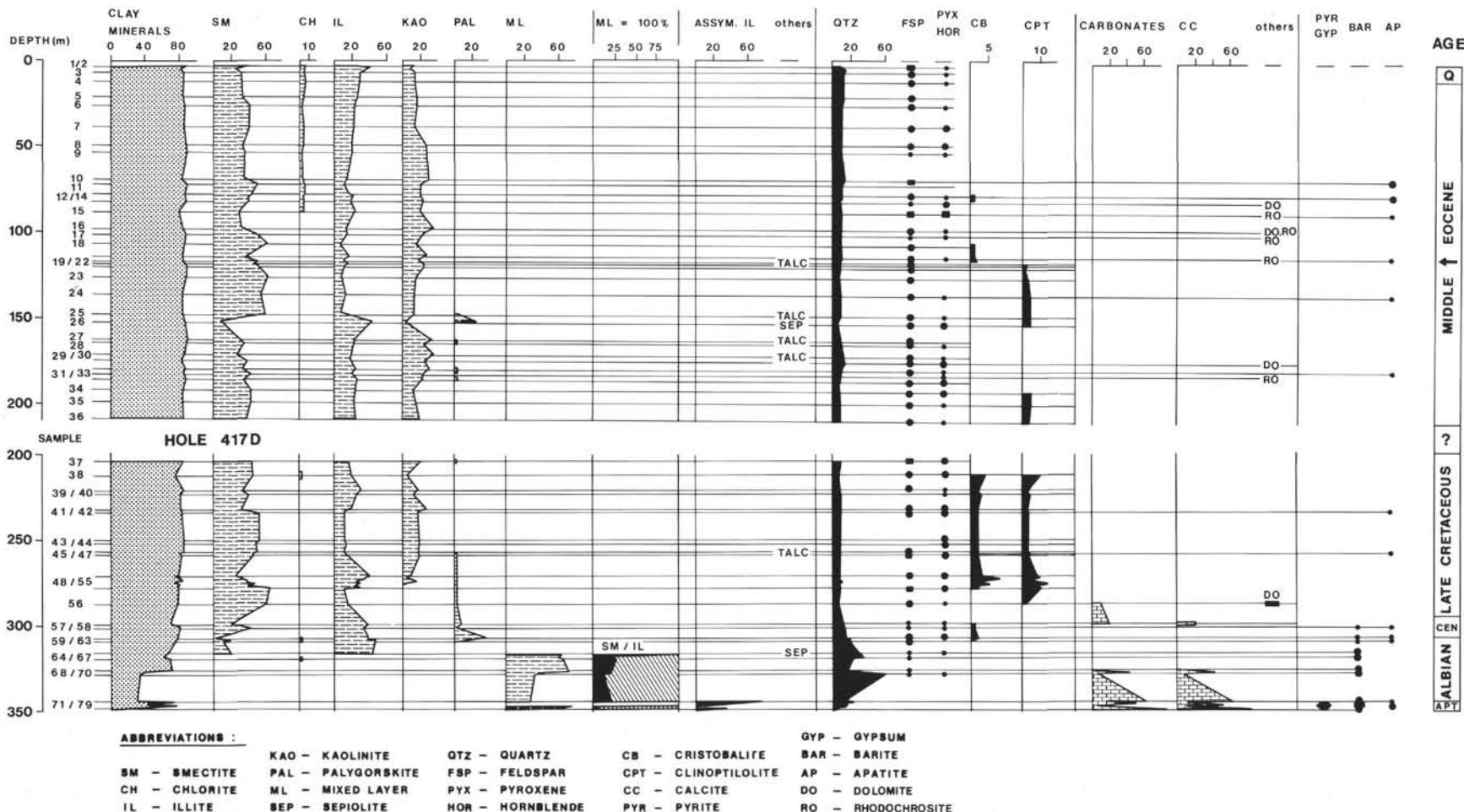


Figure 3. X-ray diffraction data, Hole 417A.

