6. SITE 471: OFFSHORE MAGDALENA BAY¹

Shipboard Scientific Party²

HOLE 471

Date occupied: 9 November 1978

Date departed: 15 November 1978

Position: 23°28.93'N, 112°29.78'W

Water depth (sea level; corrected m, echo-sounding): 3101

Bottom felt (m, drill pipe): 3115.5

Penetration (m): 823

Number of cores: 88

Total length of cored section (m): 823

Total core recovered (m): 356.4

Core recovery (%): 43

Oldest sediment cored: Depth sub-bottom (m): 741.5 Nature: Metalliferous sediment Chronostratigraphy: Middle Miocene (14.5 to 15 m.y. old)

Basement:

Depth sub-bottom (m): 741.5 Nature: Altered diabase

Principal results: Hole 471 was drilled on the distal portion of a deepsea fan west of the foot of the continental slope off Baja California. Five sedimentary units were delineated. Unit 1, from mudline to 63.5 meters, is nannofossil silty clay with minor ash that was deposited during the Quaternary and Pliocene at a rate of 15 m/m.y. Unit 2 is diatomaceous clay and silty clay and clayey diatomaceous ooze to a depth of 155.2 meters. It was deposited principally in the late Miocene at a rate of 35 m/m.y.; the Miocene/Pliocene boundary is in the uppermost part of the unit. Units 1 and 2 have densities of 1.5 g/cm3 and velocities of 1.55 km/s. Unit 3 extends to a depth of 304 meters and is porcellanite and porcellaneous silty claystone with fragments of opal-CT (cristobalite) chert and thin beds of clayey dolomite. Core recovery averaged 5% in this unit; the Density and Neutron Logs suggest the presence of softer sediment interbeds in the unit that were not recovered. The top of Unit 3 is a diagenetic break marked by a sharp increase in density to 1.6 to 2 g/cm³ for porcellanite and up to 2.8 g/cm3 for dolomite; velocity increases sharply to 1.8 to 2.8 km/s (porcellanite) and 4 to 6 km/s (dolomite). Fossils are upper Miocene, but most samples are barren.

Unit 4 comprises the main part of the deep-sea fan and extends from 304 to 735.7 meters depth. It is bioturbated silty claystone with thin interbeds of calcareous sandstone and minor clavey carbonate and vitric ash. Faunal control is poor but indicates that deposition took place during the middle Miocene, with a sharp increase in sedimentation rate from 50 m/m.y. to 200 m/m.y. at about 360 meters depth. The seismic record shows an angular unconformity at this boundary; and the biostratigraphic record is consistent with a hiatus at the Unit 3/Unit 4 boundary. Sediment densities average about 2 g/cm^3 with somewhat higher values (2.4-3 g/cm^3) for carbonates and carbonate-cemented sandstone layers. Velocities are 2 km/s, increasing downsection to 2.3 km/s; carbonate and sandstone lavers are as high as 4.9 km/s.

Unit 5 consists of hemipelagic claystone from 735.7 to 741.3 meters depth and altered sediment to the top of the diabase at 741.5 meters depth. The unit is intensely burrowed and contains microfaults and calcite veins. The altered sediment includes chalcopyrite- and sphalerite-bearing claystone, black quartzose chert, and red brown metalliferous sediment. Velocities and densities are the same as those in Unit 4 but without the high carbonate or sandstone values. Intercalations of metalliferous sediment also occur within the diabase, which is altered and consists of at least two or three sills. Fragmental texture may be the result of emplacement into soft sediment, although slickensides indicate some shearing after consolidation. Density varies from 2.3 to 2.8 g/cm³, and velocity from 3.1 to 5.4 km/s; variability is in part caused by different degrees of alteration. The age of the oldest sediments above basement is 14.5 to 15.0 m.y., considerably older than the age extrapolated from the nearest striped magnetic anomalies.

A full suite of downhole logs was run from about the top of Unit 3 (top porcellanite) to total depth (T.D.). The Density and Sonic Logs clearly show the soft sediment in Unit 3 not recovered by coring; the porcellanite and dolomite beds are high values on both logs. Metalliferous sediment interbeds in diabase are also clearly indicated. The Density Log may demonstrate a correlation with degree of alteration of diabase. The conductivity curve on the Neutron Log best indicates the resolution of thin sandy beds in Unit 3. The Neutron Log shows considerable character in the diabase and may indicate fracture porosity or degree of alteration. Two Temperature Logs and two heat-probe measurements indicate high geothermal gradient and high heat flow; temperature is 12.5°C at a depth of 95 meters and 24.0°C at 142.5 meters. Assuming a conductivity of 2.5, heat flow at the site is 1.8 heat flow units (HFU) based on downhole logs and 3.9 HFU based on heat probe.

BACKGROUND AND OBJECTIVES

The Franciscan-like terrain of the California Continental Borderland reappears on Cedros and San Benito islands, westernmost Vizcaino Peninsula, and Magdalena and Santa Margarita islands west of Magdalena Bay (Fig. 1). The rocks of the Magdalena Bay islands include sheared gabbro, serpentinite, and variegated thin-bedded chert, an ophiolite assemblage that is highly sheared and locally a melange. Blueschist is rare at Magdalena Bay, although it is fairly abundant in the islands off Vizcaino Peninsula. The terrain is characterized by a strong free-air gravity high, which extends

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south-southeast from Magdalena Bay (Figs. 1 and 2), and high-relief magnetic anomalies, which extend seaward somewhat west of the gravity high (Figs. 1-3).

East of Magdalena Bay, Paleogene shallow-water sandstone (Tepetate Formation) is overlain by Oligocene and Miocene marine strata (San Ysidro and Monterey Formations), which are themselves overlain by the late Tertiary Comondu Volcanics, which form the mountainous backbone of Baja California del Sur, the Sierra Giganta. The Mesozoic batholithic terrain that dominates the high ground of Baja California del Norte may be buried beneath the Comondu. At this latitude, it appears at the surface only in the islands of the Gulf of California. Southeast of La Paz, granitic rocks also compose the Cape massif, a mountainous uplift separated from the Sierra Giganta by a north-south lowland that may be controlled by a fault.

Farther east is the Gulf of California, which underwent most of its rifting from mainland Mexico in the last 4 m.y., although a proto-gulf may have existed



Figure 1. Bathymetric map of the sea floor west of central Baja California (from Chase et al., 1974) indicating locations of Sites 470 and 471 and seismic profile AA' shown in Figure 4.



Figure 2. Free-air gravity anomaly map of the western continental margin of southern Baja California (from Huehn, 1977). (Contour interval is 10 mgal.)

earlier. The timing of initial separation of Baja California from the mainland is important to determine, because it would indicate when at this site a large source terrain, including the main Sierra Madre Occidental of Mexico, changed to a more restricted Baja California source terrain.

West of the ophiolitic zone of Magdalena Bay, the continental shelf is underlain by broadly folded Neogene sedimentary rocks (Fig. 4) that may overlie a ductile "Franciscan" or accretionary-wedge basement. The shelf is marked by linear highs and lows on the free-air gravity map and local magnetic highs suggestive of the high-amplitude anomalies of the islands off Magdalena Bay (Figs. 2 and 3). The shelf edge is marked by a linear gravity high that is flanked by a linear gravity low at the foot of the slope (Fig. 2).

A gravity low is associated with the Cedros deep, a feature that is topographically prominent from about 29°N south to about 24°30'N (Figs. 1 and 5). At these

latitudes, the deep appears as a graben downfaulted against the continental slope on the east and against the abyssal sea floor on the west. The Cedros deep gravity low persists southeast to the latitude of the site and beyond to 23 °N (Fig. 2). However, the topographic low is not present; instead, there is broad topographic bulge underlain by abyssal sea-floor sediments that dip gently to the east. The younger flat-lying sediments of Cedros deep and its southward continuation appear to overlie the pelagic sediments of the abyssal sea floor, but this relationship is not clearly established by reflection profiles. The Cedros deep does not appear to have a magnetic signature.

Site 471 is located in an area that is quiet in terms of gravity and magnetics (Figs. 1 and 3). Striped magnetic anomalies to the southwest were correlated to the geomagnetic time scale by Chase et al., (1970); their magnetic anomaly ages are shown in Figure 3. As Chase et al. (1970) pointed out, anomalies 12 m.y. of age and



Figure 3. Total magnetic-field anomaly map of the western continental margin of southern Baja California from Huehn (1977); contour interval is 100 gammas. (Ages of magnetic stripes are after Chase et al. [1970].)



Figure 4. Line drawing of seismic-reflection profile AA' through Site 471. (Vertical scale in kilometers assumes 1.5 km/s for the velocity of sand in water [from Huehn, 1977]. Deepest short-dashed line below site is basement reflector.)

older trend north-northwest, roughly parallel to the continental slope, but younger anomalies produce a fan pattern—such that those 8 m.y. old and younger are parallel to the present spreading center at the mouth of the gulf. The fan pattern of the magnetics makes estimating the magnetic-anomaly age of the crust at Site 471 somewhat ambiguous; the best guess is 11 m.y.

The thicker sediment section closer to the continental slope must indicate an increase in terrigenous input compare with Site 470. The eastward dip of the basement surface toward the foot of the slope (Fig. 4) suggests that the Sclater age-depth curve may be unusable in this case, because basement slopes downward in the direction of younger crust, the opposite effect of the Sclater age-depth relation. (Interestingly, the *Glomar Challenger* track from Site 470 to the Cedros deep shows the basement rising eastward in the direction of younger crust, as predicted by the Sclater curve, reversing only fairly close to the trench.) The deepening of the basement toward the continental slope may be the result of the vertical load of the continent, as suggested by the gravity low at the foot of the slope.

Site 471 is located, as are the more northerly sites, near a continental margin that was once a subducting margin, indicated by the "Franciscan" and ophiolite terrain; this would imply that the east slope of the Cedros deep was once a trench slope and that the deep itself is a fossil trench. The orientation of magnetic anomalies swung from north-south to northwest by 12 m.y. ago, an effect that Menard (1978) suggests may have been caused by the Farallon plate subducting only where there was sufficient thermal contrast between the cold, sinking plate and the surrounding asthenosphere for the plate to sink gravitationally. Where the rise crest intersected the trench obliquely, the Farallon plate would have been of near zero age at the time of subduction, and it would not have sunk because of a lack of thermal and density contrast; instead, the plate and its trailing rise crest would have pivoted counterclockwise parallel to the continental slope. The subsequent clockwise swing between 12- and 10-m.y.-old magnetic stripes may have been the result of a triple junction involving a

small plate to the north, as Chase et al. (1970) suggested.

Whether the foot of the continental slope was a transform boundary 60 to 10 m.y. ago is not clear, but it is now a passive margin with low seismicity, just as it is farther north off southern California.

The oceanic front migrated approximately 10° of latitude in the northeastern Pacific during the Neogene in response to major climatic oscillations (Ingle, 1973). The location of Site 471, just south of the present-day mixing zone between distal California Current and equatorial waters and about 5° south of Site 470 (Figs. 2, 5), is well-suited for the study of the extent of southward penetration of higher-latitude assemblages during cold pulses in the paleoclimatic history of the area. The 800 meters of sediments overlying the basement were expected to contain a middle Miocene to Holocene planktonic record of mostly temperate elements, with influx of cooler, higher-latitude elements during times of climatic deteriorations and/or intensification of oceanic circulation.

OPERATIONS

The track from Site 470 to 471 was designed to determine the relations of both sites to the continental slope off Baja California. Accordingly, we took a slightly zigzag course to cross the Cedros deep twice, then steamed northeast across Site 471 to the continental slope, then returned on our track and dropped the beacon as we headed southwest down the broad apron of a deep-sea fan at the foot of the slope (Figs. 6 and 7).

Continuous coring was routine, with moderate to high recovery to 161.5 meters (Table 1). The heat-flow probe was run twice; the criterion for running the probe was that sediment be firm enough to take weight, but not so hard that the probe would not penetrate it. With a planned 50 meters between probe runs, it was possible to run the probe only twice, once at 95 meters and again at 155 meters. The sand line just above the heat probe was found to be knotted after the second run, requiring that about 150 meters of sand line be cut off before the next core.



Figure 5. Location of Leg 63 sites.

Beginning with Core 18 at 161.5 meters, recovery dropped to about 5%, and the coring rate decreased from less than 10 min per core to 20 to 50 min per core. This zone of low recovery corresponds to Lithologic Unit 3, characterized by porcellanite alternating with softer sediments that were not recovered in cores but were identified on downhole logs. Recovery improved beginning with Core 34 at a depth of 313.5 meters, the approximate top of Lithologic Unit 4-the distal turbidite fan sequence. In the past, turbidites have been difficult for DSDP to drill and to recover in cores. Our favorable experience here may have been caused by the low porosity resulting from carbonate cement and to the lack of interbedded cherts or porcellanites. Core recovery was moderately high in Unit 4, with an occasional empty core barrel (Cores 55 and 67) and some recoveries limited to core catchers (in Cores 40, 49, 54, 56, 61, 62, and 68). There is no obvious correlation between core recovery and lithology in Unit 4. Coring rates in Units 4 and 3 were about the same. Other problems included bit plugging: a piece of core would lodge in the throat of the bit, which had to be dislodged with a center bit. Some torquing was noted in Unit 3, probably caused by fractured porcellanite falling in the hole. Frequent mud flushes in this interval cleaned up the hole, and few problems occurred at greater depths in sediment. Small amounts of gas bleeding from the cores were monitored on the Carle and HP gas chromatographs.

Diabase was encountered in Core 79, and coring continued through Core 88. Recovery was fairly good, and cores were close to gauge even at the bottom of the hole. The diabase occurs as altered sills, a lithology that has much higher recovery than pillow basalts, according to experience at previous DSDP sites. The center bit had to be run after Cores 79, 80, and 85, following indications of a plugged bit; otherwise operations were fairly routine. Coring time ranged from 26 minutes for Core 86 to 197 minutes for Core 84; there was no significant decrease in coring rate with depth. Our most trying moments were a stuck core barrel at Core 88, which finally released after once shearing the pin on the fishing neck. "Clay" at the bottom of the hole was in part fill and in part altered, fractured diabase.

The hole was flushed with 30 barrels of gel mud and 20 barrels of guartec in preparation for logging. The hydraulic bit release (HBR) go-devil was pumped down and the bit released at 1600 psi. The hole was filled with 300 barrels of gel mud, and the drill pipe was pulled to 158 meters below the mudline, near the top of Unit 3 (containing porcellanite). We set the pipe this deep to avoid losing the hole in sediments as unconsolidated as those in Units 1 and 2. (We attribute the failure of the logging tool to penetrate the sediment at Site 470 to the fact that sediments below the drill pipe were so unconsolidated that the hole was lost simply by motion of the drill pipe related to heaving of the ship.)

The Sonic, Caliper, and Gamma-Ray Logs were run, and the hole was found to be clean to bottom. The



Figure 6. Challenger track approaching and departing Site 471. (Small concentric circles on track lines are satellite fixes. Bathymetry [in meters] was contoured aboard ship on the basis of data from Baja 75 and Baja 76 cruises, Oregon State University.)



Figure 7. Challenger seismic line approaching Site 471. (See Fig. 6 for location.)

Table 1. Coring summary, Site 471.

