

60. GEOPHYSICAL PROFILES AND NAVIGATION, DEEP SEA DRILLING PROJECT LEG 62, CENTRAL NORTH PACIFIC OCEAN¹

David K. Rea, Oceanography Program, Department of Atmospheric and Oceanic Science,
University of Michigan, Ann Arbor, Michigan,
Tracy L. Vallier, Pacific-Arctic Branch of Marine Geology, U.S. Geological Survey, Menlo Park, California
and
Jörn Thiede, Department of Geology, University of Oslo,
Blindern, Oslo 3, Norway

ABSTRACT

Leg 62 of the Deep Sea Drilling Project traversed about 3900 nautical miles, from the Marshall to the Hawaiian Islands, collecting bathymetric, air-gun seismic-reflection, and magnetic data. Major geological features along the track line include the Marshall Island pedestal, Central Basin, Mid-Pacific Mountains, North Pacific Basin, and Hess Rise.

INTRODUCTION

Leg 62 of the Deep Sea Drilling Project traversed about 3900 nautical miles between Majuro Atoll in Micronesia and Honolulu, Hawaii. The D/V *Glomar Challenger* left Majuro on July 29, 1978, and arrived in Honolulu on September 6.

We collected 12-kHz, 3.5-kHz, and air-gun continuous seismic-reflection profiles and total-field magnetometer measurements along the track lines. Progressive distance along the track, dates, hours, site numbers, and major bathymetric features are shown in Figure 1. Underway navigation information (Table 1) and plots prepared by the Geological Data Center, Scripps Institution of Oceanography, were used to locate the track. Satellite fixes, courses, and speeds were encoded on the ship from data in the underway geophysical log. The data were keypunched on shore, run through a navigation smoothing program, and subsequently edited on the basis of reasonable ship drift velocities and directions. The corrected navigation data were merged with the bathymetric and magnetic data. Satellite navigation errors are generally less than 1 nautical mile (Talwani et al., 1966).

Figures 2 through 8 show depth, distance along the track, dates and times, major bathymetric features, and magnetic anomalies. Figure 9 (plate in back pocket of this volume) includes photographic reproductions of original 5-sec-sweep air-gun seismic-reflection records. Vertical exaggerations, therefore, are greater than on the 10-sec records used in most previous DSDP reports. Depths in two-way travel time (seconds) are given along the edges of the profiles, and times, dates, and headings are at the bottom. Distances along the track in hundreds of nautical miles are at the top of the profiles. Sites

drilled and major bathymetric features also are printed on Figure 9.

Sound sources used in generating the seismic-reflection data were two 40-in³ air guns. Incoming signals were filtered through a Bolt filter-amplifier system and then recorded on dry-paper EDO recorders. Fire rate was at a 10-sec interval at all times, and sweep rates were 10 sec (on EDO recorder 1) and 5 sec (on EDO recorder 2). Hull-mounted transducers were the sound source for the 3.5- and 12-kHz profilers.

The magnetic-anomaly profiles in Figures 2 through 8 were derived by removing IGRF from the total-field measurements. Magnetic data produced at sea on a Geometrics magnetometer were recorded in the underway geophysical logbook at 5-min intervals. Depths taken from echo sounders (calculated at 1500 m/sec calibrated sound velocity) also were recorded in the logbook at 5-min intervals. Both types of data were edited on shore by the Geologic Data Center, Scripps Institution of Oceanography.

GEOPHYSICAL NARRATIVE

Marshall Island Pedestal

Underway geophysical records for Leg 62 begin a few miles north of Majuro Lagoon. The first of the major geological features crossed was the lower east flank of the Marshall Island chain (Figs. 2 and 9, mile 0-120). The three peaks at about 0, 45, and 90 miles are the lower flanks of Majuro, Aur, and Maloelop Atolls, respectively. Sediments over 1.0 sec in thickness underlie the sea floor between Majuro and Aur. About half of this sediment package occurs in a deeply incised, semi-transparent upper layer (Fig. 9). The lower unit has a smooth upper surface and exhibits closely spaced reflectors. Three flat-floored, channel-like features occur along this section of the profile, and the broad base of Maloelop Atoll shows very little sediment accumula-

¹ Initial Reports of the Deep Sea Drilling Project, Volume 62.

