Shipboard Scientific Party²

Date occupied: 15 January 1980; 0611 hr. (beacon dropped)

Date departed: 21 January 1980; 0342 hr. (underway; beacon close aboard 0532)

Number of holes: 1

Time on hole: 141 hr.

Position: 51°00.28'S, 46°58.30'W

Water depth (sea level; corrected m, echo-sounding): 2589

Water depth (rig floor; corrected m, echo-sounding): 2599

Bottom felt (m, drill pipe): 2602

Penetration (m): 632

Number of cores: 70

Total length of cored section (m): 632

Total core recovered (m): 385.62

Core recovery (%): 61

Oldest sediment cored:

Depth sub-bottom (m): 632

Nature: Black mudstone and nannofossil mudstone (black shales), rich in organic matter Age: Late Jurassic

Measured velocity (km/s): ~2.3

Principal results: Drilling at Site 511 gave definitive results on the nature and age of the seismically determined wedge of sediments on the Falkland Plateau. The section drilled and continuously cored in the basin province of the plateau, about 10 km south of Site 330 on Maurice Ewing Bank, consists largely of diatomaceous, nannofossil, calcareous, and zeolitic oozes of Paleocene-Eocene to early Oligocene age between seafloor and a major erosional unconformity at 195 meters depth. This unit is underlain by 14 meters of early Maestrichtian calcareous and zeolitic calcareous oozes followed by 203 meters of zeolitic clays and claystones of Turonian to Campanian age, below which are 86 meters of early Albian to Turonian claystones and nannofossil claystones and chalks. The chalks rest on mudstones and nannofossil mudstones of Jurassic to early Albian age, rich in organic matter and called the black shales, of which 134 meters were cored for a TD of 632 meters below seafloor. A depositional hiatus of some 20 m.y. separates the Jurassic from the Cretaceous sediments.

At Site 511, the first successful heat flow measurement was obtained on the plateau, through use of the Downhole Temperature Probe. The heat flow is 1.5 HFU.

BACKGROUND AND OBJECTIVES

The Falkland Plateau extending eastward off southernmost South America is bounded on the north by the Falkland Escarpment and on the south by the Falkland Trough and North Scotia Ridge (Fig. 1). The western part of the plateau is a segment of oceanic crust (or attenuated continental crust) over which have been deposited 4-6 km of sediment in a basin bounded by the Falkland Islands Platform on the west, a narrow ridge associated with the Falkland Escarpment on the north, Maurice Ewing Bank on the east, and the North Scotia Ridge on the south. Maurice Ewing Bank, a subsided continental block that received thick carbonate as well as siliceous biogenic sediment throughout most of its history, forms the eastern part of the plateau.

In the basin thick, widespread sheets of sediment dip southward from the Falkland Escarpment ridge and are terminated updip by erosional truncation (Ludwig, this volume). They lap out against the Falkland Islands Platform and Maurice Ewing Bank. The lower boundary of the depositional sequence has been disrupted through movement of the North Scotia Ridge toward the plateau, resulting in deformation and uplift of near-surface sediment layers onto the north side of the ridge and subduction of lower layers beneath the ridge.

The wedge of sediments on the plateau appears to represent basin-slope and basin-floor seismic facies units. In the parallel sheet-drape facies generally described by Sangree and Widmier (1977), "Parallel reflections... drape over contemporaneous topography with only gradual changes in thickness or reflection character and suggest uniform deposition independent of the bottom relief. This pattern is strongly indicative of deep-marine hemipelagic clays and oozes" (p. 176).

Site 511 is located in the basin province of the Falkland Plateau, on the back slope of a cuesta-type ridge and about 10 km south of DSDP Site 330 on Maurice Ewing Bank, where Leg 36 scientists drilled 567 meters of Mesozoic and Cenozoic sediment above continental basement (Shipboard Scientific Party et al., 1977). During Middle to Late Jurassic to late Early Cretaceous (Aptian) time, shallow marine sedimentation, sometimes under euxinic conditions, prevailed on the bank. Predominantly pelagic deposition was established over the bank during Albian time as the bank subsided, with subsequent accumulation of clays containing planktonic fossils.

Multichannel seismic reflection records collected during Conrad cruise 21-06 (Fig. 2) show reflections dipping steeply to the south from the vicinity of Site 330.

¹ Ludwig, W. J., Krasheninnikov, V. A., et al., Init. Repts. DSDP, 71: Washington (U.S. Govt. Printing Office). ² William J. Ludwig (Co-Chief Scientist), Lamont-Doherty Geological Observatory.

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Figure 1. Locations of Site 511 and other Leg 71 drill sites.



Figure 2. Robert D. Conrad 21-06 multichannel seismic reflection profile near Sites 511, 330, and 327. See Figure 3 for location.

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Under the ridge structure, reflections are weak owing to the effect of scattering of seismic energy. Nevertheless, several prominent reflections may be traced from Site 330 to Site 511. The unconformity that apparently marks a major erosional hiatus on the plateau is present at Site 511. Thus Site 511 has been situated to delineate and define the depositional environment and the erosion and subsidence history of the Falkland Plateau. Cores from the site are expected to help clarify the biostratigraphic, biogeographic, and sedimentological history related to major changes in circulation patterns through the Mesozoic and Cenozoic. The paleoceanographic alterations occurred in response to the breakup of Gondwanaland and the initial opening and enlargement of the South Atlantic, the development of bottom water passageways through the North Scotia Ridge, the opening of the Drake Passage, and the climatic evolution of Antarctica and the Southern Ocean.

SURVEY AND OPERATIONS

Glomar Challenger left Punta Arenas, Chile, at 1535 hr. on 10 January 1980. The route to the first drill site took the ship through the Strait of Magellan, then eastward to skirt the Falkland Islands to the south and on to Site 511 on the Falkland Plateau. The magnetometer and seismic reflection gear were streamed about midway between the mainland and the Falkland Islands. Progress was slowed first by head winds, then by low visibility in fog, and the average speed was held to 8.3 kn. As a result the ship arrived on site about 18 hr. later than anticipated.

Multichannel Seismic Line 142 (Fig. 2) obtained during Conrad cruise 21-06 was used as the basis for site location. Challenger arrived at the longitudinal location of the seismic line at 0430, 15 January, at a preselected point south of Site 511. After the ship changed course to the north and the speed was reduced to 5 kn., the site area on the back slope of a well-mapped, cuesta-type ridge was soon located. We profiled a short distance past the site to confirm its morphostructural setting, turned back, and dropped the beacon at 0611 hr., 15 January. Subsequent satellite fixes indicated that at 1700 hr. we were about one mile to the east of the approved site on the Conrad 21-06 seismic line; however, because the Challenger seismic line matched the Conrad line, the beacon position was occupied for the site (Fig. 3).

The first core was recovered from Site 511 at 1838 hr., 15 January (Table 1). Drilling and coring proceeded routinely to 0812, 20 January, when the hole was terminated at 632 meters below seafloor (the depth corresponding approximately to a 4.1 s reflector limit of penetration imposed by the JOIDES Safety Panel). Heat flow measurements were obtained at 52.5 meters and 119 meters below seafloor. Upon completion of the hole, a gamma-ray and neutron log was run through the pipe.

Our plan called for offsetting *Challenger* about 600 meters to the south and drilling an A hole to obtain a better percentage of core recovery in the 110–220-meter interval (below seafloor) than was obtained at the main

site. The Paleocene section there is either missing or was missed in the coring, and as a consequence we were unable accurately to date a major erosional unconformity at 195 meters depth. However, after the drill string was pulled above the mud line, it was found that the hydraulic release on the drill bit had been activated and the bit had been dropped. Because of increasingly heavy seas and a 7-day limitation on site occupation time south of 50°S, it became necessary to leave Site 511 and move on to the next site on Maurice Ewing Bank.

LITHOLOGICAL SUMMARY

A sequence of pelagic and hemipelagic oozes and claystones was continuously cored at Site 511 to a subbottom depth of 632 meters. Below a thin veneer of Pliocene to Recent sediments, strata ranged in age from Late Jurassic to early Oligocene. The dominant lithologies cored were nannofossil-diatomaceous oozes, zeolitic claystones, and muddy nannofossil chalks.

Lithologic units were distinguished principally on the basis of color and composition. The six major units, their ages, thicknesses, colors, and major characteristics are summarized in Figure 4 and Table 2. In general, sediments cored below Core 27 (223.5 m) were sufficiently indurated to be termed claystones and chalks, whereas those above this depth were clays and oozes.

Unit 1

Unit 1 consists of a 3-meter veneer of gray to olive gray, siliceous, gravelly sands and thin foraminiferal oozes of Pliocene-Recent age. Pebbles are subangular to angular and of highly varied lithology including quartzite, basalt, and granite. Several Mn nodules (up to 5 cm in diameter) and manganiferous zones several centimeters thick were encountered in this unit. These youngest Cenozoic glaciomarine sediments unconformably overlie Unit 2.

Unit 2

Subunit 2A

Subunit 2A consists of massive, gray to dark gray, muddy diatomaceous oozes and muddy nannofossil-diatomaceous oozes of Paleogene age recovered in Cores 3 to 20 (3-186 m sub-bottom). This subunit consists of slightly to moderately bioturbated diatomaceous oozes with variable carbonate content.

The recovered cores are characterized by severe drilling disturbance resulting in the occurrence of frequent zones of coarse sand and gravel caved in from the Pliocene-Recent strata above; it is thought that many of the large Mn nodules from the lower part of this unit may have been similarly derived. The small Mn nodules (less than 1 cm diameter) which occur commonly in the uppermost part of the unit are thought to be *in situ*.

Disseminated, very fine, sand-sized detrital glauconite is common, particularly in the lower part of the unit.

Several more indurated zones (4–10 cm thick) were encountered in Core 12. No apparent difference in composition was noted between these harder layers and the surrounding oozes.





		1	Cored nterval (m)	Lithology	Age	Benthic Foraminifers	Planktonic Foraminifers	Nanno- plank ton	Radiolarians	Diatoms	Silico- flagellates	Dino- flagellates	Macrofauna	Sponge Spicules	Palynology	Calci- spaerulids
	-0	1 2	- 5		Neo.2-Quat.	Rare Recent species	N22	Emiliania huxleyi Zone	Mixed P ₃ - Quaternary species	Neo.2-Quat.	Neo.2-Onat	Barren		Barren	Barren	
	-	3	- 24.0	<u>、</u> 、、、、、、、、、、、、、、、、、、、、、、、、、、、、、、、、、、、					berosa- ris ateuchu alent	superbus Zone	Naviculo constrict Corbise archangeli Zone					
	50 -	5	43.0				l.	Clausioco fenestrata 2	eocyrtis tu Dorcadospy tonal equiv		icta/ a Zone			ida		
	-	7 8	- 62.0 - 71.5		early Oligocene	siformis, nulata,	<i>brevis</i> Zone		early 1	<i>Rhizosolenia gravida</i> Zone	N. constr Dictyoch deflandrei	rianum, sudiae,		ss sponges re Tetraxon	rpus, eae,	
		9 10	- 81.0		254	ınsis, A. spis dənə, G. pla O. rohri, etc	erina angipo r part of G.	Blackites			N. trispinosa Zone	e of I. victo tum, H. rig compositum		iblage of gla <i>iae)</i> with ra	len <i>(Podoca</i> ibordinate gus, <i>Myrtac</i> spores	Barren
	100 -	11 12	-100.0 -109.5			with B. cube tus, G. giran icostatum, (<i>Globig</i> uppe	<i>spinosus</i> Zone	Theocyrtis tuberosa zonal	Brightwellia spiralis Zone		Assemblag K. capula C. in		Assem Hyalospong	iosperm pol etc.) and su len <i>(Nothofa</i> 9, etc.) and	
	-	13	-119.0 -128.5			Assemblage O. umbona C. tenu			equivalent	architec- turalis Zone	entalis cene to locene)			2.7	Indant gymr Dacrydium, iosperm pol Casuarin	
th (m)	150 -	16 17	-138.0 -147.5				ja			Asterolampra insignis Zone	socena occió le (upper Eo vermost Olig	um sp.			Abu	
bottom Dep	-	18 19	- 157.0		late Eocene		naperta Zon ver G. bravis Zone	oamaruensis Zone	Thyroscyrtis bromia zonal equivalent	Rylandsia inaequiradiata Zone	Me	A. ornata atosphaeridi atiacasphaer chetodiniun	Barren	semblage of etraxonida		
Sub-l	4	20	-176.0				G. Ii lov	A.				Adra B		A T		
		21 22	-185.5 -190.0 -195.0		Paleocene- Eocene	Barren	Barren	Barren	middle to early Eocene species						Barren	
	200 -	23 24 25	-204.5	1 z	late Campanian?— Maestrichtian	B. incrassata R. szajnochae G. quadratus	mblage of ibergella erohelix obigerina rinelloides	Biscutum coronum Zone	late Cretaceous species			Barren				Abundant
	-	26 27 28	-219.0	2 2 2		Impoverished	Asse Het Rugog Globige	Barren								Bare
	250	29 30	-242.5	Z		long-ranging species	Barren									
2	-	31	-252.0	Z		tof	ae, iana, des, ispinata		Barren	Barren	Barren			Barren	Sporadic pollen and	
	1	33	-271.0 -280.5	Z	Campanian	uus speciei nulina, yroidina, illenia, etc.	. plumme , G. linne obigerinoi e, S. multi	asterites atus yne				tum rifera liadema cula		4920234/12291	spores	
	300	34 35	- 290.0	Z		ly calcared ina, Margi inopsis, G inoides, Pu protalites,	age with G . cororiata nata, G. glu G. cretace	Marth furc Zo				1. la 0. po Amphid A. nu				Hare to very rare
	-	36 37	-309.0	z	14	erse, most Dental Marginul Gyroid	Assembli G. arca, G G. margli bulloides,									
		38 39	-328.0	Z 		Div	G.									



Figure 4. Columnar section of Hole 511 showing the lithology recovered and biostratigraphic correlations. Although the cored interval for Core 24 was 4.5 m, 9 m of sediment were recovered. (Refer to Ludwig et al., Introduction, this volume, for a key to the lithologic symbols.).

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Table 1. Coring summary, Site 511.

	Date		Depth from	Depth below	Length	Length	Core
Core No.	(Jan. 1980)	Time	Drill Floor	Seafloor (m)	Cored (m)	Recovered (m)	Recovered
Hole	11		(,	(()		
		1020	2002 0 2002 0			4.00	07.0
1	15	1838	2602.0-2607.0	0.0-5.0	5.0	4.89	97.8
2	15	22103	2607.0-2010.3	14 5 24 0	9.5	6.50	68 4
4	15	2320	2626 0-2635 5	24 0-33 5	9.5	3 70	38.9
5	16	0030	2635 5-2645 0	33 5-43 0	9.5	5.98	62.9
6	16	0210	2645 0-2654 5	43 0-52 5	9.5	5 41	57.0
	Heat flow	0210	2045.0-2054.5	45.0-52.5	1.5	5.41	57.0
7	16	0501	2654.5-2664.0	52.5-62.0	9.5	3.54	37.2
8	16	0610	2664.0-2673.5	62.0-71.5	9.5	0.04	0.4
9	16	0730	2673.5-2683.0	71.5-80.0	9.5	9.67	101
10	16	0848	2683.0-2692.5	80.0-89.5	9.5	0.40	4.2
11	16	1000	2692.5-2702.0	89.5-99.0	9.5	4.94	52.0
12	16	1120	2702.0-2711.5	99.0-108.5	9.5	3.12	32.8
13	16	1243	2711.5-2721.0	109.5-119.0	9.5	0.32	3.3
220	Heat flow						
14	16	1541	2721.0-2730.5	119.0-128.5	9.5	0.05	0.5
15	10	1705	2730.5-2740.0	128.5-138.0	9.5	1.16	12.2
10	10	1805	2740.0-2749.5	138.0-147.5	9.5	2.40	25.2
17	10	1913	2/49.3-2/59.0	147.5-157.0	9.5	3.52	37.0
10	10	2034	2/39.0-2/08.3	157.0-100.5	9.5	2.82	29.0
20	16	2137	2700.3-2778.0	176 0 196 5	9.5	3 79	20.7
20	17	2310	27/0.0-2/0/.3	1/0.0-185.5	9.5	0.70	39.7
22	17	0150	2792 0-2792.0	100.0-105.0	5.0	0.72	10.0
23	17	0332	2792.0-2797.0	195.0-195.0	0.5	0.60	7.2
24	17	0451	2806 5-2811 0	204 5-209.0	45	9.78	217 3
25	17	0555	2811 0-2816 0	209.0-214.0	5.0	1.65	33.0
26	17	0709	2816 0-2821 0	214 0-219 0	5.0	0.15	3.0
27	17	0921	2821 0-2825 5	219 0-223 5	45	2 42	53 7
28	17	1050	2825 5-2835 0	223 5-233.0	9.5	9 38	100
29	17	1210	2835 0-2844 5	233 0-242 5	9.5	0.45	47
30	17	1401	2844.5-2854.0	242.5-252.0	9.5	8.16	85.8
31	17	1517	2854.0-2863.5	252.0-261.5	9.5	9.24	97.2
32	17	1640	2863.5-2873.0	261.5-271.0	9.5	9.00	94.7
33	17	1810	2873.0-2882.5	271.0-280.5	9.5	9.20	96.8
34	17	1928	2882.5-2892.0	280.5-290.0	9.5	9.75	100
35	17	2100	2892.0-2901.5	290.0-299.5	9.5	2.43	41
36	17	2227	2901.5-2911.0	299.5-309.0	9.5	9.59	100
37	17	2347	2911.0-2920.5	309.0-318.5	9.5	2.43	26
38	18	0115	2920.5-2930.0	318.5-328.0	9.5	8.08	85
39	18	0250	2930.0-2939.5	328.0-337.5	9.5	5.21	55
40	18	0403	2939.5-2949.0	337.5-347.0	9.5	7.87	83
41	18	0640	2949.0-2958.5	347.0-356.5	9.5	4.59	48
42	18	0810	2958.5-2968.0	356.5-366.0	9.5	7.03	74
43	18	0932	2968.0-2977.5	366.0-375.5	9.5	6.51	69
44	18	1050	2977.5-2987.0	375.5-385.0	9.5	7.01	74
45	18	1224	2987.0-2996.5	385.0-394.5	9.5	4.20	44
46	18	1400	2996.0-3006.0	394.5-404.0	9.5	5.25	55
47	18	1510	3006.0-3015.5	404.0-413.5	9.5	8.98	94
48	18	1644	3015.5-3025.0	413.5-423.0	9.5	5.84	61
49	18	1820	3025.0-3034.5	423.0-432.5	9.5	8.05	84
50	18	1927	3034.5-3044.0	432.5-442.0	9.5	5.30	56
51	18	2045	3044.0-3053.5	442.0-451.5	9.5	8.95	94
52	10	2215	3053.3-3063.0	451.5-401.0	9.5	9.89	100
55	10	2343	3003.0-3072.3	401.0-470.3	9.5	9.20	97
54	19	0259	3072.3-3082.0	470.3-480.0	9.5	8.72	92
56	19	0532	3091 5-3101 0	480.5-400.0	9.5	7 48	70
57	19	0756	3101.0-3110.5	499 0-508 5	0.5	8.05	89
58	19	0950	3110 5-3120.0	508 5-518 0	9.5	6.18	65
50	19	1150	3120 0-3120 5	518 0-527 5	9.5	6.03	63
60	19	1350	3129.5-3139.0	527.5-535.0	9.5	8.23	87
61	19	1614	3139.0-3148 4	537.0-546.5	9.5	7.44	78
62	19	1809	3148.5-3158.0	546.5-556.0	9.5	8.69	91
63	19	1952	3158.0-3167.5	556.0-565.5	9.5	4.89	51
64	19	2130	3167.5-3177.0	565.0-575.0	9.5	7.97	84
65	19	2317	3177.0-3186.5	575.0-584.5	9.5	6.60	69
66	20	0056	3186.5-3196.0	584.5-594.0	9.5	7.94	84
67	20	0240	3196.0-3205.5	594.0-603.5	9.5	7.21	76
68	20	0425	3205.5-3215.0	603.5-613.0	9.5	3.16	33
69	20	0602	3215.0-3224.5	613.0-622.5	9.5	6.65	70
70	20	0812	3224.5-3234.0	622.5-632.0	9.5	6.78	71
Total					612.0	386 6	61
rotat		_			052.0	303.3	- 01

Subunit 2B

Subunit 2B, consisting of 72 cm of greenish gray (5GY 6/1) clays obtained in Core 21, is characterized by minor amounts of glauconite, chert, and volcanic ash. In Core 22, no sample was recovered; in the absence of core, Subunit 2B is arbitrarily continued to the base of Core 22.

