# 6. SITE 434: THE LOWER TRENCH SLOPE, LEG 56

Shipboard Scientific Party<sup>1</sup>

# HOLES 434, 434A, AND 434B

- Dates Occupied: 12 September 1977 (1800); 22 September 1977 (1300)
- Dates Departed: 18 September 1977 (1010); 30 September 1977 (1720)

Total Time on Site: 13 days 20.5 hours

- Position: Hole 434—39°44.76 'N; 144°06.12 'E Hole 434A—39°44.76 'N; 144°06.12 'E Hole 434B—39°44.87 'N; 144°06.08 'E
- Water Depth (sea level): Holes 434 and 434A-5985.8 meters; Hole 434B-5986 meters
- Water Depth (rig floor): Holes 434 and 434A-5995.8 meters; Hole 434B-5996 meters
- Bottom Felt (rig floor): Hole 434—6000.5 meters; Hole 434A— 5998 meters; Hole 434B—5996 meters

Penetration: 637.5 meters

Number of Cores: 72

Total Length of Cored Section: 662 meters

Total Core Recovery: 133.35 meters

Oldest Sediment Cored: Late Miocene

#### **Principal Results:**

Three holes were drilled at this site; the deepest penetrated to 637.5 meters sub-bottom. The drilled section consisted of a very uniform hemipelagic deposit. The oldest sediments sampled were late Miocene. The major constituents are terrigenous silty clay, biogenic silica, and vitric ash. These sediments are similar in nearly every aspect to the Neogene sediments at the other transect sites. Below 100 meters significant induration of the sediments begins. The mudstones recovered are highly fractured with slickensiding commonly discernible on fractured faces. This fracturing is no doubt the main cause of drilling and coring difficulties encountered at this site. These fissile sediments were found only in Holes 434, 440, and 441 on the lower slope of the inner trench wall.

The early Pleistocene sediments are missing from the section. Pliocene sediments begin in Core 2 of Hole 434 at 10 meters. Benthonic foraminiferal assemblages are similar to those at shallower sites on the landward trench slope and do not contain the taxa that characterize the oceanic biofacies at Site 436. Although the foraminiferal evidence does not identify sediments at Site 434 as belonging to an accretionary prism of oceanic sediments, there is biostratigraphic evidence suggesting considerable structural complexity: All fossil groups show anomalous vertical distributions. Diatom data in particular identify several repetitions of assemblages that form an undisturbed succession elsewhere. Possible causes for these distributions, other than downhole contamination, include slumping, sliding, faulting, and/or climatic oscillations.

Below 300 meters there are calcitic concretions and pods, and fractures are often healed. The concretions are thought to be diagenetic products, because some contain diatoms that are also found in the surrounding mudstone and some contain vitric shards showing the same range in values of the refractive index as shards in the surrounding mudstone. The alkalinity of the interstitial water is high, which would favor carbonate precipitation.

The sediments at Site 434 were anoxic, as evidenced by high interstitial water alkalinities, a strong  $H_2S$  odor to about 200 meters, and biogenic methane gas pockets from 130 meters to the bottom of the hole. Geochemical analyses indicate that the hydrocarbons are immature and have experienced only low temperatures.

The highly disturbed condition of most of the core materials made representative physical properties measurements difficult to obtain. The seismic velocities showed small but significant variations that may correspond to changes in the degree of fracturing along the section.

#### BACKGROUND

Site 434 is located on the lower slope of the landward wall of the Japan Trench. It has been known for several years from refraction seismic studies that the lower slope was underlain by a wedge of sediments with a relatively low seismic velocity, in the range of 1.6 to 2.5 km/s (e.g., Ludwig et al., 1966). These low velocity deposits extend from the midslope terrace to the floor of the trench. Beneath this sedimentary wedge is a layer

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with velocities typical of oceanic basement, and recent multichannel seismic profiles (see Figure 1) show that the top of this layer is continuous with the oceanic basement surface seaward of the trench. This sedimentary wedge has been described by some authors as identical with scraped off sediments found on the overthrusted leading edge of many oceanic subduction sites (see, for example, Karig and Ingle, 1975; Kulm et al., 1973). Sediments in the accreting prism show a strong degree of deformation in response to the compressive and shear stress resulting from ongoing subduction of the oceanic lithosphere. Apparently, sediments that lie on the subducting oceanic crust are scraped off and plastered onto the base of the accreting prism. Underthrusted layers must be uplifted and compressed to accommodate the continuous input of new material at the foot of the slope. The resulting style of deformation has been studied by means of multichannel seismic techniques over similar deposits along the western Central American Margin (Seely et al., 1974) and analogous deposits now exposed on land. Typically, the sedimentary layers are positioned in sheetlike overthrusted units, forming an imbricate stack. The shearing plane steepens noticeably landward from the foot of the slope.

### **OBJECTIVES**

A primary objective for Site 434 was to penetrate as deeply as possible into the accretionary wedge at the foot of the Japan Trench landward slope to sample sediments involved in the accreting tectonics. The geophysical surveys of the Japan Trench, which include 24-channel digital reflection profiling, provide only a vague indication of the configuration of the sedimentary layers in the accretionary zone, owing in large part to the enormous complexity of the acoustic stratigraphy and lack of large velocity contrasts. Penetrating the wedge with a drill might reveal the style of deformation through such features as fracturing, folding, slickensiding, orientation of bedding plane, and so forth. A direct confirmation of the hypotheses of formation and evolution of accretionary wedges would be the discovery of sediments irrefutably identified with the sediments on the Pacific Plate seaward of the Trench.