| Core No. | Date (Nov. 1978) | Time | Depth from Drill Floor (m) | Depth below Sea Floor (m) | Length Cored (m) | Length Recovered (m) | Core Recovered (%) |
|-------------|------------------------|------|----------------------------------|---------------------------------|------------------------|----------------------------|--------------------------|
| 1 | 9 | 0108 | 3115.5-3125.0 | 0.0-9.5 | 9.5 | 9.87 | 100+ |
| 2 | 9 | 0204 | 3125.0-3134.5 | 9.5-19.0 | 9.5 | Trace | 0 |
| 3 | 9 | 0259 | 3134.5-3144.0 | 19.0-28.5 | 9.5 | 0.00 | 63 100 ± |
| 5 | 9 | 0450 | 3153.5-3163.0 | 38.0-47.5 | 9.5 | 10.83 | 100 + |
| 6 | 9 | 0548 | 3163.0-3172.5 | 47.5-57.0 | 9.5 | 9.75 | 100 + |
| 7 | 9 | 0645 | 3172.5-3182.0 | 57.0-66.5 | 9.5 | 10.12 | 100 + |
| 8 | 9 | 0736 | 3182.0-3191.5 | 66.5-76.0 | 9.5 | 7.80 | 82 |
| 10 | 9 | 0840 | 3191.5-3201.0 | 76.0-85.5 | 9.5 | 3.48 | 100+ |
| 11 | 9 | 1215 | 3210.5-3220.0 | 95.0-104.5 | 9.5 | 9.82 | 100 + |
| 12 | 9 | 1320 | 3220.0-3229.5 | 104.5-114.0 | 9.5 | 9.56 | 100 + |
| 13 | 9 | 1412 | 3229.5-3239.0 | 114.0-123.5 | 9.5 | 10.26 | 100 + |
| 14 | 9 | 1507 | 3239.0-3248.5 | 123.5-133.0 | 9.5 | 7.95 | 84 |
| 15 | 9 | 1600 | 3248.5-3258.0 | 133.0-142.5 | 9.5 | 10.30 | 100 + |
| 17 | 9 | 2101 | 3267 5-3277 0 | 152.0-161.5 | 9.5 | 3.25 | 34 |
| 18 | 9 | 2217 | 3277.0-3286.5 | 161.5-171.0 | 9.5 | 0.20 | 2 |
| 19 | 10 | 0039 | 3286.5-3296.0 | 171.0-180.5 | 9.5 | 0.26 | 3 |
| 20 | 10 | 0227 | 3296.0-3305.5 | 180.5-190.0 | 9.5 | 0.50 | 5 |
| 21 | 10 | 0550 | 3305.5-3315.0 | 190.0-199.5 | 9.5 | 0.78 | 8 |
| 22 | 10 | 0720 | 3315.0-3324.5 | 199.5-209.0 | 9.5 | 0.39 | 4 |
| 24 | 10 | 1005 | 3324.3-3334.0 | 218 5-228 0 | 9.5 | 0.00 | 0 |
| 25 | 10 | 1145 | 3343.5-3353.0 | 228.0-237.5 | 9.5 | 0.03 | 0.3 |
| 26 | 10 | 1326 | 3353.0-3362.5 | 237.5-247.0 | 9.5 | 0.83 | 9 |
| 27 | 10 | 1451 | 3362.5-3372.0 | 247.0-256.5 | 9.5 | 0.20 | 2 |
| 28 | 10 | 1630 | 3372.0-3381.5 | 256.5-266.0 | 9.5 | 0.50 | 5 |
| 29 | 10 | 1839 | 3381.5-3391.0 | 266.0-275.5 | 9.5 | 0.19 | 2 |
| 30 | 10 | 2150 | 3391.0-3400.3 | 275.5-285.0 | 9.5 | 1.03 | 14 |
| 32 | 10 | 2326 | 3410.0-3419.5 | 294.5-304.0 | 9.5 | 0.94 | 10 |
| 33 | 11 | 0110 | 3419.5-3429.0 | 304.0-313.5 | 9.5 | 0.91 | 10 |
| 34 | 11 | 0233 | 3429.0-3438.5 | 313.5-323.0 | 9.5 | 3.27 | 34 |
| 35 | 11 | 0402 | 3438.5-3448.0 | 323.0-332.0 | 9.5 | 3.98 | 42 |
| 36 | 11 | 0525 | 3448.0-3457.5 | 332.5-342.0 | 9.5 | 1.78 | 19 |
| 38 | 11 | 0825 | 3457.5-3407.0 | 342.0-351.5 | 9.5 | 2.98 | 31 |
| 39 | 11 | 0952 | 3476.5-3486.0 | 361.0-370.5 | 9.5 | 6.10 | 64 |
| 40 | 11 | 1120 | 3486.0-3495.5 | 370.5-380.0 | 9.5 | 0.04 | 0.4 |
| 41 | 11 | 1235 | 3495.5-3505.0 | 380.0-389.5 | 9.5 | 6.52 | 69 |
| 42 | 11 | 1345 | 3505.0-3514.5 | 389.5-399.0 | 9.5 | 1.25 | 13 |
| 43 | 11 | 1500 | 3514.5-3524.0 | 399.0-408.5 | 9.5 | 6.63 | 70 |
| 44 | 11 | 1018 | 3524.0-3533.5 | 408.5-418.0 | 9.5 | 3.82 | 40 |
| 45 | 11 | 1907 | 3543 0-3552 5 | 418.0-427.5 | 9.5 | 2 36 | 25 |
| 47 | 11 | 2026 | 3552.5-3562.0 | 437.0-446.5 | 9.5 | 2.53 | 27 |
| 48 | 11 | 2143 | 3562.0-3571.5 | 446.5-456.0 | 9.5 | 4.05 | 43 |
| 49 | 11 | 2303 | 3571.5-3581.0 | 456.0-465.5 | 9.5 | 0.32 | 3 |
| 50 | 12 | 0045 | 3581.0-3590.5 | 465.5-475.0 | 9.5 | 5.47 | 58 |
| 52 | 12 | 0212 | 3590.5-3600.0 | 475.0-484.5 | 9.5 | 4.46 | 47 |
| 53 | 12 | 0452 | 3609.5-3619.0 | 494.0-503.5 | 9.5 | 6.82 | 72 |
| 54 | 12 | 0615 | 3619.0-3628.5 | 503.5-513.0 | 9.5 | 0.14 | 1 |
| 55 | 12 | 0735 | 3628.5-3638.0 | 513.0-522.5 | 9.5 | 0.00 | 0 |
| 56 | 12 | 0910 | 3638.0-3647.5 | 522.5-532.0 | 9.5 | 0.12 | 1 |
| 57 | 12 | 1030 | 3647.5-3657.0 | 532.0-541.5 | 9.5 | 6.75 | 71 |
| 50 | 12 | 1205 | 3657.0-3000.3 | 541.5-551.0 | 9.5 | 0.40 | 38 |
| 60 | 12 | 1450 | 3676.0-3685.5 | 560.5-570.0 | 9.5 | 2.65 | 28 |
| 61 | 12 | 1612 | 3685.5-3695.0 | 570.0-579.5 | 9.5 | 0.09 | 1 |
| 62 | 12 | 1733 | 3695.0-3704.5 | 579.5-589.0 | 9.5 | 0.12 | 1 |
| 63 | 12 | 1856 | 3704.5-3714.0 | 589.0-598.5 | 9.5 | 5.95 | 63 |
| 64 | 12 | 2031 | 3714.0-3723.5 | 598.5-608.0 | 9.5 | 7.46 | 79 |
| 66 | 12 | 2159 | 3723.0-3742.5 | 608.0-617.5 | 9.5 | 4.17 | 44 |
| 67 | 13 | 0135 | 3742.5-3752.0 | 627.0-636.5 | 9.5 | 0.00 | ò |
| 68 | 13 | 0345 | 3752.0-3761.5 | 636.5-646.0 | 9.5 | 0.10 | 1 |
| 69 | 13 | 0715 | 3761.5-3771.0 | 646.0-655.5 | 9.5 | 5.67 | 60 |
| 70 | 13 | 0930 | 3771.0-3780.5 | 655.5-665.0 | 9.5 | 6.48 | 68 |
| 71 | 13 | 1123 | 3780.5-3790.0 | 665.0-674.5 | 9.5 | 6.10 | 64 |
| 72 | 13 | 1632 | 3790.0-3799.5 | 684 0_603 5 | 9.5 | 7.55 | 80 |
| 74 | 13 | 1812 | 3809 0-3818 5 | 693 5-703 0 | 9.5 | 7.00 | 74 |
| 75 | 13 | 1948 | 3818.5-3828.0 | 703.0-712.5 | 9.5 | 6.96 | 73 |
| 76 | 13 | 2145 | 3828.0-3837.5 | 712.5-722.0 | 9.5 | 7.35 | 77 |
| 77 | 13 | 2350 | 3837.5-3847.0 | 722.0-731.5 | 9.5 | 8.63 | 91 |
| 78 | 14 | 0221 | 3847.0-3856.5 | 731.5-741.0 | 9.5 | 7.71 | 81 |
| 80 | 14 | 1105 | 3830.3-3866.0 | 741.0-750.5 | 9.5 | 2.23 | 23 |
| 81 | 14 | 1440 | 3875.5-3884.5 | 760.0-769.0 | 9.5 | 4 35 | 48 |
| 82 | 14 | 1705 | 3884.5-3893.5 | 769.0-778.0 | 9.0 | 2.12 | 24 |
| 83 | 14 | 1934 | 3893.5-3902.5 | 778.0-787.0 | 9.0 | 3.65 | 41 |
| 84 | 15 | 0000 | 3902.5-3911.5 | 787.0-796.0 | 9.0 | 2.22 | 25 |
| 85 | 15 | 0334 | 3911.5-3920.5 | 796.0-805.0 | 9.0 | 2.41 | 27 |
| 80 | 15 | 1045 | 3920.5-3922.5 | 805.0-807.0 | 2.0 | 0.50 | 25 |
| 88 | 15 | 1515 | 3929.5-3928.5 | 814.0-823.0 | 9.0 | 4.99 | 55 |
| Catal | | 1010 | | 31110 04310 | 000.0 | 266.10 | 13 |
| otau | | | | | 823.0 | 330.40 | 43 |

variable Density-Sonic (Wave-Train) Log was then run, followed by the Temperature-Density-Gamma-Ray Log; temperature was logged going down and the Density-Gamma-Ray Log was taken coming up. This was followed by the Guard-Neutron-Gamma-Ray Log, which was followed by the Induction-Gamma-Ray Log. A final Temperature Log was taken after pulling a stand of pipe while the log was in the hole, thereby allowing another 28 meters of open hole to be logged.

The hole was then cemented because of the gas shows monitored in the cores, the pipe was pulled, and we left for the next site at 1327 hours,³ 18 November, 1978.

LITHOLOGY

Sediments and Sedimentary Rocks

Site 471 is characterized by a thick section of interbedded silty claystone and sandstone overlain by a thinner section of sediments and sedimentary rocks, including porcellanite, diatomaceous clay, and nannofossil silty clay. We defined five lithologic units above altered diabase at this site (Fig. 8; Table 2).

Unit 1: Nannofossil Silty Clay (0-63.5 m)

Unit 1 is mainly composed of grayish olive green nannofossil silty clay. The abundance of nannofossils ranges from 15% to 40%. Grains of angular silt-size quartz and feldspar are less common, 3% to 12% and 1% to 6%, respectively. Siliceous microfossils are rare. In addition to nannofossil silty clay, this unit also contains layers of olive gray to dark greenish gray silty clay. Angular, silt-size grains of quartz and feldspar average 20% and 8%, respectively. Small patches of pinkish gray vitric ash and dusky yellow green calcareous ooze are scattered throughout this unit. Dark reduction spots and streaks of finely disseminated pyrite are also present. The boundary with the underlying diatomaceous sediment of Unit 2 is gradational, marked by a distinct increase in the abundance of diatoms in Core 6.

Unit 2: Variegated Diatomaceous Clay and Ooze (63.5-155.2 m)

Unit 2 consists of diatomaceous clay, diatomaceous silty clay, and clayey diatomaceous ooze. These sediments vary from dusky yellow green and olive gray to dark yellowish brown and grayish olive green. The darker colors correspond to greater proportions of diatoms in the sediment (e.g., Cores 9–10 and 16–17). Siltsize grains of quartz and feldspar decrease in abundance downhole in this unit, and diatomaceous silty clay grades into the underlying diatomaceous clay and clayey diatomaceous ooze in Core 8. Correspondingly, the abundance of diatoms increases downhole from 15% to 70%. Grayish purple reduction spots and streaks of finely disseminated pyrite as well as dusky yellowish

 $^{^3}$ Times specified in the text are local times in hours, and those in the seismic-section figures are Zulu (Z) times.

| | LENGTH RE- COVERED | CORE | GRAPHIC LITHOLOGY | LITHO- LOGIC UNIT | COCCOLITHS | RADIOLARIANS | DIATOMS | PLANKTONIC FORAMINIFERS | CHRONOSTRATIGRAPHY |
|-------------|--------------------------|----------|----------------------|-------------------------|--|------------------------------|-----------------------------|---------------------------------------|--------------------|
| 0- | | 1 2 3 | ×,° ⊥ | 1 | Emiliania huxleyi or Gephyrocapsa oceanica Gephyrocapsa caribbeanica | Quaternary | | N23 | Quaternary |
| | | 4 | ·· · · · · | | Discoaster brouweri D. surculus D. tamalis | -?Lamprocyrtis heteroporos?- | | N21 | Upper |
| 1 | | 6 | | - | | Cticheses a constring | Pliocene | | Pliocene |
| | | 8 | $\sim \sim$ | - | | (upper) | | | Lower |
| | | 9 | ~~~ | | | 2 | | | |
| 100- | | 11 | ~~~ | 2 | | Stichocorys peregrina | Nitzschia reinholdii | | |
| | | 13 | ~~ | | | (lower) | or Thalassiosira antigua | | |
| | | 14 15 | ~~~ | | | Ommatartus penultimus | | | |
| - | | 16 | | | | | | | Upper |
| | | 18 | | | | Ommatastus | | | Miocene |
| | | 20 | | | | antepenultimus | | | |
| 200- | | 21 | | | | | | | |
| | | 23 | | | | תחחחחח | | | |
| | | 24 | | 3 | | | | | |
| - 1 | | 26 | L L . | | | | | | |
| - | | 28 | Δ | | | | | | |
| | | 30 | Δ | | | | | | Upper |
| 300 - | | 31 | Δ | | | | | | Miocene |
| | | 33 | | | | | | | Middle Miocene |
| | | 35 | | | | | | | 1111111490 |
| 1 | | 36 | | | | | | | |
| Ê | | 38 39 | | | | | | | |
| pth | | 40 | | | | | | | |
| ≗ ∈ 400- | | 41 42 | | | | | | | |
| otto | | 43 | · · *** | | | | | | |
| q-qns | | 45 | | | | | | | |
| - | | 40 | | | | | | | |
| | | 48 | | | | | | | |
| | | 50 51 | | | | | | | |
| 500- | | 52 | | | | | | | |
| 500 | | 53 | | 4 | | | | | |
| | | 55 56 | | | a 58 - 54 | | | | |
| | | 57 | | | Coccolithus miopelagicus | | | | |
| 1 | | 59 | | | | | | | |
| - 3 | | 61 | | | | | | | |
| 600- | | 62 63 | | | | | | | Middle |
| 000 | | 64 | | | | | | | Miocene |
| | | 66 | | | | | | | |
| 1 | | 68 | | | | | | | |
| | | 69 70 | | | | | | | |
| - 0 | | 71 | | | | | | | |
| - 1 | | 73 | | | | | | | |
| 700- | | 74 | | | | | | | |
| | | 76 | | | | | | | |
| | | 78 | 0047470077F | 5 | Sphenolithus heteromorphus | | | | |
| | | 80 | 1 | | × × 7 × 7 × × × | > 4 2 4 7 7 7 7 | 4774F74 | 2 2 4 7 7 2 4 7 7 | |
| | | 81 | Diabase | | 2 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 | 5-1-7-2-5-4- | 7 6 7 7 1 7 7 3 | 75474 | |
| | _ | 83 | | | F7 F7 J S7 V A | × × × × × × × × | | ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ | |
| 800- | | 85 87 | 86 | | ×1 × 1 × 1 × 1 × 1 × 1 × 1 | 1 2 1 1 4 4 4 4 4 | 1727217 | 77777 | |
| | | 88 | | | 1 < 1 > V < < 1 < V | KN7K7 67 77 | Y | L L V 7 K J | |

Figure 8. Lithologic and biostratigraphic summary, Site 471.

Table 2. Summary of lithologic units, Site 471.

| Unit or Sub-unit | Core Number | Depth below Sea Floor (m) | Chrono- stratigraphy | Lithology |
|---------------------|---------------------------------|---------------------------------|-------------------------|---|
| 1 | 1-6, Section 3 | 0.0-63.5 | Quaternary- Pliocene | Nannofossil silty clay, silty clay, and scattered patches of vitric ash. |
| 2 | 6, Section 3- 17,CC | 63.5-155.2 | upper Miocene | Diatomaceous clay, diatomaceous silty clay, and clayey diatomaceous ooze. |
| 3 | 17,CC-32 | 155.2-304.0 | upper Miocene | Porcellanite and porcellaneous silty claystone with a few fragments of opal-CT chert and several clayev limestone layers. |
| 4 | 33-78 Section 3 | 304.0-735.7 | middle Miocene | Bioturbated silty claystone with thin interbeds of calcareous sand- stone. Sandstone becomes less abundant near base of unit. Several layers of clayey carbonate and vitric ash. |
| 5a | 78, Section 3- 79, Section 1 | 735.7-741.3 | middle Miocene | Chalcopyrite- and sphalerite- bearing rock, quartzose chert, |
| 5b | 5b 79, Section 1 | | | and metalliferous sediment above altered diabase. |

green patches of vitric ash also occur in this unit. Although the gradational lower boundary of Unit 2 was not recovered intact, the few fragments of porcellanite that occur with pieces of firm clayey diatomaceous ooze in Sample 17,CC mark the base of the unit.

Unit 3: Porcellanite and Silty Claystone (155.2-304.0 m)

Unit 3 is olive gray to olive black opal-CT porcellanite and porcellaneous silty claystone. A few fragments of olive black opal-CT chert and several thin interbedded layers of light olive gray silicified dolomite also occur. Recovery was extremely poor in this interval, averaging only 5%. Most cores consist of fragments of hard porcellanite that have been broken and brecciated by drilling. Downhole logs (especially the Density and Neutron Logs) and several pieces of firm silty clay within the sequence (as in Cores 26, 27, and 29-31), however, suggest that these hard siliceous rocks are interbedded with softer sediment. Slabbed surfaces of the porcellanite display abundant burrows.

In thin section, the porcellanite consists dominantly of clay cemented by silica. X-ray diffraction data show the siliceous cement to be opal-CT. Minor chalcedonyfilled molds of diatoms, radiolarians, and sponge spicules also occur, as well as scattered rhombs of carbonate. Silt-size grains of quartz and feldspar are present in rare laminae within the porcellanite fragments. A few pieces of chert occur near the top of the unit; they are conchoidally fractured, have a vitreous luster, and are composed mostly of opal-CT with minor clay minerals.

The siliceous dolomite interbeds in Unit 3 are intensely bioturbated. In thin section, the dolomite consists of micrite or sugary-textured dolomite and clay minerals that have been cemented with opal-CT.

Porcellanite grades downhole to porcellaneous silty claystone, then to silty claystone at the base of Unit 3. This gradual change makes it difficult to place a lower boundary on these siliceous rocks. The boundary between Units 3 and 4 is placed between Cores 32 and 33, where the claystone becomes noticeably less siliceous. Density and Sonic Logs also display a distinct break at about 304 meters, near the base of Core 32.

Unit 4: Interbedded Silty Claystone and Sandstone (304.0-735.7 m)

Unit 4 is a thick, well-bedded sequence of olive gray silty claystone containing numerous thin interbeds of bluish gray calcite-cemented sandstone and light gray, structureless silty claystone. A few beds of light gray altered vitric tuff and several thin layers of clayey dolomite and a limestone (Core 37) occur near the top of the unit. The dark olive gray claystone, in beds 4 to 15 cm thick, is thoroughly bioturbated. Large lenticular burrows are filled with light olive gray silty claystone. Other burrow forms include Zoophycos and possible Condrites. The claystone is uniform in composition and is composed of angular silt-size grains of quartz (3-15%) and feldspar (2-7%), clay minerals (60-80%), nannofossils (2-10%), and scattered white sponge spicules (<1%). The siliceous sponge spicules are cigarshaped and 2 to 5 mm long, have dark claystone-filled interiors, and are commonly coated with a thin film of dark organic material.

Sandstone interbeds first appear in Core 35 as thin laminae within the bioturbated silty claystone sequence and become thicker and more abundant in Core 47. These persist through Core 77 but are not present in Core 78 immediately above the claystone of Unit 5. Most sandstone layers range from 1 to 4 cm thick, although some attain a maximum thickness of 10 cm. The spacing and frequency of these layers are also variable. Generally, they occur at least 5 to 10 cm apart. A core may contain as many as 30 to 40 thin sandstone beds (e.g., Core 64).

The sandstone is mostly fine-grained and moderately well sorted. It consists of angular grains of polycrystalline and unstrained monocrystalline quartz and chert (15-20%), albite twinned plagioclase (2-15%), rock fragments (10-20%), and coated pellets (2-10%). Biotite (2%), hornblende (2%), pyroxene (trace), and foraminifers (1%) are accessory components. Rock fragments are chiefly silicic volcanics such as dacite and rhyolite. Some coarse sand-size polycrystalline quartz grains may be metaquartzite. Fine sand-size pelletal grains contain cores of angular quartz or feldspar; the coating is opaque or, less commonly, light brown and may be a mixture of phosphate and manganese oxideshydroxides. Abundant calcite, in large, optically continuous patches, is a cement and partially replaces framework grains in these calcareous sandstones.

