Shipboard Scientific Party²

HOLE 491

Date occupied: 13 April 1979

Date departed: 20 April 1979

Time on hole: 182.1 hours

Position: 16°01.74'N; 98°58.33'W

Water depth (sea level; corrected m, echo-sounding): 2883

Water depth (rig floor; corrected m, echo-sounding): 2893

Bottom felt (m, drill pipe): 2870.0

Penetration (m): 542.0

Number of cores: 59

Total length of cored section (m): 542.0

Total core recovered (m): 388.0

Core recovery (%): 72

Oldest sediment cored:

Depth sub-bottom (m): 542 Nature: Muddy siltstone and sand Age: Early Pliocene

Principal results: At Site 491 we penetrated 542 meters and recovered 59 cores that comprise three lithologic units (Table 1, Fig. 1). Unit 1 extends from 0 to 57.5 meters and consists of upper Pliocene to upper Quaternary mud. Unit 2 (57.5-437.5 m) is lower Pliocene muddy silt with minor fine sand layers. Unit 3 extends from 437.5 to 542 meters and is composed of lower Pliocene muddy silt and siltstone with interbedded fine to coarse pebbly sand.

Tilted beds and fracturing first occur at about 120 meters and continue the total depth through the lower two lithologic units. In Unit 2 dips are variable and range to nearly vertical; fracturing is present throughout, though slickensides are rare. Dip angles in Unit 3 range up to 30° and are more uniform than in Unit 2. Fracturing and slickenslides are more abundant in Unit 3 than in Unit 2, and a strong fissility develops locally. Paleomagnetic restorations of bedding in both Units 2 and 3 indicate modal dip directions to the north, though with considerable scatter.

Unit 3 and much of Unit 2 probably accumulated in a trench and/or lower slope environment. Deformation of these sediments occurred shortly after deposition and prior to being uplifted to the present midslope position.

BACKGROUND AND OBJECTIVES

Site 491 is located on the inner slope of the Middle America Trench at 16°01.74'N, 98°58.33'W in about 2.9 km of water (Fig. 2). The site lies about 2.1 km above and 14.25 km north-northeast of the adjacent trench floor on a steep average slope of about 9°. The multichannel seismic reflection profile adjacent to the site (Fig. 3) shows a thin slope blanket overlying a series of landward-dipping reflectors that are cut by a bottomsimulating reflector at about 0.75 s below the mudline.

Outcrops of a Mesozoic to Precambrian crystalline basement complex and a Mesozoic magmatic arc, anomalously close to the modern trench, argue for tectonic truncation of the Pacific margin off southern Mexico during the late Cretaceous or Paleogene (Karig et al., 1978). Results from Sites 489 and 490 demonstrate that the seaward margin of the basement complex extends to within 30 to 35 km of the trench. Seaward of the basement complex, the lower trench slope may be underlain by pelagic, hemipelagic, and trench sediments accreted during Neogene convergence in the Middle America Trench. A fundamental goal at Site 491 is to date any accreted deposits and compare their ages to similar sediment at Site 488 downslope and Site 492 upslope. If the simple accretionary model holds, then any class of offscraped sediments should increase in age landward.

An imbricate stack of landward-dipping thrust faults constitutes one of the most popular structural models of accretion for the lower trench slope (e.g., Karig, 1974; Seely et al., 1974). The data supporting this model consist primarily of landward-dipping reflectors, which are exceptionally well shown in the seismic reflection line through Site 491 (Fig. 2). Packages of reflectors with slightly discordant contacts are interpreted bedded sedimentary sequences separated by thrust faults (Seely et al., 1974). Alternatively, some of the dipping reflectors beneath trench slopes may represent faults. Because none of the previous drilling at active margins has established the origin of landward-dipping reflectors, a prime objective at Site 491 is to penetrate these surfaces and determine whether they represent faults, tilted bedding surfaces, or some other phenomena.

A prominent bottom-simulating reflector occurs beneath the lower slope on the multichannel seismic reflection profile adjacent to the site (Fig. 2) as well as on other reflection profiles in the site survey area (Shipley, this volume). Comparable bottom-simulating reflectors have been observed off northern Panama and Columbia and along the Middle America Trench margin between

²

Initial Reports of the Deep Sea Drilling Project, Volume 66. J. Casey Moore (Co-Chief Scientist), Earth Sciences Board, University of California, Santa Cruz, California; Joel S. Watkins (Co-Chief Scientist), Gulf Research and Development Company, Houston, Texas (present address: Geology and Interpretation Department, Exploration and Production Division, Gulf Science and Technology Company, Pittsburgh, Pennsylvania); Steven B. Bachman, Department of Geology, University of California, Davis, California (present address: Department of Geological Sciences, Cornell University, Ithaca New York); Floyd W. Beghtel, Phillips Petroleum Company, Bartlesville, Oklahoma; Arif Butt, Institut und Museum für Geologie und Paläontologie, Universität Tübingen, Tübingen, Federal Republic of Germany; Borys M. Didyk, Research and Development Laboratory, Empresa Nacional del Petróleo (ENAP), Concon, Chile; Glen Foss, Deep Sea Drilling Project, Scripps Institution of Oceanography, La Jolla, California; Jeremy K. Leggett, Department of Geology, Imperial College of Science and Technology, London, United Kingdom; Neil Lundberg, Earth Sciences Board, University of California, Santa Cruz, California; Kenneth J. McMillen, Geophysics Laboratory, Marine Science Institute, University of Texas, Galveston, Texas (present address: Geology and Interpretation Department, Exploration and Production Division, Gulf Science and Technology Company, Pittsburgh, Pennsylvania); Nobuaki Niit-suma, Institute of Geosciences, Shizuoka University, Oya, Shizuoka, Japan; Les E. Shephard, Department of Oceanography, College of Geosciences, Texas A&M University, College Station, Texas (present address: Sandia National Laboratories, Division 4536, Albuquerque, New Mexico); Jean-François Stephan, Département de Géotectonique, Université Pierre et Marie Curie, Paris, France; Thomas H. Shipley, Scripps Institution of Oceanography, University of California at San Diego, La Jolla, California; and Herbert Stradner, Geologische Bundesanstalt, Vienna, Austria.

Table 1. Coring summary, Hole 490.

| | Cored Interval below Bottom | Cored | Reco | vered | |
|------|--------------------------------|-------|------|-------|----------------------|
| Core | (m) | (m) | (m) | (%) | Remarks |
| 1 | 0.0-10.0 | 10.0 | 9.59 | 104 | |
| 2 | 10.0-19.5 | 9.5 | 6.34 | 67 | |
| 3 | 19.5-29.0 | 9.5 | 9.59 | 101 | |
| 4 | 29.0-38.5 | 9.5 | 7.86 | 83 | |
| 5 | 38.5-48.0 | 9.5 | 5.26 | 55 | |
| 6 | 48.0-57.5 | 9.5 | 9.62 | 101 | |
| 7 | 57.5-67.0 | 9.5 | 9.42 | 99 | |
| 8 | 67.0-76.5 | 9.5 | 9.53 | 100 | |
| 9 | 76.5-86.0 | 9.5 | 9.10 | 96 | |
| 10 | 86.0-95.5 | 9.5 | 9.80 | 103 | |
| 11 | 95.5-101.0 | 5.5 | 5.13 | 93 | |
| 12 | 101.0-105.0 | 4.0 | tr | 0 | N 127 12 1978 111200 |
| 13 | 105.0-114.5 | 9.5 | 2.07 | 22 | hole deviation 1.0° |
| 14 | 114.5-124.0 | 9.5 | 8.34 | 88 | |
| 15 | 124.0-133.5 | 9.5 | 8.18 | 86 | |
| 16 | 133.5-143.0 | 9.5 | 9.00 | 95 | |
| 17 | 143.0-152.5 | 9.5 | 9.26 | 97 | |
| 18 | 152.5-162.0 | 9.5 | 8.11 | 85 | |
| 19 | 162.0-1/1.5 | 9.5 | 7.01 | 80 | |
| 20 | 1/1.5-181.0 | 9.5 | 3.40 | 30 | |
| 21 | 181.0-190.5 | 9.5 | 4.85 | 51 | |
| 22 | 190.5-200.0 | 9.5 | 1.28 | 11 | |
| 23 | 200.0-204.0 | 4.0 | 1.94 | 49 | |
| 24 | 204.0-209.5 | 5.5 | 0.15 | 02 | |
| 25 | 209.3-219.0 | 9.5 | 9.40 | 100 | |
| 20 | 219.0-228.5 | 9.5 | 9.34 | 100 | |
| 28 | 228.3-238.0 | 9.5 | 6.02 | 73 | |
| 20 | 238.0-247.3 | 9.5 | 0.92 | 00 | |
| 30 | 257 0-266 5 | 0.5 | 0.40 | 100 | |
| 31 | 266 5-276 0 | 9.5 | 9.00 | 95 | |
| 32 | 276 0-285 5 | 9.5 | 7 94 | 84 | |
| 33 | 285 5-295 0 | 9.5 | 7.87 | 83 | |
| 34 | 295.0-304.5 | 9.5 | 8.31 | 87 | |
| 35 | 304.5-314.0 | 9.5 | 8.83 | 93 | |
| 36 | 314.0-323.5 | 9.5 | 7.62 | 80 | |
| 37 | 323.5-333.0 | 9.5 | 6.30 | 66 | |
| 38 | 333.0-342.5 | 9.5 | 2.95 | 31 | |
| 39 | 342.5-352.0 | 9.5 | 3.09 | 33 | |
| 40 | 352.0-361.5 | 9.5 | 8.52 | 90 | |
| 41 | 361.5-371.0 | 9.5 | 9.02 | 95 | |
| 42 | 371.0-380.5 | 9.5 | 7.30 | 77 | |
| 43 | 380.5-390.0 | 9.5 | 6.08 | 64 | |
| 44 | 390.0-399.5 | 9.5 | 4.98 | 52 | |
| 45 | 399.5-409.0 | 9.5 | 5.80 | 61 | |
| 46 | 409.0-418.5 | 9.5 | 1.68 | 18 | |
| 47 | 418.5-428.0 | 9.5 | 3.28 | 35 | |
| 48 | 428.0-437.5 | 9.5 | 9.22 | 97 | |
| 49 | 437.5-447.0 | 9.5 | 7.02 | 74 | |
| 50 | 447.0-456.5 | 9.5 | 8.48 | 89 | |
| 51 | 456.5-466.0 | 9.5 | 5.18 | 55 | |
| 52 | 466.0-475.5 | 9.5 | 3.52 | 37 | |
| 53 | 475.5-485.0 | 9.5 | 1.16 | 12 | |
| 54 | 485.0-494.5 | 9.5 | 3.39 | 36 | |
| 55 | 494.5-504.0 | 9.5 | 6.30 | 66 | |
| 56 | 504.0-513.5 | 9.5 | 6.18 | 65 | |
| 57 | 513.5-523.0 | 9.5 | 8.64 | 91 | |
| 58 | 523.0-532.5 | 9.5 | 5.17 | 54 | |
| 59 | 532.5-542.0 | 9.5 | 1.17 | 12 | |