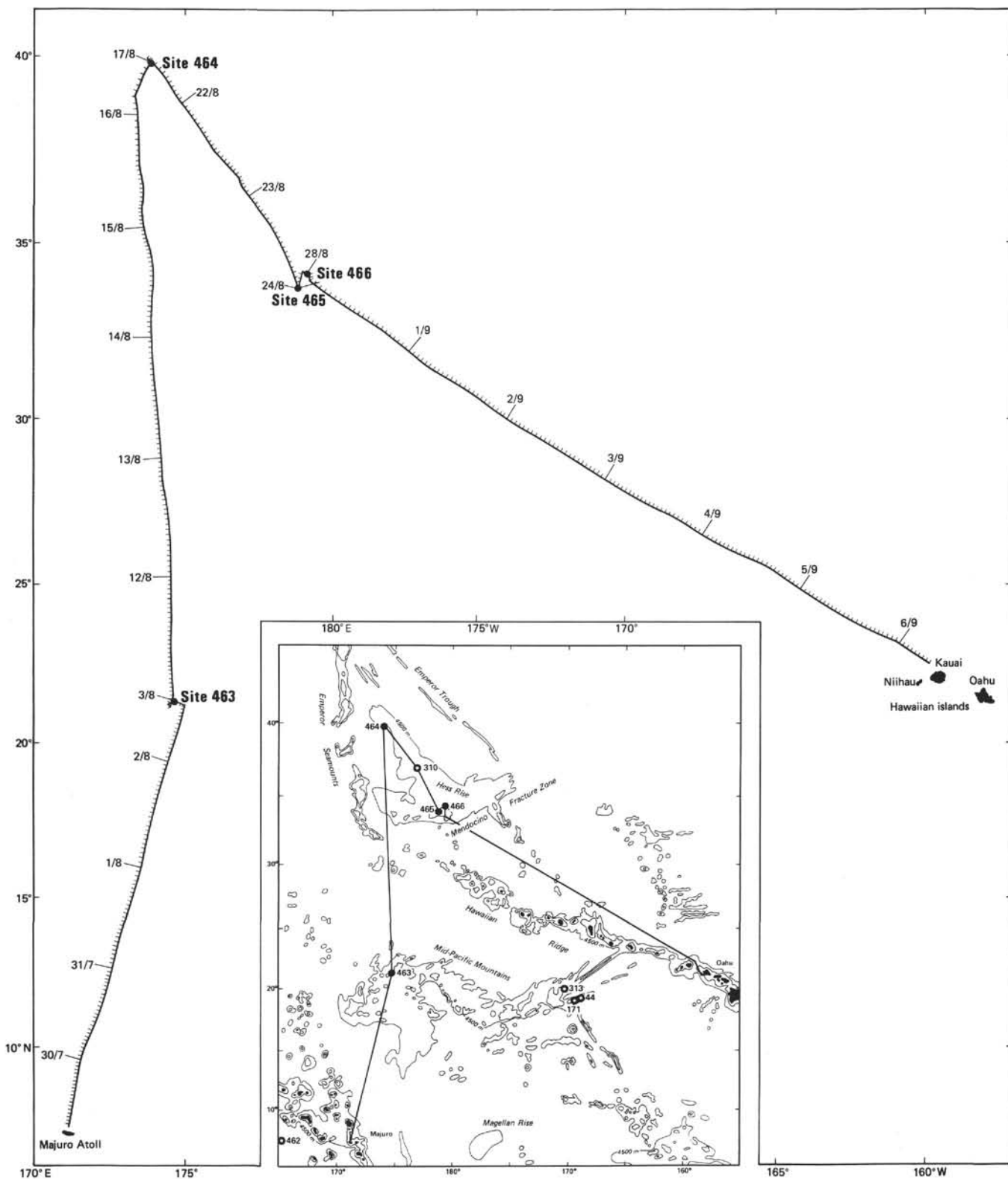


Figure 1. Map of the *Glomar Challenger* track line, DSDP Leg 62.