Many sandstone layers display Bouma d,e sequences (i.e., parallel laminations gradationally overlain by thicker and finer-grained burrowed, sandy, silty claystone). In some, an intervening layer of micro-crosslaminations (division c) is also present, forming Bouma b,c,d or c,d,e sequences. Commonly, the upper few centimeters of each sandstone bed are intensely bioturbated, creating the appearance of distorted ripple crosslaminations. Compact lenticular concentrations of dark gray, silt- and sand-size faecal pellets are abundant in the burrows in the sandstone and also in the overlying thicker, sandy, silty claystone. It is possible that some of the sandstone layers may have been ripple cross-laminated and only partially disrupted by burrowing. Some sandstone layers have load structures, and most have sharp basal contacts. The underlying claystone, which has only a few burrows, is generally lighter gray and thinner (<2 cm) than the more extensively burrowed silty claystone that immediately overlies each sandstone layer.

Thin layers of altered vitric tuff occur near the top of Unit 4. In Core 44, a 7-cm-thick, biotite-bearing, light gray tuff is present. It has a sharp basal and gradational upper contact and is graded.

Some sandstone layers are present in Cores 78 and 79 near the base of the unit, although there is a regular alternation of light and dark gray silty claystone beds, especially in Core 78. The boundary between Unit 4 and the underlying claystone of Unit 5 is placed at the last occurrence of calcite-cemented sandstone in Core 78.

Unit 5: Claystone and Sulfide-Bearing Sediment (735.7-741.5 m)

Unit 5 is a thin sequence of claystone that is hydrothermally altered at its base. We divide it into two parts: Sub-unit 5a, claystone, and Sub-unit 5b, sulfide-bearing sediment, quartzose black chert, and metalliferous sediment overlying altered diabase.

Sub-unit 5a is greenish gray claystone. It is intensely mottled by burrows and contains abundant microfaults and some calcite-filled veins. Reduction streaks and spots of finely disseminated, dark gray pyrite are common. Except for a few scattered silty layers, this unit contains significantly fewer silt-size grains of quartz and feldspar than does the overlying silty claystone of Unit 4. The clay to silt ratio is about 9:1. Nannofossils are present in the claystone at the top of this sub-unit but decrease and finally disappear at its base. Several scattered grayish blue green streaks may be altered vitric ash.

Sub-unit 5b is a potpourri of sediments and sedimentary rocks including chalcopyrite- and sphalerite-bearing claystone, black quartzose chert, and metalliferous sediment. Chalcopyrite- and sphalerite-bearing rocks form the upper part of this unit. Rock fragments recovered in the brecciated base of this sequence consist of claystone, chalcopyrite-sphalerite, and black quartzose chert. The chert is conchoidally fractured and veined by calcite. A thin (2-cm) layer of dusky yellowish brown metalliferous sediment overlying altered diabase forms the base of Unit 5. Clay, X-ray amorphous iron-rich material, and minor carbonate rhombs are dominant components of this sediment. Similar sediment occurs within fractures and between altered diabase in Section 82-1 and Section 88-1. Where metalliferous sediment fills some fractures in the diabase, it has been thermally altered to red brown jasper.

Igneous Rocks

Diabase (and its fine-grained equivalent along chilled zones) and brecciated diabase are the two igneous rock types recovered at Site 471. Thin intercalations of claystone at depths of 750.5 meters and 769.5 meters provide convenient boundaries for dividing these rocks into three sequences (Fig. 9). The upper sequence (741.5–750 m) is mainly fine- to medium-grained, aphyric diabase having a chilled margin of aphyric basalt at the upper contact with the overlying sediment. Grain size increases away from this contact. Plagioclase is the only remaining primary constituent in these rocks, occurring as euhedral laths (< 2 mm). The texture is intersertal. Clays, calcite, and zeolite are the common alteration minerals; neither primary mafic minerals nor any pseudomorphs after them are found.

The middle diabase sequence (750.6–769.6 m) coarsens slightly away from the bounding claystone layers over an interval of about 1 meter. The upper part of this sequence (Core 80) is mostly altered brecciated diabase consisting of angular to subangular fragments of altered diabase and clayey chert generally less than 2 cm across. Slickensides mark the face of one diabase fragment. The lower part of this sequence is highly altered, mediumgrained diabase composed of subhedral plagioclase and clinopyroxene with intersertal to subophitic textures. Faint purple pink tints indicate some clinopyroxenes may be titaniferous. Clay, calcite, and an unidentified zeolite are common alteration minerals.

The lower diabase sequence (769.6-820 m) may be further subdivided into two portions. The upper part (Cores 82-83) contains a distinct chilled zone with grain size gradually decreasing in the first 1.5 meters below the claystone. This part is compositionally similar to the middle sequence, except that chert fragments are absent in the brecciated diabase. The lower part is distinct in that it contains significant amounts of biotite, quartz, and K-feldspar(?). In this respect it is similar to the up-



Figure 9. Igneous rock sequence at Site 471.

per portion of the diabase sill at Site 469. The mineralogic constituents of this lower part include, in decreasing order of abundance, plagioclase, clinopyroxene, brownish green smectite, opaque minerals, biotite, quartz, K-feldspar(?), colorless amphibole, sphene, calcite, and zeolite. Intersertal to subophitic textures prevail, with euhedral to subhedral plagioclase laths partly embayed into anhedral clinopyroxenes. Plagioclase is strongly zoned and ranges in grain size from 0.5 to 4 mm. Clinopyroxene has faint purple pink tints; several grains are further fringed by pale green clinopyroxene rinds. Biotite, quartz, K-feldspar(?), and colorless amphibole all occur in the interstices between plagioclase laths and clinopyroxenes. Intense alteration of K-feldspars to clays precludes their positive identification; however, they appear to be confined to Cores 84 to 86. Despite the distinct mineralogy of this lower part of the lower diabase sequence, we make no further subdivision because of the absence of a chilled zone.

The brecciated diabase could have formed by intense weathering along incipient fractures. In addition, some of these rocks (Core 83) have a laminated fabric similar to cataclastic rocks, although microscopically, constituent minerals do not show such strain effects as marginal granulation, undulatory extinction, or bending of cleavage cracks or twin lamellae. Some fragmented crystals of plagioclase do occur, however, suggesting minor cataclastic deformation. Possibly these diabase breccias formed by autobrecciation, a fragmentation process whereby portions of the first-consolidated crusts of intrusions or flows are incorporated into the still-molten interior.

Because of the absence of pillow structure, the absence of microscopic quench texture that would indicate rapid chilling of magma against cool water, the coarse grain size, ore mineralization in the overlying sedimentary rocks, and the inclusion of sedimentary rock fragments in the brecciated diabase, the diabase sequence at Site 471 is probably intrusive. This sequence is probably a composite of three or four cooling units, with each unit representing a thin sill or sheet. Alternatively, these cooling units may correspond to offshoots of a single, larger intrusive.

BIOSTRATIGRAPHY

Pleistocene through middle Miocene sediments were recovered at Site 471. Planktonic foraminifers and coccoliths are common to abundant and provide stratigraphic control in the upper 45 meters (Cores 1–5) and 58 meters (Cores 1–6) respectively. In Cores 7 through 17, calcareous microfossils are absent, and radiolarians and diatoms provide stratigraphic control (Fig. 10). Microfossils are essentially absent from Cores 18 through 38, although sparse radiolarian assemblages provide some stratigraphic control. Coccoliths reappear in Core 39 and are sporadically present to the bottom of the sedimentary section. Benthic and planktonic foraminifers are present but they are extremely sparse in this interval. Coccoliths suggest placement of the Pliocene/Quaternary boundary just below Sample 3,CC and the lower Pliocene/upper Pliocene boundary between Cores 5 and 6. Figure 8 summarizes zone determinations for Site 471. The top of the *Sphenolithus heteromorphus* Zone is in Core 78, about 3 meters above the uppermost igneous rock. The age of the oldest sediment at Site 471 is thus estimated to be about 15 m.y.

Coccoliths

Coccoliths recovered at Site 471 represent only short intervals of time. In Core 79, just above basalt, the oldest assemblages belong to the upper Sphenolithus heteromorphus Zone (approximately 14 m.y. of age; Bukry, 1975). The overlying Coccolithus miopelagicus Subzone (approximately 13.4 to 14 m.y. old) extends through a thick interval from Core 39 to Core 78 (368-737 m), but the short time span involved yields an exceptionally high sedimentation rate (383 m/0.6 m.y. = 640)m/m.v.). Reworked Cretaceous and Paleogene coccoliths occur (Fig. 11) in this interval (Lithologic Unit 4), which is characterized by turbid flow sedimentary structures. An equally thick interval in Cores 6 to 39 (55-368 m) is barren of coccoliths. Pleistocene to lower Pliocene coccoliths are sparse to abundant in the upper cores. Sample 1,CC to Core 6 (9.5-52 m), representing the interval from approximately 0.5 m.y. to 4 m.y. ago. The upper part of Core 1 was not investigated.

All sediment layers intercalated with the igneous rocks of Cores 80 to 88 that were examined for coccoliths are barren. The lower middle Miocene Sphenolithus heteromorphus Zone assemblages of Cores 78 and 79 are similar to overlying assemblages, except for the presence of Sphenolithus heteromorphus Deflandre and the more common Cyclicargolithus floridanus (Roth and Hay); and these assemblages show the only notation for overgrowth (+2 to +3; Bukry, 1973) at Site 471, in Sample 471-79-1, 7 cm (at a depth of 751 m). This sample also contains some fragments of pyritized centric diatoms.

Dark olive clay-rich siltstone directly above blue gray turbiditic sandstones of Cores 39 to 78 are in the *Coccolithus miopelagicus* Subzone and yield the largest, most diverse coccoliths. Lighter-color claystone from higher in the turbidite sequences yields mostly smaller, less diverse coccoliths. Many sediment layers are barren or poor, producing an uneven record of abundance and preservation through the subzone (Fig. 10). Reworked Cretaceous coccoliths in Cores 39 to 79 appear to represent Campanian to Maestrichtian horizons (Fig. 11). Similar coccoliths occur onshore near San Diego, California (Bukry and Kennedy, 1969). Paleocene coccoliths, such as *Discoaster multiradiatus*, are from upper Paleocene strata.

Lower Pliocene coccoliths of Cores 5 and 6 are sparse to common and etched. The assemblages are assigned to the Sphenolithus neoabies Subzone, because Reticulofenestra pseudoumbilica (Gartner) and Sphenolithus abies Deflandre are present and Discoaster tamalis



Figure 10. Plots of relative abundances of planktonic microfossils at Site 471.

Kamptner and Amaurolithus spp. absent. One reworked A. delicatus Gartner and Bukry occurs in Sample 471-5, CC.

The Pliocene/Pleistocene boundary may be in a condensed section or cut out by a hiatus, because the *Gephyrocapsa caribbeanica* Subzone of Sample 471-4-1, 0 cm (28.5 m depth) overlies the *Discoaster surculus* Subzone of Sample 471-4-2, 52-53 cm (32 m depth).

Pleistocene assemblages contain Gephyrocapsa oceanica Kamptner with high-angle bars and Coccolithus pelagicus (Wallich), indicating temperate conditions.

Silicoflagellates

Lower Pliocene and upper Miocene silicoflagellates are sparse to common and well preserved in Cores 6 to 18. They are absent in the calcareous Quaternary and Pliocene silty clay of Cores 1 to 5 and in the Miocene

silty clays of Cores 19 to 78, although sparse fragments of diatoms Coscinodiscus marginatus Ehrenberg and Thalassiothrix longissima Cleve and Grunow as well as pyritized centric diatoms occur through Cores 20 to 28. The upper silicoflagellate assemblages of Cores 8 and 9 are especially diverse, and some reworking is suggested by the presence of Distephanus mesophthalmus (Ehrenberg) (upper middle to lower upper Miocene, according to Dumitrica, 1973) and the diatom Craspedodiscus coscinodiscus Ehrenberg (lower or middle Miocene). Terrestrial opal addition to the sediment is indicated by the presence of sparse panicoid opal phytoliths in Sample 471-8, CC. Mesocena sp. aff. M. quadrangula Ehrenberg ex Haeckel and Dictyocha sp. (naviculopsoid) in Core 8 suggest correlation with the upper Miocene upper Discoaster quinqueramus Zone of coccoliths at DSDP Site 158 in the eastern equatorial Pacific



Figure 11. Upper Cretaceous and Paleogene reworked coccoliths at Site 471.

(Bukry, 1973). Although silicoflagellates are generally common in Cores 10 to 18, diversity is low and assemblages are composed mainly of *Dictyocha brevispina* (Lemmermann), *D. fibula* Ehrenberg, and *Distephanus speculum* (Ehrenberg). In most samples *Dictyocha fibula* predominates over *D. brevispina*, suggesting the warm-water *Dictyocha fibula* Zone of Martini (1971). Samples 471-13,CC and 471-15,CC, however, have more *D. brevispina* than *D. fibula*, possibly indicating the top of the *D. brevispina* Zone, a unit correlative with the upper *Discoaster quinqueramus* Zone or upper *D. neohamatus* Zone of coccoliths in the eastern equatorial Pacific (approximately 7 m.y. old) (Bukry and Foster, 1973; Bukry, 1973).

Radiolarians

At Site 471, the interval between Cores 1 and 5 is nearly barren of radiolarians. Sections 3 and 5 of Core 1 contain only rare, nondiagnostic radiolarians, and Section 1 yielded *Amphirhopalum ypsilon* and (?)Ommatartus tetrathalamus, two Quaternary to Pliocene species. Because the latter section does not contain any older radiolarians other than rare individuals of Stichocorys peregrina, it is tentatively assigned to the Quaternary. The boundary between the Quaternary and Pliocene is placed immediately above the youngest fossiliferous layers (Core 5, Section 6) with Lamprocyrtis *heteroporos*. It must be emphasized, however, that the true position of that boundary at Site 471 cannot be found by means of radiolarians.

Deposition during the late Pliocene is indicated for Core 5, Section 6 through Core 6, Section 1 by rare occurrences of *L. heteroporos* and the absence of *Eucyrtidium matuyamai*. The extinction of *Stichocorys peregrina* marks the boundary between the upper Pliocene *L. heteroporos* Zone and the lower Pliocene *S. peregrina* Zone (upper part), which comprises the interval between Core 6, Section 3 and Core 10, Section 1. Mostly few to common, moderately to well preserved radiolarians typify the lower Pliocene succession. Its lower limit is indicated by the first appearance of *L. heteroporos*. Besides rare to abundant *S. peregrina*, *Ommatartus penultimus* is another rare to common species that survived above the Miocene/Pliocene boundary.

The upper Miocene cores of Hole 471 contain a radiolarian assemblage rather rich in species that are almost exclusively warm-water varieties. The equatorial zonation is therefore applicable. The presence of S. peregrina, S. delmontensis, S. wolffii, O. penultimus, and O. antepenultimus and the absence of L. heteroporos suggest that Core 10, Section 3 through Core 17, Section 1 are upper upper Miocene (lower part of the S. peregrina Zone). A separation of the upper upper Mio-

cene S. peregrina Zone and the middle upper Miocene O. penultimus Zone is not possible at Site 471, because the first appearances of S. peregring and O. penultimus nearly coincide, and because the occurrences of S. delmontensis and S. peregrina overlap considerably. There are no radiolarian species remaining to define the upper limit of a possible O. penultimus Zone. The lower boundary can be defined exactly, however, by the extinction of Ommatartus hughesi and the first appearance of O. penultimus and can be placed below Core 17, Section 1, provided that the two events are not affected by dissolution or reworking. Besides several nondiagnostic species, Sample 471-17,CC contains S. delmontensis, S. wolffii, and O. antepenultimus, and Core 23, Section 1 yielded rare to few specimens of O. antepenultimus, O. hughesi, and Cannartus petterssoni (O. antepenultimus Zone). Samples 471-23,CC and 471-28,CC contain very rare and poorly preserved radiolarian assemblages deposited during the early late Miocene. All of the remaining cores of Site 471 lack radiolarians or bear only indeterminable fragments, frequently with heavy crystalline overgrowths.

Diatoms

Few to abundant Pliocene to upper Miocene diatoms are present in Cores 6 through 18 at Site 471. Nearshore diatoms, such as *Actinocyclus ehrenbergii*, *Rhaphoneis* spp., and *Thalassiosira* spp., are present in Core 1 but are not age-diagnostic. Diagenesis, which is first evident in Cores 17 and 18, removed all but robust, non-agediagnostic diatoms below Core 18.

The first occurrence of *Thalassiosira oestrupii* in Sample 471-9-1, 58-60 cm is indicative of a horizon slightly higher than the Miocene/Pliocene boundary. Sample 471-10-1, 20-22 cm contains relatively common *Dictyocha navicula*, a silicoflagellate with an acme across the Miocene/Pliocene boundary.

Planktonic diatoms that are used in biostratigraphy are relatively rare and sporadic in the upper Miocene of Site 471. *Thalassiosira miocenica* and *Nitzschia miocenica* in Sample 471-11-1, 20-22 cm suggest correlation with paleomagnetic Epoch 5 or uppermost Epoch 6. The last occurrence of *Cussia praepaleacea* in Sample 471-13-2, 20-22 cm indicates equivalence with the upper part of Epoch 6. *Nitzschia miocenica* is recorded from Sample 471-17,CC and argues for correlation with a level no older than the upper part of Epoch 7. Benthic diatoms are especially common in Core 12, where they mask many of the planktonic forms.

Foraminifers

Planktonic foraminifers are common and well to moderately well preserved in the upper 44.8 meters of Hole 471 (Samples 471-1,CC through 471-5-5, 80 cm). Cores 1 through 3 are assigned to the Quaternary because of the occurrence of *Globorotalia truncatulinoides* in Sample 471-1,CC, and *Neogloboquadrina dutertrei* in Samples 471-1,CC through 471-3,CC. An upper Pliocene assemblage containing *Neogloboquadrina humerosa*, *N.* aff. *N. atlantica*, and *Globorotalia inflata* was recovered from Sample 471-4,CC, whereas Sample 471-5-2, 80-82 cm, with *Globoquadrina altispira* and transitional forms between *Globorotalia* puncticulata and *G. inflata*, is considered close to the Zone N21/Zone N19 boundary.

Planktonic assemblages in Samples 471-1,CC through 471-5-2, 80-82 cm are diverse, and tropical to sub-tropical taxa such as *Pulleniatina obliquiloculata* and *Sphaeroidinella dehiscens* as well as keeled *Globorotalia* and *Globigerinoides* occur throughout this interval.

