### Shipboard Scientific Party<sup>2</sup>

## **HOLE 490**

Date occupied: 8 April 1979

Date departed: 13 April 1979

Time on hole: 115 hours

Position: 16°09.56'N; 99°03.39'W

Water depth (sea level; corrected m, echo-sounding): 1761

Water depth (rig floor; corrected m, echo-sounding): 1771

Bottom felt (m, drill pipe): 1777

Penetration (m): 588.5

Number of cores: 64

Total length of cored section (m): 586.5

Total core recovered (m): 341.4

Core recovery (%): 58

#### Oldest sediment cored:

Depth sub-bottom (m): 588.5 Nature: Muddy siltstone Age: Late Miocene

Principal results: Drilling at Site 490 penetrated 588.5 meters of mud, silt, and muddy siltstone ranging in age from Quaternary to late Miocene (Table 1). Figure 1 shows properties of these sediments. Fractures and slickensides indicate normal faulting during the early Pliocene. The strike of these faults was roughly perpendicular to the trend axis. We found no evidence of compression stress.

Pre-middle Pliocene sediments were deposited below the CCD (carbonate compensation depth). Subsequent rates of uplift are difficult to establish in detail because of poor microfossil assemblages but are roughly 200 m/m.y. Similarly, sedimentation rates are poorly known in detail but irregularly average about 150 m/m.y.

Methane is common and ethane ranges from below limits of detectability to 0.18% near the bottom of the hole. Gas-releasing ice inclusions and frozen sediments may indicate the presence of gas hydrates, but solution in seawater could also account for amounts of gases released.

Oldest sediments sampled in this hole could have been deposited on top of accretionary zone sediments as part of the slope apron or, alternatively, deposited off the edge of continental crust between continental schists of Site 489 and the accretionary zone.

## **BACKGROUND AND OBJECTIVES**

A major objective of the Leg 66 program was the sampling of sediments deposited during late presubduction and early synsubduction times in order to obtain information about the onset and beginnings of subduction. As noted in the Site 489 site chapter, the Leg 66 transit is unique in that upper parts of the presubduction and synsubduction sequences, relatively simple structurally and accessible beneath a relatively thin slope apron, are evident in seismic data. At Site 489 and 490 (Figs. 2 and 3) we focused on late presubduction and early synsubduction, 489 sampling sediments deposited on the seaward edge of the continental crustal block and 490 sampling those deposited in an area thought to lie immediately seaward of the continental crust edge.

Several geological models of Site 490 are consistent with the seismic reflection data. First, the principal strong reflector, at about 0.35 s sub-bottom, could be gneissic basement, as in Hole 489. The irregular surface of the reflector resembles that of gneissic basement farther shoreward. Deeper reflectors may have the same origin as those within the gneissic basement. Fault breccia along landward-dipping shear zones could account for the landward-dipping reflectors beneath the principal reflector. The seismic refraction velocity below the principal reflector is about 3.0 km/s, the same as the velocity of uppermost basement at Site 489. Although the lack of higher refraction velocities at greater depth argues against a basement interpretation, fractured gneissic basement would be then consistent with the observed velocities.

A second possibility is that rocks below the principal reflector consist of slope sediments deposited in the time interval between separation and removal of continental crust to the southwest and the beginning of active subduction—that is, late presubduction-facies. Seismic reflection data show these rocks to be somewhat deformed today, but deformation could have occurred during early stages of subduction.

The third possibility, and the one most consistent with seismic reflection data, is that rocks beneath the principal sub-bottom reflector belong to the early accretionary zone suite—that is, early synsubduction facies.

<sup>&</sup>lt;sup>1</sup> Initial Reports of the Deep Sea Drilling Project, Volume 66. <sup>2</sup> Joel, S. Watkins (Co-Chief Scientist), Gulf Research and Development Company,

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Table 1. Coring summary, Hole 490.

	Cored Interval	- ·	Reco	vered
Core	below Bottom (m)	Cored (m)	(m)	(%)
1	0.0-7.0	7.0	6.06	64
2	9.0-18.5	9.5	8.83	93
3	18.5-28.0	9.5	9.63	101
4	28.0-37.5	9.5	7.22	76
5	37.5-47.0	9.5	9.35	98
6	47.0-56.5	9.5	7.70	81
7	56.5-66.0	9.5	4.39	46
8	66.0-75.5	9.5	6.05	64
10	/5.5-85.0	9.5	3.20	34
10	85.0-94.5	9.5	6.98	13
17	94.5-104.0	9.5	1.20	13
13	113 5-123 0	9.5	2.05	20
14	123 0-132 5	9.5	0.30	0
15	132 5-142 0	9.5	6.47	68
16	142 0-151 5	9.5	5 35	56
17	151.5-161.0	9.5	2.92	31
18	161.0-170.5	9.5	8.49	89
19	170.5-180.0	9.5	0.90	9
20	180.0-189.5	9.5	9.43	99
21	189.5-199.0	9.5	8.27	87
22	199.0-208.5	9.5	8.60	91
23	208.5-218.0	9.5	7.10	75
24	218.0-227.5	9.5	7.52	79
25	227.5-237.0	9.5	9.17	97
26	237.0-246.5	9.5	6.20	65
27	246.5-256.0	9.5	4.92	52
28	256.0-265.5	9.5	9.73	102
29	265.5-275.0	9.5	6.04	64
30	275.0-284.5	9.5	3.46	36
31	284.5-294.0	9.5	2.08	22
32	294.0-299.0	5.0	1.29	83
34	303 5-313 0	4.5	1.20	10
35	313 0-322 5	9.5	7 20	49
36	322 5-332 0	9.5	9.21	97
37	332.0-341.5	9.5	9.43	99
38	341.5-351.0	9.5	8.70	92
39	351.0-360.5	9.5	9.15	96
40	360.5-370.0	9.5	8.73	92
41	370.0-379.5	9.5	2.19	23
42	379.5-389.0	9.5	7.30	77
43	389.0-398.5	9.5	4.27	45
44	398.5-408.0	9.5	3.70	39
45	408.0-417.5	9.5	1.60	17
40	417.5-422.5	5.0	0.72	8
47	422.5-427.0	4.5	1.98	21
40	427.0-430.5	9.5	1.01	17
50	430.3-440.0	9.5	1.22	12
51	455 5-465 0	9.5	1.65	17
52	465 0-474 5	9.5	1.05	12
53	474.5-484.0	9.5	1.30	14
54	484.0-493.5	9.5	6.03	63
55	493.5-503.0	9.5	6.05	64
56	503.0-512.5	9.5	8.14	86
57	512.5-522.0	9.5	9.16	96
58	522.0-531.5	9.5	9.02	95
59	531.5-541.0	9.5	7.25	76
60	541.0-550.5	9.5	5.35	56
61	550.5-560.0	9.5	5.21	55
62	560.0-569.5	9.5	5.97	63
63	569.5-579.0	9.5	3.01	32
64	579.0-588.5	9.5	1.88	20

Velocities are consistent with those expected in moderately indurated sediment, and weak landward-dipping reflectors resemble those downslope in younger parts of the accretionary zone.

Discrimination among the three models is fundamental to an understanding of the evolution of the subduction system. Samples from Site 490, together with samples from adjacent holes, were expected to contribute to the understanding of the early evolution of the subduction mechanism in this region.

### **OPERATIONS**

Site 490 is located 13 km south-southwest of Site 489. Transit time between Site 489 and beacon launch at 490 was one hour.

Operations proceeded routinely, and the seafloor was found with a punch core at 1777 meters, 22 meters deeper than the PDR reading. A mudline core test of the pressure core barrel was conducted before continuous coring resumed. Good hole conditions and generally good core recovery persisted throughout the hole with the exception of an interval of fractured siltstones and loose sand strata from about 405 to 475 meters below seafloor (BSF). Unstable hole conditions were again encountered beginning at about 570 meters. Despite mud flushes to clean the hole, excessive fill caused the drill string to stick temporarily following the retrieval of Core 490-64 from 588.5 meters BSF. The hole was cleaned out, with some difficulty, to total depth.