Unit 3

This unit consists of gray, calcareous, and zeolitic foraminiferal oozes in Cores 23 and 24 (195-209 m, sub-

bottom). These sediments are early Maestrichtian to late Campanian in age and are characterized by minor to moderate bioturbation, the presence of occasional chert pebbles (up to 5 cm diameter), and disseminated minor glauconite.

Unit 4

Unit 4 (Cores 25-47, 209-412 m sub-bottom) consists of zeolitic clays and claystones commencing near the bottom of Core 24. These units are Maestrichtian-Campanian to Turonian in age. Though the contact with Unit 3 is transitional, marked by an alternation of zeolitic clays and calcareous oozes, it is placed at 209 meters (bottom of Core 24), below which the zeolitic clays and claystones predominate. In general the darker zones (5Y 4/1) are more zeolitic than the lighter sections (5Y 6/1), which tend to be somewhat more calcareous. Dark greenish black mottles, lenses, and thin laminae (1-2 mm) were common throughout the unit.

The generally gray zeolitic clays and claystones are characterized by moderate to very intense bioturbation which has, in many sections, completely obliterated any evidence of stratification. The tremendous concentration of burrows of different colors (gray, light gray, dark gray, greenish gray, and dark greenish gray) in some sections gives the core a "conglomeratic" appearance. Many types of burrows were identified, including solid, ring, Zoophycos, Chondrites, Planolites, and Teichichnus.

Pyrite is frequently encountered, usually in burrows. In Core 42, however, pyrite occurs as a vein-filling along vertically and obliquely oriented fractures. Irregularly shaped, slightly brownish concretions of barite, ranging in diameter from $\sim 0.5-4$ cm, occur frequently throughout the unit.

Below Core 41, Section 3, fragments of the pelecypod *Inoceramus*, up to 3 mm thick and 3-4 cm long, were locally abundant.

Unit 5

Cores 48-56 (413.5-492 m sub-bottom) are early Albian to Turonian sediments dominated by reddish brown zones that frequently have a "blotchy" appearance. The unit is marked by minor to intense bioturbation with burrows frequently filled with pyrite.

Layers of skeletal carbonate debris, principally pelecypods (Aucellina and Inoceramus), occur frequently (Fig. 5) as muddy microcoquinas. Single, thin, whole pelecypod valves occur as well throughout the unit and Inoceramus "needles" are frequently concentrated in thin olive gray claystone bands. Scattered solitary corals also occur in this unit.

Small barite nodules (approximately 0.5 cm), similar to those in Unit 4, occur in Cores 54 and 55.

In Core 51, Section 1, several very dark gray, highly calcareous sapropelic claystone layers (0.5 cm thick) are interstratified with reddish brown calcareous claystones and muddy calcareous chalks (muddy microcoquinas).

Unit 6

The lowest lithologic unit cored (Cores 57-70, 498-632 m sub-bottom) is Late Jurassic to early Albian in

Table 2. Characteristics	of	lithologic units.	
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Depth belo	N
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3 186 195 209	1, Sect. 2 1, Sect. 3 20 21 22 23 24 25	1 2A 2B 3 4	Siliceous gravelly sand and minor foram ooze Muddy diatomaceous oozes and muddy nannofossil- diatomaceous oozes Clay Calcareous ooze and zeolit-	5Y 6/4 (pale olive) 5Y 5/1 (gray) 5Y 5/2 (olive gray) near top to 5Y 4/1-5Y 7/1 (dk, gray- lt, gray) in most of unit 5 GY 6/1 (greenish-gray) 5Y 6/1 (gray)	Subangular-angular pebbles (up to 4 cm), varied lithologies. Mn nodules (up to 5 cm), thin Mn-rich zones, interstrati- fied foram ooze and gravelly sand Minor-moderate bioturbation, sparse Mn nodules, lenses of fine glauconitic sand, occasional pebbles of chert, granite, basalt, and sandstone Minor glauconite, chert, ash
186 195 209	20 21 22 23 24 25	2A 2B 3 4	Muddy diatomaceous oozes and muddy nannofossil- diatomaceous oozes Clay Calcareous ooze and zeolit-	5Y 5/2 (olive gray) near top to 5Y 4/1-5Y 7/1 (dk, gray- lt, gray) in most of unit 5 GY 6/1 (greenish-gray)	Minor-moderate bioturbation, sparse Mn nodules, lenses of fine glauconitic sand, occasional pebbles of chert, granite, basalt, and sandstone Minor glauconite, chert, ash
195 209	20 21 22 23 24 25	2B 3 4	Clay Calcareous ooze and zeolit-	5 GY 6/1 (greenish-gray)	Minor glauconite, chert, ash
195 209	21 22 23 24 25	2В 3 4	Clay Calcareous ooze and zeolit-	5 GY 6/1 (greenish-gray)	Minor glauconite, chert, ash
209	22 23 24 25	3 4	Calcareous ooze and zeolit-	5V 6/1 (aray)	
209	23 24 25	4	ic calcareous ooze and zeom-		Minor moderate bioturbation chart peb-
209	24 25	4		51 0/1 (gray)	bles (4.5 cm) minor glauconite
	25	4	Zealitie alove and alove	ST //I (IL. gray)	Moderate yery intense bioturbation py
			stones	Magor: 5Y 4/1-5Y 6/1 (dk. gray) to (gray) 2.5Y 4/0-5/0 (dk. gray- gray)	rite common, minor glauconite in upper part. <i>Inceramus</i> fragments common in lower part. Barite (?) nodules
412	47			5Y 5/2-5Y 6/2 (olive- gray to lt. olive-gray) 5GY 2/1-6/1 (dk. green- ish gray to greenish gray) 2.5 YR 3/4-4/4 (dk. red- ish brown to reddish brown)	
	48	5	Claystones, nannofossil claystones, and muddy nannofossil chalks	Major: 5YR 3/3-5/3 (dk. red- dish brown to reddish brown) 5YR 5/4 (reddish brown) Secondary: 2.5Y 4/0-5/0 (dk. gray- gray) 5G 6/1 (greenish gray) Minor: 7-5YR 5/2 (brown) 5 RP 4/2 (grayish purple) 5GY 4/1-8/1 (ltdk.	Minor-intense bioturbation, pyrite com- mon, thin pelecypod fragments locally very abundant. Solitary corals, small barite (?) nodules, thin, dark gray, sapropelic layers (Core 51-1)
498	56 Sect 4			syreenish gray)	
770	56, Sect. 5	6	Mudstones and nannofossil mudstones	Major: SY 2.5/1 (black) Minor: 2.5Y 5/1 (gray) SY 5/1 (gray)	Black shales; petroliferous; often finely laminated and fissile, minor bioturba- tion, pyrite very common, thin zeo- lite claystones, minor chert, cone-in- cone calcite. Hard calcareous zones
(22				2.5 Y 8/0 (white)	with local very high concentrations of

age and consists of black, massive, thinly laminated mudstones and nannofossil claystones which in places are very fissile. These mudstones (black shales) are highly petroliferous below Core 60 and display only minor bioturbation. The monotonous sequence of black mudstones is interrupted by (1) occasional layers of concentrated pelecypod debris (Aucellina and Inoceramus) forming muddy microcoquinas similar to those in the overlying unit; (2) soft gray (5Y 5/1) layers (0.5-3.0 cm thick) of zeolitic claystone; (3) hard, muddy calcarenite layers; and (4) occasional layers of coarsely crystalline cone-in-cone calcite.

Pyrite is a common constituent, occurring as masses up to several centimeters in diameter, in burrows, and finely disseminated along thin laminae in zones 1-2 cm thick. Belemnites (up to 1 cm diameter and several cm long) occur scattered throughout this unit (Fig. 6).

Minor slumping and microfaulting of laminae were noted in the unit (e.g., Core 56, Sections 4-5; Core 62, Section 4). Dips of strata are generally 0.5° throughout the unit except in Core 52, in which dips are consistently 10-15° despite dip test results of less than 1° from the vertical for the inclination of the hole.

PALEONTOLOGY

Biostratigraphy Summary

The 70 cores taken continuously at Site 511 ranged from Holocene to Late Jurassic in age. Nearly all cores contained microfossils and those in the lower portions of the hole contained fossil mollusks and solitary corals as well. A broad range of microfossil groups are represented in various intervals of the section. Those dealt with in these reports include benthic and planktonic foraminifers, calcareous nannofossils, diatoms, silicoflagellates, radiolarians, pollen, spores, dinoflagellates, sponge spicules, calcisphaerulids, onychites, and calcareous microorganisms of unknown affinity (Upper Cretaceous).

One of the most significant results of this site was the collection of a 154-meter diatomaceous-ooze sequence of early Oligocene age, with exceptionally well preserved and diverse diatoms, radiolarians, and silicoflagellates along with sufficient numbers of calcareous nannofossils and foraminifers to provide temporal control by correlation with lower-latitude zonations and the wellestablished New Zealand stages. Although siliceous se-



Figure 5. Layers of skeletal carbonate debris, principally pelecypods, occur frequently as microcoquinas in Lithologic Unit 5 (Sample 511-48-2, 77-102 cm).



Figure 6. Belemnites in Unit 6 (Sample 511-60-2, 15-35 cm).

quences of this age have been sampled before by DSDP (for example, DSDP Site 274, Leg 28), the critical time control by calcareous planktonic faunas and floras has, until now, always been lacking. The section at Site 511, combined with the upper Oligocene to lower Miocene section recovered at Site 513, provides a complete Oligocene section rich in siliceous microfossils. This composite section is the most complete record available in the southern high latitudes for Oligocene diatom stratigraphy. Gombos and Ciesielski (this volume) have defined 12 diatom biostratigraphic zones for the interval from lower Miocene to upper Eocene based on a study of the composite section.

A second notable result of coring at Site 511 was the sampling of a long (174 m) Campanian-Coniacian section in the Cretaceous. Most of this unexpectedly thick section was not encountered at the nearby DSDP Site 327, Leg 36.

In addition to these results, a continuously cored section was taken through the upper portion of the Lower Cretaceous-Jurassic black shale sequence of the Falkland Plateau. This long, continuous section was characterized by high core recovery, and therefore makes possible further subdivision of the black shales by the calcareous microfossils, palynomorphs, and mollusks.

Cenozoic

The top 3 meters of the Cenozoic section contains a glacial marine sequence characteristic of this area (see Ciesielski, 1978). Approximately 10 cm of Pleistocene to Holocene calcareous ooze (repeated by a double hit of the core barrel) overlie sandy or gravelly diatomaceous sediments, dated from the Brunhes to the Gauss chrons. These are separated by an erosional disconformity from the long lower Oligocene sequence. Upper age limits for the Oligocene section are imposed by the presence of the coccoliths Reticulofenestra umbilica in Core 2 and Ismolithus recurvus in Core 4; their extinctions (according to Hardenbol and Berggren, 1978) fall well within the early Oligocene. This age is supported by co-occurring planktonic foraminifers which are dated within Zones P18-P20. The luxuriant and well-preserved assemblages of siliceous microfossils offer the opportunity for a finer siliceous zonation of this interval in addition to the description of many new taxa. Pollen and spores are also relatively diverse and numerous in the lower Oligocene sediments.

The Oligocene sequence spans two radiolarian zones, four silicoflagellate zones, and five diatom zones. Many of these siliceous microfossil zones or datums within these zones have been revised or defined as new in these reports (see Gombos and Ciesielski, this volume; Weaver, this volume; Shaw and Ciesielski, this volume). The lowermost Oligocene and the underlying upper Eocene section, however, fall within the same silicoflagellate zone; thus, the group cannot at present be used to distinguish the Oligocene/Eocene boundary. This boundary has been placed at the top of Core 17 because of the occurrence of a single specimen of the planktonic foraminifer, Globigerapsis index and the presence of rare specimens of the coccolith Discoaster saipanensis. Dinoflagellate assemblages also suggest placement of the boundary there. Interestingly, two distinct assemblages of glass sponge spicules were noted, one characteristic of the Oligocene sediments (Cores 2-16) and the other present in Cores 17-20.

Calcareous microfossils do provide good control for the upper Eocene diatomaceous oozes which were sampled in Cores 18 and 20. Unfortunately, a complete sampling of this section and of much of the lower Oligocene was precluded by drilling problems. Coccoliths, however, are exceptionally well preserved in Cores 18 and 20 and correlate well with the *Reticulofenestra oamaruensis* Zone of New Zealand. The planktonic foraminifers belong to Zones P15 and P16 and can also be tied in closely to the classic New Zealand sections. It is not certain if the uppermost Eocene zone for planktonic foraminifers was sampled in Hole 511, primarily because of drilling problems.

Dating of the Eocene section is supported by the radiolarians, which are assigned to the Thyrsocyrtis bromia zonal equivalent. When the Eocene and Oligocene radiolarian assemblages are considered together, an overall cooling trend is noted. The Eocene populations indicate relatively warm (temperate) conditions whereas a cooling is evident in those from the Oligocene Cores 16-13. Radiolarians from Cores 12-11 are characteristic of proximity to Polar Front upwelling conditions and those from Cores 9-2 suggest a cool subantarctic/antarctic water mass. This overall cooling trend may reflect the northward migration of major water mass boundaries across Site 511 during the Eocene-Oligocene. A transgression over the Falkland Plateau of the front separating temperature and subantarctic/antarctic water masses may account for the upwelling and temperature change noted in Cores 13-2.

A high productivity of siliceous planktonic organisms throughout the late Eocene–early Oligocene is indicated by the exceptionally high sediment accumulation rate of about 44 m/m.y. recorded for this interval (Fig. 7).

Core 21 sampled a yellowish pelagic clay (Lithologic Subunit 2B) with some poorly preserved radiolarians and a few specimens of the diatom *Trinacria simulacrum*. This core is thought to be late Paleocene to early Eocene in age. Unfortunately, Core 22 contained no sediment so little more can be said about this poorly dated interval. Low accumulation rates plotted in Figure 7 suggest a disconformity between these cores and the Cretaceous sediment of Core 23.

Mesozoic

Cores 23 and 24 recovered an early Maestrichtianlate Campanian calcareous ooze which contained wellpreserved foraminifers and a high-latitude coccolith assemblage so far observed only from DSDP Hole 327A. DSDP cores of this age from comparable latitudes in the Pacific sector of the Southern Ocean contain no carbonate. Benthic foraminifers dominated by calcareous species which lived above the lysocline data Cores 25-27 as early Maestrichtian-Campanian. The next core down, however, is nearly barren of all microfossils, owing to dissolution; only a few benthic foraminifers and lower Campanian coccoliths are present. Below that begins the long (19-core) section of zeolitic clays and claystones which are dated primarily by planktonic foraminifers. Sedimentation rates in this interval were exceptionally high (Fig. 7). Core 29 to Sample 511-41-3, 55-57 cm are dated as Campanian by somewhat rare Globotruncana. Coccoliths further suggest an early Campanian age. The interval from Sample 511-41,CC to Sample 511-43,CC

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Figure 7. Sedimentation rates at Site 511.

is Santonian (probably late Santonian because of the absence of the coccolith *Lithastrinus floralis*), whereas Cores 44-46 are considered undifferentiated Coniacian-Santonian. Planktonic foraminifers date Core 47 down to the topmost sample of Core 48 as Turonian in age; the remainder of that core is Cenomanian.

In general, the occurrence of planktonic foraminifers in the Upper Cretaceous sediments is quite sporadic because of dissolution. Preservation varies from moderate to poor. Two rather distinct assemblages of benthic foraminiferal assemblages were noted in this interval. An upper assemblage (Cores 31-41) includes diverse calcareous forms (up to 150 species) which inhabit bathyal depths above the CCD, whereas the lower assemblage (Cores 41-49) is dominated by agglutinated species. Benthic foraminifers are generally rare and poorly preserved in the lower section, with pronounced alternations in the percentages of calcareous and arenaceous species. This suggests numerous fluctuations of the CCD above and below the site of deposition.

Calcisphaerulids are found throughout the Maestrichtian-upper Cenomanian interval, and calcareous microorganisms of unknown affinity, first described by Sliter (1977), are present in Cores 44-47.

Cores 50-56 contain abundant calcareous microfossils dated as middle to early Albian (*Prediscosphaera prediscosphaera* calcareous nannofossil Zone). The change in sedimentation rate from the Turonian to the middle Albian is striking (Fig. 7) and indicates a highly condensed section with hiatuses in and about the Coniacian-late Albian interval. This was interpreted by DSDP Leg 36 scientists as a time of rapid subsidence of the Falkland Plateau, from shelf-break depths down close to its present depth below sea level, as a result of seafloor spreading.