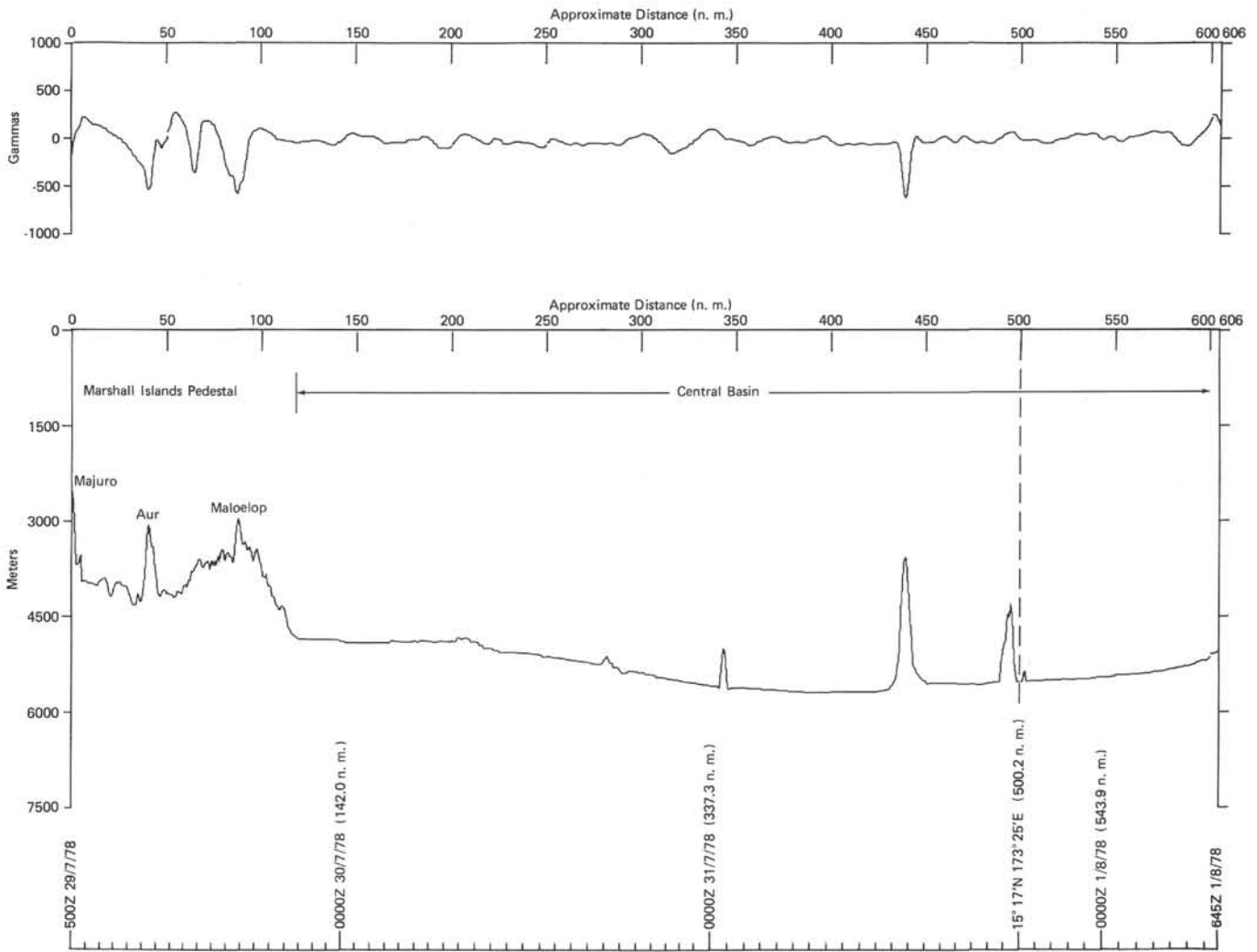


Figure 2. Bathymetric and magnetic-anomaly profiles along the track line of Leg 62. The bottom line shows hourly tic marks, days, and the positions of drill sites and some greater course changes. Distance along track line taken from Table 1 is given for each day, course-change, or site annotation. Numbers along the top of the profile give approximate distance along track line and are not in exact agreement with actual distance values from Table 1. Geological provinces, major bathymetric features, and DSDP drill sites are labeled.

ward to the base of the Mid-Pacific Mountains near mile 600. The gentle relief is pierced by volcanic peaks 600 to 2000 meters high, all of which appear to exhibit current scour around their bases (miles 340, 440, and 490). An offset of about 0.2 sec (fault?) occurs across the largest peak, at mile 440. A strong negative magnetic anomaly is associated with this largest peak.

Mid-Pacific Mountains

The Mid-Pacific Mountains (Figs. 3 and 9, mile 600-980) are the northern boundary of the Central Basin. This bathymetric feature exhibits two broad levels of topography, one at about 5.5 sec (4100-m deep), and the shallower one at about 3.5 sec (2600-m deep). Along our track line, the southern edge of the Mid-Pacific Mountains is marked by a large guyot which separates the lower plateau level (5.5 sec) of the Mid-Pacific Mountains from the Central Basin floor (Figs. 3 and 9, mile 620). The guyot rises to 1.7 sec (about 1300 m deep) and is approximately 21 km across at the top.

Sediments underlying the lower platform level are as much a 1.2-sec thick, exhibit alternating transparent and reflective intervals, and overlie a distinct acoustic basement. Portions of this sediment, especially near mile 750, are offset by closely spaced, near-vertical normal faults that have as much as 0.1-sec displacement. A thin layer of acoustically transparent sediment overlies the faulted section and fills in lows between elevated features.

The flat, lower platform (5.5 sec) gives way to the northern, upper level (3.5 sec) near mile 815 (Fig. 9). This part of the Mid-Pacific Mountains displays hummocky topography, with numerous hyperbolic side echoes, until mile 880; after which a sediment platform extends to the northern edge of the Mid-Pacific Mountains. On this sediment platform is Site 463 (Fig. 3).

North Pacific Basin, I

A part of the large North Pacific Basin lies between the Mid-Pacific Mountains and the southwestern exten-