During Leg 56 the maximum penetration capacity of the *Glomar Challenger* was approximately 6800 meters; thus it was impossible, as would have been desirable, to drill into the toe of the accretionary wedge. Site 434 was at 5986 meters water depth on the last well-defined terrace before the Trench floor. The goal was to drill and core 800 meters sub-bottom.

Although no clear acoustical stratigraphy can be discerned below the selected site, nearly horizontal discontinuous reflectors can be followed over short distances in a zone 0.3 to 0.5 second of two-way travel time subbottom (Figure 2). Below 0.5 second, other vague reflectors with a perceptible landward dip can be traced. Further description and discussion of the geophysics in the vicinity of Site 434 are in the Site 441 site chapter.

# **OPERATIONS**

We approached Site 434 on a west-to-east line that closely followed earlier geophysical lines made during site surveys. The *Glomar Challenger* obtained geophysical data at about seven knots along the line, and the morphological features and sub-bottom acoustic stratigraphy defined by the surveys were easily identified. The site was a small terrace on the steep lower slope on the landward wall of the Trench. Seismic reflection profiles across this feature suggested that the terrace may be underlain by slumped material (Figure 2). We decided, therefore, to place Site 434 a little farther downslope on another, deeper, terrace at about 5990 meters.

The bottom beacon was dropped on the second pass over this terrace in a water depth of 5986 meters. The drill pipe was lowered and touched bottom with 6000.5 meters of pipe out. The first hole at 434 was drilled and cored to a depth of 301 meters sub-bottom. Over the entire length of this hole recovery was only about 18 per cent. Of the 33 cores cut, 9 contained 0.5 meter or less of sample. In many of these the sediment was highly disturbed by drilling, and very few contained original structures.

Initial hole angle surveys yielded ambiguous results. A survey on Core 33 showed an angle of 20° to 25° from vertical; as a consequence we decided to pull up and respud a new hole, 434A, about 50 feet from the first site without retrieving the drill string.

Hole 434A was spudded in on 17 September at 0615 hours at a water depth of 5988.5 meters. A mudline core was taken and the hole deepened with the center bit in place to about 150 meters, where an angle survey was made and a second core taken. The indicated departure from vertical was low, so we continued drilling. However, tropical storm Emma, which had been stalled about 300 miles due south of us, began to move northward at increasing speed. By evening the northern edge of the storm reached us, and we decided to recover the drill string after only 185 meters penetration subbottom. The pipe was hauled up under extremely difficult sea and wind conditions.

On 22 September, after completing drilling at Site 435, we returned to Site 434 to begin Hole 434B. The beacon we had dropped on 12 September was still operating, so we located over it and began running in. A mudline core was cut at 5996 meters below the drill deck, and then the center bit was inserted and we drilled to 295.5 meters sub-bottom. The solid bit drilling went very slowly. A second core was cut at 295.5 meters and an angle survey made. The survey showed only a 4° departure from vertical.

As with Holes 434 and 434A, drilling and coring conditions were difficult, with low penetration and low recovery rates. We assume the extremely fractured and fissile nature of the semilithified sediments to be the principal cause of the difficulties. What sections of intact core were obtained were highly fractured mudstone. We also noted a cyclic variation in the core recovery: Relatively good recovery was obtained for two or three



Figure 1. Multichannel seismic profile across Sites 434, 440, and 441 extending from the trench floor to the midslope terrace. Lithologic subdivision is superimposed on Site 434. O designates the oceanic basement reflector and M the Mohorovicic discontinuity. The multichannel profile was provided courtesy of JNOC.

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Figure 2. Enlargement from Figure 1 of the multichannel seismic profile at Site 434.

cores in succession, then recovery dropped to nearly zero for three or four cores in a row. This performance suggests that we penetrated alternating zones of relatively competent material and highly fractured material, respectively. Sometimes when the bit was raised free of the bottom, sediment would sluff from the hole walls and fill the hole below the bit.

After achieving a total of 637 meters sub-bottom we began retrieving the string on the evening of 28 September. The bottom hole assembly became stuck after we had pulled up approximately 300 meters of pipe, and we spent nearly 16 hours trying to free the pipe. Apparently the hole had caved above the bottom hole assembly. Finally, we decided to sever the pipe just below the mudline with an explosive charge. This was done successfully on the first attempt. We left the bottom hole assembly and about 150 meters of pipe in the hole.

# LITHOLOGY

The sediments at Site 434 are dominantly a uniform diatomaceous clay. Variation in ash content is the only principal change downhole. Two units are differentiated on the basis of ash content as determined by smear slide descriptions: The sediment changes from an overlying section with less than 10 per cent ash (Unit 1) to an underlying section with over 10 per cent ash (Unit 2). Two sub-units are recognized within Unit 1: diatomaceous clay (Sub-unit 1A) overlying diatomaceous claystone (Sub-unit 1B). Unit 2 is subdivided into three sub-units differentiated by the abundance of virtic ash: Sub-unit 2A is diatomaceous vitric claystone with between 10 and 30 per cent ash, Sub-unit 2C, like 2A, is diatomaceous vitric claystone with 10 to 30 per cent ash (see Figure 3).

# Unit 1, Sub-unit 1A (Cores 434-1-11; 434A-1; 434B-1, 0-101.5 m sub-bottom, Pleistocene-lower Pliocene)

Sub-unit 1A is a very uniform grayish-olive-green diatomaceous clay. X-ray diffraction data (Mann and Müller, this volume) shows that proportions of all mineral components in the sediment are relatively constant throughout the dominant lithology in this unit—much more constant than the smear slide descriptions suggest. Amorphous biogenic silica ranges approximately between 10 and 20 per cent; clays, between 40 and 60 per cent; volcanic glass is less than 10 per cent; quartz, about 10 per cent; feldspar, about 4 per cent; and there are minor and trace amounts of mica, heavy minerals, palagonite, glauconite, pyrite, forams, and lithic ash. The biogenic silica is dominantly diatoms, with varying amounts of radioloarians and spicules. Patches of ash and some ash beds are present in Cores 4 and 9. There are whispy laminae, convoluted by drilling disturbance, in much of the sediment, some dark pyritized spots, and a few small pumice fragments.