This interpretation seems to be substantiated by the invertebrate macrofossils which became common in the shallow-water deposits below the Campanian-Coniacian interval. Fragile bivalves are present in Core 49, Section 3, and become very abundant in the Albian section below, where the genus *Aucellina* is identified. Solitary corals appear in Sections 4, 5, and 6 of Core 49 and also in Core 50. Fragments of the bivalve *Inoceramus* are the only mollusk shells found with any regularity in the Santonian-Coniacian section. These first appear in Core 41 and are present down the hole to Core 70.

Cores 50-70 sampled continuously an Albian-Upper Jurassic section nearly identical to that which was cored on an interval basis in DSDP Holes 327A and 330. Rich benthic foraminiferal faunas indicative of shelf-break depths (100-400 m) are present down to Core 57, which is dated by foraminifers as Albian (the *Ticinella roberti* Zone). Coccoliths occur sporadically in Core 57 and are assigned to the *Parhabdolithus angustus* Zone, which Thierstein (1973) dates as late Aptian to early Albian. Coccoliths become notably less common in the black shales, which are first encountered downhole in Core 56 (Lithologic Unit 6).

Planktonic foraminifers occur sporadically in Cores 58-62 but date this portion of the black shales as Barremian to Aptian. Mollusks date Cores 57 and 58 as Aptian and Cores 59-62 as Barremian. From Core 58 down to the fourth section of Core 62, abundant fragments of the shallow-water coccolith Micrantholithus obtusus (mid-Aptian or older) occur at selected intervals, but few other age-diagnostic calcareous nannofossils are present. No coccoliths positively indicate the presence of Hauterivian or older Cretaceous strata. Palynomorphs do indicate an Early Cretaceous age for Cores 58-61. Although spores are not numerous, species present are also found in Lower Cretaceous rocks of Argentina and South Africa which have been dated with various degrees of certainty as later Valanginian to Barremian in age. For this report, we adopt an age of Aptian to Neocomian for Cores 58-62.

Palynomorphs, coccoliths, and mollusks all indicate a Late Jurassic age for Cores 63-70. The pollen and spore assemblages, which abound with *Classopolis* pollen, are conditionally attributed to the Tithonian. The mollusks, which include the bivalves *Malayomaorica* and *Jeletzkyella*, the perisphinctid ammonite *Virgatosphinctes*, and the belemnite *Hibolithes*, indicate an early late Tithonian to a probable early Kimmeridgian age for Cores 63-1 to 67-3. Cores 67-4 to 70-5 did not yield any diagnostic macrofossils. Cores 63-70 contain the coccolith *Watznaueria britanica* with a well-developed small central area and bar. The co-occurrence of *Steph-anolithion bigoti* (short lateral spines) and *Vekshinella stradneri* in Cores 65-67 indicates an Oxfordian to early Tithonian age for that interval. The absence of *V. stradneri* and the presence of *S. bigoti* and *S. hexum* in Core 70 may suggest an early Oxfordian to Callovian age for that core. Dinoflagellates suggest an early Oxfordian to late Callovian age for Core 70.

In addition to the above, onychites (arm hooks associated with the belemnite animal) were encountered in foraminiferal washings from Cores 66–70. These had also been encountered in the Jurassic strata of DSDP Hole 330 (Cores 5–10).

Sedimentation Rates

The accumulation rates indicated in Figure 7 testify to the highly episodic nature of sedimentation on the Falkland Plateau. At any given locality, episodes of exceptionally high sedimentation separated by hiatuses or condensed intervals seem to be the norm for this area (compare with DSDP Leg 36 results). The Neogene and Mesozoic time scales used here were taken from Vail and Hardenbol (1979). The Plio/Pleistocene boundary in the Neogene was set at 1.8 Ma. The Paleogene time scale is from Hardenbol and Berggren (1978).

Foraminifers

Both planktonic and benthic foraminifers occur through the entire Site 511 section in almost all stratigraphic subdivisions of the Mesozoic and Cenozoic deposits, but their species diversity and abundance show significant fluctuations (Fig. 8). Foraminifers are missing or extremely rare at the base of the section (Jurassic-Aptian), in the dissolution facies at the Tertiary/ Cretaceous boundary, and in some beds within the upper Cenomanian-Campanian interval. Benthic foraminifers are rather scarce compared to planktonic forms. Nevertheless, their distribution through the section permits interpretation of the Mesozoic-Cenozoic paleoenvironmental changes in the area under study.

Cenozoic

Quaternary foraminifers are found in Core 1 (Samples 511-1-top to 511-1-2, 85-87 cm). This assemblage consists of abundant Globorotalia inflata, Globigerina pachyderma, G. bulloides, common Globorotalia truncatulinoides, and rare G. scitula. G. truncatulinoides is represented by low-conical specimens which are considered a cool-water form (subspecies?) of this species (Herb, 1968). A low-diversity benthic foraminiferal assemblage is present, consisting of the deep-water species Pyrgo murrhina, Hoeglundina elegans, Gyroidina soldanii, Laticarinina pauperata, and Cibicidoides wuellerstorfi; all are characteristic of the lower bathyal zone north of the Antarctic Convergence (Herb, 1971). Foraminifers are absent in the remaining siliceous oozes of Cores 1 and 2 (except Sample 511-2,CC).

Lower Oligocene assemblages of planktonic foraminifers are found in the diatomaceous oozes and nannodiatomaceous oozes of Core 2 (Sample 511-2,CC) to



Figure 8. Abundance, preservation, and species diversity of foraminifers at Site 511. (Abundance: R, rare; F, few; C, common; A, abundant. Preservation: P, poor, M, medium; G, good.)

Core 15. They are sparse and preservation is poor except in several samples (Cores 9, 11, 15) from the lower part of the section. They are represented by Globigerina angiporoides, G. aff. linaperta, G. prasaepis, G. officinalis, G. praebulloides, G. ouachitaensis, G. brevispira, Globorotaloides suteri, Globigerinita martini, Chiloguembelina cubensis, Globorotalia permicra, G. munda, and G. gemma. According to Blow's (1969) zonal scale, this assemblage limits the age of the sediments and is restricted to the *Globigerina tapuriensis* Zone through the *G. ampliapertura* Zone. Impoverished assemblages of planktonic foraminifers reflect cold-water conditions at high latitudes in the Southern Hemisphere. They resemble the lower Oligocene foraminiferal fauna of the southern part of New Zealand (the *G. angiporoides* Zone and the upper part of the *G. brevis* Zone according to Jenkins's [1971] zonal scale).

Samples of Core 16 contain very poor assemblages of planktonic foraminifers: *Globigerina* aff. *linaperta, G. labiacrassata, G. angiporoides,* and *Globorotaloides suteri*. This interval is tentatively assigned to the lower Oligocene because of the absence of the typical upper Eocene species.

In the interval from Sample 511-17-1, 90-92 cm to Sample 511-18-2, 27-29 cm, sediments also include an impoverished assemblage of planktonic foraminifers: *Globorotaloides suteri, Globigerina linaperta, G. angiporoides, G.* aff. galavisi, G. praebulloides, Globigerinita martini, and Globorotalia munda. In Sample 511-17-1, 90-92 cm, however, one specimen of Globigerapsis index was found. If it is in situ, sediments should be attributed to the uppermost part of the upper Eocene.

Sample 511-18, CC and Core 20 are characterized by the low-diversity planktonic foraminiferal assemblage composed of *Globigerina angiporoides*, *G. linaperta*, *G. tripartita*, *Globigerinita pera*, *Globorotaloides suteri*, *Chiloguembelina cubensis*, and *Globigerapsis index*. The presence of the latter testifies to the late Eocene age of the sediments. The upper Eocene/lower Oligocene boundary is placed just above Sample 511-17-1, 90-92 cm, the last sample containing *G. index*. These deposits are correlated with the tropical *G. semiinvoluta*, *Globorotalia cocoensis*, and *G. cunialensis* zones and with the *Globigerina linaperta* and *G. brevis* zones (lower part) of New Zealand.

The upper Eocene-lower Oligocene portion of the section is characterized by a rather diverse assemblage of benthic foraminifers consisting of about 80 species. The quantity of specimens is low in these sediments owing to dissolution and dilution by radiolarians and diatoms; most species are very rare and irregular in occurrence. Preservation is poor.

The dominant benthic species occurring throughout the whole upper Eocene/lower Oligocene section are Fissurina sp., Guttulina adhaerens, Pullenia bulloides, P. quinqueloba, Cibicidoides sp., Oridorsalis umbonatus, Nonion havanense, Anomalinoides spissiformis, Stilostomella curvatura, S. bradyi, S. antillea, and Bulimina sp. Common are also Karreriella sp., Pyrulina cylindroides, Gyroidina planulata, G. zealandica, Orthomorphina rohri, and O. glandigena.

The presence of these foraminifers is consistent with the accumulation of Paleogene diatomaceous and nanno-diatomaceous oozes at lower bathyal water depths. Poor preservation of benthic and planktonic species shows that the foraminiferal lysocline probably occupied the same position.

In Cores 21 and 22 and Sample 511-23-1, 3-4 cm, foraminifers are absent (dissolution facies).

Mesozoic

Comparatively rich planktonic foraminiferal assemblages characterize calcareous sediments in the interval from Sample 511-23-1, 27-29 cm to Sample 511-26.CC: Globigerinelloides multispinatus, G. impensus, Schackoina multispinata, Hedbergella holmdelensis, Rugoglobigerina pilula, R. rotundata, R. pustulosa, Heterohelix reussi, H. pulchra, H. globulosa, and H. glabrans. All these species are well developed in the Campanian-Maestrichtian. The influence of Austral climates is evidenced by the absence of representatives of the Globotruncana and by many species of heterohelicids. As a result, neither zonal subdivision of these sediments nor an exact age determination within the Campanian-Maestrichtian interval are possible. Taking into account the age of the underlying deposits, which are Campanian, it is probable that the calcareous oozes of Cores 23-26 belong only to the Maestrichtian. Calcisphaerulids are common in many samples.

The upper Campanian(?)-Maestrichtian calcareous sediments contain a very monotonous, highly diverse assemblage of benthic foraminifers mainly dominated by Gaudryina healyi, G. pyramidata, Spiroplectammina complanata, Ramulina aculeata, Marginulinopsis texaensis, Dentalina gracilis, D. legumen, Globulina subsphaerica, Lenticulina muensteri, Pullenia coryelli, P. americana, Valvulineria brotzeni, V. allomorphinoides, Gyroidinoides gudkoffi, G. nitidus, G. quadratus, Quadrimorphina camerata, Gavelinella beccariiformis, Globorotalites spineus, and Reussella szajnochae. The abundance of these moderately well preserved benthic foraminifers dominated by calcareous species leads us to conclude that they inhabited depths similar to their modern counterparts, that is, depths above the foraminiferal lysocline.

Sediments of Core 27 and most of Core 28 (to Sample 511-28-5, 90-92 cm) represent a dissolution facies and contain a very low diversity assemblage of Upper Cretaceous benthic foraminifers, consisting of rare Dentalina gracilis, Ramulina aculeata, Valvulineria brotzeni, Osangularia corderiana, Globorotalites michelinianus, and Bandyella aff. greatvalleyensis.

Comparatively diverse and quite different assemblages of planktonic foraminifers were encountered in sediments of Campanian age (from Core 29 to Sample 511-41-3, 55-57 cm). They include Globotruncana linneiana, G. cretacea, G. marginata, G. bulloides, G. globigerinoides, Globigerinelloides asperus, G. bollii, Archaeoglobigerina blowi, Heterohelix globulosa, H. reussi, H. pulchra, H. ramsayensis, Hedbergella crassa, H. spp. Specimens of Globotruncana arca and G. plummerae are rare but very important for age determinations. In general, hedbergellids and heterohelicids, together with some representatives of the genus Globigerinelloides, strongly predominate and are accompanied by comparatively rare Globotruncana.

Sediments within the interval from Sample 511-41,CC to Sample 511-43,CC belong to the Santonian. Among planktonic foraminifers are numerous low- and high-spired specimens of *Archaeoglobigerina bosquensis*, ac-

companied by Whiteinella baltica, Globotruncana pseudolinneiana, G. marginata, G. cretacea, Globigerinelloides asperus, Hedbergella crassa, H. spp., Heterohelix

reussi, H. globulosa, and H. pulchra. To the undifferentiated Coniacian-Santonian are assigned sediments within the interval from Sample 511-44-1, 44-46 cm to Sample 511-46, CC. Whiteinella baltica and Archaeoglobigerina bosquensis occupy a subordinate position; the genus Globotruncana is represented only by G. pseudolinneiana and G. marginata. Very common are specimens of Hedbergella spp., Heterohelix globulosa, H. reussi, and Globigerinelloides asperus. Schackoina cenomana is present but very rare in almost every sample.

The planktonic foraminiferal assemblage from Sample 511-47-3, 44-46 cm to Sample 511-47-6, 44-46 cm includes rare specimens of small Globotruncana pseudolinneiana (the G. lapparenti type) together with Schackoina cenomana, Hedbergella aff. agalarovae, Globigerinelloides asperus, Heterohelix globulosa, and Globotruncanella inornata. If these sediments correspond to the level of the evolutionary appearance of Globotruncana pseudolinneiana, they belong to the upper Turonian.

In sediments within the interval from Sample 511-47,CC to Sample 5-48-1, 35-37 cm, double-keeled Globotruncana species are absent. The assemblage of planktonic foraminifers consists of rather numerous Praeglobotruncana aff. oraviensis (with a spinose test), together with Hedbergella planispira, H. bornholmensis, H. aff. holtzi, Schackoina cenomana, Globigerinelloides asperus, Heterohelix globulosa, H. reussi. The age of the sediments is Turonian, probably early Turonian.

Sediments within the interval from Sample 511-48-1, 70-72 cm to Sample 511-49-5, 102-104 cm are characterized by Hedbergella praehelvetica, H. planispira, H. delrioensis, H. portsdownensis, H. infracretacea, Heterohelix sp. (a very small form), Globigerinelloides eaglefordensis, and Schackoina cenomana. In Sample 511-49-5, 102-104 cm, very rare specimens of Praeglobotruncana turbinata were found. This last species indicates a late Cenomanian age for the sediments (although an early Turonian age is not excluded).

The distribution of planktonic foraminifers in the Upper Cretaceous sediments is very inconsistent. Many samples from Cores 30, 32–36, 40–44, and 47–49 contain few to common tests of moderate preservation; in other layers, poorly preserved tests or barren intervals (dissolution facies) were observed.

The Coniacian-Santonian assemblage of planktonic foraminifers at Site 511 closely resembles that of the high latitudes of the Northern Hemisphere (the lower Senonian microfauna of Bornholm Island, Denmark [Douglas and Rankin, 1969]).

Calcisphaerulids occur throughout the Maestrichtian-upper Cenomanian section, and calcareous microorganisms of unknown affinity (Sliter, 1977, plate 14) were found in Cores 41-47.

The upper (Cores 28-37) and lower (Cores 38-48) parts of the Campanian-upper Cenomanian section are characterized by two different ecological assemblages of

benthic foraminifers. Cores 28-37 contain more than 150 species which usually inhabit bathyal water depth above the CCD. Many of these species occur in sediments very rarely and irregularly; others, such as Dentalina legumen, Lenticulina muensteri, Gyroidinoides nitidus, Globorotalites michelinianus, G. sp., Gavelinella stephensoni, Valvulinelia brotzeni, and Pleurostomella obtusa, are present constantly. Common are also Dentalina gracilis, D. catenula, D. basiplanata, Astacolus jarvisi, P. subnodosa, and Bolivina watersi. In Cores 31-37 the assemblage of benthic foraminifers is even more diverse. In addition to the species just mentioned, Dorothia trochoides, Lingulina pygmaea, Globulina lacrima, Charltonina sp., Conorotalites sp., P. torta, Ellipsopleurostomella curta, and Bulimina reussi are common.

The lower part of the section (Cores 38-48) is characterized by an assemblage which includes most species in the overlying sediments, but agglutinated species are dominant. Glomospira corona, G. gordialis, G. irregularis, Ammodiscus sp., Silicosigmoilinella sp., Hyperammina elongata, H. gaultina, H. friabilis, Rzehakina epigona, Ammobaculites echinatus, and Spiroplectammina sp. are also present sporadically. The species composition and diversity are quite variable in this interval, especially in the lower part. Assemblages where calcareous species dominate, testifying to conditions above the CCD, alternate with assemblages consisting mostly of arenaceous species, indicative of sedimentation below the CCD. The quantity of benthic foraminifers and their state of preservation also vary throughout the section. These foraminifers are usually rare or very rare and poorly preserved; only in some samples of Cores 30-34, 36, 41, 42, and 47 are they common to few, with moderate preservation.

Rather rich assemblages of planktonic foraminifers are characteristic of Albian sediments recovered in the interval from Sample 511-49-5, 120-122 cm to Sample 511-57-6, 11-13 cm. They include moderately preserved tests of Hedbergella delrioensis, H. trocoidea, H. planispira, H. infracretacea, H. brittonensis, H. amabilis, and H. globigerinellinoides. Sediments from Sample 511-55-6, 34-36 cm down to Sample 511-57-6, 11-13 cm contain the lower Albian species Ticinella roberti, T. aff. primula, H. trocoidea, H. infracretacea, H. delrioensis, H. planispira, and Globigerinelloides gyroidinaeformis; these belong to the T. roberti Zone. Foraminifers are abundant to common with some exceptions in Cores 49-54 and 56; preservation is generally moderate. In Sample 511-49-5, 120-122 cm, among the common Albian planktonic foraminifers, very rare specimens of T. raynaudi were identified; these indicate late Albian age for the sediments.

Calcisphaerulids are abundant to common in Cores 50-51 (down to Sample 511-51-2, 60-62 cm), rare in Samples 511-51-3, 60-62 cm to 511-51,CC, disappear completely in Cores 52-53, and again are present in some sample of Cores 54-55. They are not present in Cores 56-57.

Albian sediments (Cores 49-57) contain a moderately well preserved and diverse benthic foraminiferal fauna consisting mostly of shallow-water species ("shallowwater" is defined as 100-400 m, shelf-break fauna, after Sliter, 1977): Dorothia trochoides, Glomospira corona, G. gordialis, Clavulina gabonica, Lenticulina sp., Uvigerinammina jankoi, Osangularia utaturiensis, Patellinella australis, Spirillina elongata, Gyroidinoides primitiva, Anomalinoides indica, Gavelinella intermedia, and Pleurostomella obtusa. Less common are D. gradata, Gaudryina dividens, L. turgidula, Hoeglundina cretacea, S. minima, Conorboides minutissima, Tribrachia australiana, Marginulina sp., and Praebulimina sp. This assemblage is typical for the Albian deposits of the Southern Hemisphere (Austral fauna) (Scheibnerová, 1974; Sliter, 1977).

