

4. SITE 440: JAPAN TRENCH MIDSLOPE TERRACE, LEG 57

Shipboard Scientific Party¹

HOLE 440

Date Occupied: 10 November 1977
Date Departed: 10 November 1977
Time on Hole: 1 day
Position: 39°44.33'N; 143°55.74'E
Water Depth (sea level): 4509 corrected meters, echo sounding
Water Depth (rig floor): 4519 corrected meters, echo sounding
Bottom Felt: 4517 meters, drill pipe
Penetration: 73.0 meters
Number of Cores: 8
Total Length of Cored Section: 73.0 meters
Total Core Recovery: 50.4 meters
Core Recovery: 69 per cent
Oldest Sediment Cored:
Depth sub-bottom: 73.0 meters
Nature: Diatomaceous silty clay
Age: Pleistocene
Measured velocity: 1.5 km/s

HOLE 440A

Date Occupied: 11 November 1977
Date Departed: 11 November 1977
Time on Hole: 1 day
Position: 39°44.13'N; 143°55.74'E
Water Depth (sea level): 4509 corrected meters, echo sounding
Water Depth (rig floor): 4519 corrected meters, echo sounding
Bottom Felt: 4517 meters, drill pipe
Penetration: 139.5 meters
Number of Cores: 7

Total Length of Cored Section: 66.5 meters
Total Core Recovery: 33.5 meters
Core Recovery: 50 per cent
Oldest Sediment Cored:
Depth sub-bottom: 139.5 meters
Nature: Diatomaceous clay
Age: Pleistocene
Measured velocity: 1.5 km/s

HOLE 440B

Date Occupied: 12 November 1977
Date Departed: 19 November 1977
Time on Hole: 8 days
Position: 39°44.13'N; 143°55.74'E
Water Depth (sea level): 4509 corrected meters, echo sounding
Water Depth (rig floor): 4519 corrected meters, echo sounding
Bottom Felt: 4517 meters, drill pipe
Penetration: 814.0 meters
Number of Cores: 71
Total Length of Cored Section: 814.0 meters
Total Core Recovery: 401.9 meters
Core Recovery: 49 per cent
Oldest Sediment Cored:
Depth sub-bottom: 814.0 meters
Nature: Diatomaceous claystone
Age: Late Miocene
Measured velocity: 2.1 km/s

Principal results:

Site 440 is on the Japan Trench inner slope, 28 km from its axis and on the midslope terrace. The core recovery and biostratigraphic zonation were good, and a suite of downhole logs contributed greatly to interpretation of lithology and structure. Beneath the midslope terrace is a flat-lying nonfolded continental-slope section which is probably not material accreted at the leading edge of the upper plate. The sediment section is of fairly uniform hemipelagic mudstone similar in composition to that at sites farther upslope. Over 800 meters were penetrated; sediment ranges in age from late Miocene through late Pleistocene.

The only breaks in continuous sedimentation are three or more periods of slumping in the Miocene and Pliocene. Ponding of turbidites and a twofold increase in sedimentation rate in the Pleistocene are consistent with the present topography of the midslope terrace.

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The section was deposited at bathyal depths near or below the carbonate compensation depth (CCD). A stress environment possibly of pre-Pleistocene age has tectonically brecciated the section at depths where the rock deforms brittlely. Although microfaulting was observed at other sites in the transect, it was not as intense and pervasive at relatively shallow depths. The reflective section above acoustic basement is at least 1600 meters thick. The estimated sediment age at this depth is lower Miocene, if sedimentation continued without hiatus and at a nearly constant rate during the early Neogene, as it did at sites upslope.

The fractured but nonfolded block 1.6 km thick and at least 6 km wide may have slumped from upslope. The seismic records and cores show a continuous Neogene continental slope section from Site 440 landward across the deep sea terrace almost to the shore. Therefore active Neogene accretion seems to have been restricted to the lower trench slope.

BACKGROUND AND OBJECTIVES

A principal objective of the Japan Trench Transect was sampling of the accretionary wedge in two or three places to establish trends of diagenesis and changes in the physical state of off-scraped sediments. Two sites had been drilled on Leg 56 to make such observations (Figure 1). At the Leg 56 Site 435 on the upper trench inner slope, undeformed Pliocene slope sediment was recovered, and the site was interpreted to be on a large slump block from farther up the slope. At Site 434 on the lower slope, deformed sediment was encountered. However, weather and unstable hole conditions forced termination of drilling at 600 meters, when the drill stem became stuck in a highly fractured Miocene diatomaceous mudstone section. Thus there was no evidence that an accreted section of oceanic sediment had been sampled.

Site 440 was placed between the two Leg 56 sites on a midslope terrace of the Japan Trench inner slope at a depth of 4500 meters. A series of faint subhorizontal reflections below the flat sea floor of the terrace indicated a relatively less deformed section and possibly more favorable hole conditions than those at Site 434. At the base of the faint reflections is a low frequency event that is similar to the acoustic basement reflection at Site 439. Thus the reflection might be a continuation of the unconformity between Cretaceous and Tertiary rock. On the other hand, this reflection was interpreted by the Leg 56 scientific staff to be oceanic crust uplifted by imbrication. This acoustic basement was the target objective at Site 440, although it was recognized that it was probably beyond the reach of a single bit hole and possibly within reach of a multiple re-entry hole. The decision on re-entry was deferred until the results of the single bit pilot hole could be evaluated, indicating whether the rocks encountered were part of an accretionary sequence. A primary consideration was whether deep penetration could be achieved, because previous experience showed caving of highly fractured material rather than drill bit life to be the primary cause of hole termination.

OPERATIONS

The 100-km transit from Site 439 to Site 440 was made during a gathering storm. The *Challenger* turned onto a course of 090° about 12 km west of the site. As was to be expected in these latitudes, the navigation aids at night were insufficient to correct for a set from a beam sea and strong wind, so there was no assurance that the ship was making the desired track after an hour of steaming along a dead-reckoned course. In addition, the seismic system and the 3.5-kHz records were essentially only noise because of sea conditions, so the beacon was not dropped during the first pass over the midslope terrace (Figure 2). A satellite position was not available during the following 4-hour period, and with sea conditions beyond the point where drill pipe would be lowered we decided to wait for a more opportune time to drop the beacon. This occurred about 5½ hours later, when the ship had been brought to a heading that provided a lee wind for personnel during beacon-dropping operations. At the time of the drop the position was recognized to be slightly south of the seismic record track line, and we planned to offset north of the beacon, essentially to the desired position.

After the ship had hoisted and pipe was run, the first beacon signal became intermittent and the partially buried bottom hole assembly had to be pulled to the mudline, after only 8 cores had been recovered, until the ship's position-keeping ability could be improved. This was soon accomplished, and Hole 440A was spudded in; but after only 7 cores had been recovered, the weather became too rough to drill with incomplete burial of the bottom hole assembly and this hole was also abandoned. Hole 440B was washed down to the previously attained depth and then cored until yet another storm forced termination after 70 cores had been recovered. However, the ship's position held well enough through the storm to keep pipe in the upper 200 meters of the hole. After about 24 hours, coring was resumed until the bit released unintentionally. A decision to log and abandon this site had already been made because of the desire to drill at least one more site during this leg, and partial logging was accomplished in the less than ideal conditions during a short interval between two storms.

LITHOSTRATIGRAPHY

Site 440 was drilled at a water depth of 4515 meters, 28 km west of the Japan Trench, on the trench inner slope. The site lies on a prominent midslope structural terrace that is about 5 km wide and extends parallel with the trench more than 50 km without structural offset. In seismic reflection records the site is underlain by nearly flat-lying reflectors that are only slightly deformed, whereas the outer edge of the midslope terrace marks the edge of an apparently much more deformed and seismically diffractive area that extends down to the trench floor.

Three holes were drilled at Sites 440 (Holes 440, 440A, and 440B) (Figures 1 and 2). Except in their intervals of overlap, the holes were cored continuously, and together the samples from them provide a relatively con-

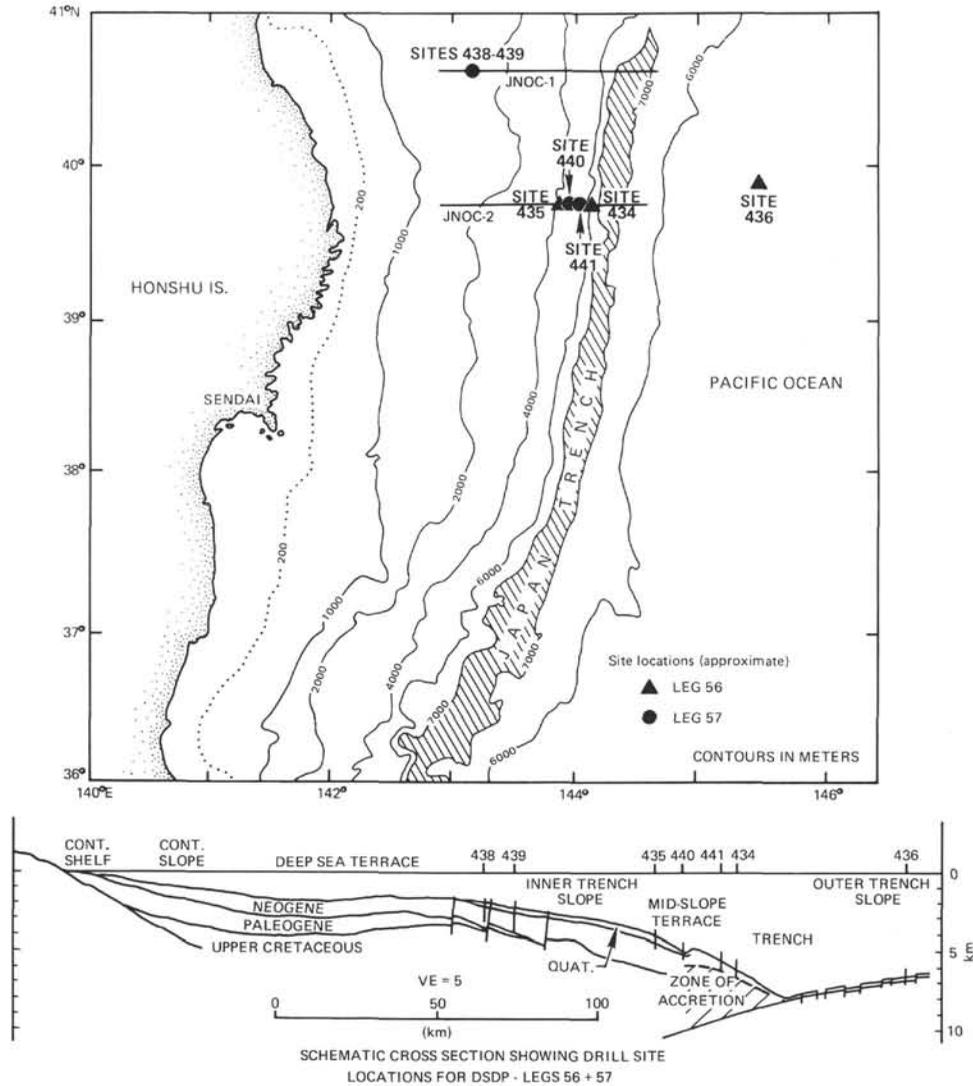


Figure 1. Location map showing position of sites drilled on Legs 56 and 57 and of JNOC multichannel seismic reflection profiles. Diagrammatic cross section of continental margin shown below (based largely on Ishiwada and Ogawa, 1976).

tinuous record of the stratigraphy and structural deformation of strata from the sea floor to the base of the lowest core in Hole 440B, which was taken from a sub-bottom depth of 808 meters.

The rocks have been divided into three units on the basis of their lithology (see Figure 3; Sites 435/440 Summary Chart [back pocket]; and Table 1). Unit 1 (0–38 m sub-bottom) consists of graded beds of silt, sand, and gravel with interbeds of clayey diatom ooze. Unit 2 (38–380 m) consists of claystone and diatomaceous claystone containing several zones of randomly dispersed pebbles. Unit 3 (380–808 m) consists of highly fractured claystone and diatomaceous claystone containing some folded beds deformed during mass movement.

Lithostratigraphic Unit 1 (Cores 440-1-440-5-2, 0-38 m sub-bottom, Holocene to upper Pleistocene)

The top 15 meters of Unit 1 consist of interbedded poorly sorted greenish-black to olive-gray (5G 2/1-5Y

4/2, wet) clayey sand and gravel, and firm dusky-yellow-green to olive-gray (5GY 5/2-5Y 3/2) diatomaceous clay and sandy diatom ooze. The remainder of the unit consists of grayish-olive (10Y 4/2) diatomaceous silty and sandy clay. Both parts contain graded beds. In the sand and gravel, these beds consist of matrix-supported units up to 80 cm thick, in which the largest clasts are 5 cm in diameter. In the sandy diatomaceous clay and ooze, the graded units are sand and silt beds about 2 cm thick that are spaced about 1 meter apart on the average. These show characteristics which are typical of turbidity current deposits.

The pebbles consist of subrounded to well-rounded clasts of light-colored plutonic rocks, andesite or dacite porphyry, mafic porphyry, altered metamorphic rocks, reddish-brown chert, pumice and concretionary claystone containing diatoms of earliest Pleistocene age. The diatoms in the claystone clasts are of the same time interval, represented at Site 438 (and Site 435) by a

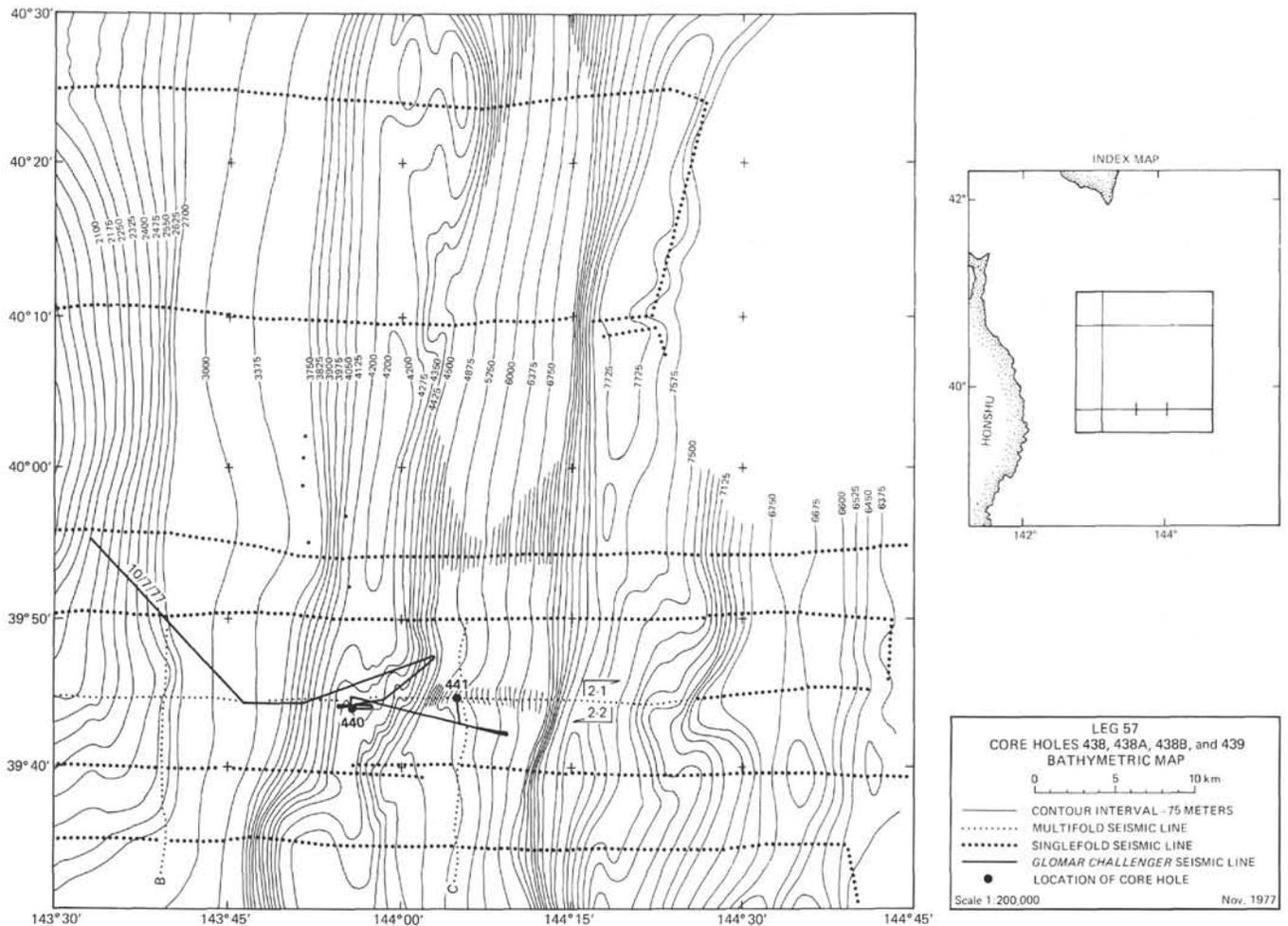


Figure 2. Bathymetric map of portion of Japan Trench inner slope showing location of Site 440, tracklines of Glomar Challenger, and traces of other available seismic lines in area. Contours based on network of geophysical data as shown.

hiatus. The claystone clasts probably were derived from farther upslope and redeposited here in deeper water. Successively deeper seismic reflectors appear to crop out at the slope above Site 440 (Figure 10, Geophysics); slumping and/or erosion has probably occurred on this slope.

Within the finer-grained beds of Unit 1, clay minerals average about 30 per cent, diatoms 21 per cent, quartz plus feldspar 20 per cent, sponge spicules 8 per cent, volcanic glass 6 per cent, heavy minerals 2 per cent, glauconite 2 per cent, and authigenic pyrite 0.5 per cent (Site Summary Chart, back pocket). Beginning in Core 4 at about 25 meters and extending into Unit 2 to a depth of 161 meters, the deposits are stained black by an ephemeral mineral, possibly the iron sulfide mackinawite. After a few hours of exposure to air, the black stain disappears.

Lithostratigraphic Unit 2 (Cores 440-5-2-8, 38-73 m sub-bottom, upper Pleistocene; Cores 440A-1-7, 82-139.5 m sub-bottom, upper to lower Pleistocene; Cores 440B-1-26, 139.5-380 m sub-bottom, lower Pleistocene to upper Pliocene)

Unit 2 is generally dark olive gray to olive gray (5Y 3/2-4/2). The upper half of the unit consists of diatomaceous clay and claystone, the lower half claystone. Its age is early Pleistocene and late Pliocene. At a sub-bottom depth of 168 meters, the cores could no longer be split with a wire, and material below that depth, cut with a power saw, by convention is considered to have been lithified into rock.

Randomly dispersed rounded pebbles, probably ice-rafted dropstones, occur at three intervals within Unit 2: at 91 to 96 meters, at 202 to 211 meters, and at 275 to

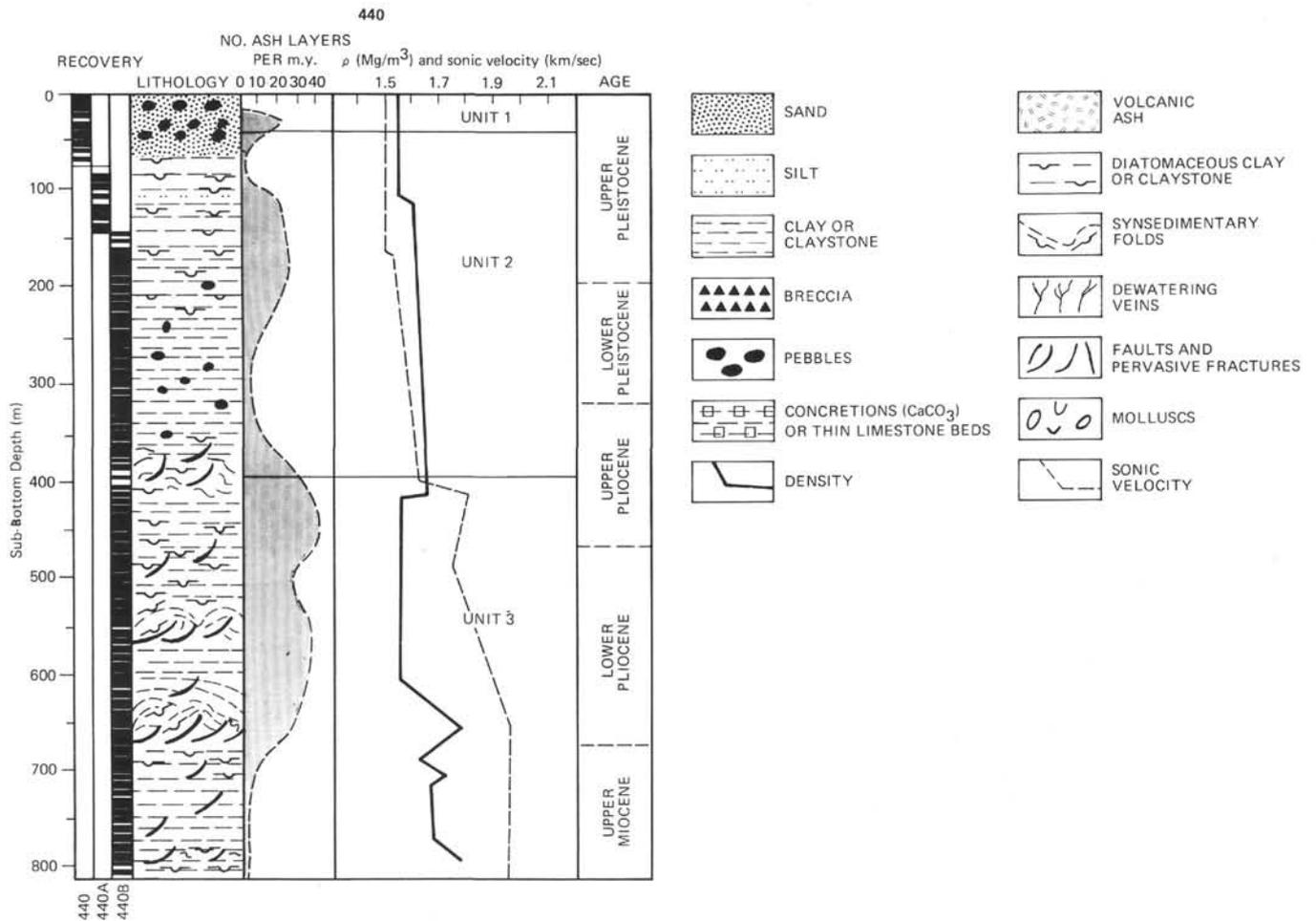


Figure 3. General lithology, lithologic units, core recovery, ash layer frequency, and physical properties from geophysical logs for Site 440.

TABLE 1
Lithologic Units, Coring Summary, Ages, and Description of Sequence at Site 440

Hole	Lithologic Unit	Core Numbers	Depth (m)	Age	Lithology
440	1	1 to 5	0.0-38.0	Holocene-upper Pleistocene	Greenish black to olive gray (5G 2/1-5Y 4/2) clayey sand and gravel, poorly sorted, from dusky yellow green to olive gray (5GY 5/2-5Y 3/2) diatomaceous clay, sandy diatom ooze, and grayish olive (10Y 4/2) diatomaceous silty and sandy clay; some graded bedding in coarser-grained intervals. Rounded pebbles of various igneous and sedimentary rocks common. Some thin ash layers.
440	2	5 to 8	38.0-480.0	Lower Pleistocene-upper Pliocene	Mostly dark olive gray to olive gray (5Y 3/2-4/2) diatomaceous clay and (below 168 m) diatomaceous claystone. Randomly dispersed rounded pebbles common; burrow mottling increases downhole. Minor thin ash layers.
440A	1 to 7				
440B	1 to 26				
440B	3	27 to 71	380.0-808.0	Lower Pliocene-upper Miocene	Generally olive gray to dark olive gray (5Y 4/2-3/2) interbedded claystone and diatomaceous claystone. Zones of folded beds (due to mass movement), veins, and microfaults. Thin (1-cm) ash layers common throughout.

352 meters. The clasts are of rock types similar to those that are concentrated near the top of Unit 1. The pebble intervals of Units 1 and 2 correspond roughly with strata containing cold-water planktic foraminiferal assemblages (see Keller, this volume). They may record

past glaciations; if so, several are recorded here, the oldest in the late Pliocene.

In typical beds of Unit 2, clay minerals average about 60 per cent, quartz plus feldspar 15 per cent, diatoms 10 per cent, volcanic glass 6 per cent, sponge spicules 3 per

cent, heavy minerals 1 per cent, glauconite 1 per cent, and pyrite 1 per cent (Site Summary Chart, back pocket).

Unit 2 is slightly to intensely mottled, the average level of mottling being moderate. Bedding disturbances identifiable as burrows increase toward the base of the unit. This may in part be due to the method of core-slicing. Sawing and careful cleaning exhibits structures much more clearly. Sponge fragments are seen in most cores.

Lithostratigraphic Unit 3 (Cores 440B-27-440B-71, 380-808 m sub-bottom, lower Pliocene to upper Miocene)

Unit 3 is generally olive gray to dark olive gray (5Y 4/2-5Y 3/2) and consists of interbedded claystone and diatomaceous claystone. A distinguishing feature of this unit is the presence of several zones of folded beds that may have been caused by mass movement. A typical example occurs in Core 43 (541-544 m sub-bottom), which is characterized by a series of recumbent folds. These folds are all associated with veins and faults that were clearly formed after lithification of the rock, suggesting that more than one episode of deformation may have occurred along certain zones within the unit. Veins, fractures, and faults occur nearly throughout Unit 3 to the bottom of the hole. Most of these features are rehealed, but many open fractures in cores tend to parallel the directions of the healed fractures (see Arthur, Carson, and von Huene this volume).

In typical beds of Unit 3, clay minerals average 70 per cent, quartz plus feldspar 12 per cent, diatoms 11 per cent, volcanic glass 4 per cent, sponge spicules 4 per cent, heavy minerals 0.5 per cent, glauconite 0.5 per cent, and pyrite 0.5 per cent (Site Summary Chart, back pocket).

Mottling is commonly intense in Unit 3, and burrows occur throughout the unit. Sponge fragments occur in most cores near the top of the unit; they are somewhat less common near the base.

Ash and Tuff Layers

Core recovery was sufficiently good at Site 440 that the number of ash and tuff layers per million years can be plotted with reasonable accuracy (Figure 4; see also Cadet and Fujioka, this volume) despite the effects of bioturbation on disruption of layers and diagenesis deeper in the hole. The record of volcanogenic deposition is roughly similar to that which has been noted at previous drilling sites in this area (see Kennett et al., 1977). Volcanic activity increased near the end of the Miocene and continued at a high level through the first two thirds of the Pliocene. The activity waned during the last part of the Pliocene and in the early Pleistocene. Renewed activity apparently characterizes the remainder of the Pleistocene.

Ash and tuff layers are rarely more than 1 cm thick at this site. Most of the ash probably settled through the water column after having been transported to the area through the atmosphere. Some graded beds, however, are intimately mixed with terrigenous material and were probably transported and redeposited as turbidites.

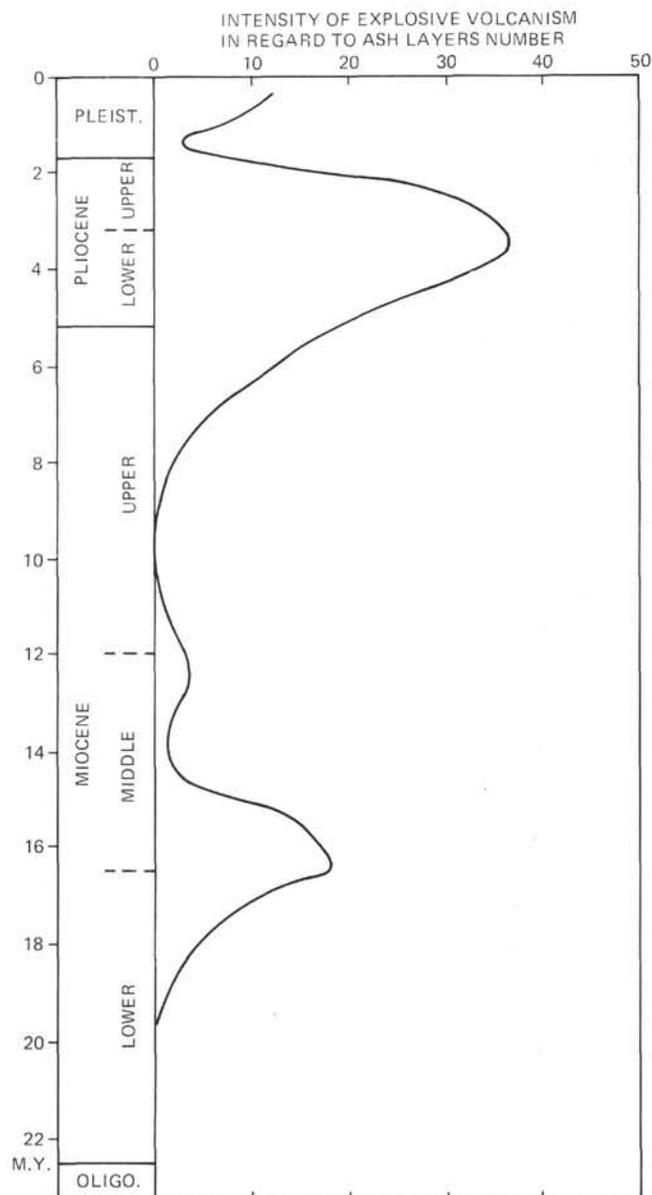


Figure 4. Ash layer frequency through time at Leg 57 sites (see Cadet and Fujioka, this volume).

Pumice clasts, which presumably first floated to this area, became waterlogged, and subsequently sank are dispersed throughout the section. Their abundance roughly mirrors the frequency of eruptions as recorded by the ash and tuff layers.

Structural Geology

In addition to the previously discussed synsedimentary small-scale folding that has affected the rocks of Unit 3, that unit has also been extensively faulted and fractured (see Arthur et al., this volume). Several different types of displacement have taken place at different times. One of the most characteristic types consists of healed microfaults, along which mottles and thin beds are offset from 1 mm to a few centimeters. In places where this type of faulting, along with other

types, has been especially intensive, there is some brecciation of the rock.

Closely associated with the healed microfaults are black-stained veins. Some of these veins are older than the microfaults, and the older ones, which are probably dewatering conduits, are joint-like and lack discernible displacement. Other black-stained veins contain gouge and follow fractures that offset the microfaults. Both types of veins branch, and many of them are sinuous.

A third set of fractures consists of offsets along which parting occurs. These seem to be the youngest structures, and they usually result in systematically broken core fragments. They commonly exhibit faint slickensides. Normal, reverse, and strike-slip movement occurs along the microfaults, black veins, and slickensided faults with inclinations ranging from horizontal to vertical.

The faults and syndimentary folds tend to be concentrated in several thick zones within Unit 3. Core recovery is poorer within these intervals. Anomalously high levels of ethane from 250 to 400 meters across the uppermost zone of fracturing may indicate that the fractures are interconnected (see Whelan, Hydrocarbons, this volume).

Sedimentation History and Depositional Environments

The gross lithology and stratigraphic succession at Site 440 is very similar to that of the other sites along the Japan Trench margin transect; hemipelagic sediment predominates. Diatoms compose a volumetrically important portion of sediment from the upper Pliocene through the Pleistocene but generally decline in importance downhole. According to biostratigraphic and paleomagnetic data, there is a continuous stratigraphic sequence from late Miocene through Pleistocene (~8 m.y.B.P. to recent; only an uppermost Miocene hiatus breaks the continuity). There is no evidence for tectonic repetition of units, and a series of subhorizontal seismic reflectors underlying the midslope terrace is evident in the multichannel seismic records.

The sedimentation rate during the late Miocene-Pliocene interval at Site 440 is comparable to that at Site 438 on the deep sea terrace (~110 m/m.y.); however, during the Pleistocene the sedimentation rate at Site 440 exceeds 230 m/m.y. which is much greater than that at any other site along the transect. In fact, upper Pliocene-lower Pleistocene hiatuses are found at Sites 438, 435, 441, and 434. Possibly the high sedimentation rates on the midslope terrace during the Pleistocene represent ponding of sediment behind a growing ridge. However, coarse turbidite sands occur only in the upper 38 meters of the section at Site 440. Redeposition of sediment from upslope is suggested by claystone clasts in some upper Pleistocene intervals bearing diatom floras of the late Pliocene-early Pleistocene. The presence of a few syndimentary slump folds indicates some redeposition from upslope during the early to late Pliocene.

The tectonically induced fractures, faults, and veins found below about 370 meters in the section apparently represent a tectonic jostling of the strata, but, as previ-

ously stated, the intense brecciation has not disrupted the stratigraphic continuity of the sequence.

BIOSTRATIGRAPHY

Introduction

An essentially continuous sequence of Quaternary through upper Miocene sediment was cored at Site 440 (Holes 440, 440A, and 440B). Microfossil assemblages provide excellent biostratigraphic control — comparable to that at Site 438. Diatoms are the most abundant microfossil group present; radiolarians range from very common to very rare; foraminifers are common above the upper Pliocene but rare to absent in the lower Pliocene and upper Miocene; and nannofossils range from abundant to rare in Pleistocene, upper Pliocene, and upper Miocene.

The relationship between the zonal assignments made for the different microfossil groups observed at Site 440 is shown in Figure 5.

The average sediment accumulation rate curve for Site 440 is shown in Figure 6. Sediment accumulation is greater above the 0.9 million year datum (~230 meters/million years) than before that datum (as low as ~100 meters/million years) and may reflect, in part, compaction of sediments.

No major hiatuses are recognized, although diatom evidence suggests a brief upper miocene hiatus between Cores 60, CC and 61, CC of Hole 440B.

Benthic foraminiferal assemblages indicate bathyal depths (greater than 2000 m) throughout the Pleistocene to upper Miocene sequence. The sparsity and generally poor state of preservation of calcareous plankton suggest deposition at or below the CCD, presently at approximately 3500 meters.

Diatoms

The *Denticula seminae* Zone (0–0.26 m.y.B.P.) is present down to Sample 440-5,CC. Reworking of earlier Pleistocene forms such as *Nitzschia fossilis*, *N. reinholdii*, *Rhizosolenia barboi*, *R. curvirostris*, and *Thalassiosira nidulus* is common in Samples 440-1,CC and 440-2,CC.

The *Rhizosolenia curvirostris* Zone (0.26–0.9 m.y. B.P.) occurs from Samples 440-6, CC through 440B-6, CC. The last occurrence of *Nitzschia reinholdii*, which marks the top of Subzone *a* and the base of Subzone *b*, is in Sample 440B-2,CC and has an estimated age of 0.63 m.y.B.P. (Burckle, 1977). The last occurrence of the silicoflagellate *Mesocenica elliptica*, which is estimated at 0.79 m.y.B.P. by Burckle (1977), occurs in Sample 440B-5,CC.

The *Actinocyclus oculatus* Zone (0.9–~1.4 m.y.B.P.) corresponds to Samples 440B-7,CC through 440B-7,CC. Following the criteria of Haq et al. (1977), the Pliocene/Pleistocene boundary is approximated by the base of this zone (Barron, this volume).

The *Denticula seminae* var. *fossilis* Zone (~1.7–2.43 m.y.B.P.) occurs from Samples 440B-18,CC through about 440B-25,CC. The base of this zone is placed at the

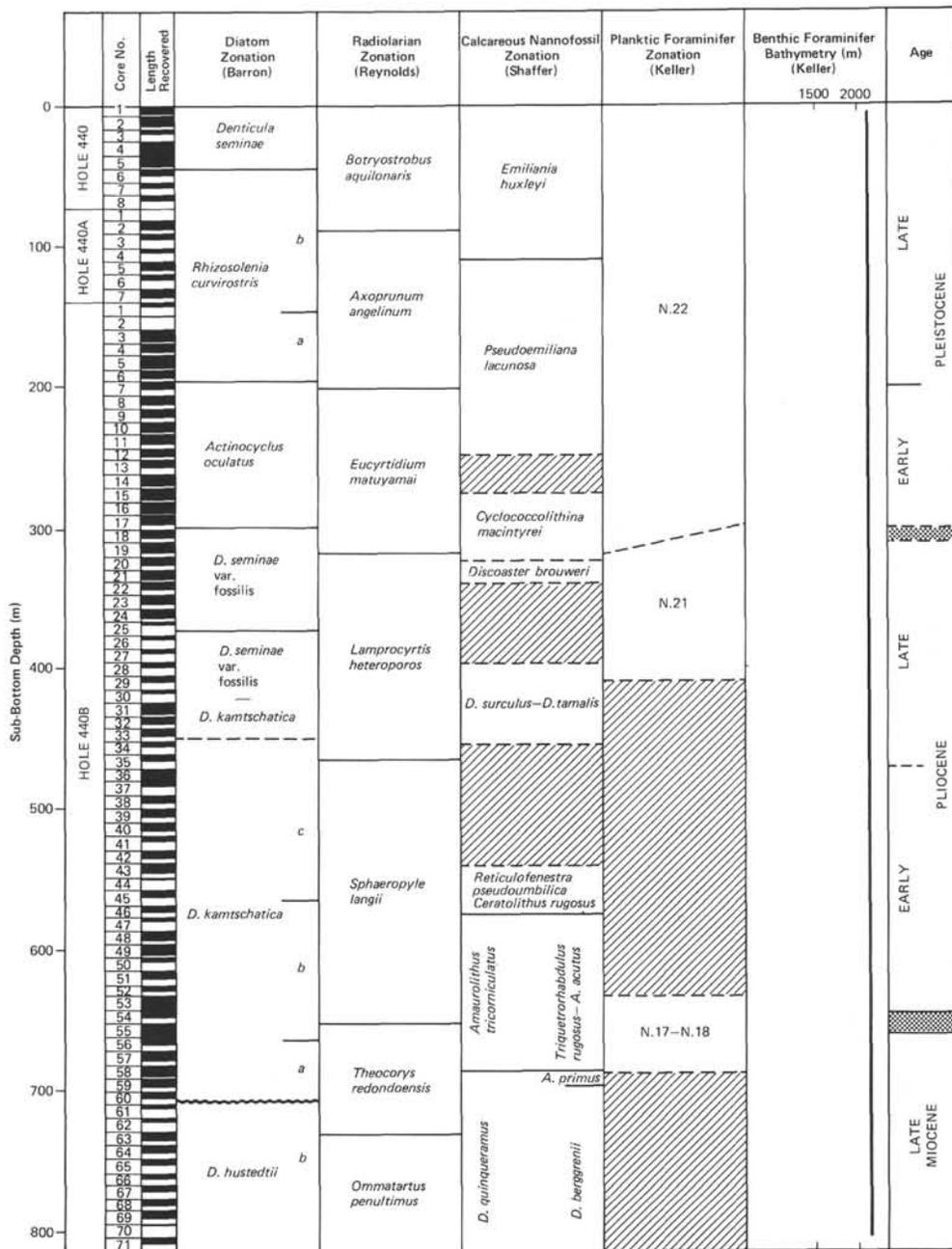


Figure 5. Planktic biostratigraphy and generalized paleobathymetry at Site 440.

last common *D. kamtschatica*; reworked *D. kamtschatica* is present throughout the uppermost Pliocene.

The *D. seminae* var. *fossilis*-*D. kamtschatica* Zone (2.43-~3.0 m.y.B.P.) is present from Samples 440B-25, CC through 440B-33, CC. The last occurrence of *Nitzschia jouseae* in Sample 440B-29, CC is assigned an absolute age of approximately 2.5 m.y.B.P. following Burckle's and Opdyke's (1977) correlation of this datum level with the uppermost part of the Gauss Magnetic Epoch. The last occurrence of *Thalassiosira nativa* in Sample 440B-33, CC corresponds to the base of the zone, and because it correlates well with its last occurrence in Hole 438A, it can be used to approximate

the base of the *Denticula seminae* var. *fossilis*-*D. kamtschatica* Zone (Barron, this volume).

The *D. kamtschatica* Zone occurs from Samples 440B-34, CC through 440B-60, CC (4.4-6.2 m.y.B.P.). The same three subzones identified in Hole 438A occur in the proper sequence. The Miocene/Pliocene boundary (~5.1-5.2 m.y.B.P.) is approximated by the last *Rouxia californica* (top of Subzone *a*) in Sample 440B-56, CC. The first *Thalassiosira oestrupii* in Sample 440B-56, CC lends support to this assignment (Barron, this volume).

The uppermost Miocene (Messinian event?) hiatus recognized in Core 42 of Hole 438A seems to be present

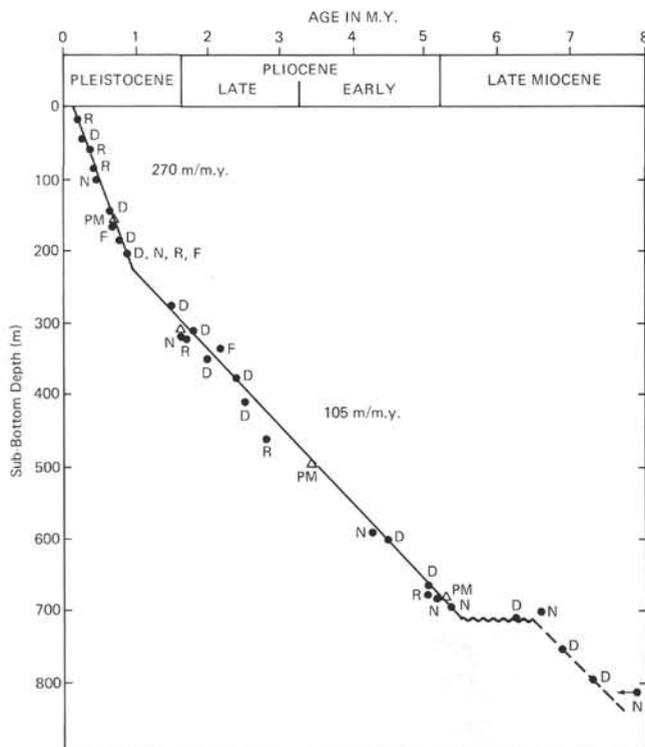


Figure 6. Sedimentation rates for Site 440, uncorrected for compaction, based on diatom (D), nannofossil (N), radiolarian (R), planktic foraminifer (F), and paleomagnetic (PM) data. Paleomagnetic data represent the top of the Olduvai Event, the Gauss/Gilbert boundary, and the Gilbert/Epoch 5 boundary (Smetzer and Hall, this volume).

also between Samples 440B-60,CC and 440B-61,CC. The first occurrences of *Nitzschia reinholdii* and *Denticula kamtschatica* s.str. are in Sample 440B-60,CC, and specimens of *Denticula* cf. *kamtschatica* below that level are primitive in morphology.

Assemblages from Samples 440B-61,CC through 440B-71,CC (base of hole) are correlated with Subzone *b* of the *Denticula hustedtii* Zone (Barron, this volume). The last occurrences of *Thalassionema hirosakiensis* in Sample 440B-66-3, 8–11 cm, and of *Distephanus pseudofibula* in Sample 440B-70,CC are supportive.

Preservation is good to moderate down to Sample 440B-38,CC; poor to moderate between Samples 440B-39,CC and 440B-60,CC; and poor to the base of the hole. Diatoms are common to abundant down to Sample 440B-38,CC, few to common from Samples 440B-39,CC to 440B-63,CC, and generally rare below that level.

Radiolaria

The expanded upper Miocene to Recent section of Site 440 contains common to rare radiolarian assemblages with good to poor preservation. Preservation as well as radiolarian abundance decreases with depth as noted herein:

Preservation

good — 1 (Section 440-1-1) to 409 (Section 440B-29-3) meters

good to moderate — 415 (Section 440B-30-1) to 516 (Section 440B-40-5) meters

moderate to poor — 521 (Section 440B-4-1) to 602 (Section 440B-49-5) meters

poor — 606 (Section 440B-50-1) to 814 (Section 440B-71,CC) meters

Abundance

common — 1 (Section 440-1-1) to 333 (Section 440B-21-3) meters

common to few — 336 (Section 440B-21-5) to 521 (Section 440B-41-1) meters

few — 530 (Section 440B-42-1) to 716 (Section 440B-68-1) meters

rare to few — 780 (Section 440B-68-3) to 814 (Section 440B-71,CC) meters

Radiolarian zones recognized from the cores taken at Site 440 are represented by the following intervals: 440-1-1 to 440-2-5 (*Botryostrobus aquilonaris*), 440A-3-1 to 440B-7-3 (*Axoprimum angelinum*), 440B-20-1 to 440B-35-1 (*Lamprocyrtis heteroporos*), 440B-35-3 to 440B-55-1 (*Sphaeropyle langii*), 440B-55-3 to 440B-63-1 (*Stichocorys peregrina*), and 440B-63-3 to 440B-71,CC (*Osmatartus penultimus*). The Pleistocene/Pliocene boundary occurs between Sections 440B-19-3 and 440B-20-1, and the Pliocene/Miocene boundary between Sections 440B-16-5 and 440B-17-1.

Calcareous Nannofossils

The three holes that make up Site 440 represent an apparently complete late Neogene stratigraphic sequence. Nannofossils are moderately to poorly preserved and abundances range from very rare to abundant. The generally low species diversities probably reflect extratropical environmental influences and selective dissolution. Sporadic occurrences throughout much of the Pliocene and a portion of the lower Pleistocene may indicate sedimentation within the calcite compensation range during these times. Because of these poor to barren intervals, a continuous zonation based on nannofossils was not possible, and zone and stage boundaries sometimes could not be delineated. However, nannofossil recovery was enhanced by selective sampling of light-colored calcareous mottles occurring occasionally in otherwise barren matrices. Approximately 25 per cent of these mottles, probably because of bioturbation, yielded adequate, sometimes abundant nannofossil assemblages.

All of Hole 449 (Cores 1-8) and Cores 1-3,CC of Hole 440A lie above the *Pseudoemiliana lacunosa* datum (*Emiliana huxleyi* Zone). Core 1 of Hole 440 contains common lower to middle Pleistocene nannofloras mixed with those of the late Pleistocene age. This is attributed to reworking or to mixing due to slumping. The interval between Core 4,CC in Hole 440A to Core 14,CC in Hole 440B encompasses the *Pseudoemiliana lacunosa* zone. The lower boundary of this zone is not delineated, because the samples between

Cores 8 and 14 (Hole 440B) are either barren or contain rare occurrences. The *Gephyrocapsid* assemblage in Core 7, CC is characterized by small individuals and may correspond to the upper portion of Gartner's (1977) small *Gephyrocapsa* Zone. Because of the (rare) occurrences of that taxa, we estimate that Core 15, CC (Hole 440B) is within the lower Pleistocene *Cyclococcolithina macintyreii* Zone. Samples between Cores 12, CC and 20-4 are barren or show rare, sporadic occurrences. The presence of *Gephyrocapsa caribbeanica* down to Core 19, CC suggests a Pleistocene age at least down to that level.

Cores 20 and 21 are assigned to the uppermost Pliocene *Discoaster brouweri* Zone; Cores 22 through 27 are barren or contain very rare, nondiagnostic nanofloras; and Cores 28 to 30 are within the *Discoaster surculus-D. tamalis* Zones. The top of the *Reticulofenestra pseudoumbilica* and *Sphenolithus abies* have apparent last occurrences. This may not represent the true top of the zone, as samples between Cores 35 and 42 were mostly barren. Ceratoliths, which are instrumental for lower Pliocene zonations, were too poorly represented to permit zonal assessments of Cores 47 through 57. This interval is characterized by rare to barren nanofossil occurrences. It was not possible to determine the Miocene/Pliocene boundary precisely, although the uppermost Miocene *Discoaster quinqueramus* zone top (5.6 m.y.) is well documented in Core 58-5. *D. berggrenii*, which characterizes the lower portion of the zone, is present throughout Cores 63 to 71. Hole 440B terminated near the lower boundary of the *D. quinqueramus* zone (about 7.0 m.y.), as indicated by the projected sediment accumulation rates.

No significant hiatuses could be determined from the biostratigraphic composite of Site 440 holes, and there were no repeated zones or anomalous thickening as a result of imbrication.

Foraminifera

Foraminifers at Site 440 are common to few in upper Pliocene to Pleistocene cores and absent throughout the lower Pliocene to upper Miocene, with the exception of rare occurrences in Hole 440B, Cores 54 to 64 and 68. Holes 440, 440A, and 440B contain very similar planktonic foraminiferal assemblages from the late Pleistocene. Similar to Site 438, the Plio/Pleistocene boundary falls within a dissolution interval between Cores 17, CC and 18, CC, Hole 440B. Well preserved late Pliocene faunas are present in Hole 440B, Cores 19-3 to 23, CC. Planktonic foraminifers are rare to absent in early Pliocene to late Miocene sediments because of dissolution. Samples 53-3 and 58-3 contain rare upper Miocene (N.17) forms of *Neogloboquadrina continua*, *N. pachyderma*, and *Globigerina praebulloides*.

Benthic foraminiferal faunas of Site 440 are similar to those of Site 438 but include a number of deeper-water species reflecting the deeper-water location of Site 440. Throughout this late Miocene to Pleistocene sequence the depositional environment appears to have been in lower bathyal depths (2000 meters+). A number of shallow-water species present in these sam-

ples indicate that downslope transport occurred throughout this sequence.

GEOCHEMISTRY

Shipboard geochemical studies at Site 440 consisted of analysis of interstitial water, characterization of hydrocarbon gases, and evaluation of organic richness as it relates to hydrocarbon potential. The methods of extraction and analysis were the same as those used at Sites 438 and 439.