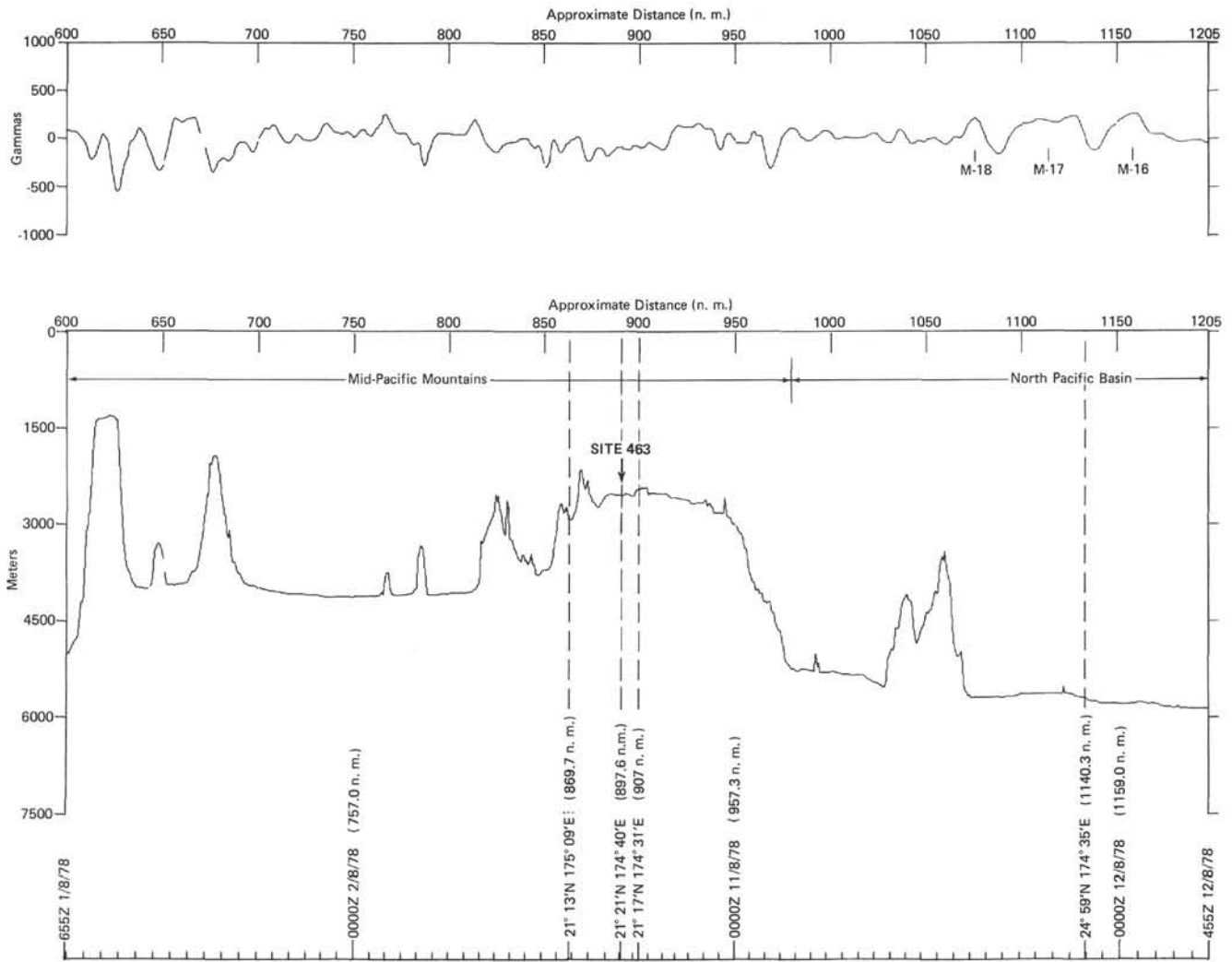


Figure 3. Bathymetric and magnetic-anomaly profiles along the track line of Leg 62. Anomalies labeled M-18, M-17, and M-16 are a portion of the Mesozoic magnetic-anomaly sequence. Annotations are the same as for Figure 2.

sion of Hess Rise (Figs. 3, 4, and 9, mile 980–1650). The basin floor slopes northward from the base of the Mid-Pacific Mountains to a point near mile 1300, where the sea-floor depth reaches nearly 8.0 sec (~ 6000 m). From mile 1300, the sea floor rises northward to the base of Hess Rise, interrupted by the Mendocino Fracture Zone at mile 1430 to 1510. Our track line fell between the seamounts near the Hawaii–Emperor bend, passing just east of Kanmu Seamount. The unnamed seamount near mile 1460, which has about 3.0 sec (~ 2250 m) of relief, lies near Jenkins Seamount.

Seismic-reflection profiles across the North Pacific Basin show three sedimentary layers: two well-stratified ones that underlie a thin transparent layer (e.g., near mile 1200). Maximum sediment thickness exceeds 0.5 sec. DSDP Site 311 at $28^{\circ}07.5'N$, $179^{\circ}44.3'E$ lies about 560 km east of our track line, in 5780 meters of water. At Site 311, 15 meters of red clay overlies a sequence of volcanic turbidite sands (Larson, Moberly, et al., 1975).

Linear magnetic anomalies between miles 1070 and 1400 (Figs. 3 and 4), south of the Mendocino Fracture Zone, correlate with anomalies M-11 to M-18 of the

Mesozoic magnetic anomaly sequence. These anomalies are believed to range in age from 126 to 139 m.y. (Larson and Hilde, 1975).

Southwestern Hess Rise

The southwestern extension of Hess Rise is a northwest-trending ridge that lies between mile 1650 and 1765 of the track line (Figs. 4 and 9). Hess Rise appears as a broad arch on the profiles, rising from the basin floor to 4.5 sec (~ 3400 m); a small peak rises above that general upper level of the rise. A series of normal faults with up to 0.5 sec of offset (mile 1660) characterize the southern slope of Hess Rise. Seismic-reflection profiles show two layers that display acoustic stratification. The upper layer, which reaches 0.4 sec in thickness, is only locally present. It becomes more acoustically transparent toward its base. The lower unit displays another 0.4 sec of sediment that shows continuous, well-developed acoustic reflectors.

North Pacific Basin, II

That part of the deep-sea floor bounded by the Emperor Seamounts to the west and Hess Rise on the