A "wiper trip" was made to prepare the hole for logging. Considerable fill was again encountered at the bottom of the hole. Torquing and plugging of the drill string prevented cleaning the hole below 580 meters. A go-devil was then pumped down the pipe to actuate the hydraulic bit release. The release failed to shift after repeated pressuring up with the mud-pump. The godevil was retrieved with the sandline and a second godevil was pumped down. Again the bit failed to release and the overshot was lowered to retrieve the go-devil. The go-devil stuck at the bit and a second wireline was required to work it free after the safety release pin in the overshot was sheared.

Plans for logging were abandoned, and the hole was filled to about 280 meters BSF with weighted drilling mud. The bit was then pulled to the top of the mud fill, and cement slurry was emplaced up to about 120 meters BSF before the drill string was recovered. The *Challenger* departed for Site 491 at 1400 hours, April 13.

## LITHOLOGIC SUMMARY

Drilling at Site 490 penetrated 588.5 meters into argillaceous sediments of upper Miocene to Quaternary age. Sedimentary and structural attributes suggest division of sediments into three units (Fig. 4, back pocket).

Unit 1, Quaternary (0-~142 m sub-bottom), consists of muddy silt to mud, for the most part olive gray in color, becoming olive green in the uppermost 20 meters and olive black between 94.5 and 114 meters. Although predominantly homogeneous, subtle color changes sug-



Figure 1. Summary of age, nannofossil and radiolarian zones (1 = Pterocanium prismatium, 2 = Spongaster pentas, and 3 = Stichocorys peregrina, B = barren), magnetic polarity zones (black = normal, white = reversed), lithology, structures, porosity, organic geochemistry ( $C_2/C_1$  and  $C_{3-5}$ ), age-depth relationships for recovered sediments. Radiolarian and nannofossil boundaries based on Berggren and Van Couvering [1974] and paleomagnetic ages based on Ryan et al.

gest bedding in several places. Bioturbation is only discernible in a small area of Core 12. From about the 85 to 107 meter levels the mud is diatomaceous. Occasional minute white spots of sponge remains occur throughout the unit, and scattered blebs of fine sand (2-8 mm in diameter) occur below the 56-meter level.

[1974].)

Two bored limestone nodules, each measuring  $8 \times 2$  cm, occur in the topmost 10 meters (Sample 490-1-4, 80-90 cm and 490-1-1, 20-30 cm). In the shallower sample, the borings are concentrated in the upper half of the nodule. Carbonate elsewhere in the unit is limited to occasional diffuse olive brown spots and indurated concretions, never more than a few centimeters in diameter.

Glauconite is locally important in this unit, occurring both in scattered concentrations and as a common trace component in the upper 60 meters of mud. A 10-cm graded bed of glauconitic muddy sand at the 47-meter level has a sharp, irregular (possibly scoured) base and is probably redeposited.

Unit 2, Pliocene and lowermost Quaternary (~142-~399 m), comprises some 250 meters of olive gray muddy siltstone, for the most part bioturbated. Only small variations in composition and grain size are evident; these include one mudstone horizon (Core 21), one siltstone horizon (Core 35), and a siliceous component in the muddy siltstone in Core 23.

The essential uniformity of this unit is broken in places by minor thin interbeds and lenses of ash and limestone. Millimeter-scale laminations and thin beds of silt and fine sand, in some cases graded, are relatively abundant in several horizons. The small calcareous areas (both diffuse and relatively well indurated) and



Figure 2. Site location and index. A-A' shows location of Figure 3.



Figure 3. Portion of multichannel seismic Line OM-7N with hole locations 489 and 490.

occasional millimeter-scale fine sand blebs observed in Unit 1 also occur in Unit 2. Glauconite is a common trace of the sediment, with one major concentration occurring between  $\sim 260$  and 280 meters. Pyrite nodules up to several millimeters across are common through much of the unit. A 30-cm mud chip breccia occurs at 231 meters (Sample 490-25-4, 0-32 cm); plant material is concentrated in the muddy silt immediately underlying and overlying this unit.

Ash and vitric beds occur in several places. At 145.5 meters (Sample 490-16-3, 50 cm) horizontal lamination is pinpointed in a 25-cm vitric silt by light to dark gray color changes. At 166.5 meters (Sample 490-18-4, 50-85 cm) a 35-cm ash bed, with similar horizontal lamination, is coarser grained in the basal 4 cm. These ash beds were very cold when first removed form the core liner. In Core 18, thermister probe readings varied between  $-0.8^{\circ}$  and  $-2^{\circ}$ C. Furthermore, an irregular 3-cm fragment of bubbling ice was recovered in Core 15. The implications of these observations are discussed in the Organic Geochemistry section. Several thinner ash and vitric silt beds occur in the lower half of Unit 2.

Limestone and chalk occur in thin lenses and discrete beds, between 6 and 40 cm thick. The sediment varies in color from pale greenish gray to light olive gray to dusky yellow green and contains no clear biogenic component. Several beds contain irregular microveins, and the thickest bed (Sections 490-29-4, and 490-29-5) is internally brecciated.

Bedding, which in most places is easily discernible in Unit 2, infrequently deviates from the horizontal. Oblique open fractures become increasingly common toward the base of the unit; many of these are thought to be of tectonic origin, especially where slickensiding is observed. In Unit 3, bedding is inclined throughout, and slickensided fractures are more or less ubiquitous. For convenience, the boundary between Units 2 and 3 has been placed above the first of two breccia horizons (Core 44). It might equally well be placed above the first appearance of pervasive slickensided fracture planes (Core 42).

Unit 3 (~399-588.5 m) differs little from Unit 2 in overall sedimentary characteristics (Fig. 4). Thin chalk and ash beds are present, although glauconite is largely missing. However, the muddy siltstone of Unit 3 is well indurated throughout and characterized by abundant fracture planes varying in orientation from zero to near vertical. Slickensides, often showing dip-slip movements, are common on the fracture surfaces. Bedding, often demarcated by concentrations of flattened burrows, dips at shallow to moderately steep angles.

Breccia units at the top of the unit consist of a mixture of loose, muddy siltstone chips and some medium sand. The granule- and pebble-sized clasts of muddy siltstone are highly indurated, angular, and up to 5 cm across. The thicker breccia indicated in the generalized column was observed in a zone of poor recovery (Cores 45 and 46) and may or may not occupy the entire zone.

### Interpretation

The central question arising from Site 490 data is whether to interpret the deformed sediments of Unit 3 as accreted trench or ocean floor deposits below a relatively undeformed lower slope sequence (Units 1 and 2) or to consider the whole section a slope sequence progressively more deformed at the base. Given that trench deposits are not invariably sandy, especially during periods not associated with lower stands of sea level, the fine-grained sediments of Unit 3 give little clue to their paleogeographic setting, other than that they were probably deposited below the CCD (see Biostratigraphy section). They could be terrigenous ocean floor sediments, fine-grained trench fill, or deposits of the lower portion of the lower slope.

The somewhat abrupt appearance of intense fracturing might indicate a structural break around 380 meters sub-bottom (around Core 42), arguing for an accretionary origin for Unit 3. On the other hand, dip-slip motions indicated by slickensides are not what would be expected in a compressional (i.e., underthrusting and offscraping) regime. Dip-slip motions have been recorded on fractures and microfaults in lower slope sediments from the Japan (Leg 57, Hole 440B, e.g., Core 45) and Marianas (Leg 60, Site 459) forearcs. The origin of this deformation is further discussed in the Paleomagnetism and Summary and Conclusions sections.

On balance, the whole section recovered from Hole 490 is best interpreted as a lower slope sequence of upper Miocene through Quaternary age.