The holes at Site 440 were drilled on the midslope terrace of the trench inner slope of the Japan Trench at a depth of 4515 meters. The strata range in age from late Miocene to Holocene and consist of 808 meters, chiefly claystone and diatomaceous claystone, in which the diatom content averages about 11 per cent. The rock is fractured and brecciated below about 350 meters. The rate of deposition in the upper 230 meters (Quaternary) is 230 m/m.y., and below this level to the bottom it is 100 m/m.y. Such rapid sedimentation at this site would minimize sea floor oxidation of organic matter and maximize the availability of the more easily decomposed fraction of the organic compounds. This labile fraction is the one most utilized during anaerobic processes.

A strong hydrogen sulfide odor was detected in the shallowest samples of this site, and a maximum H₂S analysis of 140 parts per million was measured in the gas from the first cores. In the depth interval from 26 to 160 meters, the deposits are stained black by an iron sulfide mineral, probably mackinawite, whose color disappears after exposure to air for several hours. The authigenic-pyrite content of the core ranges from 0.5 to 1 per cent.

Interstitial Water Analysis

Water samples at this site (Figure 7) do not show the strong decline in salinity and chloride-ion content that was measured at Sites 438 and 439 on the Japan Deep Sea Terrace. Site 440 seems to be too far seaward to have been influenced by the fresh-water artesian system inferred to have been associated with the Oyashio ancient landmass to which the anomalous previous analyses were ascribed.

Values of pH are higher at this site, and the alkalinity is exceptionally high (Figure 8). The alkalinity for the section as a whole at Site 440 is higher than that of any other site studied by the Deep Sea Drilling Project. A value of 91.8 milliequivalents per liter in diatomaceous claystone at a sub-bottom depth of 89 meters, which is nearly 100 times the commonly measured values, was exceeded only by a single previous sample, 92.9 meg/1 at Site 262 in the Timor Trough (Heirtzler, Veevers, and others, 1974), and the values there declined more rapidly above and below this record value than they do at Site 440. The very high alkalinity at Site 440 probably results from unusually favorable conditions for bacterial sulfate reduction and attendant bicarbonate production, as is also attested to by the very high levels of near-bottom H₂S and by the black staining of the strata. The high sedimentation rate, which guarantees burial protection of reactive organic constituents, is clearly an important factor in creating the favorable conditions.

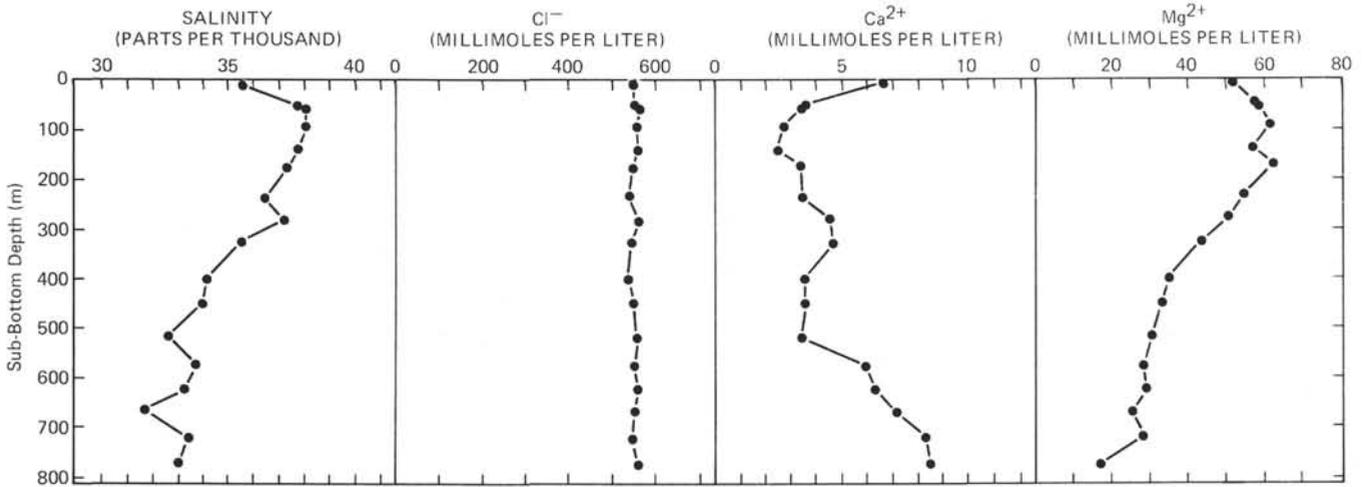


Figure 7. Relationship between sub-bottom depth and salinity, Cl^- , Ca^{2+} , and Mg^{2+} of interstitial water, Site 440.

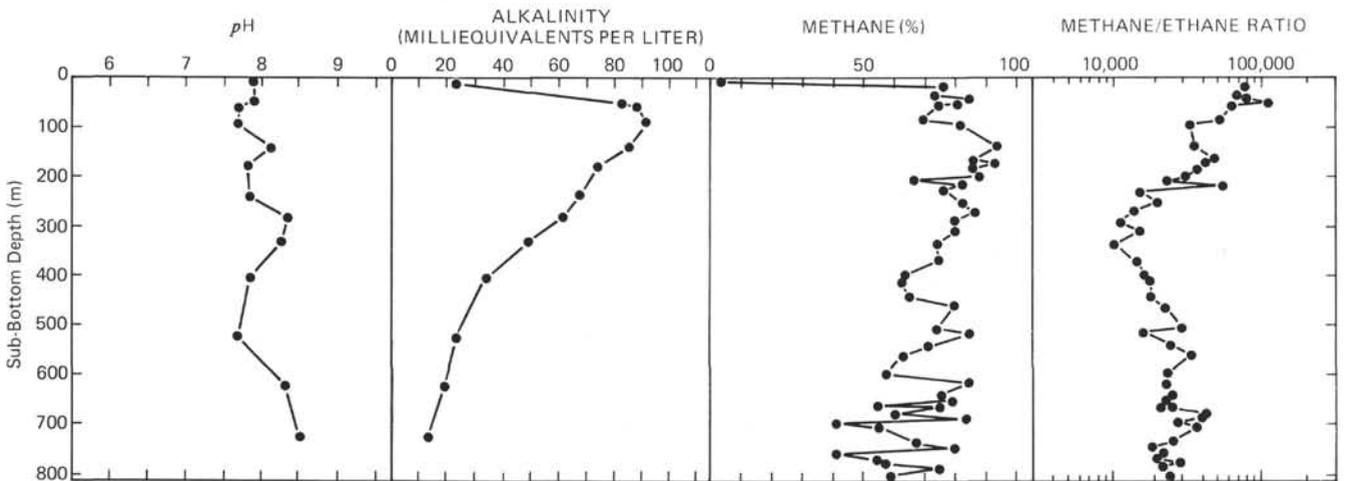


Figure 8. Relationship between sub-bottom depth, pH, and alkalinity of interstitial water, and methane and methane/ethane ratio of core gas, Site 440.

Organic Carbon and Hydrocarbon Gas Analysis

Results of organic-geochemical analyses at this site support the evidence for preservation of organic matter by rapid sedimentation, because the organic carbon content of the sediment is relatively high. But, despite the high organic carbon content, low pyrolysis-fluorescence values were obtained, and the intercept of the plot (Figure 9) indicates that about 0.5 per cent of the carbonaceous matter is free of hydrocarbons. Based on 0.28 to 1.17 per cent organic carbon and 0.2 to 13 per cent pyrolysis-fluorescence, the sediment at this site would be considered low in hydrocarbon richness with respect to oil generation. Moreover, extrapolation from the temperature log measurement of $9^{\circ}C$ at 498 meters, 26 hours after circulation of the drilling fluid ceased, suggests that the equilibrium temperature at that depth is about $9^{\circ}C$. Combined with a sea floor temperature of $1^{\circ}C$, this gives a gradient of $1.6^{\circ}C$ per 100 meters and a temperature of $14^{\circ}C$ at 808 meters, indicating that the

cores at this site have not been exposed to temperatures required for petroleum generation.

At about 20 meters depth the methane content reaches 77 per cent, and it is relatively constant through the remaining cores below. As indicated by changes in the methane/ethane ratio (Figure 8), the ethane content first increases with depth, as is normal (Claypool and others, 1976), but then decreases below a depth of about 350 meters. This is approximately the depth at which faults and fractures become common in the cores and also the depth at which a $1^{\circ}C$ step occurs in the high-resolution temperature log. We believe that the older accumulations of diagenetic gas in this zone have partly escaped along fractures and that subsequently produced biogenic methane has not been subjected to diagenetic processes long enough to restore the more usual amounts of ethane.

A characteristic feature of this site is an unusually high content of propane, butane, and pentane, especially in the depth range of 20 to 140 meters (Sato and

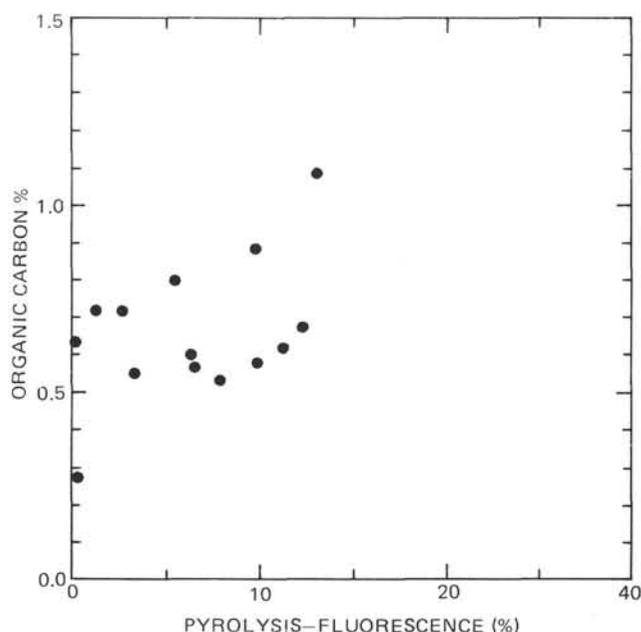


Figure 9. Relationship between organic-carbon content and pyrolysis-fluorescence of core samples from Site 440.

Whelan, this volume). Maximum volumetric values for these gases in the methane are measured at 94 meters, where propane, butane, and pentane, respectively, are 3.89, 4.56, and 1.68 parts per million. It is especially noteworthy that the isobutane/normal butane and isopentane/normal pentane ratios are high and that neopentane was not detected at these levels. This suggests that gas may have migrated into this zone from below or from another structural block, a possibility that is also suggested by the fact that the ethane gradient in the upper part of the hole is higher than that of Site 438, which has a higher geothermal gradient. Neopentane does first appear at a depth of 225 meters. The neopentane is not believed to be related to thermochemical processes of petroleum generation but to originate through normal diagenetic reactions.

PHYSICAL PROPERTIES

As is evident from the seismic reflection records, the sediments sampled at Site 440, predominantly diatomaceous claystones and claystones, have accumulated continuously with a nearly horizontal attitude. The structural configuration and uniform sedimentation rates suggest orderly accumulation and burial of these deposits until at least the Pleistocene. The faulting, veining, and brecciation observed in the lower half of the column, however, imply that a regional tectonic stress may have been superimposed on normal compaction and consolidation (Arthur, Carson, and von Huene, this volume). Previous studies (Lee, et al., 1973; Bouma and Moore, 1975; Carson, 1977) suggest that such a tectonic overprint may be reflected in the physical properties of the sediments.

Wet bulk density, porosity, water content, shear strength, thermal conductivity, and sonic velocity were measured on sediments recovered at Site 440. Hole 440B

was logged for sonic velocity and formation (bulk) density. The data are plotted on Site Summary Chart 435/440 (back pocket) and discussed more fully in Carson and Bruns (this volume).

Sonic Velocity

Velocities range from about 1.5 km/s at the surface to 2.0 km/s at 800 meters. Near-surface (< 250 meters) sonic velocities are approximately the same (~1.60 km/s) as those at Site 438 (although control is very poor above 180 meters owing to signal attenuation in gas-rich sediment). At depths greater than 250 meters, however (and particularly below 400 meters), velocities are significantly higher at Site 440 than at Site 438.

Both the *in situ* sonic log and the laboratory determinations indicate gradually increasing velocities from the surface to a depth of 390 to 400 meters (Lithologic Units 1 and 2). At this point there is a marked increase in velocity from 1.65 km/s to 1.78 km/s. Below 400 meters (Lithologic Unit 3) velocities decrease slightly to minimum values (~1.75 km/s, *in situ*; 1.53–1.70 km/s, laboratory) at 485 meters and then increase regularly to the base of the hole.

The 390- to 400-meter depth marks the top of the brecciated zone at Site 440 as well as a section (~380–420 m) of apparent sediment influx by mass movement (possibly as slump deposits). The velocity increase at 400 meters may reflect either a local lithologic (textural?) variation associated with the slump unit or a general change in fabric related to brecciation, or both. The fact that relatively high velocities (>1.70 km/s) as well as brecciation persist to the base of the hole, however, implies a general change in the consolidation state below 400 meters. It may be that this depth marks a front of consolidation or induration above which stress is accommodated by plastic deformation and below by brittle fracture.

Bulk Density, Water Content, and Porosity

The bulk density profile consists of three distinct sections: 0 to 410 meters; 410 to 620 meters; and 620 meters to the bottom of the hole. The uppermost section (encompassing Lithologic Units 1 and 2) is characterized by densities which range from ~1.55 Mg/m³ (surface) to ~1.70 Mg/m³ (410 m), with a high degree of small-scale variability. There is a density high, defined by a lack of low values, rather than exceptionally high densities (~1.75 Mg/m³; 90–140 m, laboratory determinations, Hole 440A; 113–140 m, *in situ*, Hole 440B) within this section, which reflects the occurrence of a silty, diatomaceous clay unit. Thinner intervals of higher density within this section correlate primarily with carbonate cementation (e.g., 220–227 m; 257–265 m). Between 260 meters and 400 meters, the *in situ* log defines a regular decrease in bulk density (from 1.7 Mg/m³ to 1.58 Mg/m³, which is not recorded in laboratory — GRAPE and gravimetric — determinations. The cause of this discrepancy is unknown, although it may indicate the presence of fracturing on a scale (>1 m?) which affects the log but is too large to be resolved by the laboratory methods.

The second section, 410 to 620 meters (upper Lithologic Unit 3), encompasses the upper half of the observed brecciated zone. The density within this section is nearly uniform from top to bottom and averages 1.55 Mg/m³. It forms a low-density zone beneath the upper section. Both the *in situ* log and laboratory determinations show a low degree of variability.

Below 620 meters (lower Lithologic Unit 3), the density increases sharply and is apparently lithologically as well as depth-controlled. Two density maxima (648 m, 1.80 Mg/m³, and 704 m, 1.75 Mg/m³) reflect more vitric and/or calcareous deposits. The upper maximum coincides with a slump deposit. The lower maximum may indicate partial carbonate cementation.

Unlike Sites 438 and 439, Site 440 shows no consistent relationship between density and sonic velocity sections. In particular, the section above 390 meters contains density maxima which are not reflected in the velocity profile (see, for example, the section from 113 m to 138 m). The borehole diameter, lithology, and the bulk density do not appear to contribute to the sonic maxima. Furthermore, the density minimum from 410 meters to 620 meters, in the zone of brittle fracture, is characterized by constant but relatively high (~1.77 km/s) velocity. It is inferred that the extensive fracturing in this zone contributes to these divergent properties.

The sediment porosity at Site 440 shows a profile which mirrors that defined by bulk density. In the upper 400 meters (Lithologic Units 1 and 2) porosities decrease fairly regularly, from ~67 per cent at the surface to ~63 per cent at the top of the brecciated zone. From 400 meters to 590 meters the porosity remains relatively constant at about 60 per cent, suggesting that fractures and veins maintain a constant (apparent) porosity whereas true intergranular porosity, within breccia clasts, may continue to decrease with depth (under increasing load). Below 590 meters, porosities again decline with increasing depth to about 54 per cent at the bottom of the hole. This zone, like the overlying section, is brecciated. The porosity and density data, however, suggest that either the fractures are progressively annealed with increasing depth in this lowermost section or lithologic variations exert a dominant control on void volume reduction.

Water content decreases from the surface (~47 per cent) to the bottom of the site (~30 per cent) in a fashion analogous to the decrease in porosity.

Shear Strength

The shear strength of the sediment cored at Site 440 increases rapidly and nonlinearly with depth, from average surface values of ~15 kPa to 50 kPa at 130 meters. An enveloping curve (which covers all values) is defined by a surface value of ~24 kPa and increases to 44 kPa at 50 meters and to >85 kPa at 100 meters. These latter values are considerably greater than those given for typical hemipelagic-terriginous sediments by Bouma and Moore (1975). It is unclear at this time whether the shear strengths at Site 440 are anomalously high and reflect tectonic loading of the near-surface

sediments or are normal and simply expand the enveloping curve defined by Bouma and Moore (1975) for diatomaceous clays.

GEOPHYSICS

Seismic Reflection

Site 440 was positioned along JNOC multichannel record section 2, between Sites 434 and 435 on the trench slope (Figure 1) drilled previously during Leg 56. Site 440 was selected for drilling to define the landward edge of the accretionary prism. This boundary had been presumed to occur at the midslope terrace by Honza and others (1977). At Site 435, on the upper inner slope of the Japan Trench, the cores were interpreted as slumped terrigenous sediment. The cores from Site 434, on the lower trench inner slope, were interpreted during Leg 56 as an imbricated section, repeated along at least two faults. This interpretation strengthened the argument of Honza and others (1977), because the boundary was inferred to lie between these sites. The midslope terrace is the most persistent topographic feature of the slope and it might therefore correspond to the continental/oceanic boundary.

Drilling at most sites, particularly at Site 434 (see Langseth et al., this volume) had terminated because of hole-caving and stuck pipe. The highly fractured claystone encountered below 300 meters was generally difficult to hold open. Because reflections at the midslope terrace appeared to be less faulted than at other places along the trench inner slope, this site looked attractive.

The relatively flat topography of the midslope terrace is a seismic window in which the reflected seismic signal is not scattered or defracted by surface roughness on the sea floor (Figure 10). Therefore the subsurface reflections in this area are stronger than elsewhere along the trench inner slope. The terrace is bounded on either side by pronounced zones of defractions in the seismic record which obscure the structural relations of the terrace to the upper and to the lower slopes.

The *Challenger* single-channel record shows possible ponding of the sediment on the midslope terrace to 0.15 sec (110 m) at Site 440. Below this level no structural indication of ponding is recorded. Upslope from the pond, a short sequence of landward dipping reflections was recorded on the single channel record which are obscured in the multichannel record.

The midslope terrace reflections are of low amplitude and have a fragile continuity. This is particularly true of the reflections above 350 meters. Nevertheless, they outline a sequence of subparallel horizons that dip slightly landward. The faint reflective character is consistent with the uniform lithology of the section recovered here, and the upper 350 meters may correspond to the less lithified part of the section. The base of the sequence is a diffracted reflection of lower frequency and higher amplitude than the overlying sequence. The basal reflection may continue west — it is implied to be continuous in Figure 10 — but its continuity is somewhat uncertain. Parts of this reflection may originate from beds to the side of the plane of the section.

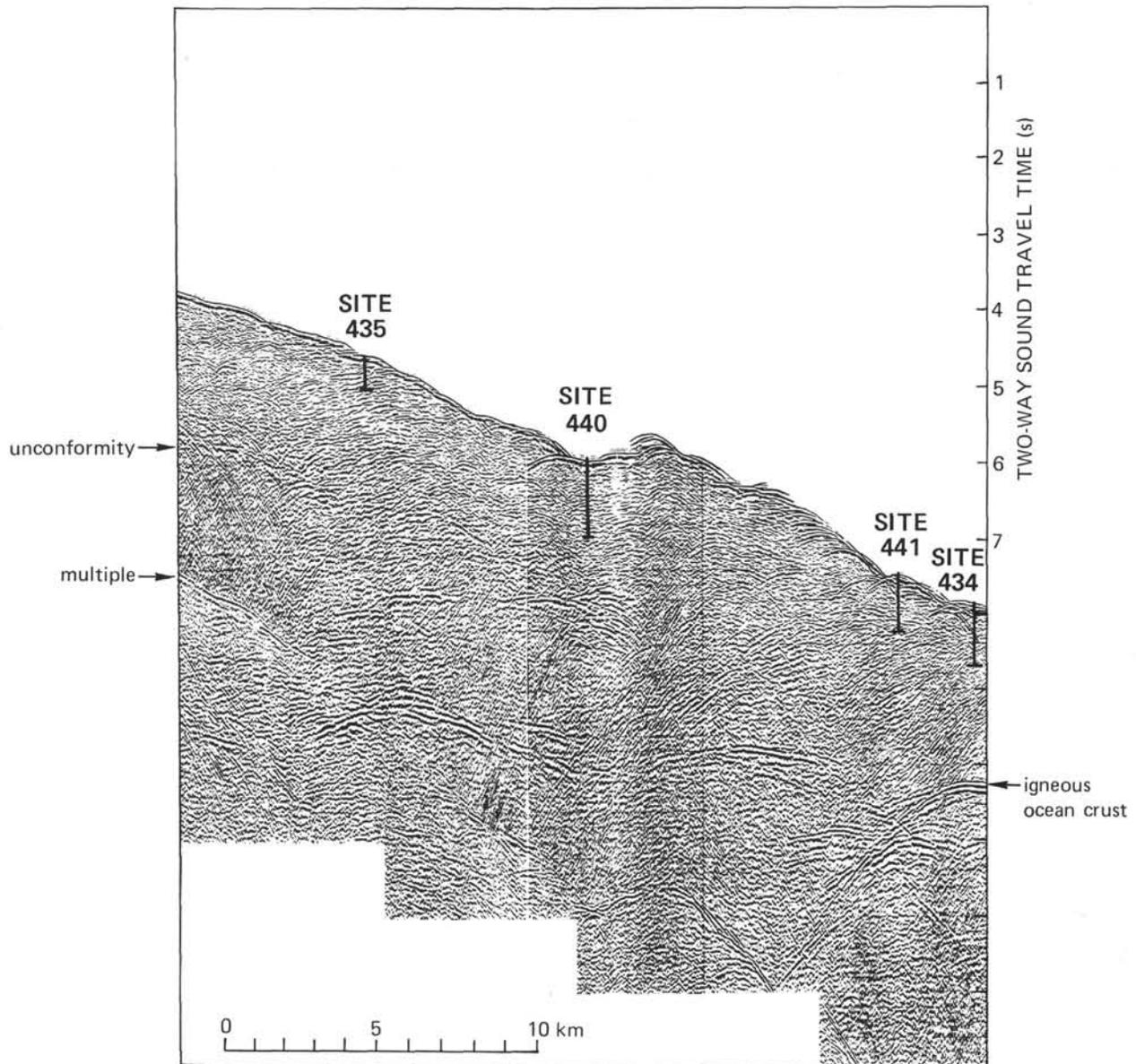


Figure 10. Section of JNOC-2 multichannel seismic reflection record showing mid-slope terrace and reflective sequence penetrated at Site 440.

According to a projection of the velocity gradient determined by downhole logging, the estimated depth of the reflective section is about 1600 meters. The estimated age of the base, assuming no hiatus, is 16 m.y. if the rate of sedimentation in the late Miocene is projected to this depth.

The interpretation (Figure 11) of how the reflective sequence in the mid-slope seismic window relates to other parts of the continental slope is not clearly indicated in the seismic records. Because the mid-slope terrace is part of a convergent margin, compressional structures would be expected. Most faults within the mid-slope terrace reflective sequence are steep seaward-dipping apparent reverse faults that have about 10 to 30 meters of vertical displacement. Using the same criterion, there are also some apparent steep normal faults. But the landward-dipping thrust faults indicated in this position in many

convergent margin models are absent. A tectonic mechanism, which is not clear from the seismic data, is needed to bring the section found at Site 435 more than 1000 meters down to the elevation of Site 440. Because folding appears unlikely, faults or large slumps, including the rotation of bedding attitude, are inferred in the areas obscured by diffractions.

Logging

Except for the laterolog, Site 440 was logged with the same set of tools used at previous sites. Conditions were less than desirable because the large swells that caused the ship to heave superimposed this motion in turn as an oscillating motion during the logging run. The ship had held position through a storm with gale winds, and both the logging and the retrieval of the drill stem had to be accomplished in the brief period prior to another rapid-

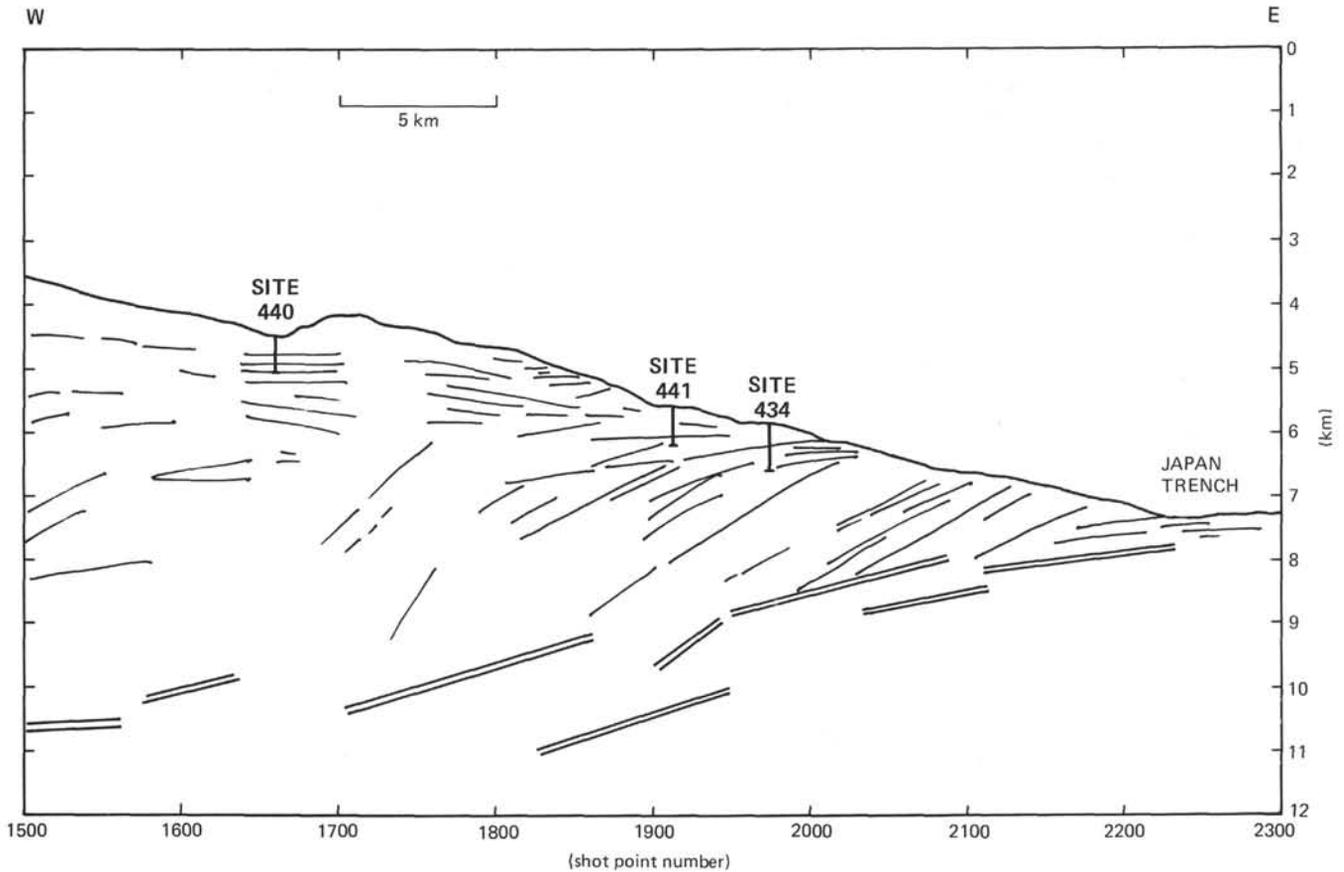


Figure 11. Interpretive depth-section of JNOC-2 multichannel seismic reflection profile normal to trench inner slope. Vertical exaggeration 2:1. Double lines show reflection from oceanic crust.

ly approaching storm. During logging the ship lay in the troughs, because it had to head into a current stronger than either the wind or the sea. This precarious situation did not allow sufficient time for full logging.

Temperature. The high resolution temperature (HRT) tool was run last, 24¾ hours after circulation of mud in the hole. Table 2 gives measurements of the three maximum reading thermometers included in logging runs, and it shows that the bottom hole temperatures were less than those of surface sea water. Unfortunately the hole had bridged just before the HRT run, so that only the upper 500 meters were open. Readings were taken on the way down, and jumps in the trace suggest that because of hole deviation the tool was probably riding against the side of the hole in the lower area. However, we cannot explain a sudden rise in temperature at 383

meters. Two stationary readings, five minutes apart, were made at the bottom.

The uncorrected temperature gradient, about 1°C/100 meters, is one-third of the previously run gradients, and the bottom hole temperature is at least 8°C less than the reading at an equivalent depth in Hole 439. The low temperature gradient confirms the low heat flow reported at accreting margins from sea floor heat probe measurements.

Gamma Ray Log. Two runs were made with the gamma-ray tool. Comparison of the two traces shows many small differences which may reflect the effects of motion from the heave. Features with wavelengths greater than about 10 meters are reproducible. Areas washed out show increased gamma-ray values.

Caliper. This tool showed that hole conditions were good. Local "washouts" correspond well to lithologic boundaries.

Formation Density Log (FDC-CNP). The trace has much character; however, some of it may be caused by motion from heave. Because the low values above 325 meters may be incorrect, the high values should be favored. The heave frequency is very clearly superimposed on the trace in some places.

Sonic Log (BHC). This log is particularly noisy from 250 meters to 325 meters. The general velocity gradient is comparable to the log from Site 439 except that the

TABLE 2
HRT Temperature Data, Hole 440B

Run	Depth below Rig Floor (m)	Total Sub-bottom Depth (m)	Temperature (F°)	Time since Mud Circulation (hr.)
1 1st BHC	5280	763	< 60 (15.5C)	6¾
2 FDC-CNP	5275	758	< 60	11¾
3 2nd BHC	5250	733	< 60	17¾
4 HRT	5015	498	< 60	24¾

velocity increases below 400 meters to values 100 to 140 m/s greater.

Taken together, the logs show changes corresponding to lithologic changes, noted visually, particularly with regard to the zone of mass movement. Velocity and density jumps corresponding to folded areas are the most obvious large features of the log traces. Washouts show the locations of contacts and provide an explanation for their absence in the cores. In a gross sense the logs show a distinct upper section to 400 meters, a uniform section from 400 meters to 600 meters, and a lower section to the total depth (see Physical Properties). Numerous smaller features in lithology that were not observed visually because of lack of core recovery may nevertheless be reflected in the logs.

SUMMARY AND CONCLUSIONS

Summary

Site 440 is on the midslope terrace of the Japan Trench inner slope, a persistent tectonic structure that extends parallel with the trench for 360 km with some intermittent short breaks. In the area of Site 440, it is about 5 km wide in water 4500 meters deep. The terrace is the most notable topographic feature of the trench inner slope, as is typical of many other trench slopes. For that reason, it was selected for sampling by the JOIDES Active Margins Panel and the Japan IPOD Committee.

During the Japan Trench transect, the first two sites of Leg 56 were on the lower and upper trench inner slope. From the shipboard study of samples during that leg, the lower-slope site was inferred to be on either a thrust-faulted or a slumped terrigenous slope deposit and the upper-slope site on a large block of downslope terrigenous sediment. Thus the imbricate wedge and the apron of downslope terrigenous sediment must have a common boundary between the sites, and this boundary may be the midslope terrace. In addition, because the failure of holes due to caving of fractured rock had occasionally stopped drilling during both Legs 56 and 57, the absence on seismic records of large folds or faults below the terrace increased the attractiveness of the deep drilling originally planned there.

The predrilling site survey and positioning of the beacon at Site 440 were made during a storm. Intermittent stormy weather persisted during much of the 10 days on site, twice causing abandonment of holes. Because yet another storm was approaching and because we felt that many of the original objectives at Site 440 had been accomplished, it was abandoned at 814 meters, or at about 60 per cent of the targeted penetration, despite a good hole and bit. Logging and retrieval of the drill string were accomplished in a short break between storms. Logging was possible only because of the skill and persistent efforts of the Global Marine crew and the Cruise Operations Manager in holding the ship's position through the last storm.

The relatively flat sea floor of the midslope terrace affords a seismic window that reveals a 1.6-s sequence of straight reflections dipping gently landward above a defracting acoustic basement. This section ends on

either side in overlapping defractions that obscure indications of the structural relation between the midslope terrace section and the strata upslope and downslope. The uniform sequence of reflections are of low amplitude and fragile continuity, particularly above a 350-meter depth, consistent with the uniform lithology cored and the results of the downhole logging. A single-channel record made during the predrilling site survey indicates ponded sediment in the upper 110 meters of the section near the drill site. Projection of velocities below the bottom of the downhole acoustic log indicates a stratified section at least 1600 meters thick.

Downhole logging was carried out under the adverse condition of strong ship's heave, but despite degradation from the heave, the logs record a much greater variability in density and velocity than is apparent in visual observation of the lithology. The velocity gradient at Site 440 is the same as at 438 and 439 in the upper 400 meters but is 100 to 140 m/s more at greater depth. This may be caused partly by a sparser diatom content and greater compaction. Probably the velocity increase and peculiar intervals of corresponding density decrease have a diagenetic or tectonic origin that may involve overpressuring of interstitial fluids. The caliper log showed pronounced washouts at lithologic contacts that were not recovered. The gamma-ray trace reflected the uniformity of the sediment composition. A significant result came from the temperature log, which indicates bottom hole temperatures 8°C below those at equivalent depths of the outer two sites and a gradient of only 1°C/100 meters.

Physical properties measurements in conjunction with downhole logs indicate a normal increase of density, porosity, and water content to 400 meters; a constancy of these properties for the next 200 meters; and then an increase below 600 meters. This is thought to be related to the microfracturing that begins at about 400 meters and continues to the total depth. Apparently the section dewateres normally above that depth and then develops water-filled fractures that may not dewater until 600 meters, where they are gradually forced closed by the lithostatic pressure.

Biostratigraphic study gave much the same zonation as at Sites 438 and 439, documenting a continuous sequence of Quaternary through upper Miocene sediment to rocks with an age somewhere between 6 and 7 m.y. at the deepest penetration. Diatoms are again the most abundant microfossil group; radiolarians range from common to rare; foraminifers are common above the upper Pliocene; and nannofossils are more abundant than at the other sites. Sediment accumulation occurred at a very constant rate in the Pliocene and is comparable to Sites 438, 439, and 435 without a correction for water content. The Pleistocene section, however is 200 to 250 meters thicker than at any other site in the Japan Trench transect. This is probably because of ponding of sediment on the terrace. Displaced shelf foraminiferal assemblages in the upper 70 meters are consistent with the ponded structure noted in seismic records and suggest that partial closure first developed here in the upper Pleistocene.

Foraminiferal assemblages indicate bathyal depth and deposition below the local CCD. They also reflect the same climatic changes observed at other North Pacific sites. The cyclic recurrence of the diatom *Denticula seminæ fossilis* in the Pliocene might also reflect climatic cycles.

The lithologies cored at Site 440 are similar to those at most of the other sites on the Japan Trench transect. Below thin graded beds interspersed with clayey diatom ooze is claystone and diatomaceous claystone which in its lower part is marked by intervals folded during local mass movement. The uppermost unit, 38 meters thick has poorly sorted graded beds of clay, sand, and gravel between layers of diatomaceous clay and ooze. Ice-rafted dropstones are common. The second unit (38–380 m) is composed in the upper half of diatomaceous claystone and in the lower half of claystone. Mottling is slight to intense; lithification begins at 168 meters and increases downward. The third unit (380–808 m) is also diatomaceous claystone and claystone, but it is distinguished by several zones with syndepositional folds that probably resulted from local mass movement. Mottling is commonly intense.

The zones of mass movement can be matched to jumps in the downhole log traces, which can in turn be matched to reflections. The downhole logs indicate that mass movement occurred in zones 5 to 10 meters thick. The amplitudes of reflections change laterally, consistent with the interpretation that the mass movement is a local intraformational feature.

Ash and tuff layers afford a record of volcanism much like that at Site 438. Volcanic ash begins to increase in the late Miocene and remains high during the first two-thirds of the Pliocene. After a waning period, the ash layers again become numerous in the late Pleistocene.

A primary objective at Site 440 was to examine structural geology in the seaward portion of an accreting margin. The general impression obtained from examination of the sediment is that this sequence has been stressed more than the section at Site 438, which is on the deep sea terrace 50 km farther landward. From shipboard observations, the first major change in the cored section is the appearance of brittle deformation in the form of fractures or joints at a sub-bottom depth of 386 meters (Core 27). Below this point, many healed microfractures, joints, and veins are observed; microfaulting is common, and fault planes of all inclinations have both normal and reversed senses of displacement. A large part of the section below 400 meters exhibits tectonic brecciation. This brecciation occurred in place after some degree of lithification of the original sediment. Rehealing occurred after small relative displacements between the pieces of the breccia. The brecciated appearance seems due to the rather random, anastomosing, and branching nature of veins and fractures. However, not all fracturing is interpreted to have resulted solely from tectonic stress. Some of it could have been induced along anastomosing veins which originally acted as dewatering conduits; these could then have propagated preferentially in a direction favorable to the

ambient stress field and become planes of strain release that evolved into a series of closely spaced joints and microfaults. In turn, some dewatering must have continued along existing joints and fractures. The orientation of fractures seems generally random, and although a preferential steep trend would be expected, this cannot be established from the core without additional information on core orientation. The overall impression is that at the scale of a DSDP core, the section was subjected to mild stress fields of various orientations or to tectonic jostling without a prolonged single strong orientation of any one stress field. Possibly this is one of an initial stage of tectonic consolidation at a converging margin. However, the consolidation is much less intense and systematic than one might expect, just 28 km from the trench axis.

Conclusions

The stratified sequence beneath the mid-slope terrace at Site 440 consists of a nonfolded upper continental slope block rather than of material accreted at the leading edge of the continental plate. The sediment section is uniform hemipelagic mudstone containing few turbidites except in the upper Pleistocene. The basic sediment composition is similar to that at Sites 435, 438, and 439 farther upslope. A uniform hemipelagic lithology is defined well by cores, by downhole logs, and by seismic reflection records; its character is consistent with the very uniform rate of sedimentation from the late Miocene through the early Pleistocene. The block is more than a superficial slump, because it is at least 1600 meters thick, and if the rate of sedimentation is assumed to have continued without hiatus from an earlier period through the late Miocene, as it did at Sites 438 and 439, the estimated age of the sediment at the acoustic basement is early Miocene. Thus it can be inferred that the small sediment wedge seaward of the mid-slope terrace represents all the accretion that has occurred here since the late Neogene.

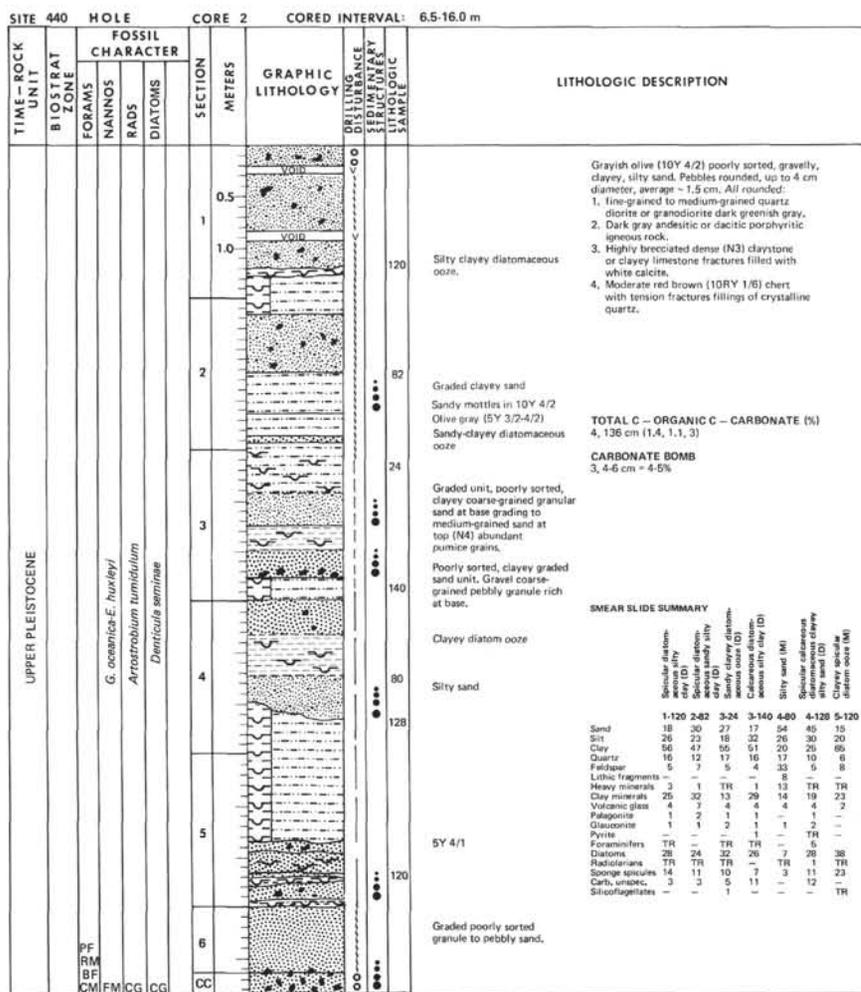
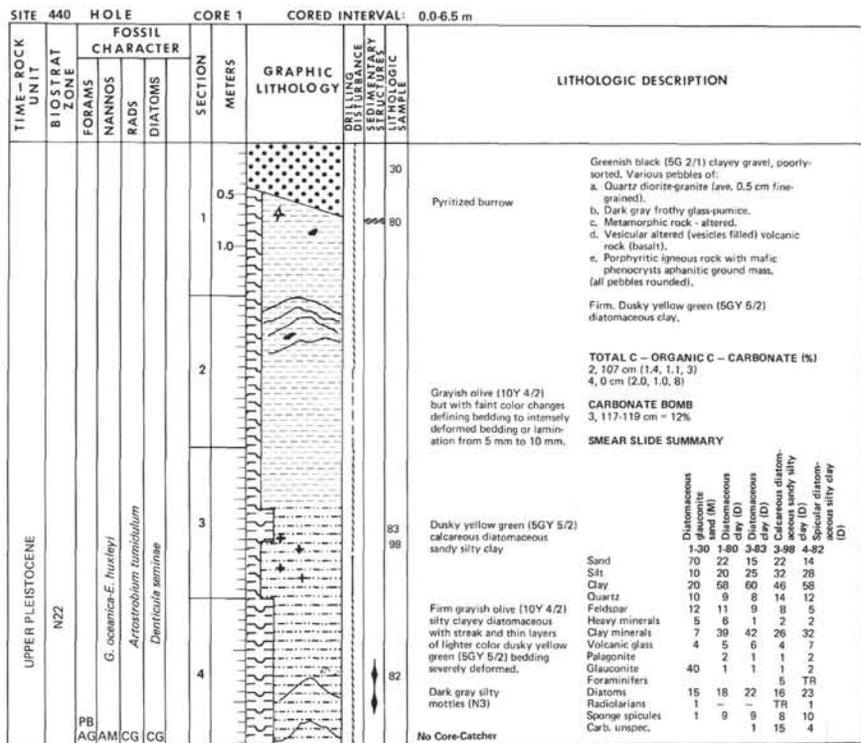
The only breaks in continuous hemipelagic deposition are three or more periods of local slumping in the Miocene and Pliocene. In the late Pleistocene, ponding of turbidites, consistent with the present topography of the mid-slope terrace, is recorded by the litho- and biostratigraphic histories. The sampled section has been at bathyal depths (greater than 2000 m) and near or below the CCD since at least the late Miocene. No tectonic event that would change sediment source or dispersal paths is obvious from the history recorded here. However, a stress environment of unknown age has tectonically brecciated the section at depths where the rock deforms brittlely. Although microfracturing was observed at other sites on the transect, it was not as intense and pervasive at the relatively shallow depths at which microbrecciation was observed here except at Sites 434 and 441. Fault and joint planes generally have a random orientation, and systematic study would be required to see if there was a dominant or preferred stress direction.