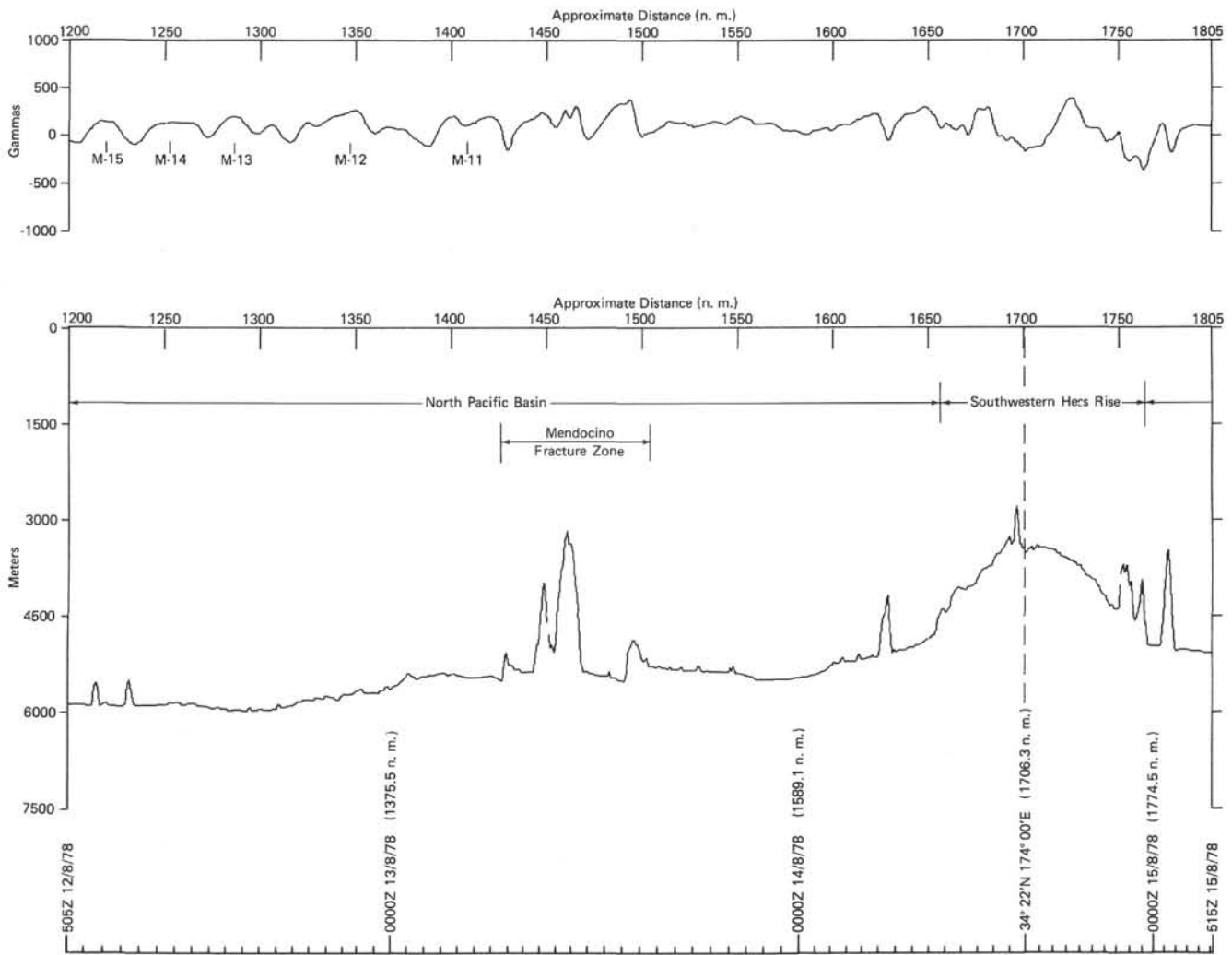


Figure 4. Bathymetric and magnetic-anomaly profiles along the track line of Leg 62. Anomalies labeled M-15 to M-11 are a portion of the Mesozoic magnetic-anomaly sequence. Annotations are the same as for Figure 2.

north, east and south (Figs. 4, 5, and 9, mile 1765–1985) lies at a regional depth of about 7.0 to 7.5 sec (5200–5600 m). The seismic-reflection profiles show a well-developed transparent layer about 0.1-sec thick overlying 0.25 sec of material having closely spaced, continuous reflectors. Beneath this layer is a unit as thick as 0.3 sec, characterized by strong, but more widely spaced reflecting horizons. Near mile 1950, a series of features characterized by less-coherent internal reflectors protrudes through the more-reflective horizons; the transparent layer conforms to the upper surfaces of these features.

Hess Rise

Hess Rise is a large plateau that stands 1000 to 3000 meters above the floor of the central North Pacific Ocean (Figs. 5, 6, and 9, mile 1985–2620). Based on our data and those from previous seismic lines, there are two distinct large-scale morphological trends on Hess Rise: (1) a north-northwestern-trending region at about 4000 to 4500 meters depth that extends from 34°N to 40°N; and (2) a west-southwest-trending southern edge.

This southern part stands about 2000 meters above the more northerly regions. Mellish Bank, the shallowest point of Hess Rise, comes to within 117 meters of the sea surface. The south-facing scarp that forms the southern margin of Hess Rise has a total relief of 4.5 sec (~3400 m) and coincides with the northern boundary of the Mendocino Fracture Zone.

Block faulting occurs across much of the rise and is responsible for most of the surface morphology. The offset along these faults is as much as 0.5 sec (~400 m); they are exceptionally well defined on Figure 9. Seismic stratigraphy of the northern part of the rise shows a transparent layer up to 0.1-sec thick overlying a well-stratified unit with more widely spaced reflectors. At Site 464, this lower, well-stratified unit is limestone. The lower acoustic layer becomes more transparent near Site 310. On the southern, upraised part of Hess Rise (Fig. 9), the seismic units are a surface layer 0.15 sec thick, exhibiting closely spaced, continuous reflectors, overlying a more-transparent unit about 0.25-sec thick near Site 465 (mile 2500). Sediments appear to be ponded in several small, fault-bounded troughs.