### **Drilling Artifacts**

Horizontal laminations, 2 to 4 mm thick on average, are common in the better-lithified cores retrieved so far during Leg 66. Their extreme regularity (2-4 cm spacing) engendered debate among the shipboard sedimentologists as to their origin: sedimentary or mechanical.

### BIOSTRATIGRAPHY

At Site 490 we penetrated a Quaternary to upper Miocene section of hemipelagic middle-slope sediments. Carbonate dissolution in the lower half of the hole, reworking of upslope Miocene to Quaternary microfossils, and microfossil dilution by terrigenous sediments hamper biostratigraphic dating.

Figure 1 presents the biostratigraphic summary of Site 490. The miscorrelation between the nannofossil NN11 Zone and the radiolarian *Stichocorys peregrina* Zone represents coccolith reworking. In view of the fact that calcareous nannoplankton preservation is very poor in the lower part of Hole 490 and that nannoplankton are absent at the base of the hole, it is possible that all preserved nannoplankton below Core 34 (310 m) were redeposited and buried before dissolving. Pliocene nannofossils, which accumulated under ordinary pelagic conditions, are mostly dissolved as well. Reworked Miocene and Pliocene radiolarians occur throughout the Pliocene and Quaternary sections of Hole 490.

### **Calcareous Nannoplankton**

The assemblages of coccoliths and other nannofossils found in the sediments of Hole 490 are assigned to Quaternary, Pliocene, and upper Miocene. No nannofossils older than Miocene were found, not even reworked ones.

## Quaternary Nannoplankton (Sections 490-1,CC-490-15,CC)

With the exception of Section 490-1, CC only moderately preserved coccoliths are present. In 7, CC there is the lowest occurrence of *Gephyrocapsa oceanica* with the central bridge preserved. The last occurrence of *Emiliania annula (Pseudoemiliania lacunosa)* was found in 10, CC. The last discoasters (*Discoaster brouweri*) occur in 16, CC, marking the Pliocene/Pleistocene boundary at about the top of the Olduvai (1.6 m.y. ago). The following datum levels were observed: Core 7: NN20 Zone bottom-0.6 m.y.; Core 15: NN19 Zone bottom-1.6 m.y.

## Pliocene Nannoplankton (Sections 490-16,CC-490-27,CC)

The Discoaster brouweri Zone NN18 extends from Sections 490-16 to 490-25, CC. The D. pentaradiatus Zone NN17, a comparatively narrow nannoplankton zone, is found in 26, CC and 27, CC. The following datum levels occur: Section 490-16, CC: NN18 top-1.6 m.y.; Sections 490-25 to 26, CC: NN17 top-2.5 m.y.; Section 490-27, CC: NN17 bottom-2.7 m.y. Core catchers 28, CC and 29, CC are barren.

## **Upper Miocene Nannoplankton**

From Sections 490-30,CC down to 490-59,CC the nannoplankton assemblages have few coccoliths. The dominant species are Discoasters guingueramus, which are five-rayed discoasters with distinct knobs on either sides of their central areas, and common sphenoliths (Sphenolithus abies and S. neoabies). There are also rare D. calcaris and D. variabilis in 30 to 32, CC and D. brouweri and D. surculus. No ceratoliths were recorded. Whether these assemblages are reworked or not is an open question. From 30,CC to 32,CC one might still consider reworking of NN11 nannoflora to a poor NN17 nannoflora. Deeper down, however, D. quinqueramus and the sphenoliths are the main constituents of the poor assemblages. Nannoplankton zones NN16 to NN12 are lacking and should be recognizable by common D. surculus and Ceratolithus tricorniculatus. The NN11 D. quinqueramus range zone comprises a long interval of 4.5 m.y. in the upper Miocene-that is, from 5 to 9.5 m.y. ago.

## **Siliceous Microfossils**

Silicoflagellates were found only sporadically except in Sections 490-5, CC and 490-30, CC, where the following species occur: Section 490-5, CC: Dictyocha stapedia, D. fibula, and Mesocena elliptica. This places Section 490-5, CC in the M. elliptica Zone (Bukry and Foster, 1973) or the M. quadrangula Zone (Martini, 1976), in the lower to middle Quaternary (top NN19). In Section 490-30, CC we found D. fibula, D. perlaevis, D. rhombica, D. speculum, and D. crux. This places 30, CC in the D. fibula zone (Martini, 1976) in the upper Miocene (NN11) to middle Pliocene (top NN15). Actiniscus tetraterias, which has an upper Pliocene to lower Pleistocene range (Dumitrica, 1973), was found in 16, CC. Hole 490 penetrates Miocene to Quaternary glauconitic muds, muddy silts, and siltstones. The sediment in the lower part of the section is indurated and could not be fully disintegrated for foraminiferal study.

Cores 1 to 12 contain abundant well-preserved upper Quaternary foraminifers. Cores 13 to 28 contain poorly preserved lower Quaternary and Pliocene foraminifers. Cores 29 to 64 are nearly barren of foraminifers except for reworked upper slope species.

Scarcity of the foraminifers in the sedimentary section precludes application of standard foraminiferal zonation. Nevertheless, abundance and preservation of foraminifers are useful in distinguishing the depth of deposition of inner slope sediments.

Cores 1 through 16 contain the following Quaternary foraminifers: Globorotalia flexuosa, G. tumida, G. menardii, Globigerinoides ruber, G. triloba, G. sacculifer, G. quadrilobatus, Neogloboquadrina dutertrei, Pulleniatina obliquiloculata, and Orbulina universa.

The planktonic foraminiferal assemblage in the section indicates tropical to subtropical watermasses. Benthic foraminifers in the upper Quaternary sediments are abundant and diverse and correspond well to the modern water depth (lower bathyal). Some of the typical forms include *Melonis pompilioides, Planulina wuellerstorfi, Gyroidina soldanii, Hoeglundina elegans, Oridosalis umbonatus, Pullenia bulloides, Cassidulina subglobosa, Uvigerina proboscidea, U. aculeata, U. hispida, Lagena, Stilostomella*, and, occasionally, *Pyrgo.* 

The Pliocene section, in Cores 17 through 30, contains the following species: *Globigerinoides fistulosus?*, *G. ruber, Hastigerina aequilateralis, Pulleniatina obliquiloculata*, and *Neogloboquadrina humerosa*. The Pliocene-lower Quaternary assemblage also contains shelf or upper slope reworked forms that include *Cancris, Hanzawaia*, and *Bolivina*. Certain sandy layers in the section—for example, Sample 490-24-1, 18-20 cm shows a high concentration of upper slope foraminifers that were transported down by turbidity currents.

## **Depositional Environment**

The absence of foraminifers in the Miocene sediment at Site 490 indicates dissolution below the CCD. Relatively abundant partially preserved foraminifers in the Pliocene and lower Quaternary sections reveal an upward shallowing trend through the CCD, whereas the abundant and diverse well-preserved upper Quaternary foraminiferal assemblage indicates deposition above the CCD.

## Radiolarians

Two factors affect the radiolarian record at Site 490: dilution by a fairly high sedimentation rate and contamination by upper Miocene through Quaternary radiolarians reworked from upslope outcrops. Contamination is the more serious problem, because extinction levels are difficult to recognize and many seem to be in error.

Approximately the upper one-third of Site 490 is Quaternary. The Quaternary/Pliocene boundary is difficult to locate because *Pterocanium prismatium* (Riedel and Sanfilippo, 1971) is absent in this area. Within the Quaternary, the extinction level of *Axoprunum angelinum* (0.4 m.y. ago, Hays, 1970) occurs within Core 2 at 15 meters, and the extinction level of *Anthocyrtidium angulare* occurs within Core 18. Pliocene radiolarians are moderately to poorly preserved, and a largely barren interval occurs between Cores 33 and 50 (305-445 m). The only recognizable datum plane is the evolution level of *A. angulare* (280 m), which occurs midway through the *P. prismatium* Zone (Dinkelman, 1973).