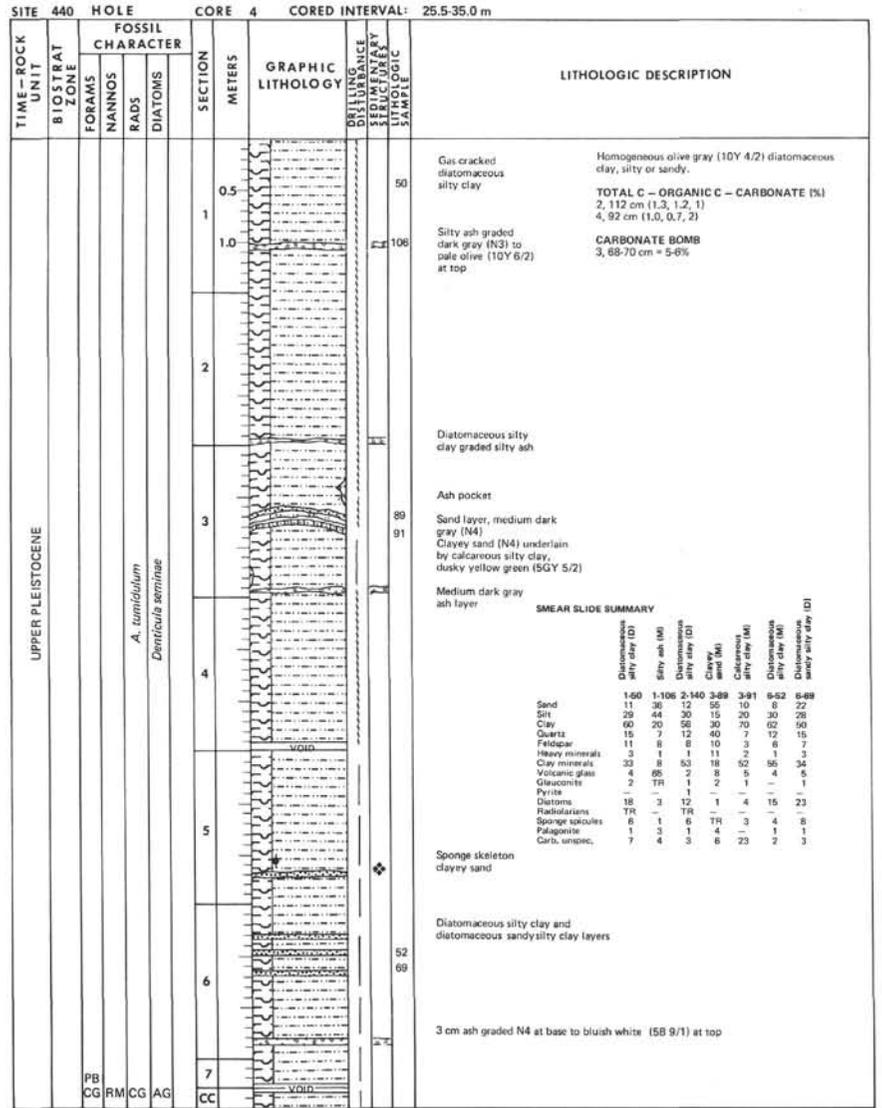
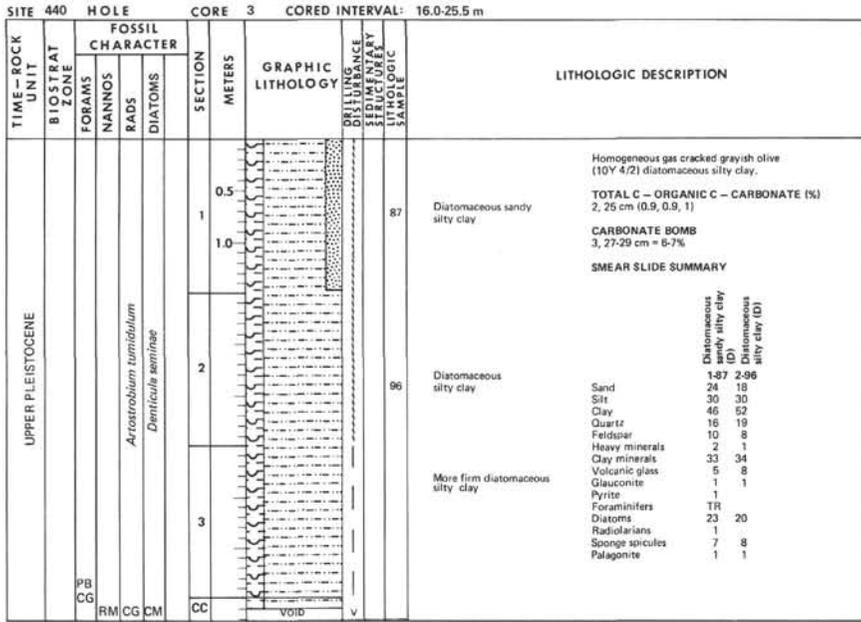
The tectonic history during which a continental slope block has remained coherent so far down the trench in-

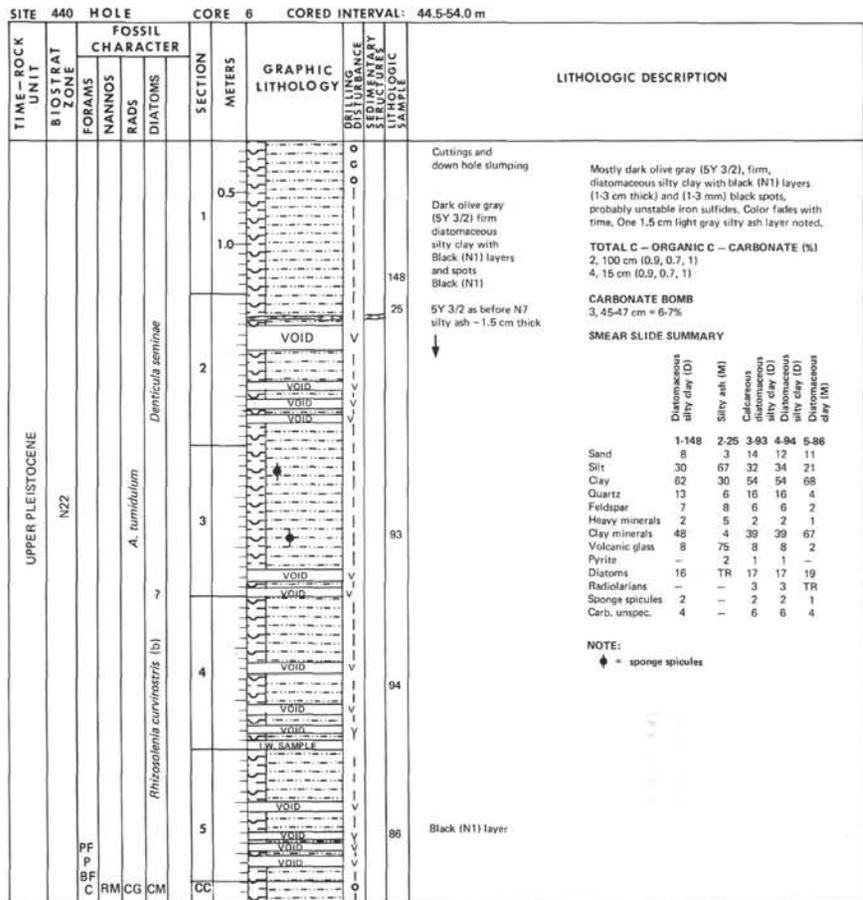
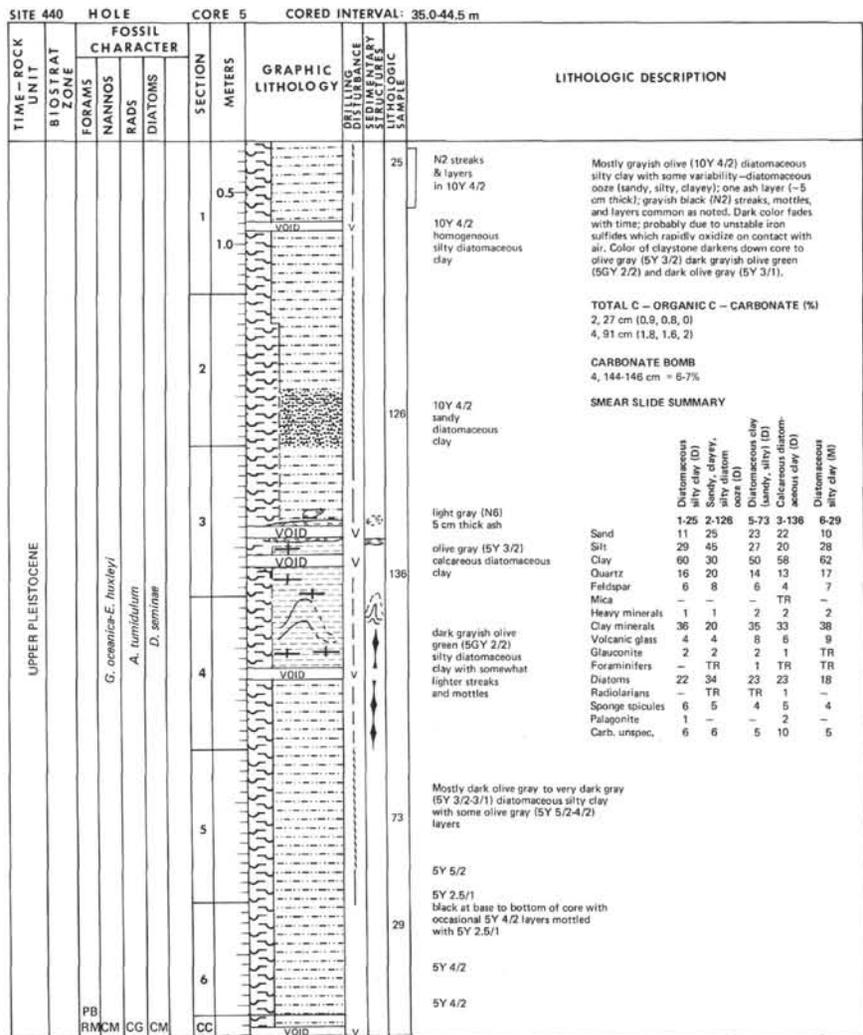
ner slope of a convergent margin is not fully apparent from the seismic records or the drill cores. Multichannel seismic records show that the deep sea terrace ends at a broad arch. The seaward flank of this arch is the trench upper slope above the midslope terrace. The uniformly thick section on the deep sea terrace continues over the arch, begins to descend the slope, and then is cut into four major segments. Between each segment is a zone of defractions that might mark a fault zone. Two faults have sea floor expression, indicating continuing activity. Within blocks, most of the clear reflections have been tilted landward from their original seaward dip. This pattern is suggestive of large back-rotated slump blocks, which seem very anomalous on a trench slope, or of steep imbricate faults. The slump interpretation rather than accreted sediment and brecciation rather than imbricate slices and isoclinal folding are consistent with the occurrence of continental slope deposits far down the trench wall. This seems the simplest explanation of the present data, although it is not the usual one in convergent margin models. It appears that the intense convergent stress at this particular margin is communicated for only a short distance from the trench in the upper plate.

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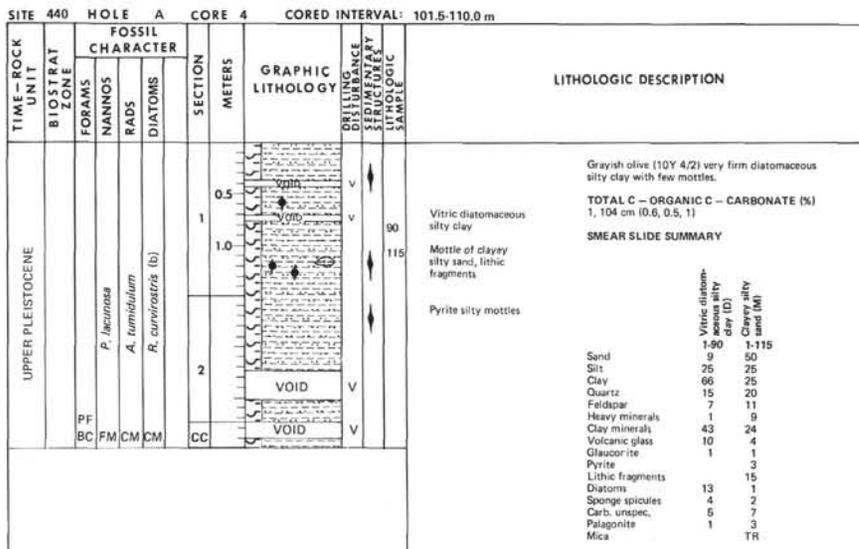
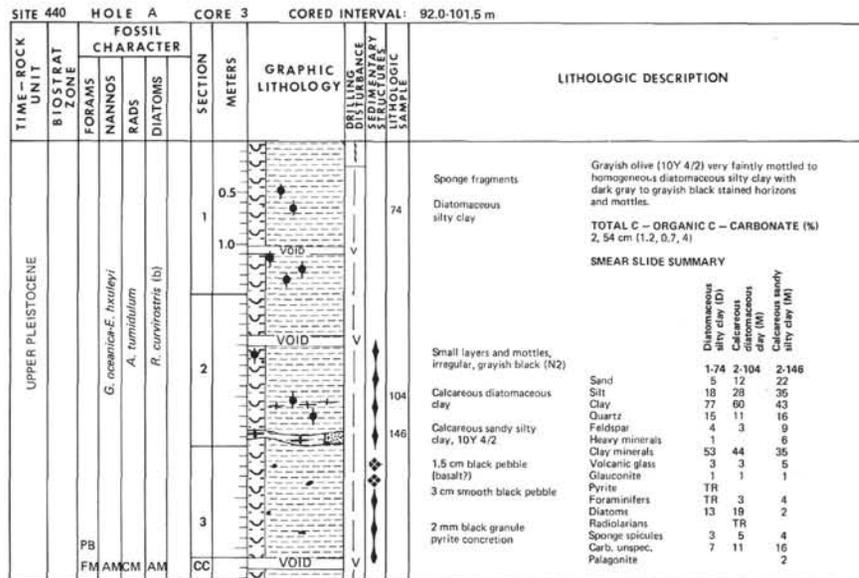
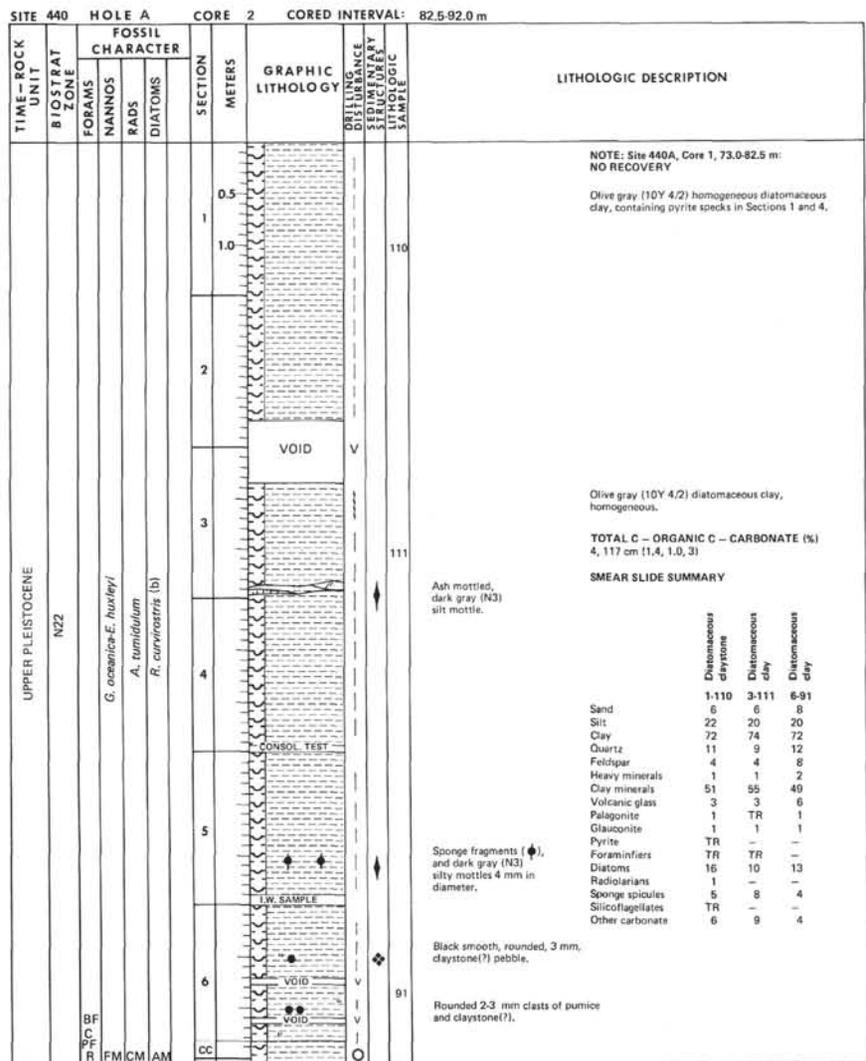
SMEAR SLIDE SUMMARY

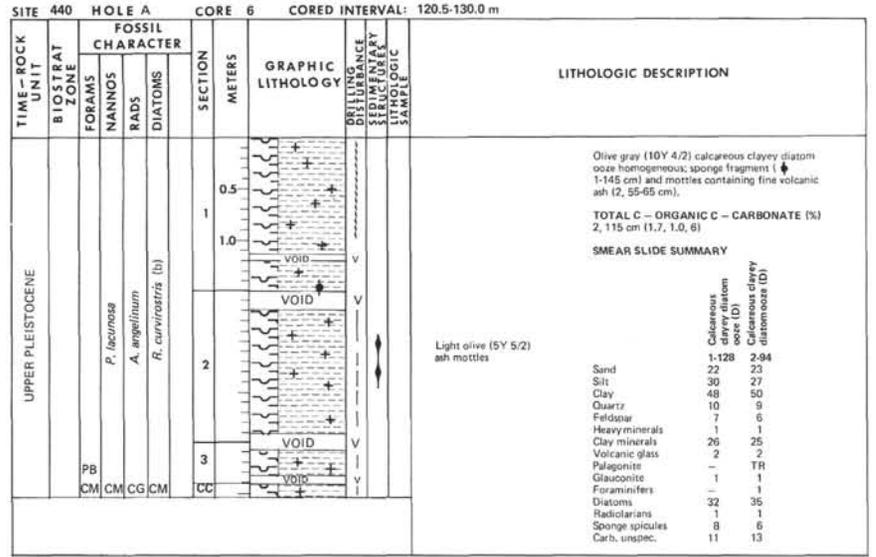
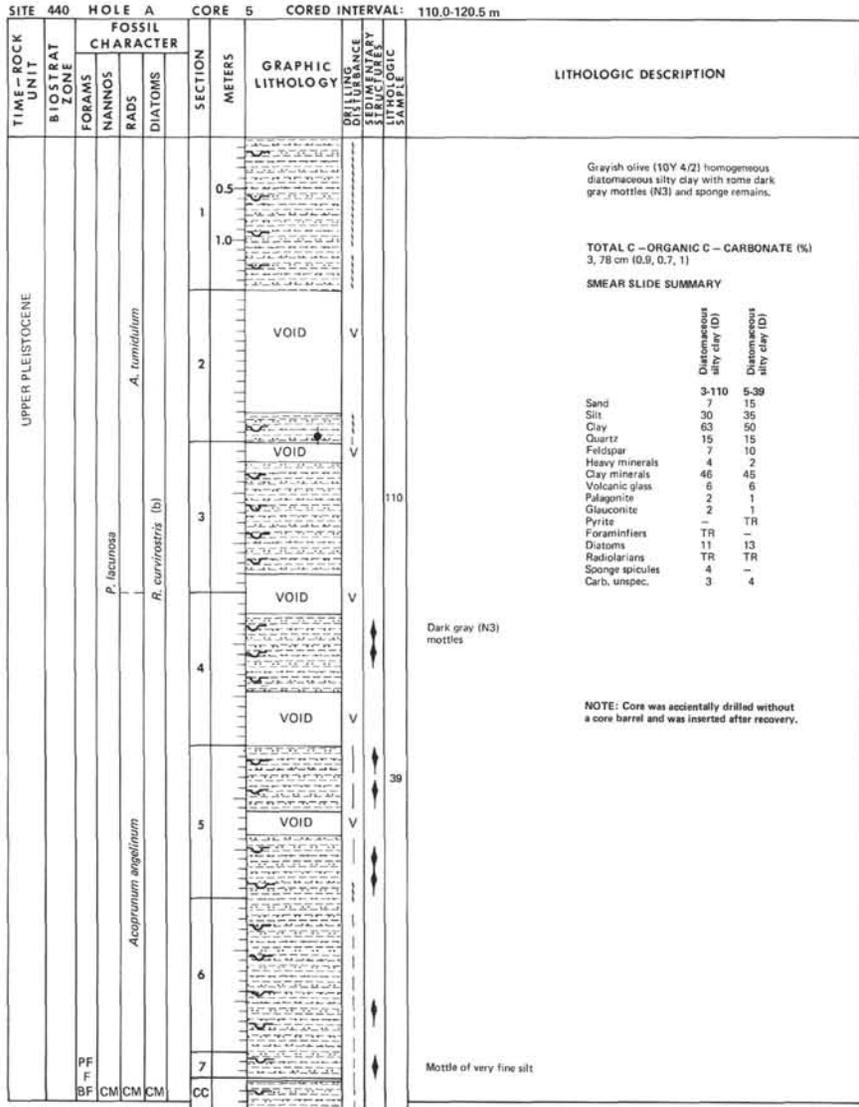
	1-148	2-25	3-93	4-94	5-86
Sand	8	3	14	12	11
Silt	30	67	32	34	21
Clay	62	30	54	54	68
Quartz	13	6	16	16	4
Feldspar	7	8	6	6	2
Heavy minerals	2	5	2	2	1
Clay minerals	48	4	39	39	67
Volcanic glass	8	75	8	8	2
Pyrite	-	2	1	1	-
Diatoms	16	TR	17	17	19
Radiolarians	-	-	3	3	TR
Sponge spicules	2	-	2	2	1
Carb. unsp.	4	-	6	6	4

NOTE:
♦ = sponge spicules

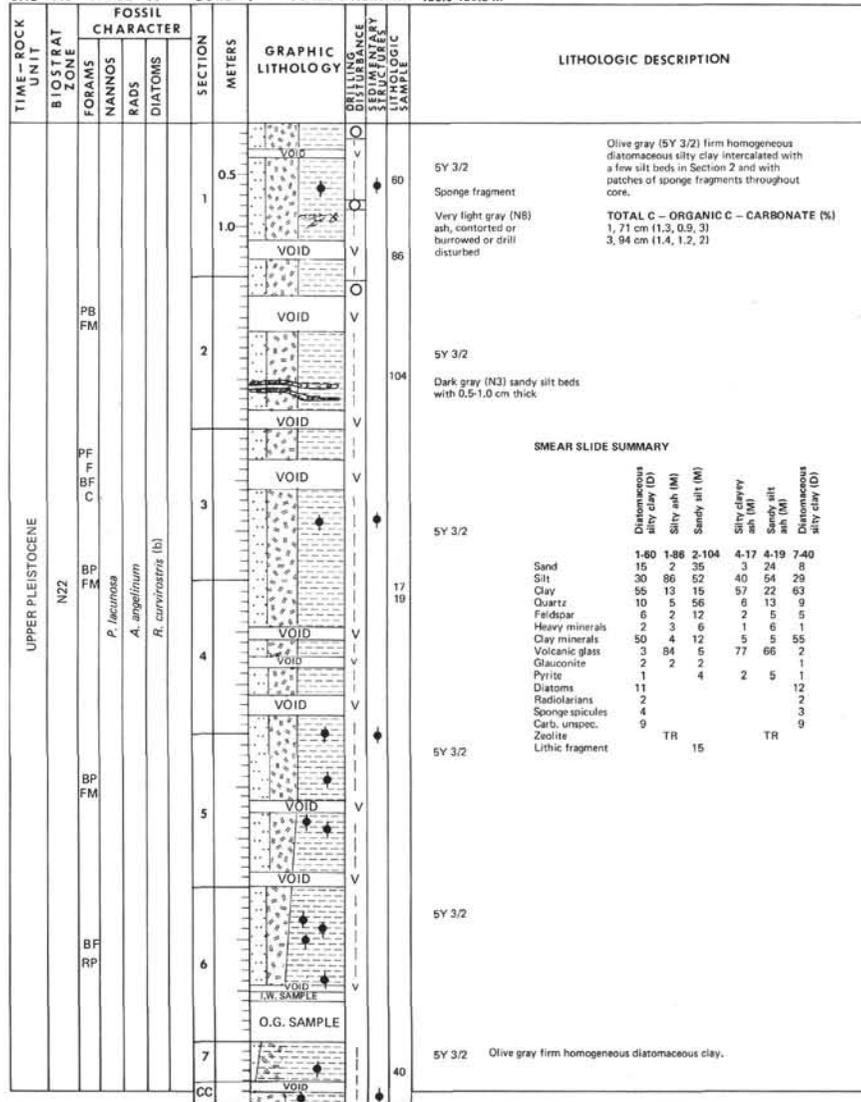
SITE 440 HOLE		CORE 7		CORED INTERVAL: 54.0-63.5 m																																																																																
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SEDIMENTARY STRUCTURES	LITHOLOGIC DESCRIPTION																																																																											
		FORAMS	NANNOS	RADS						DIATOMS																																																																										
UPPER PLEISTOCENE	P1 P1b	AM	CG	CM	1				Drilling cut																																																																											
					2				107	121	Sponge spicule																																																																									
									<p>Olive gray (5Y 3/2) diatomaceous silty clay with grayish blacks (N2) patches and black (N1) spots throughout core.</p> <p>TOTAL C - ORGANIC C - CARBONATE (%) 2, 100 cm (0.9, 0.7, 1)</p> <p>CARBONATE BOMB 2, 110-112 cm = 2-3%</p> <p>SMEAR SLIDE SUMMARY</p> <table border="1"> <thead> <tr> <th></th> <th>Diatomaceous silty clay (D)</th> <th>Diatomaceous silty clay (W)</th> <th>Sandy silty ash (M)</th> <th>CC</th> </tr> </thead> <tbody> <tr> <td>Sand</td> <td>1-107</td> <td>2-121</td> <td></td> <td></td> </tr> <tr> <td>Silt</td> <td>3</td> <td>5</td> <td></td> <td>13</td> </tr> <tr> <td>Clay</td> <td>20</td> <td>29</td> <td></td> <td>67</td> </tr> <tr> <td>Quartz</td> <td>77</td> <td>86</td> <td></td> <td>20</td> </tr> <tr> <td>Feldspar</td> <td>11</td> <td>5</td> <td></td> <td>5</td> </tr> <tr> <td>Heavy minerals</td> <td>5</td> <td>2</td> <td></td> <td>4</td> </tr> <tr> <td>Clay minerals</td> <td>3</td> <td>1</td> <td></td> <td>2</td> </tr> <tr> <td>Volcanic glass</td> <td>50</td> <td>54</td> <td></td> <td>12</td> </tr> <tr> <td>Glaucinite</td> <td>4</td> <td>3</td> <td></td> <td>70</td> </tr> <tr> <td>Zeolite minerals</td> <td>2</td> <td></td> <td></td> <td>2</td> </tr> <tr> <td>Diatoms</td> <td>15</td> <td>18</td> <td></td> <td>3</td> </tr> <tr> <td>Radiolarians</td> <td>3</td> <td></td> <td></td> <td>TR</td> </tr> <tr> <td>Sponge spicules</td> <td>1</td> <td>2</td> <td></td> <td></td> </tr> <tr> <td>Carb. unspc.</td> <td>6</td> <td>5</td> <td></td> <td></td> </tr> </tbody> </table>		Diatomaceous silty clay (D)	Diatomaceous silty clay (W)	Sandy silty ash (M)	CC	Sand	1-107	2-121			Silt	3	5		13	Clay	20	29		67	Quartz	77	86		20	Feldspar	11	5		5	Heavy minerals	5	2		4	Clay minerals	3	1		2	Volcanic glass	50	54		12	Glaucinite	4	3		70	Zeolite minerals	2			2	Diatoms	15	18		3	Radiolarians	3			TR	Sponge spicules	1	2			Carb. unspc.	6	5		
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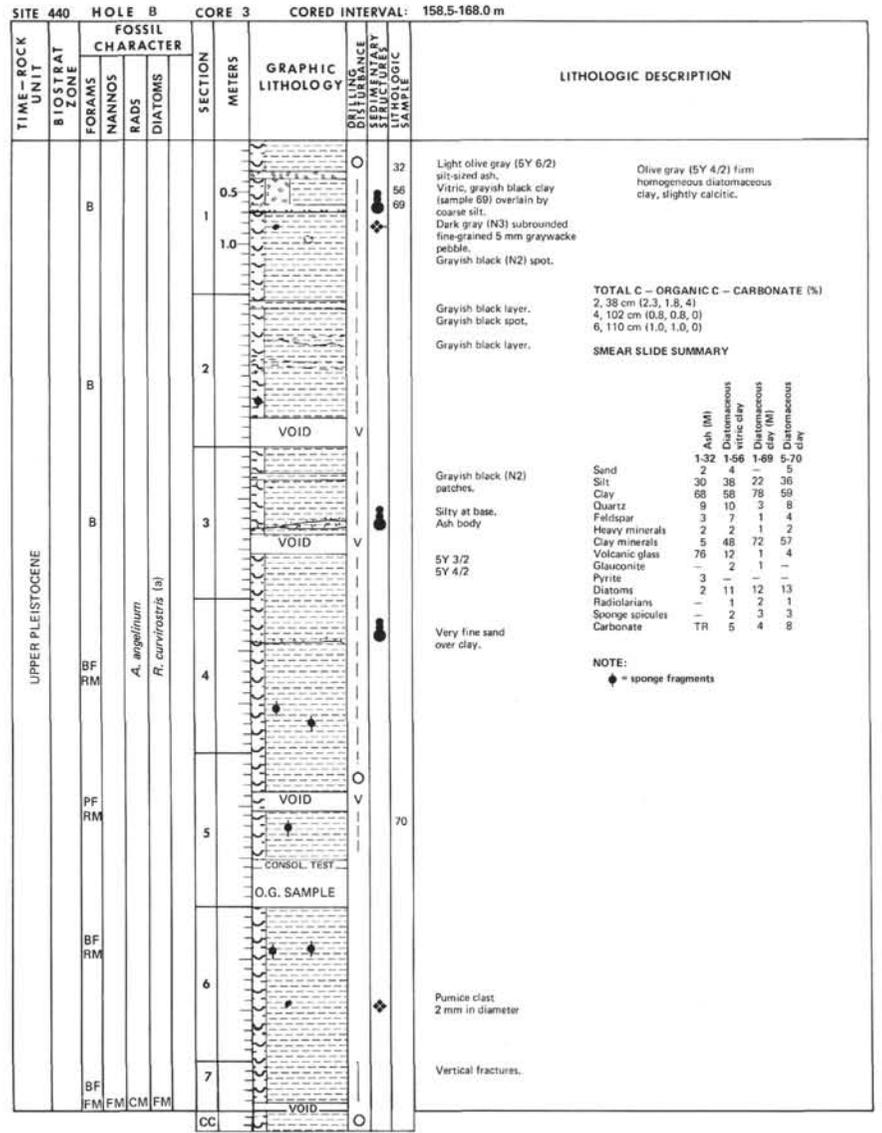
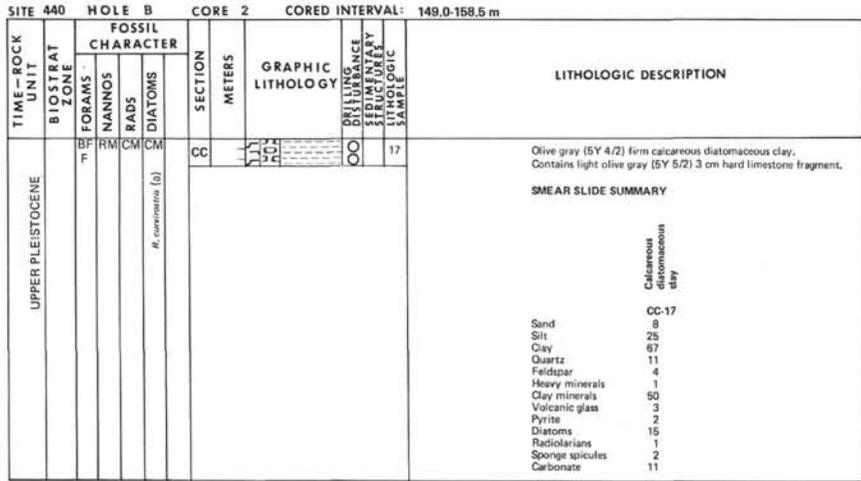
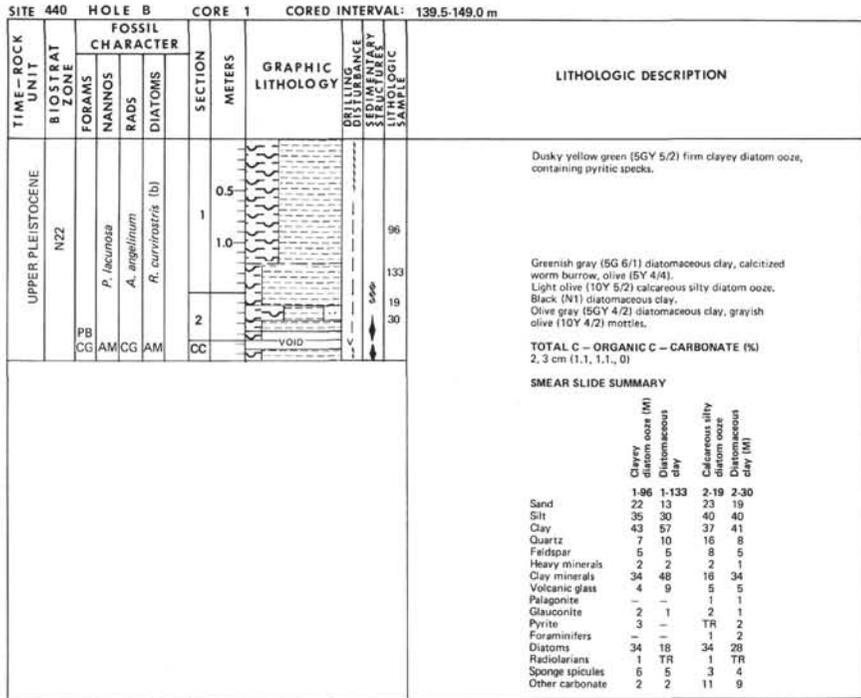
SITE 440 HOLE		CORE 8		CORED INTERVAL: 63.5-73.0 m																																																		
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SEDIMENTARY STRUCTURES	LITHOLOGIC DESCRIPTION																																													
		FORAMS	NANNOS	RADS						DIATOMS																																												
UPPER PLEISTOCENE	PB CM	CP	CG	CM	1				Olive gray (5Y 3/2) homogeneous diatomaceous clay throughout core.																																													
					2				34	62	Sponge fragments																																											
									<p>CARBONATE BOMB 2, 25-27 cm = 5-6%</p> <p>SMEAR SLIDE SUMMARY</p> <table border="1"> <thead> <tr> <th></th> <th>Diatomaceous clay (D)</th> <th>Diatomaceous silty clay (D)</th> </tr> </thead> <tbody> <tr> <td>Sand</td> <td>1-34</td> <td>2-62</td> </tr> <tr> <td>Silt</td> <td>4</td> <td>5</td> </tr> <tr> <td>Clay</td> <td>19</td> <td>20</td> </tr> <tr> <td>Quartz</td> <td>77</td> <td>75</td> </tr> <tr> <td>Feldspar</td> <td>11</td> <td>10</td> </tr> <tr> <td>Heavy minerals</td> <td>6</td> <td>5</td> </tr> <tr> <td>Clay minerals</td> <td>3</td> <td>2</td> </tr> <tr> <td>Volcanic glass</td> <td>52</td> <td>56</td> </tr> <tr> <td>Glaucinite</td> <td>3</td> <td>3</td> </tr> <tr> <td>Pyrite</td> <td>1</td> <td></td> </tr> <tr> <td>Diatoms</td> <td>17</td> <td>18</td> </tr> <tr> <td>Radiolarians</td> <td>2</td> <td>1</td> </tr> <tr> <td>Sponge spicules</td> <td>1</td> <td>1</td> </tr> <tr> <td>Carb. unspc.</td> <td>4</td> <td>3</td> </tr> </tbody> </table>		Diatomaceous clay (D)	Diatomaceous silty clay (D)	Sand	1-34	2-62	Silt	4	5	Clay	19	20	Quartz	77	75	Feldspar	11	10	Heavy minerals	6	5	Clay minerals	3	2	Volcanic glass	52	56	Glaucinite	3	3	Pyrite	1		Diatoms	17	18	Radiolarians	2	1	Sponge spicules	1	1	Carb. unspc.	4	3
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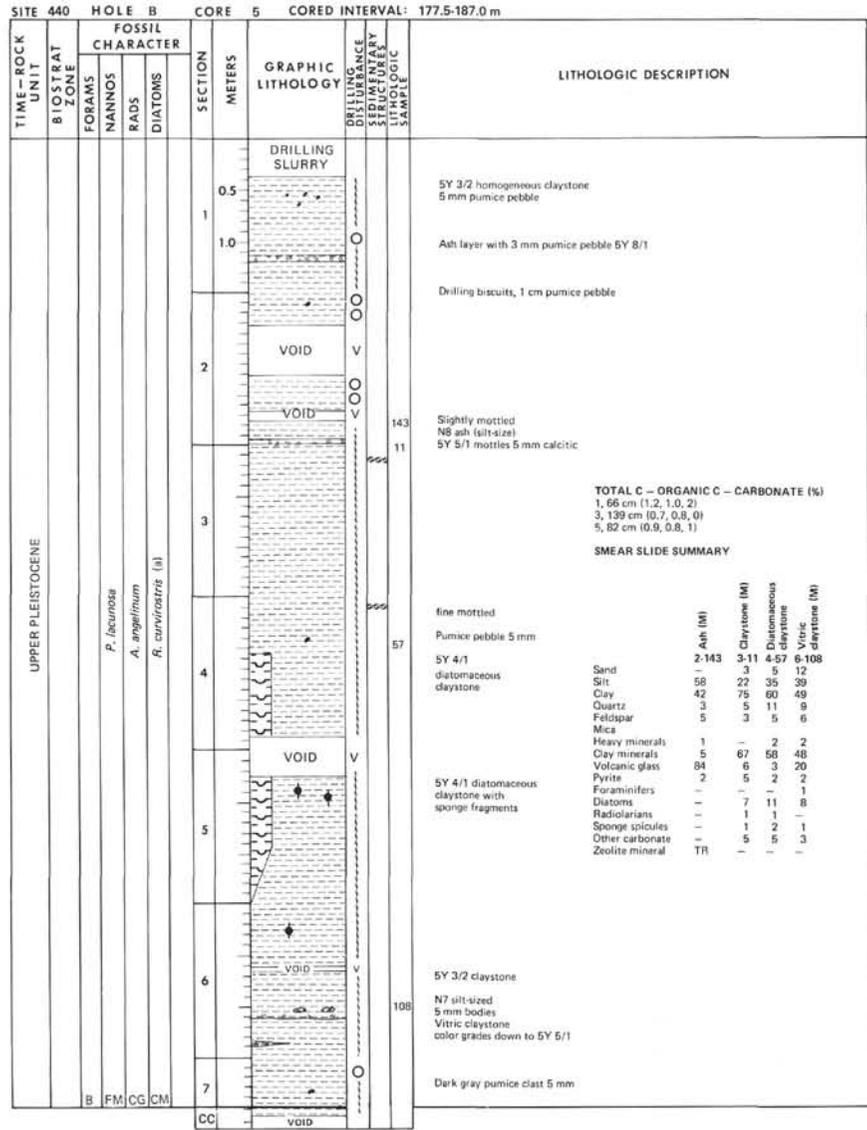
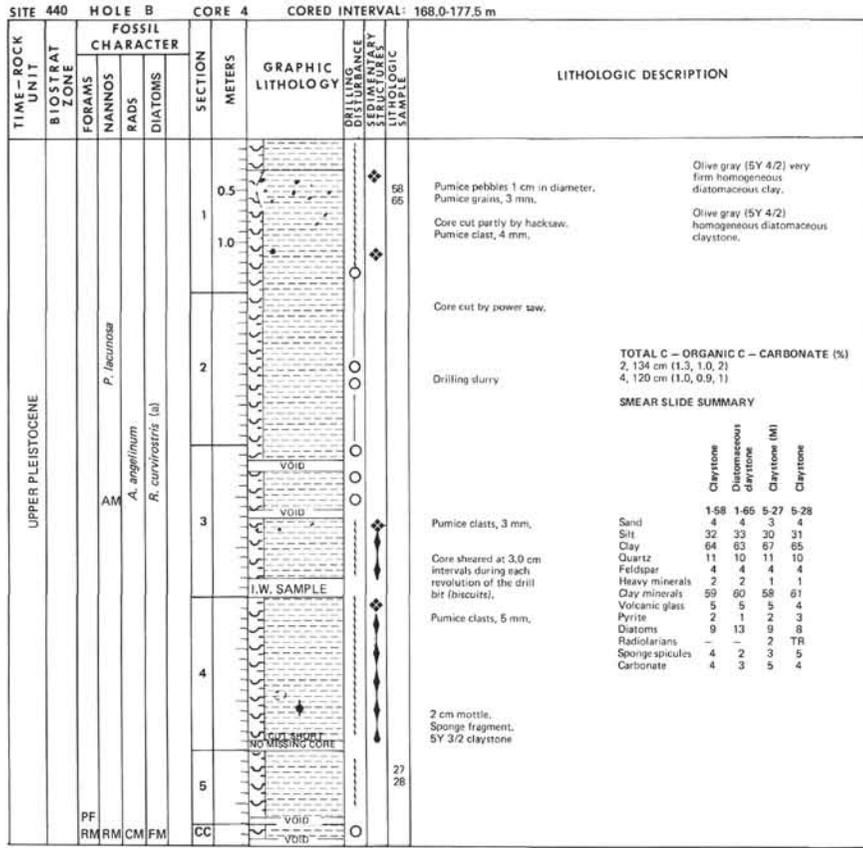


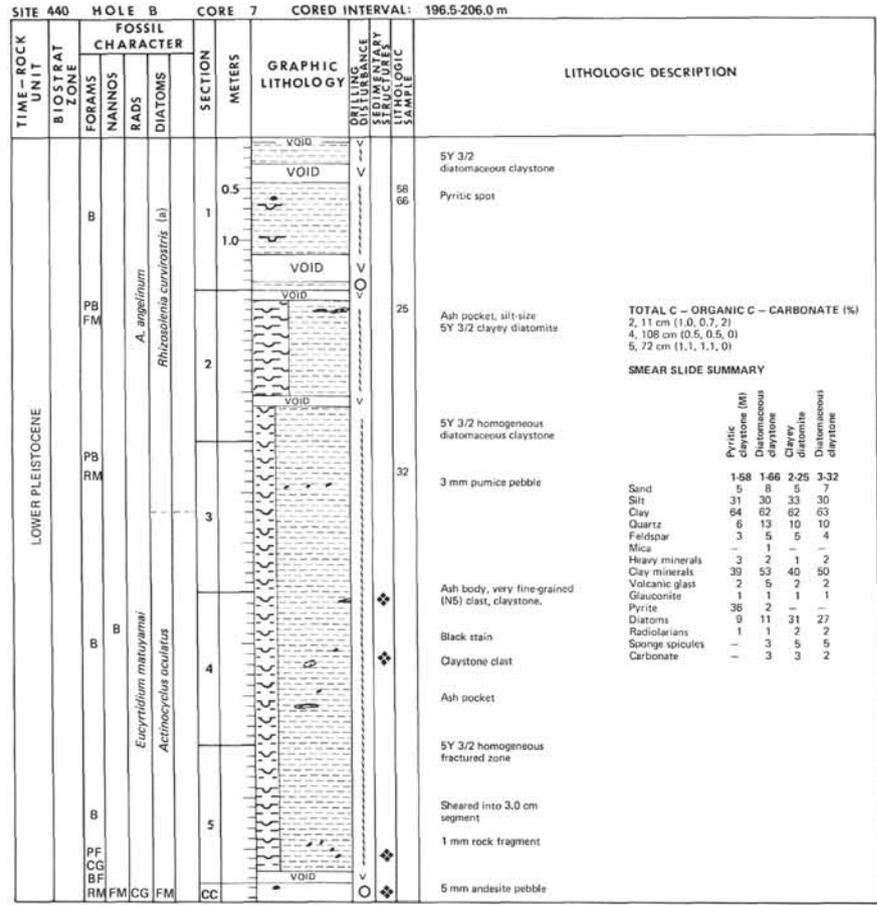
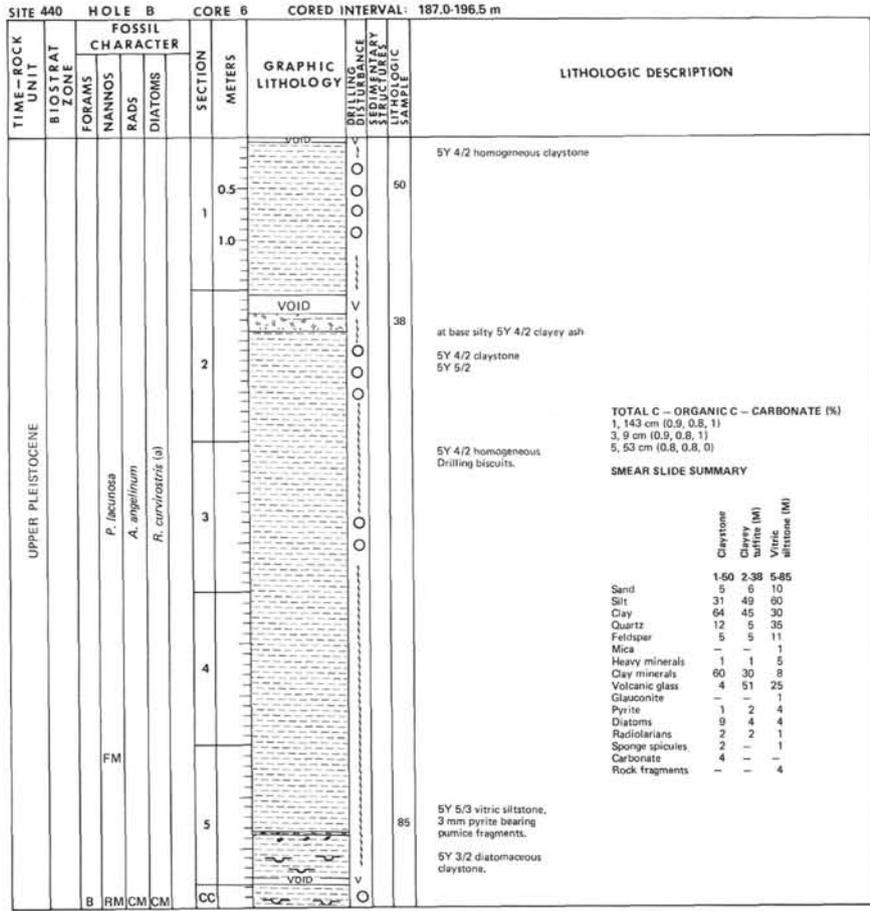


SITE 440 HOLE A CORE 7 CORED INTERVAL: 130.0-139.5 m







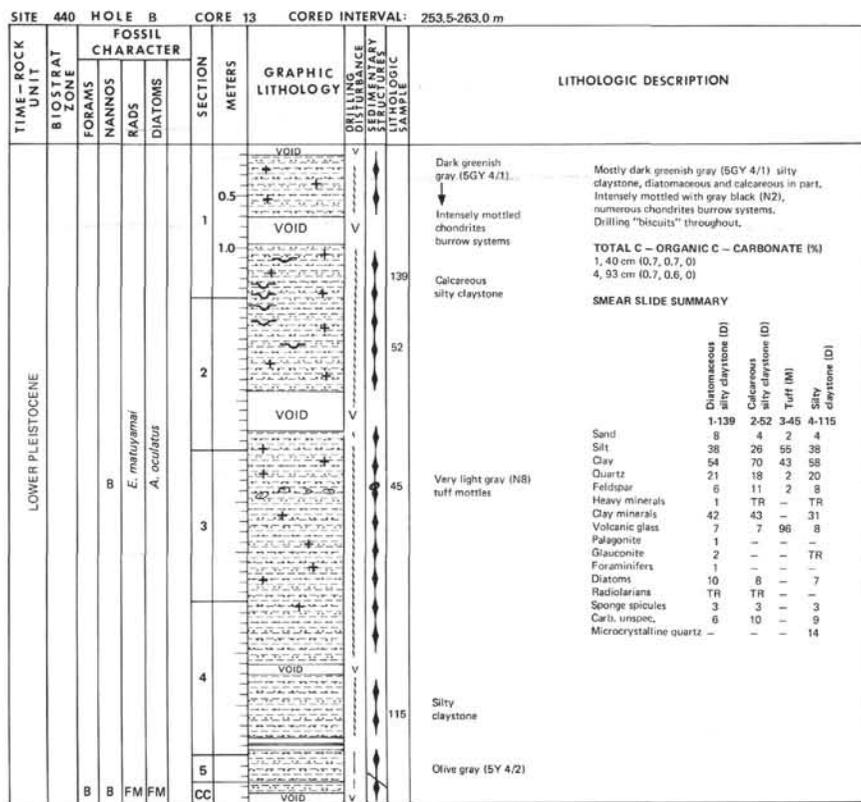
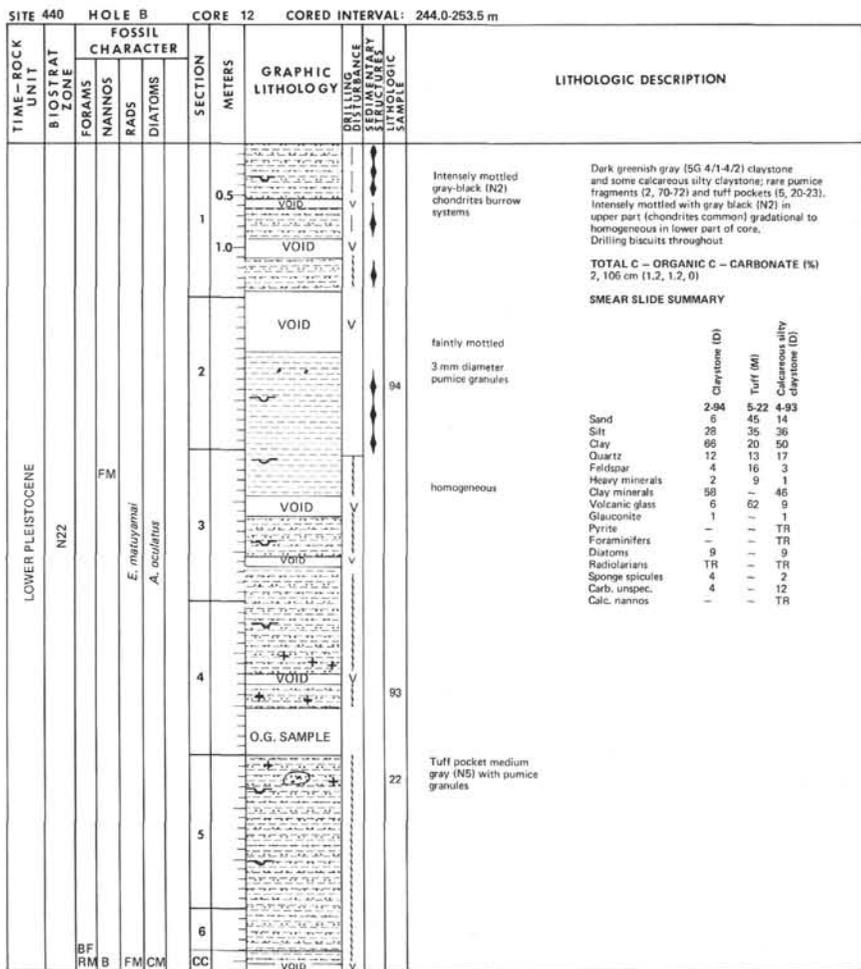


