5. SITE 5141

Shipboard Scientific Party²

Date occupied: 8 February 1980; 1709 (beacon dropped) Date departed: 13 February 1980; 0007 (beacon close aboard)

Number of holes: 1

Time on hole: 103 hr.

Position: 46°02.769'S; 26°51.297'W

Water depth (sea level; corrected m, echo-sounding): 4318

Water depth (rig floor; corrected m, echo-sounding): 4328

Bottom felt (m, drill pipe): 4322

Penetration (m): 150.8

Number of cores: 35

Total length of cored section (m): 150.8

Total core recovered (m): 139.4

Core recovery (%): 92

Oldest sediment cored:

Depth sub-bottom (m): 150.8 Nature: Diatomaceous mud Age: Early Pliocene (Gilbert Epoch) Measured velocity (km/s): ~1.6

Principal results: Site 514 is situated on the lower flank of the Mid-Atlantic Ridge to the east of the Argentine Basin and about 250 mi. north of the present-day position of the Polar Front. Pliocene-Quaternary sediments were hydraulic piston cored to a depth of 150.8 meters below the seafloor. They consist of 130.3 meters of muddy diatomaceous ooze and diatomaceous clays underlain by 7.3 meters of gray and greenish gray muds and nanofossil muds, which are gradually replaced downward by 13.2 meters of stiff diatomaceous mud having 35 cm of hard muddy nannofossil chalk.

All high-latitude Pliocene-Quaternary diatom zones and all but one radiolarian zone were recognized. Calcareous plankton species are of low diversity but useful for paleoenvironmental reconstructions. Paleomagnetic measurement of the cores enabled recognition of the Brunhes, Matuyama (with Jaramillo and Olduvai events), Gauss (with Kaena and Mammoth events), and Gilbert (with Cochiti event) epochs. Correlation of the paleomagnetic scale with siliceous microfossil zonations is an important scientific achievement made by possible by the cores drilled at Site 514. The sequence of Pliocene-Quaternary is probably continuous, except for one hiatus in the middle Pliocene spanning approximately 0.8 m.y. and another possible hiatus in the Quaternary of less than 300,000 y. Fluctuations of the Polar Front because of climatic changes were very pronounced. The most southerly positions of the Polar Front occurred in the Pliocene, during warm later Gilbert and middle Gauss epochs. Between these two intervals the Polar Front occupied a more northerly position, the uppermost suggesting cooler conditions during the Gilbert-earlier Gauss epochs. Late Pliocene-Quaternary time is marked by deterioration of climatic conditions with brief warmings near the Pliocene/Quaternary boundary (upper Matuyama) and at the end of the Quaternary (uppermost Matuyama, upper Brunhes).

The sedimentation rates partially reflect these fluctuations of the Polar Front. They decrease markedly from the earliest Pliocene through the Quaternary. The highest rate, 180 m/m.y., occured in the early Pliocene (Gilbert Epoch): the lowest rate, 2.3 m/m.y., occurred in the Quaternary (early Brunhes).

Higher sedimentation rates result from higher clay content in the sediments as well as higher fossil content. The scarcity of reworked microfossils in the clay suggests transportation and deposition by bottom currents. The unusually high sedimentation rates result in elevated concentrations of organic carbon because the material is protected from oxidation through fast burial.

BACKGROUND AND OBJECTIVES

Site 514 lies about 250 mi. north of the present position of the Polar Front (Fig. 1). Together with Site 513, this site provides a transect of the Polar Front at its most northerly inferred position during the late Cenozoic. The two sites make it possible to reconstruct the late Cenozoic history of the Polar Front, including its influence on biogenic productivity, stable isotopes, and biogeographic fluctuations. Detailed correlations between the two sites will permit us to compare the history of development of this water mass boundary with that established in the Southwest Pacific, where major sediment changes that occurred during the early Neogene reflect the migration of the Polar Front toward its present position. In the Southwest Pacific, calcareous and siliceous biogenic sediments of Oligocene age are replaced by purely siliceous biogenic sediments by the early Miocene. Siliceous biogenic productivity increased during the Neogene as upwelling rates increased. Sedimentation rates monitor the increase in turnover of the oceans toward a peak in the Quaternary. In the South Atlantic, Sites 513 and 514 are well located to study this evolution in a different sector of the Southern Ocean.

SURVEY AND OPERATIONS

The transit from Site 513 to Site 514 was completed in 20 hr. Two icebergs were passed at a safe distance and were the last to be seen on the voyage. A beacon was dropped at the new location, 130 mi. to the northwest at 1709 hr., 8 February, after a 4-hr. survey.

The *Challenger* arrived at the proposed site, located on the east-west-oriented *Conrad* 12-13 seismic line, at 1830 hr. (Fig. 2). Here the *Challenger* seismic line approaching the site (Fig. 3) did not resemble the *Conrad*

 ¹ Ludwig, W. J., Krasheninnikov, V. A., et al., *Init. Repts. DSDP*, 71: Washington (U.S. Govt. Printing Office).
 ² William J. Ludwig (Co-Chief Scientist), Lamont-Doherty Geological Observatory,

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Figure 1. Locations of Site 514 and other Leg 71 drill sites.





Figure 2. Robert D. Conrad 12-13 seismic reflection profile near Site 514. See Figure 4 for location.

West

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Figure 3. Glomar Challenger 71 seismic reflection profile near Site 514. See Figure 4 for location.

cross line, indicating navigational uncertainty arising from efforts to position the ship by dead reckoning between fixes. Rather than survey to match the *Conrad* line in an area of complex sub-bottom topography, we reversed course (Fig. 4) to a location passed earlier, where the sediment reflector pattern consists of a simple, subparallel configuration of reflectors draped over oceanic basement at a depth of about 0.50 s two-way travel time (425 m) below seafloor.

Our coring plan for Site 514 called for hydraulic piston coring of the Neogene sediment section, which was expected to be no more than 200 meters thick. After 43 hr., 150 meters had been cored (Table 1) and the base of the Pliocene sediments had not been reached. We then decided that, given the limited operating time remaining, scientific objectives could best be met by coring the entire section to basement. A round trip was begun for the conversion to a rotary coring bit and bottom hole assembly (BHA).

By the time the bit had reached spud-in position, however, weather conditions had deteriorated to a marginal level, with wind gusts that exceeded 35 kn. Operations were halted in anticipation of better conditions following the passage of the weather front. After 5 hr., the wind had not abated and growing seas were causing the vessel's motion to approach operational limits. The drill string was retrieved, although we continued to hope



Figure 4. Glomar Challenger track approaching and departing Site 514.

Table 1. Coring summary, Site 514.