Panama and Mexico (Shipley et al., 1979). These data all show increasing depth of the reflection with increasing water depth. They have high-amplitude reflections generally showing a reflection phase reversal and occurring at pressures and temperatures approximating the methane/hydrate, gas/solid phase boundary. Limited reflection amplitude and phase studies by Shipley et al. (1979), though restricted to areas where the seafloor is horizontal, suggest that the BSR along the Oaxaca-Guerrero transect may be a gas hydrate. However, 5° by 5° averages of thermal gradients (Langseth and Von Herzen, 1970) in the case of Mexico put the BSR well into hydrate stability field and not at the phase boundary. The bottom-simulating reflector may also represent a diagenetically induced mineral transition (e.g., Hein et al., 1978). In view of the uncertainty regarding the origin of the bottom-simulating reflector, we proposed to sample the sediments above it with a pressure core barrel to test for the presence of hydrate.

OPERATIONS

The Glomar Challenger steamed 9¼ nautical miles southeast from Site 490 to Site 491 in 2.25 hours. After dropping the beacon and lowering the pipe, we offset 850 meters to the north-northwest to attempt to locate on the multichannel reflection profile Line MX-16. Hole 491 was spudded at 0120 hours, 14 April and a completely full core obtained after the bit had been lowered to 2880 meters. One joint of pipe was set back, and a "punch core" extending to 2870.5 meters was attempted without recovery. Water depth was set at 2870 meters, 13 meters shallower than PDR depth.

The pressure core barrel was utilized in two unsuccessful attempts to recover cores under pressure in hopes of capturing gas hydrates. Otherwise, coring proceeded smoothly through a long and monotonous section of muddy silts and siltstones. Because of use of a special core bit with the pressure core barrel, the core diameter was somewhat reduced. Recovery was excellent and no hole problems were encountered, even in the lower 70 to 80 meters of loose, coarse sand. We terminated coring at 542 meters below the seafloor because of scheduling considerations.

We made 15-stand wiper trip to clear the hole to total depth. Unfortunately, the bit did not release, which precluded logging. The hole was filled with barite-weighted mud and the drill string recovered.

LITHOLOGIC SUMMARY

Site 491 is located midslope in a water depth of 2877 meters. Drilling penetrated 542 meters of early Pliocene to upper Quaternary argillaceous to sandy sediments. On the basis of the sedimentary characteristics, three units can be distinguished (Fig. 4, Chart 1, back pocket).

Unit 1, upper Pliocene to upper Quaternary (Cores 1-7, 0-67.0 m sub-bottom depth), consists mainly of grayish olive green (5GY 3/2) mud and muddy silt with minor thin lenses (1 mm-2 cm) of ash (Core 1, 4, 6) and calcareous concentrations (up to 6 cm, Cores 3-7). Sponge spicules (1-2 mm) and pyrite nodules are also present. The gas content of the sediment is high enough to produce numerous expansion cracks.

Bedding has not been observed, and fracturing is not apparent.

Unit 2, lower Pliocene (Cores 8-48, 67.0-437.5 m sub-bottom depth), is grayish olive green muddy silt locally indurated into muddy siltstone. Minor thin, fine-



Radiolarian zones 1. *Pterocenium prismatism* 2. *Spongaster pentus* B. Barren Magnetic epochs 1. Brunhes 2. Matayama 3. Gauss 4. Gilbert immud and muddy silt immuds sand ▲ frozen sediments Low dips-0 -20, moderate dips-20 -40, steep dips-40

Figure 1. Site summary diagram.

grained sand layers occur irregularly throughout the section (Figs. 5 and 6). The thickness of the sands is generally from 1 mm to 2 cm but ranges locally to 10 cm. Various normal graded sand layers have been observed, which indicate upright bedding. Also, disseminated in the muddy silt are ash beds or lenses (top of the unit, Cores 8 and 16), pyrite nodules, and calcareous concentrations (Cores 28-31). The size ranges for ash, pyrite, and calcareous concentrations are the same as for Unit 1. Gas expansion cracks are also common in Unit 2 down to Core 17. Low temperatures (1°C and 1.5°C) and gas bubbling were encountered in Core 19 at about 170 meters.

Bedding is manifested in this unit by thin, silty laminations and fine sand layers. The dips average 20° and range from 0° to 85° (mostly apparent dips); in individual cores variation is generally high. The variable dips with the presence of probable slump folds can be partly explained by original synsedimentary disturbance prior to tectonic deformation. Fracturing is present throughout the unit; however, slickensides are rare. Normal faults offsetting drilling laminations are clearly due to



Figure 2. Site location map and index. A-A' indicates segment of seismic profile Line MX-16 shown in Figure 3.

drilling disturbance and complicate analysis of the fractures. Nevertheless some fractures seem to be essentially tectonic features—for example, the normal faults in Cores 15, 48, 14, 15, 33, and 37 and reverse faults in Cores 22 and 43.

Unit 3, early Pliocene (Cores 49-59, 437.5-542 m sub-bottom depth), consists of muddy silt to muddy siltstone interbedded with fine to coarse pebbly sand. The thickness of the sandy layers ranges from a few centimeters up to 40 cm in Core 50. Sand is the dominant lithology in Cores 52 and 53 and probably accounts for the poor recovery in this interval (Fig. 7). Normal graded sandy beds were observed in Cores 49, 58, and 59, indicating lack of overturning. A worm cast in Core 54 also suggests upright bedding.

Structural deformation exists also in Unit 3 but is obscured by both drilling disturbance and relative abundance of sand. Nevertheless, the bedding has a mean value of 15° , ranging from 4° to 31° (mostly apparent dips); dips seem to be more uniform in individual cores than was the case in Unit 2. Fracturing and slickensides are more abundant than in Unit 2. Drilling faults are also present but account for only a small fraction of the deformation. A normal fault appears in Core 37.

Conclusions

Sedimentologic Data

The sedimentary sequence of Site 491 is very similar to that of Site 488, with the mean grain size increasing downhole, from mud through muddy silt, to coarse pebbly sand. However, the sediments at Site 488 are Quaternary, whereas those at Site 491 are mainly Pliocene. In this terrigenous sequence a genetic distinction can be made between the sands at the base (Unit 3) and the overlying muddy silt/mud (Units 1 and 2), as was determined for Site 488. Much of the interpretive discussion for site conclusions applies here. The sedimentologic arguments alone do not allow us to assign the sandy unit (Unit 3) with certainty to a former lower Pliocene trench environment; such coarse clastic sediments could have been deposited for example in a slope basin or channel cutting the slope. Definitive sedimentologic proof of origin of the sands as trench deposits would be the oc-



Figure 3. Migrated multichannel seismic reflection profile Line MX-16.

currence of pelagic clay lower in the sequence. Alternatively, we might expect to obtain an approximate paleodepth for the sandy unit of Site 491 from the microfossils below Core 4 (2 Ma ~ Pleistocene/Pliocene boundary): planktonic foraminifers disappear and nannofossils are much less abundant; this probably indicates passage through the CCD (at approximately 3200 m) at this time. Lower in the sequence, in Core 27, a trace fossil assemblage suggests a depth of 4 km or greater (see Paleobathymetry and Vertical Tectonics), which favors deposition either in the trench or in a lowermost basin or channel.

Structural Data

Bedding dips of more than 10° and true tectonic fractures appear in Core 14, at about 120 meters in Unit 2 (lower to upper Pliocene). At greater depth, deformation becomes more intense. Variation in bedding dip in Unit 2 is probably due to synsedimentary disturbance before tectonic deformation. No unconformity (or disconformity) indicating the superposition of two different deformed units has been observed in this sequence. On the contrary, the increase of fracturing density is progressive; moreover, the first deformation appears about 300 meters above the top of coarse pebbly sands of Unit 3, in the so-called "slope apron deposits" (muddy silt of Unit 2). As for Site 488, we detected no tectonic repetition by paleontologic data in the whole drilled sequence.