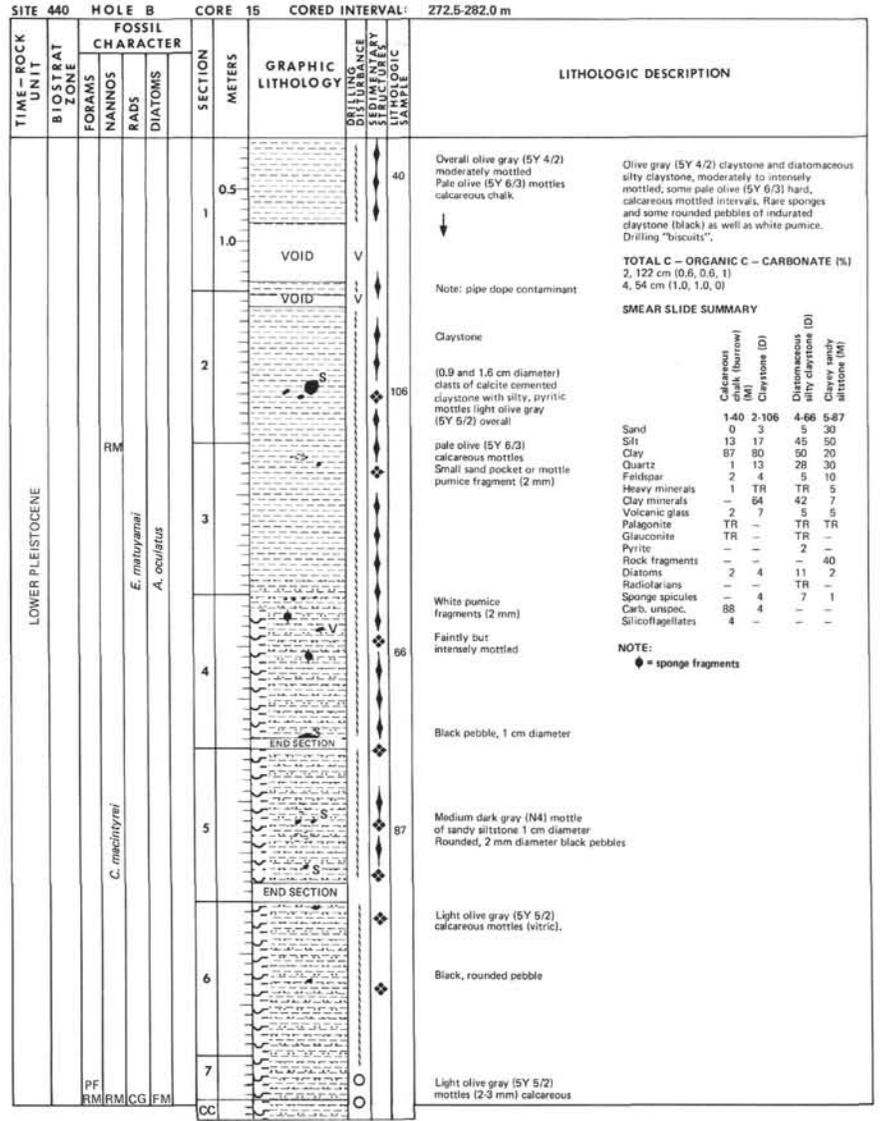
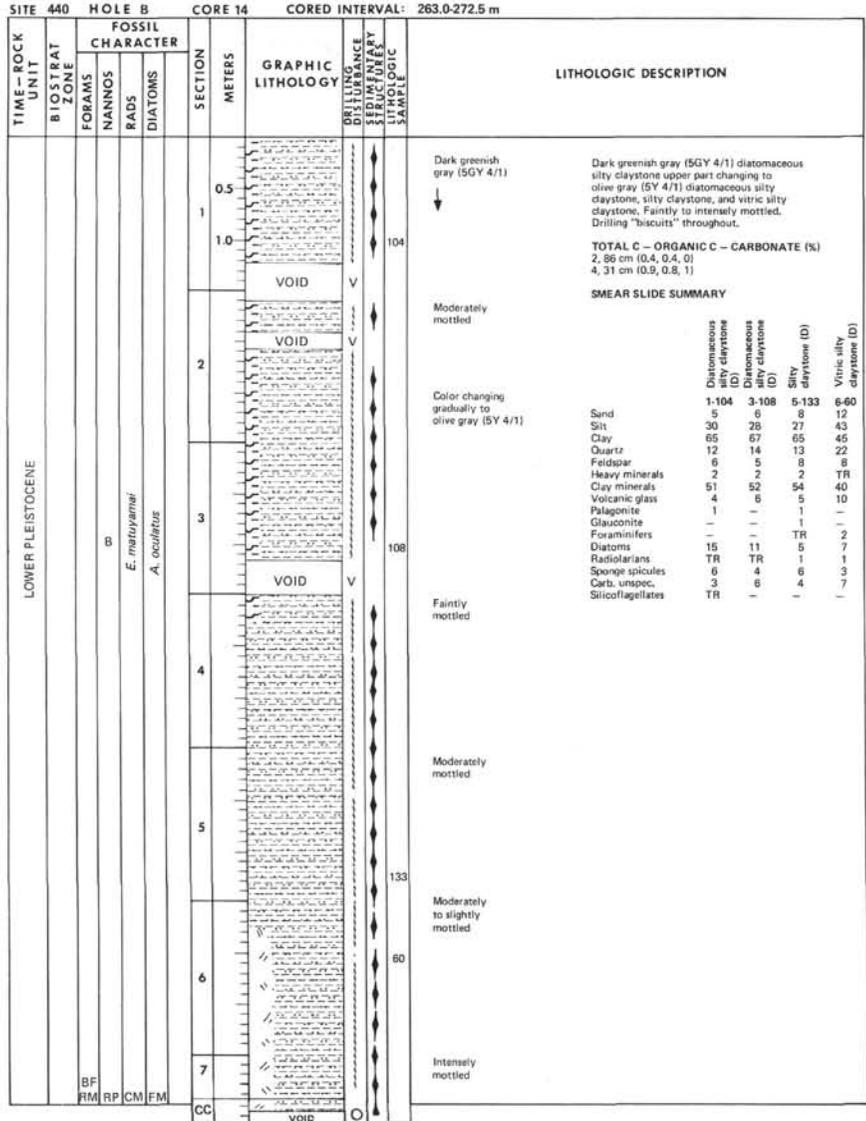
SITE 440		HOLE B			CORE 8		CORED INTERVAL: 206.0-215.5 m	
TIME - ROCK UNIT	BIOTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS				
LOWER PLEISTOCENE	B	RP	CG	FM	0.5	VOID	67	Homogeneous to faintly streaked and mottled grayish olive (10Y 4/2) silty claystone to silty diatomaceous claystone with sandy spots (quartz). (Sand disseminated as coarse white angular quartz grains.)
					1.0	VOID		
					2	VOID		Faintly streaked and mottled grayish olive (10Y 4/2) diatomaceous claystone. Noticeable sand content.
					25	VOID	25	TOTAL C - ORGANIC C - CARBONATE (%) 3, 27 cm (0.7, 0.7, 1)
					36	VOID	36	Soft sandy silty claystone (10Y 4/2). Whole section angular granules to sand grains disseminated in matrix.
					3	VOID	Angular clayey gravel of white pumice fragments 0.3-0.5 cm. Subround fragment of 0.7 cm diameter granitic rock.	
					4	VOID	Slightly streaked and mottled grayish olive (10Y 4/2) diatomaceous clay. 1 cm angular fragment of 10R 2/8 very dusky metasediment or altered sediment with angular quartz granules.	
					5	VOID	Faintly streaked and mottled grayish olive (10Y 4/2) silty claystone.	
					78	VOID	Siltier sandier interval.	
					CC	VOID		

SMEAR SLIDE SUMMARY				
	Silty claystone	Pumice (M)	Sandy clayey siltstone	Silty claystone
Sand	1-67	3-25	3-36	6-78
Silt	12	2	28	8
Clay	48	-	30	50
Quartz	28	2	24	28
Feldspar	5	2	15	6
Heavy minerals	2	1	4	1
Clay minerals	46	-	30	43
Volcanic glass	7	94	7	8
Palagonite	1	-	1	1
Glaucinite	1	-	1	1
Pyrite	-	-	2	2
Diatoms	5	1	9	5
Radiolarians	-	-	TR	TR
Sponge spicules	2	-	4	4
Carbonate	3	-	3	1

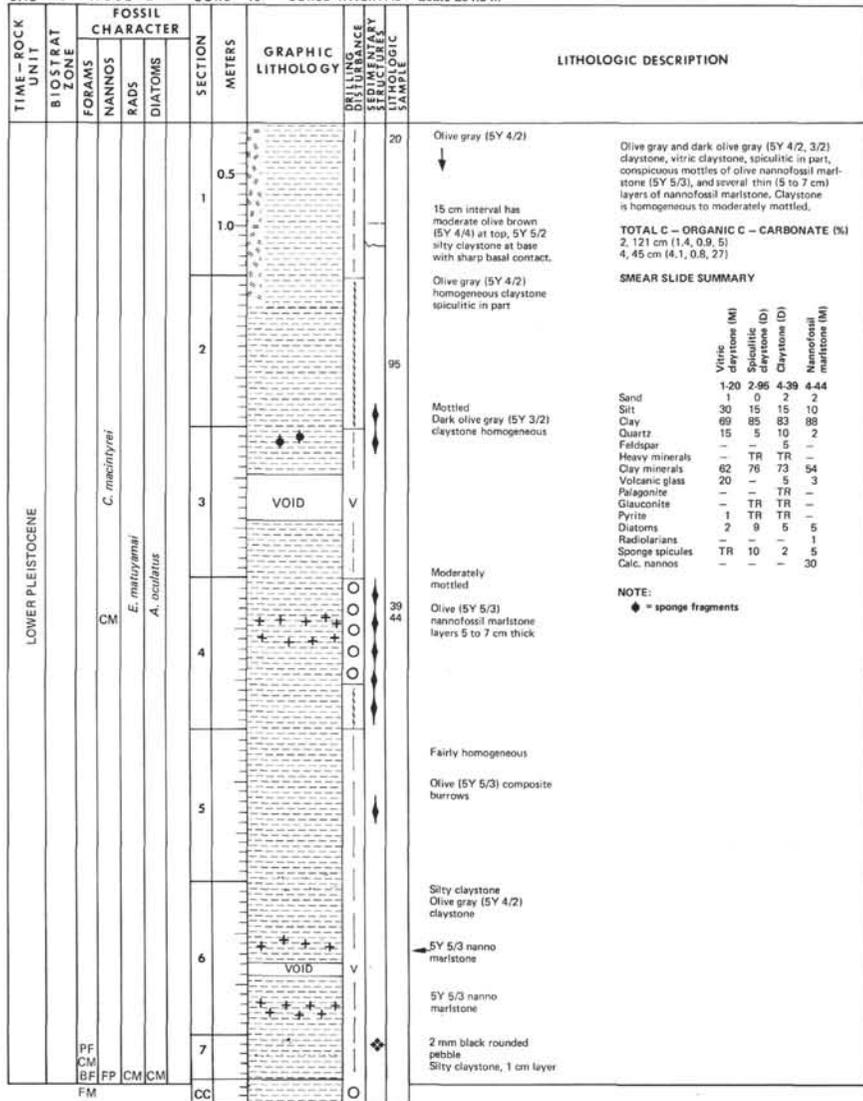
SITE 440		HOLE B			CORE 9		CORED INTERVAL: 215.5-225.0 m	
TIME - ROCK UNIT	BIOTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS				
LOWER PLEISTOCENE	B	B	FM	FM	0.5	VOID	41	Drilling biscuits throughout. Grayish olive (10Y 4/2) burrow mottled claystone (silty).
					1.0	VOID		
					2	VOID		Grayish olive (10Y 4/2) silty claystone faintly mottled.
					3	VOID		
					4	VOID	Light yellow gray olive (5Y 6/2) calcareous chalk. 2 cm laminations vitric clayey siltstone.	
					CC	VOID		

SMEAR SLIDE SUMMARY				
	Silty claystone (D)	Silty claystone (D)	Chalk (M)	Vitric clayey siltstone (M)
	1-41	2-62	4-13	4-34
Sand	4	4	6	22
Silt	30	28	23	42
Clay	66	68	71	36
Quartz	25	16	2	27
Feldspar	6	11	1	8
Heavy minerals	1	3	-	4
Clay minerals	51	43	-	32
Volcanic glass	4	8	4	11
Palagonite	1	-	1	1
Glaucinite	1	1	-	1
Diatoms	6	8	5	6
Radiolarians	TR	-	TR	TR
Sponge spicules	3	4	2	3
Carbonate	2	1	85	7
Rock fragments	-	5	-	-

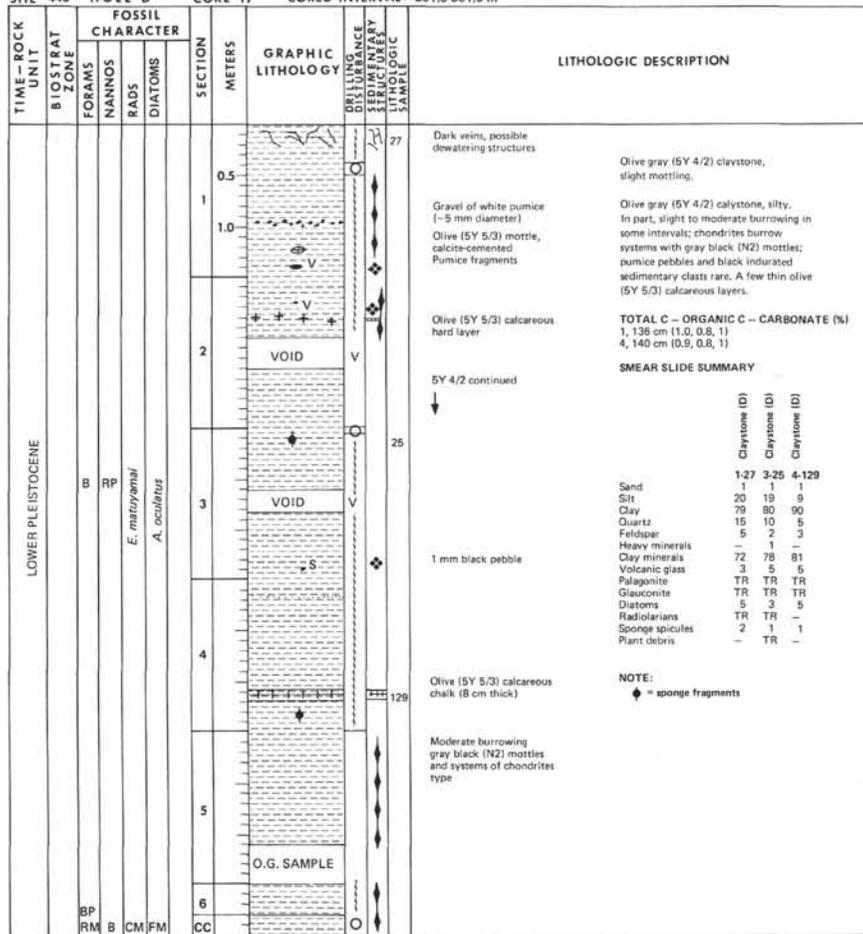


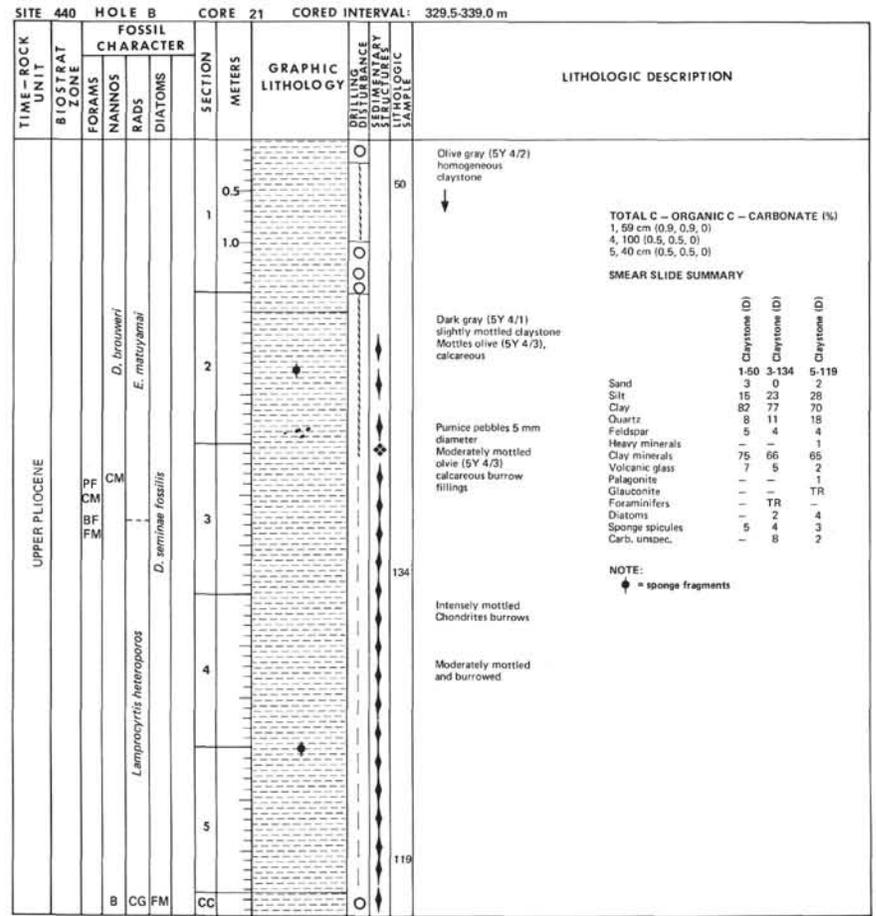
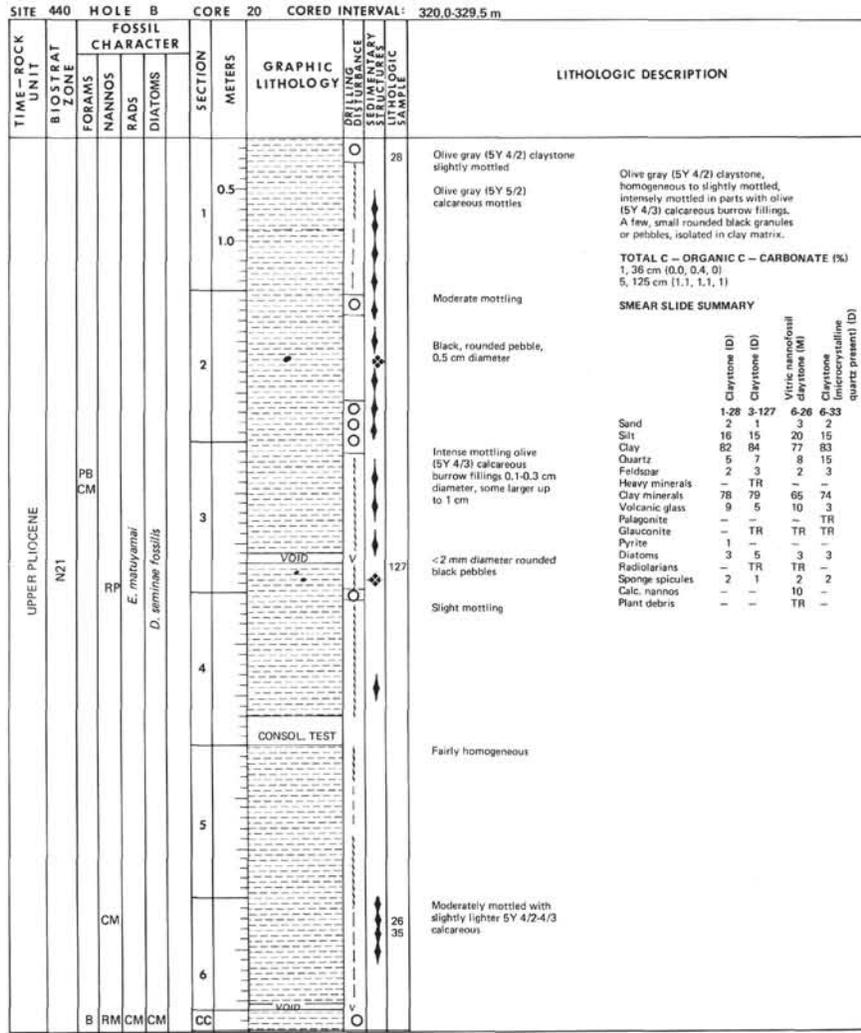


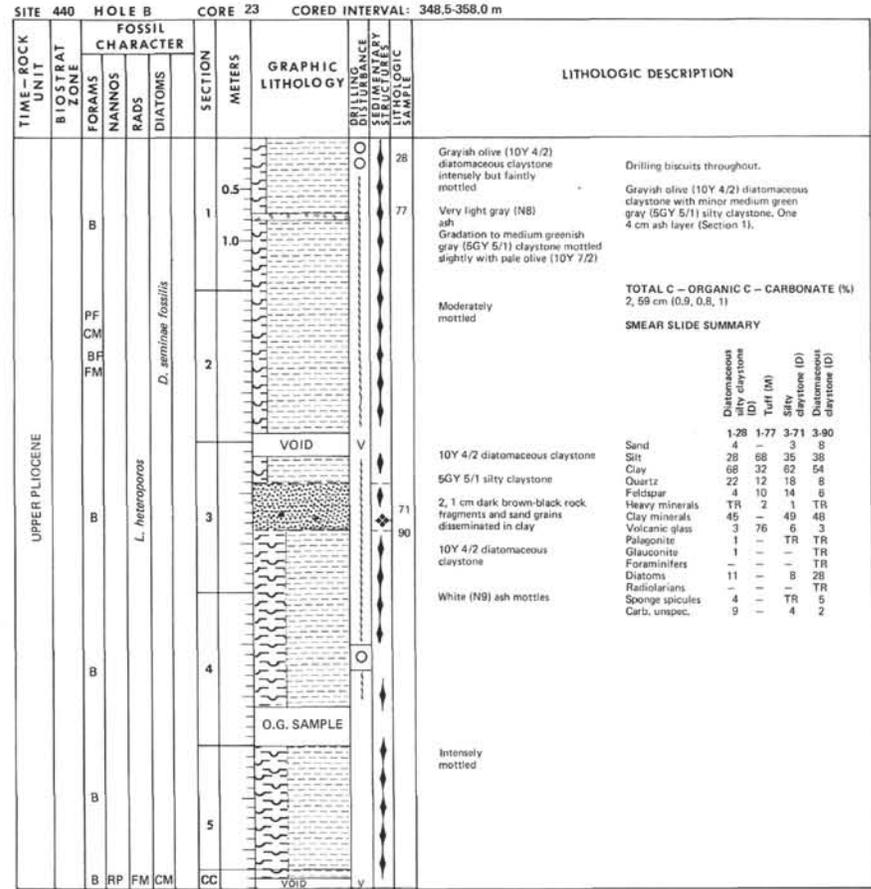
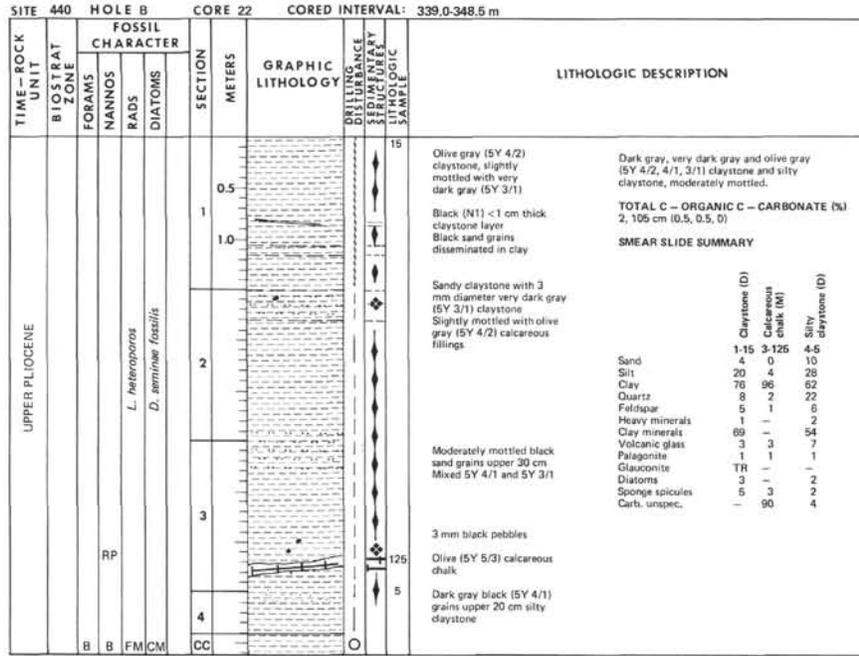
SITE 440 HOLE B CORE 16 CORED INTERVAL: 282.0-291.5 m

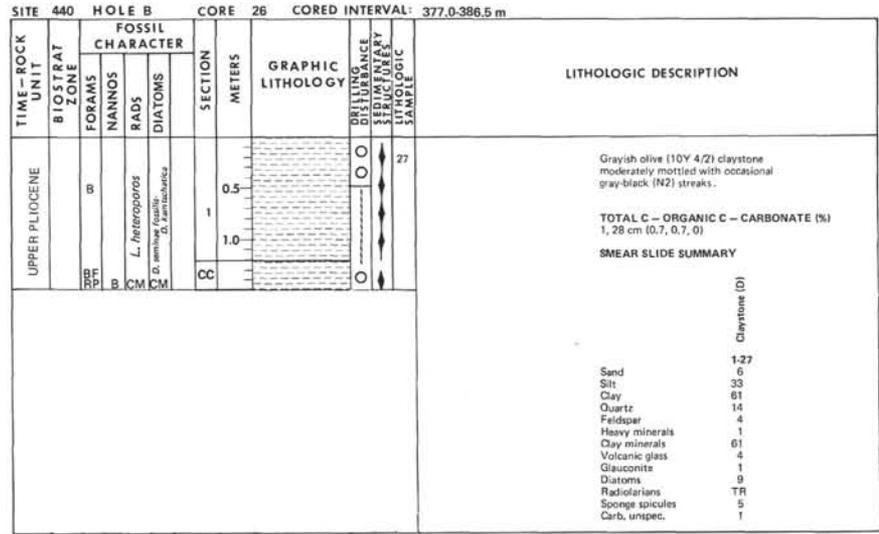
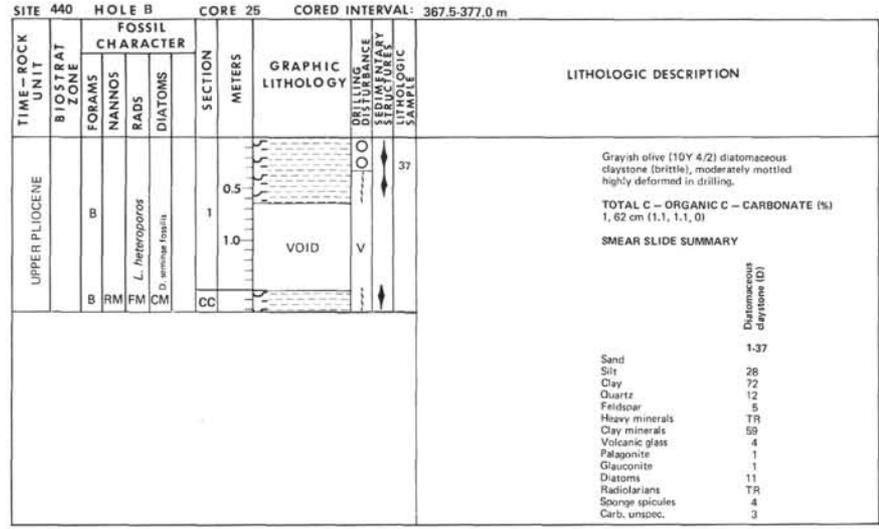
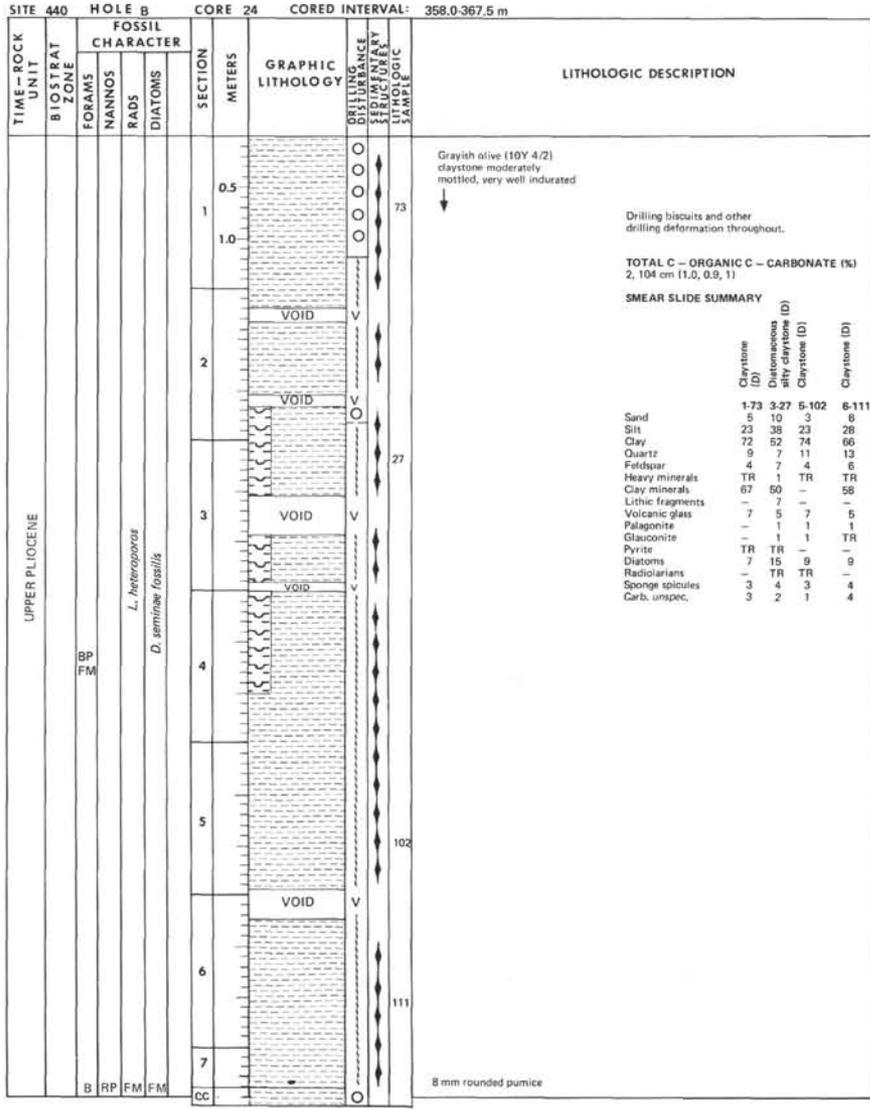


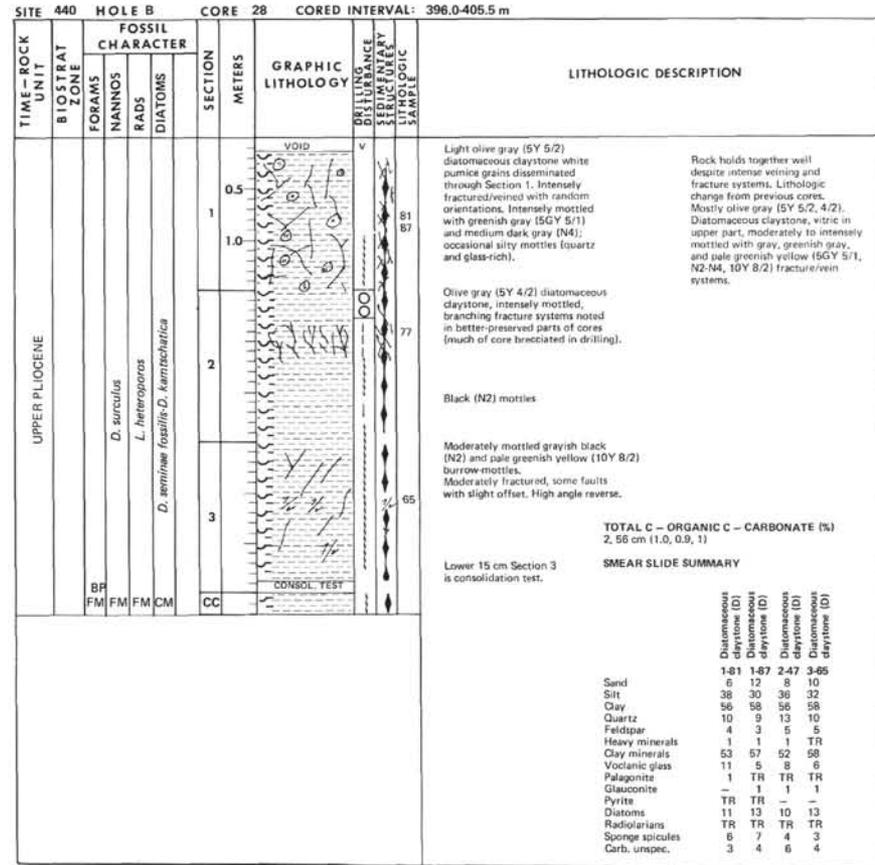
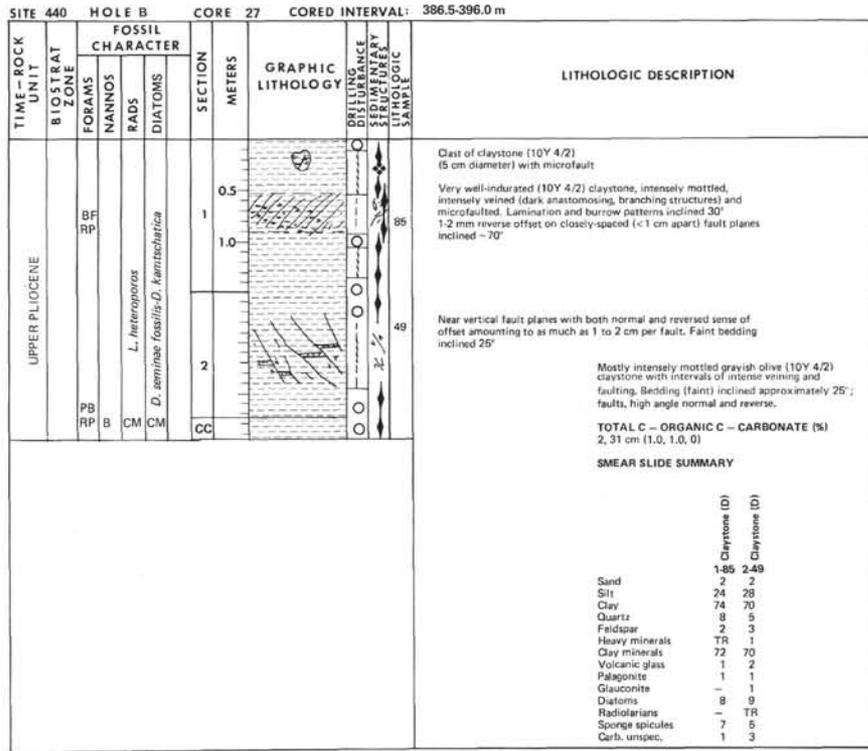
SITE 440 HOLE B CORE 17 CORED INTERVAL: 291.5-301.0 m

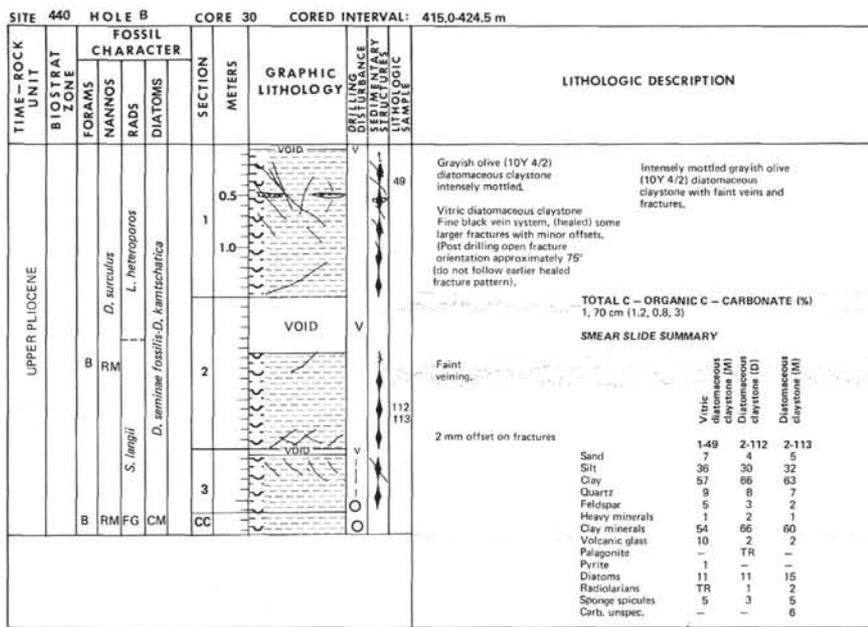
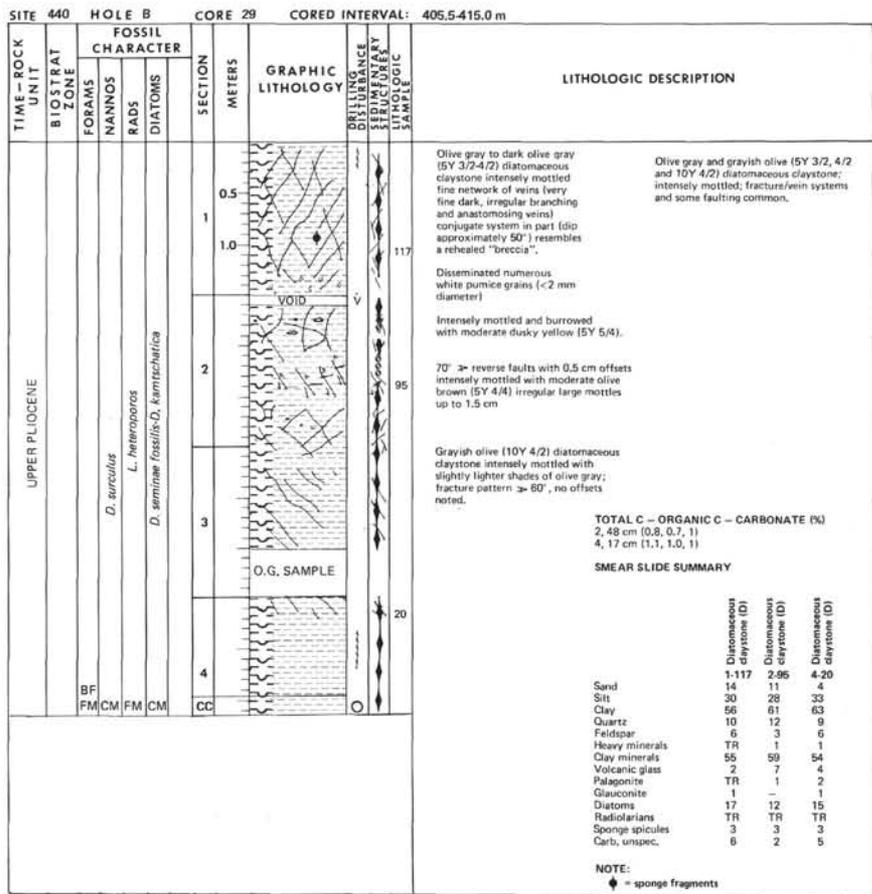












SITE 440		HOLE B		CORE 33		CORED INTERVAL: 443.5-453.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
UPPER PLIOCENE	B	G. langii	D. sominae fossilis-D. karmscharica	1	0.5	VOID	Reverse fault with slickensides, dip-slip motion, plane dips 28°. Silty claystone, dip 3° (hole deviation 2.8°)
				1	1.0		Sponge fragments
				2			Moderate mottling
				3			Claystone
				3			Horizontal burrows
				24		Tuff	
				4	73	Light gray tuff	
				5		Faint veins	
				CC 5		VOID	

TOTAL C - ORGANIC C - CARBONATE (%)	
2, 49 cm (0.7, 0.7, 0)	
4, 107 cm (1.0, 0.8, 2)	

CARBONATE BOMB	
3, 42-44 cm = 0-1%	

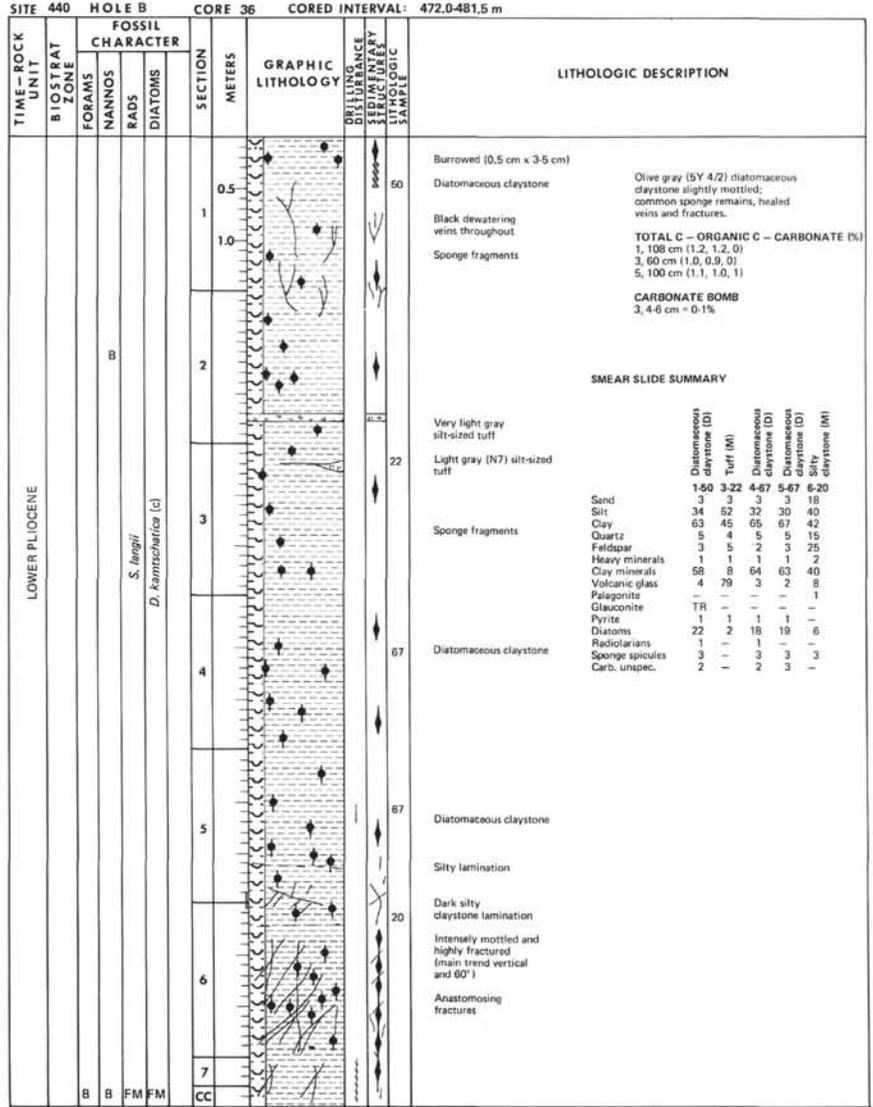
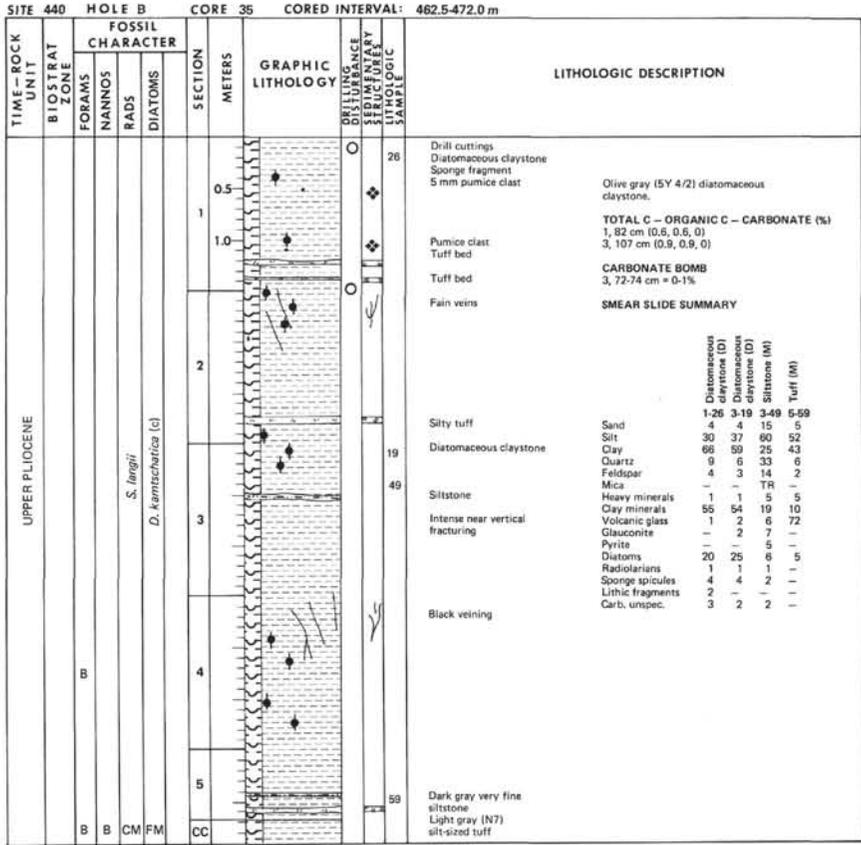
SMEAR SLIDE SUMMARY				
	Silty claystone (M)	Claystone (D)	Tuff (M)	Diatomaceous claystone (D)
Sand	1-67	3-33	4-24	4-73
Silt	10	4	2	9
Clay	40	43	44	30
Quartz	23	14	5	13
Feldspar	8	6	5	5
Heavy minerals	2	1	1	1
Clay minerals	48	50	12	59
Volcanic glass	7	7	73	3
Pyrite	6	1	4	2
Diatoms	8	9	10	10
Radiolarians	1	1	1	1
Sponge spicules	2	5	3	3
Carb. unspec.	2	6	1	4

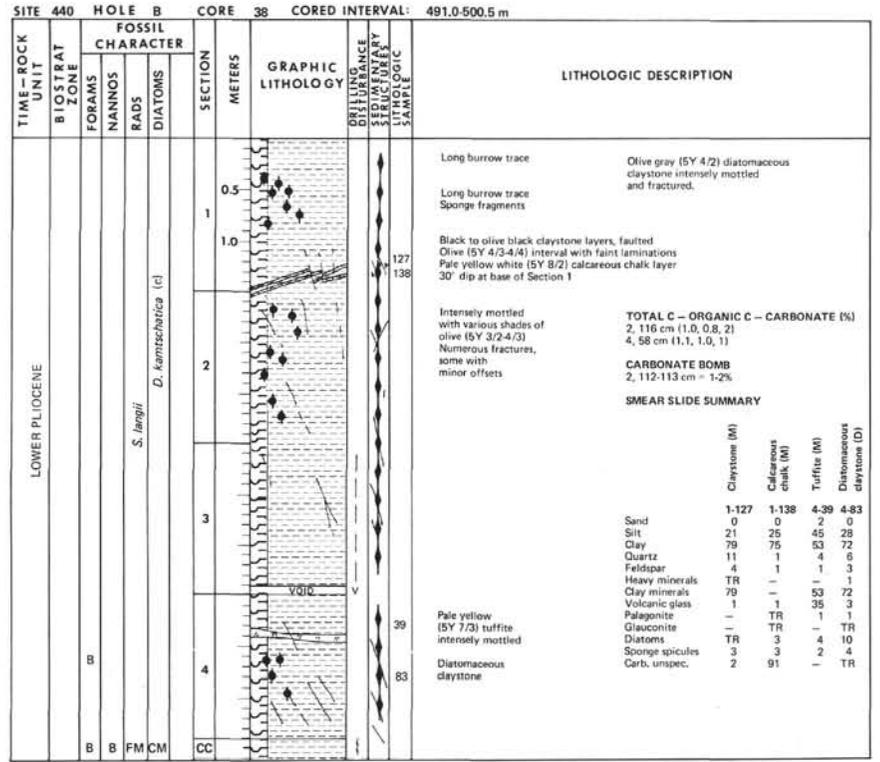
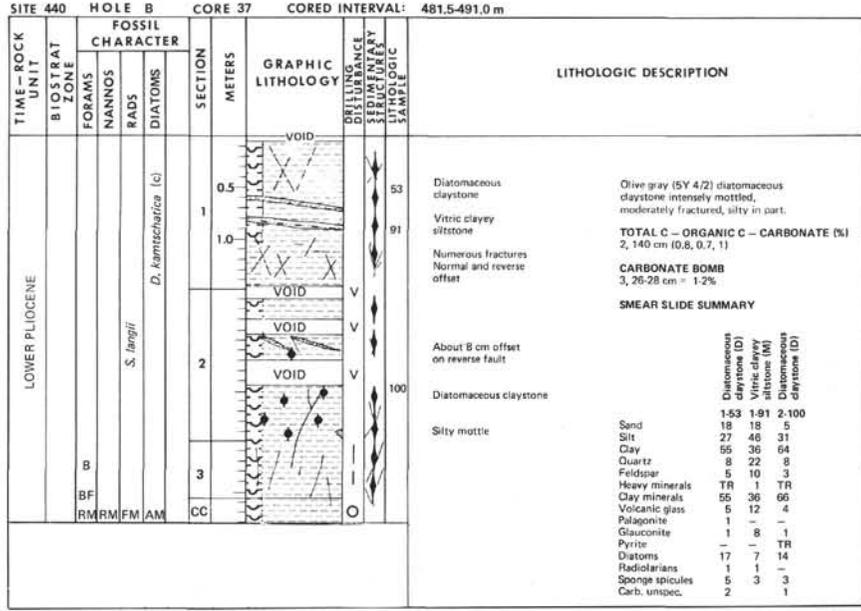
SITE 440		HOLE B		CORE 34		CORED INTERVAL: 453.0-462.5 m		
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION	
		FORAMS	NANNOS					RADS
UPPER PLIOCENE	B	R. pseudumbillica?	S. langii	D. karmscharica (c)	1	0.5	VOID	Dark olive (5Y 3/2) diatomaceous claystone with persistent +75° partings throughout.
					1	1.0		3 mm pumice clast
					2			Light gray silt-size tuff
					2			Light gray (N7) tuff layer silt-size
					2			Light gray very fine-grained tuff
				27		Sponge fragments		
				67		Very light gray (N8) silt-sized tuff		
				3		Ash pocket		
				3		Light gray (N7) fine-grained tuff		
				4		Diatomaceous claystone		
				4		Tuff layer (5 mm thick)		
				4		Pumice fragment		
				4		Dewatering veins		
				CC				