Cores 58-62 are characterized by a low-diversity assemblage of planktonic foraminifers composed of *Hedbergella infracretacea*, *H. globigerinellinoides*, *H. aff. delrioensis*, *H. similis*, *H. gorbachikae*, *H. sigali*, *H.* sp., and *Globigerinelloides ferreolensis*; the assemblage suggests a Barremian-Aptian age for the sediments. The Aptian/Albian boundary is drawn below the first occurrence of *Ticinella roberti* and *T.* aff. *primula* in Sample 511-57-6, 11-13 cm. Benthic foraminifers are practically absent in Cores 58-62 sediments except in Samples 511-61,CC and 511-62,CC, where tiny, thin-walled Astacolus, Gyroidinoides, Gavelinella, Frondicularia, and Marginulina are recognized.

Cores 63-70 (Upper Jurassic section) are barren of foraminifers.

Calcareous Nannofossils

Coccoliths were present in sparse to moderate numbers throughout most of the Cenozoic-Mesozoic section of Hole 511, becoming abundant only in the Maestrichtian-Campanian and mid-Cretaceous intervals. Preservation was generally poor to moderate because of dissolution, although exceptionally well preserved assemblages were noted in the upper Eocene, Campanian, and Albian. Diversity was generally low because of the high latitudine of the site, the problem of dissolution, and in the case of the Aptian-Jurassic, a highly restricted (euxinic) depositional environment.

A short (10-cm) interval (repeated by a double hit of the drill string) at the top of Section 511-1-1, belongs to the *Emiliania huxleyi* Zone; the rest of that core is barren of coccoliths. Cores 2 and 3 contain a limited flora characterized by the presence of *Reticulofenestra umbilica* and belong to the *R. daviesii* Zone whereas Sections 1-3 of 511-4 contain *Clausiococcus fenestratus* and are assigned to the zone of that name. These two zones are equivalent to the *R. placomorpha* Zone of New Zealand (Edwards, 1971). Rare discoasters are present as well as a few *Sphenolithus moriformis* and *Zygrhablithus bijugatus*. Other taxa commonly present are *Chiasmolithus altus*, *R. bisecta*, and *Coccolithus pelagicus*.

Sample 511-4,CC to Core 16 contains an assemblage similar to that just noted plus *Ismolithus recurvus* and can be assigned to the early Oligocene *Blackites spinosus* Zone of Edwards (originally called the *B. rectus* Zone; see Edwards, 1971). A few rhabdoliths are seen in this interval, as well as *Pontosphaera multipora* and *P. pulcheroides*.

Cores 17-20 contain common and exceptionally well preserved coccoliths of late Eocene age. These are assigned to the *Reticulofenestra oamaruensis* Zone of Edwards (1971) because of the presence of *R. oamaruensis* and *Discoaster saipanensis*. Other members of this assemblage include *Ismolithus recurvus*, *R. bisecta, Chiasmolithus oamaruensis*, *D. tani* (few), *C. altus*, and *Coccolithus pelagicus*. Core 21 is barren of coccoliths.

Core 23 contains an early Maestrichtian nannoflora dominated by the cool-water Biscutum coronum and can be assigned to Wind's (1979) proposed zone of the same name. Other species include B. dissimilis, Arkhangelskiella specillata, Gartnerago costatum, Eiffellithus turriseiffeli, Prediscosphaera cretacea, Nephrolithus corystus (mostly rims), Kamptnerius magnificus, Ahmuellerella octoradiata, and members of the Lucianorhabdus/Phanolithus plexus. Core 24 contains a somewhat different assemblage with common Monomarginatus pectinatus, Misceomarginatus pleniporus, Marthasterites inconspicuus, and Biscutum coronum. This assemblage is also assigned to the B. coronum Zone of Wind (1979), but the core may be early Maestrichtian or latest Campanian in age (see Wind and Wise, this volume).

The interval from Core 25 through Section 511-27-2 was barren of coccoliths. Section 511-28-1 through Core 42 contained Marthasterites furcatus but no Lithastrinus floralis and, considering the foraminifers present, is mostly dated as early Campanian in age. L. floralis is present downhole starting in Core 43, indicating an age of mid-Santonian or older. Other common species within this interval are Watznaueria barnesae, Micula decussata, Biscutum constans, Eiffellithus eximius, E. turriseiffeli, Seribiscutum primitivum, Kamptnerius magnificus, Cretarhabdus conicus, and Prediscosphaera cretacea. E. trabeculatus makes a brief appearance in Cores 39 and 40 (warm interval?). Thiersteinia ecclesiastica n. gen., n. sp., is present from Section 511-46-1 to Section 511-48-1 and is used to define a new zone.

Micula decussata and Kamptnerius magnificus are present down to the uppermost sample of Core 48 (middle Turonian or younger); the other core sections are barren. These species are absent in Core 49 (lower Turonian or older).

Section 511-50-2 to Section 511-56-1 contain an assemblage similar to Core 49 except that *Eiffellithus turriseiffeli* is absent; thus this interval is assigned to the *Prediscosphaera cretacea* Zone of Thierstein (1973). Core 51 contains common *Sollasites falklandensis* and overlies the subzone of the same name.

Prediscosphaera cretacea is absent below Core 56, Section 1, but Lithastrinus floralis is present down through Section 511-58-2; therefore this assemblage belongs to the Parhabdolithus angustus Zone which Thierstein (1973) dates as middle Aptian to early Albian in age. Cores 58 to 62, Section 1, do not contain L. floralis but in selected intervals are flooded with fragments of Micrantholithus obtusus. The rest of Core 62 is essentially barren of coccoliths except for the core catcher. Sample 511-62, CC to Core 70 contains Watznaueria britanica (small central area and bar generally present), Zygolithus erectus, and other forms described from the Jurassic of DSDP Hole 330, Leg 36 (Vekshinella stradneri, Ethmorhabdus gallicus, and Stephanolithion minutum with short lateral spines). This assemblage is found in the Volgian of the Russian platform (Medd, in press). S. hexum and S. bigoti co-occur in Core 70, which would indicate a Callovian-early Oxfordian age for the base of the drilled sequence.

Radiolarians

Radiolarians are common and well preserved in Cores 1 through 20 at Site 511. Cores 21, 23, 56, and 57 also contain radiolarians but they are sparse and generally recrystallized. All other cores examined are barren of any siliceous microfossil remains. Only core-catcher samples were studied at Site 511.

In the absence of any well-established, high-latitude, Paleogene radiolarian zonation, the cored sequence between Cores 2 and 20 is equated to Riedel and Sanfilippo's (1978) standard low-latitude zonal scheme.

Age Determinations

Age determinations can be summarized as follows:

Core 23:	Late Cretaceous
Core 21:	middle to early Eocene
Cores 17-20:	late Eocene
Cores 5-16:	early Oligocene
Cores 2-4:	early Oligocene-possibly early late
	Oligocene
Core 1:	mixed Pleistocene and Oligocene

Core 23 is dated as Late Cretaceous because several radiolarian genera known to be restricted to this time period occur in it. No identifications to the specific level are possible because the specimens are very poorly preserved. Most radiolarians are represented as internal casts composed of glauconite or are entirely replaced by this mineral.

Core 21 is tentatively dated as middle to early Eocene, because the *Amphymenium splendiarmatum* and *Amphicraspedum prolixum* group co-occur in the presence of many specimens of *Buryella* species. Nearly all specimens are recrystallized to some degree and most cannot be identified to the specific level.

Cores 20–17 are late Eocene in age, equivalent to the *Thyrsocyrtis bromia* zonal equivalent at low latitudes. This age determination is based upon the occurrence of *Theocyrtis* sp., the ancestral form of *Theocyrtis tuber*-osa known only to occur in the upper half of the *T. bromia* Zone at low latitudes. In addition, the co-occurrence of *Cryptoprora ornata, Theocyrtis diabloensis*, and *Lychnocanoma amphitrite* within these cores indicates a late Eocene age.

Cores 17 through 5 are early Oligocene, equivalent to the *Theocrytis turberosa* Zone at low latitudes. This determination is based upon the occurrence of *Calocycletta acanthocephala* throughout the sequence of cores. This radiolarian is known only from sediments of the early Oligocene age (*T. tuberosa* Zone) in mid- to lowlatitude regions (DSDP, Legs 22 and 31). Johnson (1974) found *C. acanthocephala* restricted to *T. tuberosa* Zone sediments in the Indian Ocean (DSDP Site 216, Leg 22) and Ling (1975) also encountered this species in the lower half of the *T. tuberosa* Zone at DSDP Site 292, Leg 31, in the Northwest Pacific-Japan Sea area.

Cores 4 to 2 contain *Calocycletta* cf. *parva*. This species is believed to be the ancestral form of *C. parva*, which appears in the late Oligocene in mid- to low-latitude areas. Dinkelman (1973) reports *C.* cf. *parva* restricted to the early Oligocene *Theocyrtis tuberosa* Zone, but Moore (1972) believes this species may range into the basal *Dorcadospyris ateuchus* Zone, or early late Oligocene. Therefore, Cores 4 through 2 are dated as probable early Oligocene, but possibly early late Oligocene in age.

Core 1 contains a homogenized mixture of Pleistocene and early Oligocene radiolarians, indicating that the integrity of this interval is disrupted.

Paleoclimatic Interpretation

The presence in Cores 20 through 16 of *Calocycletta* species of Orosphaerid and Collosphaerid radiolarians, and of *Cryptoprora ornata, Theocyrtis* sp. (ancestral *T. tuberosa*), *Lychnocanoma amphitrite, Lithocyclia crux,* and *Thyrsocyrtis bromia* (very rare) are indicative of relatively warm conditions, probably cool temperate waters, in the vicinity of Site 511 during the latest Eocene and early Oligocene interval.

However, by Core 12 the dominant species in the overall radiolarian assemblage have changed. Most warm-water elements such as the *Calocycletta* have disappeared, and the assemblage becomes dominated by cooler-water *Spongoplegma*, *Prunopyle*, and Eucyrtid species. This assemblage persists through Core 11 and is considered transitional, indicative of significant upwelling conditions during climatic deterioration within the earliest Oligocene. The convergence of cool temperate and subantarctic waters and the eventual transgression of subantarctic waters over Site 511 are proposed. However, the exact timing of this event cannot be ascertained from radiolarian data.

Between Cores 11 and 9 further climatic deterioration is evidenced by the radiolarian taxa encountered. The radiolarian assemblage in Cores 9 through 2 generally resembles the early Oligocene assemblage recorded by Chen (1975), from Leg 28, Site 274, at 69°S on the Ross Sea continental margin. Cyclampterium(?) longiventer, Spongomelissa fenestrata, Lithomelissa challengerae, S. sphaerocephalis, S. sp., and Eucyrtidium sp. are significant constitutents of this radiolarian assemblage. Chen (1975) encountered this assemblage in Cores 31 through 39 at Site 274, and owing to the proximity of Site 274 to the Antarctic continent, it must be considered antarctic in character. Radiolarians in Cores 9 through 2 at Site 511 appear to represent the northerly elements of a broad antarctic/subantarctic paleobiogeographic province encompassing 18° of latitude in the circum-Antarctic seas during the early Oligocene.

By Core 5 at Site 511, a slight moderation of climate is evidenced by the reappearance of *Calocycletta* and warm-water Eucyrtid species. A slight southerly shift of warmer subantarctic waters over Site 511 seems feasible to explain this faunal modification.

In summary, climate through the latest Eocene and early Oligocene at Site 511 fluctuates from cool temperate to broadly subantarctic/antarctic in character. The exact timing of this transition cannot be ascertained from radiolarian data; it is, however, an early Oligocene event.

Diatoms

Diatoms are well preserved and abundant in Core 1. A detailed examination of this core revealed the presence of three disconformities. The siliceous gravelly sand and foraminiferal ooze from 0-37 cm of Section 1 is assigned to the *Coscinodiscus lentiginosus* Zone of the Brunhes Magnetic Chronozone.

The lithology change at 37 cm marks a disconformable boundary separating the upper *Coscinodiscus lentiginosus* Zone from the *C. elliptipora/Actinocyclus ingens* Zone encountered at 511-1-1, 52-53 cm. Sample 511-1-1, 91-92 cm is assigned to the *C. kolbei/Rhizosolenia barboi* Zone of the uppermost Pliocene.

No zonal designation is given to Sample 511-1-1, 120-121 cm because of considerable downhole slumping of younger microfossils and reworking of lower Pliocene and upper Miocene microfossils. This sample is probably close to an apparent disconformity separating lower Matuyama Chronozone sediments (511-1-1, 91-92 cm) from upper Gauss Chronozone sediments (511-1-1, 142-143 cm). Sample 511-1-1, 142-143 cm contains a flora characteristic of the *Nitzschia weaveri* Zone. Samples 511-1-2, 31-31, 65-66, and 89-90 cm are assigned to the *N. interfrigidaria/Coscinodiscus vulnificus* Zone of the mid-Gauss and are apparently conformable to the *N. weaveri* Zone of basal Section 1.

The third disconformity in Core 1 probably coincides with the sharp lithology change at Section 2, 144 cm. This sharp boundary between the siliceous gravelly sand above and the diatom ooze below separates the middle Pliocene of Sample 511-1-2, 89–90 cm from the lower Oligocene of Sample 511-3-1, 4–5 cm.

Abundant and diverse diatom assemblages of early Oligocene to late Eocene age occur in sediments recovered in Core 1, Section 3 through Core 20, Section 3. Six diatom zones have been identified in this interval (see Gombos and Ciesielski, this volume). The interval from Sample 511-1-3, 132-134 cm through Sample 511-4-3, 10-12 cm contains part of the Coscinodiscus superbus group Zone; that from Sample 511-5-1, 10-12 cm through Sample 511-11-2, 30-32 cm contains the Rhizosolenia gravida Zone; Sample 511-11-3, 30-32 cm through Sample 511-12-1, 11-13 cm contains the Brightwellia spiralis Zone; Sample 511-12-2, 11-13 cm through Sample 511-15-1, 13-15 cm contains the Melosira architecturalis Zone; Sample 511-16-1, 22-24 cm through Sample 511-16-2, 22-24 cm contains the Asterolampra insignis Zone; and Sample 511-17-1, 83-85 cm through Sample 511-20-3, 23-25 cm contains the Rylandsia inaequiradiata Zone.

In Hole 511, the Eocene/Oligocene boundary occurs somewhere within or between Cores 16 and 17; that is, close to the top of the *R. inaequiradiata* Zone. This boundary cannot be closely defined on the basis of any definite diatom appearance or extinction datum. However, a significant change in the relative abundance of the genus *Pyxilla* was observed in the vicinity of the boundary. In Core 20 (below the boundary) the abundance of this genus is about 1500-2000 specimens per 2 traverses at $\times 400$ magnification. Above Core 20 the abundance decrease by an order of magnitude. It is not certain if this dramatic abundance change reflects actual oceanographic conditions or if it is a preservational phenomenon.

Sporadic specimens of reworked Eocene diatoms were noted throughout the hole.

Silicoflagellates

Silicoflagellates are common and well preserved in most sediments examined from Cores 1–20; below Sample 511-20,CC, they are absent. Species diversity is low in Pliocene to Recent sediments and high in lower Oligocene to upper Eocene sediments.

Samples 511-1-4, 15-16 cm, through 511-3-3, 65-66 cm, are assigned to the revised Naviculopis constricta-Corbisema archangelskiana Interval Zone of the lower Oligocene (see Shaw and Ciesielski, this volume). The dominant silicoflagellate flora throughout this interval are Naviculopsis trispinosa, N. biapiculata, Distephanus crux, and Dictyocha deflandrei. Other species commonly present in this zone include Corbisema apiculata, C. flexuosa, C. geometrica, Dictyocha aspera martinii, Distephanus crux loeblichii, and others.

Core 3 below Sample 511-3-4, 65-66 cm, Cores 4 through 7, and Core 9 through Sample 511-9-4, 80-81 cm contain a newly defined Naviculopsis constricta/ Dictyocha deflandrei concurrent Range Zone. The wellpreserved and diverse assemblage in these lower Oligocene sediments includes N. constricta, N. biapiculata, N. trispinosa, Corbisema geometrica, C. apiculata, Dictyocha aspera martinii, Mesocena occidentalis, and others. Sediment recovery in Core 8 was poor and the core was not sampled for study.

Sample 511-9-5, 80-81 cm through Sample 511-11-4, 5-6 cm is placed in the newly defined *Naviculopsis trispinosa* Zone of the lower Oligocene and is characterized by the abundant occurrence of *N. trispinosa*. The first consistent occurrence of *Dictyocha deflandrei* marks the top of this zone; only few to rare *D. deflandrei* occur sporadically within the zone. Other species common in this interval include *Corbisema hastata globulata*, *C. triacantha*, *Mesocena apiculata*, *M. occidentalis*, *N. biapiculata*, and *N. constricta*.

Sample 511-11,CC through Sample 511-18,CC is placed in the newly defined *Mesocena occidentalis* Zone. The rich silicoflagellate assemblage of this upper Eocene to lowermost Oligocene interval includes common *Corbisema hastata globulata*, *C. triacantha*, *Distephanus boliviensis*, *D. crux crux*, *M. apiculata*, *M. occidentalis*, *Naviculopsis biapiculata*, and *N. constricta*. *N. trispino-* sa is present (<10%) but is significantly less abundant than in younger strata of this site.

Sponge Spicules

At Site 511 sponge spicules were recognized in Paleogene sediments (Cores 2-20); in Mesozoic deposits they are totally absent. In Cores 2-20 the assemblage is represented mostly by spicules of glass sponges (class Hyalospongiae) with a compact skeleton. In much smaller numbers are isolated macroscleres and microscleres. Rare are spicules of tetraradiate sponges (order Tetraxonida), the number of which increases considerably in the upper Eocene deposits.

When morphological types are compared, two assemblages of sponge spicules are distinguished. The first characterizes the lower Oligocene deposits (from Sample 511-2-1, 98-100 cm to Sample 511-16,CC) The second association is present in upper Eocene sediments (Cores 17-20). Rather numerous are representatives of the order Tetraxonida: dichotriaenes, anatriaenes, oxeas, strongyles, styles. Fossil microscleres such as sigmas occurred rarely, and only in the upper Eocene.

Pollen and Spores

Cenozoic

In Cores 1-3 (Quaternary to the uppermost layers of the lower Oligocene) spores and pollen are absent.