BIOSTRATIGRAPHY

Site 491 penetrated a midslope lower Pliocene through Quaternary sedimentary section of sand and hemipelagic mud (Fig. 1). Calcareous fossils are partially or completely dissolved below Core 4 (40 m), but radiolarian and diatom preservation is generally good throughout. Reworked lower Miocene, upper Miocene, and Pliocene microfossils occur throughout the section, commonly in upper Pliocene and Quaternary sediments.

Radiolarians

Moderate to well-preserved radiolarians occur in nearly all cores from Site 491, although abundance varies markedly. Radiolarians are abundant and well preserved in Cores 1 to 9 but rare and poorly preserved in Cores 10 to 59, hampering biostratigraphic dating. This change in preservation and abundance coincides with an increase in sedimentation rate within Core 8. Reworking of lower and upper Miocene and Pliocene species, which occurs throughout the section at 491, hinders age determinations.

The good radiolarian preservation and reduced reworking in the upper part of Site 491 permit the identification and location of Quaternary and upper Pliocene radiolaria datum planes. As in other Leg 66 sites, mixtures of North Pacific and equatorial Pacific forms occur. The last occurrence of *Axoprunum angelinum* (0.4 Ma) occur at the base of Core 1 (10 m), and the last occurrence of *Anthocyrtidium angulare* (1.2 Ma) occurs in the lower part of Core 3 (28 m).

The lower part of Core 4 through Core 7 is within the *Pterocanium prismatium* zone, as suggested by the occurrence of *Lamprocyclas neoheteroporos* (Hays, 1970) and the first occurrences of *Amphirhopalum ypsilon* and *Anthocyrtidium angulare*, although *P. prismatium* itself is absent. The top of the *Spongaster pentas* zone,



Figure 5. Very thin discontinuous silt beds of Unit 2, Sample 491-15-2, 130-145 cm.

Figure 6. Fine-grained sand bed (Sample 491-19-4, 40-60 cm) representative of thickest sand beds observed in Unit 2. Note reverse fault of small displacement.



Figure 7. Coarse sand from Unit 3.

indicated by common Stichocorys peregrina and the appearance of Spongaster pentas shortly below, occurs within Core 9. Cores 10 to 59 may be within the S. pentas zone, as S. pentas is sporadic and Ommatartus penultimus never becomes common (Riedel and Sanfilippo, 1971). The early Pliocene age of Cores 10 to 59 is supported by calcareous nannoplankton.

Foraminifers

At Site 491, Cores 1 through 59 penetrate a Quaternary through Pliocene sedimentary section. Planktonic foraminifers are few to common in Cores 1 through 3, rare to very rare in Cores 4 through 6, and absent in Cores 7 through 59. Benthic foraminifers show almost the same pattern of distribution in the samples. Some samples in Cores 7 through 59, however, contain a few dwarf shelf-derived planktonic and benthic foraminifers. The planktonic specimens are usually juvenile and give no age indication.

Cores 1 through 3 include the following Quaternary foraminifers: Globorotalia tumida, G. menardii, G. flexuosa, Globigerinoides ruber, G. sacculifer, G. triloba, Neogloboquadrina dutertrei, and Orbulina universa. Although no index Pliocene species are encountered, occurrence of Globorotalia cf. acostaensis and Neogloboquadrina humerosa in Cores 4 and 5 may suggest an upper Pliocene age for these cores.

Paleoecology and Depositional Environment

Absence of in situ planktonic and benthic foraminifers in most of the Pliocene section indicates deposition below the CCD and dilution due to terrigenous influx. The occurrence of some shallow-water foraminifers in the samples, such as Bolivina, Ammonia, and Gavelinella, suggests reworking of the shelf assemblage. This Pliocene faunal assemblage corresponds to the modern trench fill fauna at Site 488. In contrast to the Pliocene environments, the upper Quaternary benthic foraminifers indicate a shallower depth (lower bathyal) corresponding to the modern water depth at this site. They include Melonis pompilioides, Planulina wuellerstorfi, Hoeglundina elegans, Uvigerina proboscidea, U. hispida, Bulimina striata, and Pyrgo depressa. The Quaternary assemblage contains both well-preserved and partially corroded foraminifers, indicating deposition within the foram lysocline or local excursions of the CCD.

Calcareous Nannoplankton and Silicoflagellates

The nannofossil assemblages of the sediment column from Section 491-1,CC to 491-59,CC can be assigned to four different biostratigraphic zones (Fig. 1):

1) Middle to upper Quaternary—Nannoplankton Zone NN20 in 1,CC to 2,CC (0.2–0.6 Ma): *Gephyrocapsa* oceanica, G. caribbeanica, Helicosphaera carteri, Cycloccolithus leptoporus, and Ceratolithus cristatus.

2) Upper Pliocene—Nannoplankton Zone NN18 in Core 4 (1.8-2.5 Ma): Discoaster brouweri, Gephyrocapsa doronicoides, and Helicosphaera carteri. This sediment was deposited after the extinction of D. pentaradiatus.

3) Upper Pliocene—Nannoplankton Zone NN17 in Cores 5-7 (2.5-2.7 Ma): Discoaster pentaradiatus, D. brouweri, and Helicosphaera carteri. Between Cores 7 and 8 there might be a hiatus, because the NN16 nannoplankton zone *D. surculus* was not found.

4) Lower Pliocene—Nannoplankton Zone NN15 in Cores 8-59 (3.5 Ma-?): Reticulofenestra pseudoumbilica, Cyclococcolithus macintyrei, Discoaster pentaradiatus, and Sphenolithus neoabies. Because of unfavorable deposition within a mostly siliceous sediment, the nannofossil assemblages from Core 9 down are impoverished, with only the more sturdy species preserved. As all the listed species of this zone have a long range from the Miocene up, no subdivision of this core interval can be given. No species restricted to the Miocene are found. Therefore a lower Pliocene age is the most probable one for the reduced nannofossil assemblages from Core 8 down. Ages are those used by Perch-Nielsen (1977).

Silicoflagellate assemblages were studied in Sections 491-1,CC and 491-4,CC. They constitute only a minor part of the siliceous sediment fraction consisting of mainly diatoms and radiolaria.

1) Core Catcher 1—middle to upper Quaternary, NN20: Dictyocha stapedia stapedia and D. fibula, 97%; Distephanus octogonus, 3%. This is a typical upper Quaternary assemblage, which fits into the D. epiodon Zone of Bukry and Foster (1973) as well as into the D. octangulatus Zone of Bukry (1973) for the eastern and eastern north Pacific or the Dictyocha aculeata aculeata Zone of Bukry (1978), respectively.

2) Core Catcher 4—upper Pliocene, NN18: Dictyocha fibula s.l., 96%; Distephanus speculum s.l., 4%. No Mesocena elliptica, or D. octogonus is found to indicate a date older than Quaternary which suggests that the associated nannofossils are not reworked. The occurrence of Actiniscus tetrasterias also gives evidence of an age near the Pliocene/Pleistocene boundary.

SEDIMENT ACCUMULATION RATES

The sediment accumulation rate curve (Fig. 8) is based on biostratigraphic datum points and paleomagnetic datum levels; the exact method used to construct the curve is outlined in the introductory chapter. No corrections have been made for compaction or deformation. Sediment accumulation rates in the upper part of Site 491 vary between 34.2 m/m.y. and 5 m/m.y. Below 70 meters, the minimum rate is 1055 m/m.y., if all sediment below 81 meters accumulated within the *Spongaster pentas* and NN15 zones. The higher accumulation rate in the lower part of Site 491 may result partly from tectonic thickening of the section which occurred at the lower slope (see Paleobathymetry) as well as from higher initial sedimentation rates. After 3 Ma, slope sediments were deposited at a much slower rate.

PALEOBATHYMETRY AND VERTICAL TECTONICS

Sedimentologic and paleontologic data permit us to construct a paleobathymetric curve for Site 491 (Fig. 9). The three points used are the present water depths, the point where sediments of Site 491 pass through the CCD at 3.2 km at 2 Ma (see McMillen, this volume); the basal sandy sediments, which were probably deposited on the



Figure 8. Sediment accumulation rate, Site 491. Rate uncorrected for compaction or tectonic tilt.

lower slope or trench at 5 km about 3.7 Ma; and the *Teichichnus-Zoophycos-Chondrites* trace fossil assemblage occurring at 4 km or greater water depth in Core 27.

These three points define the paleobathymetric curve for Site 491 (Fig. 8). The most striking feature of the curve is that the uplift began at a high rate of about 800 m/m.y. and slowed to a rate of about 135 m/m.y. from 2 Ma onward. The higher rate prior to 2 Ma probably represents the emplacement of lower slope or trench sediment into the zone of accretion, and the slower rate after 2 Ma probably represents regional uplift of the slope due to sediment underplating of the accretionary zone.

PALEOMAGNETISM

Paleomagnetic analyses at Site 491 established magnetostratigraphy and determined dips of certain beds and faults. Sediments in Site 491 can be correlated mainly with the Gilbert reversed polarity epoch. Paleomagnetically oriented cores dip generally northward.



Figure 9. Paleobathymetry (dashed line) and vertical tectonics (solid line) for Site 491. Control points: CCD is the location of the carbonate compensation depth at 2 m.y., TF is the location of the *Zoophycos-Chondrites-Teichichnus* trace fossil assemblage indicative of water depths 4 km or greater, and SF is the location of the sandy facies interpreted as uplifted trench deposits. The tectonic curve is derived from the bathymetric curve by removing the sediment accumulation at any given time.