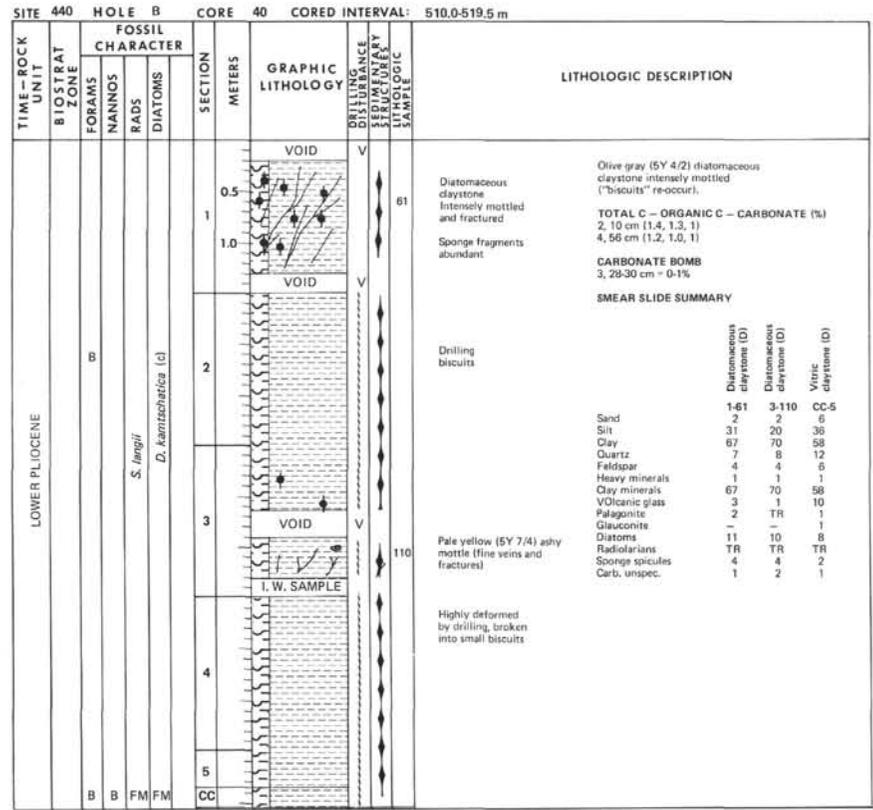
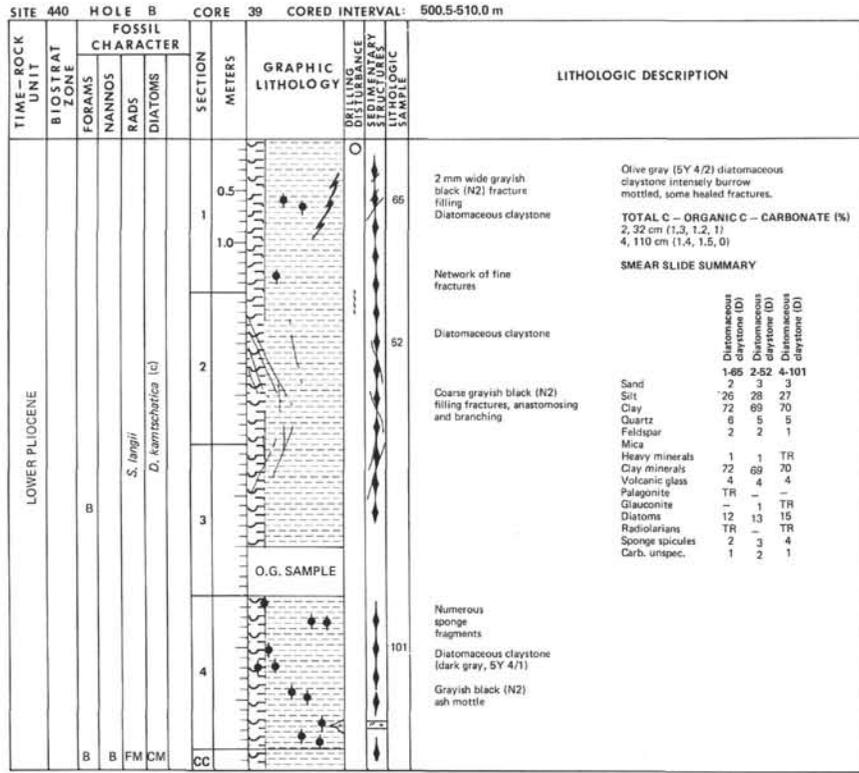
TOTAL C - ORGANIC C - CARBONATE (%)	
1, 114 cm (0.7, 0.7, 0)	
2, 98 cm (0.8, 0.8, 0)	

CARBONATE BOMB	
3, 3-5 cm = 1-2%	

SMEAR SLIDE SUMMARY				
	Tuff (M)	Silty tuff (M)	Diatomaceous claystone (D)	
Sand	1-146	2-27	3-67	
Silt	0	15	5	
Clay	22	35	32	
Quartz	78	52	63	
Feldspar	2	5	10	
Heavy minerals	4	12	4	
Clay minerals	5	1	1	
Volcanic glass	10	10	51	
Pyrite	89	58	2	
Diatoms	1	6	2	
Radiolarians	1	4	18	
Sponge spicules	1	2	5	
Carb. unspec.	1	3	5	
Zeolite	1	3	1	





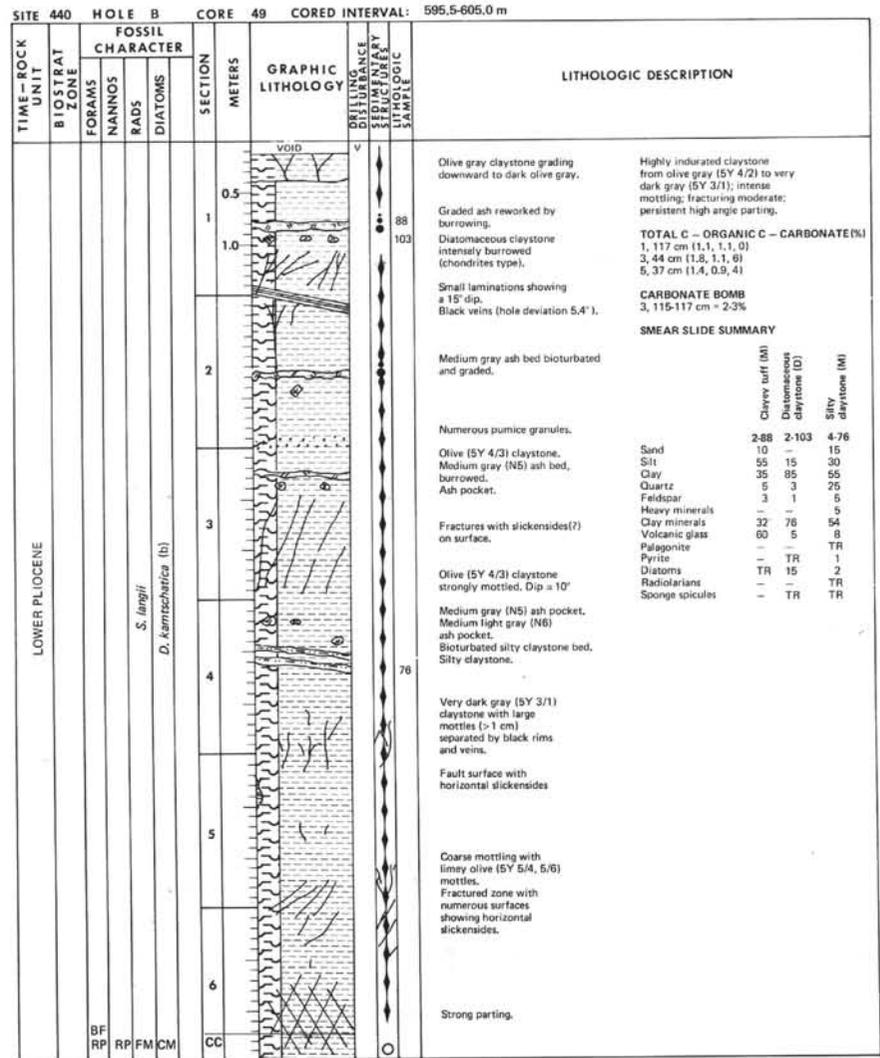
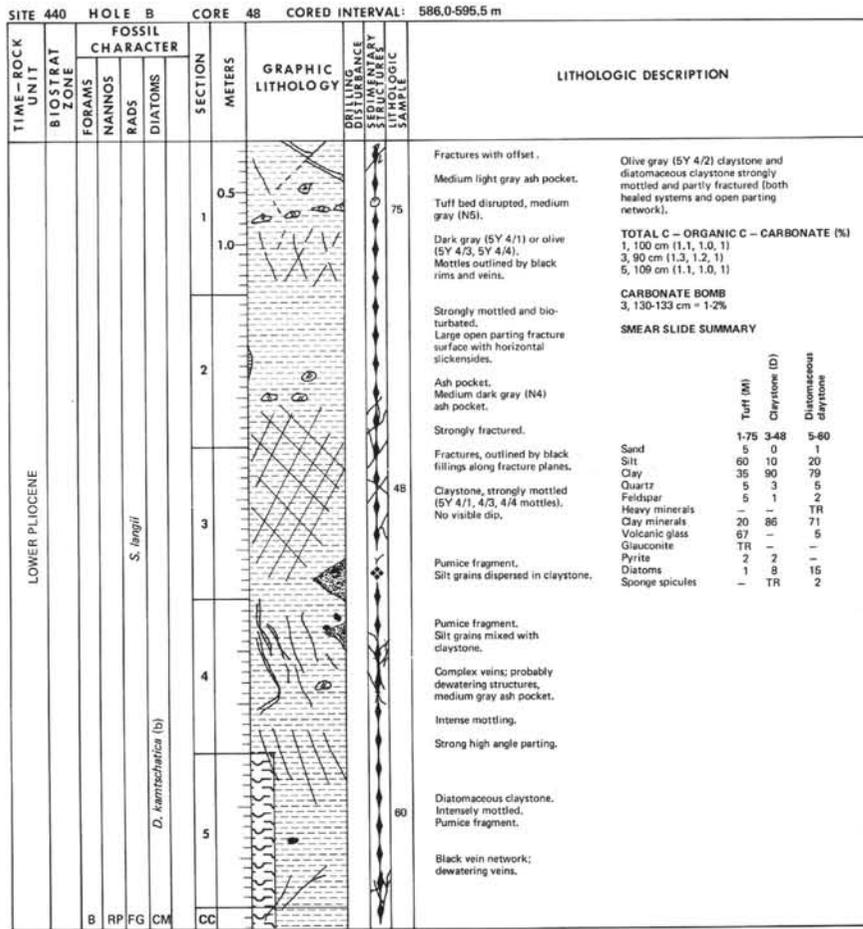


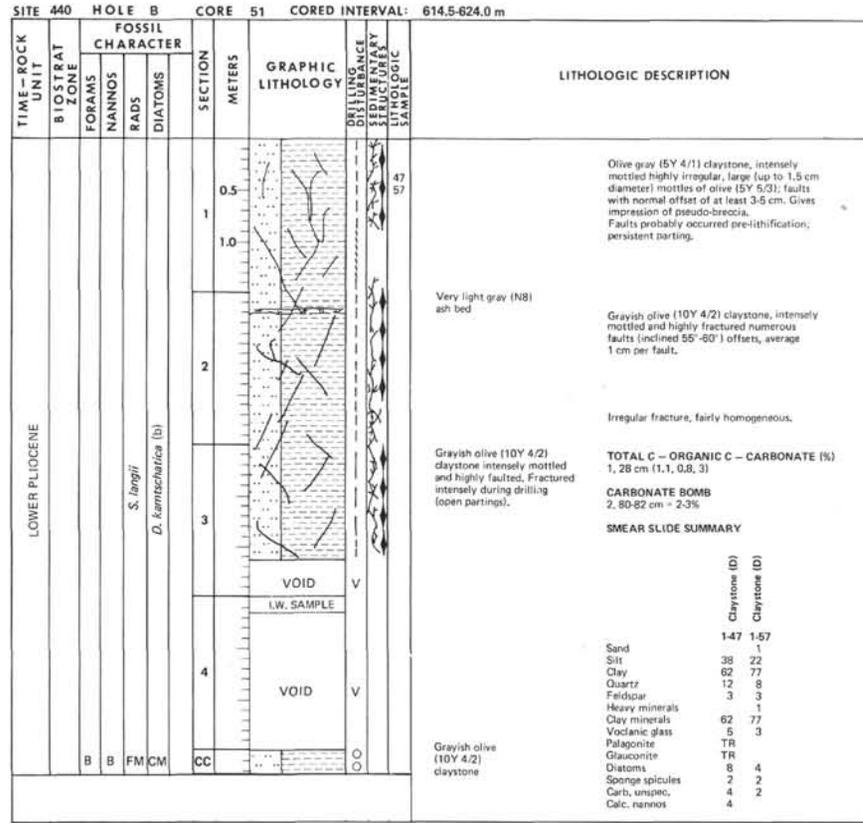
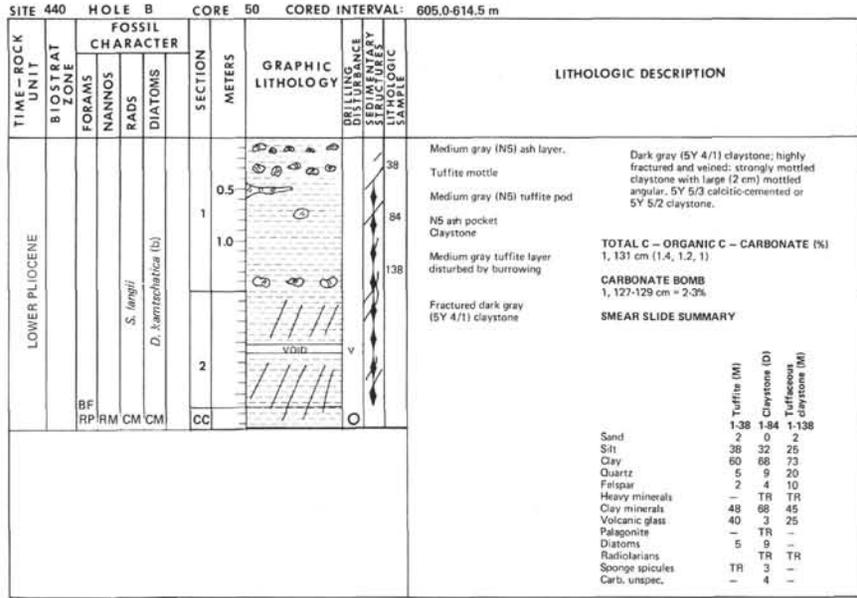
SITE 440		HOLE B		CORE 44		CORED INTERVAL: 548.0-557.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS			
LOWER PLOCIENE		BF RM	FM	FM	CC		Diatomaceous claystone, olive gray (5Y 4/2).
			<i>S. langii</i>				
			<i>D. kamtschatica</i> (c)				

SITE 440		HOLE B		CORE 45		CORED INTERVAL: 557.5-567.0 m																																																									
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION																																																								
		FORAMS	NANNOS	RADS				DIATOMS																																																							
LOWER PLOCIENE		B	B	FP	CM	CC	<p>Fractured (normal fault) light gray (N7) tuff layer offset > 15 cm.</p> <p>Dip of layer 40°.</p> <p>Small graded beds offset by fractures.</p> <p>Medium dark gray (N4) disrupted vitric sandy siltstone grading to claystone.</p> <p>Mottled diatomaceous claystone.</p> <p>Ash lenses.</p> <p>Large burrow (Zoophycus)</p> <p>Medium light gray (N6) ash pockets.</p> <p>Pumice fragment.</p>																																																								
			<i>S. langii</i>																																																												
			<i>D. kamtschatica</i> (c)																																																												
							<p>TOTAL C - ORGANIC C - CARBONATE (%) 1, 107 cm (0.9, 0.9, 0) 2, 103 cm (0.9, 0.9, 0)</p> <p>CARBONATE BOMB 1, 4.6 cm = 0%</p> <p>SMEAR SLIDE SUMMARY</p> <table border="1"> <thead> <tr> <th></th> <th>1-23</th> <th>1-83</th> <th>1-123</th> </tr> </thead> <tbody> <tr> <td>Sand</td> <td>30</td> <td>30</td> <td>1</td> </tr> <tr> <td>Silt</td> <td>40</td> <td>40</td> <td>25</td> </tr> <tr> <td>Clay</td> <td>10</td> <td>30</td> <td>74</td> </tr> <tr> <td>Quartz</td> <td>5</td> <td>32</td> <td>29</td> </tr> <tr> <td>Feldspar</td> <td>10</td> <td>10</td> <td>2</td> </tr> <tr> <td>Heavy minerals</td> <td>TR</td> <td>1</td> <td>1</td> </tr> <tr> <td>Clay minerals</td> <td>10</td> <td>25</td> <td>72</td> </tr> <tr> <td>Volcanic glass</td> <td>75</td> <td>25</td> <td>5</td> </tr> <tr> <td>Palagonite</td> <td>-</td> <td>TR</td> <td>TR</td> </tr> <tr> <td>Glaucconite</td> <td>-</td> <td>2</td> <td>1</td> </tr> <tr> <td>Pyrite</td> <td>1</td> <td>1</td> <td>-</td> </tr> <tr> <td>Diatoms</td> <td>TR</td> <td>5</td> <td>15</td> </tr> <tr> <td>Sponge spicules</td> <td>1</td> <td>-</td> <td>1</td> </tr> </tbody> </table>		1-23	1-83	1-123	Sand	30	30	1	Silt	40	40	25	Clay	10	30	74	Quartz	5	32	29	Feldspar	10	10	2	Heavy minerals	TR	1	1	Clay minerals	10	25	72	Volcanic glass	75	25	5	Palagonite	-	TR	TR	Glaucconite	-	2	1	Pyrite	1	1	-	Diatoms	TR	5	15	Sponge spicules	1	-	1
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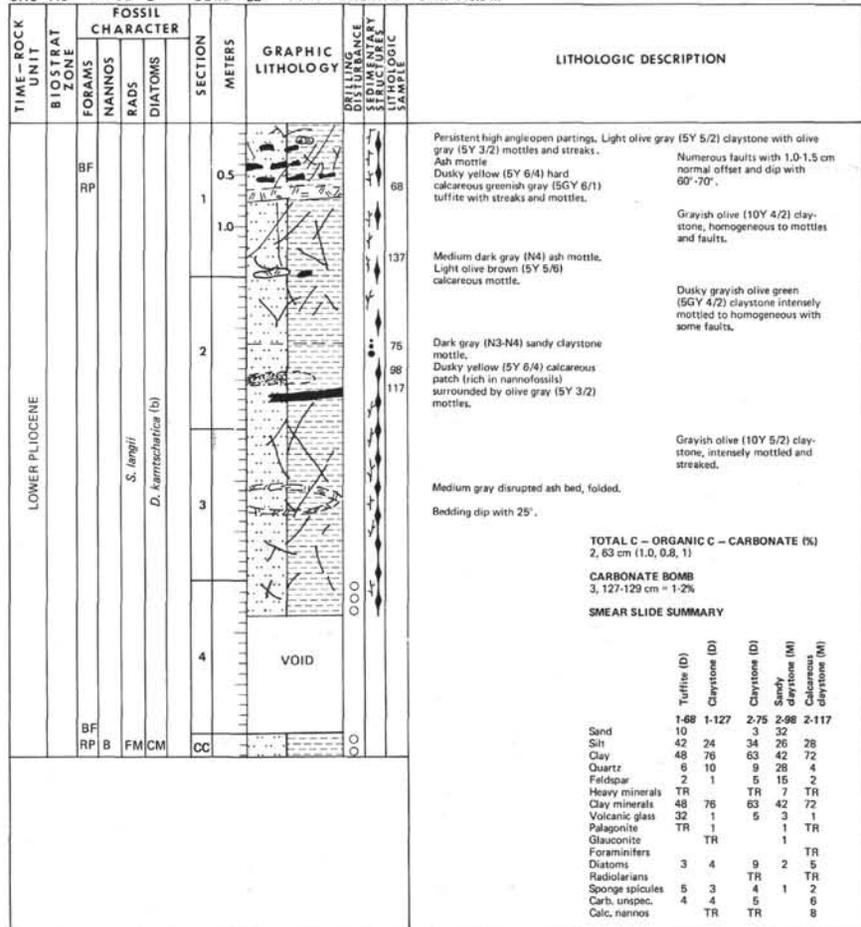
SITE 440		HOLE B		CORE 46		CORED INTERVAL: 567.0-576.5 m																															
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION																														
		FORAMS	NANNOS	RADS				DIATOMS																													
LOWER PLOCIENE		BF RM	CM	FP	CM	CC	<p>Silty pocket</p> <p>Fine olive (5Y 5/3) and dark gray (5Y 4/1) laminations cut by fracture.</p> <p>Pumice fragments.</p> <p>Silty ash pocket.</p> <p>Net of conjugate fractures, open, with possible horizontal slickensides(?)</p> <p>Diatomaceous claystone.</p> <p>Contorted olive (5Y 5/4) clayey diatomite bed.</p> <p>Dark olive gray mottled claystone layer.</p> <p>Ash pocket.</p> <p>Highly fractured (healed fractures and open partings).</p> <p>Medium gray (N5) silty pocket.</p> <p>Offset of silty bed along faults.</p> <p>Highly fractured, moderate mottling.</p>																														
			<i>S. langii</i>																																		
			<i>D. kamtschatica</i> (b)																																		
							<p>TOTAL C - ORGANIC C - CARBONATE (%) 1, 103 cm (1.2, 1.1, 1) 2, 77 cm (1.4, 1.3, 1)</p> <p>CARBONATE BOMB 3, 34-37 cm = 1-2%</p> <p>SMEAR SLIDE SUMMARY</p> <table border="1"> <thead> <tr> <th></th> <th>2-51</th> <th>2-60</th> </tr> </thead> <tbody> <tr> <td>Sand</td> <td>0</td> <td>0</td> </tr> <tr> <td>Silt</td> <td>25</td> <td>30</td> </tr> <tr> <td>Clay</td> <td>75</td> <td>70</td> </tr> <tr> <td>Quartz</td> <td>5</td> <td>4</td> </tr> <tr> <td>Feldspar</td> <td>2</td> <td>1</td> </tr> <tr> <td>Clay minerals</td> <td>72</td> <td>52</td> </tr> <tr> <td>Volcanic glass</td> <td>5</td> <td>3</td> </tr> <tr> <td>Diatoms</td> <td>15</td> <td>40</td> </tr> <tr> <td>Sponge spicules</td> <td>1</td> <td>TR</td> </tr> </tbody> </table>		2-51	2-60	Sand	0	0	Silt	25	30	Clay	75	70	Quartz	5	4	Feldspar	2	1	Clay minerals	72	52	Volcanic glass	5	3	Diatoms	15	40	Sponge spicules	1	TR
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Sponge spicules	1	TR																																			

SITE 440		HOLE B		CORE 47		CORED INTERVAL: 576.5-586.0 m																																											
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION																																										
		FORAMS	NANNOS	RADS				DIATOMS																																									
LOWER PLOCIENE		B	CM	FM	CM	CC	<p>Olive gray (5Y 5/3) laminations cut by normal fault.</p> <p>Diatomaceous claystone with 1 cm blocks with horizontal slickensides.</p> <p>Ash.</p> <p>Pumice fragment.</p> <p>High drilling disturbance. No structure visible.</p>																																										
			<i>S. langii</i>																																														
			<i>D. kamtschatica</i> (b)																																														
							<p>TOTAL C - ORGANIC C - CARBONATE (%) 1, 15 cm (1.2, 1.1, 1)</p> <p>CARBONATE BOMB 1, 79-81 cm = 2%</p> <p>SMEAR SLIDE SUMMARY</p> <table border="1"> <thead> <tr> <th></th> <th>1-70</th> <th>CC</th> </tr> </thead> <tbody> <tr> <td>Diatomaceous claystone (M)</td> <td>0</td> <td>-</td> </tr> <tr> <td>Pumice clast (M)</td> <td>25</td> <td>25</td> </tr> <tr> <td>Sand</td> <td>75</td> <td>75</td> </tr> <tr> <td>Silt</td> <td>25</td> <td>25</td> </tr> <tr> <td>Clay</td> <td>75</td> <td>70</td> </tr> <tr> <td>Quartz</td> <td>5</td> <td>2</td> </tr> <tr> <td>Feldspar</td> <td>2</td> <td>-</td> </tr> <tr> <td>Clay minerals</td> <td>71</td> <td>20</td> </tr> <tr> <td>Volcanic glass</td> <td>5</td> <td>74</td> </tr> <tr> <td>Palagonite</td> <td>-</td> <td>TR</td> </tr> <tr> <td>Pyrite</td> <td>1</td> <td>-</td> </tr> <tr> <td>Diatoms</td> <td>15</td> <td>3</td> </tr> <tr> <td>Sponge spicules</td> <td>1</td> <td>1</td> </tr> </tbody> </table>		1-70	CC	Diatomaceous claystone (M)	0	-	Pumice clast (M)	25	25	Sand	75	75	Silt	25	25	Clay	75	70	Quartz	5	2	Feldspar	2	-	Clay minerals	71	20	Volcanic glass	5	74	Palagonite	-	TR	Pyrite	1	-	Diatoms	15	3	Sponge spicules	1	1
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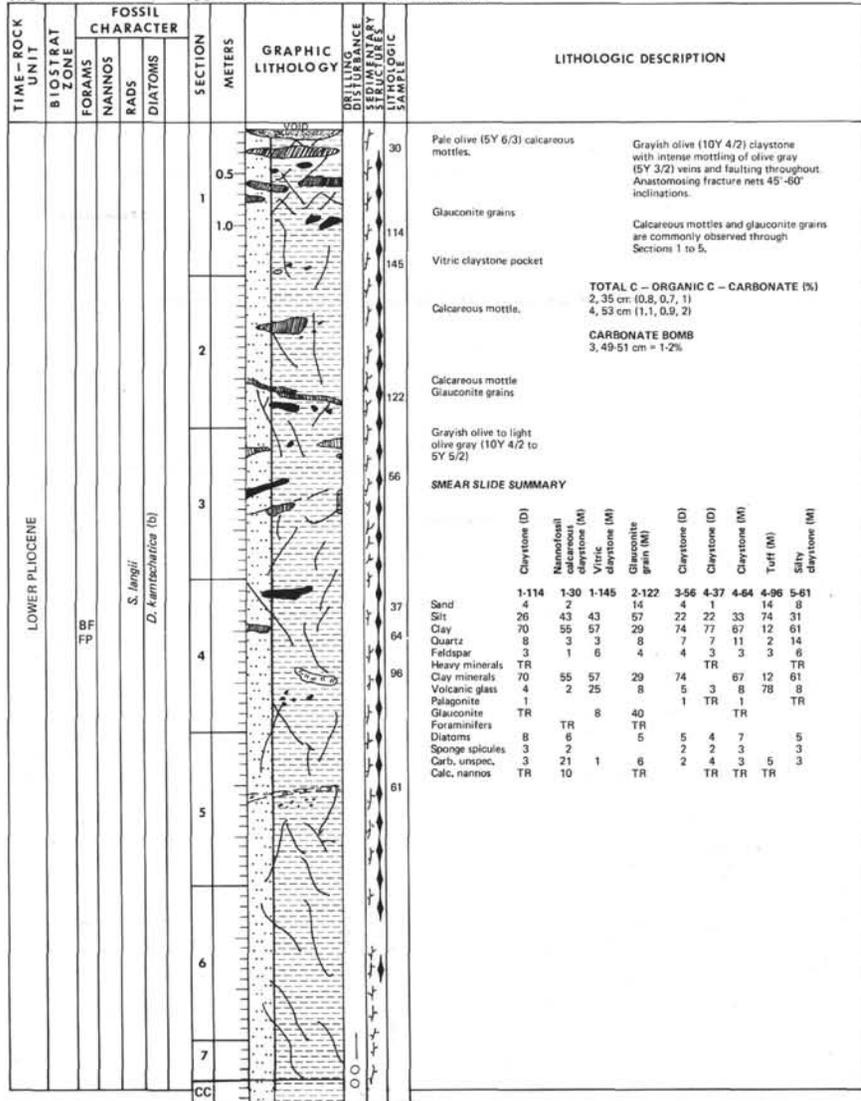


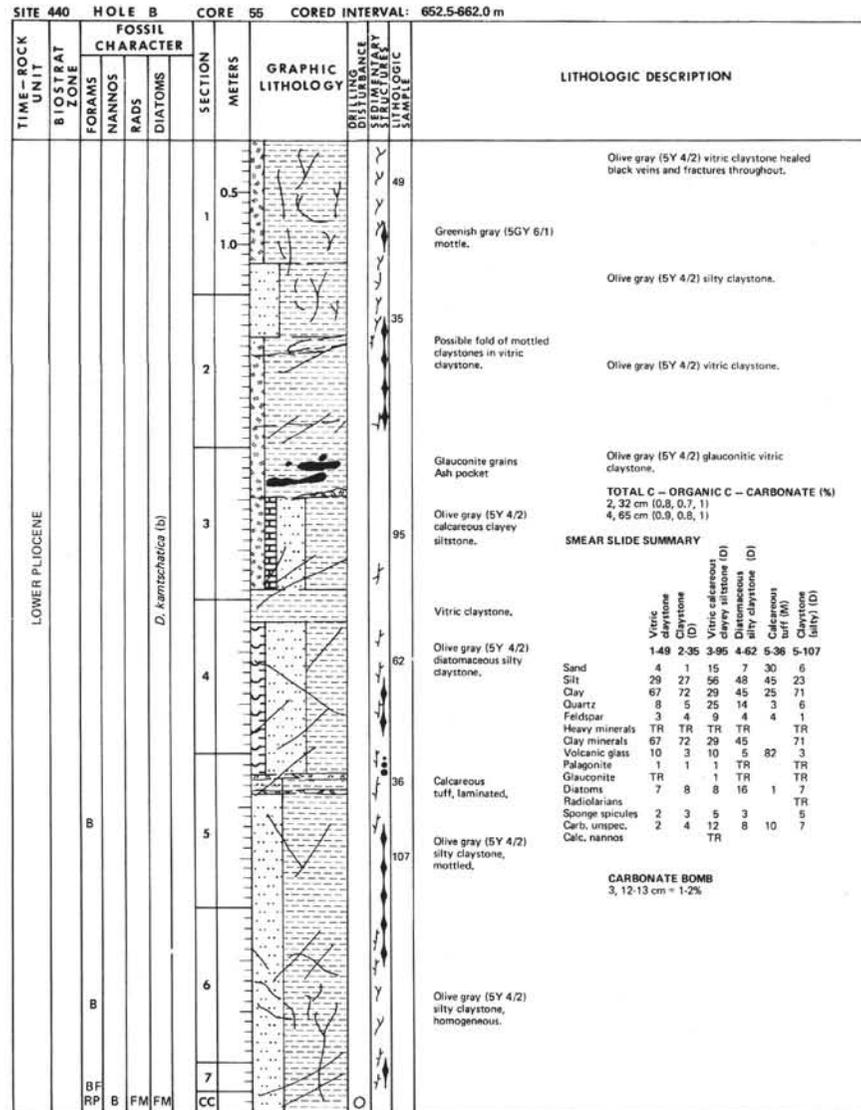
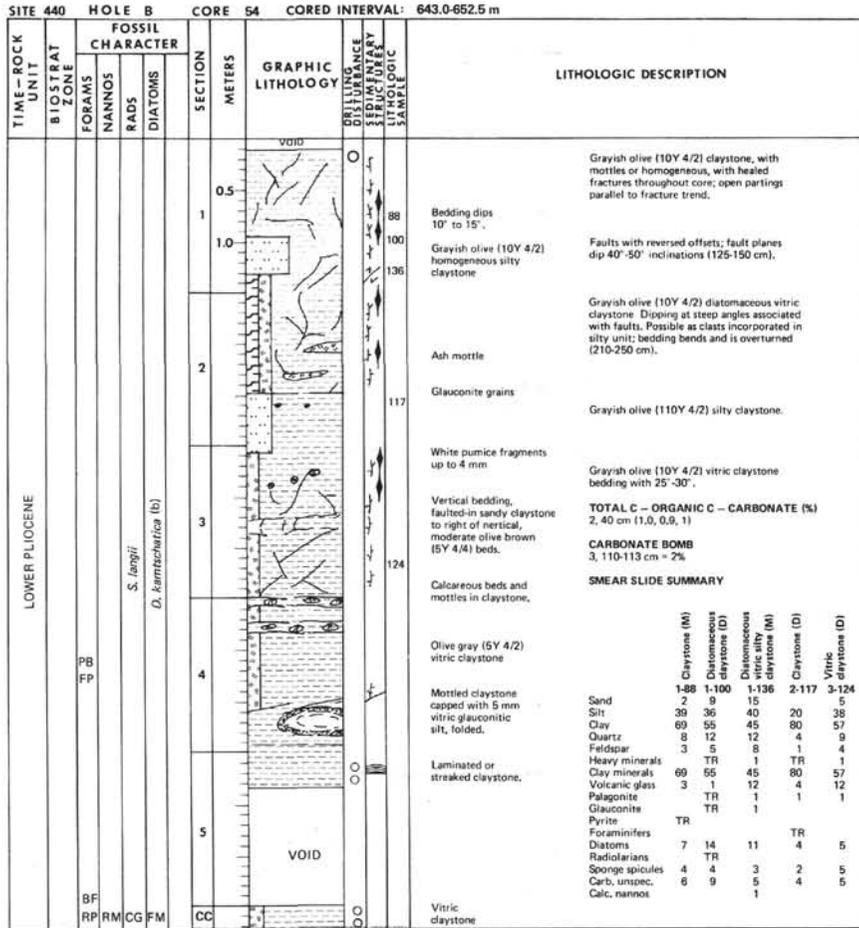


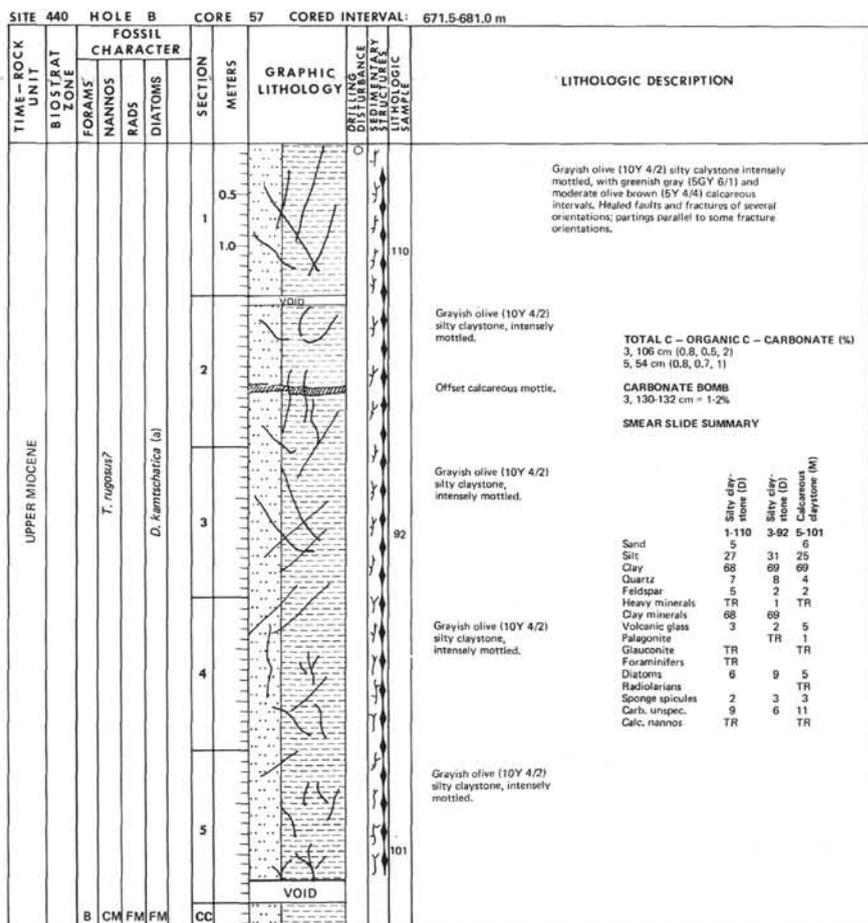
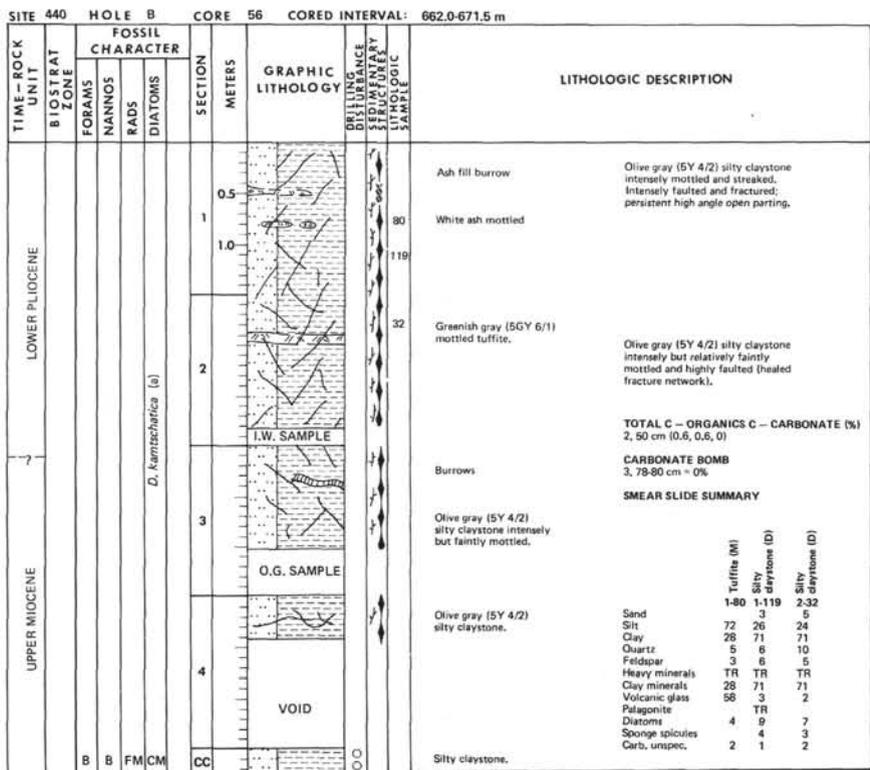
SITE 440 HOLE B CORE 52 CORED INTERVAL: 624.0-633.5 m

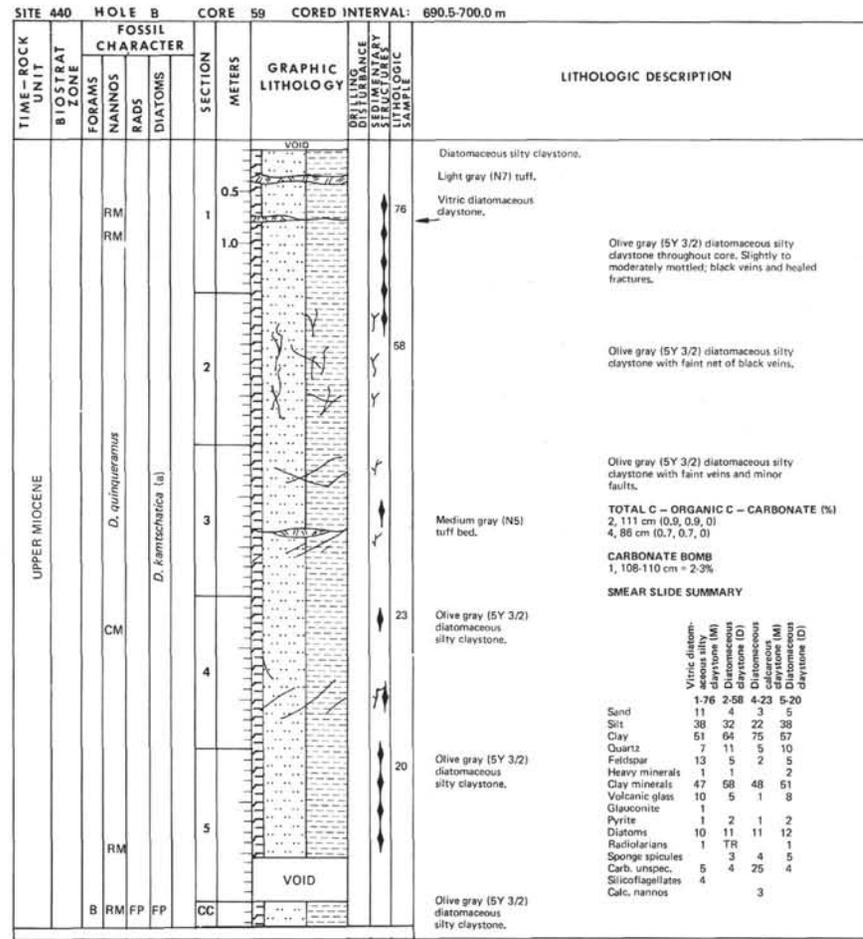
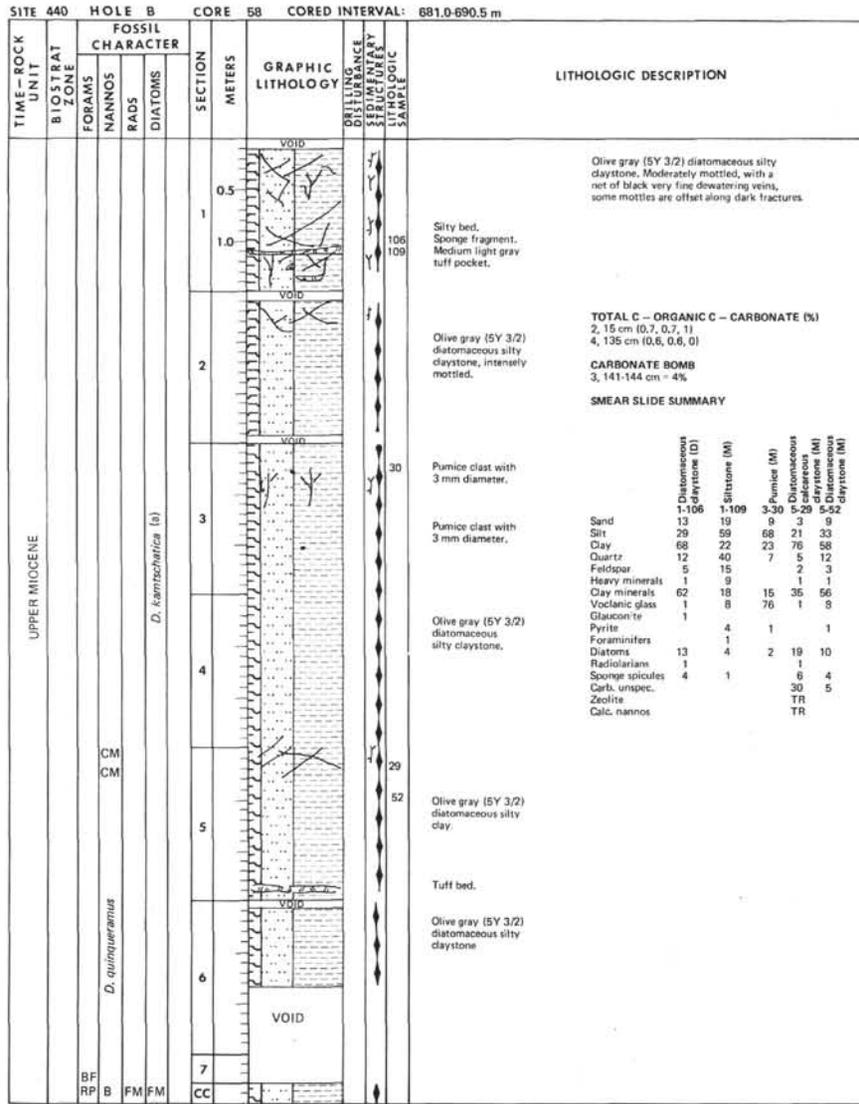


SITE 440 HOLE B CORE 53 CORED INTERVAL: 633.5-643.0 m

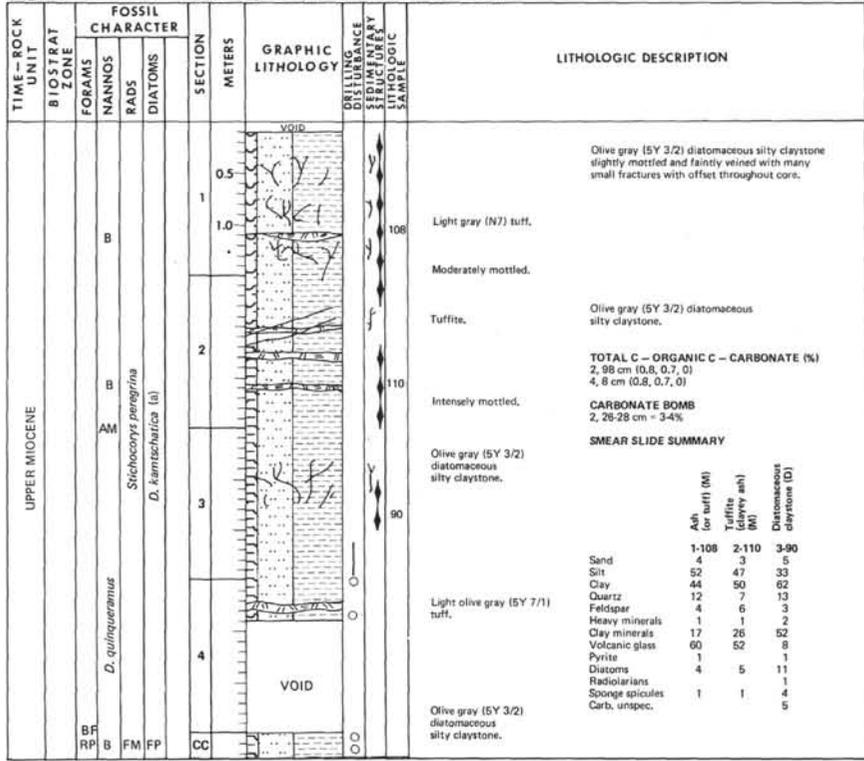




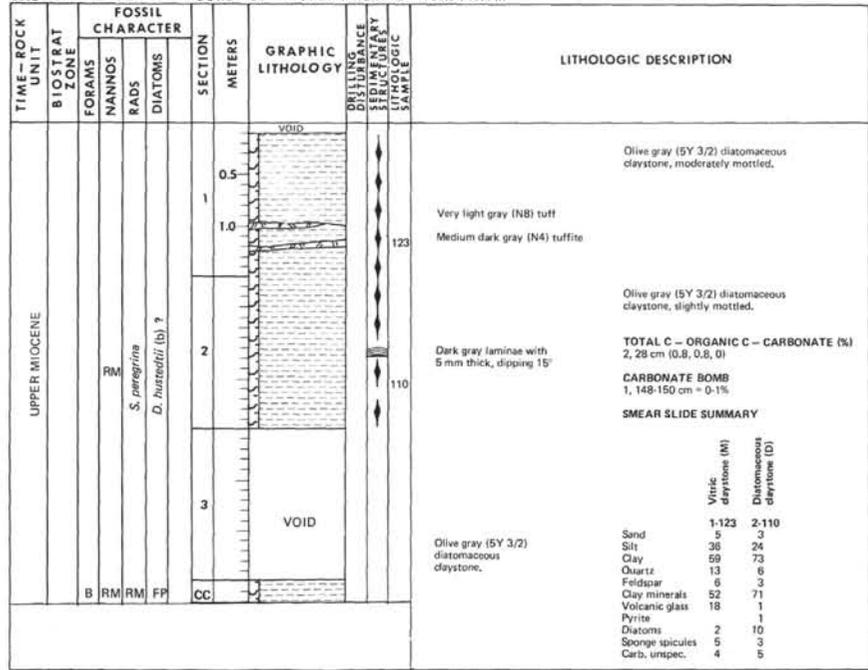


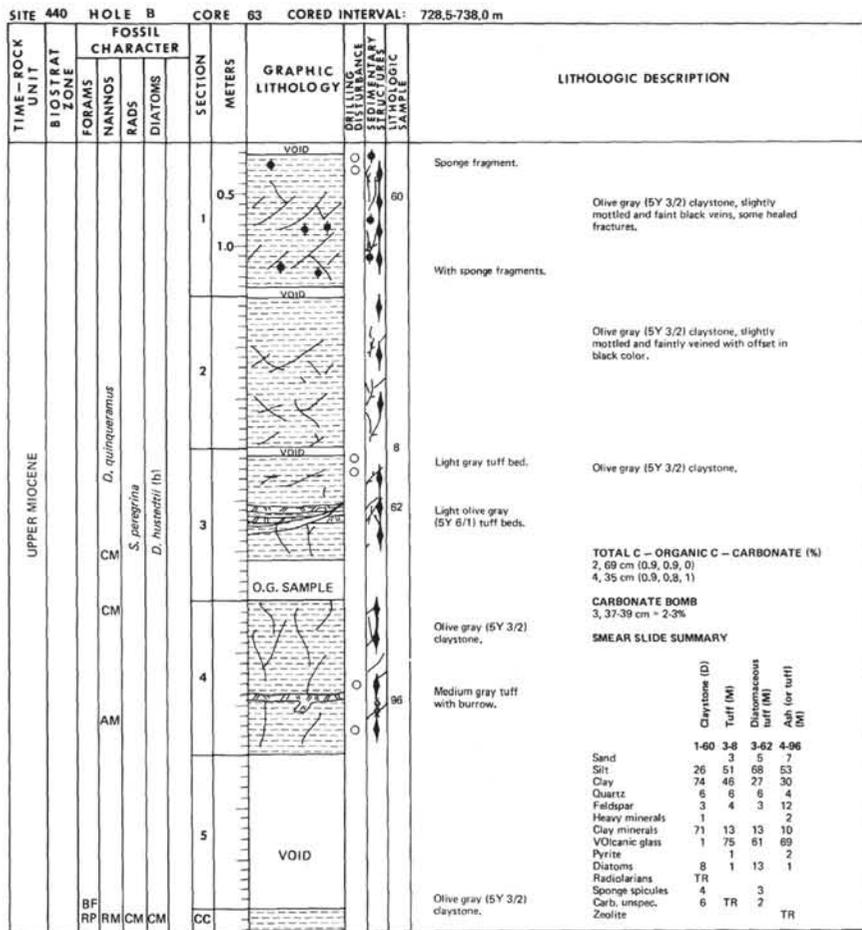
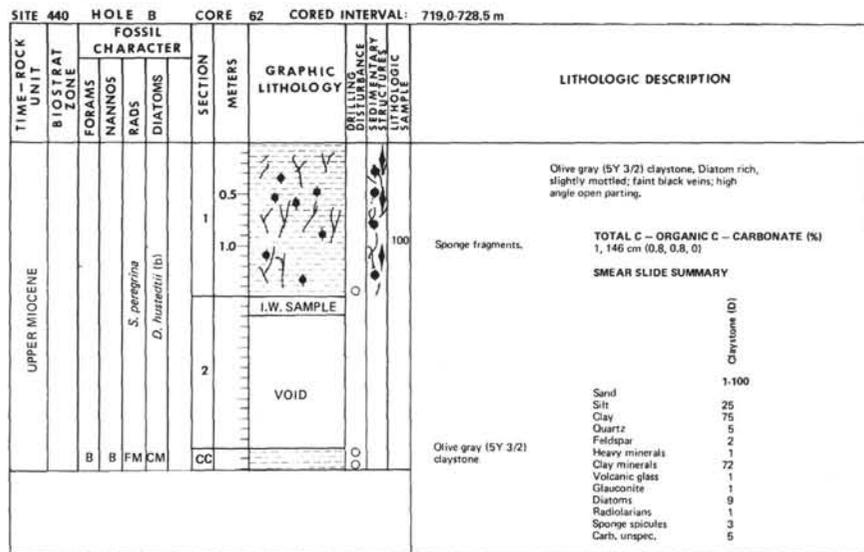