In Cores 4, 5, 6, 9, 11, 12 (lower Oligocene), pollen and spores are relatively diverse and numerous. Abundant are Gymnosperm pollens: Podocarpus sp., Dacrydium sp., D. cupressinum, Phyllocladus sp., P. mawsonii, Microcachryidites antarcticus, and Pinus sp. Angiosperms are represented by the pollens Nothofagus (brassi type and fusca type), Casuarina sp., Myrtaceae, Proteaceae, Combretaceae, Tricolpites sp., T. reticulata, Rhoipites minusculus, R. baculatus, Psilatricolporites sp. Among spores were identified Cyathea sp., Gleichenia sp., Anogramma sp., Matonisporites sp., Cyatheacidites sp., Leiotriletes regularis, Cingutriletes australis, Cicatricosisporites sp., Polypodiaceae.

In Cores 15-17 palynomorphs are few: Podocarpus sp., Dacrydium sp., D. cupressium, Phyllocladus sp., Nothofagus sp., N. sp. (fusca type), N. sp. (brassi type), Tricolpites sp., Triatriopollenites sp., Psilatricolporites sp., Polypodiaceae, Cyathea sp., Leiotriletes sp., Gleichenia sp., and Cicatricosisporites sp.

In Cores 18-20 only very rare grains of pollen and spores were found: *Podocarpus* sp., *Dacrydium* sp., *Phyllocladus* sp., *Nothofagus* sp., *Tricolpites* sp., *Cyathea* sp., *Leiotriletes* sp., and Polypodiaceae. Impoverished assemblages of pollen and spores testify only to the Eocene-Oligocene age of the sediments.

Mesozoic

In Upper Cretaceous deposits (Cores 29-49), spores and pollen grains are sporadic. They are missing in the greater part of Albian sediments (Cores 49-55). In underlying deposits of the Lower Cretaceous and Upper Jurassic, considerable amounts of satisfactorily preserved spores and pollen make it possible to subdivide sediments into three units with different palynological assemblages.

In Cores 56–58 Gymnosperm spores and pollen are present in relatively equal quantities. Sporadic grains of Angiosperm pollen of the genus *Clavatipollenites* were recognized. Among spores, *Gleicheniidites* spp. and *Cyathidites minor* are predominant. Very common are spores such as *Cicatricosisporites* spp., *Ceratosporites distalgranulatus*, *Coronatispora valdensis*, *Crybelosporites* spp., *Cyatheacidites tectifera*, *Muricingulisporis annulatus*, *Perotrilites linearis*, *Taurocusporites* sp., *Polypodiaceoisporites elegans*. *Classopollis* pollen predominates within the Gymnosperms. Pollen of *Cyclusphaera psilata* is also present.

This assemblage has some species common to a palynological association from the Baqueró continental formation (Santa Cruz Province, Argentina), the age of which is palynologically determined as Barremian-Aptian (Archangelsky and Gamerro, 1967). It is also similar to an assemblage, tentatively assigned an Aptian-Albian age, from deposits off the South African coast penetrated by Hole 361, Leg 40 (McLachlan and Pieterse, 1978) and to an Albian assemblage from Hole 327, Leg 36, on the Falkland Plateau (Hedlund and Beju, 1977). It is noteworthy that according to planktonic foraminifers, the age of the sediments at Site 511 is dated as upper Aptian (Core 58)-lower Albian, the *Ticinella roberti* Zone (Cores 57 and 56).

In Cores 59–62 the palynological assemblage is characterized by a predominance of Gymnosperm pollens; among them *Classopollis* is dominant. Pollens from *Cyclusphaera psilata* and *Inaperturopollenites limbatus* are constantly present. Angiosperm pollens are missing. Spores are not numerous, though represented by diverse genera and species such as *Antulsporites saevus*, *A.* granulatus, Contignisporites cooksoni, Cicatricosisporites spp., Dictyotosporites complex, Distaltriangulisporites sp., Foveosporites subtriangularis, Ischiosporites volkheimeri, Interulobites algoensis, I. sinosus, I. aff. triangularis, Matonisporites crassiangulatus, Nevesisporites sp., Polypodiaceoisporites elegans, Staplinisporites caminus, Taurocusporites segmentatus.

Though most spore species of this assemblage are peculiar to Upper Jurassic and Lower Cretaceous deposits of Argentina, South Africa, and Australia, presence of such typically Lower Cretaceous forms as Cyclusphaera psilata, Cicatricosisporites spp., Inaperturopollenites limbatus, Polypodiaceoisporites elegans, and Taurocusporites segmentatus enable us to date deposits as Early Cretaceous. The assemblage can be correlated with spores and pollen from the Agrio and Ortiz formations of Neuquén Territory in Argentina (Volkheimer and Sepulveda, 1976; Volkheimer et al., 1977), the age of which was determined by ammonites in the Agrio Formation as Hauterivian-Barremian. In addition, a similar assemblage was observed in deposits of the Sundays River Formation (Alboa Basin, South Africa). Its age was tentatively determined as late Valanginian-early Hauterivian on the basis of ostracodes, benthic foraminifers, and mollusks (Scott, 1976).

In Cores 63-70 the palynological assemblage abounds with Classopolis pollen. Spores are not numerous, being represented mainly by the same species and genera as in the previous assemblage. A peculiar feature is the absence of species typical of the Lower Cretaceous deposits only. From this standpoint the age of the assemblage is Late Jurassic. Most similar to this assemblage are those from the Vaca Muerta Formation (Argentina) of Tithonian age (Volkheimer and Quattrocchio, 1975) and the Kirkwood Formation (the Uitenhage Group, Algoa Basin, South Africa), supposedly of Late Jurassic-Early Cretaceous age (Scott, 1976). Similarity with the above formations is suggested by the absence of Cyclusphaera psilata and presence of some species of the genus Interulobites. As there are no data on Oxfordian and Kimmeridgian palynoassemblages from Argentina and South Africa, and no spores of Interulobites have been identified from the Oxfordian-Kimmeridgian of Australia (Filatoff, 1975), the assemblage at Site 511 in the interval from 556 to 632 meters is conditionally attributed to the Tithonian.

Mesozoic Dinoflagellates³

The interval between Sample 511-29-1, 0-3 cm and Sample 511-40-5, 34-36 cm is dated as Late Cretaceous (early Campanian to Santonian). Restricted to this interval are Isabeladinium latum and Chatangiella granulifera, Odontochitina porifera, Amphidiadema denticulata, and A. nucula. Xenascus ceratioides occurs at and below Sample 511-34-4, 124-126 cm; the last appearance datum (LAD) of Conosphaeridium striatoconus is in Sample 511-40-5, 34-36 cm. The underlying interval between Sample 511-41-3, 51-53 cm and Sample 511-47-1, 42-44 cm is Late Cretaceous (Coniacian to Turonian). C. striatoconus occurs throughout the interval, and the LAD of Palaeohystrichophora infusorioides is in Sample 511-42-4, 110-112 cm. Actinotheca aphroditae occurs only in Sample 511-42-4, 110-112 cm. No dinoflagellates were recovered from the interval between Sample 511-48-3, 42-44 cm and Sample 511-53,CC; the age is therefore indeterminate. The interval between Sample 511-54-3, 46-48 cm and Sample 511-58-2, 52-54 cm contains a low-diversity assemblage with Ovoidinium scabrosum, Angustidinium acribes, and ?Oligosphaeridium asterigerum. The interval is interpreted as Early Cretaceous (Cenomanian to Albian) because of this association. Samples from Cores 59 and 60 were not examined. The interval between Sample 511-61-1, 69-71 cm and Sample 511-67-2, 80-82 cm contains Endoscrinium luridum (in Sample 511-61-1, 69-71 cm only), Leptodinium mirabile (LAD in Sample 511-64-5, 45-47 cm), and Gonvaulacysta jurassica (LAD in Sample 511-67-2, 80-82 cm). The interval is dated as Late Jurassic (early Kimmeridgian to late Oxfordian) by this assemblage. Samples from Cores 68 and 69 are nondiagnostic. The deepest sample examined is at 511-70-1, 50-52 cm; it contains Wanaea digitata, Rigaudella sp. cf. R. aemulum, Endoscrinium galeritum; and Gonyaulacysta juras-

³ Additional report by D. K. Goodman and M. E. Millioud, Exxon Production Research Company, Houston, Texas.

sica, which indicate an early Late to late Middle Jurassic (early Oxfordian to late Callovian) age. Based on dinoflagellate data, the Jurassic/Cretaceous boundary is in the interval between Sample 511-58-2, 52-54 cm and Sample 511-61-1, 69-71 cm. No Cretaceous sediments older than Albian and no Jurassic sediments younger than early Kimmeridgian were identified.

Macrofauna

Cores 41-43 contain numerous but fragmentary remains of *Inoceramus* s.l., a few other indeterminate pelecypods, and locally, a number of generically indeterminate solitary corals (evidently redeposited, at least in part). The sediments were deposited in the bathyal to outermost neritic environment, judging by the general absence of such shelf-bound fossils as belemnites and thick-shelled pelecypods in the presence of depth-tolerant organisms such as *Inoceramus*.

In Cores 50-56 Aucellina species range from extremely numerous (coquinas) to rare—A. sp. cf. radiatostriata, A. sp. cf. andina, and A. sp. indet. The coquinas were deposited under the influence of strong bottom currents. The presence of other, apparently shelfbound pelecypods suggests an inner neritic origin, although an alternate interpretation can be made in favor of an outer neritic to bathyal environment.

Rare aconeceratid, cheloniceratic, and ancyloceratic ammonites of Aptian affinities are found in Sections 57-3 to 58-1. These are all depth-tolerant or nectoplanktonic ammonites which suggest a bathyal to outermost neritic environment similar to that of Cores 41-49.

Section 58-4 to Core 62 did not yield any immediately diagnostic macrofossils, although belemnites and buchiid pelecypods were present. No shells comparable to Aucellina were present; this fact could suggest a pre-Aptian age. A significant fluctuation in water depth is seen in this interval. Sections 58-4 to 60-4 were definitely deposited in an inner neritic to littoral environment in close proximity to the basin shoreline. This is indicated by abundant thick-shelled shallow-water pelecypods (often forming coquinas), including numerous oysters. Belemnites, which come from inner- to median-shelf environments at depths of less than 150 meters, are common, whereas ammonites, which favor outer-shelf to upper bathyal environments, are nearly totally absent. The alteration of belemnite-bearing and oyster-rich units indicates a frequent alternation between a normally saline environment and a distinctly brackish (presumably delta-influenced) environment.

The interval from Section 60-5 to Core 63, on the other hand, was deposited in the same bathyal to outermost neritic environment as Section 41-3 to Core 49 and Sections 57-2 to 58-3. This is suggested by the general paucity or complete absence of most fossils (pelecypods, belemnites, and large ornate gastropods). The depositional environment is one of very low energy and reducing conditions.

The hiatus between Core 62 and Core 63 probably includes the upper Tithonian as well as the Berriasian to Hauterivian (or possibly even Barremian) rocks. This is strongly suggested by the occurrence of the exclusively Late (but not Latest) Jurassic Malayomaorica sp. and/ or Jeletzkyella sp. in Core 63. A perisphinctid ammonite in Sample 511-63-3, 49-50 cm is identified as Virgatosphinctes (Pseudoinvoluticeras) sp. of early Tithonian affinities. The appearance of fairly common belemnites and a general increase of fossil content in Sections 64-1 to 65-2 suggest a shallowing trend which apparently brought the bottom of the basin into a mid- to outer neritic environment. This trend culminates in the appearance of numerous interbeds of a presumably inner neritic to littoral pelecypod coquina in sandy siltstone containing common to abundant pelecypods and some Hibolites-like belemnites (Sample 511-65-2, 90 cm). This abrupt lithological and faunal change marks the second episode downhole of inner neritic to littoral deposition at Site 511.

This second interval of shallow depositional environment is seen from Sections 65-2 to 65-3 in a distinctly cyclical alternation of fossil-rich sandy to pure siltstones with very thin interbeds of pelecypod coquina. The faunal composition of these units is exactly the same as in the Cretaceous (Sections 58-4 to 60-4), including numerous oysters in some coquinoid interbeds. The environment was much the same as in the Cretaceous.

The lowermost cores from Site 511 (Sections 67-4 to 70-5) consist of siltstone (black and petroliferous), rich in belemnite of the genus *Hibolites* but lacking thick-shelled shallow-water pelecypods. This suggests a low-to very low energy, mid-neritic (50-?150-meter) depositional environment.

PALEOMAGNETISM

Paleomagnetic measurements were made on Leg 71 sediments using samples taken by pushing plastic cylinders (2.5 cm in diameter and approximately 2.5 cm in length) into the sediment. These samples are oriented by a notch pointing in the uphole direction. Additional samples were taken in the harder sediments of Site 511 using a diamond-tipped drill, after an orientation arrow was inscribed on the sediment. Measurement of natural remanent magnetization (NRM) and postdemagnetization remanence were carried out using the shipboard Digico fluxgate spinner magnetometer, integrating the readings over up to 29 spins. Demagnetization was performed on board the Glomar Challenger by means of a Schonstedt alternating field demagnetizer. Samples were demagnetized along three mutually perpendicular axes at the peak field; then one further demagnetization was carried out at half the peak field with the third axis reversed with respect to the demagnetizer coils. This additional demagnetization reduces the effect of any anhysteretic remanent magnetization (ARM) which may build up in the sample.

At Site 511, 135 samples were taken for paleomagnetic study at intervals of approximately 3 meters in the more complete sections of the hole. Pilot demagnetization shows that the NRM is stable and probably represents a primary magnetization. Samples with stronger NRM intensities were demagnetized at 150 oersteds (Oe). Demagnetization produced changes in inclination of usually less than 10°. Gaps in sampling in the upper part of the hole prevent the identification of magnetic epochs within the Cenozoic. The NRM inclinations of Mesozoic sediments are shown in Figure 9. The majority of the samples are normally magnetized (i.e., have negative inclinations), representing the long Cretaceous Normal Interval. Three samples with low or reversed inclination occur in the lower part of the Cretaceous sediments. Only the zone between Samples 511-60-6, 44-46 cm and 511-60-6, 51-53 cm has been confirmed by resampling; however, it is possible that these sediments were deposited during the Serra Geral Mixed Interval (as defined by van Hinte, 1976). The NRM of these samples may comprise a reversed primary magnetization with a normal secondary overprint. Demagnetization of these samples was not carried out on board the Glomar Challenger because the NRM intensities were too low. Two reversed samples occur in Jurassic sediments (in Cores 63 and 68), but these have not been confirmed by resampling.

The upper boundary of the long Cretaceous Normal Interval is marked by the top of Core 27. The reversed sample in Core 35 has a very weak intensity and is not considered to represent a true geomagnetic reversal. The mean absolute inclination for the Mesozoic sediments is $53.5 \pm 21.9^{\circ}$, which would indicate a paleolatitude of 34.0° for the Falkland Plateau.

Intensity variation in Hole 511 sediments reflects lithology (Fig. 10). The Cenozoic and Maestrichtian oozes (Subunits 2A and 2B, and Unit 3) were predominantly low in intensity, ranging from 0.2 to 0.9 μ G. Intensity in the zeolitic clays is low (0.5–2.0 μ G) to a depth of about 320 meters, below which values increase to between 20 and 140 μ G. Sediments have a marked red color below 350 meters, reflecting a high iron oxide content; this red coloration is also marked in the nannofossil claystones and chalks of Unit 5, which have an intensity of 10–140 μ G. Intensities within the black shales (Unit 6) are low (0.1–1.0 μ G).



Figure 9. Variation in natural remanent magnetization (NRM) inclinations within the Mesozoic at Site 511.



Figure 10. Correlation between natural remanent magnetization (NRM) intensity and lithology at Site 511.

ORGANIC GEOCHEMISTRY

Sediments encountered in Site 511 were analyzed onboard ship for organic carbon, gases in gas pockets, and fluorescence and were subjected to a pyrolysis/fluorescence procedure (Heacock et al., 1970) to provide an estimate of hydrocarbon occurrence. Gas samples were taken wherever possible, although the gases were not severely pressured in any of the cores taken. Results of these shipboard procedures and analyses are discussed hereafter, followed by General Observations that take into account both the shipboard and shore laboratory studies (see Schaefer et al.; von der Dick et al.; Copelin and Larter; Deroo et al.; all this volume). In some instances, the more sophisticated shore laboratory analyses changed significantly the impressions of the maturation and migration patterns derived from the shipboard analyses (Ludwig et al., 1980).

Organic Carbon

Samples for organic carbon analyses were selected mainly in fine-grained and dark material. The values obtained should therefore reflect zones of enrichment in organic carbon. Samples were dried, crushed, treated with HCl, washed, weighed, and analyzed on a Hewlett Packard 185B CHN analyzer. Prior to this, the HP 185B was calibrated by different standards, and integrator counts were plotted against weight in milligrams of organic carbon. Errors and deviations in organic carbon analyses may be due to resistant carbonates (e.g., dolomite) and especially to weighing on a rolling and pitching vessel. Both the organic carbon values and the carbonate contents (determined by the carbonate bomb technique) are listed in Table 3.

Pyrolysis/Fluorescence

Pyrolysis/fluorescence should give a background for both the presence of heavy free hydrocarbons in a sediment and the generative potential of the sediment for releasing bitumen at stages of higher maturity when related to organic carbon content. The procedure was improved by injecting N_2 into the pyrolysis tube, thus preventing hydrocarbons from burning. The values obtained by this method seem to be more reliable.

The fluorescence units of pyrolyzed sediments of different types and ages are plotted against carbon content in Figure 11.

Gas Analyses

We noticed no highly pressured gas or large gas pockets in any of the cores. Gas analyses were run on a Carle-GC equipped with a thermal conductivity cell and a much more sensitive Hewlett Packard 5710 AGC, equipped with a dual FID system.

Fluorescence

Random samples were taken to look for fluorescence, but even the black shales showed no indications.

Discussion of the Shipboard Results

Organic carbon values of Site 511 range from 0.1% to 4.1% and can be roughly divided into two units.

Table 3. Organic carbon, nitrogen, and carbonate content, Hole 511.