Cores from the upper 50 meters are disturbed by drilling showing laminations concave downward along core margins. Cores deeper than 220 meters in the hole are separated into pieces of several centimeters in thickness which are rotated relative to one another. Several cores from the middle of the hole have irregular bedding, possibly caused by slumping.

Using a plastic tube in the upper 480 meters and a minicore drill below 480 meters 67, oriented samples were collected from less disturbed portions of cores with regular bedding direction. Stability of remanent magnetization of selected samples was examined with stepwise AF demagnetization (Niitsuma, this volume). The magnetization of the samples shows stable orientation and a slow decrease in intensity with AF demagnetization. The noise level of magnetometer is $10^{-7.6 \pm 0.5}$ emu/cc. Structural corrections have been applied to the inclination values. Average intensity is $10^{-5.8} \pm 0.7$ emu/ cc after 15 mT AF demagnetization, although the intensity of upper 70 meters of the hole is one order weaker $(10^{-6.7 \pm 0.7} \text{ emu/cc})$. The sudden downhole increase in intensity at 70 meters coincided with lithologic change from calcareous nodule-bearing mud to mud with fine sand patches. Inclinations in the upper 140 meters of the hole changes sign frequently. Cores below 140 meters have mainly negative inclination except for three intervals of significant positive inclination. We believe that the Brunhes/Matuyama boundary occurs at about 24 meters, the Matuyama/Gauss boundary at about 53 meters, and the Gauss/Gilbert boundary at about 110 meters (Fig. 1). Below 100 meters the sediments probably are entirely in the Gilbert polarity epoch.

Sediment at Site 491 dip from 4° to 64°. Since the drilling core axis is nearly vertical, orientation of bedding plane can be calculated from magnetic inclination and declination (Niitsuma, this volume).

Dip direction of bedding plane are generally northnortheast, which is the opposite direction of submarine topographic slope. Clockwise direction of change of dips appears cyclically from westward to eastward depth (Niitsuma, this volume). Northward dip direction is accompanied with steeper dip angle and westward dip direction with shallower dip angle. This kind of regularity and cyclicity in dip angle and direction suggests either slump folding or small-scale tectonic deformation.

Several conjugate faults sets were observed in the lower portion of the cored section at Site 491 (Niitsuma, this volume). Paleomagnetic orientation of these faults suggests that the tensional axis is horizontal and almost parallel to the bedding dip direction.

ORGANIC GEOCHEMISTRY

The shipboard organic geochemistry monitoring program consisted of analysis of gases released in core liners, determination of organic carbon, hydrogen, nitrogen, and carbonate content of selected sediment samples, and visual inspection for fluorescence in split core.

Gases

Moderate amounts of gas were released in core liners from depths of about 20 meters and greater. The gas initially contained methane, CO_2 , and small amounts of H_2S . The last, detectable by its distinctive odor, was present down to depths of about 60 meters. Methane content remained fairly constant with depth (Fig. 10) except for minima in the vicinity of 100, 175, 225, 350, and 450 meters.

Shallower than about 50 meters, ethane content was below the detection limit of the Carle gas chromatograph but increased gradually with depth and maintained a concentration of about 0.03% by volume except at the methane minima mentioned in the foregoing. The correspondence of low values in ethane + methane



Figure 10. Gas concentrations, Site 491.

suggest air dilution, probably due to washout. The ethane to methane ratio (Fig. 10) seldom varied from a value of about 3×10^{-4} throughout the cored section.

 CO_2 content in core liner gases varied from 0.06% to 2.04%, with both the highest and lowest concentrations occurring in the upper portion of the cored sequence.

Hydrocarbons in the C_3 to C_5 range were monitored on the Hewlett-Packard 5710-A gas chromatograph from a depth of 75 meters to TD. Their abundance was found to vary erratically with depth, showing a maximum of 381 ppm at 214.7 meters (Fig. 10).

Upon splitting Cores 10 and 19, gas-releasing ice inclusions or frozen sediment were observed from depths of 88.9 to 89.0 meters, 162.7 to 163.15 meters, and 168.3 to 168.35 meters, respectively (Figs. 1 and 11). One ice inclusion at 89 meters released gas equivalent to 7.2 times its volume, suggesting the presence of hydrate. The evolved gases consisted of methane and CO_2 , with no ethane detectable on the Carle gas chromatograph.

Fluorescence

Split cores showed no evidence of fluorescence due to crude oil or bitumen impregnation.



Figure 11. Probable hydrate-bearing ice inclusion at 89 meters (Sample 491-10-2, 145-147 cm).

Organic Carbon, Hydrogen, Nitrogen, and Carbonate

Samples for CHN and carbonate analysis were taken from selected cores. Within the cored sequence, the organic carbon content varied from 0.89% to 1.98% and the total nitrogen content from 0.08% to 0.17%. The organic potential of the sediments remained relatively constant and at an intermediate level throughout the hole.

The C/N ratio varied from 10.0 to 14.9, remaining approximately constant throughout the hole. This is in the range for organic matter associated with recent sediments (Fairbridge, 1972) and suggests that the organic matter present in these sediments has a low degree of thermal maturation. Low concentration of carbonates were detected, particularly in the lower portion of the hole.

Conclusions

Gases, mainly of biogenic origin, were detected throughout this hole, causing a low to moderate degassing of the cores. The organic potential of the sediments stayed at an intermediate level through the hole, being marginally higher for the upper section. The C/N ratio of the organic matter suggests a low degree of geothermal maturation. No evidence of petroleum or bitumen impregnation was detected.

Heavier hydrocarbons C_3 to C_5 varied erratically with depth without evidence of an increase in geothermal maturation. The origin of these gases is not known. They may originate in geothermally more matured sediments and be emplaced in shallower depths by migration, but a biogenic origin cannot be excluded.

The gas-generating ratio of one frozen sample at 88.9 to 89 meters suggests the presence of gas hydrate at Site 491.

PHYSICAL PROPERTIES

Physical property analyses of Site 491 sediments included porosity, water content, wet bulk density, and undrained shear strength (Fig. 12) (see Boyce, 1976, for procedures). Gas attenuation prevented compressional sound velocity measurements. Poor core recovery and increased core disturbance limited sampling density below 250 meters sub-bottom. The lack of logging precluded correlation with *in situ* properties.

Uniformly gradational changes in physical properties occur to 410 meters sub-bottom. Below 410 meters, variations in physical properties and the necessity to increase pump pressure to maintain drilling rates suggest increased induration. This increase may result from mass movement and/or tectonic processes or may reflect the influence of more permeable sand beds below which facilitated dewaterting of the interbedded muds at the time of deposition.

Porosity, Water Content, and Bulk Density

Porosity decreases uniformly from 66% at 0.50 meters to 33% at 525 meters (Fig. 12). Water content de-

creases from 44% at 0.50 meters to 15% at 525 meters. Both porosity and water content decrease more rapidly below 410 meters. Bulk density increases from 1.58 Mg/ m^3 to 2.07 Mg/m³ from 6.60 to 520 meters, respectively (Fig. 12).

Shear Strength

Shear strength increases regularly from 13.5 kPa at 0.5 meters to 122.5 kPa at 69 meters (Fig. 12). Good correlation exists between vane shear and torvane results.

INHOLE TEMPERATURE MEASUREMENTS

Of three downhole temperature measurements at Site 491 with the Uyeda Probe, none yielded typical temperature curves. The measurement at 304.5 meters yielded a bottomhole temperature of 14.5°C and a gradient of 2.5°C/100 m. The measurement at 304.5 meters yields a thermal gradient which is about twice as high as predicted from methane hydrate-phase relationships (Shipley et al., 1979) and three times conventional heat flow work of Langseth and Von Herzen (1970). The postulated relationship between a prominent reflection and the base of gas hydrates of Shipley et al. (1979) is based on the pressure and temperature of hydrate stability. If the gradient is as high as 4°C/100 m, then the base of the hydrate stability occurs much shallower than the depth to the so-called "bottom-simulating reflection" (BSR). See Shipley and Shephard (this volume) for a complete discussion of the temperature data.

CORRELATION OF SEISMIC REFLECTION DATA AND DRILLING RESULTS

Site 491 coincides with a zone of landward-dipping reflectors about halfway up the inner slope of the trench between Sites 488 and 490. A wave-equation-migrated multichannel profile about 1000 meters northwest of Site 491 shows a zone of discontinuous and indistinct reflections, extending to perhaps a 0.4 s sub-bottom and clearly separate from the deeper landward-dipping, more coherent reflectors (Fig. 13). The actual boundary between the two zones is not well defined (Fig. 14). A high-amplitude reflection at 0.7 s sub-bottom may correlate with the base of the gas hydrate layer (Shipley et al., 1979).

The multichannel seismic data reveal a fairly high velocity (2.2 km/s) for the interval from the seafloor to the gas hydrate(?) reflection. Using this velocity, the total drilled section of 542 meters corresponds to at least 0.49 s sub-bottom (Fig. 13). The base of the discontinuous and indistinct reflectors occurs at about 440 meters sub-bottom, coinciding with the top of Lithologic Unit 3, which is defined by the first appearance of significant sand beds downhole. However, thick sand beds, which are probable reflectors, do not occur above 470 meters. The higher velocity for the sediments, compared to Site 488, is manifested in both the greater age and fissility of these rocks versus any significant lithologic changes.