Core No.	Date (Feb. 1980)	Time	Depth from Drill Floor (m)	Depth below Seafloor (m)	Length Cored (m)	Length Recovered (m)	Core Recovered (%)
1	9	0528	4322.2-4323.4	0-1.2	1.2	1.19	100
2	9	0652	4323.4-4327.8	1.2-5.6	4.4	4.29	97.5
3	9	0802	4327.8-4332.2	5.6-10.0	4.4	3.99	90.7
4	9	0920	4332.2-4336.6	10.0-14.4	4.4	4.53	100
5	9	1030	4336.6-4341.0	14.4-18.8	4.4	4.86	100
6	9	1140	4341.0-4345.4	18.8-23.2	4.4	4.53	100
7	9	1258	4345.4-4349.8	23.2-27.6	4.4	Trace	0
8	9	1415	4349.8-4354.2	27.6-32.0	4.4	3.61	82.0
9	9	1530	4354.2-4358.6	32.0-36.4	4.4	4.36	98.6
10	9	1647	4358.6-4363.0	36.4-40.8	4.4	4.73	100
11	9	1803	4363.0-4367.4	40.8-45.2	4.4	2.36	53.6
12	9	1912	4367.4-4371.8	45.2-49.6	4.4	3.25	73.9
13	9	2030	4371.8-4376.2	49.6-54.0	4.4	4.66	100
14	9	2138	4376.2-4380.6	54.0-58.4	4.4	3.64	82.7
15	9	2255	4380.6-4385.0	58.4-62.8	4.4	4.36	98.6
16	10	0007	4385.0-4389.4	62.8-67.2	4.4	4.54	100
17	10	0124	4389.4-4393.8	67.2-71.6	4.4	4.61	100
18	10	0235	4393.8-4398.2	71.6-76.0	4.4	4.52	100
19	10	0350	4398.2-4402.6	76.0-80.4	4.4	4.37	99.3
20	10	0501	4402.6-4407.0	80.4-84.8	4.4	4.14	94.1
21	10	0604	4407.0-4411.4	84.8-89.2	4.4	4.55	100
22	10	0737	4411.4-4415.8	89.2-93.6	4.4	4.50	100
23	10	0847	4415.8-4420.2	93.6-98.0	4.4	3.42	77.7
24	10	1002	4420.2-4424.6	98.0-102.4	4.4	4.10	93.2
25	10	1110	4424.6-4429.0	102.4-106.8	4.4	4.45	100
26	10	1230	4429.0-4433.4	106.8-111.2	4.4	4.01	91.1
27	10	1345	4433.4-4437.8	111.2-115.6	4.4	2.90	65.9
28	10	1456	4437.8-4442.2	115.6-120.0	4.4	4.18	95.0
29	10	1610	4442.2-4446.6	120.0-124.4	4.4	4.50	100
30	10	1805	4446.6-4451.0	124.4-128.8	4.4	4.31	97.7
31	10	1913	4451.0-4455.4	128.8-133.2	4.4	4.35	98.9
32	10	2028	4455.4-4459.8	133.2-137.6	4.4	4.37	99
33	10	2140	4459.8-4464.2	137.6-142.0	4.4	4.58	100
34	10	2250	4464.2-4468.6	142.0-146.4	4.4	4.23	96.1
35	11	0013	4468.6-4473.0	146.4-150.8	4.4	4.40	100

that conditions would improve and the trip could be reversed. No real signs of improvement were noted until the drill pipe had been recovered and only the BHA remained below the vessel. After another 3-hr. wait, wind, vessel motion, and station-keeping conditions had improved sufficiently for drilling operations to resume. The wind and seas continued to decrease for an additional 3 hr. while the drill pipe was run. Then, within a short time, the wind changed direction nearly 180° and regained its former strength. This caused a confused sea and swell condition which resulted in unavoidable rolling of the ship and disrupted positioning. With the BHA and 104 stands of pipe suspended, a stand of extra-highstrength pipe was picked up and another wait began. When over 4 hr. had elapsed and conditions had not improved enough for spudding, we finally conceded that too much time had been lost to weather conditions for further scientific objectives to be attained.

The drill string was recovered for the final time, and the rig was secured for getting under way to Santos, Brazil. Ironically, when the vessel departed Site 514 at 2328 hr., 12 February, operating weather conditions were quite good and remained so during the 8-day transit to port.

LITHOLOGICAL SUMMARY

General Statement

The Plio-Pleistocene sediments sampled by continuous hydraulic piston coring to a depth of 150.8 meters consist of a single unit of predominantly diatomaceous muds and muddy diatomaceous oozes. The monotonous sequence of gray and greenish gray sediments is characterized by minor but frequent variations both in clay content and in intensity of bioturbation. The site is notable in that it has somewhat higher quartz contents than Holes 513 and 513A. The major features of the lithologic subunits are summarized in Figure 5.

Subunit 1A

This subunit of diatomaceous muds and muddy diatomaceous oozes extends to a sub-bottom depth of 130.3 meters.

The surficial part of the subunit (Core 1 to Sample 514-2-1, 130 cm) consists of an olive brown (2.5Y 4/4) to dark grayish brown (2.5Y 4/2) soft diatomaceous mud grading downward into a muddy diatomaceous ooze; it is characterized by minor to moderate bioturbation increasing with depth, disseminated manganese micronodules, and 5-7% disseminated quartz sand and silt. Olive, olive brown, and dark grayish layers, 3-6 cm thick and with sharp upper and lower contacts, occur throughout.

The surficial zone grades downward into greenish gray (5G 6/1 and 5G 4/1) to gray (5Y 5/1) diatomaceous muds and muddy diatomaceous oozes which extend to the base of Core 22 (93.6 m, sub-bottom). These sediments are soft, becoming firmer with depth, and are characterized by sections, from tens of centimeters to meters in thickness, in which abundant, well-defined, green and dark gray laminae (0.5-1 cm) alternate with homogeneous sections, similar in thickness but with little evidence of stratification. Within the well-stratified sequences, thin laminae (0.5 cm) of greenish gray, dark gray, and dusky purple are common; they occur both individually and as pairs. The paired laminae are always greenish gray over dark gray. The homogeneous sections are usually more intensely bioturbated than the well-stratified sections; solid burrows of very dark gray are most abundant, whereas "ring burrows" are encountered only occasionally.

Occasional pebbles, manganese nodules, manganiferous sediments, and disseminated manganese micronodules and sand occur in the upper part of this subunit. A very small lens (approximately 1 cm in diameter) of nearly pure, fine quartz sand was encountered in Sample 514-5-3, 106 cm, and occasional small lenses of very sandy mud occur in Core 8.

Contacts throughout the subunit are generally sharp. The sediment above 28 cm is an olive gray muddy diatomaceous ooze; below this sharp contact the sediment is a greenish gray muddy diatomaceous ooze. Immediately below the contact (28–29 cm), the sediment is somewhat stiffer than the material both above and below.

The clay content increases with depth in Subunit 1A, and the sediments below Core 22 (93.6–130.3 m, subbottom) are all diatomaceous muds that appear to be transitional to the underlying subunit of muds and nannofossil muds. In Cores 28 and 29, clasts (<1 cm), small lenses, and layers (up to 4 cm thick) of pale olive (5Y 6/3) and white (5Y 8/2) hard, calcareous sediment, rich in nannofossils and "unspecified carbonate," were encountered.