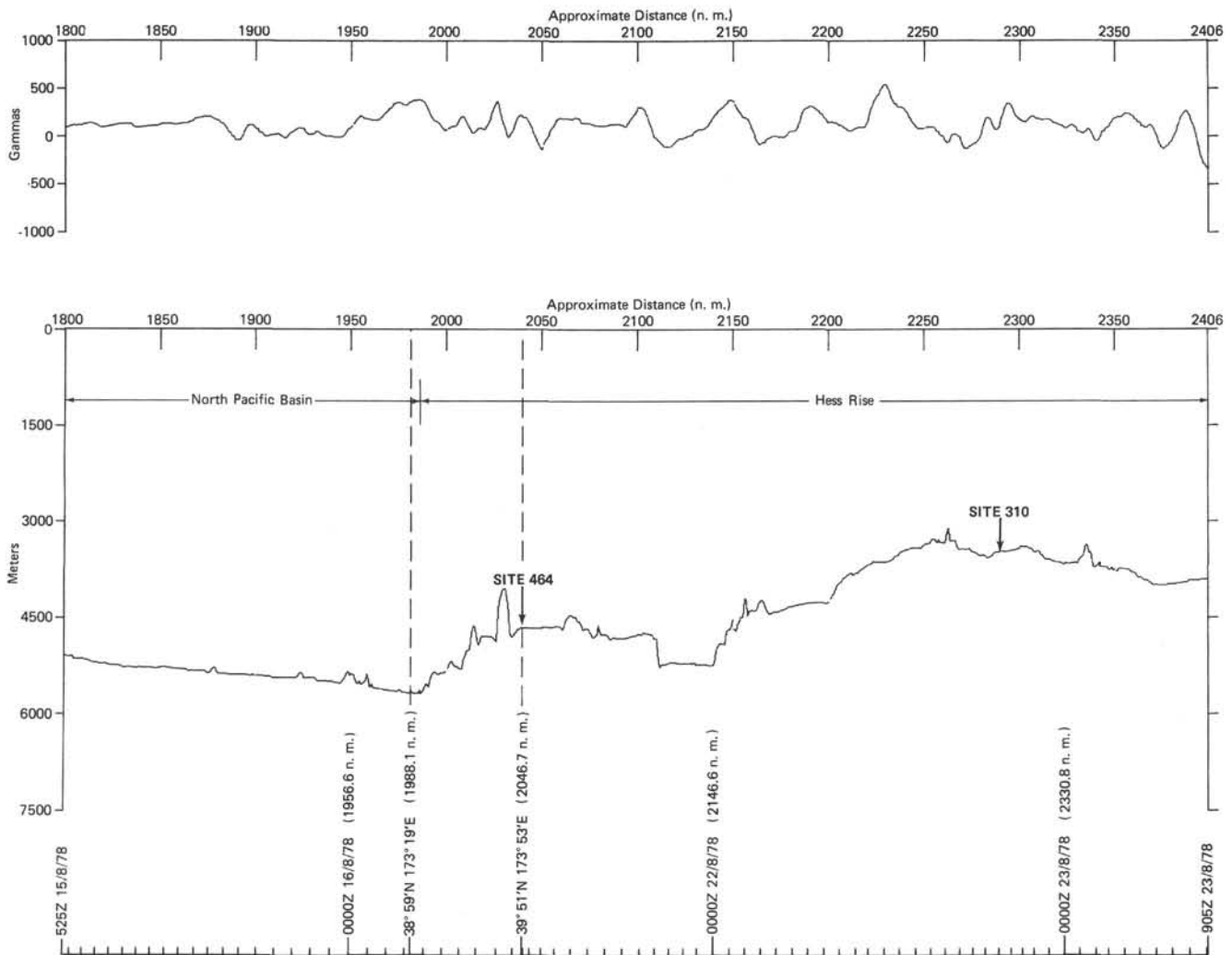


Figure 5. Bathymetric and magnetic-anomaly profiles along the track line of Leg 62. Annotations are the same as for Figure 2.

North Pacific Basin, III

Between Hess Rise and Hawaii, the ship traversed the North Pacific Basin, north of the Hawaiian Ridge (Figs. 6–9, mile 2620–3870). West of the seamount chain that extends northwest from Gardner Pinnacles to Hess Rise, the sea floor displays rough and irregular abyssal-hill topography (Figs. 6, 7, and 9, mile 2620–3020). Much of this topography, which has a relief of 0.2 to 0.3 sec with occasional larger features, appears to be fault controlled. From mile 3020 to the end of the track line, we crossed the Hawaiian Arch and Deep and the lower-most part of the Hawaiian Ridge, which are characterized by much smoother topography.

Two of the major North Pacific fracture zones, the Mendocino and the Murray, appear on our track line. The Mendocino (Fig. 6, mile 2620–2720) forms the southern boundary of the Hess Rise. The Murray fracture zone (Fig. 7, mile 3290–3385) is characterized by a single physiographic peak and a series of large magnetic anomalies. This magnetic signature is characteristic of the Murray fracture zone west of the Musicians Seamounts (Rea, 1970).

Seismic-reflection profiles reveal three sediment provinces. The sequence just south of Hess Rise (Fig. 9, mile 2640) is similar to that on the northern part of the rise. A thin, transparent layer overlies a weakly stratified unit and a coarsely stratified basal layer. Farther to the southeast, the upper layering gives way to a thin, transparent layer that overlies the coarsely stratified basal layer in the region of abyssal hills (Fig. 9, mile 2800). As the line approaches the outer flank of the Hawaiian Arch, the transparent layer disappears, and sediments form a single, flat-lying unit with strong internal reflectors. Pinnacles of acoustic basement occasionally penetrate the sedimentary layer. This simple acoustic layering continues to the end of the track line.

ACKNOWLEDGMENTS

We would like to thank our shipboard colleagues for interesting discussions of the North Pacific regional geology. Gary H. Greene and Gordon Hess reviewed the manuscript, and their comments and suggestions are appreciated.