			Corg Corrected	
Core/Section	N	Corg	for CO3	CO3-
(interval in cm)	(%)	(%)	(%)	(%)
2,CC, 7-9	0.08	0.165	0.12	30
2-4, 10-12	0.06	0.06	0.06	2
3-4, 40-41	0.1	0.7	0.7	
4-2, 118-120	0.1	0.9	0.9	
5-4, 90-92	0.1	0.62	0.62	
6-3, 90-92	0.95	0.44	0.43	2
11-3, 110-111	1.1	0.68	0.65	5
12-1, 80-82	0.12	0.95	0.94	1
15-1, 72-74	0.09	0.44	0.43	2
16-1, 90-92	0.09	0.59	0.58	2
17-2, 60-62	0.04	0.44	0.41	7
18-1, 70-72	0.07	0.37	0.33	11
20-2, 64-66	0.08	0.51	0.44	13
21-1, 15-17	0.05	0.13	0.13	0
24-3, 98-99	0.06	0.32	0.15	53
25-1, 103-104	0.04	0.03	0.03	1
28-1, 78-80	0.06	0.26	0.25	2
28-7, 0-1	0.08	0.62	0.61	1.5
31-3, 144-146	0.1	0.8	0.78	2
32-5, 146-148	0.1	0.85	0.82	3
33-2, 92-94	0.09	0.62	0.60	2.8
33-5, 106-109	0.07	0.6	0.59	2.0
34-2, 113-115	0.09	0.71	0.70	2.0
34-7, 51-54	0.08	0.67	0.54	20.0
36-2, 141-143	0.07	0.57	0.54	6
38-2, 41-42	0.09	0.69	0.67	2.9
39-3, 81-82	0.08	1.9	1.14	40
40-2, 10-11	0.09	0.63	0.63	
41-3, 36-37	0.06	0.18	0.18	0.5
43-1, 23-25	0.08	0.26	0.26	0.4
44-3, 51-53	0.08	0.34	0.25	25.5
45-2, 93-95	0.07	0.31	0.31	
46-3, 61-63	0.1	0.42	0.42	
50-3, 62-64	0.06	0.26	0.21	21.0
51-1, 141-143	0.07	0.1	0.1	0
53-5, 65-66	0.03	0.11	0.08	25.5
55-5, 140-142	0.06	0.27	0.26	3
55-6, 127-129	0.06	0.68	0.68	
56-2, 74-76	0.47	0.15	0.14	8
56-5, 56-58	0.5	0.51	0.51	
57-3, 100-102	0.1	1.7	1.7	
58-1, 72-74	0.07	0.63	0.63	
59-1, 86-88	0.2	4.1	2.8	35.0
61-2, 81-82	0.16	3.8	3.7	3
62-3, 54-56	0.14	2.55	2.55	=1
63-3, 16-17	0.19	3.3	3.3	
66-3, 92-94	0.2	4.2	4.1	2
68-2, 71-72	0.18	3.2	3.2	
69-3, 75-77	0.19	2.95	2.95	
70-4, 96-98	0.16	2.85	2.85	

Samples from the Tertiary and Cretaceous down to the Aptian contain only small quantities of organic carbon. This is in agreement with the general features of sediments exhibiting brighter and reddish colors and indications of bioturbation. Probably this type of organic matter belongs to a kerogen Type III originating from continental runoff of higher terrestrial plant debris and perhaps reworked coaly and/or oxidized sedimentary particles. If this is so, the low content of organic carbon is apparently a result of a low but relatively constant ratio of the flux of organic and mineral matter through the water column (von der Dick, 1979). Extreme low values at the Tertiary/Cretaceous boundary can be assigned to more variable conditions. Pyrolysis/fluorescence data (Fig. 11) are very low and seem to reflect primarily the crudity of the method. The sedimentary sequence completely changed when the Lower Cretaceous black shales were encountered. Organic carbon values average 3.5%.

Pyrolysis/fluorescence data exhibit an extreme increase compared to the claystones and muddy nanno-



Figure 11. Fluorescence of pyrolyzed sediments of different types and ages plotted against carbon content (shipboard measurements).

fossil chalk in the upper part of the section (Fig. 11), thus reflecting the high potential of the black shales for generating oil. Although very crude, this quick method seems to be a useful indicator for any hydrocarbonbearing sediments and accumulations, whether free or bound to the kerogen matrix.

When Core 60 was drilled at about 530 meters and retrieved, an intense odor of oil was present. This odor was most intense in Cores 60–63 and dropped to moderate intensity in the subsequent cores. Methane was detected first in Core 41, and ethane was initially encountered in Core 54, where it exhibited a C_1/C_2 ratio of 400. This ratio decreased to a minimum value of $C_1/C_2 = 93$ in Core 61 (Fig. 12) and fluctuated between 130 and 210 in the deeper parts of the section. The same trend could be traced in the C_2/C_5 and C_2/C_3 ratios (Fig. 12). In general, low C_1/C_2 , C_2/C_3 , and C_2/C_5 ratios together with strong odor and high fluorescence values indicate a high degree of maturity. However, an objective determination of the thermal evolution is necessary; this was done in shore laboratory studies described hereafter.

The upper part of the shales (Core 60-63) seems to be of special interest for hydrocarbon generation and migration. Indications for a high hydrocarbon potential and/or a migration of released low-molecular organics within this part include (1) the intense odor of oil; (2) the minimum values for C_1/C_2 ; C_2/C_3 ; C_2/C_5 (Fig. 12); inversion of iC_4/nC_4 (Table 4); and absolute maximum fluorescence values (Fig. 11).

Gaseous hydrocarbons encountered in the hole consist predominantly of methane up to 30%, followed by ethane up to 0.4% and higher hydrocarbons, each decreasing in concentration by about one order of magnitude per increasing carbon number. Ratios of C_1/C_2 , C_1/C_2 , and C_2/C_5 follow the same pattern below Core 63, but are different above Core 60.

These features of Cores 63-70, reflecting "in phase" behavior (Whelan, 1979), were also recognized in Site 397: the behavior is apparently a result of an in situ reaction. Migrating hydrocarbons, if leaving the source rock, are not in phase because of the now discernible separation process involved in migration and diffusion. The ratio of iC₅/nC₅ is always 1, except in Cores 60-63 (Table 4). Indeed, this inversion is a result either of a different organic facies or of migration toward barren sediments across a border. If a migration has taken place, light hydrocarbons generated below this border (in phase) are selectively separated when reaching the denser chalks. According to their molecular size and form, volatility, water solubility, and polarity (which are important for migration and diffusion processes to take place), these compounds are gradually allowed to pass the border.

General Observations from Shipboard and Shore Laboratory Studies

Amount and Type of Organic Matter

The black shales of Site 511 (Lithologic Unit 6) exhibit high organic-carbon contents, on average about 3.5% per dry weight. Both microscopy and geochemical analyses reveal that the high organic content is the result mainly of an accumulation of liptinitic (to a great part marine) organic matter. Because of the high content of organic carbon and high amounts of liptinitic ("bituminous") organic matter, the black shales have an excellent potential to generate hydrocarbons from the source rock at stages of a higher maturity. Based on Rock-Eval pyrolysis and microscopy, the type of organic matter present in the shales can be defined as a kerogen Type II.

The organic gross composition of the Aptian-Upper Jurassic shales is relatively homogeneous throughout the cored sections. Nevertheless, a detailed geochemical analysis shows a general decrease in the contribution of land-derived organic matter from the Jurassic toward the lower Albian. This trend is in agreement with the occurrence of marine chalks at the top of the shales.

The change from euxinic conditions toward a betterventilated sedimentary system on the Maurice Ewing Bank is reflected by the rapid drop of organic carbon within a "transitional zone" (Deroo et al., this volume) to lower and very low values at the boundary of the Aptian/Albian. The drop of organic carbon accompanies a general shift from a bituminous organic facies to a nonmarine and residual type of organic matter (von der Dick et al.; Deroo et al.; both this volume).



Figure 12. Gaseous hydrocarbon ratios within the Mesozoic sediments (shipboard measurements).

Table 4. Analysis	of	gas	pockets,	Hole	511,	Cores	52-70.
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	52	54	55	56	57	58	59	60	61	63	64	65	66	67	68	69	70
C1/C2 C2/nC3 C2/nC5 iC4/nC4 neoC5/nC5	./. ./. ./. >10	430 21 >ag3 ./. >10	422 23 >a63 ./. >1	353 13 >10 ³ ./. >10	236 30 >10 ³ <1 ./.	286 29 >10 ³ <1 >10	160 27 660 <1 1	120 7 54 <1 0.5	95 5 44 <1 0.6	152 5 45 >1 1.1	188 6 86 >1 1.4	200 5.4 67 >1 1.4	176 5.7 83 >1 1.2	193 6 100 >1 1.1	158 5 64 >1 1	209 6.5 133 2 1.3	114 4 42 >1 1.3
	-	— Zon	e of Diff	usion an	d Scatter	ing —		Aco	Zone of cumulation	on?	-		Zone	of Genera	ation —		

^a Limits of detection.

The remarkable change in the composition and quantity of organic matter at the Aptian/Albian boundary can be related to the improvement of the water circulation on the Falkland Plateau and in other southern parts of the early South Atlantic Ocean (Reyment, 1980). Apart from a few exceptions, a residual type of organic matter represents the organic fraction of the Albian-Recent sediments. The low organic content (about 0.2-0.3% for Units 1, 3, and 5 and Subunit 2B, and 0.4-0.5% for Subunit 2A and Unit 4) and the dominating inert type of organic carbon indicate environmental conditions opposite to those in the black shales. The slightly higher

accumulation of organic carbon in Subunit 2A and Unit 4 may be related to the high sedimentation rates of these units, which hinder through rapid burial the almost complete oxidation of organic matter. Microscopic analyses made on shore and chemical analyses of the extracts suggest that *no* liquid compounds have migrated upward in the section.

The redistribution of light hydrocarbons has been minimal and is discernible only at major stratigraphic or lithologic boundaries (i.e., the Tertiary/Cretaceous unconformity, or the transition from the black shales to the Cretaceous chalks and claystones).

Thermal Evolution of the Black Shales

The vitrinite reflectance, as an indicator of the thermal history of organic matter, is 0.44% Rm downhole at a depth of 630 meters. Therefore, the shales are still in an immature stage of oil generation. The low amounts of extractable compounds (bitumen or extract) and of light hydrocarbons (Schaefer et al., this volume), as well as the patterns of the hydrocarbons all indicate low maturity. Thus, the distribution of hydrocarbons reflects mainly the type of organic matter rather than a diagenetic transformation.

The gradients of several parameters as indicators of an increasing thermal evolution with depth of burial are quite high, apparently in response to the relatively high thermal gradient of about 7°C/100 m (Langseth and Ludwig, this volume). On the Northwest African continental margin, which has a much lower thermal gradient of 4.2°C/100 m, a comparable vitrinite reflectance is reached at a depth of 1400 meters in sediments of an equivalent age (Cornford et al., 1979). Decreasing atomic N/C and O/C ratios of bitumen point to splitting off of functional groups and correspond to the slow increase of hydrocarbons (normalized to organic carbon) with depth of burial. We expect, therefore, that the zone of oil formation will be quickly reached in the black shales in regions of somewhat deeper burial and comparable thermal gradient (either updip on Maurice Ewing Bank or downdip in the "basinal" province of the Falkland Plateau).

An estimation of the vitrinite reflectance in the Tertiary claystones is 0.26% Rm. The only slightly higher degree of maturity at the top of the Cretaceous black shales does not allow for significant additional overburden at the Cretaceous/Tertiary disconformity of Site 511.

PHYSICAL PROPERTIES

Sound velocities, gravimetric densities, vane shear penetration, and thermal conductivity measurements were taken. In general, sound velocity and gravimetric density were measured on identical samples. Penetration and vane shear were taken as close to the sample as possible without disturbing it. From the basic measurements, acoustic impedance, water content, and porosity have been computed without corrections for salt content. For the wet-bulk density samples, determined by the cylinder technique, a raw, estimated shrinkage coefficient was determined. The procedures for determination of sound velocity and vane shear strength followed Boyce (1976). The results of the physical properties measurements are given in the core descriptions at the end of this chapter.

Wet-Bulk Density

The wet-bulk density data discussed here were determined by gravimetric methods, either by cylinder technique (in soft sediments) or chunk technique. For the uppermost few meters only the GRAPE densities must suffice, because no gravimetric determinations were possible in this section (Table 5). Table 5. GRAPE densities, Cores 1 and 2.

Core/Section	Dist. from Top of Section (m)	GRAPE Density
1-1	0	2.2
	0.4	2.15
	0.5	2.25
	1.5	1.9
1-2	0	1.9
	1.4	1.9
1-3	0	1.8
	1.4	1.6
2-1	0	1.4

sity decreases again until it reaches a value of 1.4 g/cm^3 in Core 2 at the boundary of Oligocene sediments.

The gravimetric data show stable values of 1.5 g/cm^3 down to 100 meters; values then decrease to 1.2 g/cm^3 between 100 and 190 meters below seafloor. The reduction corresponds to an increase in diatom content, although drilling disturbance could also contribute.

In association with the Tertiary/Cretaceous boundary and a change in lithology from diatomaceous oozes to clays, the wet-bulk density changes from 1.2 to 1.6 g/cm³. In the Cretaceous section down to the bottom of the hole, the bulk density of the clays and claystones increases continuously up to 1.9, even in the black shales below 500 meters. Some low porosity values, up to 2.6, are due to calcareous and sandy layers within the black shales.

Water Content and Porosity

The values for water content and porosity discussed here and plotted in the core descriptions are not corrected for salt water. Because of the homogeneous lithology (at least over large parts of the site), water content and porosity correlate quite well.

Porosity decreases from 80% at the top of the section to 75% at 100 meters and then increases again to 80% at the Tertiary/Cretaceous boundary. In this section the normal inverse relation with bulk density is obvious. In the uppermost Cretaceous, the porosity drops within 50 meters down to 55% and then increases rapidly again to 68%. There is some evidence that an increase in calcium carbonate may in part cause the drop in porosity. A slight deviation of the bulk-density curve from the consolidation trend can be recognized during the same interval.

From 250 meters to the bottom of the hole, the porosity of the fine-grained sediments follows a normal consolidation curve similar to bulk density. Some low porosities, especially in the black shales, are due to carbonates, carbonate sands, and sandy layers.

Sound Velocity

The sound velocities are very stable within the Tertiary sediments, with a mean value of 1.6 km/s for the

upper 200 meters. Between 200 and 250 meters a strong increase to 2.5 km/s occurs: it is parallel to the observed drop in porosity in the same interval. The sound velocity is again stable between 250 and 300 meters (1.6-1.8 km/s) and increases to about 2.2 km/s between 300 and 350 meters. From 350 meters down to the bottom of the hole, the fine-grained sediments show a continuous increase in sound velocity from 1.7 to 1.9 km/s. High velocities, mainly in the upper part of the black shales, are caused by calcareous layers and the associated drop in porosity. Large differences between horizontal and vertical velocities are due to shell layers. Down to the Tertiary/Cretaceous boundary the acoustic impedance (velocity \times density) is stable, whereas in the Cretaceous it increases continuously owing to the increase of bulk density; all peaks of the sound velocity curve are apparent in the acoustic impedance data.

Penetration and Shear Strength

Penetration was measured down to 400 meters and shear strength down to 300 meters. The penetration values are very high in the upper 130 meters of the section; this may in part be caused by drilling disturbance. Some lower values are due to higher $CaCO_3$ percentages.

From 150 to 210 meters penetration values are stable near 5 mm, whereas below that point penetrations near 2 mm were measured. The change occurs within the interval with significantly lower porosities. The vane shear values are widely scattered and give no more than a general trend since only a few measurements are valid; most of them did not show a clear shear point.

GEOPHYSICS

Correlation of Seismic Reflectors with Lithology

Site 511 is located on the back slope of a cuesta-type sedimentary ridge in the basin province of the Falkland Plateau. The general configuration of reflectors in this province is that of thick sheets of sediments dipping to the south and terminated updip by erosional truncation (Fig. 2) (Ludwig, this volume).

Correlation of geologic units cored with reflectors of the *Conrad* 21-06 seismic line is shown in Figure 13.

Drilling at Site 511 provided definitive information on the nature and age of the wedge of sediments filling the Falkland Plateau Basin. From analyses of vertical reflection and sonobuoy reflection and refraction data, Ludwig and Rabinowitz (1980) conclude that after the inception of seafloor spreading off the north flank of the Falkland Plateau, the plateau was a broad depression well below sea level, bordered on the west by the Falkland Islands Platform, on the north by a continental rim, on the east by the emerged microcontinental Maurice Ewing Bank, and on the south by the North Scotia Ridge. The basin floor thus formed was underlain either by oceanic crust or attenuated continental crust. The configuration of apparently conformable uniform foresets shown by seismic profiles and the configurations shown by Ludwig et al. (1978) strongly indicate deposition as a single, continuous event. Therefore, Ludwig and Rabinowitz (1980) postulate that the wedge of sediments may represent some post-breakup, prograded slope-type deposition of sediments eroded from the continental rim followed by basin-slope and basinfloor seismic facies units of deep-marine, pelagic, and hemipelagic sediments. The reflection geometry and lithology at drill sites indicate shallow-marine coastal downlap of sediments on a paleoshelf edge, followed by slope front fill, and by draping of pelagic sequences.

Heat Flow

At Site 511, a successful heat flow measurement was obtained for the first time on the Falkland Plateau using the DSDP Downhole Temperature Probe (DHTP), consisting of an Uyeda electronic package and von Herzen thermistors. More than 50 previous attempts by Lamont-Doherty scientists to obtain heat flow measurements through use of thermistor probes attached to a conventional piston corer had failed, owing to a seafloor cover of manganese and glacial material which destroyed the probes.

At Site 511 the DHTP was lowered twice, to sub-bottom depths of 52.5 and 113 meters. Readings of thermistor resistance sampled every minute are stored in the probe and are played back and displayed upon retrieval. Resistance readings are then converted to temperature data using the thermistor calibration data. The heave compensator was used during the drilling; this usually results in good stability of the bottom hole assembly in the unconsolidated sediments penetrated by the probe.

The temperature versus time data of the two lowerings were then plotted. The temperatures in the sediments at their depths sub-bottom define a gradient of 0.074 °C m⁻¹ (Langseth and Ludwig, this volume).

Thermal conductivities were measured on core samples at irregular intervals aboard ship and on shore. These measurements gave a mean value of 0.842 W $^{\circ}C^{-1}m^{-1}$. The heat flow indicated by these observations is 62.3 mW m⁻² (1.49 HFU), a value compatible with the geological evolution of the plateau.

SUMMARY AND CONCLUSIONS

Site 511 is located about 10 km south of Site 330 (DSDP Leg 36) in the basin province of the Falkland Plateau. The site was drilled and continuously cored to a sub-bottom depth of 632 meters. Heat flow measurements were made in unconsolidated sediments at sub-bottom depths of 52.5 and 113 meters.