Figure 12. Physical properties summary profiles, Site 491.

SUMMARY AND CONCLUSIONS

At Site 491 we penetrated 542 meters and recovered 59 cores comprising three lithologic units. Unit 1 extends from 0 to 57.5 meters and consists of upper Pliocene to upper Quaternary mud and muddy silt with minor concentrations of ash and calcareous mud. Lower Pliocene muddy silt and muddy siltstone with minor fine sand layers constitutes Unit 2, cored between 57.5 and 437.5 meters. The most significant lithologic break occurs with Unit 3, which extends from 437.5 to 542 meters and is composed of lower Pliocene muddy silt and muddy siltstone with interbedded fine to coarse pebbly sand layers up to 40 cm thick.

Near-surface cores at Site 491 include a diverse microfossil assemblage; however, below 40 meters the calcareous components are partially or completely dissolved. Diatom and radiolarian preservation is good throughout the section, with radiolarians providing the best age control for deeper cores. Sediment accumulation rates, uncorrected for structure or compaction, are low (5-34 m/m.y.) above 70 meters but increase

substantially (to about 1000 m/m.y.) below this depth to the base of the hole.

At Site 491 tilted beds and fracturing occur initially at about 120 meters and continue to total depth through the lower two lithologic units. In Unit 2 dips are variable and range to nearly vertical; fracturing is present throughout though slickensides are rare. Dip angles in Unit 3 range up to 30° and are more uniform than in Unit 2. A strong fissility develops locally. Paleomagnetic restorations of bedding in both Units 2 and 3 indicate modal dip directions to the north, though with considerable scatter. A separate deformational history for Unit 3 is suggested by discrete jumps in bulk density and porosity below 425 meters, though this variation may be simply the result of lithologic changes.

The multichannel seismic reflection profile through Site 491 shows a series of discontinuous and indistinct reflectors above about 0.4 s, overlying more coherent landward-dipping reflectors. At an interval velocity of 2.2 km/s the base of the discontinuous and indistinct reflections occurs at 440 meters near the top of Unit 3, which is defined by the first occurrence of sand. The



Figure 13. Detail of seismic profile Line MX-16.



Figure 14. Interpretive line drawing of Figure 13.

thick sandbeds, the probable reflectors, are not apparent until 470 meters sub-bottom. The mean low, northerly dip of bedding in Unit 3 suggested by the paleomagnetic restorations is subparallel to the dips of the seismic reflectors.

The depositional environment of Site 491 sediment is most directly interpreted by reference to the modern sedimentary regime as well as to the Quaternary sequence of Site 488. Coarse sands similar to those cored in Unit 3 at Site 491 occur in the modern trench and presumably in the large submarine canyon feeding the trench but have not been recovered in piston cores from the slope (McMillen and Haines, this volume). As such, we prefer to interpret the sand-bearing Unit 3 as trench sediment, since the profiles through the site show no evidence of buried canyons (Shipley, this volume). Alternatively, the coarse sands may have accumulated in a slope basin, though the seismic data show no evidence of remnants of such a basin. Much of the muddy silt of Unit 2 at Site 491 accumulated below the CCD at sedimentation rates comparable to those at Site 488. Thus Unit 2 sediments were probably deposited as lower slope hemipelagic muds.

Deformation occurs primarily in sands and muddy silts accumulated at high sedimentation rates, probably in the trench and or lowermost slope environments. This deformed sequence is overlain by about 120 meters of largely undisturbed sediment, most of which accumulated at a low rate at or above the CCD. Thus the structural history of this site indicates deformation in the trench and/or lower slope environment associated with rapid uplift from about 4 to 3 Ma, followed by an interval of slow deformation and uplift. Apparently the toe of the trench slope is a zone of concentrated tectonism that diminishes in intensity upslope.

Hydrocarbon gases from C_1 to C_5 were detected throughout the hole, causing low to moderate degassing of the cores. Gas-releasing frozen sediments were observed in the 88.9 to 89.0 meter, 162.7 to 163.15 meter, and 168.3 to 168.35 meter intervals. The gas-generating ratio (7.2) of one of the samples (88.9–89 m) suggests the presence of hydrated sediments. Failure of the pressure core barrel prevented an attempt to sample the gas hydrates *in situ*.

REFERENCES

- Boyce, R. E., 1976. Definitions and laboratory techniques of compressional sound velocity parameters and wet water content, wet bulk density and porosity parameters by gravimetric and gamma ray attenuation techniques. *In* Schlanger, S. O., Jackson, E. D., et al., *Init. Repts. DSDP*, 33: Washington (U.S. Govt. Printing Office), 931-958.
- Bukry, D., 1973. Coccolith and silicoflagellate stratigraphy, Leg 18, Deep Sea Drilling Project, Eastern North Pacific. In Kulm, L. D., von Huene, R., et al., Init. Repts. DSDP, 18: Washington (U.S. Govt. Printing Office), 817-832.
- _____, 1978. Cenozoic coccolith, silicoflagellate, and diatom stratigraphy, Deep Sea Drilling Project, Leg 44. *In* Benson, W. E., Sheridan, R. E., et al., *Init. Repts. DSDP*, 44: Washington (U.S. Govt. Printing Office), 807-863.
- Bukry, D., and Foster, J. H., 1973. Silicofagellate and diatom stratigraphy, Leg 16, Deep Sea Drilling Project. *In* van Andel, Tj. H., Heath, G. R., et al., *Init. Repts. DSDP*, 16: Washington (U.S. Govt. Printing Office), 815-872.
- Fairbridge, R. W., 1972. The Encyclopedia of Geochemistry and Environmental Sciences: New York (Von Nostrand), pp. 136-141.
- Hays, J. D., 1970. Stratigraphy and evolutionary trends of radiolarian in North Pacific deep-sea sediments. *Mem. Geol. Soc. Am.*, 126: 185-218.
- Hein, J. R., Scholl, D. W., Barron, J. A., et al., 1978. Diagenesis of late Cenozoic diatomaceous deposits and formation of the bottom simulating reflector of the southern Bering Sea. Sedimentology, 25, 155-181.
- Karig, D. E., 1974. tectonic erosion at trenches. Earth Planet. Sci. Lett., 21:209-121.
- Karig, D. E., Cardwell, R. K., Moore, G. F., et al., 1978. Late Cenozoic subduction and continental margin truncation along the northern Middle America Trench. *Geol. Soc. Am. Bull.*, 89:265– 276.
- Langseth, M. G., and Von Herzen, R. P., 1970. Heat flow through the floor of the world oceans. *In Maxwell*, A. E. (Ed.), *The Sea* (Vol. 4): New York (Wiley-Interscience), 299–352.
- Perch-Nielsen, K., 1977. Albian to Pleistocene calcareous nannofossils from the western South Atlantic, Deep Sea Drilling Project. In Supko, P. R., Perch-Nielsen, K., et al., Init. Repts. DSDP, 39: Washington (U.S. Govt. Printing Office), 699-823.
- Riedel, W. R., and Sanfilippo, A., 1971. Cenozoic radiolaria from the western tropical Pacific, Leg 7. *In* Winterer, E. L., Riedel, W. R., et al., *Init. Repts. DSDP*, 7, Pt. 2: Washington (U.S. Govt. Printing Office), 1529–1672.
- Seely, D. R., Vail, P. R., and Walton, G. C., 1974. Trench slope model. In Burk, C. A., and Drake, C. L. (Eds.), The Geology of Continental Margins: New York (Springer Verlag), pp. 249-260.
- Shipley, T. H., Houston, M. H., Buffler, R. T., et al., 1979. Seismic reflection evidence for widespread occurrence of possible gashydrate horizons on continental slopes and rises. *Bull. Am. Assoc. Pet. Geol.*, 63:2204–2213.

| SITE | 491 | н | OLE | | (| OR | E 1 | CORED | INTERVA | L 0.0-10.0 m | SITE | 4 | 91 | HOL | E | CORE | 2 CORED | INTE | VAL 9.5-19.0 m | |
|---------------------|--------------------------|--------------|---------------------------|-----|------------|--|--------|--------------------------------------|---|---|---------------------|------------------|-------|--------------|----------------------------|-------------------------------------|----------------------|----------------------------|---|---|
| TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS | RADIOLARIANS HADIOLARIANS | SIL | TO A LOCAL | actillus | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES | LITHOLOGIC DESCRIPTION | TIME - ROCK UNIT | BIOSTRATIGRAPHIC | ZONE | HANNOFOSSILS | SSIL RACTER SWOINAND | SECTION METERS | GRAPHIC LITHOLOGY | DISTURBANCE SEDIMENTARY | SI LITHOL SI LITHOL | OGIC DESCRIPTION |
| QUATERNARY | NN 20 | RP F | AC | | | 0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 | 5 | V010 V010 V010 V010 V010 | | One on diffuse dusky vellow green (SGY 5/2) sith spot MUDDY SILT, soft, grayish olive green (SGY 3/2) with spotadic white sponge specks. SMEAR SLIDES § § | QUATERNARY | Inter Pol | D7 MM | RP | AG | 0.5 1 1.0 2 2 3 3 | | | GZ GZ GZ GZ GZ GZ GZ GZ GZ GZ GZ GZ GZ G | Y SILT, soft, grayish olive green (SGY 3/2) moder- formation, sponge spicules sponade. Silty zonest are iv consentrations in burrows. Black species of common. R SLIDES |