An upper Miocene/lower Pliocene fauna consisting of Ommatartus antepenultimus, O. penultimus, Stichocorys peregrina, and A. angelium defines the S. peregrina Zone (Dinkelman, 1973) below Core 61. The paleomagnetic data suggest that Hole 490 bottoms in the uppermost Miocene, which is consistent with this faunal zone interpretation. Poor preservation above Core 61 precludes accurate location of the Miocene/Pliocene boundary, but extrapolation suggests location near 560 meters in Core 62.

Deposition of reworked radiolarians began at Site 490 in the mid-Pliocene and continues actively today. The mid-Pliocene initiation of Miocene contaminant microfossils may mark the onset of erosion of middle and upper Miocene sediments upslope.

# SEDIMENT ACCUMULATION RATES

The age versus depth curve is based on biostratigraphic datum levels and some paleomagnetic reversal boundaries (Fig. 5). The presence of the *Axoprunum angelinum* extinction level within Core 1 suggests a slow accumulation rate of 10 m/m.y. for the uppermost portion of Site 490. The carbonate nodule with animal borings found in Core 1, Section 4, suggests an erosional and/or nondepositional episode within Core 1. Rates for the rest of Site 490 show no consistent trend but vary between 90 and 405 m/m.y.

## PALEOBATHYMETRY AND VERTICAL TECTONICS

Paleobathymetric determinations based on assemblages of benthonic foraminifers is impossible at Site 490 because of the poor foraminiferal preservation below 90 meters and total absence of large foraminifers below 280 meters. As a result, the paleobathymetry is estimated by relating the carbonate preservation at Site 490 to the level of the CCD. This trend suggests that late Miocene deposition at Site 490 began below the CCD, then passed through it in the early Pliocene time (4 m.y. ago) (Fig. 5) (McMillen and Bachman, this volume). Uplift rates calculated by removal of sediment accumulation of 580 meters at Site 490 are roughly 200 m/m.y. since the late Miocene.

## PALEOMAGNETISM

Paleomagnetic analyses at Site 490 established magnetostratigraphy and determined dips of certain beds and faults. Sediments in Hole 490 can be correlated with



Figure 5. Age vs. depth and inferred paleobathymetry at Site 490.

polarity epochs to Epoch 5. Dip direction and stress field are concordant with submarine topography around the site.

We collected 105 oriented samples from the sedimentary sequence of Hole 490. Stability of remanent magnetization of selected samples was checked with stepwise AF demagnetization (Niitsuma, this volume).

Average intensity of the samples is  $10^{-6.6\pm0.9}$  emu/cc after 15 mT AF demagnetization. Intensities in sandstones are one or two orders of magnitude less, suggesting that sandy material originated in a different source area than sediments. Changes in inclination suggest seven magnetozones.

The magnetozones can be correlated with the Brunhes normal polarity epoch (1), the Matuyama reversed polarity epoch (2), the Gauss normal polarity epoch (3), the Gilbert reversed polarity epoch (4), the normal events of a and  $C_2$  in the Gilbert Epoch, early part of Gilbert Epoch (4), and Epoch 5, as shown in Figure 1. This correlation is consistent with nannoplankton fossil data.

Sediments from lower part of Site 490 dip from  $10^{\circ}$  to 81°. If the drilling core axis is vertical, orientation of bedding plane can be calculated from magnetic data (Niitsuma, this volume).

Most bedding planes dip westward; a few dip northward. Mean and standard deviation of westward dips are toward N80°  $\pm 25$ °W. This direction corresponds to the topographic slope direction in the area of Site 490.

The change in dip angle with subsurface depth is represented as follows: dips of sediments below 350 meters gradually increase with depth until they are cut by normal faults, with slickensides dipping from 10° to 75° toward N84°  $\pm$  17°W between 400 and 430 meters. Dips gradually decrease with depth below this fault zone. The change suggests that the steep dips are related to the normal faults.

Many faults with slickensides were observed in the lower part of sediments in Hole 490. The stress field associated with the conjugate set can be calculated using the fault plan orientations and fault displacements. We find that the compressional axis is almost vertical and the tensional axis horizontal with east-west direction. Direction of the tension axis is concordant with the dip direction of bedding plane. Thus the steep dip of bedding plane may have been caused by drag folds associated with westward-dipping normal fault. The westward slope of submarine topography can also be explained by this tectonic movement.

## **ORGANIC GEOCHEMISTRY**

The shipboard organic geochemistry monitoring program consisted of analysis of gases released in core liners, degasification of selected sediment samples, and visual inspection for fluorescence in split core.

## Gases

Moderate amounts of gas were released in core liners from depths of 30 meters and below. The gas initially contained methane,  $CO_2$ , and small amounts of  $H_2S$ . The last, detectable by its distinctive odor, was present down to depths of about 60 meters. Methane content remained fairly constant with depth (Fig. 6) except for a minimum in the 420 to 470 meter interval. This interval was associated with a high sand content in the sediments that resulted in substantial gas and sediment loss from washout and dilution of core liner gas with air gases.

Ethane content was below the detection limit of the Carle gas chromatograph at depths shallower than 155 meters but increased gradually with depth to reach a maximum of 0.18% by volume. The methane to ethane

ratio also increased with depth, reaching a maximum of  $26.3 \times 10^{-4}$  near the base of the cored section (Fig. 6).

 $CO_2$  content in core liner gases varied from 6.5% to 0.24%, being higher in the upper portions of the cored sequence and decreasing with depth.

Hydrocarbons in the  $C_3-C_5$  range were monitored on the Hewlett-Packard 5710-A gas chromatograph from a depth of 165 meters to T.D. Their abundance was found to vary with depth and showed a maximum in the 300 to 400 meter interval (max. 879 ppm, Fig. 6).

Upon splitting Core 15 (Sample 490-15-5, 110 cm at a sub-bottom depth of 139.6 m), gas-releasing ice inclusions were observed near a gas expansion void of some 8 to 10 cm in length (Fig. 7). These ice fragments released gas in the ratio of 0.91 volume of gas per volume of ice. The evolved gases consisted of methane (76, 77%) and  $CO_2$  (23, 23%) with no ethane detectable on the Carle gas chromatograph. Degasification of a sample at a sub-bottom depth of 155 meters by high-speed blending showed a gas-generating ratio of  $4.2 \times 10^{-2}$  volume of gas per volume of sediment and with a similar composition (methane, 74.65%, and  $CO_2$ , 25.34%).

An *in situ* interstitial water sample was retrieved from near this depth (189.5 m) and released gas in the ratio of 0.47 volume of gas per volume of water (GGR =  $47 \times 10^{-2}$ ); the composition of the gas was 87.10% methane and 12.90% CO<sub>2</sub>.

Gas released from samples involving pore fluids and sediments showed a  $CO_2$  content considerably higher than did the core liner gas. This indicates that core liner gas composition is significantly affected by the water solubilities of the different gases at the pressures and temperatures prevailing during sampling and result in  $CO_2$  depletion of the gas released in core liners. Core temperatures of  $-2^{\circ}C$  to  $8^{\circ}C$  have been observed, and the gas pressure depends on the degree of induration of the core material and on the presence or absence of closure of the core against the liner. There is no control of these parameters.

We were unable to establish the origin of the ice in Core 15; it could have originated in the freezing of a mixture of interstitial and seawater trapped between the core and core liner. The decrease in core temperature necessary for ice generation may not be adequately accounted for by gas release and expansion; possibly a phase change would be required as well. Degasification of *in situ* interstitial water indicates that the amount of dissolved gas is of the order of 1 ml gas/ml of interstitial water.

#### Fluorescence

Split cores showed no evidence of fluorescence due to crude oil or bitumen impregnation.