The texture of the sediments ranges from dominantly clayey silt at the top of the unit to dominantly silty clay at the base. As a result of the extreme disturbance by the drilling process there are no sedimentary structures preserved in Sub-unit 1A; the sediments are dominantly soupy to intensely disturbed, with a few sections moderately disturbed.



Figure 3. Lithologic and biostratigraphic zonation of Site 434.

The boundary between Sub-units 1A and 1B is a gradual transition from unconsolidated clays to claystone. The boundary was pinpointed between Cores 11 and 12 because there appeared to be a significant difference in the degree of lithification. This change is apparently supported by the physical properties data (see Carson and Bruns, this volume).

# Unit 1, Sub-unit 1B (Cores 435-12-435-33; 434A-2; 434B-2-435B-7, 101.5 – 352.5 m sub-bottom, lower Pliocene)

The dominant sediment in Sub-unit 1B is compositionally the same as Sub-unit 1A but differs in being more lithified (see Figure 3). Proportions of all components remain relatively constant downcore in the dominant lithology. The amount of amorphous biogenic silica is slightly higher than in Sub-unit 1A but still ranges between only 10 and 20 per cent. Clays range from 40 to 60 per cent; volcanic glass is less than 10 per cent; quartz, about 10 per cent; feldspar, about 4 per cent; and there are rare to trace amounts of mica, heavy minerals, pyrite, glauconite, forams, carbonate unspecified, and lithic fragments. The biogenic opal is dominantly diatoms, with varying amounts of radiolarians and sponge spicules. The sediment is gravish-olivegreen, dusky yellow green, grayish-green, and light olive gray. Ash patches and layers are present in Cores 18, 20, 24, 27, 28, and 33 in Hole 434 and in Cores 2, 3, 4, and 7 of Hole 434B. These are compositionally acidic-that is, in the silica range from dacite to rhyolite. Pebbles and cobbles, angular to subrounded and dominantly dacite with some marlstone, were found throughout the section. (The petrology of the pebbles is discussed in a subsequent section.) Cores 25 and 32 in Hole 434 contain layers of silty sand enriched with rounded to subrounded grains of quartz, feldspar, lithic fragments, and diatomaceous claystone fragments. A strong odor of H<sub>2</sub>S was present in all cores to a sub-bottom depth of 250 meters.

The texture of these sediments is dominantly silty clay with many intervals of clayey silt. The sediment is moderately to intensely disturbed by drilling, with some soupy zones; thus no structures are preserved. Zones of shearing may be indicative of *in situ* shearing and faulting, but because of similarities to drilling-induced shearing it is difficult to be definitive about their origin.

Transition from Sub-unit 1B to Sub-unit 2A is pinpointed on the basis of an increase in the abundance of ash to over 10 per cent.

# Unit 2, Sub-unit 2A (Cores 434B-8–434B-18, 352.5 – 456 m sub-bottom, lower Pliocene)

The dominant sediment in Sub-unit 2A is diatomaceous vitric claystone. It differs from Unit 1 only in increased ash content. The ash content varies but generally increases downsection in Sub-unit 2A, ranging from 10 to 30 per cent; thus the sediment type actually changes from vitric diatomaceous claystone at the top to diatomaceous vitric claystone at the base of the unit. Coincidentally, biogenic silica decreases, ranging from 20 to 10 per cent; quartz is about 9 per cent; feldspar, about 4 per cent; and there are rare and trace amounts of mica, heavy minerals, glauconite, carbonate unspecified, and calcareous nannos. The biogenic opal is dominantly diatoms with varying abundances of sponge spicules and rare amounts of radiolarians and silicoflagellates. The color ranges between dark greenish-gray, greenish-black, grayish-olive-green, and olive gray. There is only one patch of dacitic ash in Sub-unit 2—in Core 16. Marlstone, limestone, and pumice pebbles are found in Cores 8, 11, 13, 15, and 17. Patches of marlstone are present in Core 9, and there are diatomaceous marlstone patches in Core 16. A vitric siltstone layer is present in Core 9.

The texture of the sediment is dominantly silty clay with many intervals of clayey silt. The sediment is moderately disturbed to soupy with a few zones that are relatively undisturbed in Cores 8 and 9. There are some sedimentary structures preserved within Cores 8, 9, 10, 15, and 16 which show healed areas of microfaulting, what may be bedding, and some shear planes. There is a weakly graded zone of volcanoclastic sand in Core 8 and intervals that may indicate bioturbation in Core 16.

Transition from Sub-unit 2A to Sub-unit 2B is based on an increase in the abundance of ash to over 30 per cent. The boundary is designated between Cores 18 and 19, arbitrarily pinpointed on a trend of increasing ash which has significant short period variation.

# Unit 2, Sub-unit 2B (Cores 434B-19-434B-34-1, 456 – 601 m sub-bottom, upper Miocene)

Compared to the remainder of the section at Site 434, Sub-unit 2B is particularly enriched in vitric ash. The sediment is tuffite and over 30 per cent olive gray, dark greenish-gray, greenish-black, and grayish-olive-green ash. Ash content ranges between 30 and 60 per cent; amorphous biogenic silica is less than 10 per cent; clay ranges from 20 to 60 per cent; quartz is about 8 per cent; feldspar, about 3 per cent; and there are rare and trace amounts of mica, heavy minerals, pyrite, carbonate unspecified, and calcareous nannos. The biogenic silica is dominantly diatoms, with sponge spicules and radiolarians commonly present.