| SITE | 491 | | HOL | E | | CC | DRE | 12 CORED | INTER | VAL | 100.5105.0 m | |
|--------------------|-------------------------|--------------|--------------|--------------|---------|---------|--------|----------------------|--|---------|--|--|
| | PHIC | | CHA | OSS | TER | | | | | | | |
| TIME - ROCH | BIOSTRATIGRA ZONE | FORAMINIFERS | MANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | LITHOLOGIC DESC | RIPTION |
| PLIOCENE | NN 15 | RP | FP | CG | | | | SMEARS | | | Smears on core liner | . Deviation of the hole: 1.0°. |
| SITE | 491 | | но | .E | | cc | DRE | 13 CORED | INTER | VAL | 105.0–114.5 m | |
| × | THIC | | CHA | OSS RAC | TER | | | | | | | |
| TIME - ROC UNIT | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | LITHOLOGIC DESC | RIPTION |
| LOWER PLIOCENE | NN 15 Spongaster pentas | В | В | RG | | 1 | 0.5 | | | | MUDDY SILT, gra tling, minor conc layer, Gas expandi drilling deformation SMEAR SLIDES Diffuse fine sand layer TEXTURE: Sand Sit Clay COMPOSITION: Quartz Feldspar Mica Heavy minorals Pyrite Clay Correlation Heavy minorals Pyrite Clay Engraminates Namofosils Radiolarians Diatoms Sopone spicules | vish olive green (SGY 3/2); slight m Intrations of pyrite and fine SA on cracks and voids common, Inte |



| 10 | 2 | Г | F | OSS | SIL. | | T | | CORED | T | | ŕ | Ť | | | | - | |
|----------------|------------------|--------------|-------------|-------------|---------|----|---------|---------------------------|----------------------|----------|---------------------------|----------|---------|--|---|---|--|--|
| 5 | APH | | CHA | RA | СТІ | ER | | 10 | | | | | | | | | | |
| UNIT UNIT | BIOSTRATIGE | FORAMINIFERS | NANNOFOSSIL | RADIOLARIAN | DIATOME | | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY STRUCTURES | CAMOR DE | NAMPLES | N | LITHOLOGIC DESC | RIPTIO | N | |
| | | | | | | | 1 | 0.5 | | | , | | | Fault plane Fault plane | MUDDY SILT, gra mottled; some fine One ASH bed and olive gray (5Y 3 biscuits in Section I SMEAR SLIDES | vish oliv r sand la MUD, gr /2), ver 6, | e green minatic rayish o y firm | (5GY 3/2), firm, slightly ans (1 or 2 per section). live green (5GY 3/2) and mottled; some drilling |
| | | | | | | | 2 | and the second second | | | 1 | | | Fault plane | TEXTURE: Sand | Appny 1-10 4 | pnW 6-113 (M) | euij - 124 (M) |
| | as? | | | | | | | - Toronto | | | 1 | T | T | Zeolites (yellow patch) Fault plane | Silt Clay COMPOSITION: Quartz Feldspar Mica | 40 45 7 2 | 44 55 38 7 2 | 30 10 71 12 2 |
| LOWER PLIOCENE | Spongaster penti | | | | | | 3 | confictor from | | | | | - | Sandy lens Zeolite | Heavy minerals Pyrite Clay Zeolite Carb. unspec. Radiolarians Diatoms Sponge spicules | 1 35 3 TR 3 3 | TR 1 50 1 2 TR TR 1 | 2 1 10 - 2 TR TR TR |
| | NN 15 | | | | | | 4 | and market | | | | 0 | 3Z | Possible fault | GRAIN SIZE Sand Silt Clay | 4-70 11.8 83.1 25.1 | | |
| | | | | | | | 5 | the set of a set of a set | | | | | | Drilling biscuits? Zeolite | | | | |
| | | | | | | | 6 | | | | | | | ASH BED: apparent dip | 1; 40" | | | |
| | | в | RP | R | G | | cc | | VOID | | | | | | | | | |



| SITE | 491 | н | LE | | CO | RE | 18 CORED | NTER | VAL | 152.5-162.0 m | | | SITE | 491 | HOL | E | | CORE | 19 | CORED I | NTERVA | L 162.0-171.5 m | | |
|---------------------|--------------------------|--------------|--|---|----------------------------------|--------|----------------------|--|---------|---|---|--|---------------------|--------------------------|--------------|----------------------|---|-------------|----|----------------------|---|---|--|--|
| TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS | FOSSIL ARACTE SWULUIG SWULUIG | R | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | LITH | IOLOGIC DESCRIPTION | | TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | NANNOFOSSILS | DIATOMS RADIOLARIANS | 8 | SECTION | L | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE STRUGTURES STRUGTURES | | LITHOLOGIC DESCRIPTION | N |
| LOWER PLIOCENE | NN 15 Sporgaster peritas | 8 8 | RM | | 1 2 3 4 5 6 CC | 1.0 | | / III / IIII / III / IIII / III / II | • | Development of polished surfaces; slickenuides in soft mud; fractures 0-80° MUU (5Y Som Som SME Fine sand layer (deformed by drilling) Pyrite Fine sand layer (deformed by drilling) Pyrite TEX Sand Sit Clay COM Qua Pyrite TEX Sand Sit Clay COM Qua Fid Som Diffuse sand Sport dip: 12°1 Clay Com Diffuse sand Sport dip: 12°1 Clay Com Diffuse sand Sport dip: 27° Diffuse sand Section 1 Apparent dip 26° Apparent dip 27° Diffuse sand Sicks emud, possible Tad sequence. Time to very line sand, sporent dip 40° (possibly due to drilling) | JUDDY SILT, gravith oilve great Y 3/21; becomes randier from me very fine to fine SAND. IEAR SLIDES V IEAR SLIDES IEAR SLIDES I | (5GY 3/2) to olive grav top to base (10–25%). | LOWER PLIOCENE | Spongaster pentas | B FM | RG | | 4 5 6 | | | G • • | Very fine sand, 10GY 5: Very fine sand, 1°C jas babbling off Sequences Fine sand Fine sand, apparent dip 11* Sequences Apparent dip 11* Sequence Disturbed sand beds Sequence; 12* apparent Disturbed bedding | 2 MUDDY SILT, grayish ol SAND, more abundant an cores, sequences observe Sequences = | ive green with very line to line d in thicker layers than previous 5. Only traces in Core-Catcher. SGY 3/41 over sand with sharp graditional contact, base of sand sharp contact. bridd k mucht sit, slightly laminated grify diffuse Sili frame your with the same state of the same sta |

| SITE | 491 | но | LE | (| ORE | 20 | CORED | INTER | VAL | 171.5–181.0 m | | | | SITE | 491 | | IOLE | | C | ORE | 21 CO | RED INT | ERV | AL 181.0-190.5 m | | | |
|---------------------|--------------------------|--------------|------------------------------------|---|---------------------------|----|---------------------|----------------------------|---------|--|--|--|--|---------------------|--------------------------|--------------|------------|-----|------------------------|--------|-------|-----------|--------------|--|---|---|--------|
| TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS | FOSSIL ARACTE SNUINANDIOIDAN | R | METERS | L | GRAPHIC ITHOLOGY | DISTURBANCE SEDIMENTARY | SAMPLES | | LITHOLOGIC DESCRI | PTION | | TIME - ROCK UNIT | BIDSTRATIGRAPHIC ZONE | FORAMINIFERS | FORA SILES | SIL | SECTION | METERS | GRAPH | DIRILLING | SEDIMENTARY | SAMPLES | LITHOLOGIC DESCR | PTION | |
| LOWER PLIOCENE | Sponguster pentas? | 6 8 | RG | | 0.5 1 1.0 2 3 | | | 0 0 | GZ | Fine sand Fractured, polished and slickensided surfaces in soft muddy slit Bedding defined by color shade change 97 separent dip 28° true dip Highly contorted Slump fold? Highly contorted VOID 5 mm medium sand 10° apparent dip | MUDDY SILT, of medium SAND. Section 2: SMEAR SLIDES SMEAR SLIDES TEXTURE: Sand Silt Clay COMPOSITION: Cuart Radiolarians Diatoms Sponge spicules GRAIN SIZE Sand Silt Clay Radiolarians Diatoms | ve grav (6 Only trace of deformed. Jo cm " " w No 15 cm " Approved. 15 cm " Approved. 128 1 (0) 5 2 4 3 6 7 1 1 7 1 1 7 1 1 7 1 1 7 1 1 7 1 1 7 1 1 7 1 1 7 7 1 1 7 7 1 7 7 1 7 | SY 3/2) with minor fine to is in Core-Catcher. Whole , upday and darker unday sit where contact not shown (free boundary b) 1130 D) 15 15 15 15 11 15 15 11 15 17 11 11 15 17 11 | LOWER PLIOCENE | NN 15 Spongester pentes? | в | RP FA | 10 | 1 2 3 4 cc | 0.5 | | | 1 - 22 2 2 M | Near horizontal fract slickensides and polished surfaces 24" true dip sand pods 24" true dip sand pods Darker muddy silt Overturned fold Toreard of olded sand pods Sand; apparent dip 51" Chaotic discontinuous sand beds 20" apparent dip Fine sand; 22" appar dip, diffuse top, sharp base Parallel laminations | Aures, MUDDY SILT, gray to medium SAND. Section 3, 43–48 cm Section 3, 43–48 cm Section 3, 107–109 Section 3 | ish olive green (5GY 3/2) with Sin Durk modify sit: Tand Modey sit: and Burriows with sand failing 0 20 45 5 1 1 45 5 1 1 45 5 1 1 45 5 1 1 45 5 1 1 45 5 1 1 45 5 1 1 45 5 5 1 1 45 5 5 1 1 45 5 5 1 1 45 5 5 1 1 45 5 5 1 1 45 5 5 1 1 1 45 5 5 1 1 7 7 7 7 7 7 7 7 7 7 7 7 7 | i fine |