### Organic Carbon, Hydrogen, Nitrogen, and Carbonate

Samples for CHN and carbonate analysis were taken from selected cores and analyzed as indicated in the Site 487 Site Chapter. The results of these analyses are shown in Figure 6. Throughout the cored sequence, the organic carbon content varied from 0.80% to 3.0% and the total nitrogen content from 0.06% to 0.27%. The



Figure 6. Methane and methane/ethane ratios, ethane and higher hydrocarbon abundance, and carbon and nitrogen abundances.

organic potential of the sediments ranged from good for the upper 200 meters to intermediate for the lower section. The carbon content decreases with depth, being higher in upper sediments and decreasing in the lower part of the sequence. The nitrogen content also decreases with depth and parallels the decrease in organic carbon.

C/N varied from 11.9 to 16.8, remaining approximately constant throughout the hole (Fig. 6). This is in the range for organic matter associated with recent sediments (Fairbridge, 1972) and suggests that the organic matter present in these sediments has a low degree of thermal maturation. Low amounts of carbonates (up to 4%) were detected in the upper 250 meters of the hole, decreasing below the detection limit of the carbonate bomb analysis.

## Conclusions

Gases, mainly of biogenic origin, were detected throughout this hole, causing a low to moderate degassing of the cores. The organic potential of the sediments varied from intermediate to good, but the C/N suggests a low degree of geothermal maturation. No evidence of petroleum or bitumen impregnation was detected. Further, suspected hydrates consisting of gas-releasing ice inclusions and frozen sediments were detected in the 145.4 to 145.7, 166.5 to 166.8, and 364.0 to 364.5 meter intervals. Heavier hydrocarbons,  $C_3-C_5$ , increased with depth without evidence of an increase in geothermal maturation. The origin of these gases is not known. They could have originated in more geothermally matured sediments and been emplaced in shallower depths by migration, but a biogenic origin cannot be excluded.

#### PHYSICAL PROPERTIES

Site 490 physical property analyses included porosity, water content, wet bulk density, compressional sound velocity, and undrained shear strength determined by standard DSDP techniques (Boyce, 1976). Significant changes in Site 490 physical properties with depth correspond directly with changes in lithology. From 0 to 140 meters (Unit 1, see lithologic column) these properties change very slowly (Figs. 8 and 9), typical of sediments deposited in a lower slope environment at moderate sediment accumulation rates. Between 140 meters and 400 meters (Unit 2) the change in physical properties corresponds to the transformation from soft sediment to mudstone. Below 400 meters (Unit 3) a small but significant change in these properties may reflect an increase in deformation. A general summary of these individual properties is presented in the following sections.

#### Porosity, Water Content, and Bulk Density

Porosity decreases from 68% at 2.3 meters to 35% at 540 meters (Fig. 8). This fairly uniform decrease suggests gravitational rather than tectonic forces as the primary dewatering mechanism. Water content decreases regularly from 48% at 2.3 meters to 16% at 540 meters.



Figure 7. Gas-releasing ice inclusions (C15 SS-110).





Bulk density decreases from 1.53 Mg/m<sup>3</sup> at 2.3 meters to 2.11 Mg/m<sup>3</sup> at 540 meters.

## **Compressional Sound Velocity**

Only two velocity measurements were performed above 350 meters sub-bottom because of gas attenuation. Below 350 meters velocity varies, ranging from 1.71 km/s at 427 meters to 1.88 km/s at 485 meters.

## INHOLE TEMPERATURE MEASUREMENTS

Three inhole temperature measurements yielded usable data only at 227.5 meters. The typical curve, with a slow increase in temperature as the probe and sediments equilibrate, is observed for this run at 227.5 meters, giving a gradient of  $2.5^{\circ}$ C/100 m, which agrees well with the minimum gradient of  $2.1^{\circ}$ C/100 m predicted on the basis of reflection depths and hydrates phase relationships (Shipley et al., 1979).

Thermal conductivity measurements using sediment slabs yielded values of 2.3 to 2.8  $\mu$ cal/cm s°C. These values result in heat flow values of about 6 or 7  $\mu$ cal/ cm<sup>2</sup> s, much higher than predicted. We suspect conductivity measurements are in error, probably owing to the lack of needle probes normally used to measure conductivities of soft sediment. Shipley and Shephard (this volume) discuss thermal measurements made during Leg 66.

## CORRELATION OF SEISMIC REFLECTION DATA AND DRILLING RESULTS

Site 490, 13 km seaward of Site 489, is interpreted to be seaward of the continental basement. A migrated



Figure 9. Shear strength vs. depth.

seismic reflection line crosses the drilling site at the position indicated in Figure 10.

Seaward of the site several landward-dipping reflections in the deeper part of the section are interpreted as part of the accretionary zone. In detail, Site 490 lies within an anomalous zone along the seismic line where the normally strong sub-bottom reflection at 0.38 s (340 m) is discontinuous (Fig. 3). Above 0.38 s several reflections in the slope apron dip gently northeast and appear to be truncated at or near the seafloor. The lack of an unconformity suggests that the seafloor "truncation" results from a depositional process and that dipping reflections are part of a prograded slope facies.

The drilled section has a velocity of 1.8 km/s, as derived from analysis of multichannel seismic data. No refractors were detected from the several sonobuoy experiments within the upper slope area. A velocity of 1.8 km/s and a penetration of 585.5 meters is equivalent to 0.65 s (Fig. 10).

Correlation of reflectors to the cored section is difficult, because the number of reflections exceeds observed lithologic and physical property changes. However, one reflection (0.30 s, 270 m), may correspond to thin limestone beds and/or to the beginning of fractures.

#### SUMMARY AND CONCLUSIONS

Site 490 penetrated 588.5 meters of mud, silt and muddy siltstone ranging in age from latest Miocene to Quaternary. Locations of age boundaries are vague because of carbonate dissolution in the lower cores and reworking evident in upper cores.

Fracturing with slickensides, first observed between 300 and 400 meters, became pervasive below 400 meters. A significant increase in induration accompanied the development of pervasive fracturing. Paleomagnetic data and slickenside orientations suggest northwest-dipping normal faults with minor reverse faults between 400 and 430 meters. We found no evidence of the compressive stress that would be expected to accompany accretion.

Paleontologic data show that the pre-middle Pliocene part of the section was deposited below the CCD. Subsequent uplift raised the section through the CCD and thence to its present position. We estimate uplift rates to be 200 m/m.y. Sedimentation rates show no zonation except for the extreme top of the hole, where 10 m/m.y. is observed. Rates in the lower part of the hole are irregular, averaging about 150 m/m.y.

Reworked Miocene and later fauna evident in middle Pliocene and younger parts of the section probably derive from upslope erosion, as evidenced by a pronounced unconformity at Site 489.

Methane was evident throughout the section. Ethane ranged from lower limits of detectability to a maximum of 0.18% near the bottom of the hole. Smaller amounts of  $C_3-C_5$  were also detected. A gas-releasing ice inclusion at 137 meters produced a 0.91 ml CH<sub>4</sub>/ml H<sub>2</sub>O. The gas could be either an *in situ* hydrate or incorporated in seawater frozen because of cooling from gas expansion during core recovery. Organic carbon content ranged from 0.8 to 3.0%, with highest percentages occurring in the upper 200 meters. All organic carbon showed a low degree of thermal maturity. Traces of more mature hydrocarbons ( $C_2-C_5$ ) probably migrated from rocks below the hole.

We conclude the following:

1) Oldest sediments observed in Hole 490 were deposited in a lower slope environment. Younger rocks were deposited in progressively shallower environments as the section was uplifted, probably owing to underplating during accretion.

2) The section was normally faulted in middle(?) Pliocene, resulting in uplift of rocks to the east and formation of a small basin to the west. The coincidence of uplift and the upslope hiatus at Site 489 suggests a cause and effect relationship. Redeposition of microfossils after the early Pliocene probably correlates with the middle Miocene-Quaternary unconformity observed in Hole 489A.