There are no patches or layers of ash in Sub-unit 2B. Limestone and marlstone pebbles and fragments are common. The texture of the sediments is either clayey silt or silty clay. They are intensely disturbed to soupy and contain no preserved structures.

Transition from Sub-unit 2B to Sub-unit 2C is defined by the volcanic ash content decreasing to less than 30 per cent. The transition is relatively distinct and the boundary has been placed between Sections 1 and 2 of Core 34.

# Unit 2, Sub-unit C (Cores 434B-34-2-434B-36, 601 – 628 m sub-bottom, upper Miocene)

Sub-unit 2C is dominantly diatomaceous vitric claystone, compositionally defined by the decrease in ash content from that of the overlying unit. The relative composition of the other components remains the same. The vitric ash ranges between 10 and 25 per cent in abundance but averages 20 to 25 per cent. Amorphous biogenic silica makes up about 10 per cent of the sediment; clay ranges from 25 to 50 per cent; quartz is about 9 per cent; feldspar, about 4 per cent; and there are trace and rare amounts of mica, heavy minerals, pyrite, and carbonate unspecified.

There are limestone and marlstone pebbles in Cores 35 and 37. The texture of the sediment is silty clay. No sedimentary structures are observed because the recovered section was soupy.

# **Petrology of Pebbles**

Pebbles of andesite, dacite, volcanic breccia, gabbro, marlstone, and calcareous vitric tuff were recovered. Most of the igneous rock pebbles are subrounded to well-rounded and occur isolatedly.

Among andesitic pebbles, Sample 434-25-2, 86-90 cm, is hornblende-augite andesite. The phenocryst consists of plagioclase, augite, and hornblende. Plagioclase is euhedral to subhedral, generally 0.8 mm  $\times$  1.5 mm in size, An 35-37 per cent in composition, and replaced mainly by kaolinite and partly by epidote. Augite, about 0.6 mm long, is euhedral to subhedral and is altered to epidote and chlorite. Hornblende, 0.9 mm long, is also mostly altered to epidote. The groundmass, composed of plagioclase laths (<0.2 mm), magnetite, chlorite, and augite, shows mostly hyalopilitic and partly intergranular texture. Small apatite crystals are common as accessories.

Sample 434-27-1, 65 cm, is hypersthene-augite (Figure 4). Phenocrystal plagioclase,  $0.7 \times 1.3$  mm at the maximum, is sporadic in occurrence and mostly replaced by kaolinite. The composition of plagioclase falls in at 35 per cent. Pyroxene phenocrysts (0.3 - 0.6 mm) are altered to chlorite. The groundmass consists of plagioclase laths (< 0.2 mm), magnetite, augite, and glass, showing hyalopilitic texture. It is as a whole intensely silicified and marked by chalcedony veinlets which are partly replaced by calcite.

Sample 434-25-2, 86-90 cm—from the same interval as the aforementioned andesite pebble—is intensely altered dacite (Figure 5). Phenocrysts include plagioclase,



Figure 5. Photomicrograph of altered dacite pebble (Sample 434-25-2, 86-90 cm) crossed Nichols. Bar scale equals 0.1 mm.

hornblende, and biotite, of which plagioclase (0.6–1 mm) is altered to kaolinite and hornblende is replaced by magnetite and biotite along the margin. The intensely silicified groundmass may have been originally vitreous, in which plagioclase laths (0.2 mm) are rarely traceable, though much altered.

A pebble of volcanic breccia—again at the same level as the andesite and dacite (Sample 434-25-2, 86-90 cm [3]; Figure 6) consists of fragments of basalt (2-4 mm), pumice (0.6-3 mm), and dacite (6 mm), in that order of abundance. Basalt bearing plagioclase phenocrysts (0.5-1 mm long) shows pilotaxitic texture with augite, magnetite, and olivine. Some of the olivine grains are altered to serpentine and magnetite; pumice contains brown hornblende (0.3 mm) and augite (0.2-0.6 mm); dacite has fluidal subvitreous texture containing oligoclase microlites, in which a few phenocrysts of plagioclase (0.6 mm) and augite (0.2-0.6 mm) are set.

Tuff breccia (Sample 434B-24-1, 87-90 cm) shows that dark-colored fragments of diatom-rich, vitric tuff are cemented by sparry calcite with lots of coarse glass



Figure 4. Photomicrograph of hypersthene-augite andesite pebble (Sample 434-27-1, 65 cm) plane-polarized light. Bar scale equals 0.1 mm.



Figure 6. Photomicrograph of volcanic breccia (Sample 434-25-2, 86-90 cm) plane-polarized light. Bar scale equals 0.1 mm.

shards. Sparry calcite veins are developed. Another tuff breccia sample (434B-31,CC, 16–17 cm) shows that large and small fragments of vitric tuff are set in micrite.

Pebbles of tuff consist of hornblende-biotite dacitic tuff (Sample 434-15-1, 120-124 cm), plagioclase vitric tuff (434-19,CC), augite-plagioclase-biotite tuff (434A-1-4, 62-64 cm), and diatom-bearing, calcareous vitric tuff (434B-25-1, 97-104 cm; 434B-28-1, 115-117 cm; 434B-29,CC, 0-3 cm; 434B-34-1, 7-8 cm).