## LEG 52 HOLE 418A

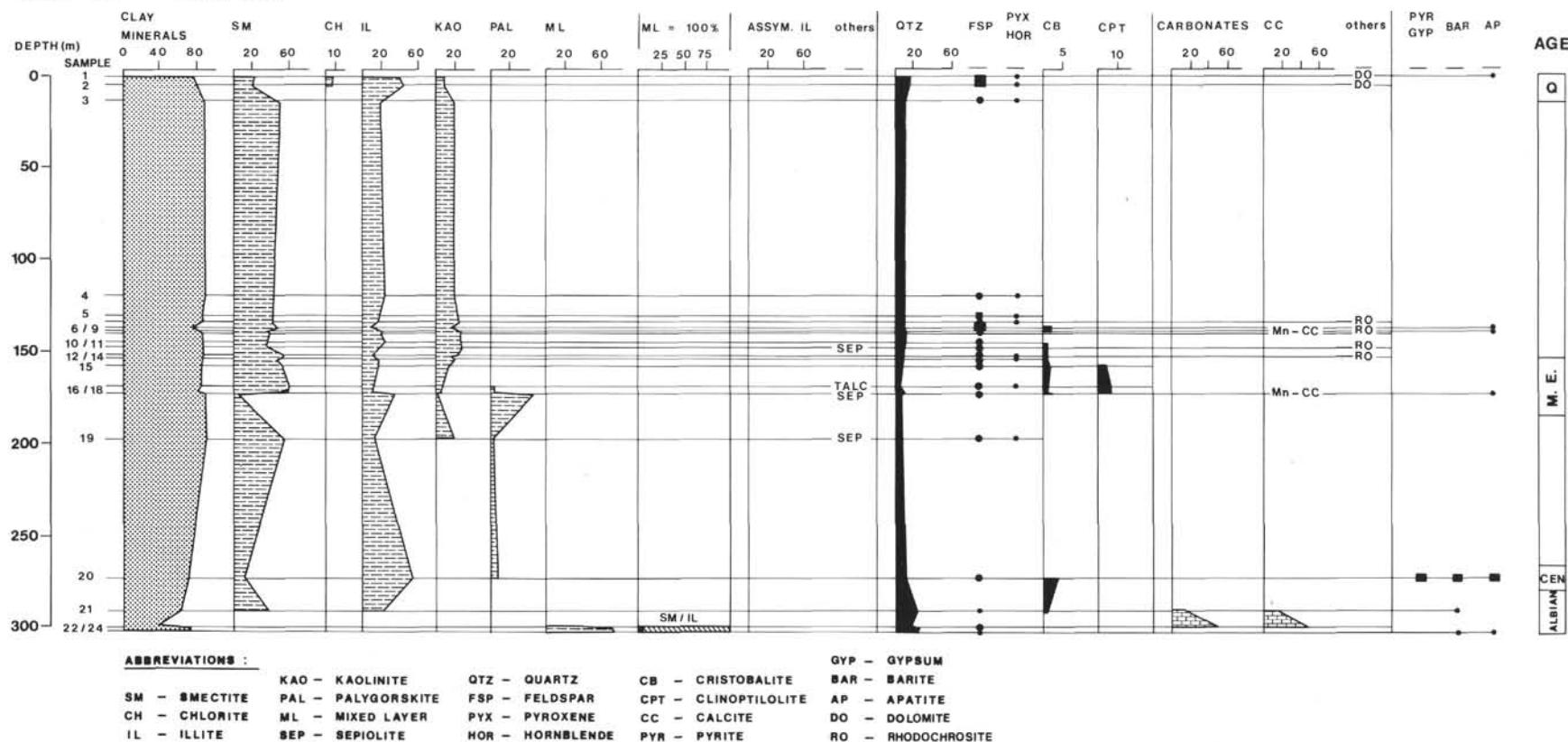


Figure 4. X-ray diffraction data, Hole 418A.