Subunit 1B

This 7.3-meter subunit consists of muds and interstratified nannofossil muds ranging in color from gray (5Y 5/1) to dark greenish gray (5G 4/1). Three intervals in the subunit are apparent. The uppermost and lowermost 3 meters (Sections 514-31-2-3, and 514-32-2-3) are a very firm gray (5Y 5/1) mud with minor bioturbation and faint greenish black and gray horizontal laminae. The sediments contain minor concentrations of nannofossils and foraminifers.

Between these muddy intervals is a dark greenish gray (5G 4/1) stiff nannofossil mud (Section 514-32-1) with faint horizontal greenish black and dark gray laminae throughout.

Subunit 1C

The lowermost 13.2 meters cored at this site consist of dark greenish gray (5G 4/1) stiff diatomaceous muds with an intervening 35-cm layer (Sample 514-33-3, 65-100 cm) of hard muddy nannofossil chalk.

The diatomaceous muds are characterized by moderate to intense bioturbation with faint greenish black and gray laminae common throughout. Small clasts of nannofossil mud occur in Section 514-34-1.

The zone of nannofossil chalk is white (5Y 8/1) and is separated from the over- and underlying sediments by sharp contacts; the lower contact is bioturbated. Bioturbation, evident as pale olive mottling, is moderate throughout the chalk.

PALEONTOLOGY

Biostratigraphic Summary

At Site 514 a nearly continuous Pliocene through Quaternary sequence of diatomaceous clays and muddy diatomaceous oozes was hydraulic piston cored to a sub-bottom depth of 150.8 meters. Site 514 is about 250 mi. north of the present-day position of the Polar Front and about 130 mi. northwest of Site 513. A major objective at both sites was to determine the late Cenozoic history of the Polar Front.

Although pertinent groups of microfossils were recovered throughout the section at Site 514, siliceous groups are dominant and occur continuously in all cores; nannoplankton and calcareous benthic and planktonic foraminifers were encountered sporadically throughout the section. Preservation is generally good to moderate for siliceous fossils but, in contrast, poor for the calcareous fauna and flora except in some portions of the lowermost part of the section (below Core 30), where moderately well preserved nannofossils and foraminifers are present.

An abundant and diverse assemblage of diatoms and radiolarians is present in all 35 HPC cores recovered at Site 514. Throughout the hole, species diversity is low among silicoflagellates, which are sparse in Cores 1 through 10 and sparse to common below Core 10. In general, calcareous nannofossil occurrences are limited to a few dissolution-resistant species. Planktonic and benthic foraminifers also exhibit limited diversity.

		Core	Cored Interval (m)	Lithology	Paleon Sc	nagnetic cale	Age	Benthic Foraminifers	Planktonic Foraminifers	Nanno- plankton	Radiolarians	Diatoms	Silico- flagellates
	- 0	2 3 4 5 6	- 5.6 - 10.0 - 14.4 - 18.8 - 23.2		Brunhes	Jaramillo Olduvai	Pleistocene	Rare agglutinated species Barren	G. pachyder- ma, G. inflata Barren		A. denticulata. Stylatractus universus Saturnalis circulus	C. lentiginosus Zone C. elliptipora/ A. ingens Zone R. barboi/N.	sparse peculum culeata adrangula olyactis
		7 8 9 10	27.6 32.0 36.4 40.8		Matuya			Deep-water M. antarctica S. bulloides etc.	uncticulata achyderma	s, es	Eucyrtidium calvertense Zone	Zone C. kolbei/ R. barboi Zone C. vulnificus	D. a D. a M. qu D. p
(m)	50 -	11 12 13 14	- 49.6 - 54.0 - 58.4 - 62.8					Almost barren	a G Barren	nt species . <i>pliopelacu</i> s some sampl		Cosmiodiscus insignis Zone	
om Depth		15 16 17 18 19	- 67.2 - 71.6 - 76.0 - 80.4		SSL		900.0	p-water ntarctica bradyi ulloides ompiliodes lobosa, etc	icticulata inflata scitula illoides	Rare resista <i>elagicus, C.</i> <i>itopora)</i> in		<i>N. weaveri</i> Zone	
Sub-bott	- - 100 -	20 21 22 23	- 84.8 - 89.2 - 93.6 - 98.0 - 102.4		Gai	Kaena	Pliocene	Deee M. ar E. br M. pc M. pc	6. pur 6. 6. 6. bu	(С. р С. lep	<i>Helotholus vema</i> Zone	Coscinodis- cus vulni- ficius/N. in- terfrigidaria Zone	<i>Distephanus</i> <i>boliviensis</i> Zone
	_	25 26 27	- 106.8 - 111.2 - 115.6 - 120.0			Mammoth		Almost barren	Almost barren			<i>N. interfrig- idaria</i> Zone	
	-	28 29 30 31	- 124.4 - 128.8 - 133.2		ilbert	Cocint			C. pupaticulato	0)	<i>N. praeinter- frigidaria</i> Zone	
	- 150 -	32 33 34 35	- 137.6 - 142.0 - 146.4 - 150.8		ö			As in Cores 15–22	G. puncticulata G. inflata G. scitula G. bulloides	D. surculus D. broweri D. variabilis		<i>N. weaveri</i> Zone	



Site 514 provides a unique record of the temporal distribution of Pliocene through Quaternary siliceous microfossils because of their high abundance and diversity, and the high sediment accumulation rates, nearly continuous deposition, and excellent paleomagnetic record. All middle Pliocene through Quaternary diatom and radiolarian zones of high latitudinal zonal schemes (Chen, 1975; Weaver, 1976; McCollum, 1975; Weaver and Gombos, 1981) are recognized.

A revised Pliocene through Quaternary diatom zonation of the southern high latitudes is presented by Ciesielski (this volume) and is based on Site 514 and conventional piston cores. The upper 11 diatom zones of Ciesielski's revised diatom zonation are present in Hole 514.

Previous correlations of the diatom, radiolarian, and silicoflagellate biostratigraphies to the paleomagnetic stratigraphy of Antarctic piston cores (Ciesielski, 1975, 1978, this volume; McCollum, 1975; Weaver, 1976; Weaver and Gombos, 1981) are used to correlate the magnetic polarity sequence of Hole 514 (see Paleomagnetism, this chapter) to the standard paleomagnetic time scale. The magnetic reversal sequence is identified as follows: Brunhes Chron, (Cores 1-3), Matuyama Chron (Cores 3-12), Jaramillo Subchron (partial Core 4), Olduvai Subchron (partial Cores 5 and 6), Gauss Chron (Cores 12-26), Kaena Subchron (partial Core 21, Core 22, partial Core 23), Mammoth Subchron (partial Core 25, Core 26, partial Core 27), Gilbert Chron (Cores 27-35), Cochiti Subchron (partial Core 27, Cores 28 and 29, partial Core 30).