We detected a well-rounded pebble of augite gabbro (Sample 434B-34-1, 6-8 cm). It shows fresh ophitic texture (Figure 7). Plagioclase consists of euhedral to subhedral, prismatic crystals (larger ones 1.8 mm  $\times$  2.1 mm; on average 1 mm long). They are An 51 per cent in composition and present simple and complex forms of twinning. Augite is subhedral to anhedral and is partly replaced by chlorite. Magnetite occurs as accessories in close association with augite.

### BIOSTRATIGRAPHY

Site 434 is dominated by diatom remains and contains a lesser amount of radiolarians and very scarce foraminifers. The composite section of Holes 434 and 434B (0 -610 m) is Pleistocene to late Miocene. All three fossil groups identify a hiatus in Hole 434 between the late Pleistocene (Core 1) and the late Pliocene (Core 2). The vertical distribution of different fossil assemblages below 54 meters displays a cyclicity which the diatom data identify as a complex repetition of early Pliocene to late Miocene age sediments. The low abundance and low age resolution of the foraminifers and radiolarians in this interval do not permit conclusive support for the diatom data. The possible causes for the assemblage repetition are discussed in the diatom section. Mineralogical data on the crystallinity of montmorillonite are discussed in Mann and Müller (this volume). Harper (this volume) also supports the cyclical nature of the sediments at Site 434, with the major breaks in crystallinity trends matching major breaks in the diatom assemblages.



Figure 7. Photomicrograph of augite-gabbro pebble (Sample 434B-34-1, 6-8 cm) crossed Nichols. Bar scale equals 0.1 mm.

#### Diatoms

The samples studied for diatom biostratigraphy from the composite section of Holes 434 and 434B reveal a complex distribution of zonal assemblages. The diatoms are abundant to common and well preserved in the upper 200 meters and become scarce and poorly preserved below. The low abundance and poor preservation of diatoms in the deeper samples lessen the possibility of accurate biostratigraphic assignment.

In Hole 434, Cores 1 through 12 are a normal sequence of progressively older diatom datum levels from late Pleistocene to early Pliocene, with a hiatus of about 1.5 m.y. between Cores 1 and 2; in the latter the *Rhizosolenia curvirostris* and *Actinocyclus oculatus* zones of early Pleistocene age are absent. The *Denticula seminae fossilis, D. seminae fossilis–D. kamtschatica, D. kamtschatica,* and *D. hustedtii* zones represent a continuous diatom sequence from late Pliocene to late Miocene. Beginning in Core 13, the vertical distribution of diatom zonal taxa is complex and difficult to interpret unambiguously throughout the rest of Hole 434 and Hole 434B. The normal age sequence of diatom zones does not occur. Below 54 meters the main zonal taxa show a repetitive alternation, as indicated in Table 1.

Because of these alternations it is impossible to locate geologic time boundaries with any meaning without assuming the cause for the repetitions. There are four probable causes: (1) downhole contamination; (2) faulting (thrusting); (3) slumping or sliding, and (4) climatic cycles. Figure 8 gives the time versus depth relations of the different probable explanations. Cause (1) places the zonal boundaries at the highest occurrence of the marker datums, whereas (3) and (4) place the younger zonal boundaries at their lowest occurrences. Cause (2) involves a complex reconstruction of each repetitive rock package. Continuous sequences of diatom marker datums occur from late to early Pliocene (8 - 54 m), early Pliocene to late Miocene (158 - 253 m), and early Pliocene to late Miocene (300 - 609 m). From 54 to 158 meters and from 253 to 300 meters, instead of continuous sequence there is an alternation between two early Pliocene assemblages on a scale of a few tens of meters.

# Radiolaria

The sediments recovered from Site 434 contain few to scarce and moderately to poorly preserved radiolarians, except for Core 1 of Hole 434, which has abundant and well-preserved specimens.

		T	ABLE	1		
Diatom	Zones	in	Cores	from	Hole	434

Core	Zone						
13	Denticula seminae fossilis-D. kamtschatica						
14	D. kamtschatica						
15-16	D. seminae fossilis-D. kamtschatica						
17	D. kamtschatica						
18-19	D. seminae fossilis-D. kamtschatica						
20-27	D. kamtschatica						



Figure 8. Plots of age versus sub-bottom depth for four models to explain the repetition of diatom zones. (Circled numerals refer to the number of repetitions of a particular biostratigraphic datum.)

Hole 434, Core 1 contains Lamprocyrtis haysi and lacks both Stylacontarium acquilonium and Axoprunum angelinum, indicating that the sediments are younger than 0.4 m.y.B.P. (Botriostrobus aquilonaris Zone). Core 2, yielding Eucyrtidium matuyamai, and underlying Cores 3 to 6 and 7 to 33 are assigned to the Eucyrtidium matuyamai, Lamprocyrtis heteroporus, and Sphaeropyle langii zones, respectively. Thus there is a hiatus between Cores 1 and 2 encompassing the upper part of the Eucyrtidium matuyamai and Axoprunum angelinum zones.

Only two cores were obtained from Hole 434A. The upper, Core 1, is assigned to the *Axoprunum angelinum* Zone and contains many radiolarian specimens, but in the lower, Core 2, they are rare and without any indication of age.

The sediments recovered from Hole 434B contain only scarce radiolarians. In the upper and lower thirds (Cores 2 through 9 and 24 through 36) radiolarian specimens are especially rare.

Stichocorys peregrina is a predominant species throughout the cores; the lowermost occurrence of Sphaeropyle langii is in Core 35 and co-occurs with Stylacontarium acquilonium. Above this core, from Cores 17 to 34, both Sphaeropyle langii and Stylacontarium acquilonium are absent, and Ommatartus avitus (Core 15), O. penultimus (Cores 15, CC and 26-1), and O. hughesi (Cores 26-1 and 28, CC) are present. Therefore, Cores 2 to 16 and 17 to 34 are assigned to the Sphaeropyle langii and Stichocorys peregrina zones, respectively, and Cores 35 to 37 to the Sphaeropyle langii Zone again, although it is possible that the presence of S. langii and Stylacontarium acquilonius is due to downhole contamination.