ORGANIC CARBON AND CARBONATE 1-84 % Organic Carbon 1.1 % CuCO₃ 0.0







| SITE 4 | 91 HOLE | _ | CORE | 2 | 7 CORED | INTER | AL 228.5-238.0 m | | | | SITE | 91 | HOL | E | _ | COR | E | 28 CORED | INTER | VAL | 238.0-247.5 m | | | | |
|---|---|---------------|-----------------------|---|----------------------|--|------------------|---|--|------------------------------------|---------------------|-----------------------------|--------------|---------|---|---------------------------------|--------|----------------------|--|---------|---|--|---|---|---|
| TIME - ROCK UNIT BIOSTRATIGRAPHIC | ZONE FORAMINIFERS NANNOFOSSILS BADIOLARILANS BS00 | SIL CTER - | SECTION | | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | Sameles | LITHOLOGIC DESC | SRIPTION | | TIME - ROCK UNIT | ZONE FORAMINIFIERS | HANNOFOSSILS | DIATOMS | R | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STRUCTURES | SAMPLES | | LITHOLOGIC DES | CRIPTIO | N | |
| LOWER PLIOCENE | spongarate jeer tas i | | 1 1 2 3 4 | | | 00 ** | GZ * | MUDDY SILT, or drilling lamination what swirling, co drilling lamination 40–50 cm; conjug laminations. Modes all prevents of the prevents of | anish olive green (SGY 3/2), firm to h is abundant. Section 3: abstrue, can her variations, truncated by camp ate set of normal faults offsetting drill Section 3: | ard mo- non 4: 10 1 | LOWER PLIOCENE | Spongaster pentas 7 1000 | RP | FM | | 3 3 6 6 6 6 7 | | PP | .4 | GZ | Dark yellowish orange concentration of carbo | 10YR 8/6) INTE grains MUDDY SILT, 6 (1–2 mm) drilli every 2–4 cm, s slight blotubation SMEAR SLIDES TEXTURE: Sand Silt Clay COMPOSITION: Quetz Peldipar Mica Heavy minerals Pyrite Clay Cl | rayish oil g tamina structure structure the structure st | ve green (50 ions of dar m. Occation 3.38 (D) 1 80 39 57 15 3 1 1 20 1 2 20 1 2 ТП ТП ТП ТП | SY 3/2], with this < gray (N3) MUD al sponge remains ponge remains sponge |
| | B FP R | G | 6 7 000 | | | 0 | | ORGANIC CARI | 80N AND CARBONATE 340 n 0.9 0.0 | | | | | | | | | | | | | | | | |



| SITE | 491 | но | LE | _ | co | RE | 31 CORED | INTER | IAVE | 266.5-276.0 m | | SITE | 49 | н | OLE | | CC | RE | 32 CORED | INTER | VAL | 276.0-285.5 m | |
|---------------------|--------------------------|--------------|--------|----|----------------------------|--------|----------------------|--|-----------------------|---|---|---------------------|--------------------------|--------------|-----------------|---------|--------------------------------------|--------|----------------------|---|---------|---|--|
| TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS | FOSSIL | ER | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY | STRUCTURES SAMPLES | | LITHOLOGIC DESCRIPTION | TIME - ROCK UNIT | BIOSTRATIGRAPHIC ZONE | FORAMINIFERS | RADIOLARIANS 25 | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING DISTURBANCE SEDIMENTARY STOLINTLIDE | SAMPLES | | LITHOLOGIC DESCRIPTION |
| LOWER PLIOCENE | Spongaster pentas? | 8 | 1 RM | | 1 2 3 4 5 6 | | PP | | • | Thin discontinuous calcaneous layer Thin subhorizontal pod of fine sand Drilling lamination Three thin (1 mm) layers of fine and 24" apparent dip Drilling lamination | MUDDY SILT, gravish olive green (GGY 3/2), soft to firm. Sections 1 and 2: broken up by drilling, infrequent small (up to 2 mm) pryite notokes. Section 3: firm, drilling laminations throughout. Section 5: slightly bioturbated. SMEAR SLIDES software gravity of the state of the state of the state of the gravity of the state of the state of the state of the State of the state of the state of the state of the state of the State of the state of the state of the state of the state of the State of the state of the | LOWER PLIOCENE | Sportgaster peritas | 8 | 6 RM | | 1 2 3 4 5 6 6 6 | 0.5 | | 0 0 0 | GZ | Eleven fracture: over 80 cm (drilling? tectonic?) Sixteen major fracture over 110 cm of stiffer material Drilling lamination Drilling lamination There and to 16' true dip Drilling laminations c | MUDDY SILT, fim to very firm, gravish olive green (SGY 3/2): appears as discrete blocks in soupy motdy silt. Section 4: soft to firm, slightly bioturbated, still very broken up. SME AR SLIDES (D) TEXTURE: Sand 3 Sith 59 Clay 38 COMPOSITION: Ourrz 7 Felsgar 15 Mica 8 Heavy minarals 2 Pyrite 2 Clay 20 10 Clas 15 Mica 8 Heavy minarals 2 Pyrite 2 Clay 20 10 Clas 17 Radiolaratint TR Diatoms 2 Sponge tpicules 1 Plant fragments TR GRAIN SIZE Sand 8.2 Sith 68.7 Clay 23.1 ver |

| SITE 491 HOLE | CORE 33 CORED INTERVAL | 285.5-295.0 m | SITE 491 HOLE CORE 34 CORED INTERVAL 29 | 95.0–304.0 m |
|--------------------------------------|--|---|---|---|
| | ER NOLLS SUB CRAPHIC SUB CLARKER SUB CLARK | LITHOLOGIC DESCRIPTION | TIME - BOURD INFERENCE CHARACTER CHARACTER RECTION RETERS RECTION RETERS RECTION RETERS RECTION RETERS RECTION RETERS RECTION RETERS RECTION RETERS RECTION RECENT RECTION RECENT | LITHOLOGIC DESCRIPTION |
| LOWER PLIOCENE Spongaster pentas? | 0.5 0.5 1 0.5 1 0.5 2 0.5 3 0.5 5 0.5 6 0.5 | MUDDY SILTSTONE, gravish olive green (5GY 3/2), firm to well-inducated internally structurates. Section 2: irregular and straight discontinuous fractures. Section 4: well inducated with minor other material batween the large blocks. Section 5: traight fractures with polished surfaces. Practured (non-coherent the large blocks. Section 6: straight fractures with polished surfaces. ORGANIC CARBON AND CARBONATE 740 % Organic Carbon 15 % Organic Carbon 15 % Organic Carbon 2.0 % Organic Carbon 3.0 % Organic Carbon 3.0 % Organic Carbon 4.0 % Organic Carbon 5.0 % Organic Carbo | | MUDDY SILT and MUDDY SILTSTONE, grayish oliver green (SGY 3/2). Chips of indurated material in soupy drilling breccia. SMEAR SLIDES 2,73 (D) TEXTURE: Sand 4 Stit 66 Clay 40 Clay 40 Clay 51 Clay 50 Clay 50 Clay 30 Clay 30 Clab, unopec 2 Formulifer TR Clay 30 Clab, unopec 2 Formulifer TR PelSopar 7 Mica 3 Heavy minerals 1 Prite TR Clay 30 Clab, unopec 2 Formulifer TR PelSopar 1 Sponge spicules 1 |
| B B RM | 7 | Horizontal slickensides | | |





Pyrite

Clay Glauconite

Diatoms

Sand

Clay

Silt

Sponge spicules

25 TR

TR TR

1-80

12.3

67.7

20.0

Fine sand

3

BB

| SITE 491 | HOLE | 0 | ORE | 40 CORED IN | TERVA | AL 352.0-361.5 m | SITE | 491 | н | OLE | | CO | RE | 41 CORED INT | ERVA | L 361.5-371.0 m | |
|---|------------------------------|---------|--------|----------------------|---|--|----------------|--------------------|--------------|------------------------------|---------|----------------------------|--------|---------------------------------|--------------------------------------|--|---------------------------------|
| DHIC | FOSSIL | ER | | | TT | | | PHIC | 6 | FOS | SIL | | | | | | |
| TIME - ROCI UNIT BIOSTRATIGRA ZONE | NANNOFOSSILS RADIOLARIANS | SECTION | METERS | GRAPHIC LITHOLOGY | DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES | LITHOLOGIC DESCRIPTION | TIME - ROCH | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS RADIOLARIAMS | DIATOMS | SECTION | METERS | GRAPHIC LITHOLOGY ONITING | SEDIMENTARY STRUCTURES SAMPLES | LITHOLOGIC DESCRIPTION | |
| LOWER PLIOCENE Sporgester perios? | BRM | 3 | | | | Fisil MUDDY SILT, grayish olive green (5GY 3/2) with fine to medium SAND. Section 1; firm to very firm. Section 2; very firm; mottled (bioturbation?); discontinuous bad, which hook like small state deformation, as in stunging. Section 3: very firm, diffing lamination; wery firm, diffing lamination; wery firm, diffing lamination; wery firm, diffing lamination, and in stunging. Section 4: moterately firm only a tree drilling lamination. Section 6: moderately firm only and drilling lamination. Section 6: moderately firm, diffing lamination; wery difficult firm, and drilling lamination. 32 SMEAR SLIDES 9 \$\$\$ and \$\$\$ 9 \$\$\$ and \$\$\$ 9 \$\$\$ and \$\$\$ 9 \$\$\$ and \$\$\$ 9 \$\$\$\$ apparent dip 9 \$\$\$ apparent dip 9 \$\$\$\$ apparent dip 9 \$\$\$\$ apparent dip 9 \$\$\$\$\$\$\$\$ apparent dip 9 \$ | LOWER PLIOCENE | Spongaster pentias | 8 1 | 8 FM | | 1 2 3 4 5 6 | | | * | MUDDY SILTSTONE, gravith olive green (50 with fire SAND and Silt Tayers, Section 1: whole fraumanian in the same drilling lamination. Section 2: numerous high angle fractures is lickensides: probabily drilling deformation. Section 2: drilling lamination. 2: drilling lamination. 2: drilling lamination. 2: drilling lamination. Section 2: drilling lamination. 2: drilling lamination. 2: drilling lamination. 2: drilling lamination. 3: drilling lamination. | Y 3/2) s section ction 4: |