3) There is no firm evidence relating Hole 490 sediments to the accretionary zone. We find no sands suggestive of deposition in a trench environment, no steeply landward-dipping bedding indicative of rotation during ing underplating, no fracture orientations indicative of north-south compression, and no landward-dipping reflectors in the immediate vicinity of 490. Thus these sediments were probably deposited on the lower slope, either as part of a slope apron or as a slope deposit in the gap between the accretionary zone and the seaward lip of continental crust. The magnitude of compressional stresses transmitted to Hole 490 sediments during underplating was evidently inadequate to completely deform the sediments, probably because underplating occurred relatively deep beneath the bottom of the hole.

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Figure 10. Seismic reflection profile near Site 490.

PHIC	СН	FOSSIL									PHIC	CH	OSSIL				П		
TIME - ROCH UNIT BIOSTRATIGRA ZONE	NANNOFOSSILS	RADI DLARIAMS DIATOMS	CENTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURDANCE SEDIMENTARY STRUCTURES	LITHOL	LOGIC DESCRIPTION		TIME - ROCH	BIOSTRATIGRA ZONE	FORAMINIFERS	RADICLARIANS DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
OUATERNARY NN 21/NN 20	AC	FG CG CG		2 2 3 4			Prominent glauconits     glauconits       Prominent glauconits     black       -concentrations between 60 and 85 cm     SMEA       2 x 1.5 cm indurated moderate olive brown calcareous claystone concertion     TEXT Clay Clay Clay Clay Composition       Diffuse concentration of glauconite     Mice Heary Pyrite concertion     Text Clay Clay Clay Clay Clay Clay Clay Clay	DY SILT, grayish elive (10Y (BGY 3/2), tot to firm, structure inite grains, dark greenish gray (BGY 2/1). <b>IX SLIDES</b>	4/2) to gravish olive refers with and sized 5GY 4/11 to greenish 19 10000000 2 100 4-30 (D) (M) 8 1 57 44 36 55 46 42 10 7 2 3 1 1 1 - 2 - 3 1 7R 1 7R 1 7R 1 7R 1 7R 1 7R 1 7R 1 7R	QUATERNARY	NN 20				2 2	OG WW		•	VOID     - Sx 2 cm limetator ?concertion, bord, same color as mud     - S0-62 cm - diffue ana of moderate olive brown (5Y 4/4) lightly calcareous mud     MUDDY SILT, grayish olive green (5GY 3/2) moderat firm, apparently structurelist, Occasional white spon gricule contemprations. SMEAR SLIDES     15 mm x 4 mm wood

VANUAL DISCUPENDEN       VANUAL DI	E 490 HOLE	DRE 3 CORED INTERVAL 18.5-28.0 m SITE 490 HOLE CORE 4 CORED INTERVAL 28.0-37.5 m	
Subject         Camport         Camport         Subject         Camport         Subject         Subject <t< td=""><td>E FOSSIL CHARACTE</td><td>U FOSSIL CHARACTER</td><td></td></t<>	E FOSSIL CHARACTE	U FOSSIL CHARACTER	
NOTE         Notestative (minutes)         Notestative	UNIT BIOSTRATIGRA ZONE FORAMINIFERS NANVOFOSSILS RADIOLARIANS DIATOMS	SR APHIC LITHOLOGY SUBJECT STREET SUBJECT STREET SUBJECT STREET SUBJECT SUBJECT SUBJEC	
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6 Giauconitic sand unit, top and base diffuse, 		Glauconitic sand unit, top and base diffuse, 	

	DHIC	C)	FOS	SIL									PHIC		FOS	SSIL							
TIME - ROCK UNIT	BIOSTRATIGRA	FORAMINIFERS	RADIOLARIAMS	DIATOMS		SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION	TIME - ROCK	BIOSTRATIGRA	FORAMINIFERS	NAWNOFOSSILS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	STRUCTURES SAMPLES	LITHOLOGIC	DESCRIPTION
QUATERNARY	NN 21/WN 20	CMC	M A	6	-	2 2 2 2 2 2 2 3 3 3 4 4 4 4 4 4 4 4 4 4			10	GZ •	-VOID MUDDY SILT, clive grav (5Y 3/2) to gravith olive (10Y 4/2), soft to moderately firm, structureless with minor gluconic SAND, greenish black (5G 2/1). Gas expansion cracks common. SMEAR SLIDES 9 9 9 9 9 9 9 9 9 9 9 9 9	QUATERNARY	NN 20	CM	CM F	26	3	0.5 1.0		1	•	B0-90 cm - concentration 	LT, olive gray (5Y 1/2) homogeneous. Gracks common. Scattered sponge remains artificially compacted lextruded from Comality. DES

	PHIC		СНА	OSS	TER	T			T					
TIME - ROCK UNIT	BIOSTRATIGRAU	FOR AMINIFERS	<b>WANNOFOSSILS</b>	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SEDIMENTARY	SAMPLES	LITHOLOGIC D	ESCRIPT	ION
						1	0.5		00		+	MUDDY SIL homogeneous smail {2-5 m common. SMEAR SLID	, olive gr scattere n) diffus S	ay (5Y 3/2) soft to moderate firm d sponge remains, scattered ver e sand blebs. Get expansion crack
DUATERNARY	NN 20					2	a contraction of	VOID				TEXTURE: Sand Silt COMPOSITIO Quartz Feldspar Mice Pyrite Clav	2-11 (D) - 30 70 4: 17 2 3 1 70	0 3-71 (M) 100 - - - - - - - - - - - - -
		FM	FP	AG		3			0			Cary Zoolite Carb, unspec. Foraminifres 5→8 mm fine → sand bleb Soonge spicul Plant fragmen ORGANIC C2	70 TR TR TR 3 2 2 2 2 3 2 8 2 8 2 8 2 8 2 8 2 8 2 8	ND CARBONATE
												% Organic Car % CaCO <sub>3</sub>	ion 2.5	

SITE 490 HOLE CORE 8 CORED INTERVAL 66.0-75.5 m APHIC TIME - ROCK UNIT METERS NOI FORAMINIFERS BIOSTRATIGR DRILLING DISTURBANCE SEDIMENTARY STRUCTURES GRAPHIC LITHOLDGY NANNOFOSSIL RADIOLARIAN DIATOMS LITHOLOGIC DESCRIPTION MPLES. 0 MUDDY SILT, olive gray (5Y 3/2) soft to firm, homogen-cous, infrequent sponge remains, Firmer towards base of core. Gas expansion cracks 0.5 - Gas expansion crack SMEAR SLIDES 1.0-Mud 3-50 (D) 1 29 70 TEXTURE: 1 Sand Silt Clay COMPOSITION: COMPOSITION: Quartz Feldspar Mica Pyrita Clay Carb. unspec. Namofosilis Radiolarians Diatoms Sponge spicules Piant fragments 2 21 2 cm lamination of lighter (grayish olive, 10Y 4/2) mud 1 \_\_\_\_\_ 1 OG 70 TH TH TH 2 2 2 IW QUATERNARY NN 19 3 PP VOID CC FM FP CG

SITE 490

SITE	490	_	HO	.E		CO	RE	9 CORED	INTER	VA	75.5-85.0 m		
6	PHIC		CHI	OSS	TER								
TIME - ROCI	BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE SEDIMENTARY	SAMPLES		LITHOLOGIC DESC	RIPTION
QUATERNARY	NN 19	RM	FP	CG		1 2 CC	0.5			GZ +	Dispersed gas expansion cracks     Very firm     Gas expansion crack	MUDDY SILT, o very small (2–3 spicules, Firmer to SMEAR SLIDES TEXTURE: Sand Clay COMPOSITION: Quartz Pelsfpar Mica Pyrite Clay Diatoms Spong spicules Plant fragments GRAIN SIZE Sand Silt Clay	live gray (5Y 3/2), toft, homogeneous mm) blebs of fine sand and sponge wards base. 2.60 (D) 2 40 60 26 2 2 2 7 R 60 6 5 2 1 1 1.110 3,1 64.8 32.0
												ORGANIC CARB	ON AND CARBONATE
												% Organic Carbon % CaCO <sub>3</sub>	2.9 0