REFERENCES

- Larson, R. L., Moberly, R., et al., 1975. Site 311: Hawaiian magnetic lineations. In Larson, R. L., Moberly, R., et al., *Init. Repts. DSDP*, 32: Washington (U.S. Govt. Printing Office), 295–309.

Larson, R. L., and Hilde, T. W., 1975. A revised time scale of magnetic reversals for the Early Cretaceous and Late Jurassic. *J. Geophys. Res.*, 80:2586-2594.

Rea, D. K., 1970. Changes in structure and trend of fracture zones north of the Hawaiian Ridge and relation to seafloor spreading. *J. Geophys. Res.*, 75:1421-1430.

Talwani, M., Dorman, J., Worzel, J. L., et al., 1966. Navigation at sea by satellite. *J. Geophys. Res.*, 71:5891-5902.

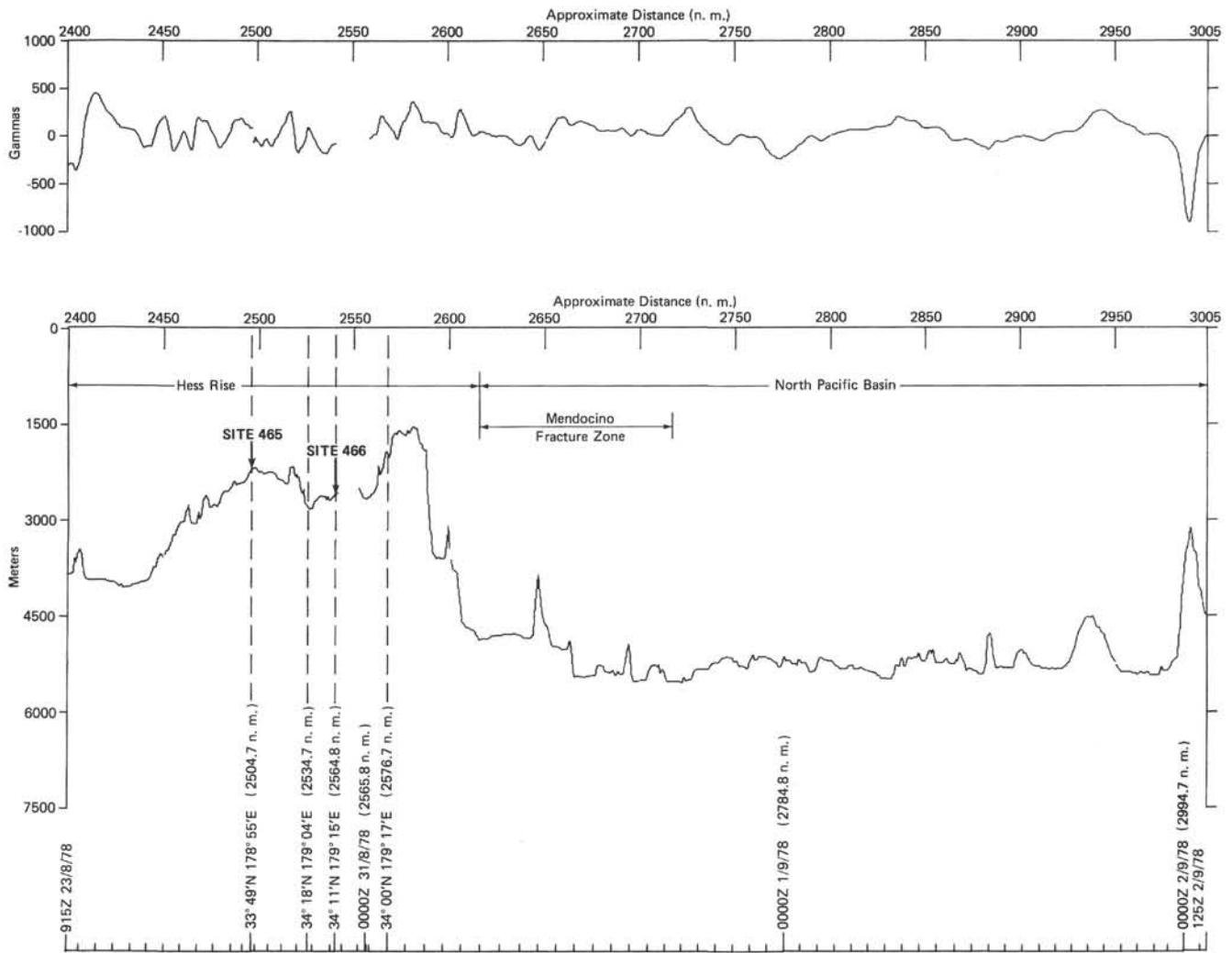


Figure 6. Bathymetric and magnetic-anomaly profiles along the track line of Leg 62. Annotations are the same as for Figure 2.