## LEG 53 HOLE 418 B

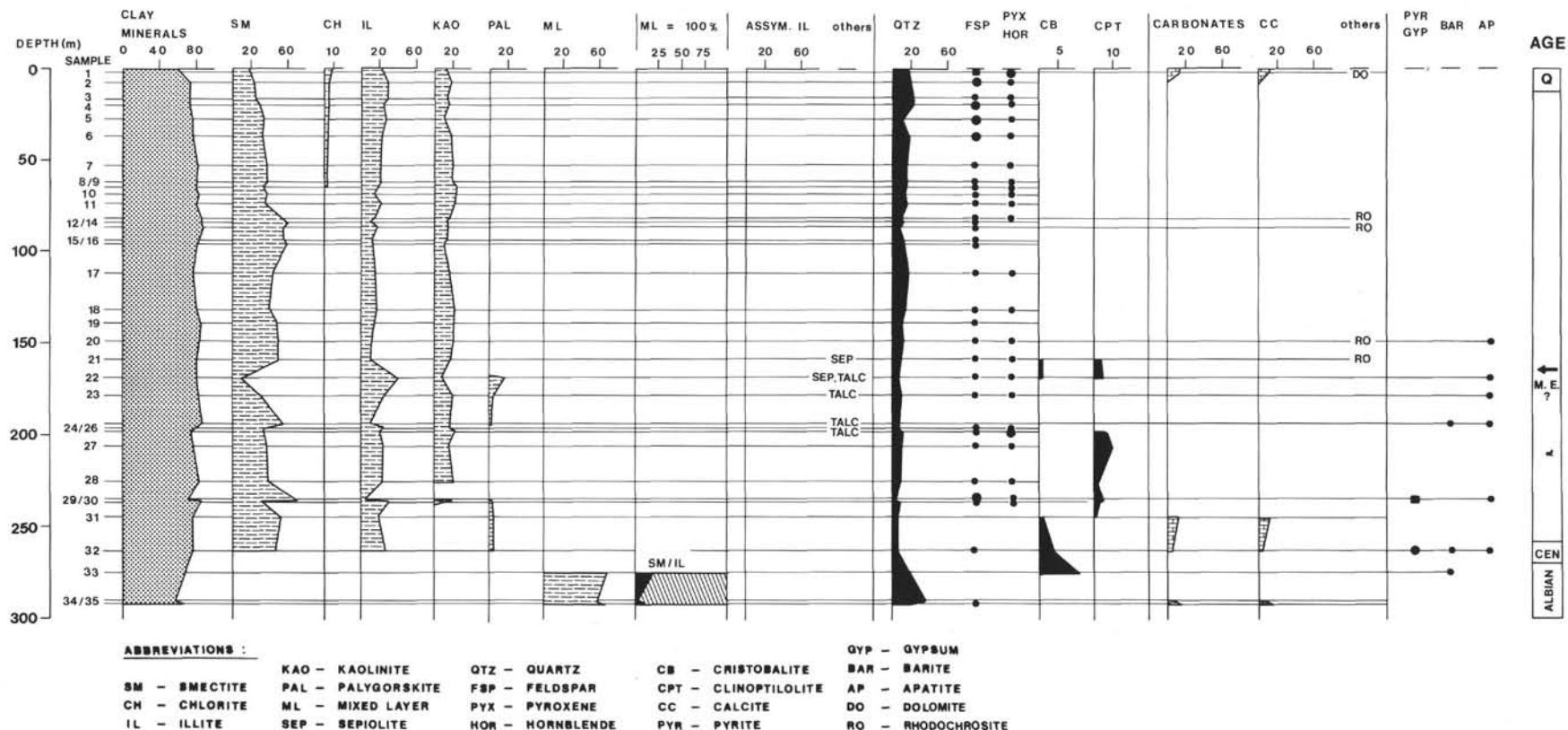


Figure 5. X-ray diffraction data, Hole 418B.