SITE 440 HOLE B CORE 60 CORED INTERVAL: 700.0-709.5 m

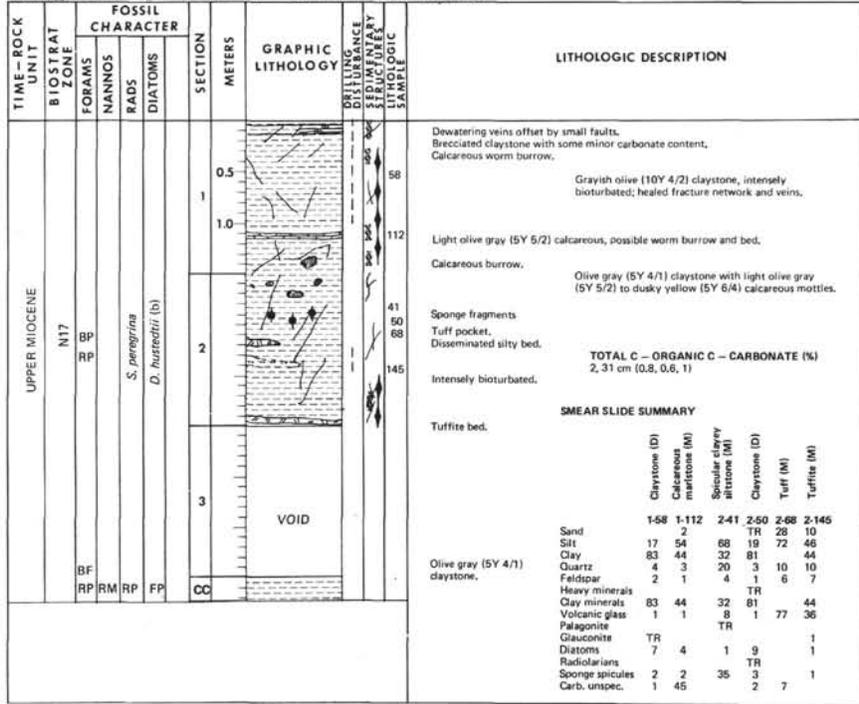


SITE 440 HOLE B CORE 61 CORED INTERVAL: 709.5-719.0 m

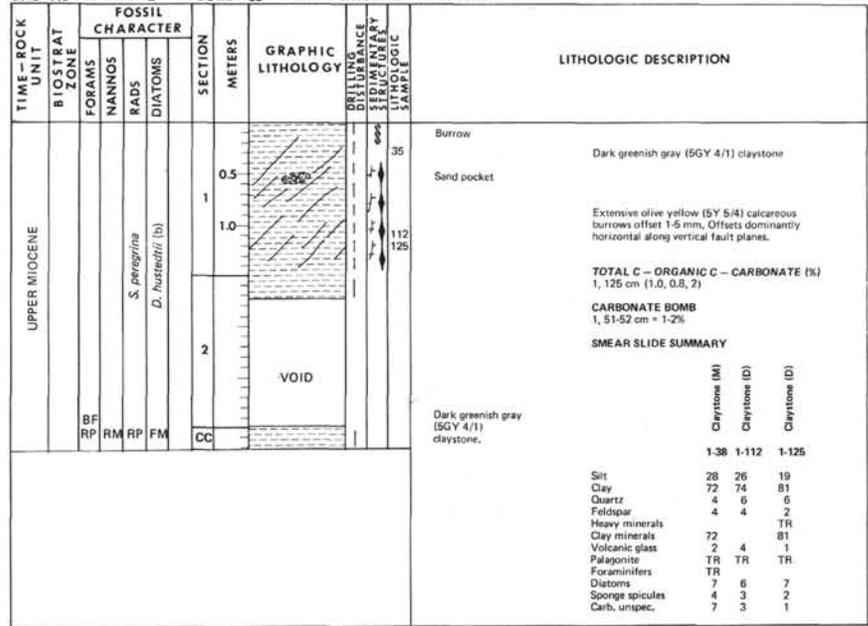


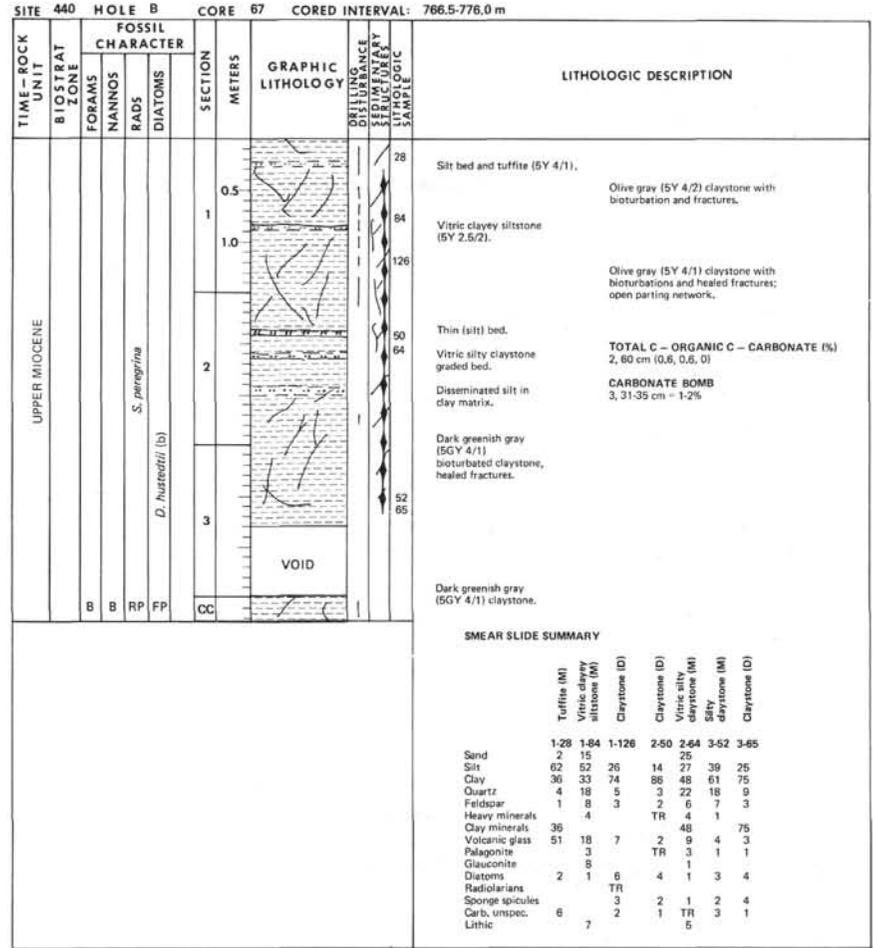
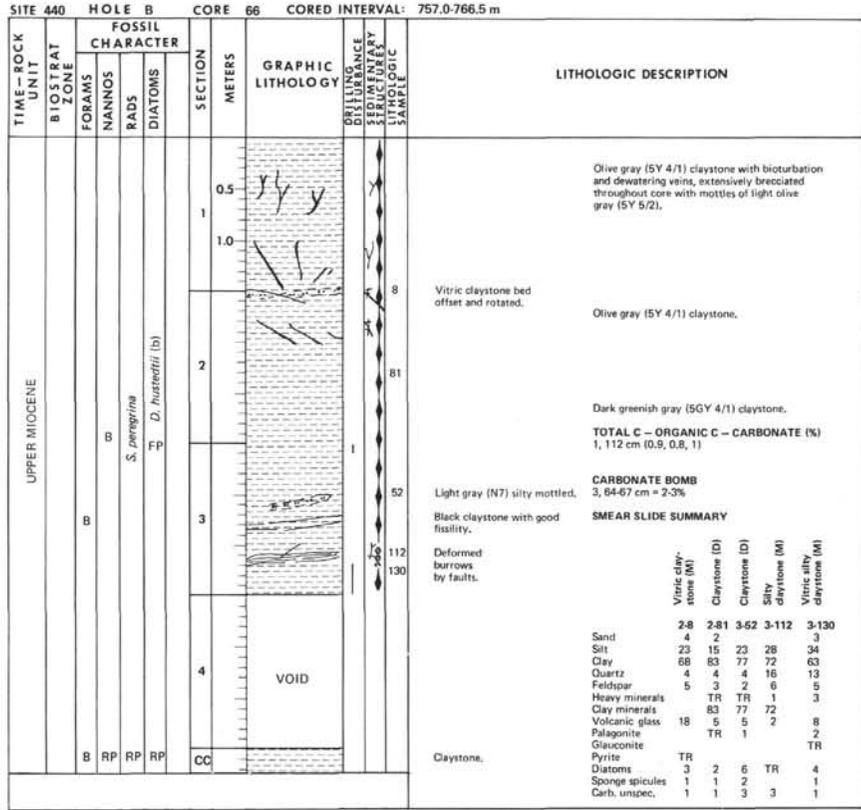


SITE 440 HOLE B CORE 64 CORED INTERVAL: 738.0-747.5 m

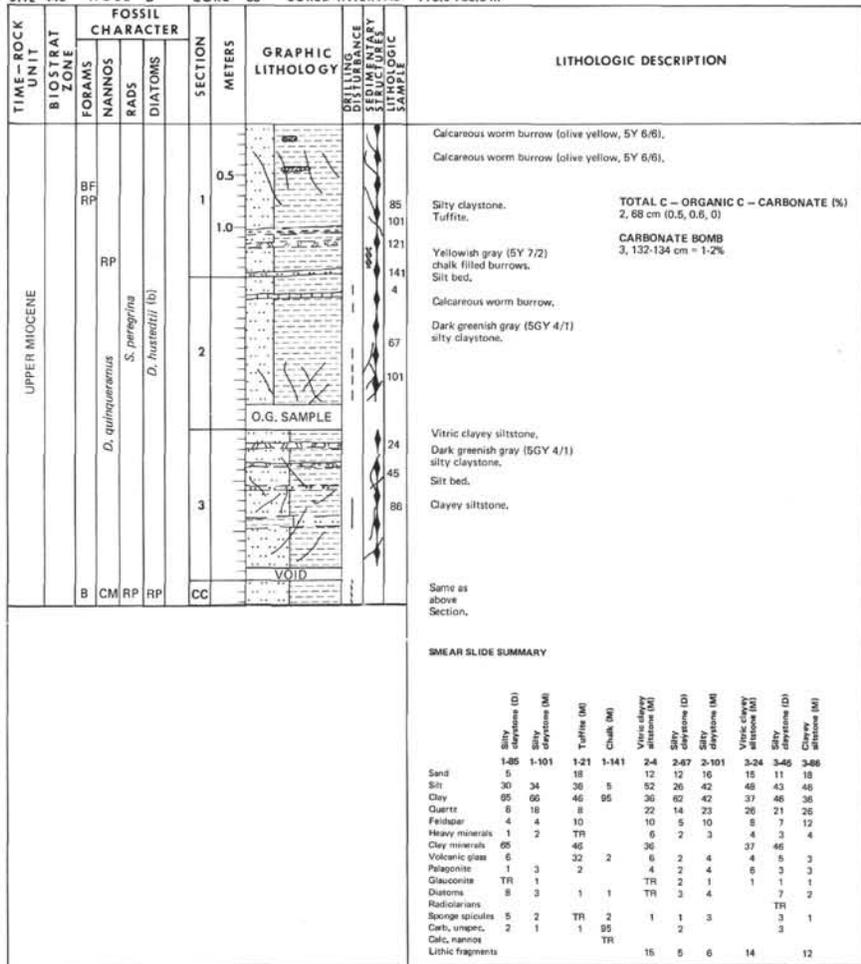


SITE 440 HOLE B CORE 65 CORED INTERVAL: 747.5-757.0 m

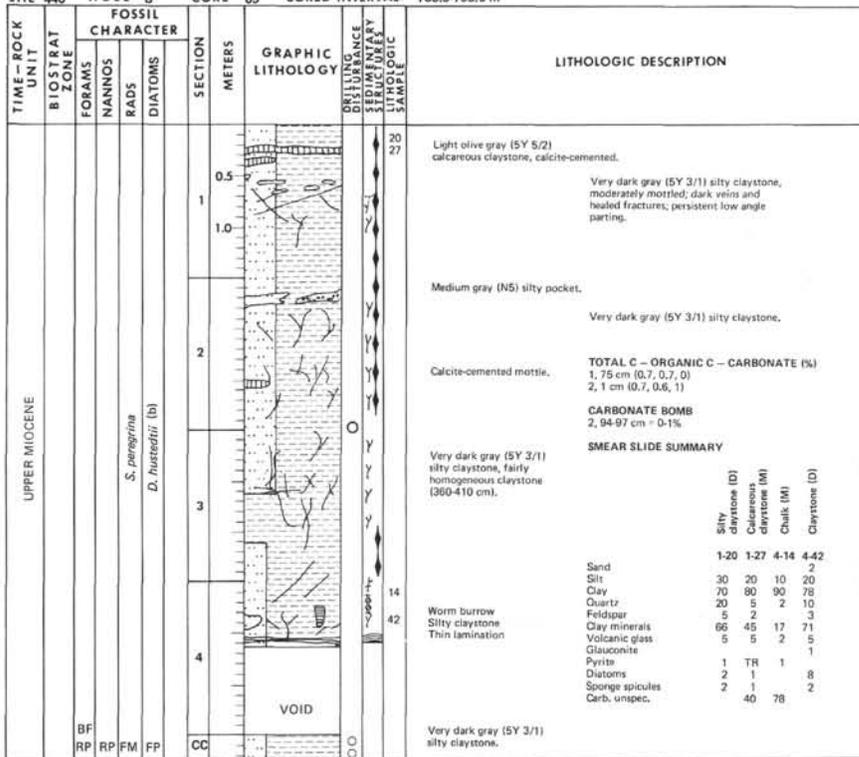


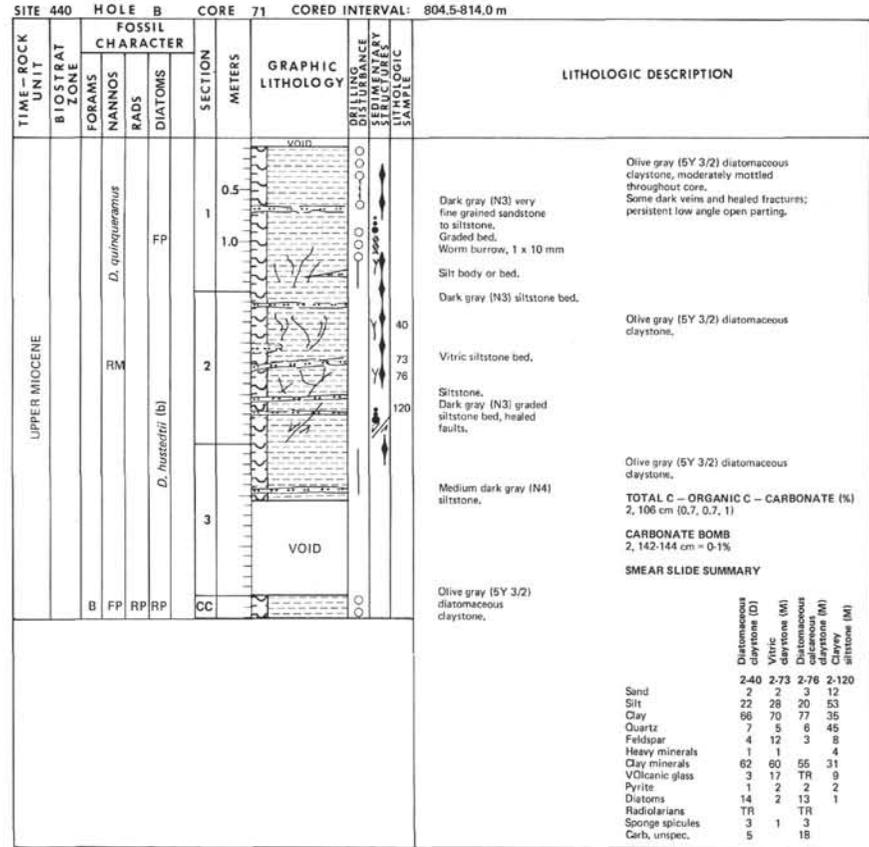
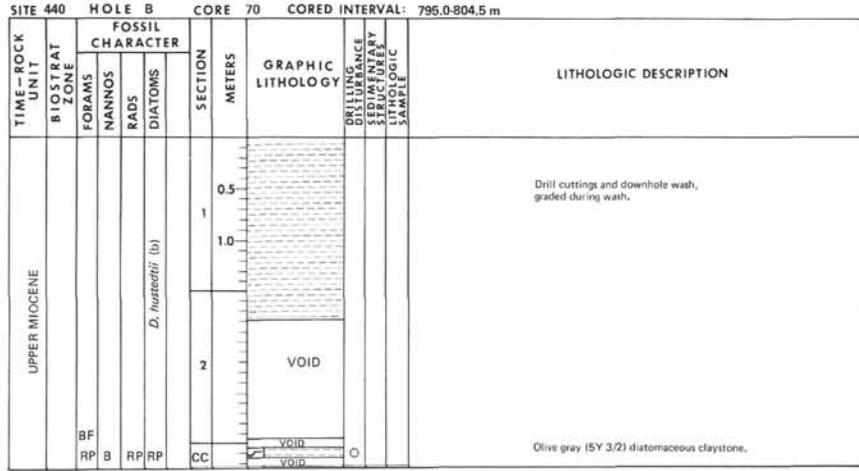


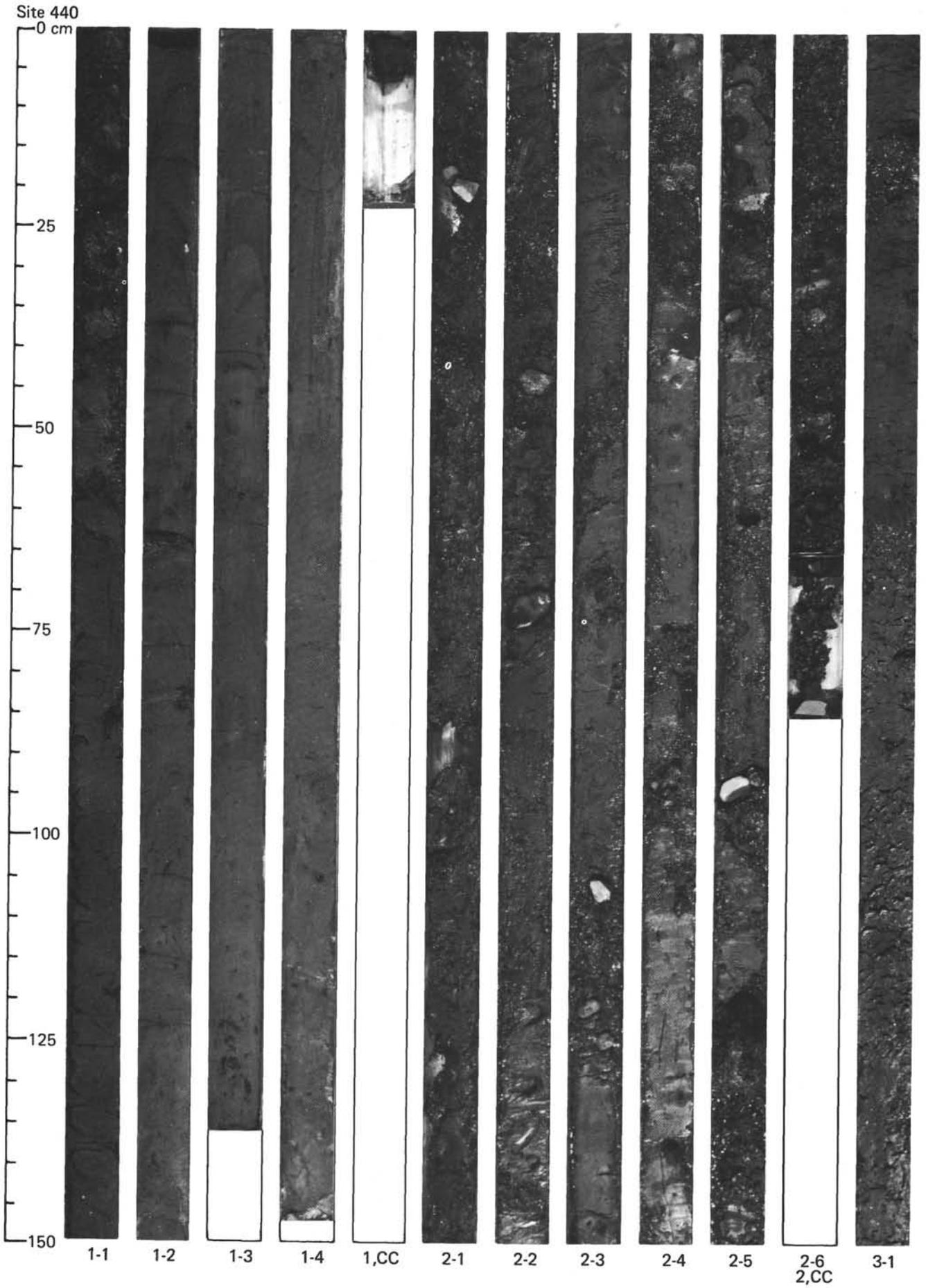
SITE 440 HOLE B CORE 68 CORED INTERVAL: 776.0-785.6 m