The section collected from Hole 514 appears to be continuous down to the boundary between Core 27 and Core 26. A hiatus is identified within the lower portion of the Nitzschia interfrigidaria diatom Zone, between Sample 514-26-3, 54-56 cm and Sample 514-27-1, 84-86 cm (Fig. 6). Characteristics of the diatom assemblage in the lower N. interfrigidaria Zone (Ciesielski, this volume) and the magnetic polarity record immediately above and below the disconformity indicate a hiatus separating middle Gauss and upper Gilbert Chronozone sediments. The missing interval represents about 700,000 yr. (~3.86-3.16 Ma) and spans a portion of the Mammoth Subchronozone, the entire lower, normal-polarity portion of the Gauss Chronozone, and most of the uppermost, reversed-polarity portion of the Gilbert Chronozone.

Abundant, reworked, siliceous microfossils are found in the upper Gauss Chronozone sediments of Core 19. In two samples from Section 1 (514-19-1, 20-22 cm and 514-19-1, 72-74 cm), uppermost Oligocene to lowermost Miocene diatoms and silicoflagellates outnumber their Gauss-age counterparts. This occurrence is particularly unusual because of the high sediment accumulation rate of upper Gauss Chronozone sediments and the scarcity of reworked forms in the remainder of this portion of the section. Deposition of these reworked microfossils occurred about 2.8 Ma, judging by the average sediment accumulation rate of the later Gauss Chron.

Paleoenvironmental Observations

One of the major objectives of drilling Hole 514 was to trace late Neogene paleomigrations of the Polar Front Zone. This has been done by an analysis of the water mass affinities of the radiolarian assemblages in Hole 514 core-catcher samples (see Radiolarians, this chapter). Figure 6 displays the position of Site 514 relative to the Polar Front through the middle Pliocene. The relative position of the Polar Front is based on Weaver's model (Weaver, this volume) of modern radiolarian biofacies distribution in the Antarctic and in subantarctic areas. The Polar Front migration curve of Figure 7 is correlated to the paleomagnetic stratigraphy of Site 514, a carbonate preservation curve, eustatic cycles, and significant climatic and glacial events.

Figure 6 illustrates that Site 514 was north of the Polar Front for most of the late Gilbert Chron—between ~4.1 and 3.86 Ma—and for the middle to late Gauss Chron—from ~3.16 to ~2.8 Ma. These two relatively warm periods were interrupted by at least two cooler intervals caused by the northward migrations of the Polar Front over Site 514. The first cool interval occurs within the newly defined *Nitzschia praeinterfrigidaria* diatom Zone (Sample 514-30,CC) and the second within the *N. interfrigidaria* diatom Zone (Sample 514-26,CC). The second of these cool intervals was accompanied by stronger Antarctic Bottom Water (AABW) flow, which caused the nondeposition or erosion responsible for forming the hiatus between Cores 26 and 27.

The relatively warm and stable conditions of the midto late Gauss Chron were ended by the last major northward migration of the Polar Front over the site, from about 2.8–2.6 Ma. This migration corresponds to the suggested onset of glaciation in the Northern Hemisphere (Berggren, 1972), to the development of the first glaciation in the Sierra Nevada Mountains in North America (2.7 Ma; Curray, 1966), and to a lowering of sea level (Vail and Hardenbol, 1979).

High sedimentation accumulation rates during the late Gauss Chron indicate that this northward migration of the Polar Front between 2.8 and 2.6 Ma was not accompanied by AABW velocities as severe as those responsible for the late Gilbert to early Gauss hiatus. The abundant, reworked, siliceous microfossils in Core 19, however, do indicate erosion of nearby upper Oligocene to earliest Miocene sediment by intense AABW flows which transported them to Site 514. Deposition of reworked microfossils in Core 19 occurred about 2.8 Ma and is correlated with the initial stages of the northward migration of the Polar Front. Possible source areas for the reworked microfossils are seen in a major scour zone identified on seismic profiles between Sites 513 and 514 (Ciesielski and Weaver, this volume). After 2.6 Ma, the Polar Front remained near or north of Site 514, except for three southward migrations during the latest Pliocene-Pleistocene to a position just south of the site. The three warming episodes responsible for these southward migrations of the Polar Front occurred just before the Pliocene-Pleistocene transition, just before the Jaramillo Subchron, and shortly before the Matuyama-Brunhes transition.

The relative preservation of calcareous foraminifers and calcareous nannofossils is given in Figure 7. Correlation of the preservation curve to the Polar Front migration curve on the same figure reveals that calcareous



Figure 6. Pliocene-Pleistocene fluctuations of the Polar Front at Site 514, based on radiolarian assemblage characteristics of core-catcher samples. (See Ciesielski and Weaver, this volume, for a more thorough discussion.) Shown also are the paleomagnetic polarity record and a foraminifer and calcareous nannofossil preservation index. Fluctuations are correlated to eustatic cycles and selected paleoenvironmental events and episodes. (In the preservation index, position 1 represents the absence of planktonic foraminifers and calcareous nannofossils, position 2 corresponds to the preservation of a low-diversity assemblage of resistant planktonic foraminifers and a calcareous nannofossil assemblage of dissolution-resistant species, and position 3 represents the presence of common planktonic foraminifers and moderately well preserved calcareous nannofossils, as well as discrete carbonate layers.)

foraminifers and calcareous nannofossils are present only during those intervals when the Polar Front was located south of Site 514. The figure shows that the best preservation of both groups occurs in the basal portion of the section below the Cochiti Subchronozone. The better preservation of the calcareous microfauna and microflora in Cores 32–35 appears to be in part a response to a deeper CCD during deposition of these sediments. The sporadic presence of calcareous foraminifers and calcareous nannofossils elsewhere in Hole 514 is also related to several factors, including Polar Front migrations, increases and decreases in the productivity of calcareous microfauna and microflora in the surface water, and fluctuations of the CCD.

For a more detailed discussion of the paleoenvironment of Site 514, see the chapter by Ciesielski and Weaver (this volume).

Diatoms

Abundant and diverse diatom assemblages of early Pliocene to Quaternary age occur throughout all of Hole 514. This continuously cored HPC section is characterized by good preservation of siliceous microfossils, high sediment accumulation rates, nearly continuous deposition, and an excellent paleomagnetic record.

Site 514 and several piston cores were used to establish a revised diatom zonation of southern high-latitude sediments (Ciesielski, this volume). Thirteen diatom zones, including four new zones and three emended zones, are used to define the Pliocene through the Quaternary. Eleven of these diatom zones are recognized in Hole 514. The stratigraphic occurrence of each zone is as follows: Sample 514-1-1, 47-49 cm through Sample 514-3-1, 72-74 cm, Coscinodiscus lentiginosus Zone;



Figure 7. Abundance, preservation, and species diversity of foraminifers at Site 514. (Abundance: VR, very rare; R, rare; F, few; C, common; A, abundant. Preservation: P, poor; M, medium; G, good.)