Sample 471-5-5, 80-82 cm yields a sparse, strongly dissolved assemblage, and no foraminifers are present in Sample 471-5, CC. Foraminifers are absent from Cores 6 through 37. From Core 38 to Core 78, occasional foraminifers were seen on cut surfaces of cores. Examination of selected samples from this interval revealed rare, poorly preserved planktonic foraminifers; only long-ranging taxa such as *Globigerina bulloides* and *Globigerinoides trilobus* were found.

Benthic foraminifers are sparse but well preserved in the first five cores of Hole 471. Uvigerina senticosa, Pullenia bulloides, and Melonis barleeanus are consistent members of the assemblage in this interval and indicate lower middle to lower bathyal water depths. Benthic foraminifers are absent or extremely rare in Cores 6 through 37.

Below Core 37, representatives of Gyroidina, Pyrgo, Stilostomella, Globulimina and (?)Bolivina were occasionally seen on cut surfaces of cores. Benthic foraminifers are sometimes concentrated toward the top of sandstone layers in the turbidite sequence of Cores 38 to 78, but they are poorly preserved, and most specimens appear to be internal casts. Occasional specimens of Pullenia bulloides and Gyroidina soldanii in samples from fine-grained sediments of Cores 37 to 78 suggest this sequence was deposited in bathyal water depths.

SEDIMENT ACCUMULATION RATES

The sediment accumulation rate curve (Fig. 12) for Site 471 was constructed from selected radiolarian (R) and coccolith (C) events. The plot indicates rates of ~ 20 m/m.y. for the Quaternary into the late Miocene, ~ 50 m/m.y. for the early late to late middle Miocene, and ~ 200 m/m.y. for the early middle Miocene. No stratigraphically diagnostic microfossils are present for more than 100 meters above the coccolith control point at 360 meters sub-bottom (Fig. 8). Thus the accumulation rate calculated for the lower portion of Site 471 represents a minimum estimate.

GEOCHEMICAL MEASUREMENTS

Interstitial Water

The salinity, chlorinity, pH, alkalinity, and calcium and magnesium concentrations of interstitial waters from nine depths at Site 471 were determined on board. The results are plotted versus depth at the site in Figure 13. No samples were taken between 125 and 360 meters sub-bottom depth; in this interval, only porcellanites and dolomites were recovered. Although the calcium concentration profile at Site 471 shows a gradual increase downcore, the magnesium concentration profile



Figure 12. Sediment accumulation rates, Site 471.

shows a marked discontinuity between normal sea water magnesium concentrations above the porcellanites and very low magnesium concentrations below the porcellanites. Salinity and chlorinity also decrease markedly over the same depth interval. The change in magnesium concentration over the interval from which porcellanites were recovered suggests that the decrease was the result of diagenetic reactions involved in the transformation of biogenic opal to opal-CT porcellanite and/or in the formation of dolomite.

Calcium Carbonate Content

The calcium carbonate concentration in samples from Site 471 was determined on board by the carbonate bomb technique. The results of these determinations are included in the core descriptions in this chapter and plotted in Figure 14. Most of the sediment column at this site is noncalcareous. There are, however, a few dolomite interbeds in Unit 3 that had carbonate contents of 99% to 100%. The silty claystones of Unit 4 are also noncalcareous, but they are interbedded with calcareous sandstones that contain 14% to 21% CaCO₃.

PHYSICAL PROPERTIES AND DOWNHOLE LOGS

Figure 14 plots the physical-properties data for Site 471, and Table 3 summarizes these data. Figure 15 displays the downhole logs obtained at this site. Relatively constant values of density ($\sim 1.5 \text{ g/cm}^3$) and velocity ($\sim 1.5 \text{ km/s}$) characterize the upper 150 meters of soft nannofossil silty clay and diatomaceous clay. A sharp increase in both parameters occurs at 155 meters, the depth at which soft diatomaceous sediment begins to convert to harder porcellanite. The density of the porcellanite at this depth is about 2.0 g/cm³ but decreases slightly to 1.7 g/cm^3 at about 317 meters. Sonic velocities show a similar trend, decreasing in this interval from 3.6 km/s to 2.1 km/s. This change corresponds to a noticeable decrease in the siliceous character of the



Figure 13. Interstitial water profiles, Site 471.



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| | Sub-bottom Depth | Density | (g/cm ³) | Velocity | Average Impedance | |
|-----------------------------|------------------|-----------|----------------------|-----------|----------------------|----------|
| Lithology | (interval in m) | Range | Average | Range | Average | Contrast |
| Silty clay | 0.0-63.5 | 1.4-1.5 | 1.5 | | 1.5 | 0.01 |
| Clayey diatomaceous ooze | 63.5-155.2 | 1.4-1.5 | 1.5 | 1.52-1.55 | 1.54 | 0.01 |
| Porcellanite/chert | 155.2-304 | 1.62-2.07 | 1.82 | 1.74-3.93 | 2.70 | 0.30 |
| Carbonate layers | 180-450 | 2.52-2.79 | 2.68 | 3.99-6.09 | 5.24 | 0.46 |
| Silty claystone | 304-741 | 1.68-2.13 | 2.07 | 1.85-2.75 | 2.13 | 0.32 |
| Calcareous sandstone layers | 445-741 | 2.12-2.98 | 2.48 | 2.80-4.71 | 3.46 | 0.32 |
| Altered diabase | 741-T.D. | 2.29-2.78 | 2.58 | 3.19-5.35 | 4.36 | 0.43 |

Table 3. Summary of laboratory measurements of velocity and density and calculated impedance contrasts of rocks and sediments recovered at Site 471.

rocks—porcellanite changes downhole to less siliceous porcellaneous silty claystone. The Sonic and Density Logs show these variations most clearly. Moreover, all logs demonstrate the occurrence of soft sediment interbedded with porcellanite between 155 and 317 meters. Recovery was only about 5% in this interval, and much of this soft sediment was not recovered.

The thick monotonous sequence of silty claystone below 317 meters has fairly uniform physical-properties profiles, except where it is punctuated by thin layers of clayey dolomite and calcite-cemented sandstone. The Neutron Log gives the best resolution of the thin sandy layers. The two peaks at about 598 meters on this log correspond to thin sandstone layers. The Density and Sonic Velocity Logs also delineate these thin layers. From a physical-properties standpoint, the pelagic claystone at the base of this section of silty claystone is indistinguishable from the overlying unit.

A sharp increase in density and velocity occurs at 741 meters, the top of the altered diabase. These two parameters fluctuate significantly with depth in the diabase and probably correspond to variations in the degree of alteration and fracturing of this unit (see also the Seismic-Spectrum Log, Figure 15). The Neutron Log is also quite variable in this unit. A thin interval of metalliferous sediment occurring at about 780 meters in the diabase shows clearly on the Density, Sonic, and Neutron Logs. Below about 812 meters, neutron, velocity, and density values decrease, perhaps marking an interval of more altered diabase extending to the bottom of the hole.

Table 3 lists impedance contrasts calculated from the laboratory determinations of density and velocity. Clearly the boundary between the porcellanite of Unit 3 and the overlying soft diatomaceous clay is a strong reflector. The carbonate layers should also be good reflectors if numerous or thick enough. The claystone/ diabase contact has a high impedance contrast and is a prominent reflector on the seismic profile.

Estimates of *in situ* temperatures and geothermal gradients at Site 471 come from three sources: (1) heat flow probe measurements made during drilling; (2) two Temperature Logs run about 18 hours apart in the open hole after completion of drilling, and (3) bottom-hole temperatures recorded with a set of three maximum temperature thermometers attached to each logging tool. Figure 16 illustrates the results of the two heat flow probe

measurements made at Site 471. The instrument used was the Uveda/Kinoshita probe first tested on DSDP Leg 60 (see Hussong, Uyeda, et al., in press for details and discusssion of this instrument). We made two measurements with the probe in the sediment at 95 meters and 142.5 meters, with additional stations in the pipe at mudline to estimate the temperature of the bottom water. Both runs show a similar pattern of decreasing temperature as the probe is lowered to the sea floor, an increase as it approaches the maximum drilled depth, a plateau corresponding to penetration into the sediment, then cooling and finally warming trends as the probe is pulled from the hole. The interval labeled "on bottom" in each run corresponds to the time the probe was actually in the sediment. Hyndman et al. (1974) note that temperature decreases (cooling curve) when frictional heating brings the temperature of the probe above that of surrounding sediment; temperature increases (warming curve) when frictional heating is insufficient to raise the temperature of the probe above that of the sediment. Throughout the measurement at 95 meters, the pipe was lowered gradually to maintain pressure on the bit. The cooling curve shown in Figure 16 suggests that frictional heating affects the temperature measurement. For the measurement at 142.5 meters, the constant temperature (prevailing 42 minutes after the last lowering of the pipe) indicates frictional heating is not important.

In situ sediment temperatures estimated from Figure 16 are about $12.5^{\circ}C$ at 95 meters and about $24.0^{\circ}C$ at 142.5 meters. Using a bottom-water temperature of 2.0°C, we determined that the geochemical gradients are $110^{\circ}C/km$ and $154^{\circ}C/km$, with an average of $132^{\circ}C/km$.

We also estimated the geothermal gradient at Site 471 using the Temperature Logs (Figure 15). Both Temperature Logs show initially steep gradients in the porcellanite, ~ 40 °C/km for the first run and ~ 74 °C/km for the second. A distinct change to lower gradients (30 °C/ km and 67 °C/km, respectively) occurs below about 345 meters. These latter values remain fairly uniform through the thick section of silty claystone, except for an interval between 520 and 570 meters, where gradients decrease slightly. This interval corresponds to the first show of gas in the cores; gas escaping into the open hole and expanding could explain this zone of lower gradients. The second Temperature Log probably more closely approximates the equilibrium gradient than does



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Figure 16. Heat-flow-probe runs and equilibrium temperatures at Site 471. (See text for explanation.)

the first run because of the added time between the last circulation and logging. From this reasoning, the average minimum gradient is 70°C/km. This value is lower than the estimates made from heat probe measurements, because the logging tool does not actually penetrate the sediment and because of the thermal disturbance created by the drilling process (Jaeger, 1965). Note that the second Temperature Log recorded lower values than did the first log in the upper 150 meters of open hole and higher values below this, an indication of how the hole re-equilibrates following drilling.

A third estimate of the geothermal gradient at Site 471 comes from extrapolating maximum bottom-hole temperatures recorded during each of the logging runs; this was done according to the method described by Timko and Fertl (1972). Figure 17 illustrates this method for Site 471, and Tables 4 and 5 list the pertinent data. From this method the equilibrium bottomhole temperature is about 68.5°C. Using a 2.0°C bottom-water temperature and a total depth of 817 meters for the hole, the geothermal gradient is 81°C/km. Again, thermal disturbance resulting from drilling and the lack of penetration of the logging tool tend to make this value a minimum.

From these three methods the range in geothermal gradient at Site 471 is 70°C/km to 154°C/km. These



Figure 17. Equilibrium bottom-hole temperature estimated from successive logging runs at Site 471. (See text for explanation.)

values are approximate because both heat probe data and Temperature Logs suggest a nonlinear gradient. A single value is difficult to estimate, because both methods have significant uncertainties—heat flow data suffer from the effects of frictional heat generated by lowering the probe and drill pipe in the sediment, and logging data are nonequilibrium values influenced by thermal disturbances associated with drilling. The average value from the four estimates is 100°C/km.

We did not measure the thermal conductivity of the sediment cored at Site 471 and therefore can only estimate heat flow. Using the range in gradients listed above and an average thermal conductivity of 2.5 mcal/cm s °C estimated from water content, we deduce that the range in estimated heat flow values at Site 471 is 1.8 to 3.9 HFU.

CORRELATION OF REFLECTION PROFILE WITH DRILLING RESULTS

Figure 18 correlates the seismic-reflection profile obtained by the *Glomar Challenger* while approaching Site 471 (Fig. 7) with the lithology at Site 471. Above acoustic basement, two acoustic units are recognized:

1) An upper acoustic unit of strong, continuous, evenly spaced, and somewhat indistinct reflectors about 0.19 s thick.

Table 4. Chronology of logging operations.

| Date | Time | Operation | | | | | |
|---------|------|---|--|--|--|--|--|
| Nov. 15 | 1730 | Last circulation of mud. | | | | | |
| | 2038 | Start down with sonic tool, | | | | | |
| Nov. 16 | 0030 | End Sonic Log, begin Wave-Train Log. | | | | | |
| | 0050 | Start up with Wave-Train Log; 1st maximum bottom hole temperature (max BHT). | | | | | |
| | 0325 | Tool on deck, log complete. | | | | | |
| | 0340 | Start down with temperature-density tool, logging temperature on way down. | | | | | |
| | 0745 | Start up with Density Log; 2nd max BHT. | | | | | |
| | 1035 | Tool on deck. | | | | | |
| | 1100 | Start down with guard Neutron Log. | | | | | |
| | 1310 | Start up with Neutron Log; 3rd max BHT. | | | | | |
| | 1550 | Tool on deck. | | | | | |
| | 1610 | Start down with Induction-Gamma-Ray Log. | | | | | |
| | 1930 | Start up with Induction Log; 4th max BHT. | | | | | |
| | 1945 | Raise pipe 1 stand (28.5 meters) to log soft sediment above porcellanite; Gamma-Ray Log run in pipe to mudline. | | | | | |
| | 2150 | Tool on deck. | | | | | |
| | 2205 | Start down with final Temperature Log. | | | | | |
| | 2304 | Begin recording at 100 meters (water column). | | | | | |
| Nov. 17 | 0015 | Repair cable. | | | | | |
| | 0050 | Continue logging. | | | | | |
| | 0200 | End Temperature Log; final max BHT. | | | | | |
| | 0406 | Temperature tool on deck. Rig down. | | | | | |

2) A darker, lower acoustic unit of very strong, con-

The upper acoustic unit correlates with the Quater-

tinuous reflectors about 0.58-s thick that becomes trans-

nary to upper Miocene nannofossil and diatomaceous

silty clay and ooze of Lithologic Units 1 and 2. The

discordant cutting of reflectors in this upper unit by the

underlying acoustic unit correlates with the marked

lithologic and diagenetic change from diatomaceous

ooze to porcellanites at Site 471. This break marks a

bottom-simulating reflector (BSR) similar to that

described at DSDP Sites 184, 185, and 188 (Scholl and

Creager, 1973). This BSR can be traced laterally on the

relates with the upper to middle Miocene porcellanite

and silty claystone and interbedded sandstone (tur-

bidite) sequence of Lithologic Units 3 and 4. A slight angular unconformity at about 0.35 s sub-bottom may

Beginning with the BSR, the lower acoustic unit cor-

approaching Challenger profile for more than 80 km.

parent in the lower part.

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correlate with the boundary between these two lithologic units.

The lower acoustic unit becomes increasingly more transparent near its base, which may be the result of a corresponding decrease in the frequency of carbonatecemented sandstone and carbonate beds downsection in Lithologic Unit 4. The corresponding Sonic and Density Logs (Fig. 15) show only minor density and velocity variations within the lower half of Lithologic Unit 4.

Acoustic basement at 0.77 s sub-bottom correlates with basalt, which can easily be traced landward on the seismic-reflection profile to the base of the continental slope. The landward thickening wedge defined by the sea-floor acoustic basement may be an older trench sequence that has undergone only mild deformation since the cessation of subduction along this portion of the margin. Its remarkably well-preserved character suggests that the ancestral East Pacific Rise never intersected the trench along this portion of the Baja California margin.

CONCLUSIONS

1. The oldest sediments overlying diabase at Site 471 are located in the Sphenolithus heteromorphus Zone and are about 15 m.y. of age. These sediments are older than a magnetic anomaly 12 m.y. old that was recognized west of the site by Chase et al. (1970). Magnetic anomalies at the site are low in amplitude and have not been dated. Extrapolation of ages eastward from areas of recognizable magnetic anomalies would predict an age of 11 m.y. at the site. It is demonstrated, therefore, that the age of oceanic crust is older eastward toward the continental escarpment off southern Baja California. This region where oceanic crust is older to the east may be a fragment of the Farallon plate that was not subducted beneath Baja California. The interaction between the East Pacific Rise and North America is more complex at this latitude than previously believed.

2. The lower sedimentary sequence contains 13 to 15 m.y. old distal turbidites (predominantly Bouma d, e sequences) that were deposited at a rate of about 200 m/m.y.; these turbidites are overlain by hemipelagic sediments deposited in the latest middle Miocene and later at rates decreasing upsection from 50 to 20 m/m.y. The siliceous component (altered diatomaceous sediments) is more pronounced upsection, with lower sedi-

Table 5. Bottom-hole temperatures from three maximum reading thermometers mounted on the logging tools, Site 471.

| | | | Distance of Thermometers above Base | Dep Measurer | th of nents (m) | - | |
|---------|------|---------------------|---|--------------------|--------------------|----------------|---------------|
| Date | Time | Tool | of Tool (m) | below Rig Floor | below Sea Floor | Temper (°F) | ature (°C) |
| Nov. 16 | 0050 | Wave Train | 5.5 | 3932 | 817 | 95,95,95 | 35,35,35 |
| | 0745 | Temperature-Density | 3.1 | 3935 | 819 | 122,122,122 | 50,50,50 |
| | 1310 | Neutron | 7.3 | 3933 | 817 | 117,122,126 | 47,50,52 |
| | 1930 | Induction | 6.1 | 3932 | 817 | 136,136,138 | 58,58,59 |
| Nov. 17 | 0200 | Temperature | 5.8 | 3932 | 817 | 144,145,146 | 62,63,64 |



Figure 18. Correlation of Challenger seismic-reflection profile with lithology at Site 471.

mentation rates. This change reflects a decrease in the supply of terrigenous material, which may have resulted from a change in vertical uplift rates of the adjacent margin or translation of the site away from major river sources. Alternatively, the increase upsection of siliceous sediments may reflect increased productivity in surface waters.

3. Chalcopyrite- and sphalerite-bearing sediments occur at the base of the sediment section; this occurrence may be an ancient analog to sulfide deposits found on the East Pacific Rise Crest and in related fracture zones.

4. The sediments are underlain by 81.5 meters of diabase with two thin intercalations of metalliferous sediment. The diabase is extensively altered, with common brecciated structures that probably resulted from the intrusion of magma into the sediment; this is indicated by the presence of clayey to cherty inclusions in the breccia. Slickensides show some faulting after consolidation. Fracturing in the diabase, as observed in cores, is substantiated by the Sonic Wave-Train Log and the Density Log.

5. The change from diatomaceous ooze to opal-CT porcellanite and porcellaneous rocks occurs at 155 meters at a present subsurface temperature of about 20°C. The transition is reflected in downhole logs and is in agreement with diagenetic changes in siliceous sediments that were observed at other Leg 63 sites (467, 468, 469, and 473).