Except for the interval between Cores 34 and 35 of Hole 434B, the distinct repetitions evident from the diatom analysis are absent from the radiolarian biostratigraphy. However, the radiolarian stratigraphy seems to agree with the scheme of the diatom data insofar as Cores 13, 18, and 19 contain more specimens with better preservation, typical of units higher in Hole 434. Other interval breaks revealed by analysis of diatoms and foraminifers are also absent.

## Foraminifera

There are three groups of foraminifers at Site 434. (1) Planktonic and middle bathyal calcareous benthonic foraminifera typically occur together, suggesting slumping, since the site is below the CCD (3500 m); (2) *Saccammina* and/or *Reophax* occur rarely near the top of the site; and (3) *Martinottiella communis* disappears at the top of Core 7 in Hole 434 near the early/late Pliocene boundary.

According to diatom and radiolarian data and a single specimen of *Neogloboquadrina eggeri*, Core 1 is late Pleistocene. Cores 2 and 5 are late Pliocene and contain *Neogloboquadrina asanoi* and *Globorotalia inflata praeinflata*. Cores 6 through 11 are barren of planktonic foraminifers. Core 13 has fauna similar to Core 2. Cores 14 and 15 are barren, and Core 16 is late Pliocene. Cores 19 through 31 are barren, and Cores 32 and 33 are late Pliocene. Although these data lend support to the hypothesis of repeated sections, the extremely low frequencies of foraminifers adds very little by way of conclusive evidence. Hole 434A is barren of foraminifers, and Hole 434B contains little more than stained and very poorly preserved specimens from Core 10.

### PHYSICAL PROPERTIES

At Site 434, compressional sound velocity, wet bulk density, water content, porosity, undrained shear strength, and thermal conductivity were measured on the majority of the cores recovered. A summary of the shipboard techniques for determining these parameters is included in the introduction to this volume, and a more detailed discussion (except for thermal conductivity) is presented by Boyce (1976). A listing of all data is in the appendix to Part 1, this volume.

Data from Holes 434, 434A, and 434B are considered together. In general, the condition of the cores at this site did not allow for frequent or high-quality determinations of physical properties. Recovery was sparse because of the fissile nature of the sediment. Many cores consisted primarily of drill cuttings and sediment slurry, and most were highly disturbed. Velocity determinations may be inaccurate, because even when samples were suitable for analysis, they frequently shattered during preparation for velocity measurements or from the slight pressure of the transducer heads during measurement. Velocity determinations were further affected by severe attenuation of the sonic signal through the sediment, except on samples less than about two centimeters thick - probably because of microfractures. Similarly, gravimetric techniques may be affected by voids due to hairline fractures that create errors in the volume determination or that allow influx of water prior to measurement of water content. Much of the GRAPE data measured on cores still in liners may also be degraded because of fracturing and the highly disturbed nature of the cores. In addition, most samples for velocity and gravimetric determinations were less than optimum size.

Plots of the physical properties data with depth are shown on the Site 434 Site Summary Chart (back pocket), and average values for each lithologic unit are given in Table 2. Although the data may not be representative of *in situ* values, some changes in physical properties with depth are apparent.

 TABLE 2

 Correlation of Physical Property Data with Lithology at Site 434

Litho- logic Sub- unit	Depth Interval (m)	Interval Velocity (km/s)	Density (Mg/m <sup>3</sup> )	Porosity (%)	Cumulative Two-Way Travel Time (s)		
1A	0-101.5	1.53 + 0.02	1.42 + 0.02	64 ± 6	0.132		
1B	101.5-352.5	$1.69 \pm 0.07$	$1.59 \pm 0.13$	$56 \pm 10$	0.429		
2A	352.5-456	$1.64 \pm 0.07$	$1.62 \pm 0.07$	$59 \pm 5$	0.555		
2B	456.0-601	$1.89 \pm 0.14$	$1.82 \pm 0.11$	$45 \pm 7$	0.708		
2C	601.0-628	$1.92 \pm 0.12$	$1.92 \pm 0.12$	$51 \pm 1$	0.736		

Note: Values for velocity, density, and porosity are averages of measured values over each lithologic unit. The two-way travel time is calculated from the interval velocity and indicates the seismic reflection time to the base of the lithologic unit.

### Velocity

From a sub-bottom depth of 0 to 80 meters, few measurements were possible; the measured velocity is between 1.50 and 1.55 km/s. Between 80 and 100 meters, there is an abrupt increase in velocity, to about 1.65 km/s, and velocity ranges from 1.65 to 1.70 km/s up to 290 meters sub-bottom. The velocity increase between 0 and 290 meters is probably due to increasing consolidation within Sub-units 1A and 1B. The second core of Hole 434B, at 300 meters, shows an abrupt increase to 2.27 km/s. The increase is suspect, however, because only material from the core catcher was recovered, and the samples measured may be erratics rather than formation sediment. However, the increase does correspond to changes in density, water content, and porosity in adjacent cores from Holes 434 and 434A and thus may indicate the presence of a relatively thin higher velocity layer. Below this layer, velocities average 1.64 km/s in Sub-unit 2, which may indicate underconsolidation of the sediment with respect to depth and reflect either the change in lithology or the effects of increasing fracturing in sediment. Velocity increases markedly, to about 1.82 km/s in Sub-unit 2B, and reaches a maximum of 1.8 to 2 km/s at the bottom of the hole. If the apparently low velocities between 352 and 456 meters are due to fracturing of the sediment, the rapid increase to velocities of 1.8 to 2.0 km/s may indicate healing of fractures with increasing overburden pressure.