Sample 514-3-2, 70-72 cm through Sample 514-5-4, 54-56 cm, C. elliptipora/Actinocyclus ingens Zone; Sample 514-6-1, 77-79 cm through Sample 514-6-2, 80-82 cm, Rhizosolenia barboi/Nitzschia kerguelensis Zone; Sample 514-6-3, 77-79 cm through Sample 514-9-2, 70-72 cm, C. kolbei/R. barboi Zone; Sample 514-9-3, 63-65 cm through Sample 514-12-1, 50-52 cm, C. vulnificus Zone; Sample 514-12-2, 26-28 cm through Sample 514-15-2, 70-72 cm, Cosmiodiscus insignis Zone; Sample 514-15-3, 70-72 cm through Sample 514-19-1, 72-74 cm, N. weaveri Zone; Sample 514-19-2, 72-74 cm through Sample 514-25-1, 95-97 cm, N. interfrigidaria/ Coscinodiscus vulnificus Zone; Sample 514-26-1, 98-100 cm through Sample 514-27-2, 84-86 cm, N. interfrigidaria Zone; Sample 514-28-1, 90-92 cm through Sample 514-32-1, 77-79 cm, N. praeinterfrigidaria Zone; and Sample 514-33-1, 75-77 cm through Sample 514-35-3, 73-75 cm, N. angulata Zone.

All index diatom species found in the subantarctic Hole 514 are also common to antarctic sediments of similar age (Ciesielski, 1978; Weaver and Gombos, in press). The stratigraphic ranges of the zonal guide species of this site are also similar in the Antarctic (Ciesielski, this volume). Previous correlations of index diatom species to paleomagnetic stratigraphy elsewhere in the southern high latitudes (McCollum, 1975; Ciesielski, 1978; Weaver and Gombos, 1981; and Ciesielski, this volume) are used to correlate the magnetic polarity sequence of Hole 514 to the standard paleomagnetic timescale. A single hiatus is identified within the lower portion of the N. interfrigidaria Zone, between Sample 514-26-3, 54-56 cm and Sample 514-27-1, 84-86 cm. The diatom assemblage (Ciesielski, this volume) and magnetic polarity record immediately above and below the disconformity indicate that the missing interval represents a portion of the Mammoth Subchronozone, the entire lower, normal-polarity portion of the Gauss Chronozone, and most of the uppermost, reversed-polarity portion of the Gilbert Chronozone ($\sim 3.86-3.16$ Ma).

Scattered occurrences of rare to few Denticulopsis hustedtii are reworked into Cores 25-30 and Core 34. Common to abundant reworked diatoms of the late Oligocene to earliest Miocene are found in upper Gauss Chronozone sediments of Core 19, Section 1 (20-22 cm and 72-74 cm). Common reworked diatoms in Core 19 include Synedra jouseana, Hemiaulus tauris, Rossiella sp. A, Rocella gelida, R. vigilans, and Asterolampra affinis. Several of these reworked species have stratigraphic ranges restricted to the R. vigilans through R. gelida zones of Gombos and Ciesielski (this volume); both of these date from the latest Oligocene to the early Miocene.

Radiolarians

Radiolarians are present in all 35 HPC cores recovered at Site 514. Abundance and diversity are high and preservation generally good. Age determinations were made using Chen's (1975) and Weaver's (1976) Southern Ocean radiolarian zonation. Thirty-five core-catcher samples from Hole 514 were examined, and in general these zonations are applicable. However, in intervals with warmwater radiolarian fauna, index species are extremely rare or absent. Below Core 30, for instance, no distinction could be made between the *Helotholus vema* and the *Desmospyris spongiosa* zones, because the appearance of *H. vema* in Hole 514 is not a first evolutionary appearance, but rather a climatically induced first appearance.

The stratigraphic occurrence of the radiolarian zones present in Hole 514, as defined by analysis of corecatcher samples, is as follows: Core 1, Antarctissa denticulata Zone ($\sim 200,000 \text{ y. ago}$); Core 2, Stylatractus universus Zone ($\sim 200,000 - \sim 400,000 \text{ y. ago}$) Cores 3-5, Saturnalis circularus Zone (0.7-1.8 Ma); Cores 6-10, Eucyrtidium calvertense Zone (1.8-2.4 Ma); Cores 11-30, H. vema Zone (2.4-3.95 Ma); Cores 31-35, lowermost H. vema Zone and uppermost D. spongiosa Zone, undifferentiated.

The section collected at Site 514 appears to be essentially continuous down through the Core 26/27 boundary, where an unconformity lasting \sim 700,000 y. separates Gauss from Gilbert Chronozone sediments. Details of this hiatus are discussed in Ciesielski (this volume).

One of the main objectives of drilling Hole 514 was to trace migrations of the Polar Front Zone through the late Neogene. It appears from a cursory examination of core-catcher samples at Site 514 that this exercise will be possible. Both warm- and cold-water Pliocene and Pleistocene radiolarian assemblages are intercalated throughout Hole 514 (to a greater degree in the glacial Pliocene and Pleistocene). Basically, four faunas can be recognized throughout Site 514: first, a dominantly antarctic assemblage; second, a dominantly subantarctic/cool temperate assemblage; third, a mixed warm- and cold-water fauna with a majority of warmer-water species; and finally a mixed fauna with a majority of cooler-water species. A generalized model of where these faunas are found in modern antarctic and subantarctic sediments and their position relative to the Polar Front Zone is shown in Figure 6. Utilizing this scheme, a plot of paleomigrations in the Polar Front Zone is illustrated. This curve is correlated to paleomagnetic stratigraphy, eustatic cycles, and significant climatic and glacial events.

Silicoflagellates

The Hole 514 silicoflagellate assemblage is well preserved but of low diversity. Silicoflagellates are sparse throughout most of Cores 1–10 and sparse to common throughout most of Cores 11–35.

Distephanus speculum s.l. is the only species consistently present in Cores 1 through 10. A few Dictyocha aculeata occur in Sample 1,CC, rare to few Mesocena quadrangula occur between Samples 514-4-1, 123-125 cm and 514-4,CC, and common Distephanus polyactis are found in Samples 514-6-1, 77-79 cm through 514-6-3, 77-79 cm.

The occurrences of Dictyocha aculeata, M. quadrangula, and Distephanus polyactis are consistent with the paleomagnetic record that Salloway and Bloemendal (this volume) identified for Cores 1-10. Dictyocha aculeata occurs only in the Brunhes Chronozone of Core 1; Ciesielski (1978) has also noted a restricted occurrence of this species in upper Brunhes Chronozone sediments of the Falkland Plateau. The range of M. quadrangula brackets a normal subchron within Core 4. At lower latitudes this species is a reliable stratigraphic marker with a restricted stratigraphic range, occurring in upper Matuyama Chronozone sediments of several ocean basins. Burckle (1977) noted that the range of M. quadrangula (1.3-0.78 Ma) brackets the Jaramillo Subchron (0.98-0.91 Ma) of the Matuyama. Thus, the range of M. quadrangula agrees with the designation of the short normalpolarity interval within Core 4 as the Jaramillo Subchron of the Matuyama Chron. Distephanus polyactis is present within and slightly below a short normal-polarity interval in lower Core 5 and upper Core 6. Ciesielski (1975, and unpublished data) noted the occurrence of D. polyactis within the Olduvai Subchronozone and lower Matuyama Chronozone, a finding that agrees with a designation of the normal-polarity zone within Cores 5 and 6 as the Olduvai Subchron.