×	PHIC		F	OSS	TER				Γ			
TIME - ROC	BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING	SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION
QUATERNARY	NN 19	F00	NA	Ra	DIA	1	0.5		DAU SIG			DIATOMACEOUS MUDOY SILT, olive gray (5Y 32 to olive black (5Y 2/1), firm, homogeneous, scatter minute (1-2 mm) blobs of fine-grained sand. SMEAR SLIDES TEXTURE: 100 Sand 2 Silt 43 Clay 55 COMPOSITION: Duartz 28 Feldspar 2 Mica 2 Pyrite 2 Clay 55 Nanofonsis TR Radiolarians TR Baldoarians TR
						5		VOID VOID VOID VOID				Gas expansion crack     Gas expansion crack     Several gas     expansion cracks



SITE 490 13 CORED INTERVAL 113.5-123.0 m HOLE CORE FOSSIL CHARACTER TIME - ROCK UNIT METERS GRAPHIC TURBANCE LITHOLOGIC DESCRIPTION BIOSTRAT RADIOLAR SECT 1 - 0 QUATERNARY RP RP CG CC MUDDY SILT, olive black (5Y 2/1), soft. NN 19 SMEAR SLIDES Mud 1-16 (D) TEXTURE: Sand Silt . 29 Clay 70 COMPOSITION: Quartz 17 Feldspar Mica Pyrite Clay 70 Nannofossils TR TR Radiolarians Diatoms Sponge spicules Plant fragments

SITE	490	но	LE		COF	RE	15 CORED I	NTERVA	L 132.5-142.0 m		SITE	490	н	OLE		COR	E	16 CORED	INTER	RVAI	142.0–151.5 m
TIME - ROCK UNIT	ZONE	FORAMINIFERS	FOSSIL ARACTE SNUDOURINA SNUTANO	R	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES		LITHOLOGIC DESCRIPTION	TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FORAMINIFERS	FOSSIL HARACT STISSOJONNAN	DIATOMS H	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE DISTURBANCE SEDIMENTARY	STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION
QUATERNARY	01 NN	82 8	P CG		1 2 3 4 5 6 CC		Voib Voib Voib Voib Voib Voib		VOID     Orac expansion cracks     Ges expansion cracks     Ges expansion cracks     Ges expansion cracks     Concertion, without     sharp margins     Gas expansion cracks     Imm lamination of very fine     Gas expansion cracks     Gas expansion cracks     Concerts	Site 490, Core 14, 123.0–122.5 m: NO RECOVERY. MUDDY SILT, olive gray (5Y 3/2) soft, structureless to MUDDY SILTSTONE, olive gray (5Y 3/2) structureless tiblow 85 cm is Section 1). Section 4: Slightly darker mottling effect in this section due to bioturbation. Several diffuse areas, up to several orn across, light olive brown (5Y 6/18) to graying olive (10Y 4/2), calcarsous — Thiopient concretions (40, 70, 84, and 100 cm). Millimeter-to-to-itons (40, 70, 84, and 100 cm). Millimeter-to-out sime, and biots between 100 and 115 cm. Yale calcarsous area, as in Section 4, 120 cm; 5 x 3 x 1 cm wood fragment, 123–125 cm; 3 cm irregular ICE fragment (canned). SMEAR SLIDES $ \begin{array}{c} & & & & & \\ & & & & & \\ & & & & & \\ & & & & & & \\ $	UPPER PLIOCENE	81 NN/61 NN	RP R	IP CG		2 3 4 CC					D-28 cm soft, muddy slit MUDDY SILTSTONE, gravish olive (10Y 4/2) to olive gray (5Y 3/2), fme SAND beds and VITRIC SILT. Frozan sedimetral (canned) Frozan (canned)



PP

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CC

Fine sand blebs

Ash areas in Core-Catcher rubble

VOID



RP CG CG

5GY 2/1







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6	APHI		CH	ARA	CTER											
LIND	BIOSTRATIGR, ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES			LITHOLOGIC DESC	RIPTIO	v
						1	0.5			¢	•		6 cm foraminiferal muddy slit (13–22 cm) Several plant fragments up to 3 cm Thin (1–3 mm) slit intervals	MUDDY SILTSTO turbated. Thin silt core but most com pyrite nodules (up ing biscuits. SMEAR SLIDES	NE, oliv layer ar imon in to 1 m	ve grav (ISY 3/2), slightly bio d lamination occur throughou Sections 1, 2, and 3. Very fev m) common to abundant drill
						2	111111	VOID						TEVTUBE	E Foraminiferal	0986 Bilittone
4							1 H H H						VOID	Sand Silt Clay COMPOSITION: Quartz	25 35 40 50	10 50 40 54
UPPER PLIOCEN	NN 18					3	the first first				•		Pale blue green (5BG 7/2) 7reduction spots?clast - slightfy more indurated than matrix, 5 mm + 1 cm diameters (5 and 20 cm)	Feldspar Mica Heavy minerals Pyrite Clay Glass Carb. unspec. Foraminifers Nannofosils Radiolarians	8 5 1 13 2 3 15 1 TR	15 5 2 2 15 2 
						4	tri hundru							Diatoms Sponge spicules Silleoflagellates Plant fragments ORGANIC CARBO % Organic Carbon % CaCO <sub>3</sub>	TR 1 TR TR 3-112 1.6 0	2 1 - 1 CARBONATE
						5	Titte				100000		VAD			









SITE 490 HOLE

TIME - ROCK UNIT



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TIME - ROC UNIT	BIOSTRATIGR/ ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING	SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
PLIOCENE?						1	0.5					MUDDY SILTSTONE (silt where intensely deformed by drilling), gravish olive (10Y 4/2) structureless.
		RP	CM	FG		CC	-		lõ			Fractured siltstone



SITE 490

Carb, unspec

Foraminifers

Nannofossils

Radiolarians

Diatoms Sponge spicules

Fracturing (probably drilling disturbance)

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1 2 1 2

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LITHOLOGIC DESCRIPTION

SMEAR SLIDES

TEXTURE: Sand Silt Clay COMPOSITION Quartz Feldgar Mica Heavy minerals Pyrite Clay Qiauconite Carb, unspec. Nannofosili Diatami Sponge spicules

PP

0

5

6

cc

MUDDY SILTSTONE, olive gray (5Y 3/2), slightly bioturbated, common to abundant drilling biscuits. Many oblique fractures with slickensides.

Muddy

3-52 (D) 10 60 30

55 15 15

1 TR TR 1

ORGANIC CARBON AND CARBONATE 4.87 % Organic Carbon 1.3 % CaCO<sub>3</sub> 0

SITE	4	90	HOL	E.		c	OR	E	41 CORE	DINTER	VAL	370.0-379,5 m		_			SITE	490	H	DLE		C	ORE	42 CORE	INTER	RVAL	379,5-389,0 m
×	APHIC		CHA	OSSI	TER		T										×	APHIC	CI	FOS	CTER				T		
TIME - ROC	BIOSTRATIGR	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	NOT DOD	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENYARY STRUCTURES	SAMPLES		LITHOLOGIC DE	ISCRI	IPTION		TIME - RO	BIOSTRATIGR	FORAMINIFERS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY	STRUCTURES	
PLIOCENE?			RP	FG		C			VOID				MUDDY SILT: bioturbated, dr ated. SMEAR SLIDE TEXTURE: Sand Silt Clay COMPOSITIO Quartz Feltspar Mica Heavy minerals Pyrite Clay Glass Glauconite Clay Glass Bailocarian Radiolarian Diatoms Sponge spicule	STON illing :S	VE. olive gray (SY 3/2). Moderatel bisouits throughout, in places separately separately and the separately sep	Y	PLIOGENE2					2	0.5	VoiD			