SITE 440 HOLE B CORE 69 CORED INTERVAL: 785.5-795.0 m

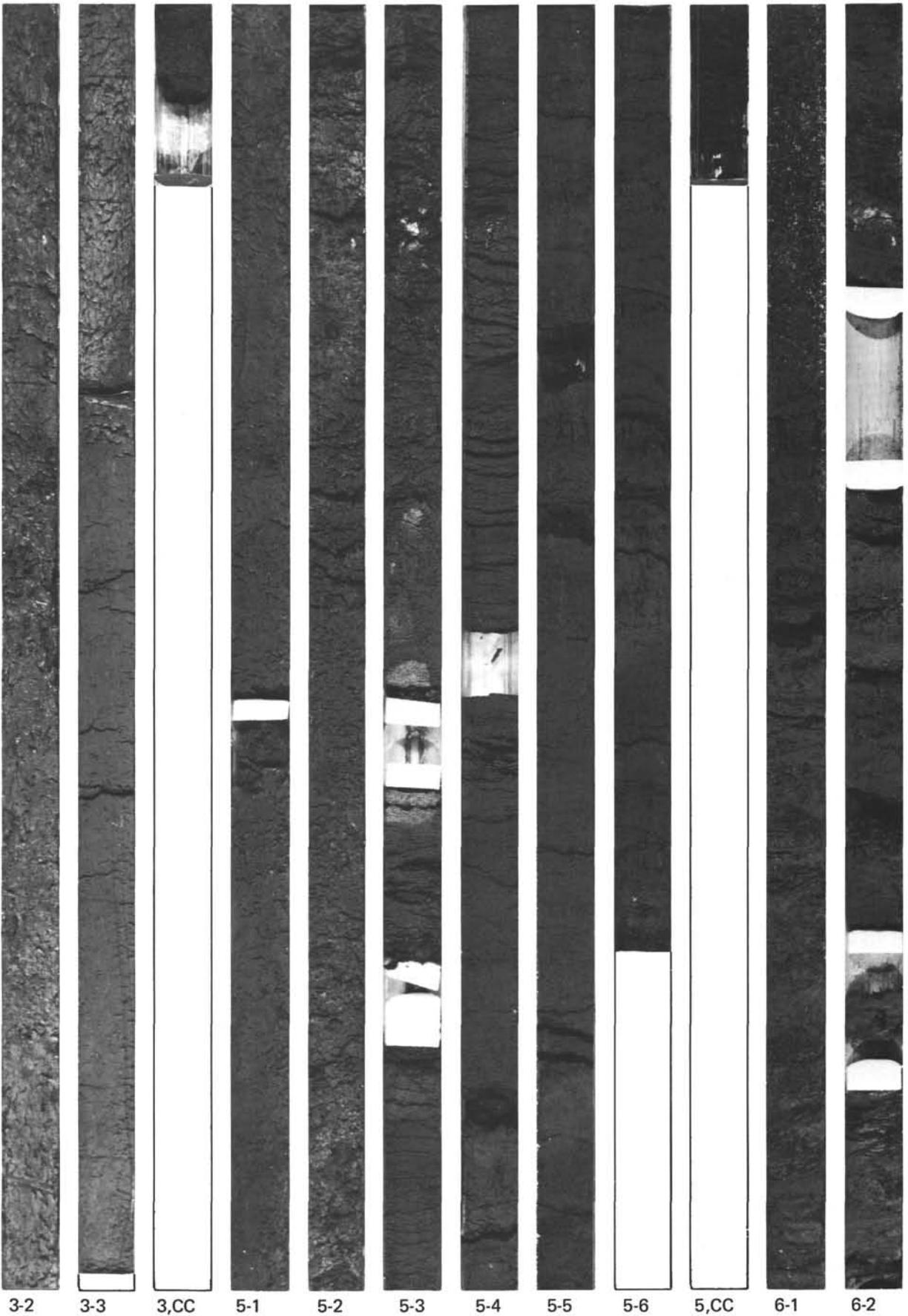


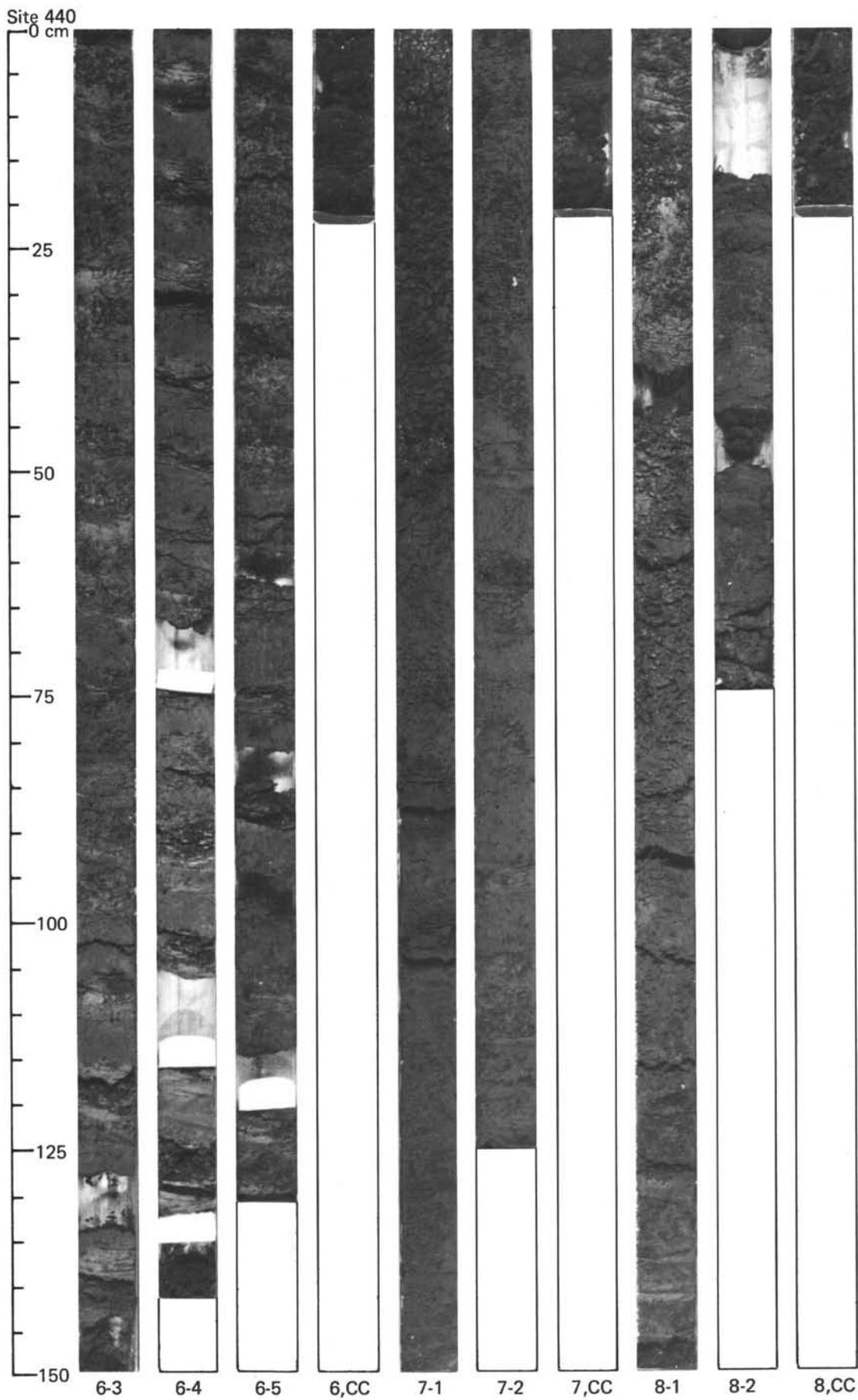


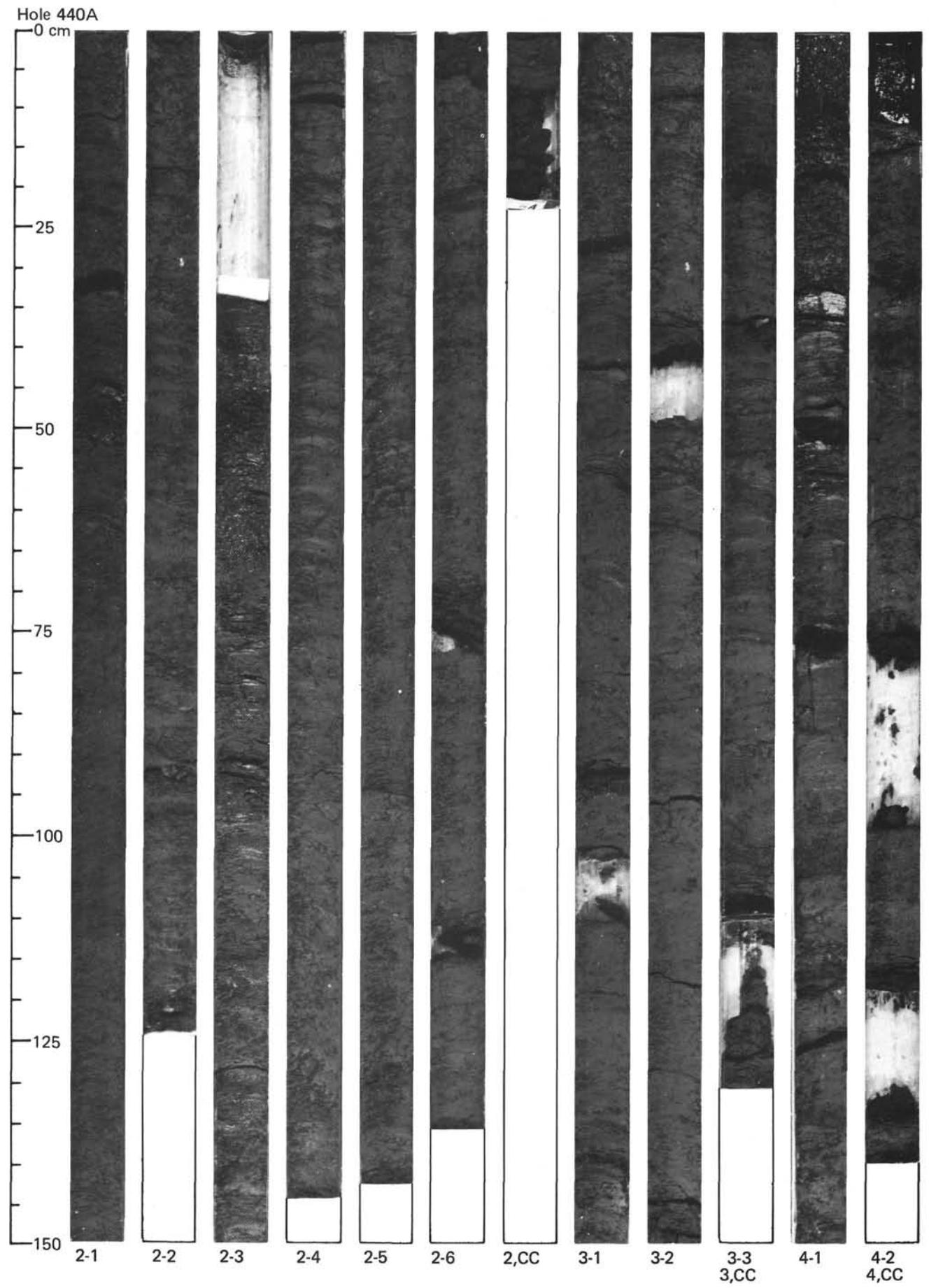


Site 440

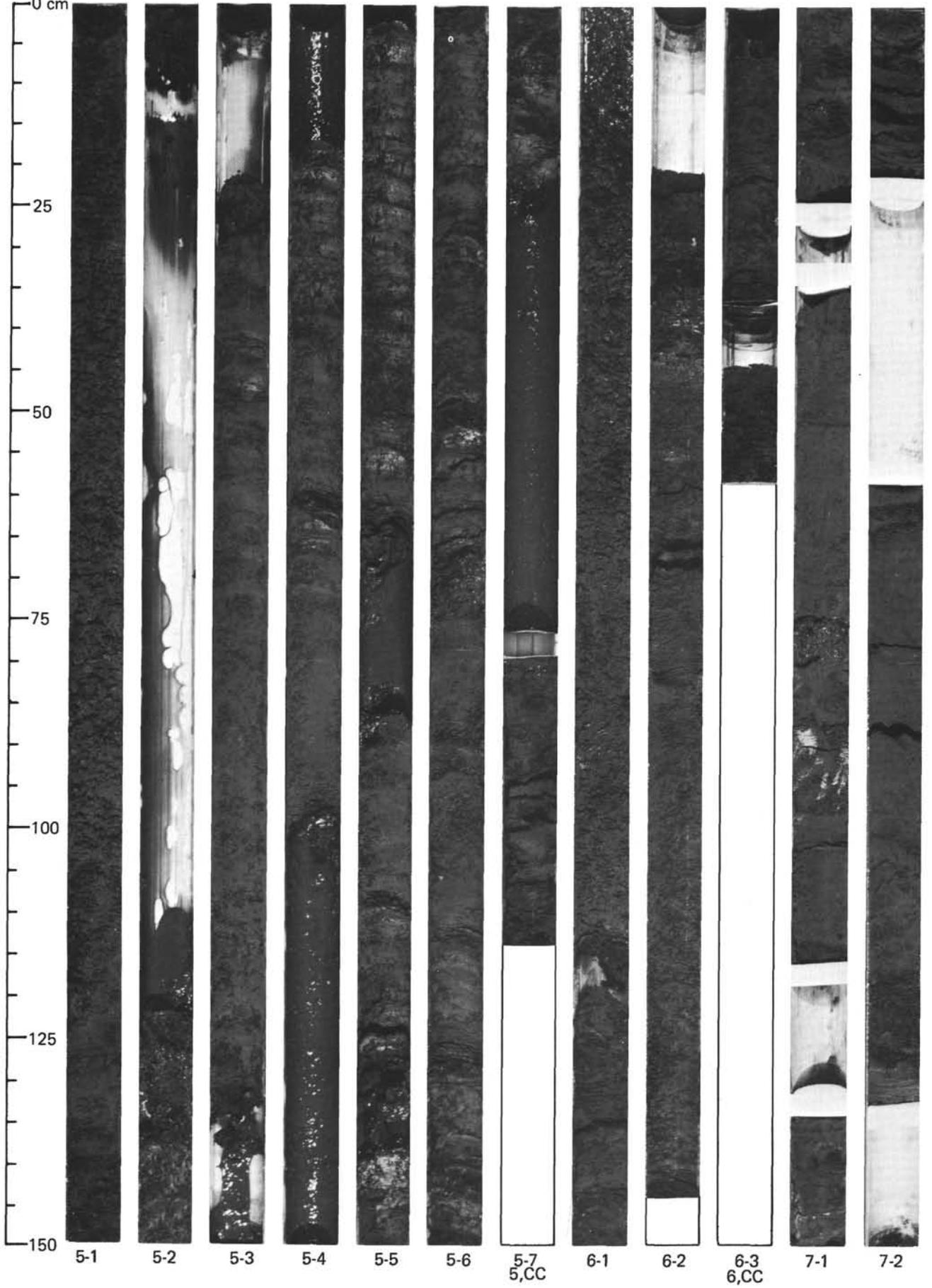
0 cm
25
50
75
100
125
150

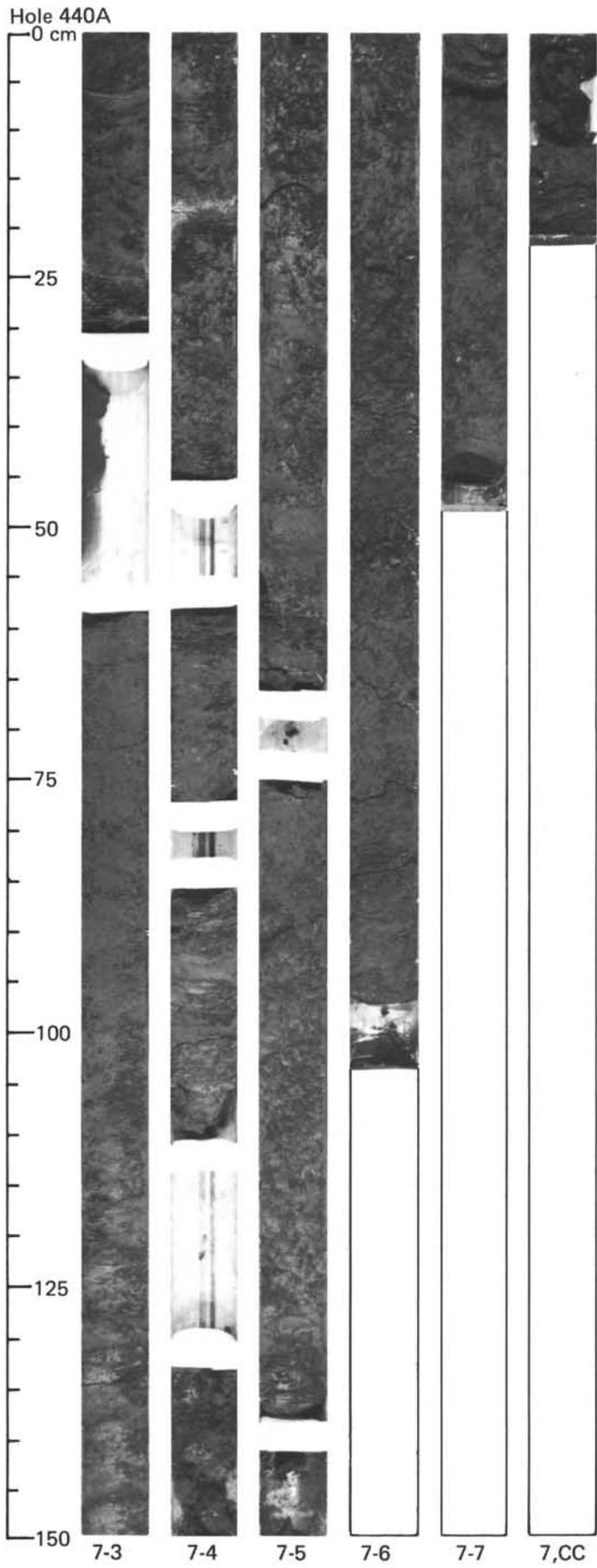




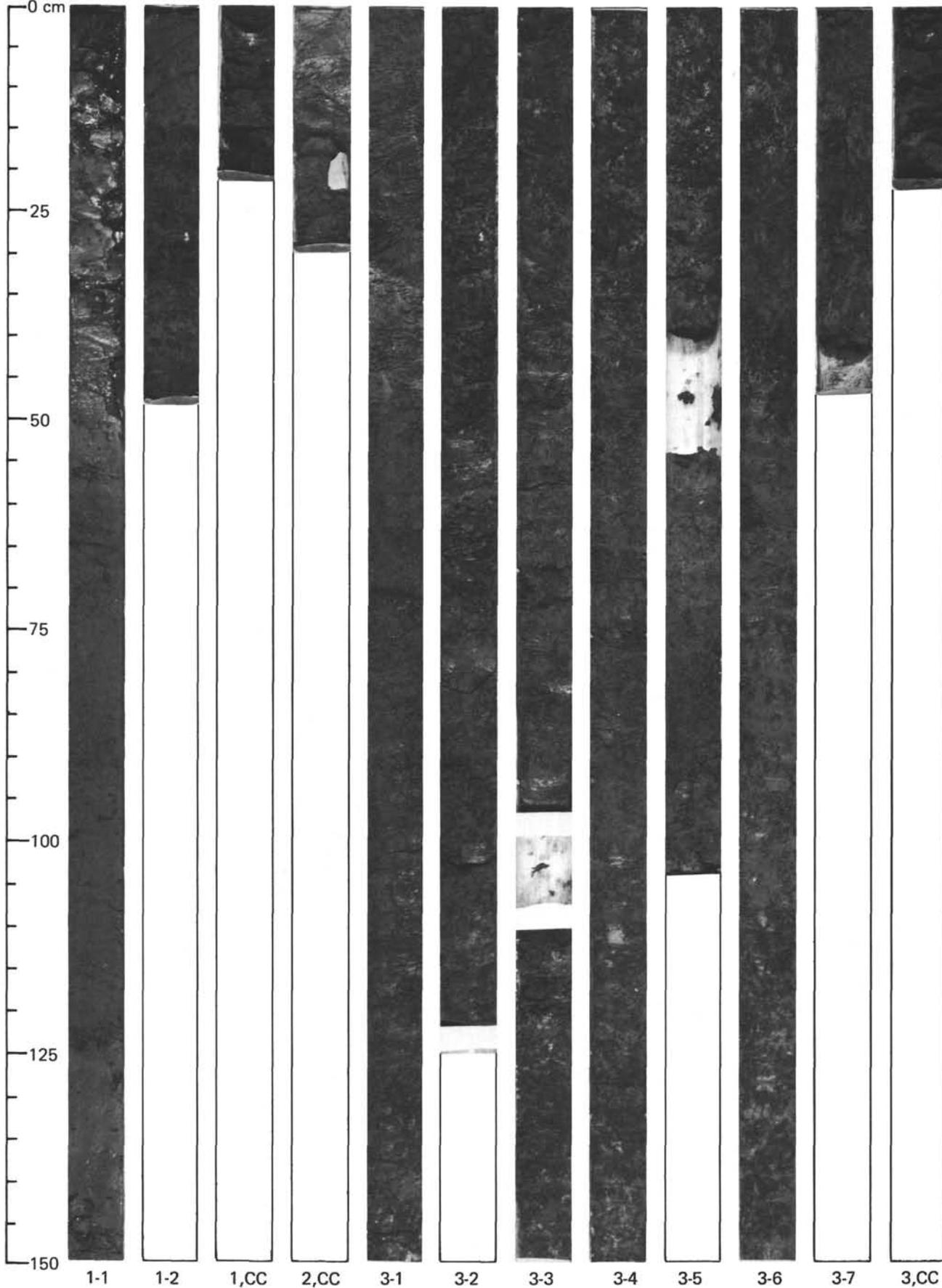


Hole 440A

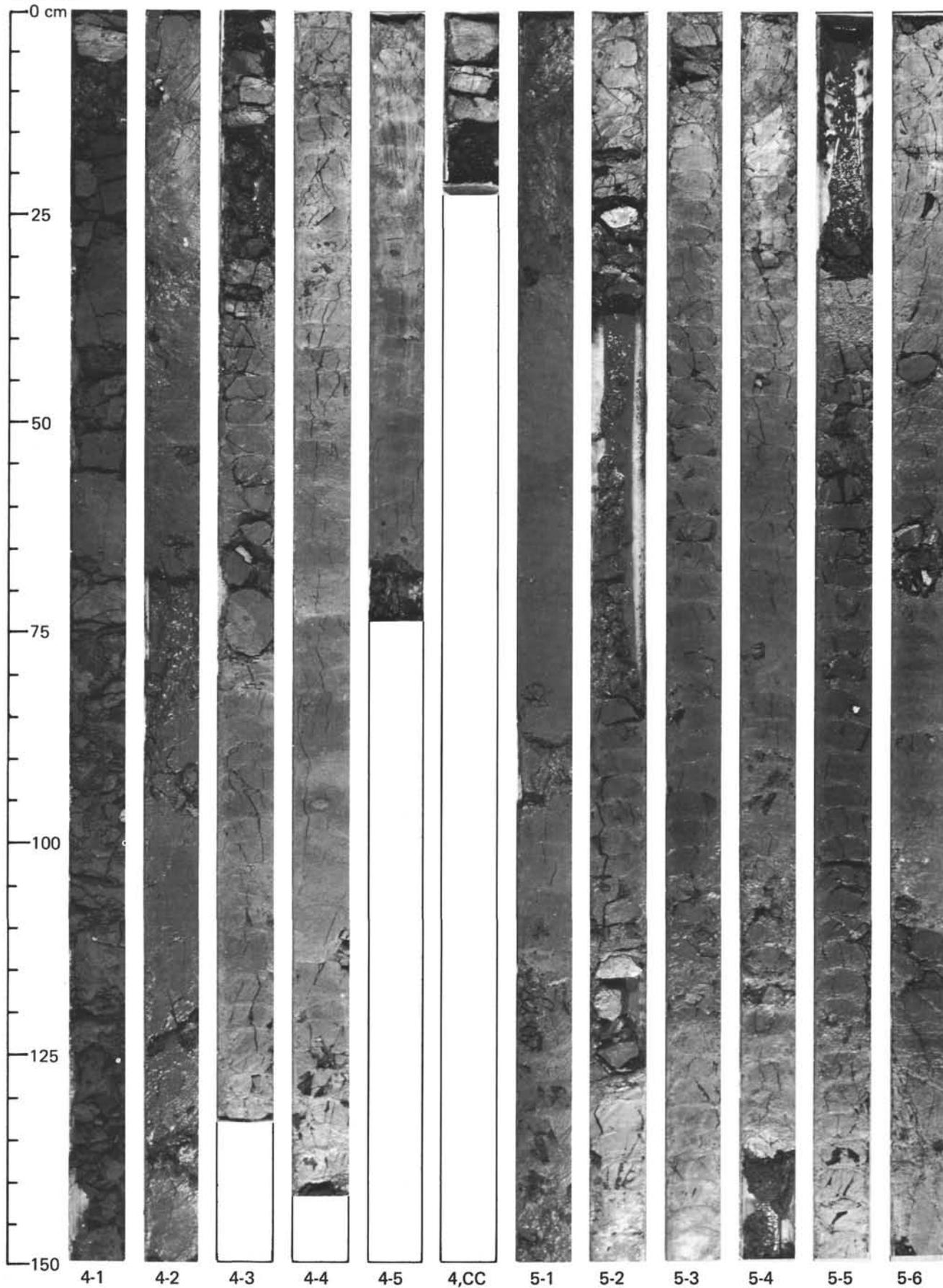




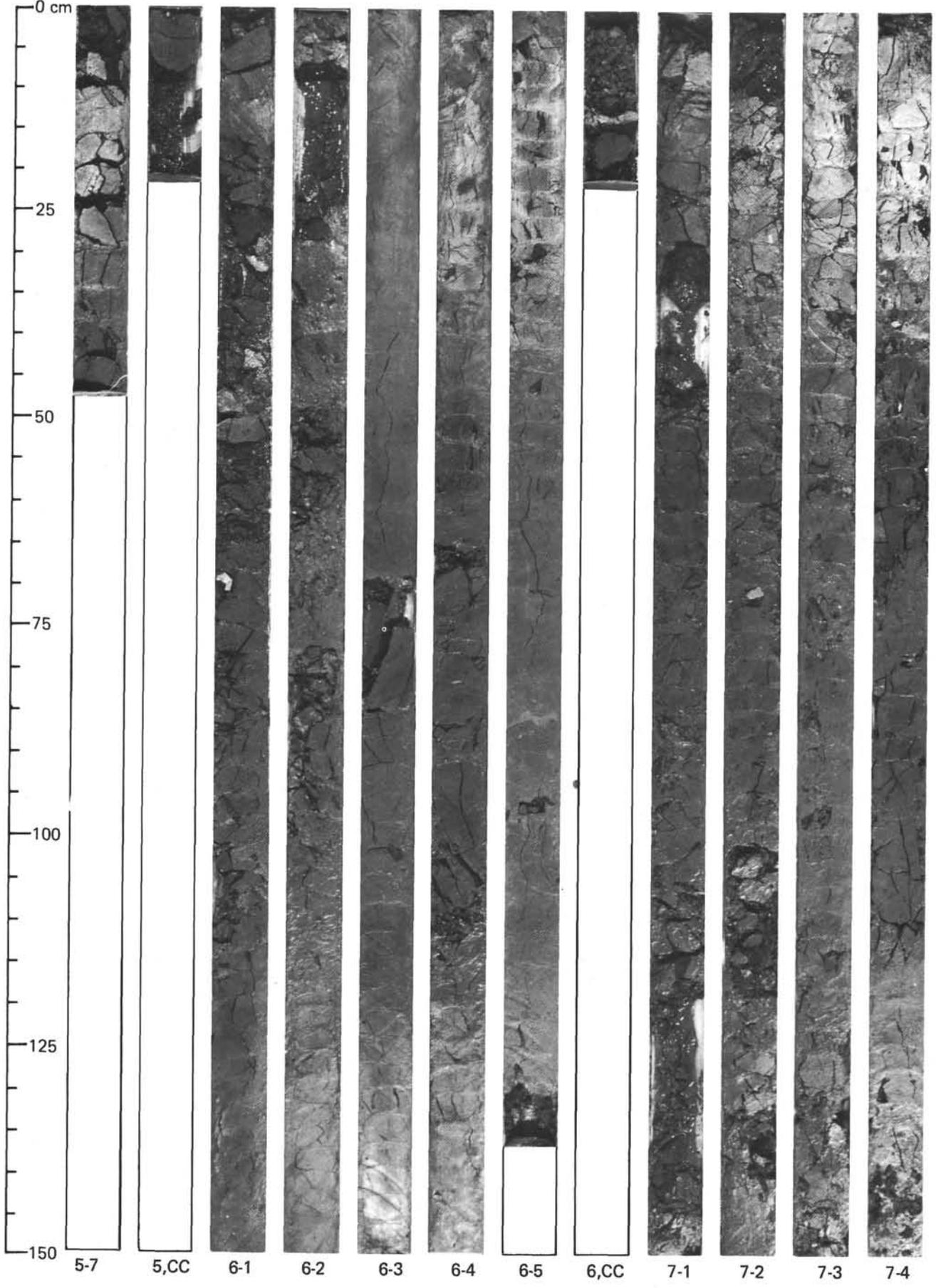
Hole 440B

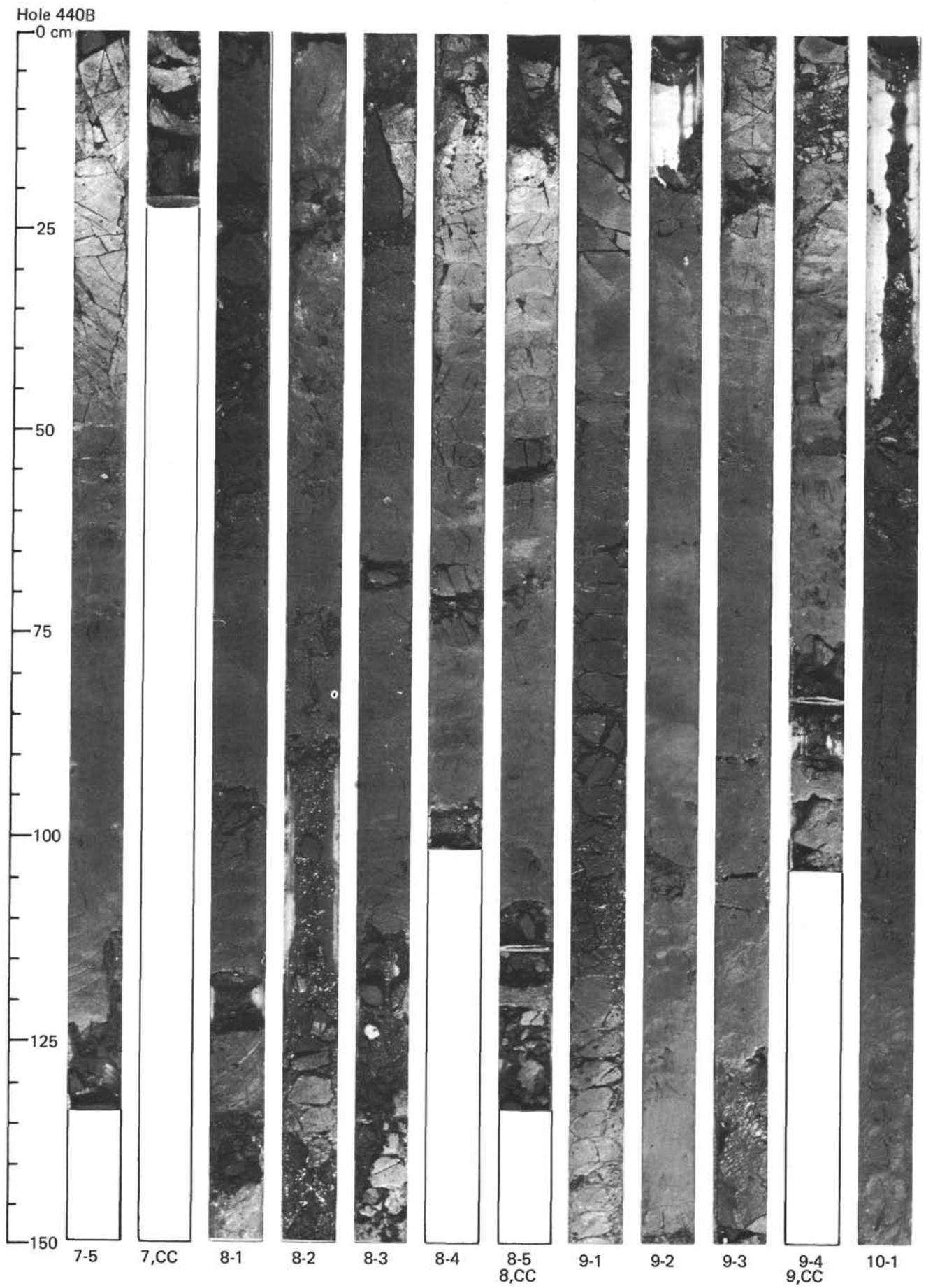


Hole 440B

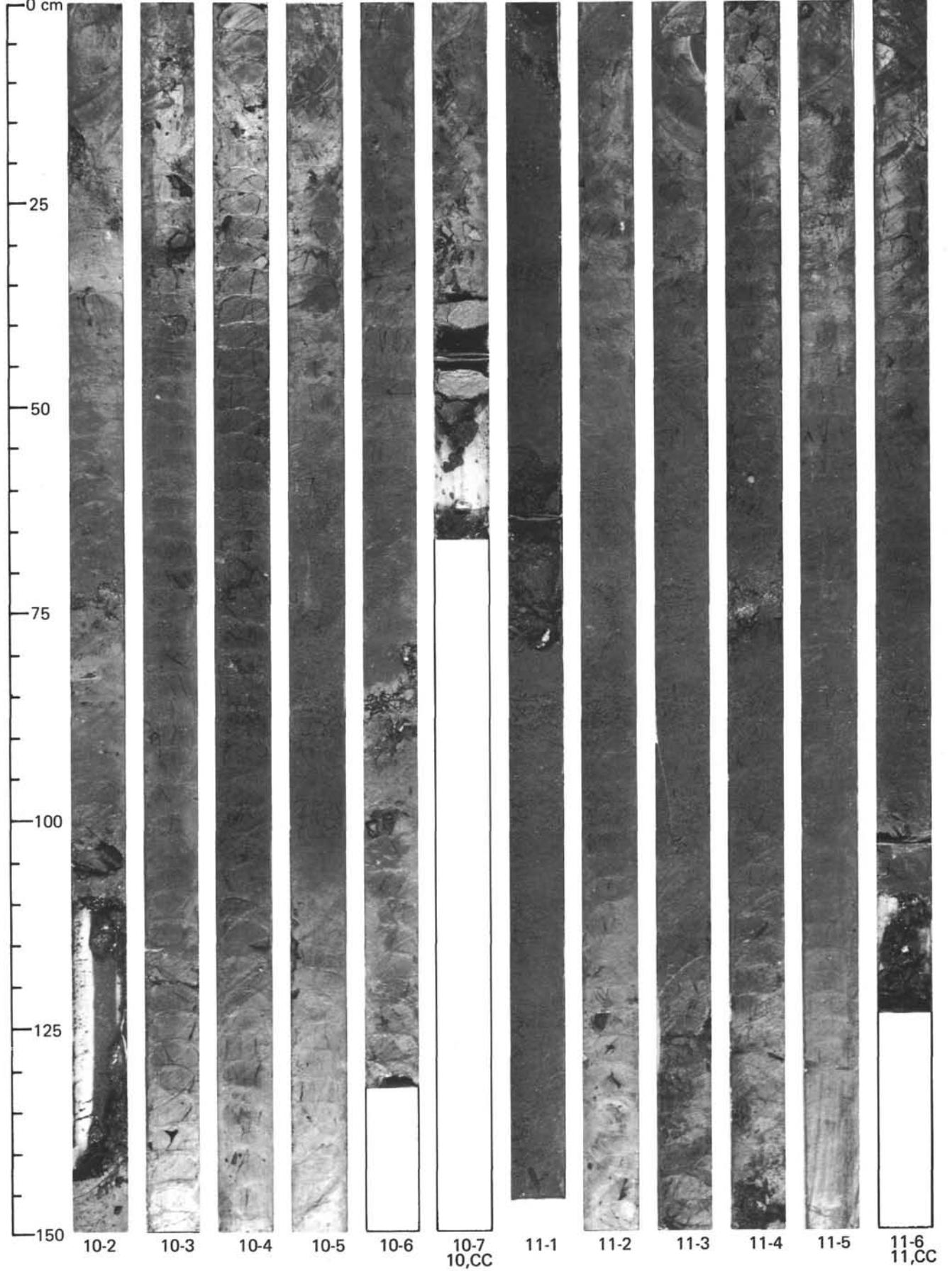


Hole 440B

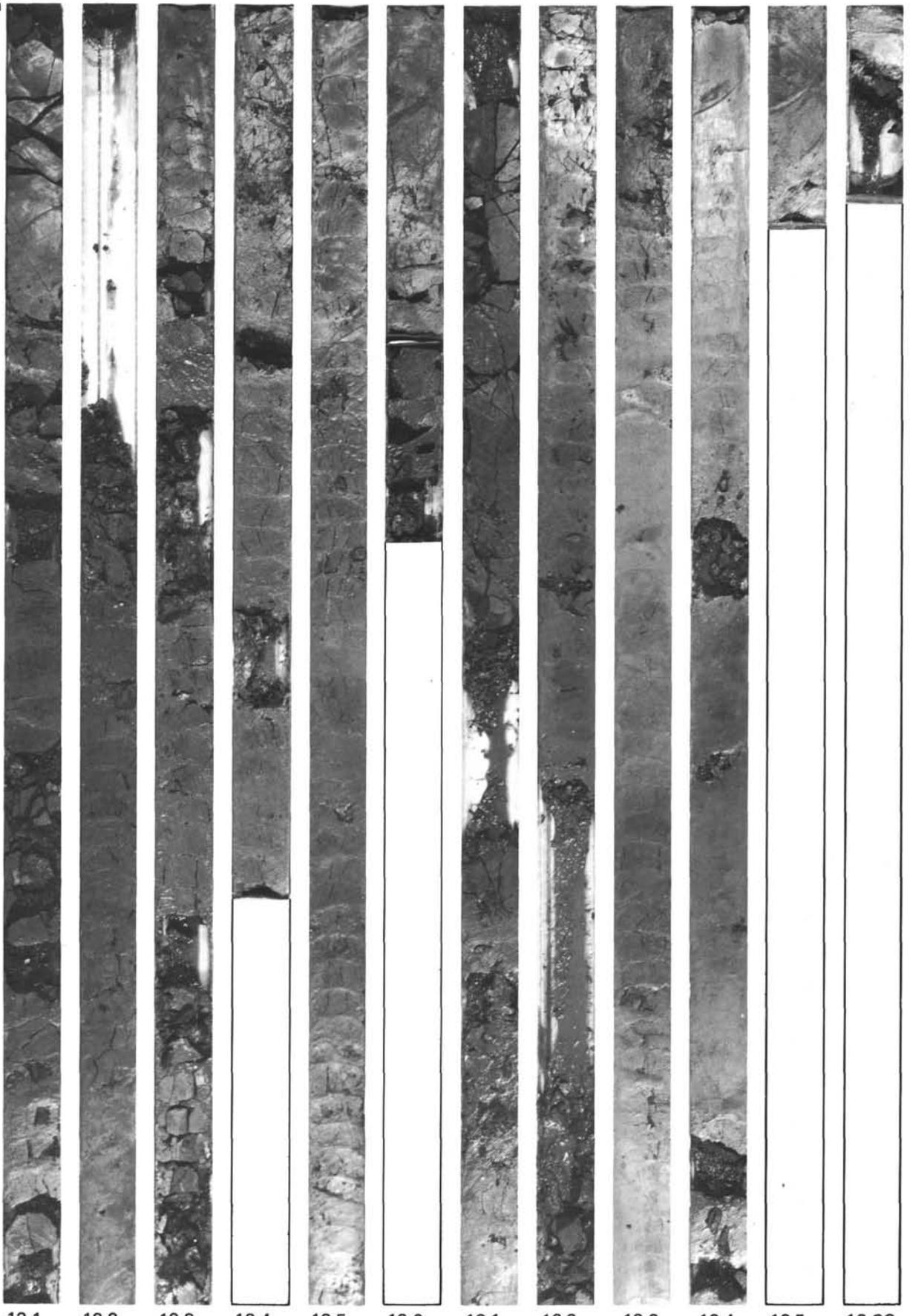
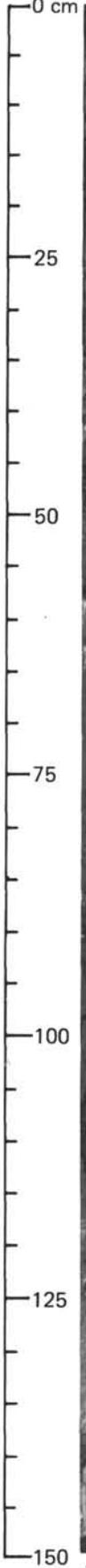




Hole 440B

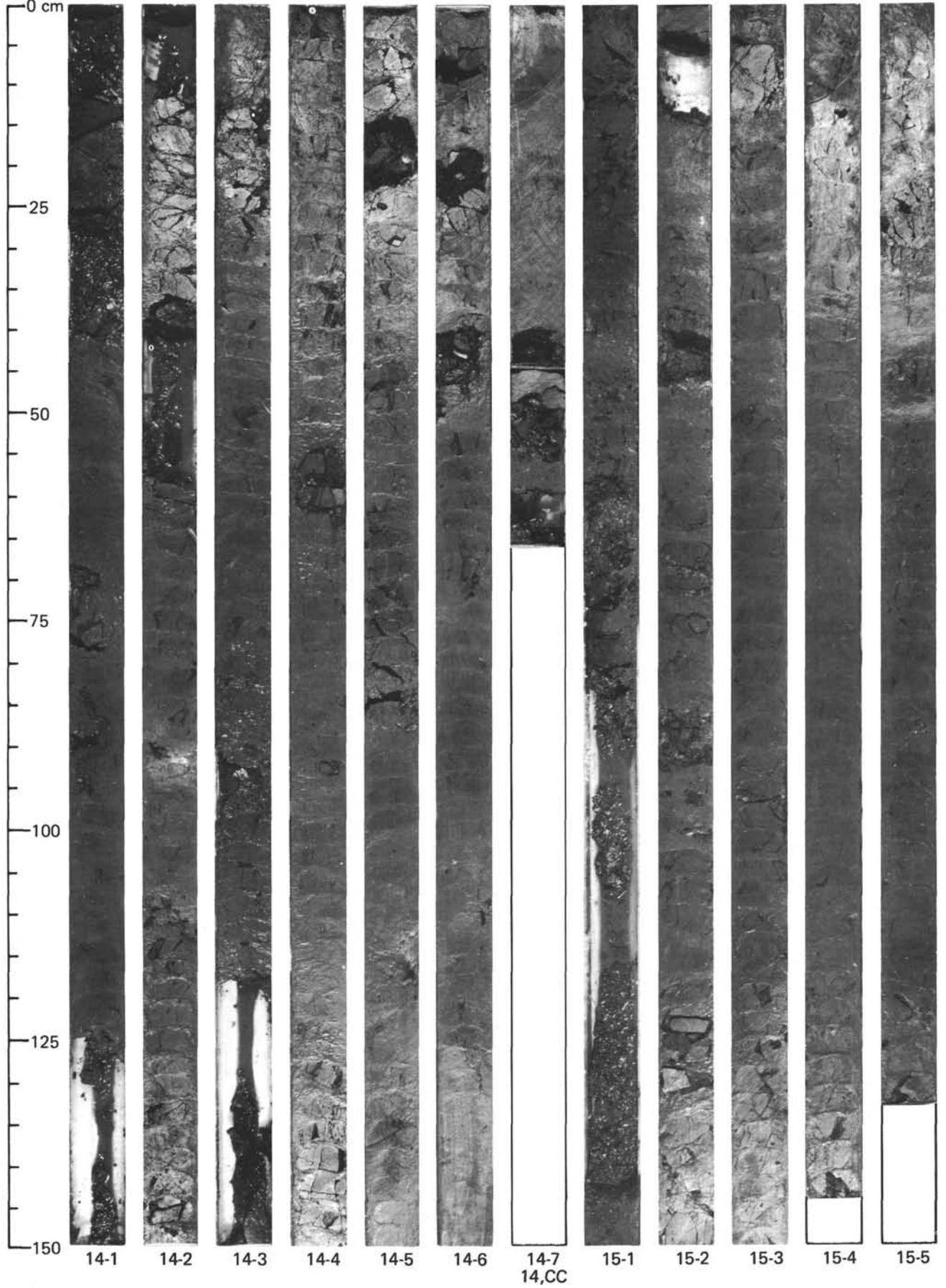


Hole 440B



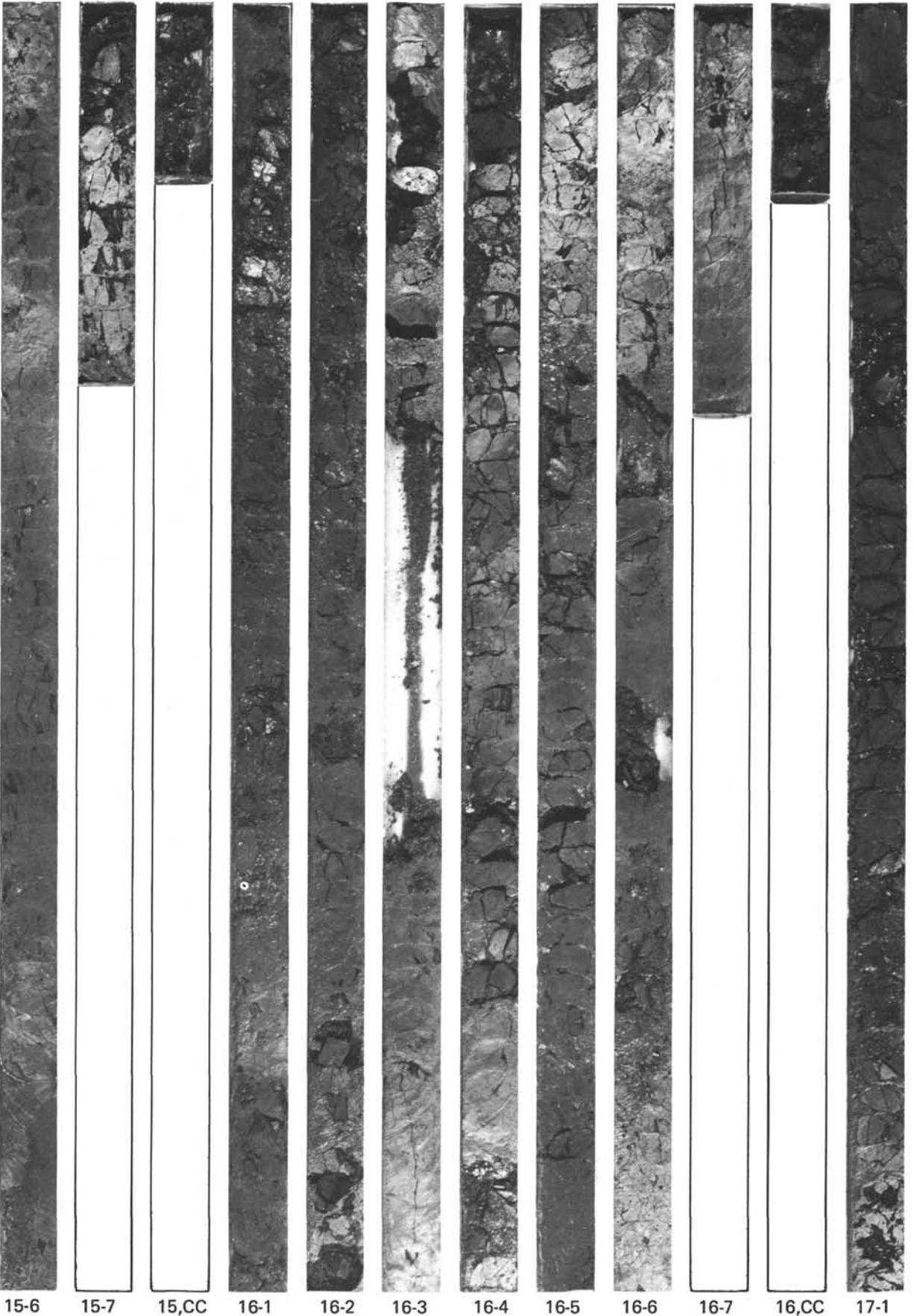
12-1 12-2 12-3 12-4 12-5 12-6 12,CC 13-1 13-2 13-3 13-4 13-5 13,CC