Lithostratigraphy

From top to bottom, the sediments cored consist of the following lithologic units:

Unit 1-3 meters of gray to olive gray siliceous oozes and thin foraminiferal oozes of Pliocene-Recent age, with manganese nodules and ice-rafted pebbles and grains.

Unit 2—192 meters of section divisible into two subunits. Subunit 2A comprises 182.5 meters of olive to dark gray muddy diatomaceous oozes and muddy nannofossil-diatomaceous oozes spanning the late Eoceneearly Oligocene interval. Subunit 2B encompasses 9.5 meters, of which only the uppermost 72 cm were recov-



Figure 13. Interpretative line drawing of Conrad 21-06 seismic reflection profile (Fig. 2) and correlation with lithology cored at Sites 327, 330, and 511.

ered. It consists of greenish gray pelagic clay with minor glauconite, chert, and volcanic ash, of Paleocene or Eocene age.

Unit 3—14 meters of gray calcareous oozes and zeolitic foraminiferal oozes with occasional chert pebbles and disseminated minor glauconite, of late Campanian(?) to early Maestrichtian age.

Unit 4—203.5 meters of gray to greenish gray zeolitic clays and claystones with intercalations of nannofossil and foraminiferal claystones, ranging in age from Turonian to Campanian-early Maestrichtian.

Unit 5-80 meters of variegated, often reddish brown claystones, nannofossil claystones, and muddy nannofossil chalk, of early Albian to Turonian age.

Unit 6—140 meters of black, massive, thinly laminated mudstones and nannofossil mudstones of Later Jurassic to early Albian age, in places highly petroliferous and indicative of anoxic conditions. Pyrite is common. Belemnite rostra are common, benthic microfossils are essentially absent, and microcoquinas of *Inoceramus* and *Aucellina* occur frequently.

Principal Results

Hiatuses recognized within this succession occur at the Quaternary/Pliocene, Pliocene/early Oligocene, Paleocene or Eocene/Maestrichtian, Cenomanian/Albian, and Upper Jurassic/Lower Cretaceous boundaries. The unconformity that truncates the southward-dipping sheets of sediment on the plateau (Figs. 2 and 13) lies between Units 2 and 3 at Site 511 and probably corresponds to the Paleocene or Eocene/Maestrichtian hiatus, the maximum duration of which was not precisely determined. The unconformity is believed to represent a major erosional surface caused by bottom-current scouring.

Microfossils occur in nearly all of the cores. The most significant biostratigraphic results at Site 511 are the collection of the following assemblages:

1) 183 meters of lower Oligocene and upper Eocene diatomaceous ooze and muddy nannofossil-diatomaceous oozes with exceptionally well preserved and diverse diatoms, radiolarians, and silicoflagellates, together with sufficient numbers of calcareous nannofossils and foraminifers to provide temporal control by correlation with lower-latitude zonations and the well-established New Zealand stages;

2) 183 meters of lower Campanian-Turonian biogenic sediments, sometimes containing rich and diverse planktonic foraminifers and nannofossils not previously described from the high latitudes of the Southern Hemisphere;

3) Continuous cores through most of the Lower Cretaceous-Upper Jurassic black shale and claystone that enabled us to determine stratigraphic subdivisions and to refine paleogeographic reconstructions based on molluskan and microfossil data;

4) Benthic foraminifers and other floral and faunal groups that demonstrate subsidence of the Falkland Plateau with a distinct acceleration at or near the Low-er/Upper Cretaceous boundary.

Accumulation rates of biogenic sediments testify to the highly episodic nature of sedimentation on the Falkland Plateau; periods of exceptionally rapid sedimentation separated by hiatuses or condensed intervals seem to be the norm for the area. High productivity of siliceous plankton during the late Eocene-early Oligocene is indicated by the very high sediment accumulation rate (about 44 m/m.y.) calculated for this interval. A rather high sedimentation rate marks the Campanian interval.

Sediments of Tertiary through Albian age contain only small quantities of organic carbon. In contrast, the underlying black shale of Late Jurassic-Aptian age contains 1.7-4.1% organic carbon (average 3.5%). Organic geochemical investigations and ratios of gaseous hydrocarbons suggest that (1) the black shale has not reached a high degree of maturity; (2) the hydrocarbons presently occurring in the black shale were formed *in situ*, that is, the shale is both the source and host rock; and (3) no significant amount of additional overburden is necessary to explain the degree of organic maturation of the black shales.

Drilling at Site 511 provided data that significantly build on drilling results from Leg 36 on Maurice Ewing Bank. They indicate that, with some exceptions, lithostratigraphic units were essentially continuous across the bank and the basin province of the plateau; interruption of units is due largely to erosion. Further, a major erosional event took place at or near the Paleocene or Eocene/Maestrichtian boundary, implying that a circum-Antarctic Australia current may have been in existance and admitted to the area *prior* to the opening of the Drake Passage in the Oligocene-middle Miocene. This, in turn, implies the existence of a passageway for currents between East and West Antarctica during the Late Cretaceous-early Tertiary, a fact which has already been determined from paleomagnetic data.

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Information on core description sheets, for ALL sites, represents field notes taken aboard ship under time pressure. Some of this information has been refined in accord with postcruise findings, but production schedules prohibit definitive correlation of these sheets with subsequent findings. Thus the reader should be alerted to the occasional ambiguity or discrepancy.



SITE 511



NIT	47	91.0	ARAC	TER							
U PLANKTO FORAMIT	FORAMINIFER	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SAMPLES	LITHOLOGIC DESCRIPTION
lower Oligocene P18-20		Blackites spinosus	Theocyrtis tuberosa	Hemiaulus incisus		3	0.5			-	$\begin{array}{l} \text{MUDDY DIATOMACEOUS OOZE; as in Cors 4, dark gray (2.5Y 4/0); }\\ \text{very slightly bioturbated throughout; diatom content increases downward.}\\ \hline \\ \textbf{SMEAR SLIDE SUMMARY} \\ \hline \\ 100 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0$

	FOS81	AN L CH	ARAG	TER						11	
PLANKTONIC	BENTHIC FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SAMPLES	LITHOLOGIC DESCRIPTION
Iower Oligocene P18-20						2	1.0				Section 1-Section 3, 65 cm: MUDDY DIATOMACEOUS OOZE; gray (2.5Y 5/b) and dark gray (7.5YR 4/0) with lense and stringers of light gray (2.5YN 7/0-2.5Y 7/0) oze containing high nanofosiil percentages. Very tilgh biotrustation throughout. Core-Catcher: MUDDY DIATOMACEOUS OOZE; gray (2.5Y 5/0). SMEAR SLIDE SUMMARY 2.99 2-120 M D Ouartz 1 1 Mica – TR Heavy minerals TR – Clay minerals TR – Clay minerals TR – Clay minerals TR – Stone stolate 1 – Stone stolate 1 – Stone stolate 1 – Stone stolate 1 – CARBONATE BOMB: 2, 61–62 (8.5)

SITE 511

53



RP RP CM CG CM G CC 1-1-SITE 511 HOLE CORE 8 CORED INTERVAL 62.0-71.5 m BIOSTRATIGRAPHIC ZONE AND FOSSIL CHARACTER TIME - ROCK UNIT FOSSILS ARIANS PLANKTONIC PORAMINIFERS FERRATINIC FERRATINIC REGARANCE REGARANCE ANANOFOSSILS READIOLARIANS FLAGELLATES FLAGELLATES FLAGELLATES FLAGELLATES METERS DRILLING DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES GRAPHIC LITHOLOGIC DESCRIPTION LITHOLOGY RM CM CGI CM FGI CC RP NANNOFOSSIL DIATOMACEOUS OOZE; Core-Catcher only; dark gray (2.5Y 4/0), containing Mn nodule (2.5 cm), subangular granite clast and light gray (5Y 7/2) clast of nannofossil-rich diatom ooze. Blackites spinosus Theocyrtis tuberosa lower Oligocene P18-20

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SITE 511 HOLE CORE 9 CORED INTERVAL	71.5–80.0 m	SITE 511 HOLE CORE 10 CORED INTERVAL	80.089.5 m
BIOSTRATIGAMUIC ZONE AND FOSSIL CHARACTER VOLUME SUBJUCTION U INTOLOGY U INTO	LITHOLOGIC DESCRIPTION	ADDRESS AND ADDRESS AD	LITHOLOGIC DESCRIPTION
ососо 1 1 1 1 1 1 1 1 1 1 1 1 1	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	Inver Oligocene	NANNOFOSSIL DIATOMACEOUS OOZE; as in previous cores; dark gray; slighty more calcarcout. High day content, Pebble (2x5x6 cm) of finely laminated fine sandstone at 38–40 cm. SMEAR SLIDE SUMMARY 1.24 D Quartz 1 Clay minerals 20 Zeolites 1 Foraminifers 2 Namoforobil 30 Diatoms 40 Sponge spicules TR
Iower Oligocene P18-20 Blackine spinouus Theorytris tubercas M	CARBONATE BOMB: 3, 111112 (9) MAGNETIC DATA: 2-60 3-60 4-60 5-60 Inclination 50.2 -0.2 -14.4 -17.4 Declination 52.1 175.5 306.7 41.8 Intensity (emu/ccc) 0.140E-06 0.640E-06 0.320E-06 0.160E-06 MAGNETIC DATA: 8-60 7-11 Inclination -26.0 -40.7 Declination 330.2 348.9 Intensity (emu/ccc) 0.150E-06 0.200E-06	SITE 511 HOLE CORE 11 CORED INTERVAL BIOSTRATICAND POSSIL CONARACTER YOOLLOWNYNY UL UNDULWYYDD UNDULWYYDD UNDULWYYDD U NDULWYYDD U NDULWYDD U NDULWYDD	89,5–99.0 m LITHOLOGIC DESCRIPTION
4	0RAIN 512E: 270 (2, 32, 66) 470 (2, 38, 60) 670 (6, 41, 53)	Iower Oligocene P18–20 P18–20 Bileckites spinous Theocyris tuberous 5 Theocyris tuberous 5 Theolyris tuberous 5	$\begin{tabular}{lllllllllllllllllllllllllllllllllll$
		R/ PM RM RP CG CM FG CC	unain 3/22: 1-71 (1, 38, 63) 3-71 (1, 38, 62)

55



× BR	BIOSTRATIGRAPHIC ZONE AND FOSSIL CHARACTER										
TIME - ROC UNIT PLANKTONIC FORAMINIFERS	BENTHIC	NANNOFOSSILS	RADICLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION
fower Otigocene		Slackites spinosus	heocyrtis tuberosa	dw	FG						GRAVEL in Core-Catcher; assorted pabbles [2-7 cm], angular to sub- rounded; granite, fine sandstone, Mn nodules.

TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE AND FOSSIL CHARACTER									Π			
	PLANKTONIC FORAMINIFERS	BENTHIC FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES SAMPLES	SAMPLES	LITHOLOGIC DESCRIPTION	
lower Oligocene	P18-20		Blackites spinosus	Theocyrtis tuberosa			1	0.5	VOID			NANNOFOSSIL DIATOMACEOUS OOZE; as in previous correr; gray (SY 5/1 to 5Y 6/1); homogeneous; slightly mottled; less indurated. Large, fine-grained sandstone obbite in top 7 cm. SMEAR SLIDE SUMMARY 1-66 D Quartz 1 Clay minerals 15 Foraminifers 1 Nannofosuls 25 Diatoms 55 Sponge spicules 1 Micronodules 2 CARBONATE BOMB: 1, 72–74 (8)	
	RP	RP	CM		CM	FG	30	-				GRAIN SIZE: 1-68 (1, 43, 56)	

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GRAIN SIZE:

1-140 (0, 9, 91) 3-5 (0, 9, 91)

5-5 (0, 11, 89)



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UNIT	PLANKTONIC	BENTHIC FORAMINIFERS	NANNOF0581L8	RADICLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING	SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
							,	0.5				• 8 •	CLAYSTONE AND CALCAREOUS CLAYSTONE: predominately dark reddah brown (SYR 3/3) with intervals of reddah brown (SYR 4/3 – Section 1, 8–22 cm), and brown (TSYR 5/2 – Section 1, 22–95 cm and Section 2, 0–95 cm), almost all of which has a light blue (58 //6) mortling. The CALCAREOUS OOZE in Section 5, 51–150 cm is highly variegated in color, in intervals from 1–17 cm thick. Colors range through tones of brown, grav, pinking grav, pale red purple, grav blue green, and pale blue green. Burrowing is slipht. Aucelfine tragments accur in Section 3 and Section 4, 73–89 cm. Incoersmut and unidentified theil material are present in Section 8, and selitary horn corals are well preserved in Section 4, 99–122 cm and throughout Section 5.
			5	- 1	1	11	2	1			1		SMEAR SLIDE SUMMARY
			no dat					- Contraction					1.82 2.28 4.40 5.59 5.78 5.118 6.20 D
	semblage		ile				3	confinen-				м	Carbonate unspec. TR 40 40 46 63 25 Foruminifers TR TR 2 5 2 TR Nanofossits TR 5 3 10 20 20 Volcanic glass TR TR TR - - Incorranue fragments TR 20 15 - -
	ella kingi As		us tourriseifi					- daries					CARBONATE BOMB: 5, 38-40 (13) 5, 99-101 (68) MAGNETIC DATA: 1,103 244 6-21
	Hedberg		Eiffellith				4	- free					Inclination -66.8 -65.6 -70.5 Declination 5.3 24.2 156.8 Intentity (armu/cc) 0.399E04 0.177E-04 0.143E-04 GRAIN SIZE:
								- tranta		1		OG	3-21 (0, 7, 92) 5-21 (1, 10, 89)
Cenomanian	RP	RP					5	and an				•	
Dian	RP	RP								1		•	
All	CM	RM	СМ				6				23	м	

SITE	51	1	HOL	E			CC	RE	50 CORED	INT	ER	VAL	432.5-442.0 m							
	BIO	STR/	TIG	TAPH	IC Z	ONE			1	Г										
×	- 1	oss	IL CH	ARA	CTE	R														
TIME - ROC UNIT	FLANKTONIC FORAMINIFERS	BENTHIC FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GR APHIC LITHOLOGY	DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	U	THOL	OGIC DI	ESCRI	PTION			
Albian	Hedbergelia planispira-Hedbergella delroiensis Assemblage		Prediscosphaera cretacea Eiffellithus tourriseiffeli				2	0.5			11	• M	CALCAREOUS CLAY Section 1, 0–94 cm: ; dith brown 12.5XR 3 Interval highly fossilite corats. Some brownia 5(4). Section 1, 94 cm—Sec calcareout shak; moj orata. Section 2, 105- and quartz grains. Yel 0–70 cm. Section 2, 70 cm—Sec careous chalk with wer and 78–81 cm. Modera Core-Catcher: as in Sec SMEAR SLIDE SUMM Quartz Fieldspar Clay minerals Zeolites Carbonst unspec.	STONE eddish (4); lig rous - red 1 tion 2, ferately lowish tion 3, ly fos tion 4, y dark taily bo tion 4, y dark taily bo tion 4, y dark 1-106 c lowish tion 3, ly fos tion 3, ly fos tion 4, y dark 1-106 c lowish tion 4, y dark 1-106 c lowish	AND h brown ht gray small : cones gr y burror m – wh brown 42 cm: gray (5) 5 5 - 24 - 35	MUDD' (2.5YF burrow < 1 cm ading i : graen wed; in the inte (10YR ; sbur (10YR ; sbur sparse 2.677 D 3 2 2 54 1 10	Y CALC 4 (4/4); r ws (57) pelec pelecyp pelecyp 2-104 D 2 - 58 2 15	AREO op 6 c 7/1) at 156 (s pelled astronation ones in 1 (2.5°) Sectio ads. 3.60 D 1 1 3.8 - 10	US CH4 m dark 6-11 sown (10 5/1) mu cypods ate gran Section (R 4/4) s Section muddy n 3, 55 3-131 D 1 - 28 1 35	ALK red- cm, tary ty/R ddy and ules and cal- 60 1 - 36 1 2
	A/ PM	FM	СМ				4	a charachara a	Void	1	11	•	Foraminifers Nannofosils Micromodules Volcanic glass CARBOATE BOMB: 1, 107–109 (11) 3, 61–62 (24) MACNETIC DATA: Inclination Declination Intensity (emu/cc) GRAIN SIZE: 1, 21 (3, 21, 76) 3, 21 (7, 24, 66)	5 25 - - - - - - - - 67.2 305.5 0.8	20 15 1 -	10 20 - - - 50.7 125.9 5 0.2	5 20 - - 42E-04	15 35 -	10 25 TR	5 55 - -

SITE 511

510	STR/	ATIGR	ID	IC Z	ONE													
FORAMINIFLES	BENTHIC FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLES	u	THOL	OGIC D	ESCRI	PTION			
						1	0.5			•	MUDDY CALCAREOL Section 1, 027 cm: in top 9 cm, light green (~1 cm) and shell debr Section 1, 27-118 c	US CHA muddy ish gray is. Shar m: intu	ALK AN chalk (5GY) p angula erbedde	D CAI (coquii 5/1) be r lowe d red	LCARE na); red low, Ab r contac to red	OUS CL (2.5Y) rundant t. brown	AYSTO R 5/6- pelecy; (5YR !)NE 4/6) xods 5/3)
								0000			chalk as above and calc Section 1, 118-143 c	areous m: red	clayston brown	e. calcare	ous clay	istone v	vith 0.5	cm
								- 00000	24		Section 1, 143 cm-S	ection	2, 16	am: re	d brow	n chalk	(coqui	na).
							1.1		1	м	Section 2, 16-129 c claystone with some lig	m: red ht gray	(1,5Y) (5Y 7/	R 4/6)) zone	to red s; fewer	brown	calcan pods.	NOUS
						2	-		1	•	Section 2, 129-149 c pelecypods.	m: ligh	t gray n	nuddy	chalk (oquina); abund	Sant
mblag									1		Section 2, 149 cm-S claystone; abundant for	ection ssils.	3, 27 c	m; red	brown	, highly	calcan	ous
is Asse						-	-	200	1		Section 3, 27-49 cm to light bluish gray (58	: mude 7/1).	tγ chaĭ	c, light	t greeni	sh gray	(5GY	8/1)
iens							1	POHPOP	$Q \in \mathbb{R}$		Section 3, 49 cm-Sec area muddy chalk loop	tion 5, uinal a	143 cm	: altern	layston	ed brov e.	vn and I	light
(elro		Cea					1	- F FF			Section 5, 143 cm-Se	ection (5, 84 cr	n: alte	mating	red/red	brown	and
14 0		reta				3	- 2	0000		•	greenish gray muddy cf	nalks (c	(sniupo	l. Same				
erge		10 0					-				Section 5, 84-130 cm:	red br	own mu	ddy na	nnotoss	il chaik	ŝ.	
edbe		ohad				1	1				Section 5, 130-135 ch	1: red b	rown m	uddy c	naik (co	Aquinar.		
Ŧ		cost					1.2				SMEAR SLIDE SUMN	ARY						
pira		edis					-					1-20	1-132	2.73	3.75	4-75	5-75	6-110
nisi		E		1	10.1		1	- 0000		1	Quartz	D	1	TR	0	TR	1	2
ã							-	0000			Clay minerals	25	36	39	25	29	38	24
118							1	0000			Zeolites	TR	1	1	2	1	1	2
ĝ						4	1	0000		•	Carbonate unspec.	25	1	5	30	20	15	2
ĝ							-	0000	11		Foraminifers	10	1	50	29	10	40	60
1 er											Organic carbon	40	45	-	-	-	-	-
							-				871.571.9728.18121							
	1					-	-	10000			CARBONATE BOMB							
							1				1, 141-143 [24.5]							
								1000			MAGNETIC DATA:	2-36		4-92	ğ	6-55		
							-				Inclination	-49.9		-60.9	3	-65.0	2	
						5			1		Declination	275.1	IRE O	12.	S ME-DE	124.0	52E_0	5
						1		10000			totensity (emu/cc)	0,4	TOE-U	0.2	140-00	0.0	WKE-U	1
							1				GRAIN SIZE:							
											2.20 (1, 13, 87)							
							-	2000	1		8-20 (15, 24, 61)							
								0000										
								-0000										
				1	L		1 3	0000	1									
						-		and the second se										