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#### CORE 44 CORED INTERVAL 398.5-408.0 m SITE 490 HOLE FOSSIL TIME - ROCK UNIT RADIOLARIANS METERS BIOSTRATIGR GRAPHIC ANCE LITHOLOGIC DESCRIPTION RADIOLARIA ...... MUDDY SILTSTONE, olive gray (5Y 3/2) intensely bioturbated. Very highly fractured. Most planes sub-horizontal in Section 1. Section 1. 30-60 cm: Brecciated unit, pos-0.5 ibly primary broccia, but not cemented. In Section 2 the fractures are steeply dipping faults as well as sub-horizontal Ē 33 tractures are steeply disping faults as well as sub-horizontal planes occur. Bedding is defined by concentrations of aligned burrows, 52–58° apparent dips. Numerous slicken-sides on faults, in two cases showing dip-slip (normal) 1.0motion. PLIOCENE SMEAR SLIDES Muddy 2-38 (D) TEXTURE: Sand Silt 8 62 Clay 30 COMPOSITION: 61 Quartz 3 Feldspar 10 Mica 2 0 cc Heavy minerals õ Pyrite Clay Zeolite 20 TR TR TR TR TR Carb, unspec. Nannofossils Rediolarians Diatoms Sponge spicules 3 ORGANIC CARBON AND CARBONATE CC-16 % Organic Carbon 1.3 % CaCO<sub>3</sub> 0











Source in the second	9	T	1		FOS	SII	-	T	-	CONC.D I				1		T	9	r-
Out_DID         Out_DID <t< th=""><th>APHI A</th><th></th><th>- 3</th><th>CH</th><th>ARA</th><th>CT</th><th>TER</th><th></th><th></th><th></th><th></th><th></th><th>11</th><th></th><th>×</th><th></th><th>PHI</th><th>1</th></t<>	APHI A		- 3	CH	ARA	CT	TER						11		×		PHI	1
Suggest Sug	TIME - HOC UNIT BIOSTRATIGRI ZONE	ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS		DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING	SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION	TIME - ROC	- INO	ZONE	FORAMINIFERS
2 cm: fault, 85° dip, 240 azimuth, silckensides pitch 15–90° azimuth 2 cm: fault, 85° dip, 240 azimuth, silckensides pitch 15–90° azimuth 25 cm: bedding, 40° dip fault, 45° dip, 200 azimuth, silckensides at 190° azimuth (læge) and 170° azimuth (imail) About 10 fractures (section cut at low angle to slickensides)	MIOCENE?							3	0.5	VOID				<ul> <li>10 cm: tedding 42°, asimuth 27° 12 cm: fault 55°, asimuth 274° 24 cm: tin working hall 3 cm x 3 mm<sup>2</sup> woody fragment</li> <li>MUDDY SILTSTONE, olive gray (5Y 3/2), moderate bio- turbation, drilling biscuits common.</li> <li>ORGANIC CARBON AND CARBONATE 3-57 % Organic Carbon 1.5 % CaCO<sub>3</sub> 0</li> <li>Seven fractures (2 along bedding plane), 1 (44 cm lavel) with sideknisides</li> <li>44 cm: fault, 38° dip, 240 azimuth 216°</li> <li>Four or five fractures within sideknisdes</li> <li>Four or five fractures within sideknisdes</li> <li>Four or five fractures within sideknisdes</li> <li>2 cm: fault, 85° dip, 240 azimuth 214° 55 cm: badding, 40° dip fact, 45° dip, 240 azimuth, sidekensides pitch 15–90° azimuth 25 cm: biologing, 40° dip fact, 45° dip, 240 azimuth, sidekensides pitch 15–90° azimuth 25 cm: biologing, 40° dip fact, 45° dip, 240 azimuth, sidekensides pitch 15–90° azimuth 25 cm: biologing, 40° dip fact, 45° dip, 240 azimuth, sidekensides pitch 15–90° azimuth 25 cm: biologing, 40° dip fact, 45° dip, 230 aximuth, sidekensides pitch 15–90° azimuth 25 cm: biologing, 40° dip fact, 45° dip, 230 aximuth, sidekensides pitch 15–90° azimuth 25 cm: biologing, 40° dip fact, 45° dip, 230 aximuth, sidekensides pitch 15–90° azimuth (large) and 17° azimuth (simul)</li> <li>About 10 fractures (section cut at low angle to silckansides)</li> </ul>	MIOCENE			

UNIT BIDSTRATIGRAPHIC ZONE FONAMINIFERS		3	CHA	RAC	TER	S															
		FORAMINIFERS	NANNDFOSSILS	RADIOLARIANS	DIATOMS		SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION								
							1	0.5					38 cm: bedding, 37° dip MUDDY SILTSTONE, olive gray (5Y 3/7), moderate turbation, drilling biscuits common, Four fractures, apparently along bedding defined by burrow concen- tions.								
								1					Muddy siltstonn								
							2	and software					1-30         3-37           TEXTURE:         (D)         (M)           Sand         5         -           Sit         75         100           COMPOSITION:         0         -           Outartz         10         -           Faidopar         10         -           Mica         2         -           Privite         3         -           Clay         20         -								
MICCINEL							3	arear have burn	OG IW				Giais – 100 Carb, unspec. 1 – Nannofossits TR – S2 cm: fault, 62° dip, 10° azimuth, Sponge spicules TR – eart, footwall facing 270°, pitch 34° to 27°. 64 cm: fault, 63° dip, 348° azimuth, silckenides pitch 34–90° (from notobas in footwall facing 270°, pitch 48° to E. suggesting normal								
							4	the forest rest					movement with right lateral components 33 cm: fault, 56° clip, 58–74° azimuth, slickensides pitch 68° NE								
							5														

SITE 490 HOLE	CORE	60 CORED INTERVAL	541,0-550,5 m	SITE	490	но	.E	CORE	61 CORED	INTERVAL	550.5-560.0 m	
TIME - ROCK UNIT BIOSTFATIGRAPHIC FORAMINIFERS NANNOFOSSIS ANNOFOSSIS ANNOFOSSIS NANNOFOSSIS ANNOFO	SECTION	GRAPHIC LITHOLOGY Wenthings Children	LITHOLOGIC DESCRIPTION	TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FORAMINIFERS	ANDIOLARIANS	SECTION	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES		LITHOLOGIC DESCRIPTION
GINIOCENE)	2 2 3 CC		S0 cm: fault, 45–59° dip slickensides, pitch 50° to NE 	MIOCENE?			66	2 3 4 0.5 1 1.0-	<ul> <li>Ale and a set of /li></ul>	2000	Only minor fracturing in Sections 2 and 3 20° true dip picked out by Mattened burrows – aligned mud chips	MUDDY SILTSTONE, olive grav (SY 3/2), fractured, alickensides, moderate bioturbation. Common drilling bicuits. SMEAR SLIDES Pyysia 273 TEXTURE: (D) Sand 5 Silt 55 Clay 40 COMPOSITION: Ouart 69 Feldspar 5 Mica 5 Heavy minerals 3 Pyrite 1 Clay 15 Glas 2 Carb. unsprec. TR Sponge spicules TR



TIME - ROCK UNIT

UPPER MIOCENE

55

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Strie

DITARATIO

SITE 490









Site 490



























![](_page_54_Figure_0.jpeg)

SITE 490

![](_page_55_Figure_1.jpeg)

![](_page_56_Figure_0.jpeg)

Site 490

![](_page_57_Figure_1.jpeg)

![](_page_58_Figure_0.jpeg)

![](_page_59_Figure_1.jpeg)

![](_page_59_Figure_2.jpeg)

![](_page_60_Figure_0.jpeg)

SITE 490

![](_page_61_Figure_1.jpeg)

![](_page_62_Figure_0.jpeg)

SITE 490

![](_page_63_Figure_1.jpeg)

![](_page_64_Figure_0.jpeg)

Site 490

![](_page_65_Figure_1.jpeg)

![](_page_66_Figure_0.jpeg)