The interval from Sample 514-11-1, 100-102 cm to the base of the hole is identified as the *D. boliviensis* Zone of Ciesielski (1975). Few to common *D. boliviensis* occur throughout the zone. More sporadically pres-

ent are *D. boliviensis* (cannopilean) and *D. quinquangellus*. Both occur more consistently and are more common near the Gauss/Matuyama Chronozone boundary and in Gilbert Chronozone sediments below the Cochiti Subchronozone. In Hole 514 *D. crux* occurs only below the Cochiti Subchronozone. A few *M. elliptica* are present in Cores 26 and 27. *Dictyocha* species occur sporadically above the upper Gilbert-lower Gauss disconformity between Core 26 and Core 27 but are more consistently present below the disconformity.

There is only one significant occurrence of reworked silicoflagellates in Hole 514. Rare *Naviculopsis biapiculata* and few *M. apiculata* occur in two samples from Core 19, Section 1 (514-19-1, 20-22 cm and 514-19-1, 72-74 cm). In both samples these reworked silicoflagellates are accompanied by common reworked diatoms from the late Oligocene and earliest Miocene (Ciesielski, this volume). These common reworked diatoms occur within the upper normal paleomagnetic polarity zone of the Gauss Chronozone, a portion of the Hole 514 sequence with an extremely high sediment accumulation rate. Deposition of the reworked late Oligocene-earliest Miocene siliceous microfossils of Core 19 immediately preceded a major northward migration of the Polar Front over Site 514 (Ciesielski and Weaver, this volume).

Dictyocha species are common in all, or portions, of Cores 2, 28, 29, 31, and 33-35. In all cases, common occurrences of D. species are accompanied by a radiolarian biofacies which indicates that the Polar Front was south of the site at the time of deposition (Fig. 2 in Ciesielski and Weaver, this volume).

Foraminifers

Although planktonic and benthic foraminifers in the Pliocene-Quaternary siliceous oozes and clays of Site 514 are more numerous and diverse than at Site 513, their stratigraphic distribution is sporadic, reflecting paleoenvironmental changes near the Polar Front during the Pliocene-Quaternary. In general, planktonic and benthic foraminifers are rare, very rare, or absent. Exceptions are Samples 514-6,CC and 514-22,CC, where benthic foraminifers are more numerous but never common; Cores 31-33 and Core 19 contain common to few representatives of both groups. Preservation is generally poor except in Cores 31-33, which contain a foraminiferal fauna (Fig. 7) that is moderately or well preserved.

The main distribution patterns of planktonic and benthic foraminifers are as follows.

Core 1 contains only two species of planktonic foraminifers, *Globigerina pachyderma* and *Globorotalia inflata*, and a low-diversity assemblage of benthic species: *Cyclammina pusilla, Martinottiella antarctica, Psammosphaera fusca, Reophax nodulosa*, and *Uvigerina* aff. *dirupta*. This fauna reflects temperate cold water and lower bathyal-abyssal depths.

Cores 2-6 (except Sample 514-6,CC) are barren of planktonic foraminifers and contain only rare fragments of agglutinated benthic species belonging to the genus *Martinottiella*.

Samples 514-6, CC and 514-7, CC include planktonic assemblages of low species diversity, consisting of *Globorotalia puncticulata*, *G. inflata*, and *Globigerina pa*-

pachyderma. Comparatively diverse benthic foraminifers are represented by the arenaceous species *M. antarctica* and *Eggerella* sp., and by deep-water dissolutionresistant calcareous forms (*Melonis affinis, Sphaeroidina bulloides, Pullenia* sp., *Alabaminoides exiguus*, and *Bradynella subglobosa*) which are characteristic of recent abyssal sediments near the CCD level (Saidova, 1976; Khusid, 1978).

Cores 8-14 do not contain foraminifers except in some samples of Cores 8, 11, and 13, where rare specimens of the benthic species *Martinottiella antarctica* were encountered; this testifies to sedimentation well below the CCD.

More diverse and abundant assemblages are recovered in sediments of Cores 15-22. Planktonic foraminifers are represented by Globorotalia puncticulata, G. inflata, G. hirsuta, Globigerina bulloides, and G. pachyderma. The rather diverse benthic assemblages consist of M. antarctica, E. bradyi, Melonis affinis, M. pompilioides, P. quinqueloba, P. bulloides, Gyroidina umbonata, Cibicidoides wuellerstorfi, Alabaminella weddellensis, Alabaminoides exiguus, B. subglobosa, Virgulina sp., and others. Within this interval, however, planktonic and benthic foraminifers are scarce in Cores 16, 17, 18, and 20 and in some samples of Cores 19, 21, and 22. Taken together, these phenomena reflect fluctuations in preservation conditions near the CCD.

Cores 23–30 are almost barren of foraminifers. Only scarce fragments of benthic agglutinated species are found in this interval. Exceptions are Samples 514-28-2, 28–30 cm and 514-28-3, 29–31 cm, which contain arenaceous species and rare specimens of calcareous benthic and planktonic foraminifers.

The most diverse, abundant, and well-preserved foraminiferal assemblages characterize the lower part of the Hole 514 (Cores 31-35). The planktonic assemblage consists of *Globorotalia puncticulata*, *G. inflata*, *G. scitula*, *G.* aff. scitula, Globigerina bulloides, *G. pachyderma*, *G.* aff. apertura, *G.* sp. The benthic forms are represented by the same assemblages as in Cores 15-22, testifying to sedimentation at lower bathyal-abyssal depths above the CCD level and near the planktonic foraminiferal lysocline. Some samples in all these cores contain only rare benthic species.

Because species diversity of planktonic foraminifers is low, this group of microfossils cannot be used for precise age determinations. On the basis of other microfossil groups, Site 514 sediments are Pliocene-Quaternary. According to Jenkins (1971) and Poore (1979), Globorotalia puncticulata disappears at the Pliocene/Quaternary contact. On this basis, the Pliocene/Quaternary boundary could be drawn just above Sample 514-6,CC. But evidence for the age of the overlying sediments is "negative" because, in Cores 2-5, planktonic foraminifers are lacking. Nevertheless other authors (Kennett and Vella, 1975) describe G. puncticulata from Pleistocene sediments. This problem evidently can be solved through an investigation of the distribution of this species in lower-latitude sections that would make it possible to determine the precise level of its evolutionary disappearance.

The alternation, in the Hole 514 section, between layers with comparatively abundant and diverse planktonic and benthic foraminiferal assemblages and layers without any foraminifers or with scarce benthic arenaceous species is evidently related to climatic changes, to increases and decreases in the productivity of calcareous microfauna and microflora in the surface waters, and to deepening and shallowing of the CCD.

Calcareous Nannofossils

Coccoliths were present in only a few of the cores examined from the Plio-Pleistocene section recovered at Site 514. They are sparse except in a few intervals near the bottom of the hole. Preservation was poor to moderate, an indication of deposition close to the CCD. Forms preserved were primarily dissolution-resistant species such as *Coccolithus pelagicus*, *C. pliopelagicus*, and *Cyclococcolithina leptopora*. Such forms were found at various intervals in Cores 6, 7, 15, 19, 22, 28, and 31-33. All other cores were barren except for occasional specimens reworked from older strata.