M

				 		 0.00	a És	SAM	
				1	0.5		1	M	MUDDY NANNOFOSSIL CHALK; brown (7.5YR 5/4) to reddish brown (5YR 5/4) with greenish gry (5GY 6/1) zones, lenses and burrow fillings increasing downwards. Highly fossiliferous (Incertainur) in Sec- tion 1, less bolevo to Section 5, 18 cm. Section 5, 18–27 cm: grayish purple (SRP 4/2) homogeneous clay- stome. Section 5, 58–66 cm: light gray (SYR 7/1) fossiliferous claystone.
			S	2	The filler		1		Nannofosii chajki in Sections 5, 6, and 7 contain abundam pelecypod shall debis, much dt 1t. Ancecamur. There chaiks are reddiich brow with greenish gray zones as above. Greenish black laminae at Section 8, 57 cm. Note: Section 7 is 67 cm long. SMEAR SLIDE SUMMARY 1.75 2.75 3.55 4.76 5.24 5.64 6
lan	iensis Assemblage	eracea		3	Truction of		1	M	D D D D M M M D Quartz 1 1 1 1 1 1 1 2 2 Feldspar - 1 1 1 1 2 1 - Clay minerals 33 32 22 20 85 81 2 2 Zeolitics 1 1 1 1 1 1 1 3 - Carbonats unsper. 5 2 10 20 - - - 5 Nanofostalis 55 55 55 55 - - 7 Micronodules - - - - 10 10 Volcanic glass - - - - - - 3
lower-middle Alt	nispira-Hedbergella delro	Prediscosphaera c		4	The firm		1		CARBONATE BOMB; 2, 22–23 (28) MAGNETIC DATA: 1-78 3-36 5-53 Indination -61.3 -54.3 -40.2 Decimation 268.4 124.4 245.9 Intensity (emu/cc) 0.389E-05 0.399E-05 0.407E-05 GRAIN SIZE;
	Hedbergella pla				interesting.		1	•	1.42 (12, 30, 58) 3.42 (12, 17, 6) 5.42 (3, 19, 79)
				3	terror terror			0G •	
	C/			6					

CORE 52 CORED INTERVAL 451.5-461.0 m



SITE 511 HOLE

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BIOSTRATIGRAPHIC ZON AND FOSSIL CHARACTER

SITE 511 HOLE

deli middle Albian

C/ PM F M CN

lower pira redisc

TIME - ROCK UNIT

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22 29

22 2 10 2 5 - 70 55 10 1 -3 - -

D 1

SITE 511 H	OLE	CORE	53 CORED	INTERV	L 461.0–470.5 m	SITE	51	1 H	DLE		co	RE	54 CORED INT	ERVAL	470.5–480.0 m
TIME - ROCK UNIT UNIT PLANKING FORAMINIFERS FORAMINIFERS	GRAPHIC 20 AND CHARACTER SING SING SING SING SING SING SING SING	FLAGELLATES A SECTION METERS	GRAPHIC LITHOLOGY	DRILCING DISTURBANCE SEDIMENTARY STRUCTURES	LITHOLOGIC DESCRIPTION	TIME - ROCK	PLANKTONIC FORAMINIFERS	FORAMINIFERS FORAMINIFERS	GRAPHI AND CHARAC SITES STUDIOLOUGH	SILICO-	SECTION	METERS	GRAPHIC LITHOLOGY ONITHING	SEDIMENTARY STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION
lower-middle Albian Kłedbergella planispira-Hedbergella okrostensk Assemblage	Previseoosphera cretacea	1 0./ 1 1.4 2 3 4 5 6 6			VOID MUDDY NANNOFOSSIL CHALK Intercalated and "blotchy" light gays to greenish gay (50Y 6/1) and reddish brown (5YR 5/4) with gay zones becoming more abundant down section. Section 1, 4-8, 30-37, and 145–150 cm: are sparsely fossil/ferous claystones, Abundant pelecypol (<i>Inceranual</i>) fragments throughout Section 1, but passen throughout the core. Burrowing moderate, "told" and "ring" variaties appearing and Choordrites and Zoophycos present in Section 5 and 6. SMEAR SLIDE SUMMARY 1	lower-middle Albian	Rectoergelta planispira-Heckorgelia deiroiendis Assemblage	FM A	M.		1 2 3 4 5 6 000	0.5		11 11 11 11 11 11 11 11 11 11 11 11 11	<text></text>

SITE	511	HC	LE		C	ORE	55	5 CORED I	NTER	VAL	480,0—489.5 m	SITE	511	н	OLE			CORE	56 CORED I	NTE	RVAL	489.5–499.0 m
ME - ROCK UNIT	FO SHITININ	SSIL C SIL	HARA	CTER	ILLATES A	METERS	1	GRAPHIC LITHOLOGY	ING RBANCE ENTARY	ES ST	LITHOLOGIC DESCRIPTION	IE – ROCK UNIT	INIFERS	TRATI	GRAP CHAR	ACTER	ILLATES	METERS	GRAPHIC LITHOLOGY	RANCE	TURES ES	LITHOLOGIC DESCRIPTION
Iower-middle Albian	Ticinella roberti Assemblege	Provisionershaura crititatia Max		Divide and the second se		0.5 1.0				M • • • • • • • • • • • • • • • • • • •	$ \begin{array}{llllllllllllllllllllllllllllllllllll$	lower-middle Albian	Tricinella roberti Assemblage	RM	Flagooliscus angustius Princiscosphaera creitacea nava Baerostalitzed Badiolonians Ado	plat	100	4 5				<section-header><section-header><section-header><section-header><text><text><text><text></text></text></text></text></section-header></section-header></section-header></section-header>

SITE	511	НО	LE		CC	RE	57 COF	REDI	NTER	IVAL	499.0–508.5 m	SITE	51	11	HOL	E		C	DRE	58 CORED	INT	ERVA	508.5–518.0 m
TIME - ROCK UNIT	ORAMINIFERS	ORAMINIFERS SOLUTION	RAPHIC ND HARAC SNVIJETOIDE	ZONE TER SWOLD	SECTION	METERS	GRAPHI	C GY	HILLING ISTURBANCE EDIMENTARY	IRUCTURES	LITHOLOGIC DESCRIPTION	TIME - ROCK UNIT	ANKTONIC B	FOSS SHAINING	AND STISSONONN	RACT	ZONE ER	SECTION	METERS	GRAPHIC LITHOLOGY	ULLING CYUDANCE	BINDRARY BINCTURES MPLES	LITHOLOGIC DESCRIPTION
lower-middle Athan	Ticinela roberti Assemblage	RP	Recrystallized Radiolarians		1 2 3 4 5 6 700	0.5				• • M	<text><text><text><text><text><text></text></text></text></text></text></text>	Barremian-Aptian	Hechergelle signin-Hechergelle infracterates Assemblage		FP			3	0.5			M M	"BLACK SHALE", as in Core 57. Sparsely laminated throughout laminas greenish black or very dark brown (10YR 22), i.e. Section 2, 139–144 cm. Gray and greenish gray zonus in Section 1, 67–68 cm, Section 2, 92–101 cm, Section 3, 114–121 cm, and Section 4, 56–70 and 70–88 cm (greenish black) and 86–89 cm. Burrowing sparse throughout core. Pyrite common in stringers, lamase, and rare small (0.5 cm) nodules. Lanse, burrow fillings and laminae of white calcilutine to fine calci- carenite common - is. Section 2, 30–32, 50–52, and 55–56 cm, Section 3, 52–53 cm; in Section 1, 1, 20–32, 90–52, and 55–56 cm; Section 3, 52–53 cm; in Section 1, 1, 20–32, 90–52, and 55–56 cm; Section 1, 22–26, 53–67, 100–103, 109–110, and 137–138 cm. Accelling(7) sparse through Sections 1, 2, and 4, Paiv-pilov (25.Y B(4) lans, Section 3, 80–62 cm, contains 40% automate. Microfaulting(7) in Section 1, 45–47 cm and Section 3, 96–97 and 101–104 cm. Coquina in Core-Catcher, 11–18 cm. SMEAR SLIDE SUMMARY 1 TR TR TR Clay minaralia 67 115 91 85 Glauconite – – – TR Zaolitik 2 – – – 7 Carbonate unprec. – 5 8 8 3 Foraminifers – 5 TR – Namofostik 3 75 1 – Undetermined minarel 7 – – 10 Volanic glass – TR TR – Namofostik 7 – – 10 Volanic glass – TR TR – Namofostik 7 – – 38.9 Declination +53.7 –48.7 –38.9 Declination +53.7 –48.7 –38.9 Declination +53.7 –48.7 –38.9 Declination +53.7 –18.7 Intensity famu/cc) 0.380E–06 0.300E–06 0.200E–06 GRAIN SIZE: 1.84 (0, 16, 84)

SITE	511	HOLE		0	ORE	59 CORED	INTER	AVA	L 518.0–527.5 m	SITE	511	но	DLE			CORE	60 CORED	INTE	RVA	L 527.5-537.0 m
30CK	FOSS	AND L CHAR	ACTER	NE SI	BS	GRAPHIC	3			OCK	BIOS FC	DISSIL C	GRAP AND HAR	CTER	SE SE	Rs	00.0000	w2		
TIME - I UNI	FORAMINIF BENTHIC FORAMINIF	NANNOFOSS	DIATOMS	FLAGELLAT	METE	LITHOLOGY	DISTURBAN	STRUCTURE SAMPLES	LITHOLOGIC DESCRIPTION	TIME - F	FORAMINIFE	FORAMINIFE	RADIOLARIA	DIATOMS	FLAGELLAT	METE	LITHOLOGY	DRILLING	STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION
Barremian-Aptian		- Chinatoosygus literarius		c	0.4 1.0 2 2			- M	 "BLACK SHALE", as in previous core. Soft greenish gray layer, Section 2, 123–125 cm. White laminae of shell debris common through core, i.e. Section 1, 78–37 cm. very common in Section 2, 19 Section 3, 98–100 and 102–110 cm they contain abundant resonotonils. Section 4, 39–51, 54–65, 69–63, and 98–100 cm again with abundant in anchosalli. Layers with large <i>Incoarcence</i> fragments and other shells in Section 3, 112–150 cm and Section 4, 0–38, 107–111, and 124–135 cm, and in CoreCatcher 0–6 and 8–9 cm. Belemite fragments sparse in Section 3 and 4 and CoreCatcher. SMEAR SLIDE SUMMARY 170 455 D M Clay minerals 3 10 Carbonate monotoxili. 3 80 Undetermined mineral 3 – CARBONATE BOMBE 1, 38 3, 105–107 (69) MAGNETIC DATA: 1-34 2.141 491 Inclination 77.9 222.2 188.6 Intensity (emu/cc) 0.230E–06 0.180E–06 0.190E–06	Barremian-Aptian						0.5 1 1.0 2 2 3 3 4 5			M • •	MUDDY CHALK AND CHALK Predominately dark gray (5Y 4/1) muddy chalk, with lesser intervals (Section 1, 43–58 cm; Section 2, 4–17 and 100–145 cm; Section 5, 0–9 and 42–80 cm) of light gray (5Y 6/1) chalk that contains a high percent of nannotosils and carbonate unspecified, and less clay minerals. The rock is highly inducteds. Lamination are common. Pale gray laminae in dark gray muddy chalk are enriched in carbonate. Balemine fragments common and incoceanous forgenetis abundant in laminee in Sections 1, 2, and 4. Bioturbation is sparse to rank. Aucellina (1) fragments sparse in Section 4. Charty layer, Section 5, 51 cm and Core-Catcher: finely to very finely law; Section 5, 80 cm -Section 6, 58 cm and Core-Catcher: finely to very finely law; marked highly pertoilferous (all table) black. (5Y 2,5/1) with brownish oversitione claystore. Section 5, 80 cm -Section 6, 58 cm and Core-Catcher: finely to very finely law; mersital 20 5 25 83 Carbonate unspect 5 93 ds 4 Carbonate unspect 5 93 ds 4 Annotosilis 67 2 30 - 1 Jonceramus fragments 8 7 CARBONATE BOMBE 1,32-34 (3B) 2,125-127 (106) MAGNETIC DATA: 154 3.129 6.52 Intensity (emu/ce) 0.410E-06 0.260E-06 0.130E-06 Optimized in the section 2,10,20 Declination 78,3, -470,0 0.0 Declination 78,3, -470,0 0.0

м

cc

76

SITE 511 HOLE CORE 61 CORED INTERVA	537.0-546.5 m	SITE 511 HOLE CORE 62 CORED INTERV	AL 546.5-556.0 m
BIOSTRATIGRAPHIC ZONE AND		BIOSTRATIGRAPHIC ZONE AND	
TIME – ROCK UNIT / UNIT	LITHOLOGIC DESCRIPTION	LIME - ROUND CONTRACTER ROUND	LITHOLOGIC DESCRIPTION
	BLACK SHALE (CLAYSTONE) Black BY 2,571), finally laminated, highly indurated, petroliferous. Constrained peerlexpood debris. Final gray layers (BY 472) containing finely com- minuted pelecyclo debris. Final gray layers (BY 571) 1.0 – <1.0 cm are present, containing sopresidable zouble in Soction 1, 1.0 – <1.0 cm are present, containing sopresidable zouble in Soction 1, 1.0 – <1.0 cm are present, containing sopresidable zouble in Soction 1, 1.0 – <1.0 cm are present, containing sopresidable zouble in Soction 1, 1.0 – <1.0 cm are present, containing sopresidable zouble in Soction 1, 1.0 – <1.0 cm are present, containing sopresidable zouble in Soction 2, 200 cm - Soction 5, 200 – 1200 cm is less car- bonaterich and very dark gray (BY 371) in color, as in Soction 2 and a. Prini lamina, Soction 4, 25 cm. Conscione calcita and pyrite gray (BY 671) to light gray (BY 771) in Soction 4, 10 cm. SOLE XINTONEX DEAR SLIDE SUMMARY Name and the social so	PRP RP FM 6	MUDSTONE AND NANNOFOSSIL MUDSTONE - Petroliferous; very dark gay: (5Y 2/1), hard, masive; locally faintly bedded, Change from more to fewer nannofosallis is subtle. Stratifica- tion varies from horizontal to 5–10". Olive gray (5Y 20) layer, 1–10 om thick liocated as shown) are zoolitic daystone. They are commonly ourlain by 0.5–1 cm dark gray ing green (52 4/1) lamina. In Section 3, 28–27 cm, is an olive gray (5Y 5/2) layer of calcaranits and from 41–43 cm, a fine to madium guartz and in a dark gray mud mutrix. Belemmite present in Section 5. MEAR SLIDE SUMMARY 166 1.07 11 2.0 10 1 20 1 11 2 12 1 13 2 140 1 14 2 15 1 14 2 15 1 16 1 17 2 18 2 19 1 10 1 120 10 15 2 16 17 140 1 141 1 20 1 </td



M

78





SITE	511	HOLE		C	ORE	69 CORE	DINT	RVAL	. 613.0-622.5 m	SITE	511	но	.E		CORE	70 COR	ED INT	ERV	AL 622.5-632.0 m
TIME - ROCK UNIT	FORAMINIFERS	ATIGRAJ AND SIL CHAR SIL CHAR	ACTER SWOLVID	FLAGELLATES	METERS	GRAPHIC LITHOLOGY	DRILLING	SERUCTURES SAMPLES	LITHOLOGIC DESCRIPTION	TIME - ROCK UNIT	PLANKTONIC FORAMINIFERS	FORAMINIFERS ST P	ARACT SNEILURIO	ZONE ER SILICO.	SECTION	GRAPHIC		SEDIMENTARY	LITHOLOGIC DESCRIPTION
Oxfordiar-lower Tithonian				3	0.5			•	MUDSTONE (BLACK SHALE) Biack (SY 25/1) and finely laminated is in previous cores but much	Callovian–Dxfordian					2 3 4				 MUDSTONE (BLACK SHALE); as in Core 69; black, laminated, with sparse shall debris laminae accept in Sections 4 and 5, whare they are moderately subundant; Al lamina of carbonate sand (Inoceentus 'needled') occurs in Section 5, 36 cm. Aucellina(1) also present in Section 5, 18–21 and 27–35 cm. Machine Section 2, 136 cm: a relatively large, elongate black (shiny) inclusion is ambadded; plant(1) ren. Grav zeolitic layers as shown. MEAR SLIDE SUMMARY 160 0 0 1 160 0 245 5.25 160 0 1 160 0 10 160 0



SITE 511 (HOLE 511)









SITE 511 (HOLE 511)





SITE 511 (HOLE 511)

	27-2	28-1	28-2	28-3	28-4	28-5	28-6	28-7	28,CC	29-1	30-1	30-2
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0 am	31,CC	32-1	32-2	32-3	32-4	32-5	32-6	32,CC	33-1	33-2	33-3	33-4
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-0 cm	33-5	33-6	33,CC	34-1	34-2	34-3	34-4	34-5	34-6	34-7	34,CC	35-1
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SITE 511 (HOLE 511)

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SITE 511 (HOLE 511)

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SITE 511 (HOLE 511)

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