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| SITE 471 HOLE CORE 1 CORED INTERVAL | L 0.0–9.5 m | SITE 471 HOLE CORE 2 CORED INTERVAL 9,5-19,0 r | 0 |
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| UDU_INIT Status GRAPHIC LITHOLOGY Status | PHIC | c | FO | RAC | L TER | | | | | | |
|--|--|--------------|--------------|--------------|----------|---------|--------|----------------------|--|---------|---|
| Image: Second | UNIT UNIT BIOSTRATIGRA ZONE | FORAMINIFERS | NANNOFOSSILS | RADIOLAHIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | LITHOLOGIC DESCRIPTION |
| 3 Heavy minerals TR 3 Clay 27 Pyrite 2 Carbonate unspec. 2 Foraminiferen TR | LOWER PLIOCENE Stichocory pregrina (upper part) (R) PLIOCENE (D) | | в | | | 2 | 0.5 | | 0 | | CLAYEY DIATOMACEOUS OOZE, gravith olive green (5GY 3/2). Drilling deformation intense. No Core-Catcher, ORGANIC CARBON AND CARBONATE 141 % Organic Carbon – % CoCO ₃ 0 SMEAR SLIDE SUMMARY 2-78 (D) TEXTURE: Sand – Slit 63 Clay 27 COMPOSITION: Quartz 2 Feldspar TR Heavy minerais TR |





| SITE 47 | 71 HO | LE | col | RE | 12 CORED | INTER | AL: | 104.5–114.0 m | SITE | 471 | HC | LE | | COR | 1 | CORE | D INTER | VAL 114.0-123.5 m | | |
|-------------------------------|---|-------------------|-----------------------------|--------|----------------------|---------|--------|--|---------------------|--|--------------|----------------------------------|----|------------------|---|----------------------|--|-------------------|---|--|
| TIME-ROCK UNIT BIOSTRAT | FORAMS P | SOSSIL ARACTER | SECTION | METERS | GRAPHIC LITHOLOGY | DELLING | SAMPLE | LITHOLOGIC DESCRIPTION | TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS | FOSSIL ARACTI SNVIUE TOIDE | ER | SECTION | | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY | SAMPLES | LITHOLOGIC D | ESCRIPTION |
| UPPER MIOCENE | Commatantus peruntrimus to Stichocorys penegrica (lower part) (H) UPPER MIOCENE (D) | FM AG | 3 4 5 6 7 tc | | | | • | DiaToMACEOUS CLAY, dusky yellow green (EGY 5/2) mattied with dark yellowith brown 110YR 4/21 patches, Gravith purple (54/2) printer-ich strakes, scattered throughout, interme diffile deformation has homogenized interlosed and the strakes scattered to compare the strake scattered to compare the strakes scattered to compare the strakes scattered to compare the strakes scattered to compare the strake scattered to compare the scattered to compare the strake scattered to compare the strake scattered to compare the scatter | UPPER MIOCENE | Ommeanter penutrimer to Stichocorps pengrine flower part) (R) UPPER MIOCENE Mitzschie mioconica to Thalessiosira convexa (D) | | | | 3 3 6 7 | | | | | CLAVEY DIATOMA (GGY 5/2) to gravith vellowish brown (10) lithology homogenizi ORGANIC CARBON % Organic Carbon % CaCO ₃ SMEAR SLIDE SUM TEXTURE: Sand Silt Clay COMPOSITION: Quartz Feldspar Mica Heavy minerals Clay Volcanic glass Pyrote Carbonate Unspec. Diatoms Radiolarians Stilooftagellates Phant debris | ACEOUS OOZE, dusky vellow green green (10GY 5/2) with some YR 4/2) mottles. Structures and ed by interve drilling deformation. I AND CARBONATE 4-70 0 MARY 2:120 4-106 (D) (D) 5-7 50 80 550 20 3 3 3 2 2 7 R 7R 7R 7R 7R 7R 7R 7R 78 7 78 7 |

SITE 471





SITE 471 HOLE

TIME - ROCK UNIT

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part)

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FORAMINIFERS NANNOFOSSILS RADIOLAPIANS DIATOMS

| ITE | 471 | | но | .E | | _ | co | RE | 16 | CORED | INTE | R | VAL | 142,5-152,0 m |
|--------------------|---------------------------------|--------------|--------------|-----------------------------|-----------|---|---------|--------------|----------------------------|---------------------------|-------------|------------|---------|--|
| × | PHIC | | CHA | OSS | TER | | | | | | | | | |
| TIME - ROC | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | | SECTION | METERS | GR | APHIC HOLOGY | DISTURBANCE | STRUCTURES | SAMPLES | LITHOLOGIC DESCRIPTION |
| | wer part) (R) | | | | | | 1 | 0.5 | 3,9,9,9,9,9,9,9, | | | | | DIATOMACEDUS CLAY, dusky yellowish green (5GY 5/2) with some dark yellow brown (10YR 4/2) streaks. Structures and likhology homoganized by drilling. ORGANIC CARBON AND CARBONATE 2-118 3-80 % Organic Carbon 1.04 – % CaCO ₃ 0 0 |
| ER MIOCENE | us to Stichocorys paragrina (ic | | | | | | 2 | | 111111111111111 | | | | * | |
| UPP | Ommetartus penultim | | | | | | 3 | | | | | | + | |
| | | | в | FM | AG | | 4 | in line line | | | | | | |
| ITE | 471 | | HOL | E | | | co | RE | 17 | CORED | INTE | R | AL | 152.0–161.5 m |
| × | APHIC | | CHA | OSSI | TER | | | | | | 11 | 1 | | |
| TIME - ROC UNIT | BIOSTRATIGR | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | | SECTION | METERS | GR | APHIC HOLOGY | DISTURBANCE | STRUCTURES | SAMPLES | LITHOLOGIC DESCRIPTION |
| OCENE | CENE (D) | | | Ommatartus penultimus to | part) (R) | | 1 | 0.5 | 333333333333 | | | | * | CLAYEY DIATOMACEOUS 002E, mottled dark yellowish brown (10YR 4/2), moderate yellowish brown (10YR 5/4) and paic oliver, homogenized by drilling, Corre-Catcher contains served pieses of greenish black (SGY 2/1) PORCELLANTE: Subbed surfaces show burrows and lenticular laminations. ORGANIC CARBON AND CARBONATE 1-80 % Organic Carbon |
| UPPER MI | UPPER MIO | | | matartus antepenultimus (R) | | | 2 | | \$\\$\\$\\$\\$\\$\\$\\$\}\ | ۵ ۲۲۲۲۲۲۲۲۲۲۲۲۲۲۲۲۲۲۲۲ | | | | % CaCO ₃ 0 SMEAR SLIDE SUMMARY (D) TEXTURE: Sand Silt 90 Clay 10 COMPOSITION: Ountria: Sand 5 Feldipar 2 Mice TR |

| | DHIC | . 8 | CHA | RAC | TER | | | | | | | | |
|--------------------|----------------------|--------------|--------------|--------------|---------|---------|--------|----------------------|-------------|-------------|---------|---|--|
| TIME - ROC UNIT | BIOSTRATIGR/ ZONE | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DISTURBANCE | SEDIMENTARY | SAMPLES | LITHOLOGIC D | ESCRIPTION |
| UPPER MIOCENE | UPPER MIOCENE (D) | | В | В | СМ | 2 | 0.5 | <u>, 8, 8</u> 0 | | ** | | Fragments of olive bi PORCELLANITE. Bi fracture: chert is vitre bedding common. Core-Catcher only. SMEAR SLIDE SUM TEXTURE: Sand Sit Clay Sit CoMPOSITION: Quartz Feldspar Clay Pyrite Carbonate unspec. Diatoms Radiolarians Sponge spicules Opal-CT | ack (5Y 2/1) CHERT AND oth lithologies have conchoidal nous. Burrows and lenticular MARY 1-24 (T) - 8 92 2 1 1 12 2 1 1 2 2 1 1 2 2 1 1 2 2 1 1 2 2 1 1 2 2 1 1 2 2 3 1 2 2 3 1 2 2 3 3 3 3 |

SITE 471 HOLE CORE 19 CORED INTERVAL 171.0-180.5 m

| × | PHIC | | CHA | OSS | TER | 3 | | | | | | | |
|--------------------|--------------|--------------|--------------|--------------|---------|---|---------|--------|----------------------|----------|-------------|---------|---|
| TIME - ROC UNIT | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIAMS | DIATOMS | | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY | SAMPLES | LITHOLOGIC DESCRIPTION |
| | | | | B | в | | cc | | | 1 | ** | | PORCELLANITE gravish plive (10Y 4/2) to plive black |
| | | | | | | - | 1 | 0.5 | | | | | (5Y 2/1); common burrows and some lenticular bedding. Scattered sponge fragments and silty layers. Core brecciated by drilling. |
| | | | | | | | | 1.0 | | | | | Core-Catcher only. |

| SITE 47 | 1 HOLE | CORE 20 CORED INTERVAL 180 | 0,5–190,0 m | SITE 471 HOLE CORE 22 CORED INTERVAL 199,5-209,0 m |
|---|---|---|--|--|
| TIME - ROCK UNIT BIOSTRATIGRAPHIC | FORMINIFERS CHARACTER NANNOFORMAN RADIOLARIANS RADIOLARIA | GRAPHIC ULTHOLOGY ULTHOLOGY SHITTING SHITING SHITTING SHITTING SHITTING SHITTING SHITTING SHITTING SHI | LITHOLOGIC DESCRIPTION | FOSSIL CHARACTER UNIT SUC SUC SUC SUC SUC SUC SUC SUC SUC SUC |
| | B B RP | | CLAYEY LIMESTONE AND PORCELLANITE, Limestone is light olive gray (5Y 5/2), intensely burrowed and partly alicified(?). Porcellanies is grayiah olive (10Y 4/2) to olive black (5Y 2/1), burrowed and fractured by drilling. In contrast limestone part of core is more intact. No Core-Catcher. ORGANIC CARBON AND CARBONATE 18 1-10 % Organic Catbon 3.98 - % CGO3 100 100 SMEAR SLIDE SUMMARY 145 | 1 AAAAAAA A A AAAAAAA A A A PORCELLANITE, medium ofive gray (5Y 4/2), intensely burrowed, Core consists of porcellanite fragments brecclated by drilling. No Core Catcher. ORGANIC CARBON AND CARBONATE 1-21 % Organic Carbon - % CarCO ₃ 0 |
| | | | (D) TEXTURE: Sand Silt 20 Clay 80 COMPOSITION: Owarz TR Feldspar TR Mica TR Heavy minerals TR | SITE 471 HOLE CORE 23 COREDINTERVAL 209.0–218.5 m |
| | | | Clay 15 Pyrite 2 Carbonate unspec, 14 Radiolarians 2 Sponge spicules 2 Opal-CT 65 | B RP RP 1 AAAAAAA 1 + OS AAAAAAA 1 1 + No Core-Catcher. |
| SITE 47 | 1 HOLE | CORE 21 CORED INTERVAL 190 | 0.0–199,5 m | St ORGANIC CARBON AND CARBONATE 133 150 % Organic Carbon 0.34 % CacO3 0 |
| TIME - ROCK UNIT BIOSTRATIGRAPH | ZONE FORAMINIFERS NANNOFOSSILS RADIOLARIANS PLATOMS | RAPHIC SHILLING SHILL | LITHOLOGIC DESCRIPTION | NOTE: Core 24, 218.5–228.0 m: NO RECOVERY. |
| | ввв | | CLAYEY LIMESTONE AND PORCELLANITE. Limestone is light oflive gray (5Y 5/2) and intensety burrowed. Percetilanite is medium olive gray (5Y 4/2) and also burrowed as in Core 20, porcellanite is completely fractured and broken by drilling while limestone is intact. No Core-Catcher, | AND |
| | | | ORGANIC CARBON AND CARBONATE 1-23 % Organic Carbon - % CaCO3 99 | B B B Concernance PonceLLanite, olive gray (5Y 3/2) to light olive gray (5Y 3/2), intensely burrowed, Conchoidally fractured. Mo Core Century No Core Century No Core Century |
| | | | | |

| | PHIC | | CHA | RAC | TER | | | | | | | |
|-----------|--------------|--------------|--------------|--------------|---------|---------|--------|----------------------|--|---------|---|-----|
| UNIT UNIT | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY | SAMPLES | LITHOLOGIC DESCRIPTION | |
| | | | B | в | 8 | 1 | 0.5 | | | • | PORCELLANITE, olive gray (5Y 3/2) to light olive gr (5Y 5/2), intensely burrowed. Most of core is highly | ray |
| | | | в | | | cc | - | | | • | brecolated by drilling, Interval between 32-42 cm is a drilling brecola of porcellanite chips and calcareous silty clay fragments. Pumice fragments also occur in Core-Catcher (e.g. Smar Side CC-29). | |
| | | | | | | 1 | | | | | SMEAR SLIDE SUMMARY | |
| | | | | | | | | | | | 1-38 CC-29 | |
| | | | | | | | | | | | (M) (M) | |
| | | | | | | | | | | | TEXTORE: | |
| | | | | | | | | | | - 1 | Silt 30 - | |
| | | | | | | | | | | | Clay 70 – COMPOSITION: | |
| | | | | | | | | | | - 1 | Quartz 5 TR | |
| | | | | | | | | | | | Feldspar 2 TR | |
| | | | | | | | | | | - 1 | Mica TR - | |
| | | | | | | | | | | - 1 | Heavy minerals TR - | |
| | | | | | | | | | | | Clay 38 - | |
| | | | | | | | | | | - 1 | VOlcanic glass - 100 | |
| | | | | | | | | | | - 1 | Glauconite TR - | |
| | | | | | | | | | | | Pyrite D - | |
| | | | | | | | | | | - 1 | Oest CT(2) 25 | |



| Non-triple No triple No t | | HIC | | F | OSS | TER | | | | TT | TT | | | | |
|---|-------------|--------------|--------------|--------------|--------------|---------|---------|--------|----------------------|--|-----------------------|------------------------|-----------|-----------------------|--|
| B B B CC Reference of the second secon | TIME - HOUN | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY | STRUCTURES SAMPLES | LITHOLOGIC D | ESCRIPT | TON | |
| 0.5- 1 | | | - | B | В | B | CC | | | | •+ | | | | |
| 0.5- 1 1 is olive gray (SY 3/2), burrowed, and brecisted by diffing. Sity class (SY 5/2) and contains some fine sandy laminations. 0.6- 1 1 0.6- 1 0.6- 1 1.0- 2 0.7- 1 0.7- 1 0.7- 1 0.7- 1 0.7- 1 0.7- 2 0.7- 1 0.7- 1 0.7- 3 0.7- 1 0.7- 1 0.7- 5 0.7- 1 0.7- 1 0.7- 7 | - 1 | | | | | | | - | London i Scholl Alto | | | PORCELLANITE A | ND SILT | Y CLAY. Porcellanite | |
| 1 1 1 1 1.0 0 1 1.0 1 1.0 1 0 1 1.0 1 | - 1 | | | | | | | 0.5 | | | 11 | is olive gray (5Y 3/2) | , burrow | ed, and brecciated by | |
| 1.0 ORGANIC CARBON AND CARBONATE CC-20 % Organic Carbon - % CaCO3 1 SMEAR SLIDE SUMMARY 1.9 1.0 1.0 2 SMEAR SLIDE SUMMARY 1.0 1.0 2 Sand 3 Carbon 3 Glay 3 Glay 3 Glay 3 Glay 3 Carbonite 3 Glay 3 Carbonite 3 Carbonite 3 Carbonite 4 1 5 5 6 - 7 Fidigas 1 1 1 Glay 3 Carbonite 4 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 | | | | | | | 1 | | | | 11 | contains some fine sa | andy lami | nations. | |
| 10 00RGANIC CARBON AND CARBONATE CC30 00RGANIC CARBON AND CARBONATE CC30 00 10 10 110 11 110 117 110 117 110 117 110 117 110 117 111 100 111 100 111 11 </td <td>- 1</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td>0.00</td> <td></td> <td>11</td> <td></td> <td>1929-10710222</td> <td></td> <td>11111111111</td> <td></td> | - 1 | | | | | | | 0.00 | | 11 | | 1929-10710222 | | 11111111111 | |
| 3 3 | | | | | | | | 1.0 | | | 11 | ORGANIC CARBON | AND CA | ARBONATE | |
| 1 3 3 1 1 SMEAR SLIDE SUMMARY 10 117 10 (D) (D) (D) 2 Sand - - Sitt 48 13 Clay 52 87 Ourtz 20 7 Feldgar 16 3 1 Haay minerals 3 TR Clay 45 87 Volcanic glass 1 1 Giauconite TR TR 3 - Catonic glass 1 1 Giauconite TR TR 3 - Catonic glass 1 1 Giauconite TR - 3 - Catonic glass 1 1 - - - 3 - - - - - - - - 3 - </td <td>- 1</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td>1</td> <td></td> <td></td> <td>11</td> <td>% Organic Carbon</td> <td>-</td> <td></td> <td></td> | - 1 | | | | | | | 1 | | | 11 | % Organic Carbon | - | | |
| 2 3 SMEAR SLIDE SUMMARY 1-9 1-17 (D) (D) 2 TEXTURE: Sand - - Site 4.6 13 3 Composition: Composition: 0 0 7 7 Feldspir 16 3 Mice 1 4 Heavy minerals 3 TR Clay 45 87 3 - Clay (as 45 87 Clay (as 45 87 Clay (as 45 87 3 - Clay (as 45 87 Clay (as 45 87 Clay (as 45 87 3 - Clay (as 45 87 Clay (as 45 87 Clay (as 45 87 4 Glauconites TR TR Clay (as 45 87 Clay (as 45 87 3 - Clay (as 75 7 - Clay (as 75 7 3 - Clay (as 75 7 - - - - 4 Clay (as 75 <td>- 1</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td>-</td> <td></td> <td>11</td> <td></td> <td>% CaCO₃</td> <td>1</td> <td></td> <td></td> | - 1 | | | | | | | - | | 11 | | % CaCO ₃ | 1 | | |
| 2 3 14 1.17 (0) (0) (0) 2 Sand - Sand - - Silt 48 13 Olay 52 87 OMPOSITION: - - Outriz 20 7 Feldigar 15 3 Mica 1 1 Heavy minaratis 3 TR Clay 45 87 Volcanic glass 1 1 Brunorite TR TR Privite 6 - Carbonatis unippec. 1 - Obtoms TR - | - 1 | | | 1 | | l l | 1 | 10 | | | 11 | | See By | | |
| 2 (D) (D) 2 TEXTURE: Sand - Sitt 48 13 Clay 52 87 COMPOSITION: Outratz 20 7 Feldigar 15 3 Mica 1 1 Heavy minerals 3 TR Clay 55 87 Volcanic glass 1 1 Gisuconite TR TR Gisuconite TR TR 3 Gisuconite TR | | | | | | | | - | | | | SMEAN SLIDE SUM | 1.9 | 1-17 | |
| 2 TEXTURE: Sind - Sint 48 Clay 52 OUMPOSITION: 0 Duartz 20 Faldgar 16 Nica 1 Heavy minarals 3 3 Glauonic glass 45 87 Volcanic glass 1 Glauoniste trapec. - Obtomist TR Diatoms TR | - 1 | | | | | | | 1.2 | | | 11 | | (D) | (D) | |
| 2 Sand - Silt 48 13 Clay 52 87 OMPOSITION: - - Quartz 20 7 Feldgaar 15 3 Mica 1 1 Heavy minarals 3 TR Clay 45 87 Volcanic glass 1 1 Glauconite TR TR Privite 6 - Cationste TR TR Objection TR TR | - 1 | | | | | | 1.0 | | 1 | | 11 | TEXTURE: | | | |
| 3 - Sift 48 13 Clay 52 87 COMOSTION: 0uartz 00 Ouartz 20 7 Feldgaar 16 3 Mics 1 1 Heavy minerals 3 T Quartz 200 7 Feldgaar 16 3 Volcanic glass 1 1 Heavy minerals 3 T Glay outs 1 1 Glay outs 1 1 Glay outs 1 1 1 1 1 Glay outs 1 1 1 1 1 Glay o | - 1 | | | | 1 | | 2 | | 1 | | 11 | Sand | - | - | |
| Clay 52 87 COMPOSITION: COMPOSITION: Outratz 20 7 Peldopar 15 3 Mice 1 1 Heavy minerals 3 TR Clay 45 87 Volcanic glass 1 1 Giauconite TR TR Pyrite 5 - Carbonite TR TR Diatomis TR - | - 1 | | | | I | | 1.1 | - | | 1.1 | 11 | Silt | 48 | 13 | |
| 3 | - 1 | | | | | | | 1.2 | 1 | 11 | 11 | Clay | 52 | 87 | |
| Clustrz 20 7 Feldgar 15 3 Mica 1 1 Henry minerals 3 TR Clay 45 87 Volcanic glass 1 1 Statistica 16 17 Prito 5 - Carbonists TR - Diatomi TR - | - 1 | | | | | | | | 1 | | 11 | COMPOSITION: | | | |
| Feldgar 15 3 Mice 1 1 Haay 3 TR City 45 87 Volcanic glass 1 1 Gisuconite TR TR Pyrite 6 - Catomite TR TR Diatomit TR - | - 1 | | | | | 11 | | | 1 | 11 | 11 | Quartz | 20 | 7 | |
| Mice 1 1 Heavy minerats 3 TR City 45 87 Volcanic glass 1 1 Gisuconita TR TR Pyrite 6 - Carbonic glass 1 1 Object 7 - | - 1 | | | | | | | | 1 | | | Feldspar | 15 | 3 | |
| 3 TR City 45 87 Volcanic glass 1 Gisuronite 3 TR 3 Carrow 4 Carrow 5 Carrow 6 Carrow 7 Carrow 8 Carrow 9 Carrow 1 Carrow | - 1 | | | | | 1 | | | | 11 | | Mica | 1 | 1 | |
| 3 Clay 45 87 Volcanic glass 1 1 Glauconits TR TR Pyrits 6 - Carbonits unspect. - 1 Diatoms TR - | - 1 | | | | | | | | 1 | | | Heavy minerals | 3 | TR | |
| 3 - Carbonats unper - 1 Diatomi TR - | - 1 | | | | | | | 1 1 | 1 | | | Clay | 45 | 87 | |
| 3 - Gluconite TR TR Pyrite 5 Carbonastungec 1 Diatoms TR - | - 1 | | | | | 11 | | | - | | 1 | Volcanic glass | 1 | 1 | |
| Pyrite 6 - Garbonats unspec 1 Diatoms TR - | - 1 | | | 1.0 | 1 | | 3 | 1 | 1 | 11 | 1 1 | Glauconita | TR | TR | |
| Carbonate unspec 1 Diatoms TR | | | | | | | | | - | | | Pyrite | 6 | 7 | |
| Diatoms TR - | - 1 | | | | | 1 | | | 1 | | | Carbonate unspec. | - | 1 | |
| | - 1 | | | | | 11 | | 1.5 | - | 1.1 | | Diatoms | TR | | |