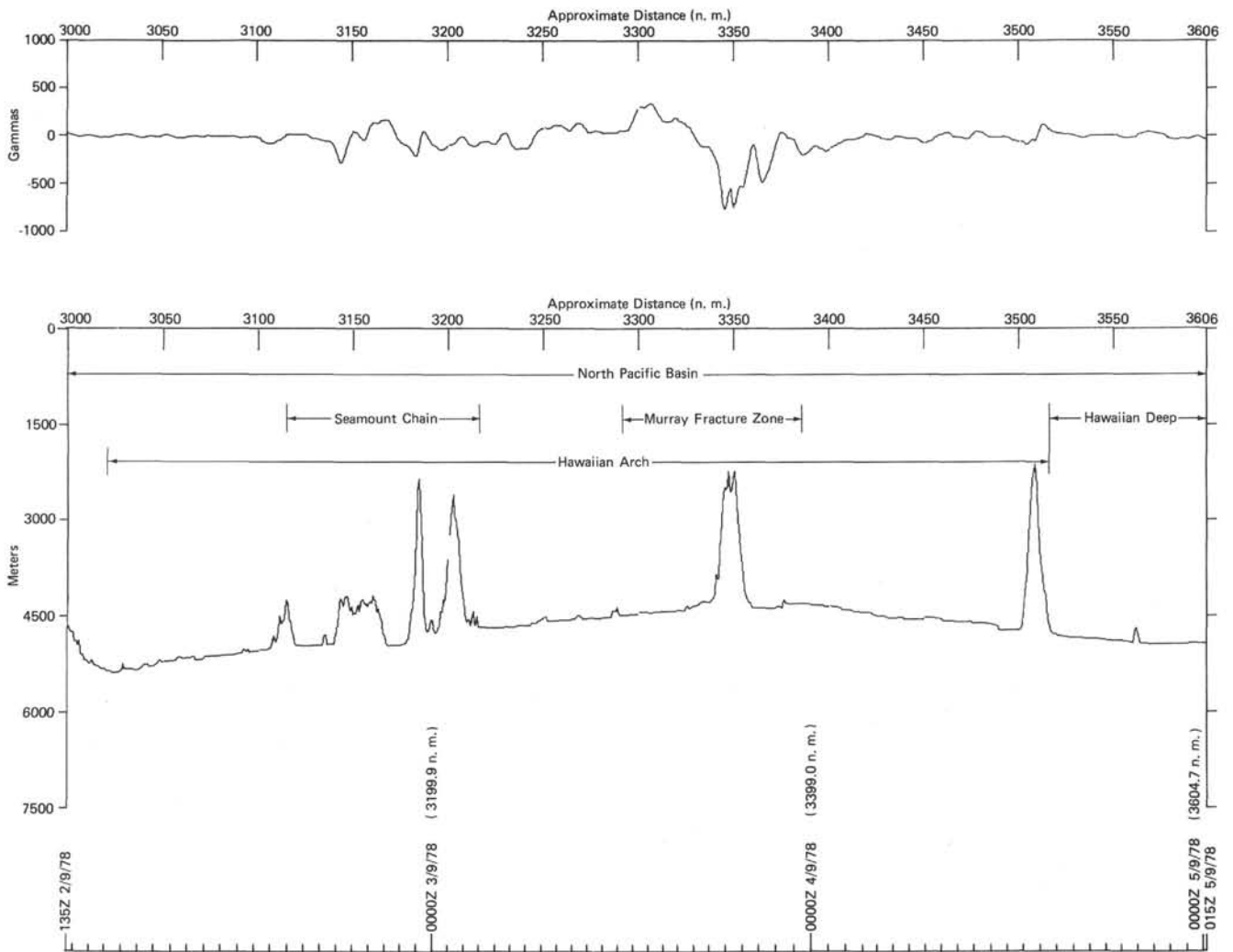


Figure 7. Bathymetric and magnetic-anomaly profiles along the track line of Leg 62. Annotations are the same as for Figure 2.

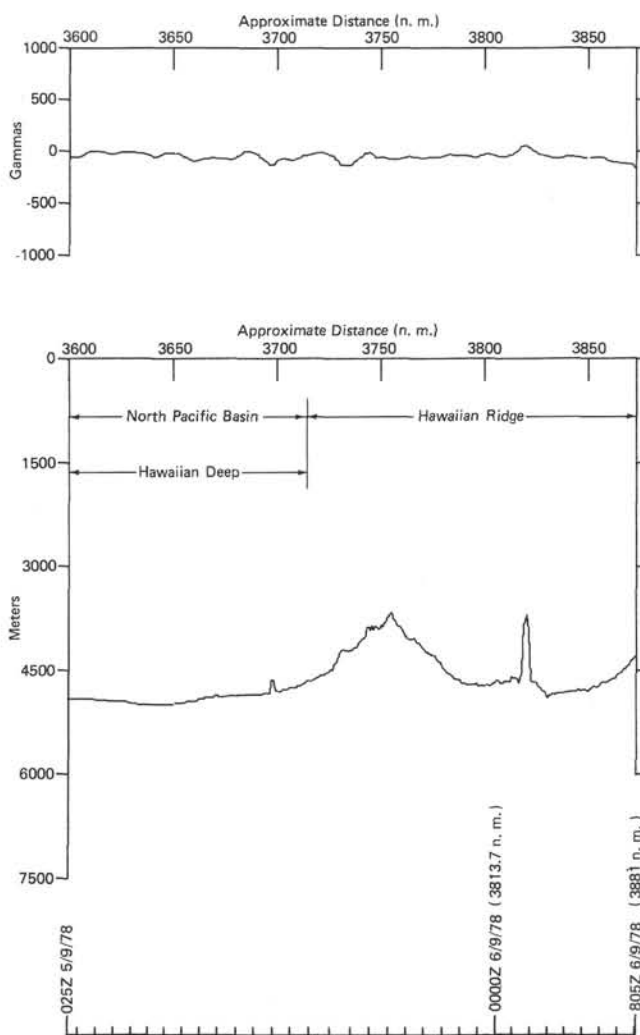


Figure 8. Bathymetric and magnetic-anomaly profiles along the track line of Leg 62. Annotations are the same as for Figure 2.