| SIT | TE 4 | 91 | HOLE | CORE | 42 CORED INTE | RVAL | 371.0-380.5 m | SIT | TE | 491 | HOL | E | CORE | 43 CORED IN | TERVA | L 380.5–390.0 m |
|-------------|---------------|----------------------------|---|------------------------|---------------------------------|-----------------------|---|-------------|----------------|-------------------------|--------------|-------------------------|---|----------------------|---|---|
| TIME - ROCK | UNIT | ZONE | FOSSIL CHARACTED SIISOJONWWN BADIOLARIANS SUISOJONNAN | SECTION | GRAPHIC LITHOLOGY UNITING | STRUCTURES SAMPLES | LITHOLOGIC DESCRIPTION | TIME - ROCK | UNIT | ZONE | NANNOFOSSILS | RADIOLARIANS BIOLARIANS | SECTION | GRAPHIC LITHOLOGY | DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES | LITHOLOGIC DESCRIPTION |
| | LOWER PLICENE | i ventual aussibuo/c ci NN | B FM | 2 3 4 5 CC | | | MUDDY SILT, soft to had, grayish olive green (GSY 32) to olive gray, GSY 320, Section 3: in place moddy siltstore, sporadic drilling laminations; in place irregular backly fracturig/indplant lissifility (polish fisile sporadic drilling laminations; framework). Section 4: both to lin occurs model siltstore. Section 4: both to lin occurs and section 4: both to lin occurs and section 4: both section 5: time model silt, in places model siltstore. Section 4: both to lin obte of siltstore section 5: time model siltstore section 5: time section 5: time model siltstore section 5: time section 5: tim | re re | LOWER PLIOCENE | NN 15 Spongeter pentez? | 8 RP | RM | 0.5 1 1.0 2 2 3 3 4 4 5 5 6 6 CC | | | Soupy: high line and centeric couldry originally received it this and introduct MUDDY SILTSTONE, grayish olive green (SGY 3/2) (in places soft to firm silt with fine SAND beds. Intense irregular fracturing Fire and: thatp benylliftus top: 15° apparent dip Discontance fine and ted, possibly displanel along a retrieve top Bedding if apparent dip Bedding sized out by worder unit in muddy allotes: 15° apparent dip Fries and: 15° apparent dip Fries and apparent dip Fries and: 15° apparent dip Fries and: 15° apparent |



| SITE | 491 | Г | F | OSS | IL. | 1 | CO | KE | NO CORED | T | ER | | 393'0—403'0 W |
|---------------------|--------------------|--------------|--------------|--------------|---------|---|--------------|-----------------------|----------------------|----------|---------------------------|---------|--|
| TIME - ROCK UNIT | BIOSTRATIGRAP | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY | SAMPLES | LITHOLOGIC DESCRIPTION |
| | | | | | | | , | 0.5 | | 0404 | | | MUDDY SILTSTONE with MUDSTONE both gr olive great (BGY 3/2) and fine to medium SAND. Se 1: fractured blocks in drilling breccia, soapy matrix. SMEAR SLIDES |
| ER PLIOCENE | Spongaster pentas? | | | | | | 2 | THE FULLER TO THE THE | | | 111 | | Sand back and muddy silationy cut by spredominant 60–65° 4-42 4-966 Cut by spredominant 60–65° TEXTURE: Sand 7 15 Sand 7 15 51 Silt 53 30 15 Clay 40 65 65 detail-1 -15° appendix COMPOSITION: Fieldspir 6 6 |
| LOWE | 51 NN | | | | | | 3 | | | | - | | Promiser and concentration Yapu diffuse concentration Yapu diffuse concentration of Rafb unspec. TR TR Inne and Rafb unspec. TR - Diatoms TR - |
| | | | 00 | - | | | 4 | trutur | | | 7 | | Bedding within muddy filtstom soberent dip 10° adding appenent dip 12° and healed fracture 50° appenent dip |
| | 401 | в | HP | FM | | | CC | | 46 | 1 | 1 | | 400.0.419.5 |
| × | APHIC | | CHA | OSS RAC | IL | , | co | RE | 40 CORED | INT | TER | | 409,0416,5 m |
| TIME - ROC UNIT | BIOSTRATIGRI | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY STRIPTIDEE | SAMPLES | LITHOLOGIC DESCRIPTION |
| LOWER PLIOCENE | Spongaster pentas? | в | RP | FM | | | 1 2 CC | 0.5 | | | 121 | | Siltstone fragments in drilling breccia soup MUDDY SILT, gravish olive green (5GY 3/2). SMEAR SLIDES VOID Healed fault: 48 apparent dip Healed fault: 48 apparent dip (145 cm) 1 cm thick fine sand bed (sharp base, diffuse top) Silt 60 Clay 30 COMPOSITION: Ouartz 64 Feldgar 7 Mica 2 Heavy minerals TR Pyrite TR Clay 25 Carb. unspec. 1 Foraminifiers TR Nancefossilis TB |



12" apparent dip

51° true dip (fissility)

 1.5 cm fine sand bed (20" true dip) Normal fault

Diatoms

Sponge spicules Plant fragments

2

TR

0

0

0

0

OG IW

3

5

CC

SITE 491 HOLE

TIME - ROCK UNIT

LOWER PLIOCENE

NN 15

BIOSTRATIGN. ZONE

FOSSIL

FORAMINIFERS NANNOFOSSILS RADIOLARIANS DIATOMS











Nannofossils

Diatoms

0

0

5

cc

FP RP

Radiolarians

Sponge spicules

TR TR TR

ORGANIC CARBON AND CARBONATE 1-99 % Organic Carbon 1.1 % CaCO3

1.0

TR TR



| | HIC | Γ | CH/ | OSS | IL | | | | | Γ | Π | Π | | | |
|-------------|--------------|--------------|--------------|--------------|---------|----|---------|---|----------------------|----------|---------------------------|---------|---|---|------------------------------|
| TIME - ROCK | BIOSTRATIGRA | FORAMINIFERS | NANNOFOSSILS | RADIOLARIANS | DIATOMS | | SECTION | METERS | GRAPHIC LITHOLOGY | DRILLING | SEDIMENTARY STRUCTURES | SAMPLES | LITHOL | OGIC DESCRIPTION | |
| | | | | | | | 1 | 0.5 | | A | 11 | | Silt lamination very f 28* apparent dip | DY SILTSTONE, grayis ine to coarse SAND. | h olive green (5GY 3/2) with |
| | | | | | | | | 1.0 | | 000 | | | Drilling breccie | SMEAR SLIDES | 4 |
| OCENE | | | | | | | | 1111 | | | | | Drilling breccia | TENTINE | 21-50 (D) |
| OWER PLI | | | | | | | 2 | The second se | | 000 | | | Soupy, fine tand, same color as muddy siltstone | Sand Silt Clay COMPOSITION: | 15 55 30 |
| - | | | | | | | | | | 400 | | | Drilling breccia | Quartz Feldspar Mica | 60 9 4 |
| | | | | | | | l | - Him | | 440 | | | Fine sand | Heavy minerasi Pyrite Clay Carb. unspec. | TR 1 25 TR |
| | | | | | | | 3 | 1100 | | 40- | | | Fine sand | Radiolarians Diatoms Sponge spicules | TR 1 TR |
| | | | | | | | | - | | | | | Graded sand, coarse to fine Fine sand Fine sand | | |
| | | | | | | cç | 4 | | | 1 | | | Very fine sand Very fine sand | | |

SITE 491 HOLE CORE 59 CORED INTERVAL 532.5-542.0 m

| NAMNOFOSSILS RADICLARIANS DIATOMS | SECTION | GRAPHIC LITHOLOGY | DRILLING | TRUCTURES | MPLES | LITHOLOGIC DESCRIPTION |
|---|---------|---|----------|-----------|--------|--|
| | - | A CONTRACTOR OF THE OWNER OWNER OF THE OWNER OWNE | _ | 20 60 | S | |
| RP RM | 1 0.5 | | 0 4 00 | 22 60 | • | Medium rand Drilling breecie MUDDY SILTSTONE and coarse to medium SAND. Coarse to medium and, probably graded |
| RP | RM | RM <u>CC</u> | RM CC | RM 1 0.3 | RM 105 | RM 103 |



























SITE 491

























SITE 491











SITE 491