SITE 471 HOLE COBE 29 COBED INTERVAL 266.0-275.5 m

| × | PHIC | | CHA | OSS | TER | | | | | | | |
|------------|----------------------|--------------|--------------|--------------|---------|---------|--------|----------------------|-------------------------|---------------------------|---------|--|
| TIME - ROC | BIOSTRATIGRA ZONE | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE | SEDIMENTARY STRUCTURES | SAMPLES | LITHOLOGIC DESCRIPTION |
| | | | В | В | | CC | | A A AI | - | | * | |
| | | | | | | 1 | 0.5 | 5 | | | | SILTY CLAYSTONE AND PORCELLANTE-PORCELANEOUS CLAYSTONE. Sitv claystone is gravito hilve (10V 4/2) and structureliss. Porcelaneous rocks are olive grav (5Y 3/2) to light olive grav (5Y 5/2), and burrowed mottled. Core brecclated by drilling. Core-Catcher only. |
| | | | | | | | 1.5 | | | | | |
| | | | | | | | | 1 | 1 | | 1 1 | SMEAR SLIDE SUMMARY |
| | | | 0.1 | | 1.1 | | - | 1 | 1 | | 1 1 | (D) |
| | | | | | | | 1.5 | | | | 1 1 | TEXTURE: |
| | | | (I I | 0.0 | 11 | | | 1 | | | 1 1 | Sand - |
| | | | h (| | | | 1 | | | | 1 1 | Silt 13 |
| | | | | | | 2 | 1 | | | | | Clay 87 COMPOSITION: |
| | | | | | | | - | | | | 1 1 | Quartz 7 |
| | | | | | | 1 | 1.2 | 1 | 1 | | 1 | Feldspar 3 |
| | | | | | | | 1.1 | 1 | | | | Mica TR |
| | | | | | | | - | 1 | | | | Heavy minerals TR |
| | | 1 | | | | | - | 1 | | | 1 | Clay 88 |
| | | | | | | | 1 3 | 1 | | | | Glauconite TR |
| | | 1 | | | | | - | 1 | | 1 | | Pyrite 2 |

| ¥ | PHIC | | F CHA | OSS | TER | | | | Π | | | | |
|-------------|--------------|--------------|--------------|--------------|---------|---------|--------|----------------------|----------|------------|---------|---|--|
| TIME - ROCI | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILE | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | STRUCTURES | SAMPLES | LITHOLOGIC DESCRIPTION | |
| | | | в | в | в | 1 | 0.5 | | | | + | PORCELLANITE-PORCELANEOUS CLAYSTORE AND SILTY CLAY. Porcelaneous rocks are olive gray (5Y 3/2) to light gray (5Y 5/2) and burrowed. Sithy clay is olive gray (5Y 3/2). Core composed of drilling fragments. No Core-Catcher. ORGANIC CARBON AND CARBONATE 1-70 % Organic Carbon – – % CaCOa 0 | |

| ~ | PHIC | 3 | CHA | OSS | TER | | | | | | |
|-------------|--------------|--------------|--------------|--------------|---------|---------|--------|----------------------|--|---------|--|
| TIME - ROCI | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY | SAMPLES | LITHOLOGIC DESCRIPTION |
| | | | в | в | в | 1 | 0.5 | | | + | PORCELANEOUS CLAYSTONE, olive gray (5Y 3/2), intensely burrowed, Most of core beecciated by drilling. No Core-Catcher. ORGANIC CARBON AND CARBONATE Sorganic Carbon – 141 Sorganic Carbon – 0 |

| 471 | | HOL | E | | C | ORE | 31 | COREC | INTE | R\ | AL/ | 285.0-294.5 m | | | _ |
|----------------------|---------------------|--|--------------|--|--|--|---|--|----------|-------------|---------|---|--|---|--|
| PHIC | | F | OSSI | L | | | | | T | | | | | | |
| BIOSTRATIGRA ZONE | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GLIT | RAPHIC | DRILLING | SEDIMENTARY | SAMPLES | LITHOLOG | GIC DE | SCRIPTION | |
| | | | B | 8 | 1 | 1.0- | | | | 1 | * | PORCELANEC CLAYEY LIMI gray (5Y 3/2) and (5Y 3/2), the (5Y 6/1) and b breccisted by d No Core-Catch ORGANIC CAI % Organic Carb % CaCO3 SMEAR SLIDE TEXTURE: Sard Sit Clay COMPOSITIO COMPOSITIO COMPOSITIO COMPOSITIO COMPOSITIO COMPOSITIO Composition Clay Mica Heavy minerals Clay Glauconite Pyrite Distoma | DUS CL and bur tureles: turrowe er, RBON kon SUMM | AYSTONE, SILTY CLAY, AND E. Porcelaneous rocks are olive rowed; silty class also olive gray to limestone is light olive gray d. Core consists of fragments AND CARBONATE 1-96 0.80 0 0 ARRY 1-55 (D) 7 3 TR 7 3 TR 7 4 1 5 5 5 7 7 3 TR 7 7 3 7 7 7 7 3 7 7 7 7 7 7 7 7 7 7 7 | |
| | | | | | | | | | | | | Opal-CT(?) | | 25 | |
| | BIOSTRATIGRAPHIC 12 | 471 SONE-VENDER SONE SONE SONE SONE SONE SONE SONE SONE | | 471 HOLE FOSSI BINO2 BIN | 471 HOLE FOSSIL FOSSIL FOSSIL STUDENTION STUDENT ST | 471 HOLE CONTRACTER NOTICE STREET OF | 471 HOLE CORE UNARADTER NOLICIAN Status Image: Status Status Image: Status </td <td>471 HOLE CORE 31 POSSIL BUDON BUDON BUDON B FOSSIL BUDON B NULL B SULL B NULL B SULL B SULL B</td> <td></td> <td></td> <td></td> <td>471 HOLE CORE 31 CORED INTERVAL FRANKLER GRAPHIC UNANCTER UNANCTER WINDOW WINDOW WINDOW WINDOW B 0.5 HINDOW HINDOW B 0.5 HINDOW HINDOW</td> <td>471 HOLE CORE 31 CORED INTERVAL 285.0-294.5 m FRANKACTER B <</td> <td>471 HOLE CORE 31 CORED INTERVAL 285.0-294.5 m FOSSIL (CRARACTER NUMBER (CRARACTER (CRARAC</td> <td>471 HOLE CORE 31 CORED INTERVAL 285.0–294.5 m POSSL VEX.00 POSSL VEX.00 POSSL VEX.00 Image: State of the s</td> | 471 HOLE CORE 31 POSSIL BUDON BUDON BUDON B FOSSIL BUDON B NULL B SULL B NULL B SULL B SULL B | | | | 471 HOLE CORE 31 CORED INTERVAL FRANKLER GRAPHIC UNANCTER UNANCTER WINDOW WINDOW WINDOW WINDOW B 0.5 HINDOW HINDOW B 0.5 HINDOW HINDOW | 471 HOLE CORE 31 CORED INTERVAL 285.0-294.5 m FRANKACTER B < | 471 HOLE CORE 31 CORED INTERVAL 285.0-294.5 m FOSSIL (CRARACTER NUMBER (CRARACTER (CRARAC | 471 HOLE CORE 31 CORED INTERVAL 285.0–294.5 m POSSL VEX.00 POSSL VEX.00 POSSL VEX.00 Image: State of the s |

| ~ | PHIC | - 0 | CHA | OSSI | TER | | | | | | | |
|------------|--------------|--------------|--------------|--------------|---------|---------|--------|----------------------|--|---|---|--|
| UNIT - ROC | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | LITHOLOGIC D | ESCRIPT | TION |
| | | | | в | в | 1 CC | 0.5 | | | SILTY CLAYSTONI burrowed, Core cons drilling breccle of san siliceous. | E, dark gr ists of sev ne litholo | eenish gray (5GY 4/1), eral intact pieces set in ogy, Claystone possibly |
| | | | | 1 | | 1 | | | | ORGANIC CARBON | AND C | ARBONATE |
| | | | | | | | | | | | 1-0 | 1-4 |
| | | | | | | | | | | % Organic Carbon | 0.86 | |
| | | | | | | | | | | % CaCO3 | 0 | 0 |
| | | | | | | | | | | SMEAR SLIDE SUI | MMARY | |
| | | | | | | | | | | | CC-18 | |
| | | 1.1 | | | 11 | 1 | | | | | (D) | |
| | | | | | | | | | | TEXTURE: | | |
| | | | | | | | | | | Sand | - | |
| | | | | | | | | | | Silt | 25 | |
| | | | | | | 1 | | | | Clay | 75 | |
| | | | | | | | | | | COMPOSITION: | | |
| | | | | | | | | | | Quartz | 15 | |
| | | | | | | | | | | Feldspar | 2 | |
| | | | | | | | | | | Mica | 1 | |
| | | | | | 11 | | | | | Clay | 70 | |
| | | | | | | | | | | Pyrite | 1 | |
| | | | | | | | | | | Carbonate unspec. | 2 | |
| | | | | | | | | | | Diatoms | 5 | |
| | | | 1.2 | | | | | | | Radiolarians | 2 | |





Abundant thin sandy layers

LITHOLOGIC DESCRIPTION

burrowed. Core-Catcher only.

SILTY CLAYSTONE, medium olive gray (5Y 4/2),

CORE 40 CORED INTERVAL 370.5-379.5 m

DRILLING DISTURBANCE SEDIMENTARY

뷶

SITE 471



ŝ FP

SITE 471

BIOSTRATIGF

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TIME - ROCK UNIT

MIDDLE MIOCENE

AM E

HOLE FOSSIL

5 5

NANNOFOSS

AM

METERS

0.5

1.0-

SECT

GRAPHIC

304
| NOTION Status Status Status Status Status Status Status Status Status Status Status Stat | DING | FOSSIL CHARACTER | | | | |
|---|----------------|--|---|--|--|--|
| BILTY CLAYSTONE, medium olive gray (5Y 4/2) bioturbased throughout with minor zones of lentitic bioturbased throughout with minor zones of lentits of lentitic bioturbased thr | UNIT - ROC | POINTINE ERIE CONTRACTOR CONTRACT | C DESCRIPT | TION | | |
| 4 Sanditione Volcanci glass 1 5(7) 2(7) Sanditione Glauconite TR TR TR TR Calcite - TR - TR - TR Calcite - TR - TR - TR - Calcite - TR - TR - | MIDDLE MIOCENE | CP 0.5 1 1.0 2 2 3 3 4 5 5 5 5 5 5 5 5 5 5 5 5 5 | NE, medium blandy-silty of medium b Sections 1 a Landy-silty Sections 1 a Core broken ON AND CA 1-128 - 0 UMMARY 1-4 (M) - 5 3 3 TR 15 5 3 3 TR 75 5 3 3 TR 75 1 2 | n olive gray of the second sec | y (5Y 4 s of lar (158 §/ lawstone ing bisc (D) - - - 2(5) 75 75 75 75 75 75 75 75 75 76 77 77 77 77 77 77 77 77 77 77 77 77 | //2), //2), //1) //2) //2) //2) //2) //2) //2) //2) |

FOSSIL TIME - ROCK UNIT SILS SILS SECTION GRAPHIC DRILLING DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES BIOSTRATIGP LITHOLOGIC DESCRIPTION NANNOFOS RP 1 4 SILTY CLAYSTONE, medium olive gray (5Y 4/2), burrowed, Nannofossils detectable in Core-Catcher. Thin sandy layer as marked. (X) MIDDLE MIOCENE 0.5 2/2 ORGANIC CARBON AND CARBONATE 1.32 1.0 % Organic Carbon % CaCO₃ BFPB Sandstone CC __4____ | | ō SMEAR SLIDE SUMMARY CC-9 (D) 30 TEXTURE: Sand Silt Clay COMPOSITION: Quartz Feldspar Mica 25 75 7 3 TR TR 70 TR 2 5 Mica Heavy minerals Cay Volcanic glass Pyrite Carbonate unspec. Foraminifers Calc, Nannofossils Lithic fragments TR 10 3

CORE 42 CORED INTERVAL 389.0-398.5 m

SITE 471 HOLE

| | APHIC | | CHA | OSS | CTER | | | | | | | | | | | |
|----------------|------------------------------|--------------|--------------|--------------|---------|---------|--------------|----------------------|----------|---------------------------|---------|--|--|---|--|---|
| UNIT | BIOSTRATIGR | FORAMINIFERS | NANNOFOSSILS | RADIOLAHIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGIC | DESCRIPT | ION | |
| | | | | | | 1 | 0.5 | 83 | | • | + | Sandstone | SILTY CLAYSTON burrowed throughos Scattered this sandy marked, SANDY IN I cm thick, medium cemented, and wavy intervals of CALCAI occur in Sections 1 contacts with non-c Com broken into diffe | E, medium it (faint le layers in 1 TERVALS bluish gra bedding (REOUS SI ind 4 as m ilcareous c | o olive gra nticular to Sections Susually y (5B 5/1 burrowed LTY CL) arked; gra laystone | edding). 1 and 5 as less than 1), calcite- 12), Two AVSTONE edational above and below. |
| | | | | | | 2 | | | | | | Institutosola. | ORGANIC CARBON % Organic Carbon % CaCO ₃ SMEAR SLIDE SUM | AND CA 1-61 0.70 0 MARY 2-15 | 4-66 | 5-12 |
| MIDDLE MIOCENE | Coccolithus miopelagicus (N) | | FM | | | 3 | real and and | | | | | | TEXTURE: Sand Clay COMPOSITION: Quartz Feldspar Mica Heavy minerals Clay Pyrite | (D) - 15 85 7 3 TR TR 70 2 | (D) - 13 87 5 3 TR 7R 87 2 | (M) 80 20 - 40 20 TR 11 - - |
| | | | в | | | 4 | red on Lone | | | | + | | Carbonate unspec. Calc, Nannofossils Lithic fragments | 2 15 2 | 1 TR 15 | 14 |
| | | | FM | B | | 5 | | | | +++ | • | Sandstone Sandstone Sandstone Sandstone | | | | |