The only significant accumulation of coccoliths occurred in discrete intervals of Cores 31-33, where some strongly etched nannofossil oozes are present. The greatest diversity was noted in Section 1 of Core 32, which contained a few astroliths, such as *Discoaster surculus*, *D. variabilis*, and a few four- to six-rayed members of the *D. brouweri* plexus. In general, the occurrences of coccoliths in Hole 514 coincided with warm-water maxima or with southerly shifts of the position of the Polar Front as recorded by the radiolarian biofacies (Fig. 6).

PALEOMAGNETISM

At Site 514 the hydraulic piston corer was used to retrieve sediments to a sub-bottom depth of 150.8 meters. Before being split, core sections were measured with the long core spinner. However, the same problems that were seen at Site 512 were encountered here, and no use is made of the long core spinner results.

A total of 477 samples were taken from the sediment in plastic cylinders. These were measured both on board and in Edinburgh, on Digico fluxgate spinner magnetometers. In Edinburgh, each measurement was duplicated; the results given hereafter are the average of two readings. Demagnetization was carried out in Edinburgh using the method described in the Site 511 report.

Natural remanent magnetization intensities are low to moderate $(0.5-5.0 \ \mu\text{G})$, with values of up to $30 \ \mu\text{G}$ at the top of the hole and in Cores 19 to 21. Occasionally single samples with high intensities (>70 μ G) occur, either at the top or bottom of the cores. These samples also have anomalous directions, and it is assumed that these are due to contamination by rust caused during the hydraulic piston coring process (see Site 512 report). Inclinations show clear changes in polarity, so that a magnetostratigraphy can be constructed. Declination values are not consistent between cores, suggesting that relative orientation has not been preserved. They do, however, show consistency within each core and change by about 180° coincident with the polarity reversals that are indicated by inclination. Demagnetization of samples indicates that the majority have stable remanences, probably primary magnetizations reflecting the geomagnetic field at the time of deposition. Inclinations after demagnetization are shown in Figure 8. Inclination averages 51.0° before and 53.7° after demagnetization, which gives a paleolatitude of 34° compared with a present-day latitude of 46° . This represents an inclination error of about 11°.

The Brunhes/Matuvama boundary is identified at a depth of 8.28 meters between Samples 514-3-2, 116-118 cm and 514-3-2, 118-120 cm. The Matuyama/Gauss boundary presumably lies in the unrecovered interval between Cores 11 and 12 (between 42.79 and 46.63 m beneath the seafloor). The Jaramillo and Olduvai events within the Matuyama Epoch are identified, although the former is represented by only two samples. The Jaramillo Event occurs in Core 4 at a depth of 13.2-13.4 meters (Samples 514-4-3, 15-17 cm and 514-4-3, 41-43 cm). The Olduvai Event occurs between Samples 514-5-4, 78-80 cm and 514-6-2, 118-120 cm and appears to be split by a short reversed interval near the beginning. Alternatively the lower normal period represents one of the two Réunion events; and a short hiatus occurs between this and the Olduvai Event. On purely paleomagnetic grounds, the boundary at 125.3 meters (between Samples 514-30-1, 87-89 cm and 514-30-1, 91-93 cm) would be assumed to be the Gauss/Gilbert boundary. However, diatom zonation places Cores 28-35 in the Nitzschia praeinterfrigidaria and N. angulata zones, which correlate with the reversed interval in the Gilbert Epoch between the Cochiti and Nunivak events. A hiatus occurs within Cores 26 and 27; the lowermost Gauss and uppermost Gilbert sediments are missing. The Kaena and Mammoth events are identified, respectively, at depths of 88.9-95.6 meters and between 103.1 meters and the hiatus at 110.6 meters (Samples 514-21-3, 112-114 cm to 514-23-2, 54-56 cm, and 514-25-1, 72-74 cm to 514-26-3, 75-77 cm). The Cochiti Event is seen between the top of Core 27, at a depth of 112.2 meters, and the previously mentioned reversal at 125.3 meters.

A short, normally magnetized event is seen in Core 32 between Samples 514-32-1, 92-94 cm and 514-32-1, 139-141 cm (depths of 134.1 meters and 134.6 m). At an estimated sedimentation rate of 16 cm/1000 y., this event lasted about 3000 y., a much shorter time than the Nunivak Event (150,000 y.; Mankinen and Dalrymple, 1979).

ORGANIC GEOCHEMISTRY

Analyses for Site 514 were the same as those for Site 513. Two "gas pockets" were analyzed, but the composition of the gas was found to be very close to atmospheric gas, with some increase of CO_2 . The gas pockets were apparently a product of core slippage.

Organic carbon values range from 0.2% (Core 2, Section 3) to 0.67% dry weight and are considered to be quite uniform throughout the section, except for Core 2, Section 3, which is close to a postulated hiatus. The uniformity of values is in good agreement with the uniform lithology of this site. The average value of the entire sequence at Site 514 is approximately 1.5 times that



Figure 8. Downhole plot of inclination after demagnetization at 200 Oe, with polarity determinations for Site 514.

of average deep sea sediments (0.3%). Despite great differences in sedimentation rates, the uniformity of carbon content throughout the section implies relatively constant environmental conditions controlling organic carbon accumulation.

Pyrolysis/fluorescence data again showed no increased values beyond the data of blank runs.

Both seawater and interstitial water exhibit higher alkalinity values than Site 513. Alkalinity, measured as liberated CO_2 as well as by titration, increased with depth.

PHYSICAL PROPERTIES

The seismic record at Site 514 indicates conditions similar to the upper seismic and lithologic unit of Site 513. Broad fluctuations in bulk density and water content are correlated with the main sedimentary components, clay minerals and diatoms. Correlation of the seismic record with physical properties and lithological structures seems possible, because all of these appear to be related to the Polar Front fluctuations, as the radiolarian paleontology indicates.

Observations

The wet-bulk density has widely scattered values in the uppermost 20 meters, corresponding to the Pleistocene. Within the Pliocene, bulk density tends to increase with depth down to 185 meters sub-bottom. Fluctuations of 25-50 meters in length are superimposed on the general trend. This pattern becomes more pronounced in water content. The overall trend follows a typical compaction curve between 30 and 140 meters below seafloor.

The porosity values repeat the general trend of the water content, but the superimposed fluctuations are not evident. Recalling that, by definition, water content is equal to (weight water)/(wet weight), that is, it is equal to $(\phi \varrho_w)/[\phi \varrho_w + (1 - \phi \varrho_s)]$, where porosity, $\phi = (\text{volume water})/(\text{wet volume})$, $\varrho_s = \text{mean grain density}$, and w = water, the observed difference between porosity and water content indicates that sinusoidal pattern is related to fluctuations in mean grain density. Some data points which are far outside the main concentration of points belong to thin calcareous layers which occur sporadically in the lower part of the section.