### Wet Bulk Density

GRAPE and gravimetrically determined wet bulk densities are plotted together on the Site 434 Site Summary Chart, (back pocket). Continuous GRAPE data were generally obtained on only one or two sections of a core and are highly interpretative averages of the maximum densities indicated on the analog records. Considerable variation is present on the GRAPE analog records, and the plotted values may be significantly in error.

In spite of the foregoing problems, there is moderately good agreement between the continuous GRAPE and the gravimetric densities; the former generally show lower density values, which probably reflects the fissile nature of the sediment and the disturbed nature of most of the cores. Only limited determinations by the twominute GRAPE method were made because of the small size of most of the available samples; these data also compare moderately well with the gravimetric data.

In general, the data show a gradual increase in wet bulk density from about 1.4 Mg/m<sup>3</sup> at the surface to about 1.85 Mg/m<sup>3</sup> at the base of the cored section, and variations from this trend correlate well with changes in velocity. There is a poorly defined increase in density from 1.4 to 1.6 Mg/m<sup>3</sup> corresponding to increasing consolidation from Sub-unit 1A to 1B and an increase between 280 and 320 meters corresponding to the possible thin high velocity zone. The average density in Sub-unit 2A between 350 and 450 meters is only slightly higher than in Sub-unit 2B, suggesting that the velocity low is accompanied by densities that are low with respect to depth. An abrupt density increase in Sub-unit 2B between 450 and 600 meters corresponds to the velocity increase. The apparent scatter between about 1.7 and 1.95 Mg/m<sup>3</sup> below 520 meters may indicate lithologic variations within Sub-units 2B and 2C.

### **Porosity and Water Content**

Scatter in the porosity and water content data is rather large, but the data generally show an inverse correlation with velocity and wet bulk density. The water content decreases gradually from 55 per cent at the surface to 20 per cent at the base of the cored section, and porosities decrease from around 70 per cent to around 40 per cent. The porosity values derived from the continuous GRAPE techniques are markedly higher than those measured by gravimetric techniques — probably because of the pervasive fissility, which imparts a secondary porosity to the sediments.

#### Vane Shear Strength

Only five vane shear strength determinations were made because of poor recovery in the upper part of the holes. They show an increase from 4 kPa to 54 kPa.

#### **Thermal Conductivity**

Thermal conductivity determinations were made on cores that allowed a valid measurement. The objective of this work was to learn more about the thermal properties of earth materials that make up the accretionary wedge of a subduction zone and to correlate these measurements with other physical properties such as seismic velocity and density. Measurements were made with both a standard heated line source needle probe device and a Showa Denko QTM (Quick Thermal Meter) device.

Most of the cores from Hole 434 had been so mechanically disrupted that it was impossible to make measurements. Nonetheless some measurements were attempted, and the results are shown in Figure 9. In the upper 10 meters, values were obtained that are typical for deep sea oozes. In general, because we had not solved some of the problems of making reliable measurements with the QTM device on water-saturated sediment, the results with this device are not as reliable as at Sites 435 or 436 and at the Leg 57 holes. The needle probe measurements are considered accurate to about  $\pm 5$  per cent.

The cored material was in such broken condition that no needle probe measurements could be made at Hole 434B. Although the sawed faces of some of the cores could be measured by the QTM device, these measurements are sparse and very little can be said about their variation with depth or correlation with lithology. The conductivity of these fractured vitric mudstones and tuffites falls within the range of most surface sediments. The lithification indicated by the increased hardness of the rocks at depth has little detectable effect on the thermal conductivity values.



Figure 9. Thermal conductivity versus sub-bottom depth, Site 434.

#### Discussion

Because of the poor core recovery and disturbed condition of the cores, measurements are sparse, scatter is large, and the similarity to *in situ* physical properties is suspect. The physical properties correlate well internally; increases in velocity are generally matched by increases in density and decreases in porosity and water content. Physical properties in the top 350 meters of the section are affected primarily by increasing consolidation. Below 350 meters, velocity and density values that are apparently low for their depth of burial may reflect either the change in lithology or the disturbed nature of the sediment. However, the abrupt increase in average density and velocity from Sub-unit 2A to 2B, with no marked change in lithology, suggests that the low average values may be due primarily to fracturing of the sediment, with the increase below 450 meters suggesting a return to normal consolidation processes or healing of the fractures with increasing overburden pressure.

The generally good internal correlation of the data suggests that, given the very low recovery, future drilling in similar conditions should utilize coring techniques that cause less disruption. Logging the holes would also aid the lithologic interpretation and is strongly advised.

Table 2 gives the lithologic units and average velocity values used to calculate two-way reflection times to lithologic boundaries. Figure 2 shows the correlation of these boundaries with the multichannel seismic line over Site 434. There is a discontinuous series of short, nearly horizontal reflectors that can be traced through Site 434, but there is no clear correlation of these reflections with changes in the physical properties.

#### GEOCHEMISTRY

#### **Inorganic Chemistry**

Shipboard interstitial water measurements of salinity, chlorinity, magnesium, calcium, and alkalinity are shown in Figure 10. Also shown are levels of methane in core gas pockets, percentage of calcium carbonate in sediments, and an estimate of hydrogen sulfide—determined qualitatively by odor.

Carbonate levels in these sediments are generally very low (1-3%), as is expected at this water depth. The paleontologists also noted a general absence of calcareous fossils in the sediments. There are irregular zones of higher values throughout the core which may represent precipitated carbonate. This hypothesis is supported by high interstitial water alkalinity values which decrease slightly with depth but remain between 7 and 5 times higher than surface sea water value.