SITE 471 HOLE CORE 45 CORED INTERVAL 418.0-427.5 m

| × | PHIC | | F | OSS RAC | TER | | | | | | | | | | |
|--------------------|---------------------|--------------|--------------|--------------|---------|---------|--------|----------------------|--|---------|---|---|--|--|---|
| TIME - ROC UNIT | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGIC D | ESCRIPT | ION | |
| DCENE | gicus (N) | | | | | 1 | 0.5 | | | • | Scattered thin sandy layers | SILTY CLAYSTONE burrowed throughout LAYERS scattered th Section 2 as marked, wavy bedding (burror (SY 6/1) CLAYEY L gradational contact w broken into drilling b | t, medium t, Thin, ca hroughour Sandy lay wed?}. Bu (MESTO) with overly biscults, | olive gra licite-cerr Section rers <1 c rrowed, i VE in Cor ring silty | w (5Y 4/2), iented SANDY 1 and most of m thick with light olive gray re-Catther has claystone. Core |
| WIDDLE MIC | Coccolithus miopels | В | AM | | | 2 | | | ▲ | + | Nannofossils Scattered thin sandy layers | ORGANIC CARBON % Organic Carbon % CaCO ₃ SMEAR SLIDE SUM TEXTURE: | AND CA 2-85 0 IMARY 1-20 (M) | 2-15 (D) | CC-15 (D) |
| | | | | в | | cc | | | | 1. | | Silt Clay COMPOSITION: Quartz Feldspar Mica Heavy minerals Clay Glauconite Pyrite Carbonate unspec. Calc. Nannofossila | 20 25 40 10 - 11 19 TR - 5 - | - 15 85 7 3 TR 70 - 2 3 15 | - - 1 16 - 2 80 - |

| , I | PHIC | | CHA | OSS | TER | | | | | | T | | | |
|--------------------|--------------|--------------|--------------|--------------|---------|---------|--------|----------------------|-------------------------|---------------------------|---------|-----------------------------------|--|--|
| TIME - ROCH | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE | SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGIC D | ESCRIPTION |
| | | | FM | | | 1 | 0.5 | | | •••• | | Scattered thin sandy layers | SILTY CLAYSTONE burrowed throughout (wavy bedding). Som Light olive gray (5Y f base of core is burrow laminations.Core brok | , medium olive gray (5Y 4/2), with scattered this sandy layers b burrows filled with fine sand, /2/2 CLAYEY LIMESTONE at ed but contains one interval of fine ten into drilling biscults. |
| | | | | | | | | | | ++ | | t | | |
| | | - | D | | - | 100 | | ليبي لمجمليين | <u>.</u> | | _ | | | |
| SITE | 471 | 1 | HOL | E | | c | ORE | 47 CORED | INTI | ERV | AL | 437.0-446.5 m | | |
| × | APHIC | | F | OSS | TER | | | | Π | | Τ | | | |
| TIME - ROC UNIT | BIOSTRATIGR | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DISTURBANCE | SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGIC DI | SCRIPTION |
| | | | | | | | | | | Î | | 1 | SILTY CLAYSTONE | , medium olive gray (5Y 4/2), |

| 2 | VPHIC | | CHA | OSS | TER | | | | Π | | | | | |
|----------------|----------------------|--------------|--------------|--------------|---------|---------|--------|----------------------|-------------|------------|---------|---|---|--|
| UNIT UNIT | BIOSTRATIGR/ ZONE | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DISTURBANCE | STRUCTURES | SAMPLES | LI | THOLOGIC DESCR | PTION |
| MIDDLE MIOCENE | | | B | | | 1 | 0.5 | | | | • | Settered thin standy layers layers | CLAYSTONE, med ed throughout; celici meley burrowed. Th meley burrowed. Th earnerhot gand occurs 5-10 cm. NIC CARBON AND 1-1 nic Carbon 0.6 03 0 R SLIDE SUMMARI 14 100 | um olive grav (ISY 4/2), recout interval in Section n, wavy layers of grav every 5–10 cm throughout ly <1 cm thick and spaced CARBONATE J9 1 7 3-455 (D) |
| | | | B | в | в | cc | - | | 11 | - | _ | Carbon Lithic f | ate unspec Iragments 25 | 3 |

| | WHIC | | CHA | OSSI | L TER | | | | | | |
|----------------|--------------|--------------|--------------|--------------|----------|---------|--------|----------------------|--|---------|--|
| LINO | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY | SAMPLES | LITHOLOGIC DESCRIPTION |
| MIDDLE MIDCENE | | | FM | | | 1 | | | | | SILTY CLAYSTONE, olive gray (5Y 3/2), burrowed throughout. Contains scattered, thin silty sandy layers. Some burrows filled with fine sand. No Core-Catcher. SMEAR SLIDE SUMMARY 1-8 (D) TEXTURE: Sand – Silt 14 Clay 88 COMPOSITION: Quartz 5 Feldspar 7 Mice TR Heavy minerals TR |



| | | | см | в | | 3 | - Translation | | | 111 | | v | | | |
|-------------|----------------------|--------------|--------------|--------------|---------|---------|---------------|----------------------|----------|-------------|---------|-----------------------------------|---|--|---|
| ITE | 471 01Hd | | HOL F | E OSS | IL | co | RE | 52 CORED | INT | ER | | 484,5-494.0 m | | | |
| TIME - ROCH | BIOSTRATIGRA ZONE | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY | SAMPLES | | LITHOLOGIC D | DESCRIP | TION |
| | | | | | | 1 | 0.5 | | | * | + | Abundanti thin sandy laytra | SILTY CLAYSTON burrowed throughou cenential sandy laye on spart. Sands are j claystone by burrow ORGANIC CARBON % Organic Carbon % CaCO ₃ | E, mediur rr. Contai ers (1–2 c partly hor ing. N AND C 2-28 – 0 | molive gray (5Y 4/2), ns abundant thin calcite- am thick) spaced - 5-10 mogenized with surrounding ARBONATE 2:137 0.55 0.1 |

CC

LITHOLOGIC DESCRIPTION

SILTY CLAYSTONE, medium olive gray (5Y 4/2),

intensely burrowed throughout; silty sand fills some burrows. Thin, calcite-cemented sandy layers (1-2

cm thick) occur every 5-10 cm throughout core. SILTY SANDSTONE near top Section 1 is laminated

1-121

0

near its base and fines upward to burrowed silty clay. Sandstone is medium gray and calcite-comented, Scattered sponge fragments also occur in this core.

ORGANIC CARBON AND CARBONATE

% Organic Carbon

% CaCO3

SITE 471

SITE 471

ATIGE

IOST

TIME - ROCK

MIDDLE MIOCENE

HOLE

FOSSIL CHARACTER

FORAMINIFER NAMNOFOSSIL RADIOLARIAN DIATOMS

2

3

cc

AW

Scattered thin sandy layers

Clay COMPOSITION:

Quartz

Pyrite

Feldspar Mica Heavy minerals

Clay Volcanic glass

Foraminifers

Carbonate unspec

Calc. Nannofossils

Lithic fragments

85

TR TR

1 10 -

20 TR 1

9

1 -

3 2

| SITI | E 471 | HC | DLE | | CORE | 5 | 3 CO | RED I | NTER | VAL | 494.0-503.5 m | | | SITE | 471 | н | OLE | | col | RE | 56 CO | RED IN | TERV | AL | 522.5-532.0 m | | |
|-------------|--|--------------|--|---|-------------------|-----------|-------------------|-----------|--|---------|-------------------------------------|--|--|---------------------|------------------------------|--------------|----------------------|---------|---------|--------|-------------------|--------------------|--|---------|----------------------------------|---|--|
| TIME - ROCK | BIOSTRATIGRAPHIC | FORAMINIFERS | FOSSIL IARACTEI SWVINVIO | R | METERS | | GRAPHI LITHOLO | IC IGY | DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGIC DESC | RIPTION | TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS | FOSSILS HADIOLARIANS | DIATOMS | SECTION | METERS | GRAPH LITHOLO | DRILLING | DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGI | DESCRIPTION |
| | | | | | 1 | | | | | + | | SILTY CLAYSTONE, m burrowed throughout. C sandy interbeds every ~5 partly mixed the two lift ORGANIC CARBON AN 1 % Organic Carbon - % CaCO ₃ 0 | edium olive gray (5Y 4/2), ontains thin calcite-comented -10 cm. Burrowing has logies. ID CARBONATE -60 | MIDDLE MIOCENE | Coccolithus miopelagicus (N) | F | PB | | 1 | 0.5 | | | 1 | | | CALCITE-CEMEN bluish gray (5B S) thin interbed of o at ~5 cm; gradatic Core-Catcher only | ITED SILTY SANDSTONE, medium 1) and parallel laminated. Contains ive gray (5Y 4/2) SILTY CLAYSTONE nal contacts. |
| | | | | | | 1 | | | ill. | | | | | | | | | | | | | | | | | | |
| | | | | | 2 | 1 | | | ill | | | | | SITE | 471 | н | DLE | | COF | RE 1 | 57 CO | RED IN | TERVA | AL | 532.0-541.5 m | | |
| DCENE | oelagicus (N) | | | | | | | | | | Abundant thin sandy interbeds | | | TIME - ROCK UNIT | BIOSTRATIGRAPHI | FORAMINIFERS | HADIOLARIANS | SWOLAND | SECTION | METERS | GRAPHI LITHOLO | DRILLLING DRILLING | DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGIC | DESCRIPTION |
| MIDDLE | Coccolithus mio | CI | M B | - | 3 4 5 00 | | | | | - | | | | NIOCENE | belagicus (7) (N) | | | | 2 | 1.0 | | | | | Abundant thin sandy layers | SILTY CLAYSTO Intensely bloturba this interfayers of comentad SILTY 3 hick and speed - parallel lamination layer; tops are bun claystone (Bouma daytone (Bouma daytone | NE, medium olive gray (SY 4/2), ad throughout. Contains abundant medium bluish gray (SB 5/1) calcite- ANDSTONE, Layers are ~ 1-4 cm 5-10 cm apart. Faint cross and s common in lower part of each owned and mixed with overhying C-E intervals). Several layers have a las marked). NN AND CARBONATE 3-98 0.55 1-2 |
| SIT | E 471 | но | DLE | | CORE | 5 | 4 co | RED I | NTER | VAL | 503.5-513.0 m | | | DLE | s mio | | | | 3 | Ŧ | | | | | | | |
| TIME - ROCK | micpellagicus (N) BIOSTRATIGRAPHIC ZONE | PORAMINIFERS | FOSSIL HARACTE SINOLUTION SINOLUTION B M B | R | L SECTION METERS | in finite | GRAPH LITHOLO | IC DGY | | SAMPLES | | LITHOLOGIC DESI SILTY CLAYSTONE, I burrowed with scattered Core-Catcher only. | CRIPTION Indium offer gray (SY 4/2), sponge fragments. | | Coccolithu | | | | 4 | | | | | + DG | | | |
| MIDDLE | Caccolithus | | | | | TILI | | | | | | | THE RECEIPTION | | | | | | 5 CC | | | | | | | | |



| PHIC | | CHA | OSS | TER | | | | | | | | | |
|---|-----------------------|------------------------------|--------------|---------|---------|--------|----------------------|----------|---------------------------|---------|--|--|---|
| UNIT UNIT BIOSTRATIGRA | ZONE | FORAMINIFERS NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY STRUCTURES | SAMPLES | LITHOLOGIC | DESCRIP | TION |
| IIDDLE MIOCENE Vitture microadanicue (N) | (N) snobeledout south | FM | | | 1 | 0.5 | | | ▲ | • | SILTY CLAYSTO calcine commented S medium oilive gray oilive gray (55 5/2, and internetly burr gray (56 5/11, qua Two lifts are parth this sandy has destroyed origi layers interval are 1 = 4 o have parallel famil 13 om thick layer is laminated near in ORGANIC CARBO | VE with all ILTY SAN (5Y 4/2) v 5Y 3/2) r wwed. Sanc tzo-feldsp r homoger nal sedime m thick, snd thistons and thistons an | Sundant thin interbads of IDSTONE. Claystone is with some light and dark nottles, non-calcereous istone is mediume buish athic and calcereous. Lised by burrowing that many contacts. Sandy sece ~ 5–10 cm. Frew faint microcross-laminations. SANDSTONE In Section 2 burrowed near its top. ARBOMATE |
| Core | 00000 | СМ | в | | 3 | | | | 449 | * | % Organic Carbon % CaCO3 SMEAR SLIDE SU TEXTURE: Sand Silt Clay COMPOSITION: Quartz Feldspar Mice Heavy minerals Glauconite Carbonets unspec. | | 2.109 0.43 0 |

| × | VPHIC | c | FO | RAC | L TER | | | | | | | | |
|--------------------|------------------------------|--------------|--------------|--------------|----------|---------|--------|----------------------|--|-----------------------|----------------------------------|--|--|
| TIME - ROC UNIT | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY | STRUCTURES SAMPLES | | LITHOLOGIC DE | ESCRIPTION |
| MIDDLE MIOCENE | Coccolithus miopelagicus (N) | F | M | в | | 1 | 1.0 | | | * * | Abundant thin sandy layers | SILTY CLAYSTONE ommind SILTY SAM medium olive gray (57 billing ray (58 6/1), intensive burrowing, 5 speed – 5–10 var Sectors, 50 CHARA SLIDE SUMM TEXTURE: Sand Clay CMAPOSITION: Quartz Clay COMPOSITION: Quartz Clay Glauconitie Pyrite Carbonite unprec. Foraminitres | with this interbeds of calcite- KDSTONE. Claystone is Y 4/20, andstore is medium Linhologies party mixed by Sendy intervals 1-4 cm thick, AND CARBONATE 241 - 0.5 MARY 184 (D) - 16 84 7 5 TR TR B4 TR B4 TR TR TR TR TR TR TR TR |



| 2 | VHHC | сн | FOS ARA | SIL | ER | | | | Π | | | | | | | |
|--------------------|--------------|--------------|--------------|---------|----------|---------|--------|----------------------|-------------------------|------------|------------------------|----------------------------------|--|--|--|--|
| TIME - ROC UNIT | BIOSTRATIGRI | FORAMINIFERS | RADIOLARIANS | DIATOMO | own trin | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE | STRUCTURES | SAMPLES | | LITHOLOGIC | ESCRIPT | ON | |
| MIDDLE MIDCENE | | В | | | | 3 | 0.5 | | | | •• • • • • | Abundant thin sandy layers | SILTY CLAYSTONI calcita-comented SIL is medium olive gray olive gray (V5 5/2-5 medium bluish gray) homogenized by inte layers have sharp, loo microcross-aminatio interval mixed with c common throughout layers as in previous; layers as in previous; Social Carbon % Organic Carbon % Organic Carbon % Organic Carbon % Composition; Clay Composition; Composition; Carbonate unspec. Calc. Nannofossite Lithic fragments | E with abu TY SAND (SY 5/4) m. (SS 5/4) (SS 5/4) | adant thin STONE C 1 with light + titles; sand wo lithol; titles; sand wo lith | interbeds of larystone and dark istone is solve and dark stone is solve and burrowed bedding uency of sandy e laminated 3-23 (D) 80 20 |



SITE 471 HOLE CORE 68 CORED INTERVAL 636.5-646.0 m

| ~ | PHIC | 8 | F | OSSI RAC | TER | | Τ | | | Γ | | Γ | | |
|-------------|--------------|--------------|--------------|--------------|---------|---|------------|--------|----------------------|----------|-------------|---------|---|--|
| TIME - ROCI | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY | SAMPLES | | LITHOLOGIC DESCRIPTION |
| | | | в | в | | Ŧ | <u> 20</u> | - | | H | 114 | F | - | |
| | | | | | | | | 0.5 | | | | | | SILTY CLAYSTONE, medium olive gray (5Y 4/2), mixed by burrowing with minor calcite-cemented medium bluish gray (5B 5/1) silty sandstone. |
| | | | | | | 1 | 1 | 10 | | 1 | | 1 | 1 | Core-Catcher only. |

| SITI | 471 | но | LE | C | ORE | 69 | COREC | D INTER | AVAL | 646.0655,5 m | | | | SITE | 471 | но | LE | | CORE | 70 | CORED | NTER | VAL | 655.5-665.0 m | | | |
|-------------|--------------------------|------------------------------|----------------------|---------|--------|-------------|----------------|--|---------------------------------------|----------------------------------|--|--|---|---------------------|--------------------------|---------------|--|---|----------------------------|----|------------------|--|---------|---|--|--|--|
| TIME - ROCK | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS NANNOFOSSILS | ARACTER SWOIDIARIANS | SECTION | METERS | GR. LITH | APHIC OLOGY | DRILLING DISTURBANCE SEDIMENTARY | STRUCTURES SAMPLES | | LITHOLOGIC | C DESCRIPTION | | TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS | FOSSIL ARACTE SNUINUNIONARI SNUINUNIONARI | R | METERS | GI | RAPHIC HOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGIC | DESCRIPTION | |
| | | В | В | | 2 | | | | · · · · · · · · · · · · · · · · · · · | Abundant thin sandy layers | SILTY CLAYSTO of calcita-cements is medium offer gr olive gray mettles burrowed through and gradel back m where it has been burrowing fana. as in previous core claystone. No Core-Catcher. ORGANIC CARBU % Organic Carbon % GaCO3 SMEAR SLIDE SI TEXTURE: Sand Silt Clay COMPOSITION: Ouertz Haavy minerals. Clay Glauconite Pyrite Carbonate unspec Lithic fragments | DNE with abundant thin is ed SILTY SANDSTONE (tray (GY 4/2) with light and (GY 3/2 - SY 6/2) and in hour, Sandstone is medium (GY 3/2 - SY 6/2) and in hour, Sandstone is medium mixed with overlying clay Frequency and thickness es. Scattered sponge fragm SON AND CARBONATE 3-97 1 0.69 0 UMMARY 1-56 (D) - - 26 74 13 5 TR TR TR TR 74 13 5 5 TR 74 13 6 | nterbeds Claystone d dark tensely n gray stone by of sandy beds sents throughout | MIDDLE CENOZOIC | | RP FP B | 8 | | 2 2 3 3 5 5 | | | ▲ | -14 | Common this sandy layers Alternating light (10% 6/2) and dark (5Y 4/2) olive gray claystone | SILTY CLAYSTOD commented SILTY S of line gray (19' 42) (19' 32-3-5' 52) in throughout. Claysto by thin sandy layer gray (19' 42) of the sandy layer gray (19') and occu- linterbed in claysto upper contact gray layers as marked. ORGANIC CARBO Soura D and Soura D and Soura D and Soura D and Soura D and Soura | E with thin interbeds of cd NDSTONE, claystone is m vith light and dark olive gr utiles and interwelv burrow na in Sections 3–5 have m and dark olive (15 4 42) o 4 cm. Each color change m is at thin (~1 cm, up to 5 4 m, Each color change m and sections of the section and the base section of the section and the section of the section and the section of the section and section of the section of the section of the section and section of the section of the section of the section and section of the section of the section of the section and section of the section of the section of the section and section of the section of the section of the section of the section and section of the | alcita- nedium ray wed oloor narked edium an, an, an, andy |