Hole 440B

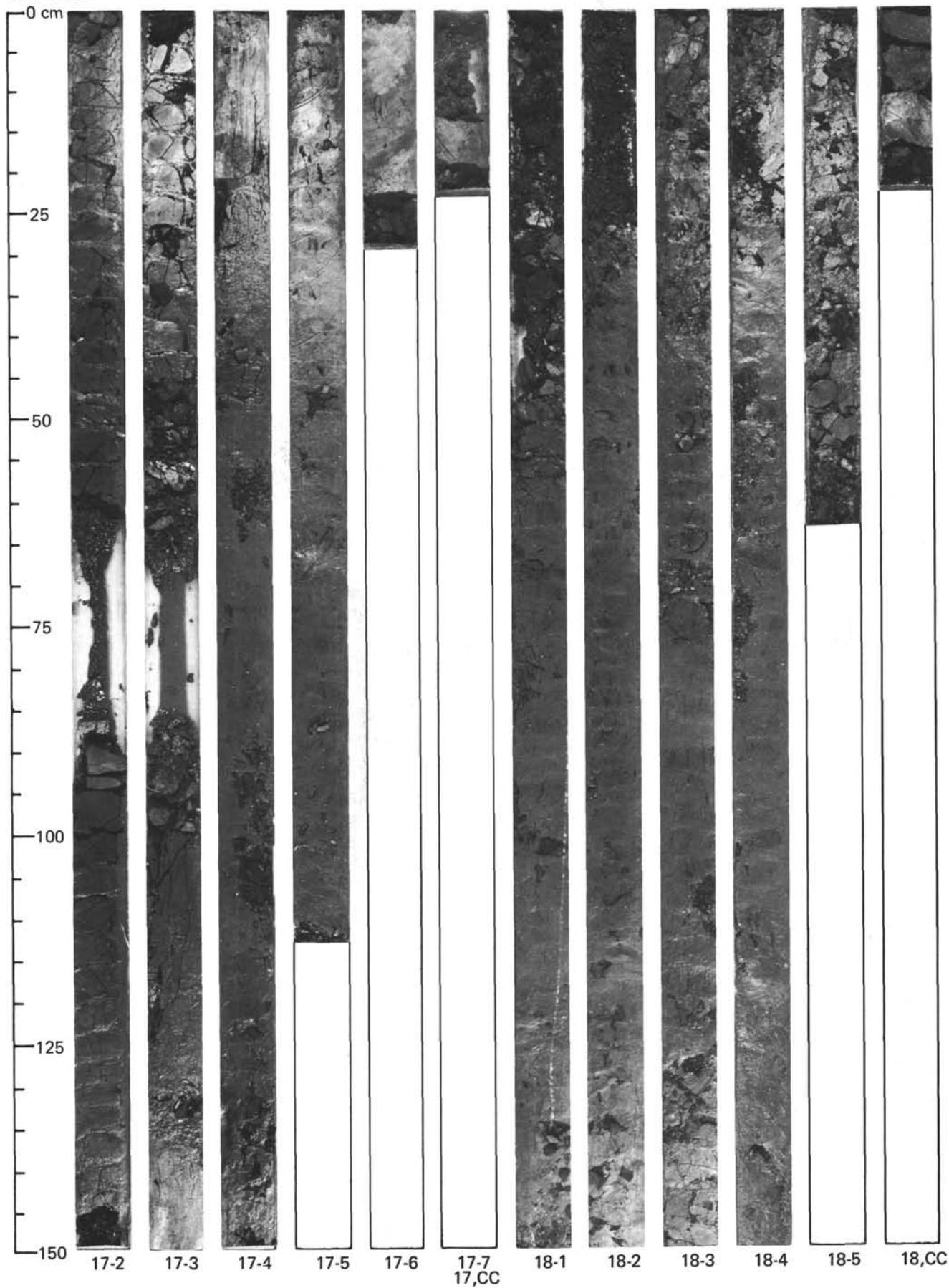


Hole 440B

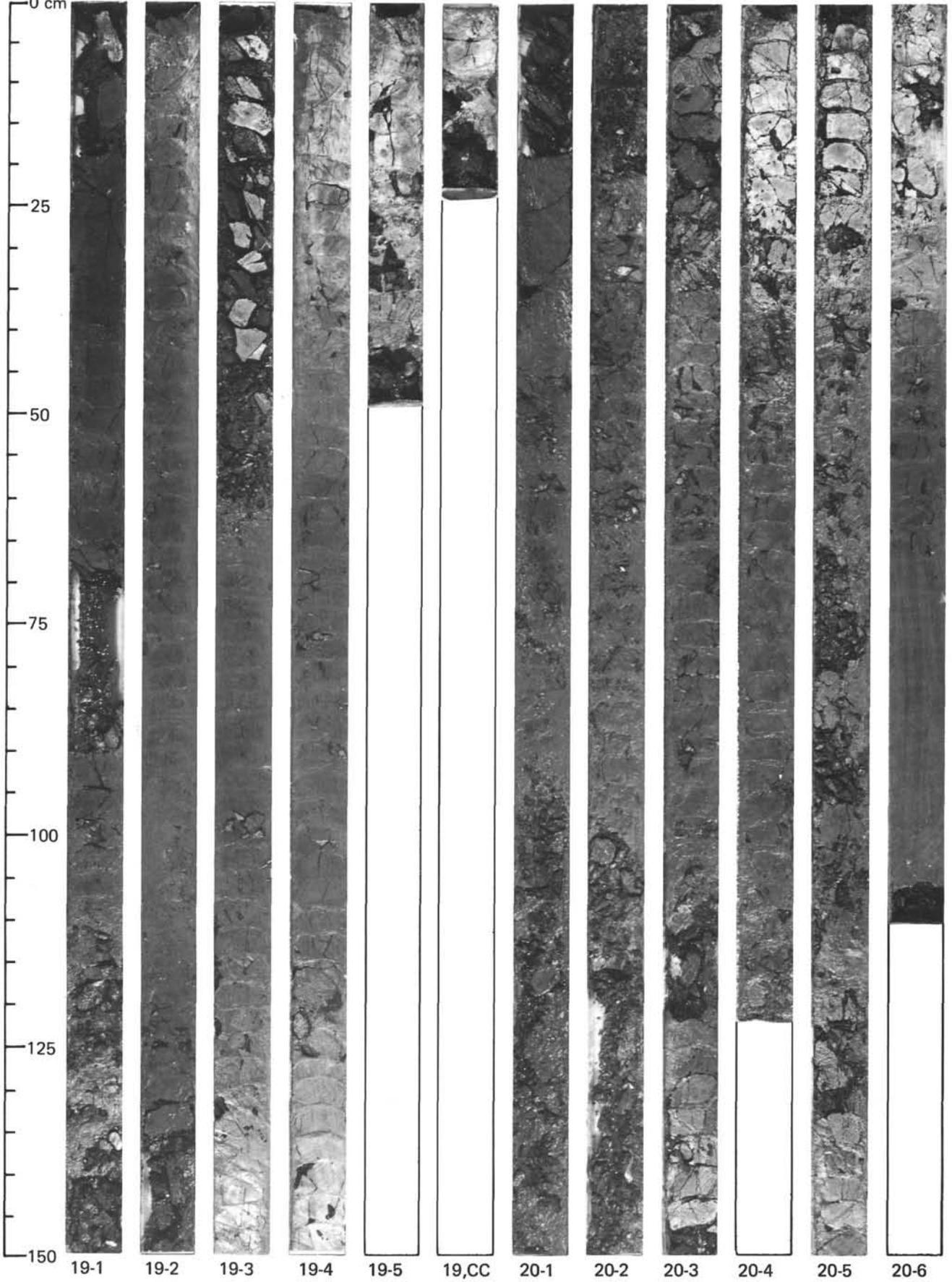
0 cm
25
50
75
100
125
150



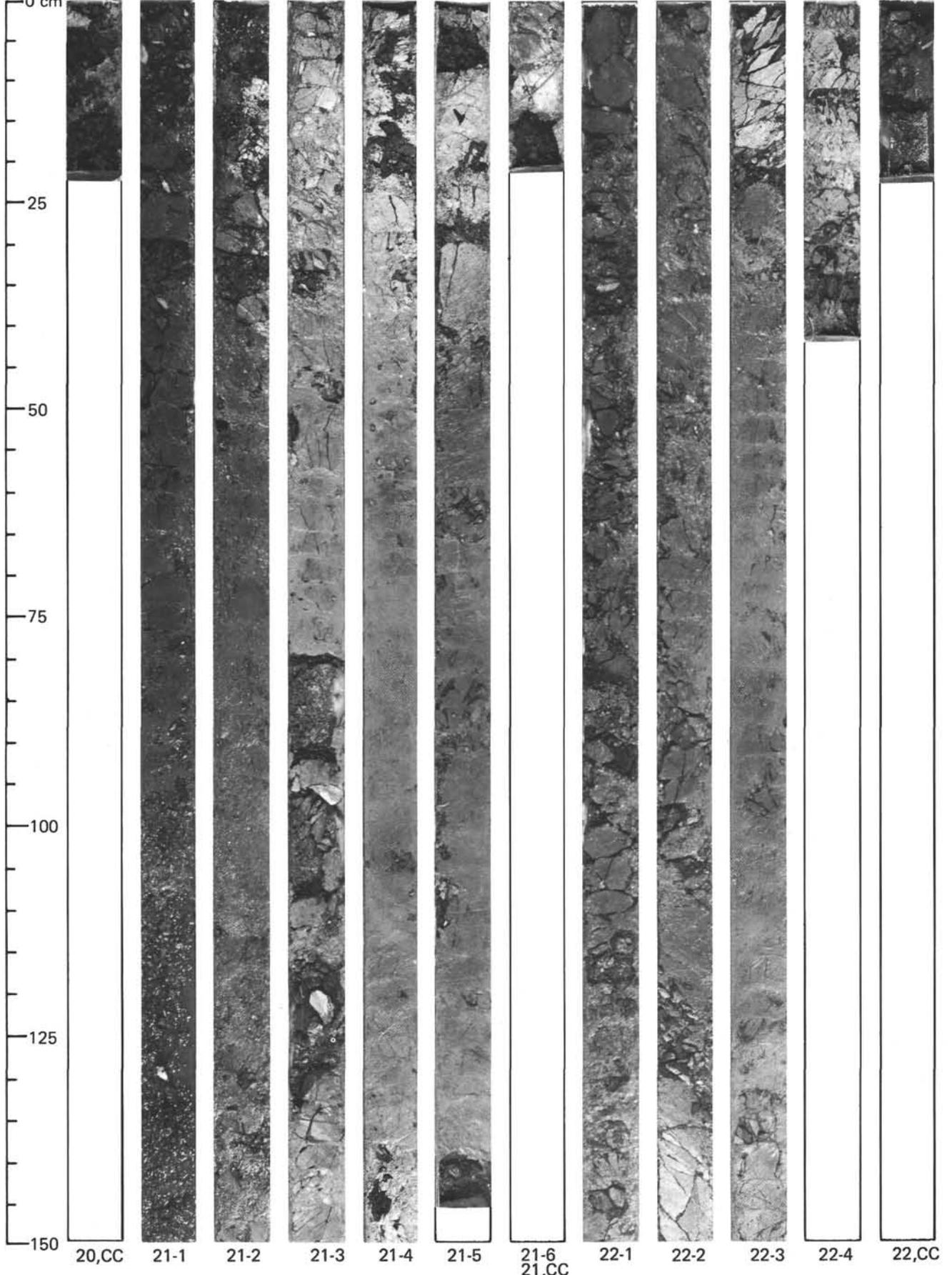
Hole 440B



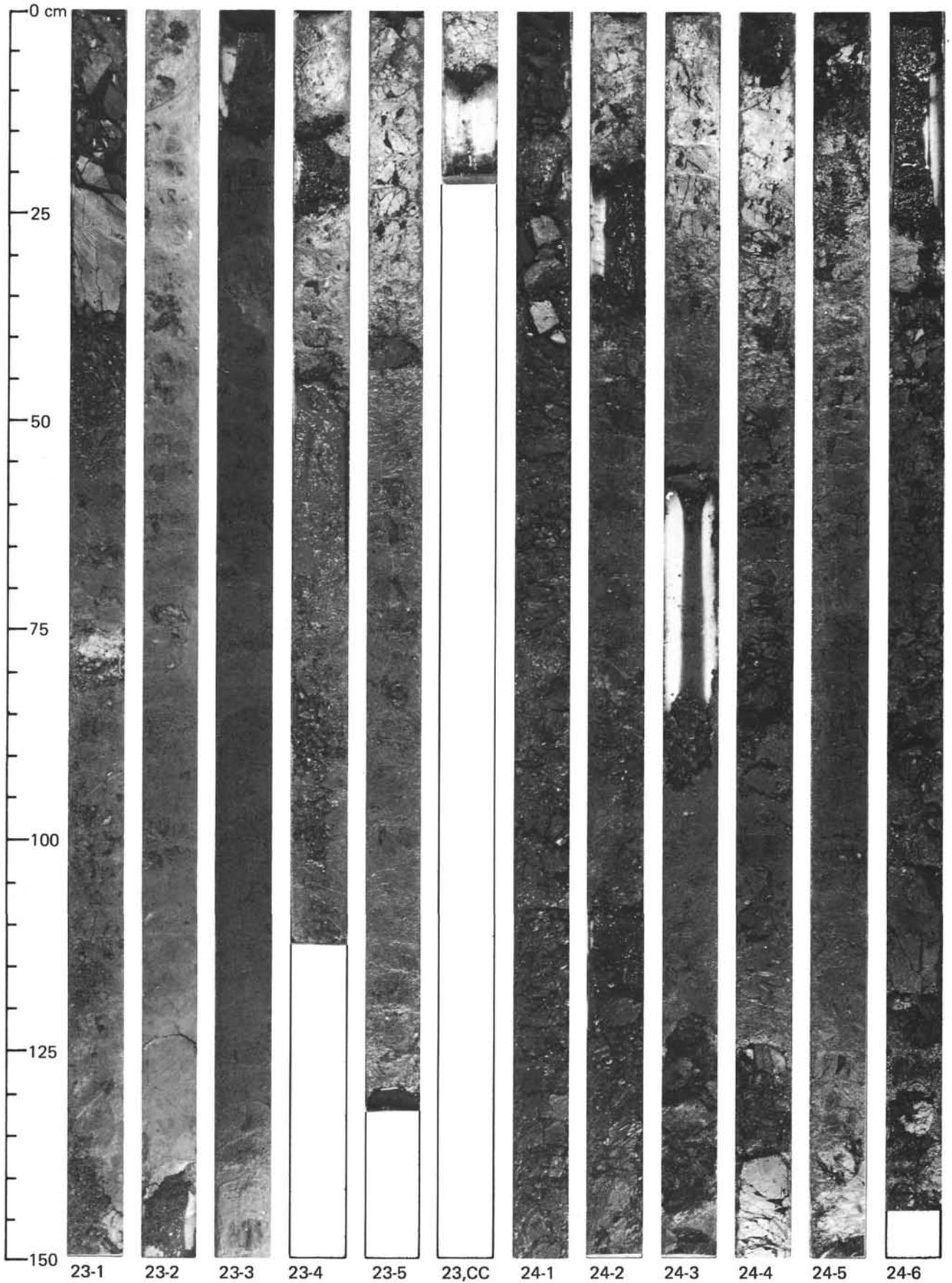
Hole 440B



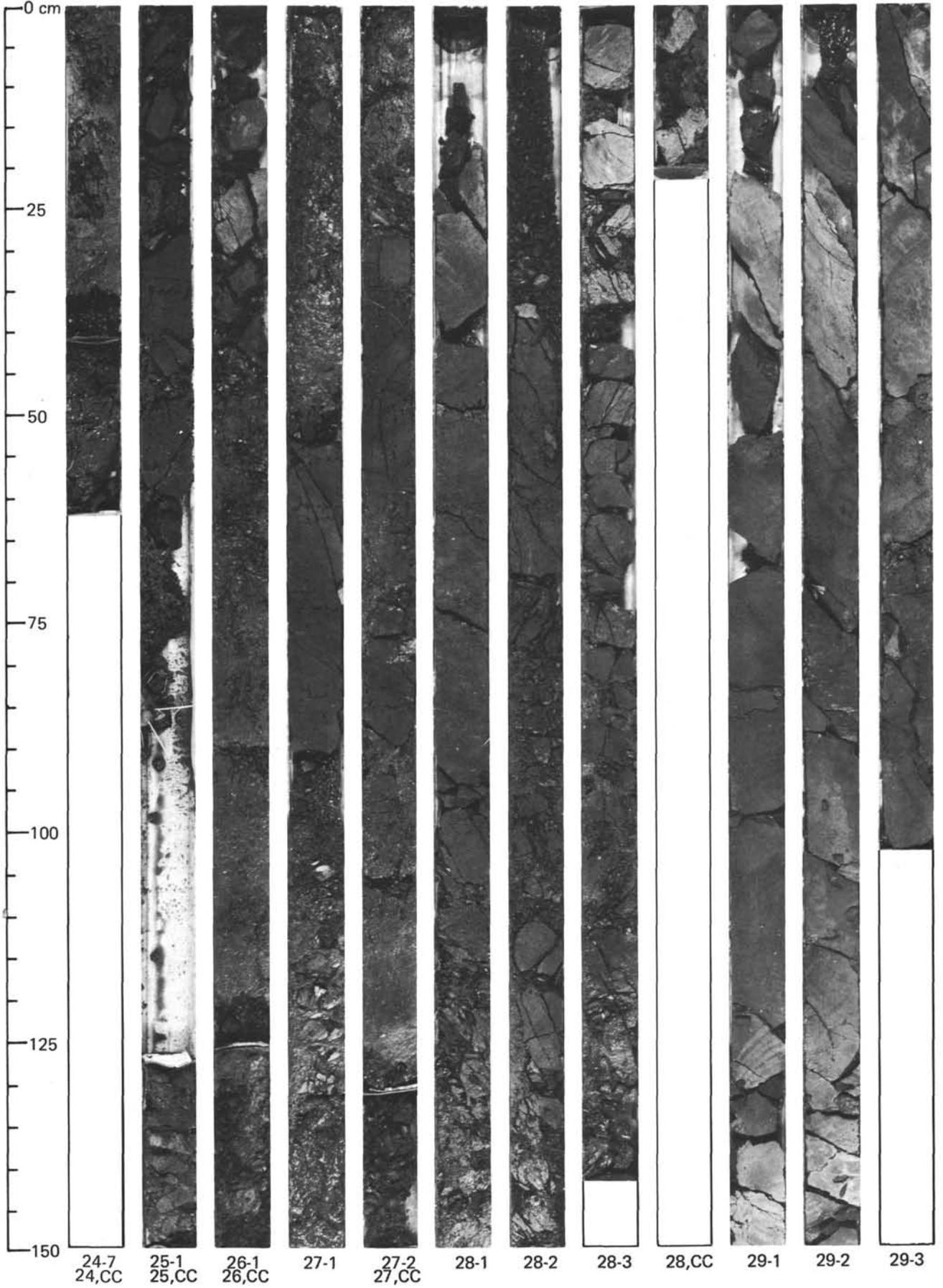
Hole 440B



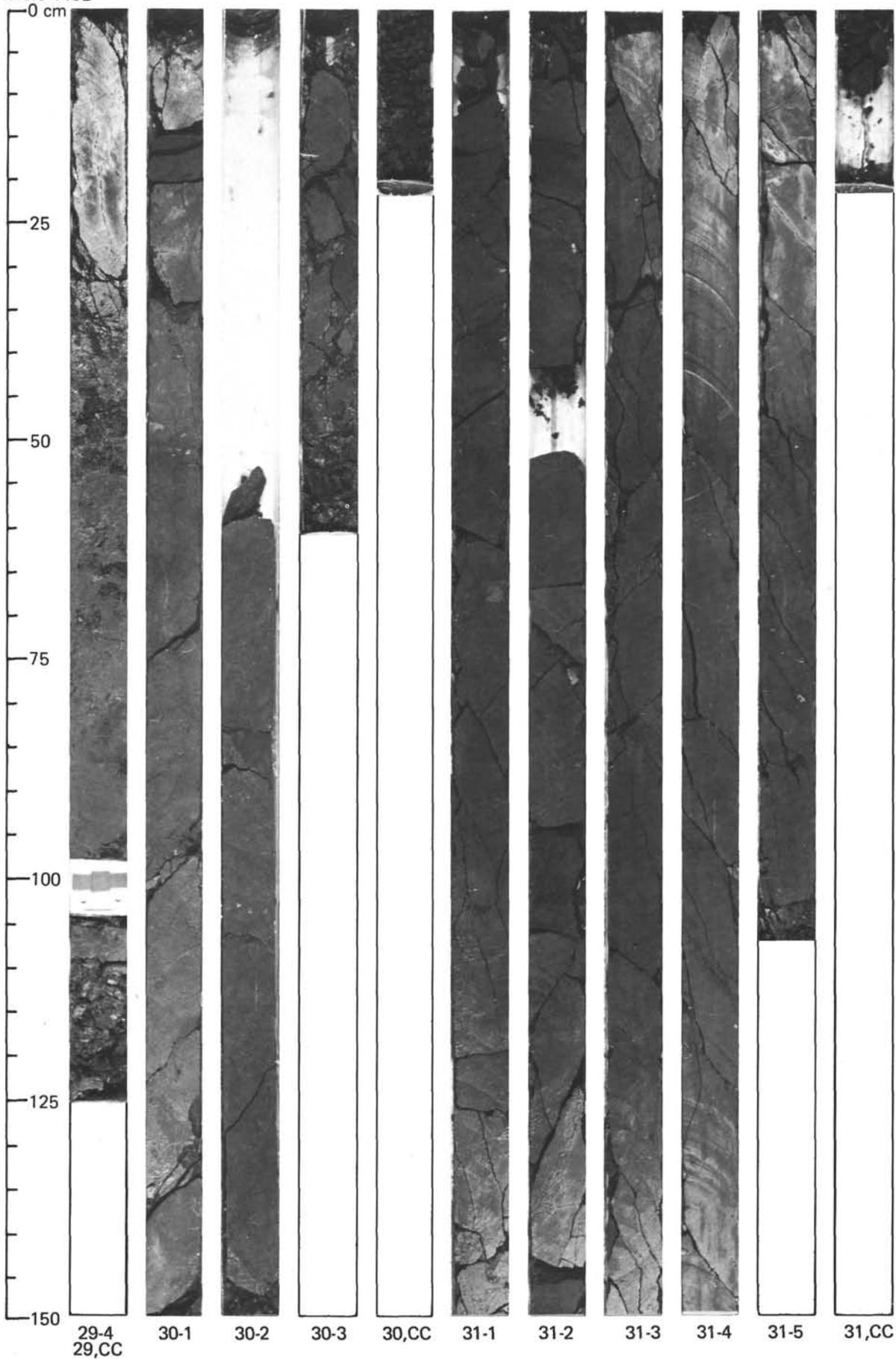
Hole 440B



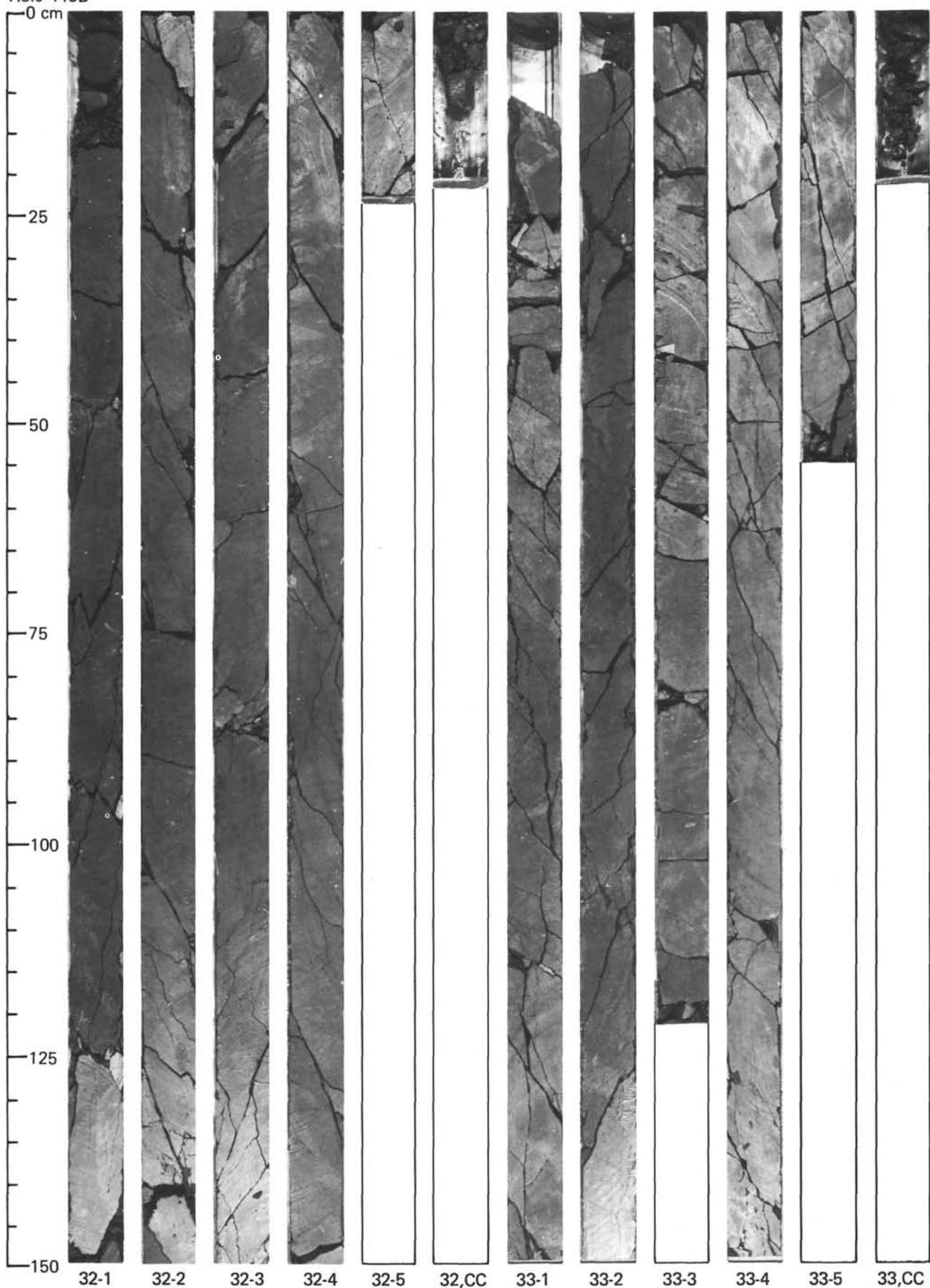
Hole 440B

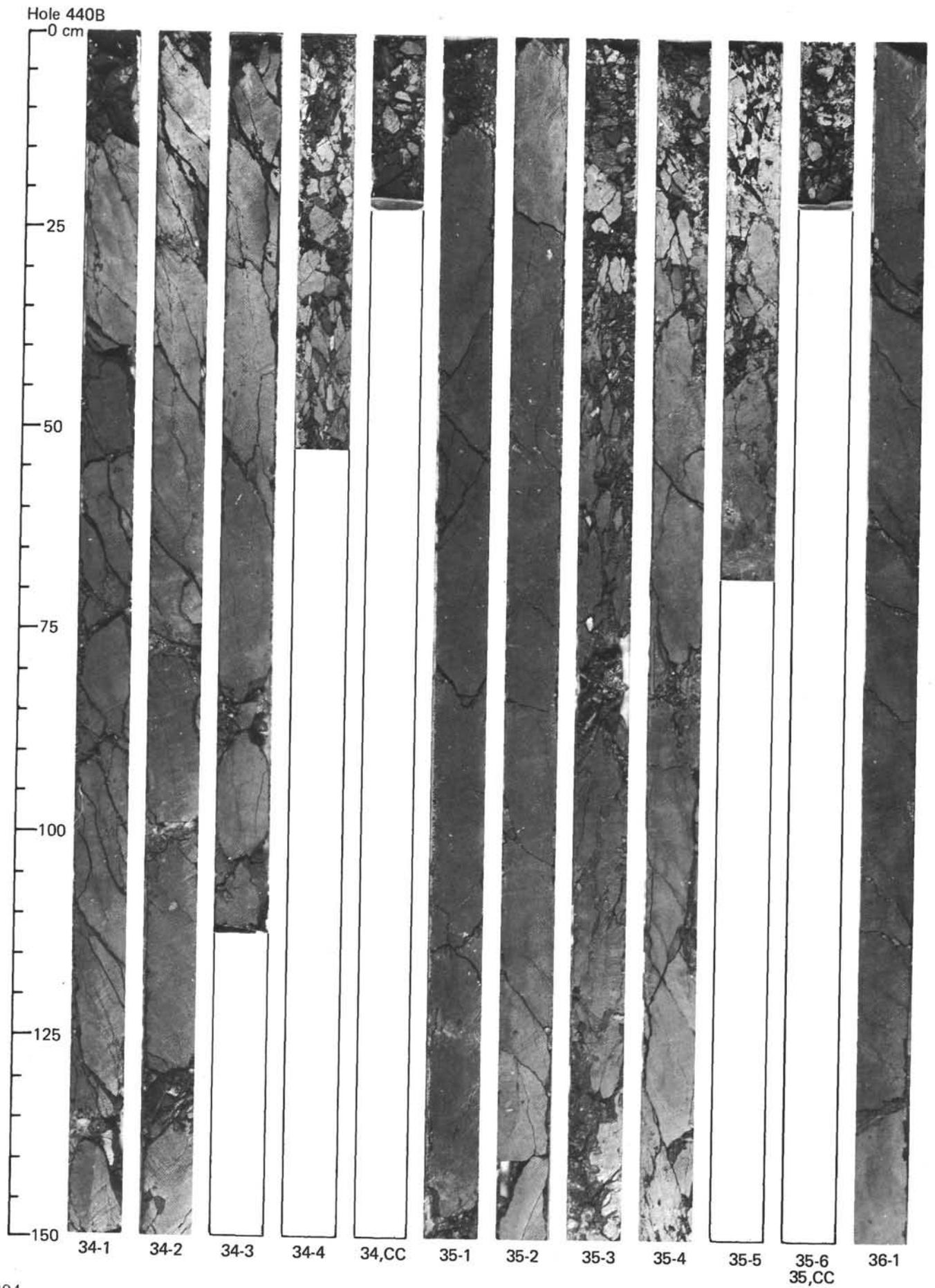


Hole 440B

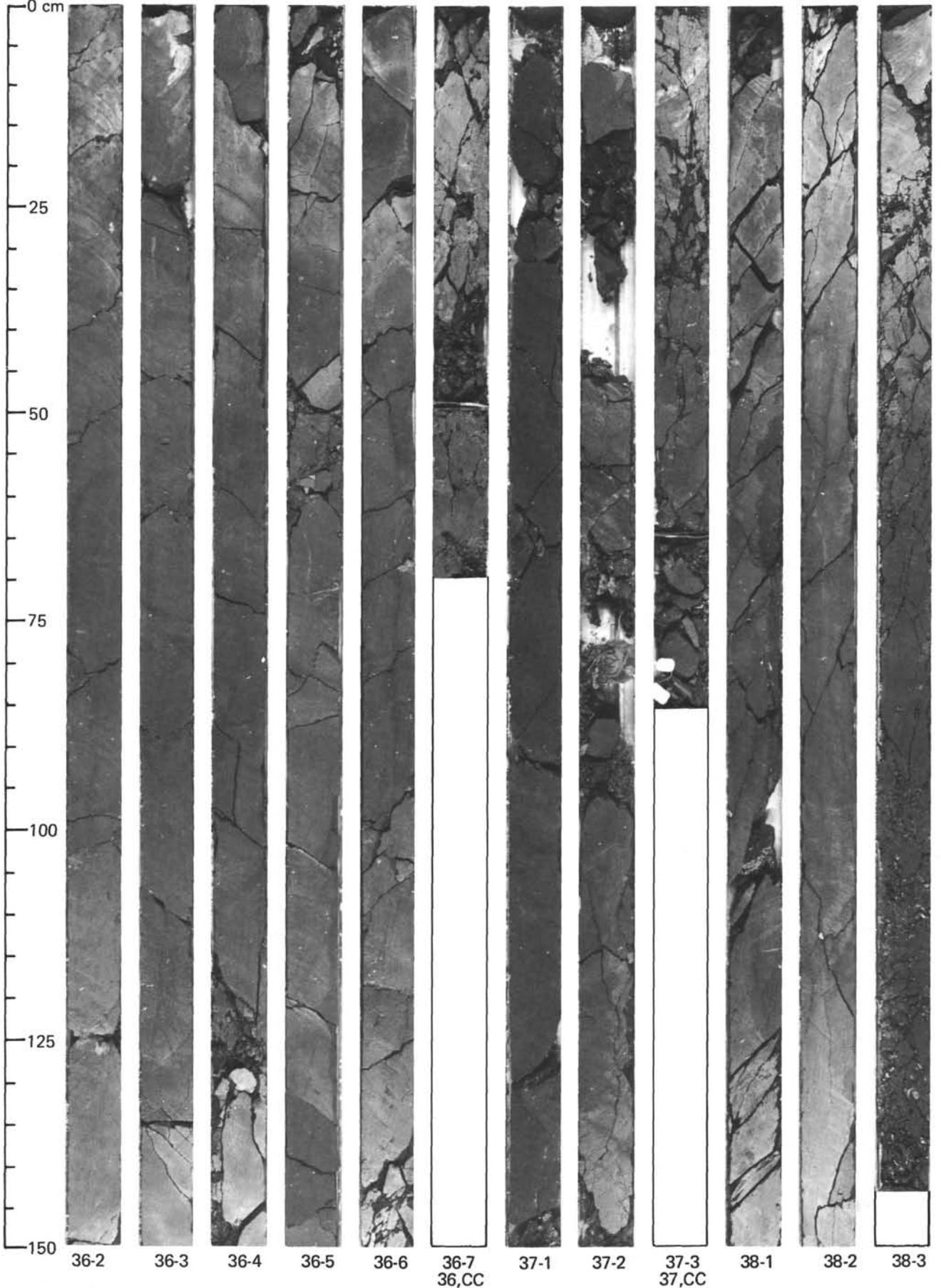


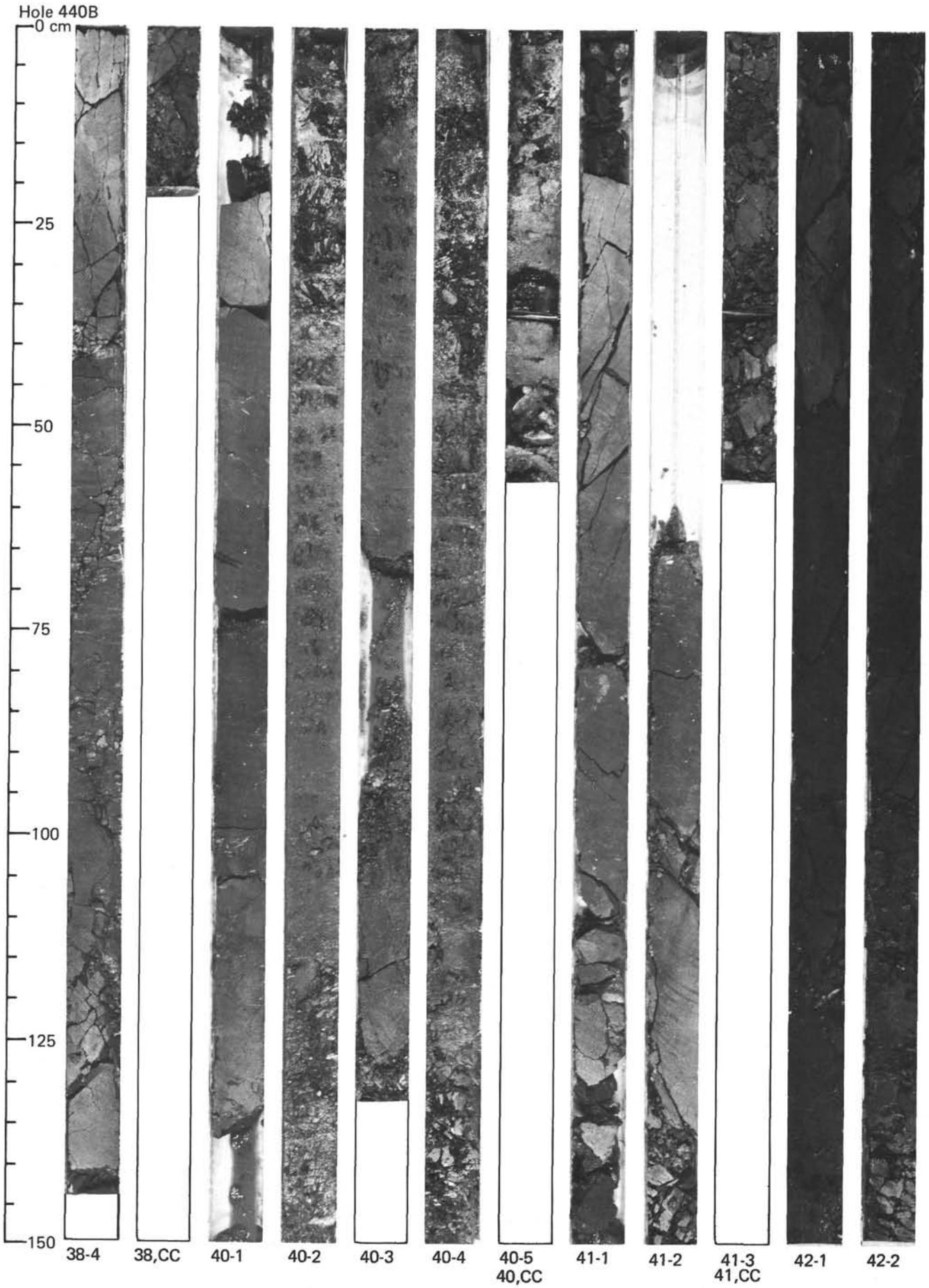
Hole 440B



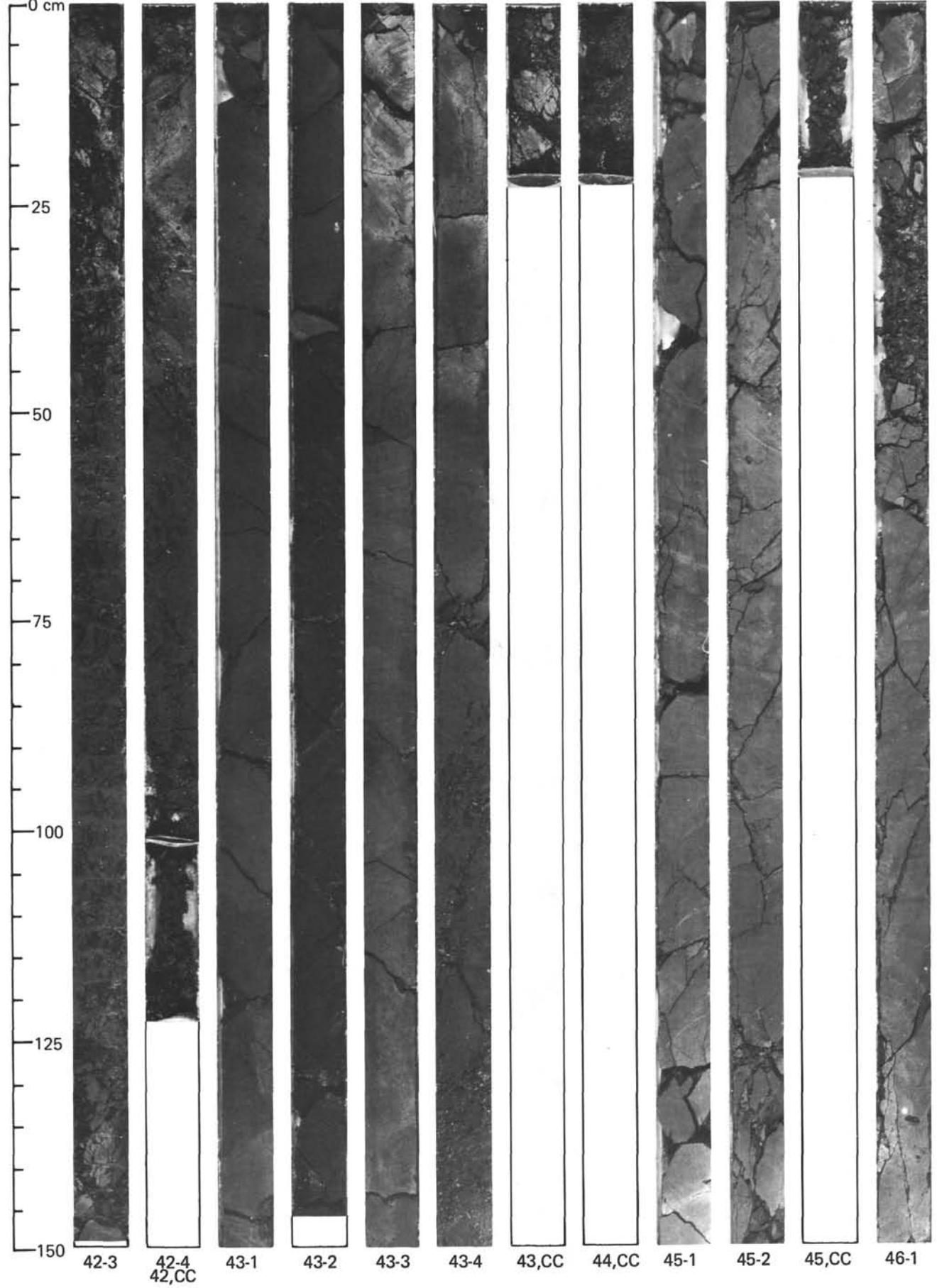


Hole 440B

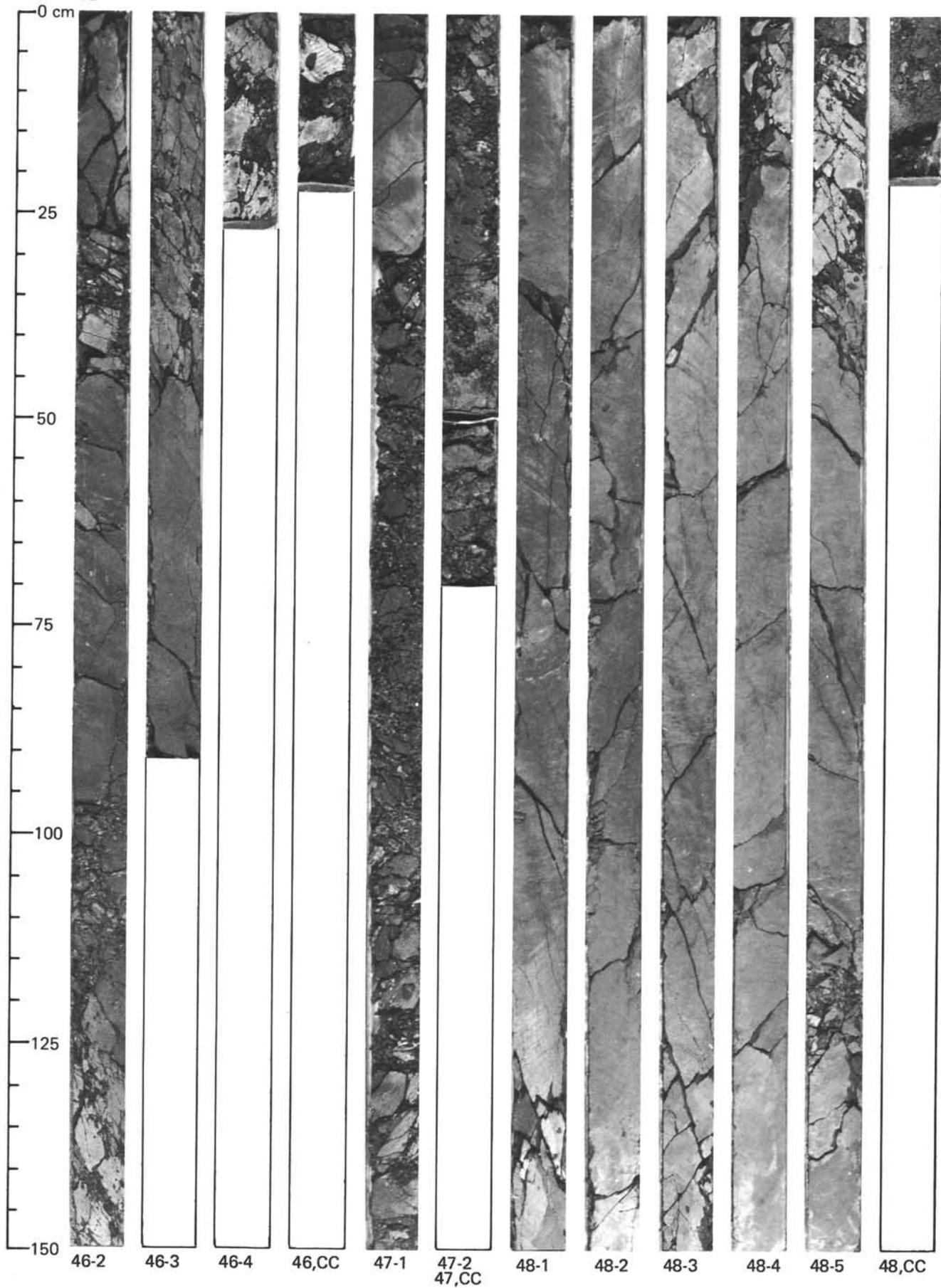




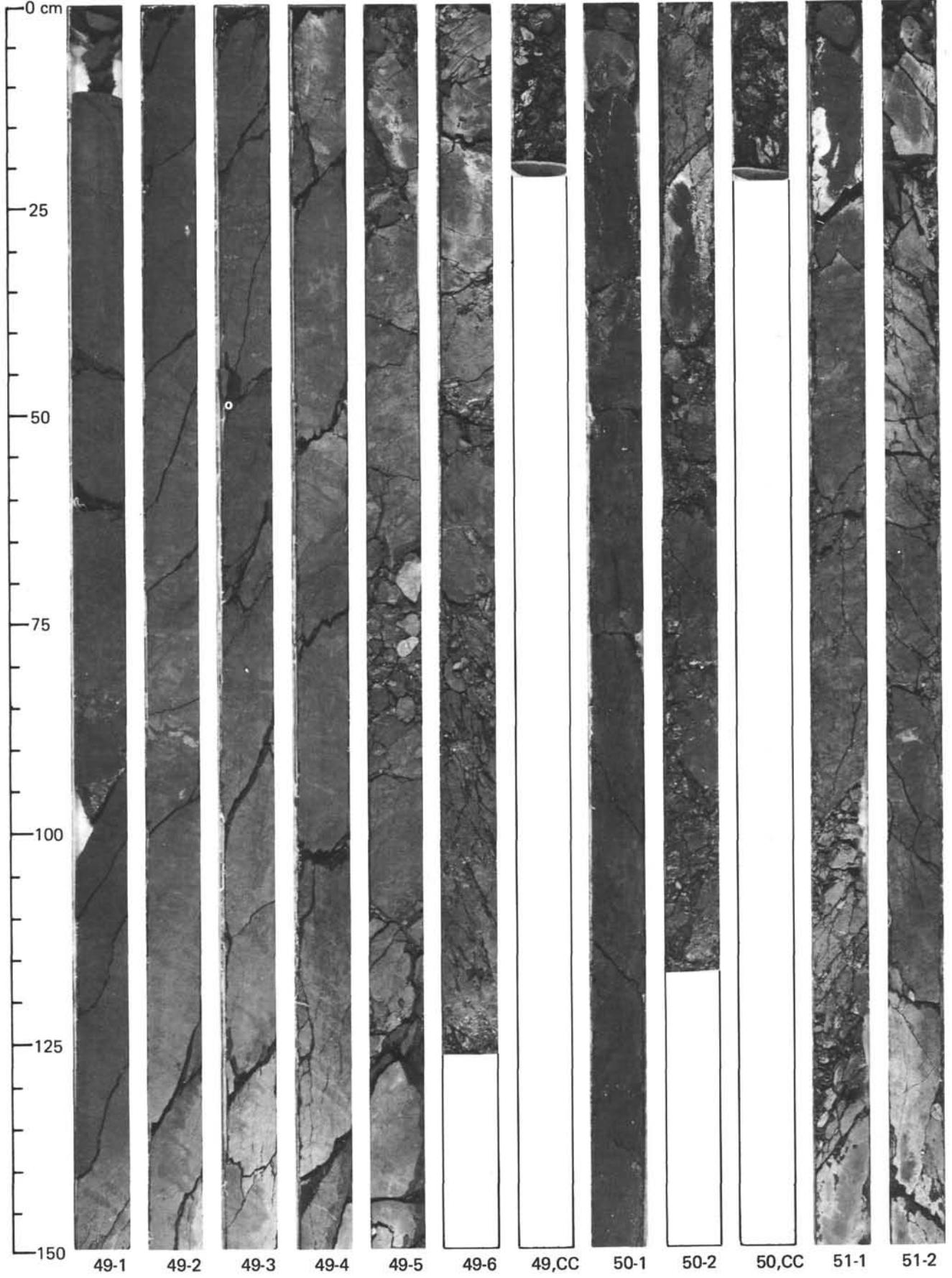
Hole 440B

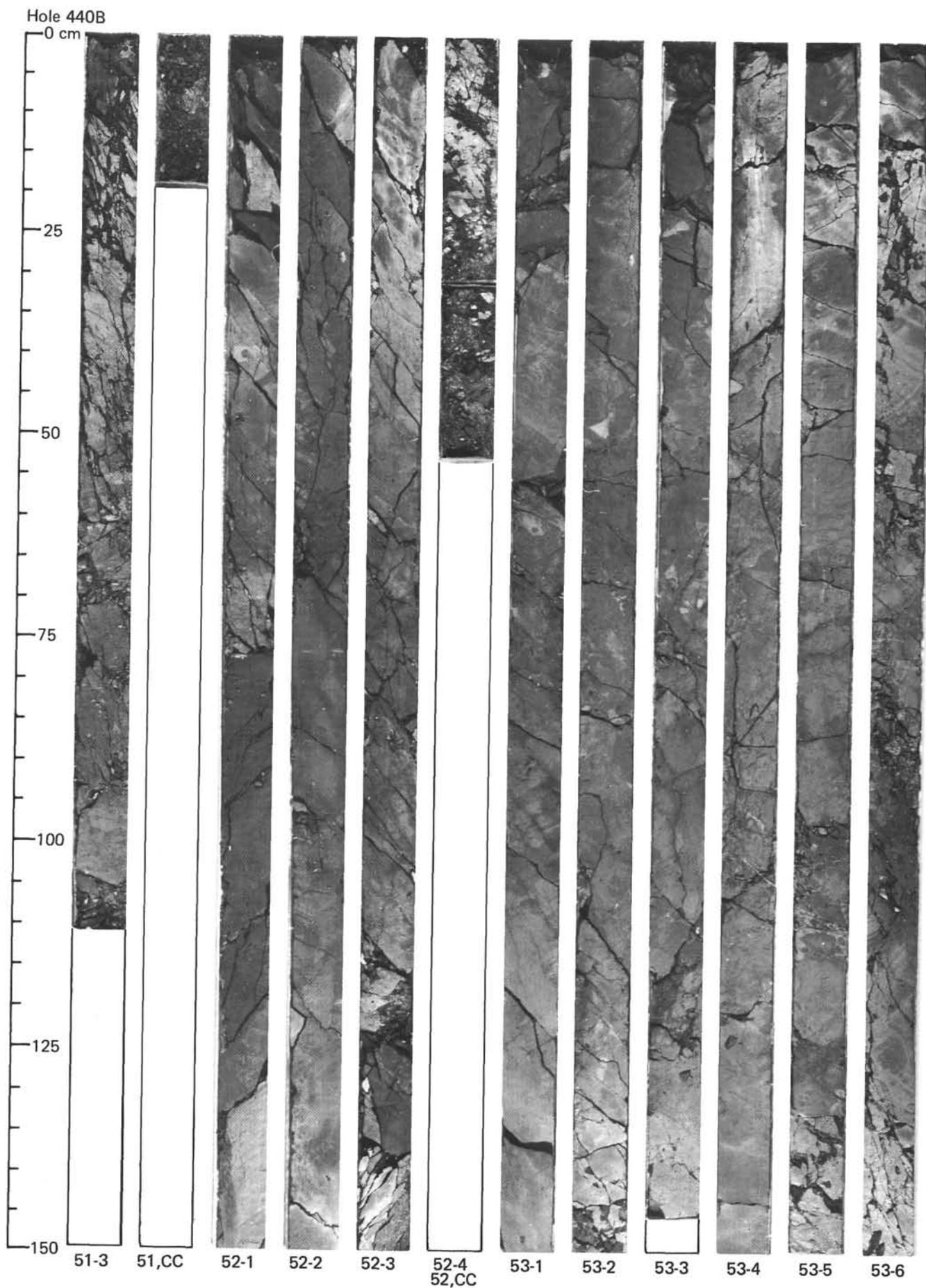


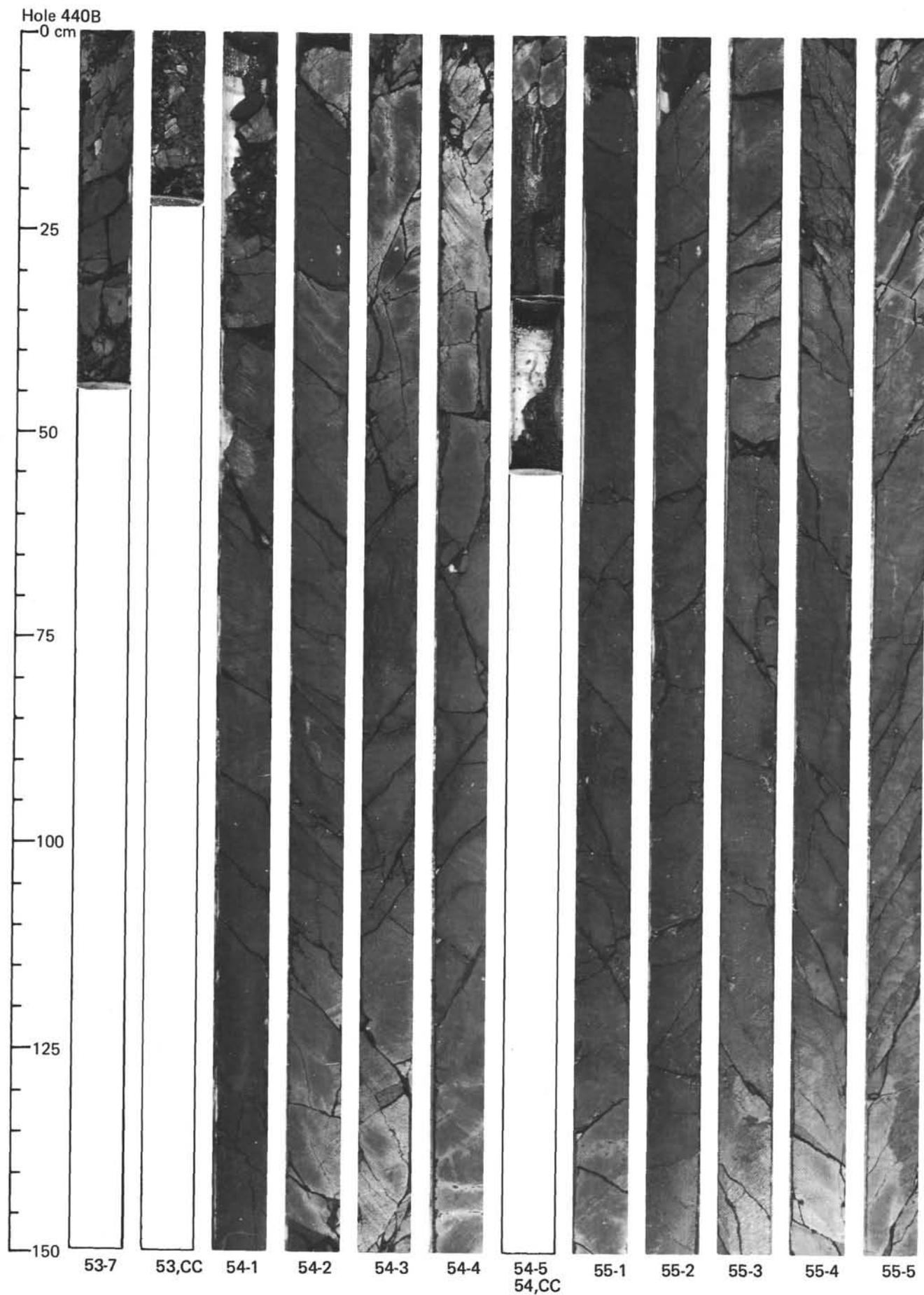
Hole 440B

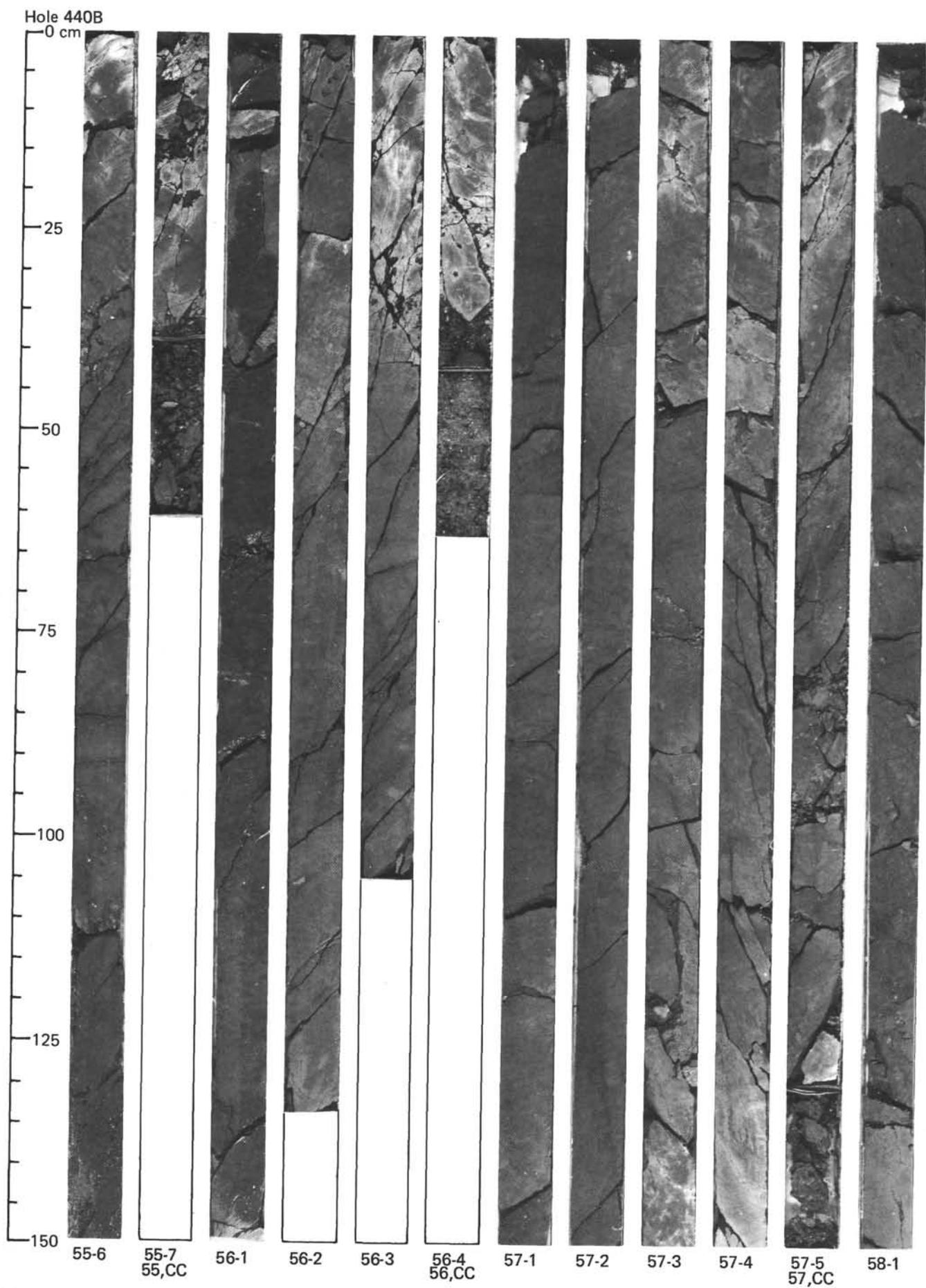


Hole 440B

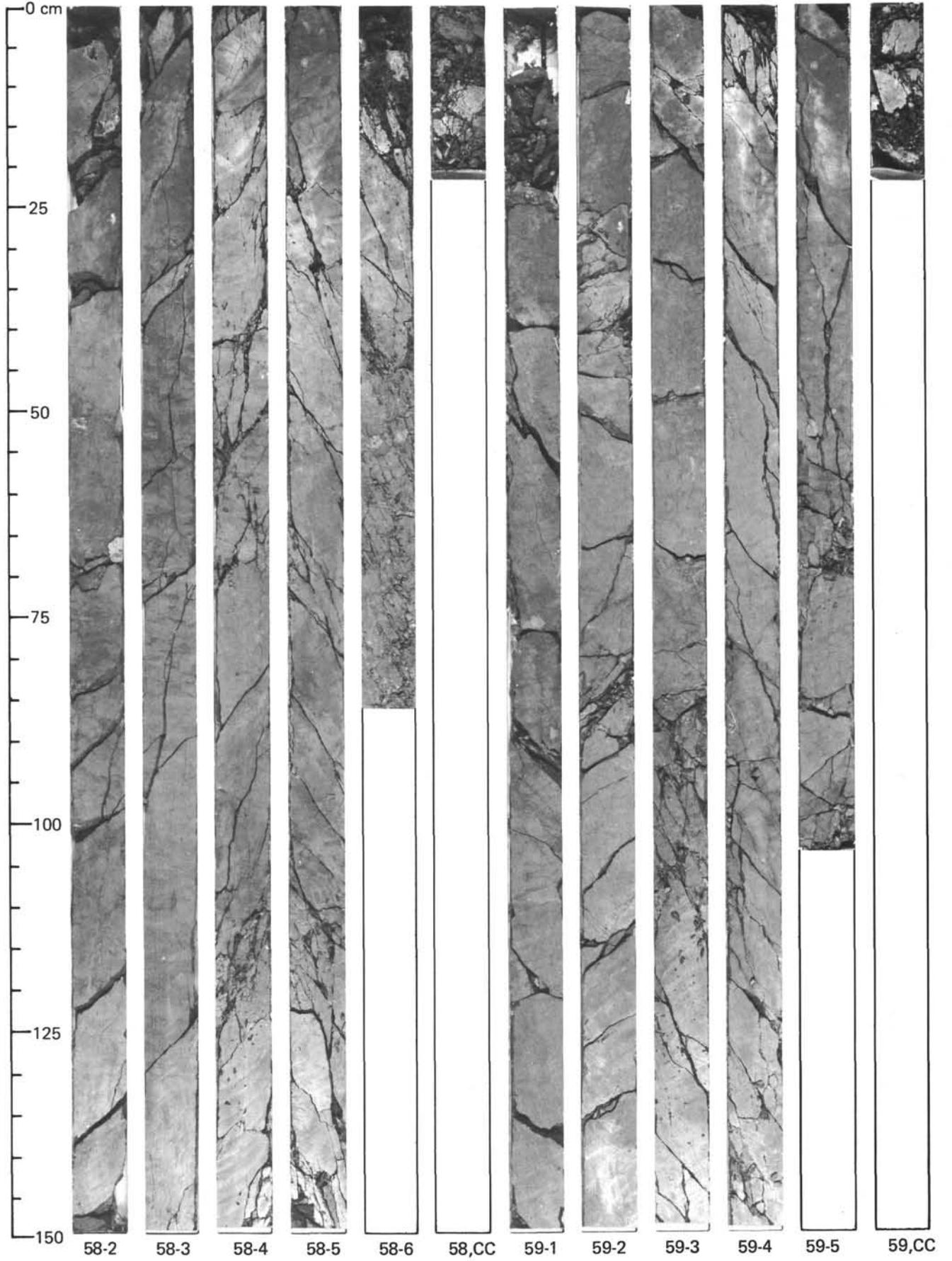




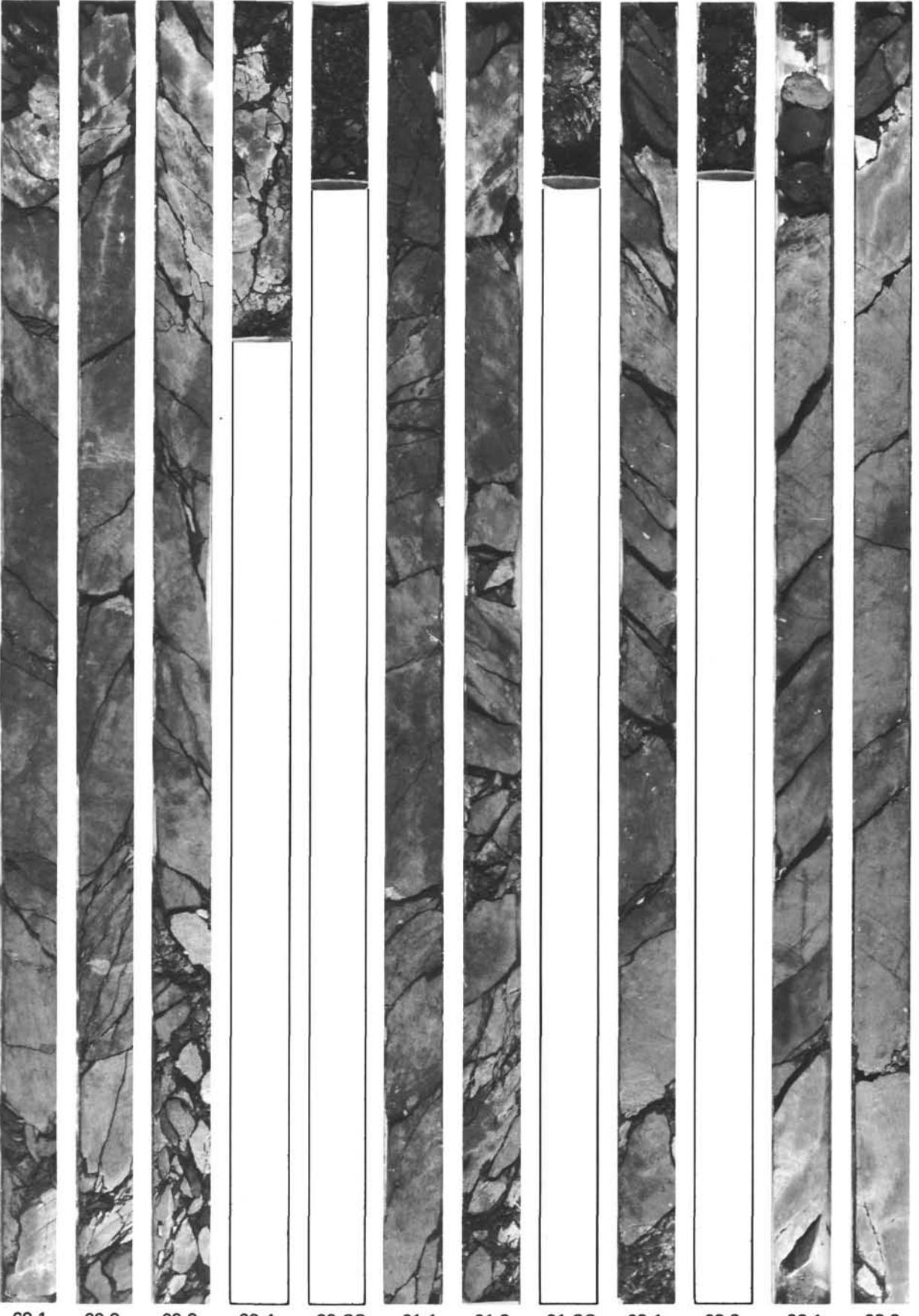
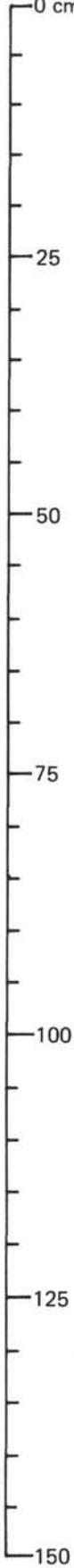




Hole 440B



Hole 440B



60-1

60-2

60-3

60-4

60,CC

61-1

61-2

61,CC

62-1

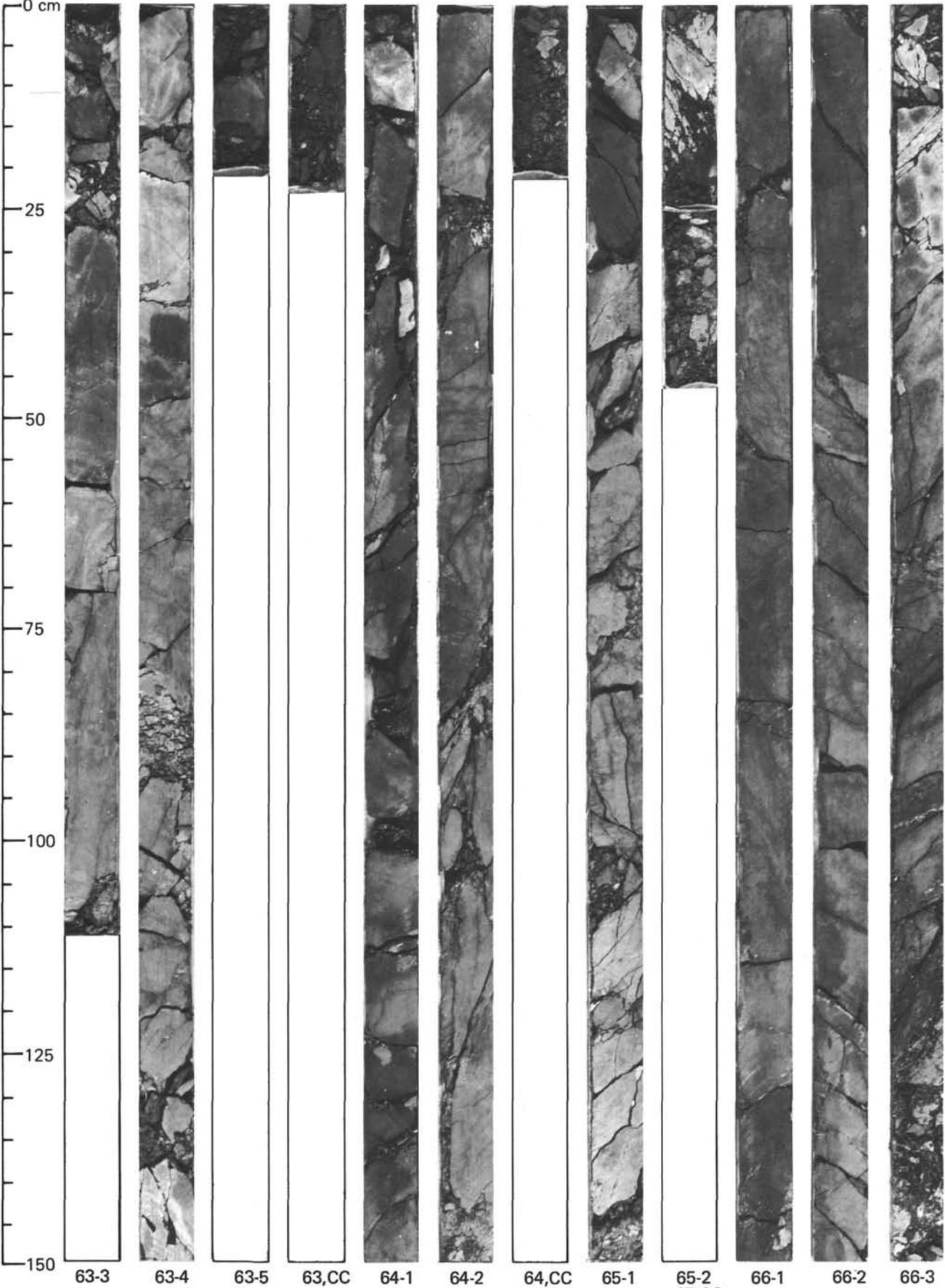
62-2

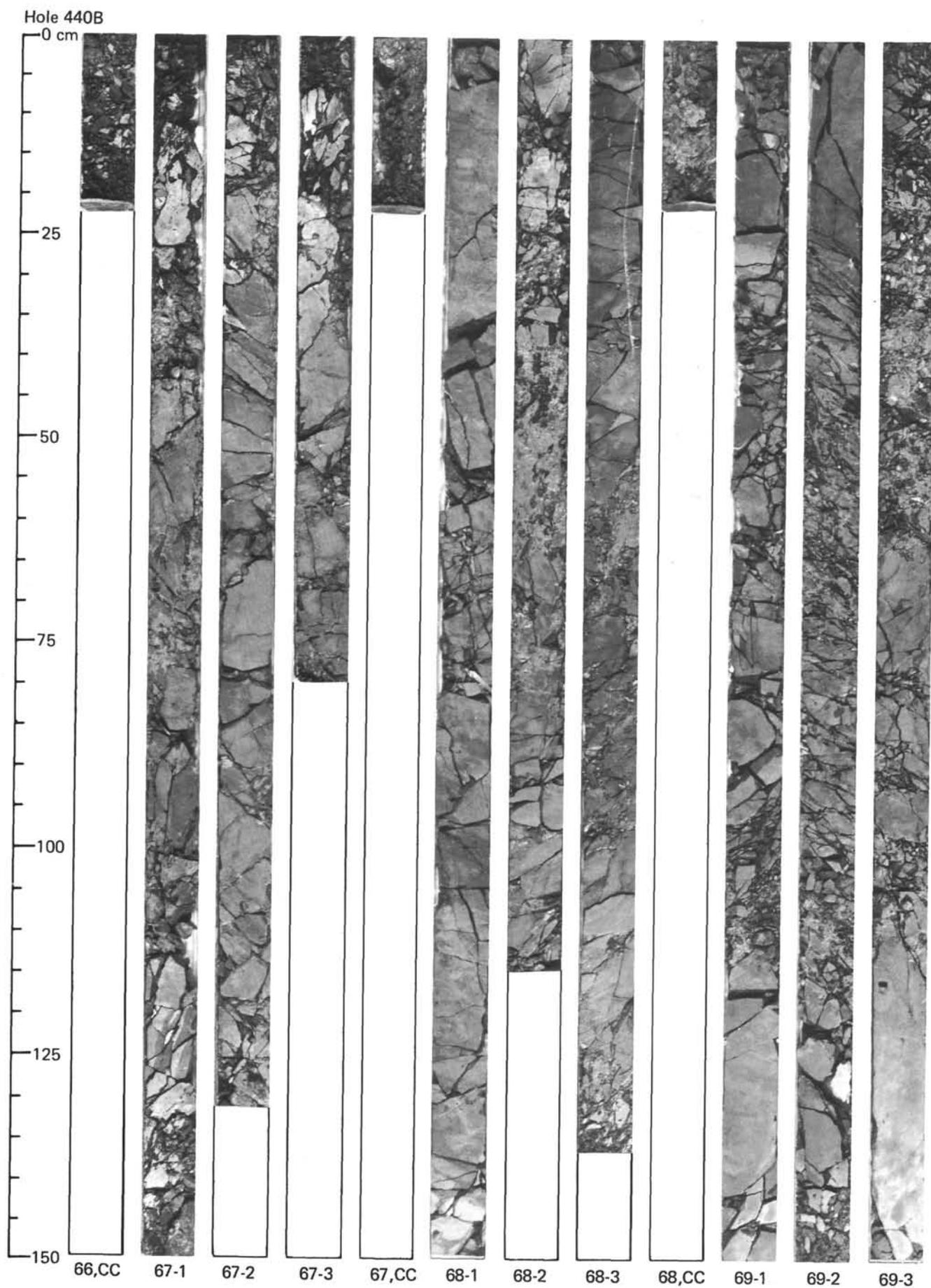
62,CC

63-1

63-2

Hole 440B





Hole 440B