There is some question about the reliability of the interstitial water chlorinity values for the holes at Site 434, since the values for surface sea water are low (see Figure 10). A normal surface sea water chlorinity value was obtained just before Hole 434B interstitial water values were determined.

It is not clear why such abrupt changes in interstitial water alkalinity should occur in the upper part of Hole 434. Considering the doubt concerning the chlorinity data, we must question whether these fluctuations are real or whether they represent a measurement problem. However, the trends shown in the Site 434B values are very clear. The high alkalinity levels (7 to 5 times surface sea water values) together with methane gas pockets and evidence of black iron sulfide minerals are consistent with the activities of sulfate-reducing bacteria and methane-producing bacteria (see Organic Chemistry section and Whelan and Sato, this volume).

Trace elements analyzed in these sediments include lead, tin, copper, nickel, zinc, silver, boron, iron, manganese, and titanium (see Kurnosov et al., this volume, and Murdmaa, this volume). The same authors also measured  $SiO_2$  both in the amorphous state and in the clay matrix.



Figure 10. Some chemical parameters of interstitial water analysis plotted versus depth, Site 434. (• = Hole 434, • = Hole 434A, × = Hole 434B, • surface sea water values — Hole 434 —  $\boxtimes$  = surface sea water values — Hole 434B.)

#### **Organic Chemistry**

The sediments were anoxic, as indicated by a strong  $H_2S$  odor present to a depth of about 180 meters, high interstitial water alkalinity despite absence of calcareous minerals, and the presence of methane in core gas pockets below 100 meters. The methane and associated traces of  $C_2$ - $C_5$  hydrocarbons are discussed later in this volume (Whelan and Sato). The methane is presumed to be biogenic for the following reasons:

1) Methane level increases at the approximate depth at which  $H_2S$  decreases. It is known that bacterial methane production normally occurs below the zone of bacterial  $SO_4^{-}$  to  $H_2S$  reduction (see, for example, Claypool and Kaplan, 1974).

2) Organic carbon levels in the sediment—generally 0.5 per cent to 1.0 per cent (see Site Summary Chart, back pocket)—are sufficient to support bacterial growth.

3)  $C_1/C_2$  ratios of 1000 to 300,000 are typical of  $C_2$  accompanying biogenic methane (Claypool, 1976). Below 130 meters, log  $C_1/C_2$  plotted as a function of depth shows a general linear decrease similar to that of biogenic methane found in many other anoxic DSDP cores (see Claypool, 1975, and Whelan and Sato, this volume).

4) The general behavior of log  $(C_2/C_5)$  level plotted as a function of depth is similar to that at other DSDP sites where these compounds accompanied biogenic methane (see Whelan and Sato, this volume).

5)  $C_{13}$  values of -67.6 to -80.9 are typical of biogenic methane (see Whelan and Sato, this volume).

6) The high interstitial water alkalinities are typical of regions where bacterial sulfate reduction and methane production have occurred. Depletion of  $SO_4^{-1}$ 

in interstitial water as a result of bacterial reduction to sulfide with concurrent oxidation of organic matter can be represented simply as:

 $2 \text{ CH}_2\text{O} + \text{SO}_4^- \rightarrow \text{H}_2\text{S} + \text{HCO}_3^-$ . (Sayles and Manheim, 1975)

7) The low geothermal gradient throughout this area and the immaturity of the sediments (see Sato and Whelan, this volume) make petrogenic methane production unlikely.

8) Occurrence of rapid deposition at this site (about 80 m/m) is favorable to preservation of reduced organic matter and, thus, to anaerobic bacterial activity.

Pyrolysis-fluorescence was measured by a method described previously (Ryan and von Rad et al., 1978). Briefly, it involves heating the sediment in an open test tube over a flame, allowing the sediment to cool, extracting with trichloroethane, and measuring the fluorescence of the resulting solution. Low fluorescence values with moderate amounts of organic carbon (0.5%-2.0%) indicate a sediment with low petroleum-generating potential. Generally, values only slightly above blank values were obtained at this site. In spite of the low values, there does seem to be a rough correlation with organic carbon, particularly in the lower part of the hole, as shown in Figure 11. Moderate pyrolysisfluorescence values representing immature sediments capable of some hydrocarbon generation were obtained from some of the shallower samples from Site 434namely 434-4,CC, 434-20,CC, 434-21,CC, and 434-23.CC.

A general decrease of pyrolysis-fluorescence values with depth occurs at this site. This decrease occurs in spite of a fairly constant organic carbon content. In the



Figure 11. Pyrolysis-fluorescence versus organic carbon for Holes 434 and 434B. (Pyrolysis-fluorescence is measured in fluorescence units per 3 ml per 0.2 g sediment.)

lower parts of the hole, pyrolysis-fluorescence minima seem to correlate to some extent with maximum neopentane levels. A possible explanation is that the fractured nature of the sediments has allowed movement of water and bacteria resulting in removal of reduced organic carbon (see Whelan and Sato, this volume).

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SITE 434







SMEAR SLIDES (%)

Quartz Feldspar Mica

Heavy mins. Clay

Volcanic glass

Glauconite Diatoms

Radiolarians Sponge spicules Pyrite 1-12 1-91 CC-20

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60 50 4

3 TR TR

10 8 3

10 10







SITE 434













SITE 434





SITE 434



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Clay

Pyrite

Volcanic glass

Glauconite

Carb. unspec.

Calc. nannos

Radiolarians

Sponge spicules

Distoms



SITE 434



SITE 434

Heavy mins. Volcanic glass

Carb. unspec. Diatoms 25 60



TR TR

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Hole 434B