SITE 471

2-11 5.72

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80 80

20 20

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35 25

TR TR 7

TR

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5 5

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| | VHIC | | CHA | OS | SIL | ER | | | | | | | | | |
|--------------------|------|--------------|--------------|--------------|---------|----|---------|---------------------|----------------------|--|-----------------------|---|--|--|--|
| UNIT UNIT | ZONE | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATONA | | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY | STRUCTURES SAMPLES | | LITHOLOGIC DESCRIPTION | | |
| | | | | | | | 1 | 0.5 | | | | Microfaults and calcite- filled fractures | SILTY CLAYSTONE with abundant thin interbeds of calcite-cemented SILTY SANDSTONE. Claystone is alternating light and medium olive gray (5Y 5/2- 5Y 4/2) and internshy burrowed throughout. Sandstone is medium gray (N4) occurring as abundant thin (~1 cm) interbeds throughout. Thicker sandstone in Section 11 pointar and microcros-laminated and broken by microfaults and calcite-filled fractures. Thin interval of calcite-filled fractures also occurs in Section 21 calcite-filled fractures also poors in Section 21 calcite-filled fractures also poors in Section 21 calcite-filled fractures have sharp | | |
| OCENE | | | | | | | 2 | | | | < | Calcite-filled fractures | basalt contacts and gradational, burrowed tops. ORGANIC CARBON AND CARBONATE 1-74 2-30 % Organic Carbon - 0.60 % CaCO ₃ 12 0 SMEAR \$LIDE SUMMARY 1-35 (D) | | |
| LOWER OR MIDDLE MI | | | FP | | | | 3 | the second second | | | | Abundant thin sandy layers | TEXTURE: Send — Sit 25 Clay 75 COMPOSITION: Ouartz 10 Feldspar 5 Mica TR Heavy minerals 1 Clay 75 | | |
| | | | | | | | 4 | terrel to the local | | | | | Glauconits TR Prrite 2 Carbonate unspec, 2 Cale, Nannofossils TR Lithic fragments 5 | | |
| | | | | | | | | | 5 | | | | | | |
| | | | FP | в | | | CC | | | | | | | | |

CORE 74 CORED INTERVAL 693.5-703.0 m FOSSIL TIME - ROCK UNIT BIOSTRATIGRAPHIC BIOSTRATIGRAPHIC FORAMINFERS MANNOFOSSILS RADIOLANIANS BIATOME SECTION GRAPHIC DRILLING DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES LITHOLOGIC DESCRIPTION SILTY CLAYSTONE with abundant thin interbeds of calcite-commond SILTY SANDSTONE. Claystone is alternating light and madium olive grav (SY 5/2–5Y 4/2) and internety burrowed throughout. Sandstone is medium gray (N4) and occurs as thin (1–2 cm) interbeds every -5-10 cm with harp basel context and burrowed tops, Indistinct parallel laminations common in sandstones. 0.5 1.0 No Core-Catcher. ORGANIC CARBON AND CARBONATE 2-3 3-98 - 0.54 0 0 % Organic Carbon % CaCO₃ EM 2 -SMEAR SLIDE SUMMARY 5-45 (D) lagicus (N) MIDDLE MIOCENE TEXTURE: Sand Silt Clay COMPOSITION: 20 30 50 20 5 Abundant thin sandy layers Quartz Feldspar lithus Mica TR 2 49 TR 1 3 20 Heavy minerals Cocc Clay Glauconite Pyrite Carbonate unspec. Lithic fragments OG 1111 111

SITE 471 HOLE







TIME - ROCK UNIT

MIOCENE OR PLIOCENE(?)

HOLE





Piece Numbe cm 0-8 1A VV 1A 1 vv D,VP 10 1B 2 00 00 2 0 00 300 2 3 3 3A 0.00 3 D. 38 V ٨ 2 44 v v Jue Call ¥. 4 т ٨ 4B v 1 40 V ٨ 50ent 40 * Sedi т D,V,P 4E 3 Nego Cl 5B 5 6A ×* 6A **6**B 1 (To) 68 6E 6F 6G 6C ٨ 7A . V ٨ 7B V D 7 1A 100-7C 00 00 4 1B T 8 B 1 т 1 1C 1D 61 0.6 0000 т 1E v 6L 200 v 1F v * 10 ~~ 1G 150-CORE/SECTION 79/1 79/2 80/1

0000000 Thin Section Descriptions Diabase - Section 1, Piece 1E, 127-128 cm Texture: intersertal Phenocrysts: -5%, 0.05 mm Alteration: calcite (35%) as vein fill; brownish green clays (20%) replace groundmass between plagloclase crystals; some plagloclase (10%) 1.0 zeolitized(?) 1. Diabase - Section 2, Piece 2, 40-42 cm Texture: intersertal 5.00 Phenocrysts: -0.1-0.2 mm plagioclase; zeolites(?) (10%) replace plagioclase 5A Shipboard Data 1 · . 0. 5 . V 1 v MINOR ROCK TYPE - CLAYSTONE VV Macroscopic Descriptions ~ 4 Calcite-filled voids in Section 1, 48-49 cm 05 v ٨ ۷ ٧ 120 Thin Section Descriptions Altered Diabase - Section 1, Piece 4D, 55-58 cm Texture: intersertal Phenocrysts: -(× ٨ Texture: clastic (×, Composition: diabase fragments (60%), <15 mm across, angular to subangular; matrix (10%), zeolites and clays altered diabase fragments; white mica (TR) 11 💿 80/2 80/3

Drier

SITE 471, CORE 79, SECTIONS 1-2, 741.5-743.0 m MAJOR ROCK TYPE - ALTERED DIABASE

MINOR ROCK TYPES - CLAYSTONE, SULFIDES, METALLIFEROUS SEDIMENT AND CHERT (= "sediment" in Section 1; descriptions on preceding core form)

Macroscopic Description (diabase only)

Medium gray (N5), fine-grained, massive, aphyric diabase. Prismatic plagioclase crystals up to 2 mm; slight increase in grain size from top to base of Section 1. Badly altered with calcite filled veins in Piece 1A, Section 1, and Piece 2, Section 2.

Groundmass: plagioclase, 30%, 0.2-1.0 mm, euhedral, lath shaped; opaques,

Groundmass: plagioclase, 20%, 0.5-2 mm, euhedral, laths; opaques, 5%,

Alteration: calcite (30%) as vein fill; brownish green clays (20%) between

| ample | D | V (II) | V(1) | P | |
|--------------------------|------|--------|------|----|--|
| ection 2, Piece 18, 8 cm | 2.47 | 3.70 | - | 16 | |
| | | | | | |

SITE 471, CORE 80, SECTIONS 1-3, 750.5-755.0 m

MAJOR ROCK TYPE - ALTERED DIABASE, DIABASE BRECCIA

Altered Diabase (Section 1, 0-100 cm; Section 3, 0-45 cm [with breccia] and 60-150 cm) - medium-grained, aphyric and badly altered to greenish clays.

Diabase Breccia (Section 1, 100-150 cm; Section 2, 0-150 cm; Section 3, 0-60 cm [with altered diabase]) - angular and subangular fragments (<5 cm) of diabase and minor sediment (clayey chert, set in an aphanitic, dark green clayey matrix. Fragments are fine- to medium-grained with plagioclase crystals up to 2 mm. Some fragments almost completely altered to clays.

Claystone (Section 1, 0-8 cm) - light olive gray (5Y 5/2), laminated, with two dark red (5Y 2/6) lenses or layers -0.5 cm thick

Groundmass: plagioclase, 45%, 0.5-3 mm, subhedral; clinopyroxene, 15%, 0.5-3 mm, anhedral; opaques, 7%, 0.2-1 mm

Alteration: brownish green clays (30%) between plagloclase laths; calcite (3%)

Diabase Breccia - Section 1, Piece 7D, 105-109 cm

Alteration: greenish clays (10-20%) form matrix and replace groundmass in

Diabase Breccia - Section 2, Piece 6G, 102-105 cm (diabase fragments) Texture: intersertal (fragments only) Phenocrysts: -

Groundmass : plagioclase, 45%, 1-2 mm, subhedral, laths; clinopyroxene, 5%, 0.5-1 mm, subhedral; opaques, 7%, 0.2-1 mm, anhedral, irregular Alteration: brownish green clays in groundmass; calcite (10%) in veins; zeolite(?) (3%) replacing plagioclase; white mica (TR)

Diabase Breccia - Section 3, Piece 3A, 21-24 cm

- Texture: clastic, intersertal (fragments)
- Composition: aphanitic green clay matrix (41%) altered diabase fragments (30%), <10 mm across, subangular; plagioclase (20%), 1-3 mm, subhedral laths, in groundmass of fragments and in matrix; opaques (2%) in fragments; clinopyroxene (TR), in fragments
- Alteration: green clays (41%) in matrix and in groundmass of diabase

fragments; zeolites(?) (7%) replaces plagioclase in veins Altered Diabase - Section 3, Piece 9, 133-135 cm

Texture: intersertal

Phenocrysts: -

Groundmass: plagioclase, 45%, 1-3 mm, subhedral, some crystals are zoned; opaques, 5%; clinopyroxene (TR), 0.2-0.5 mm, anhedral Alteration: brownish green clays (46%) and zeolites(?) (4%) replacing groundmass

Chinhaned Data

| D | V (ii) | V (i) | P |
|------|-----------|-----------------------|--------------------------------|
| 2.47 | 3,60 | - | 16 |
| | D 2,47 | D V (II) 2.47 3.60 | D V (ii) V (i.) 2.47 3.60 - |





SITE 471, CORE 81, SECTIONS 1-4, 760.0-766.0 m MAJOR ROCK TYPE - ALTERED DIABASE

MINOR ROCK TYPE - DIABASE BRECCIA

Macroscopic Descriptions

Altered Diabase – fine- to medium-prained and aphyric. Grain size decreases slightly, 0–83 cm in Section 1; slight grain size increase from top to base, Section 3. Common alteration to green clays with plagioclase laths (up to 3 mm long) often set in green clay matrix. Clay alteration occurs as irregular patches or along fractures. Apparent layering (<0.5–1 mm thick) in Piece 7B, Section 3, caused by alternations of altered (green clay) and less altered diabase. Calcite-filled veins and voids scattered throughout. Pyrite crystals at 22–26 cm, Section 3.

Diabase Breccia (Pieces 5–6, Section 1; Pieces 3–5 and 9–11, Section 4) – angular fragments of altered diabase set in green clay matrix.

Thin Section Descriptions

Altered Diabate - Section 2, Piece 7, 73-76 cm Texture: intersertal

Phenocrysts: -

Groundmass: greenish clays, 32%; zeolltes 20%; clinopyroxene, 30%, 0.5–3 mm, subhedral; plagioclase, 15%, 0.5–3 mm; opaques, 2%; sphene, 1% Alteration: greenish clays (32%) and zeolites(?) (20%) replace groundmass

Altered Diabase - Section 4, Piece 2, 20-24 cm

Texture: intersertal to subophitic

Phenocrysts: -

Groundmass: clinopyroxene, 35%, 0.5–3.5 mm, anhedral; plagioclase, 15%, 0.5–3 mm, subhedral; opaques, 5%; sphene, TR; green clays, 30%; zeolite(?), 5%

Alteration: green clays (30%) and zeolite(?) (5%) replace groundmass Shiphoard Data

| Sample | D | VO | V (i) | P | |
|---------------------------|------|------|-------|----|--|
| Section 2. Piece 6, 53 cm | 2.55 | 4.56 | | 12 | |

SITE 471, CORE 82, SECTIONS 1--2, 769.0-771.5 m

MAJOR ROCK TYPE - ALTERED DIABASE

MINOR ROCK TYPES - DIABASE (BASALTIC?) BRECCIA, CLAYSTONE

Macroscopic Descriptions

Altered Diabase — fine- to medium-grained, massive and aphyric. Grain size decrease, 0-55 cm, Section 1 then increases from 60 – 150 cm. Extensive green clay alteration of groundmass; plagioclase laths (up to 2 mm long) set in clay matrix. Green clay also fills some fractures. Slickenside on Piece 3, Section 2.

Diabase (Basaltic?) Breccia (Section 1, Pieces 6A-B) - angular fragments of altered diabase (basalt?) set in green clay matrix.

Claystone (Section 1, Piece 5) - grayish brown (5YR 3/2), massive.

Thin Section Descriptions

Diabase (Basaltic?) Breccia - Section 1, Piece 6A, 65-68 cm

Texture: clastic, hyaloophitic (fragments)

Composition: altered cryptocrystalline fragments (originally basaltic glass[7]), 60%, <10 mm across, angular to subangular; green clay, 11%; plagioclase, 10%, 0.02-0.1 mm, prismatic; clinoproxene, 10%, 0.05-0.2 mm, prismatic; opaques, 4%,

0.05-0.1 mm; zeolites(?), 5% Alteration: green clay marix (11%); zeolite (?) (5%) replacing plagloclase and

groundmass

Altered Diabase - Section 1, Piece 9, 124-127 cm

Texture: intersertal

Phenocrysts: -

Groundmass: plagioclase, 30%, 0.5–3 mm, subhedral; clinopyroxene, 30%, 0.5–3.5 mm, subhedral; clays, 30%; zoolites, 10%; opaques, 5%; sphene, TR Alteration: green clays (30%) and zeolites (10%) replace groundmass

Shipboard Data

 Sample
 D
 V (1)
 P

 Section 1, Piece 8D, 111 cm
 2.60
 4.00
 –
 12



SITE 471, CORE 83, SECTIONS 1-3, 778.0-782.1 m

MAJOR ROCK TYPE - DIABASE BRECCIA

Macroscopic Description

Diabase Breccia — angular to subangular fragments of altered, vesicular, fineto medium-grained diabase set in dark gray (N3), aphanitic clayey matrix. Most fragments are <5 cm across. In Section 2, below 110 cm and in Section 3 the long dimension of fragments is aligned subparallal to core axis (i.e. vertical), Plagioclase laths in diabase fragments are up to 3 mm long.

Thin Section Descriptions

Diabase Breccia - Section 2, Piece 5D, 132-133 cm Texture: clastic, intersertal (fragments)

Composition: plagioclase crystal fragments (angular to subangular, 0.1-0.7 mm) form about 40% of specimen; unbroken plagioclase crystals,

20%, 0.5-3.5 mm, subhedral; clinopyroxene, 3%, 0.5-1 mm, anhedral; clays, 32%; opaques, 5%, 0.2-0.5 mm; sphene, TR, calcite TR

Alteration: greenish clays (32%) replace groundmass; calcite (TR)

Diabase Breccia - Section 3, Piece 3B, 29-33 cm

Texture: clastic Composition: plagioclase, 50%, 0.1–1.0 mm, subangular, fragmented crystals; clinopyroxene, TR, 0.2–1.0 mm, subangular, fragmented;

opaques, 4%, 0.5–1.0 mm; sphene, 2%; matrix (clays 14%) Vesicles: 15%, 0.5–2 mm, irregular

Alteration: clays (10%) replace groundmass; calcite (3%)

Shipboard Data

 Sample
 D
 V (s)
 V (s)
 P

 Section 3, Piece 18, 11 cm
 2.31
 3.19
 –
 25

SITE 471, CORE 84, SECTIONS 1-2, 787.0-789.5 m

MAJOR ROCK TYPE - ALTERED DIABASE

Macroscopic Description

Altered Diabase – fine- to medium-grained, massive, and aphyric. Scattered dark green to black prismatic pyroxene crystals and plagioclase laths up to 5 mm lone

Thin Section Descriptions

Altered Diabase - Section 1, Piece 6F, 102-106 cm

Texture: intersertal

Phenocrysts: -

Groundmass: plagioclase, 32%, subhedral; clinopyroxene, 15%, 0.5–2 mm, subhedral to anhedral; opaques, 15%, 0.5–1.5 mm, anhedral; quartz, 3%, 0.1–0.2 mm, anhedral; sphene, 1%, 0.05 mm, anhedral; amphibole, TR, 0.05–0.1 mm, acicular prismatic, coloriess: biotite: TR, 0.1–0.2 mm.

platy, brown to dark green

Alteration : clays (30%) and calcite (4%) replace groundmass

Shipboard Data

| Smpboard Data | | | | |
|----------------------------|------|-------|-------|----|
| Sample | D | V (1) | V (1) | P |
| Section 1, Piece 6E, 94 cm | 2.67 | 4.41 | - | 11 |

SITE 471, CORE 85, SECTIONS 1-2, 796.0-798.6 m

MAJOR ROCK TYPE - ALTERED DIABASE

Macroscopic Description

Altered Diabase – massive, fine- to medium-grained and aphyric. Plagioclase laths and dark green to black prismatic pyroxene crystals up to 5 mm long scattered throughout.

Thin Section Description

Altered Diabase - Section 2, Piece 1A, 0-4 cm Texture: subophitic to intersertal

Phenocrysts: -

Groundmass: plagioclase, 40%, 1-4 mm, euhedral to subhedral; clinopyroxene, 35%, 1-5 mm, anhedral; opaques, 4%, 0.2-1 mm, anhedral; alkali feldspar, 3%, 0.2-1 mm, anhedral; quartz, 1%, 0.2-0.5 mm, anhedral; amphibole, rR, prismatic,

colorless; biotite, TR, 0.1 mm, platy, brown

Alteration: clays (15%) and calcite (2%) replace groundmass

Shipboard Data

Sample D V (1) V (1) P Section 1, Piece 4C, 140 cm 2.78 4.97 - 3





SITE 471, CORE 88, SECTIONS 1-4, 814.0-819.8 m

Altered Diabase - massive, fine- to medium-grained, aphyric, and holocrystalline. Many calcite- and clay-filled fractures and yeins as marked. Large fractured/brecciated zone between 44-73 cm, Section 1, contains many angular fragments of diabase cemented together with calcite. Most veins and fractures are less than 1 cm wide and are bordered by a zone of alteration. Section 4 is a drilling breccia consisting

0.5-3.5 mm, anhedral; biotite, 7%, 0.1-0.5 mm, platy, brown to dark green; opaques, 4%, 0.05-0.01 mm; amphibole, TR, 0.02-0.1 mm,

0.5-3.5 mm, anhedral; biotite, 7%, 0.1-0.5 mm, platy, brown to dark green; opaques, 4%, 0.05-0.2 mm; amphibole, TR, 0.02-0.1 mm















(a) 3.












