The sonic velocity throughout the section is very stable at about 1.6 km/s (arithmetic mean)—a feature which corresponds very well to Site 513. Some points outside the normal scattering between 1.58 and 1.62 km/s are very probably due to core disturbance. This is especially the case 30 meters below seafloor, where low velocities follow a core with zero recovery. Some higher values (1.66 km/s) at 140 meters are due to the sporadic occurrence of thin calcareous layers in this interval.

The peaks in *acoustic impedance* (velocity \times density) are due to corresponding fluctuations in bulk density, because the sonic velocity is quite uniform.

The general trend in *vane shear values* is nearly linear. A remarkable drop at 40 meters below seafloor to lower values corresponds very well with a reduction in bulk density. Similar results were obtained for penetration values. Only the uppermost 20 meters deviate from the trend, with data points widely scattered because of core disturbance.

Interpretations

Smear slide descriptions for Site 514 clearly relate the fluctuations in bulk density and water content to sediment composition. Clay content and percentage of diatoms are negatively correlated; the clay maxima correspond clearly with maxima in bulk density. In contrast, water content is reduced with increasing clay content and increases with increasing diatom content. The wide scattering of bulk density within the uppermost few meters has its counterpart in the occurrence of higher quartz contents and a narrower period of clay-diatom alternation. In general, the quartz content seems to modulate the physical properties somewhat, even when the contents are very low. The quartz peak at 117 meters appears to correlate with an unconformity that is indicated by the paleontological data. Thus, in general, higher quartz contents may indicate lower sedimentation rates.

The relationship between physical properties and sediment composition also leads to an interpretation of the shear strength values. The drop in vane shear strength at 35 meters below seafloor and the associated increase in clay content indicate a negative correlation between these parameters.

Assuming a velocity of about 1.6 km/s, the impedance peaks at 60 and 100 meters correspond to reflectors at 0.75 and 1.55 s. Similarly the impedance peaks of Site 513 can be related to reflectors at reasonable positions. For both sites, however, a question arises from examination of the impedance values, for the differences are so small that we would not normally expect to find pronounced reflectors. One explanation is that, within these extremely constant velocity profiles, short-path multiples (Anstey, 1977) may build up enough reflected energy to be recognized in the seismic record. A similar acoustic bedding pattern might be produced by alternating sequences of thinly stratified intervals and of intervals that are homogeneous as a result of bioturbation.

SUMMARY AND CONCLUSIONS

Summary

Pliocene–Quaternary sediments were hydraulic piston cored to a depth of 150.8 meters. They consist of a rather monotonous succession of gray, greenish gray, and dark greenish gray diatomaceous clays and muddy diatomaceous oozes. Fine laminations consisting of paired laminae of greenish gray over dark gray sediments are common throughout the section. Bioturbation is also common, but it is mainly of minor intensity. The section is notable for having a higher quartz sand and silt content (5–7% in the upper part of the section) than Holes 513 and 513A. Three subunits can be recognized.

Subunit 1A. This is composed of 130.3 meters of olive brown, greenish gray, and gray muddy diatomaceous oozes and diatomaceous clays. Bioturbation is minor to moderate. Disseminated manganese and quartz silt and sand are present in the upper part of the subunit; lower parts are often finely laminated. The clay content increases downsection. Near the base of the subunit, one core section contains small (~ 4 cm) pale olive clasts high in "unspecified carbonate."

Subunit 1B. This interval consists of 7.3 meters of gray to dark greenish gray faintly laminated firm mud, in which the diatom content was low and in which local concentrations of nannofossils and foraminifers were present; 150 cm of dark greenish gray nannofossil mud were found mid-unit.

Subunit 1C. About 13.2 m of dark greenish gray stiff diatomaceous mud containing one 35-cm layer of whit-

ish muddy nannofossil chalk composed this subunit. Bioturbation is moderate to intense; faint laminae are common throughout.

CONCLUSIONS

All high-latitude Pliocene (mid-Gilbert)–Quaternary diatom zones and all but one radiolarian zone were recognized at Site 514. Calcareous nannoplankton and foraminifers, although of low species diversity and limited stratigraphic value, play an important role in paleoenvironmental reconstructions. Paleomagnetic measurements of the cores identified the Brunhes Epoch, the Matuyama Epoch with the Jaramillo and Olduvai events, the Gauss Epoch with the Kaena and Mammoth events, and the Gilbert Epoch with the Cochiti Event. Correlation of the paleomagnetic time scale with siliceous fossil zonations is perhaps the most significant achievement of drilling at Site 514.

Only one hiatus (within the diatom *Nitzschia inter-frigidaria* Zone and the radiolarian *Helotholus vema* Zone) is obvious in the otherwise continuous Pliocene-Quaternary sequence. The hiatus represents the 0.7 m.y. interval between later Gilbert (3.86 Ma) and earlier Gauss (3.16 Ma) and appears to be of regional significance, since it is contemporary with the time of cooling in East and West Antarctica and with the glaciation in Argentine Patagonia.

Analyses of siliceous and calcareous planktonic groups determine paleoceanographic changes connected with latitudinal migrations of the Polar Front. The Pliocene-Quaternary section recovered from Hole 514 contains antarctic cold-water, subantarctic cool temperate, and mixed cold- and warm-water species. Their succession in time clearly records fluctuations in the position of the Polar Front. On the basis of shipboard and laboratory analyses, it seems likely that the most southerly positions (warmings) of the Polar Front during the ~4.2 m.y. record drilled at Site 514 occurred in the later Gilbert and middle Gauss epochs. These southward fluctuations of the Polar Front are tentatively correlated with oceanographic and/or climatic events in the Tasman Sea (DSDP Site 283), in the Ross Sea (DSDP Site 274), and near East Antarctica (DSDP Site 266). Sediments of these epochs are characterized by cool-temperature subantarctic assemblages of siliceous microfossils, comparatively abundant, well-preserved, and diverse nannofossils, and planktonic and benthic foraminifers.

The warm periods were separated by a cool interval (latest Gilbert-early Gauss epochs) marked by accumulation of diatomaceous clay with cold-water assemblages of siliceous microfossils and without calcareous plankton. Intensification of bottom currents and erosion resulted in the hiatus mentioned earlier. Within the cooling interval, there were other oscillations of a yet undetermined number and amplitude.

In comparison, the late Pliocene to Quaternary interval (upper Gauss, Matuyama, and Brunhes epochs) is a time of pronounced cooling, evidenced by all microfossil groups. The onset of this deterioration corresponds to the beginning of late Pliocene-Quaternary glaciation in the Northern Hemisphere, increasing glaciation in the Southern Hemisphere, and lowering of sea level. Cooling was interrupted by brief warming near the Pliocene-Quaternary boundary (later Matuyama) and at the end of the Ouaternary (latest Matuyama, late Brunhes).

The sedimentation rates during the Pliocene-Quaternary in the area of Site 514 reflect, with minor variances, shifts of the Polar Front (Fig. 9). In the early Pliocene Gilbert Epoch, the rate was enormously high (180 m/ m.y.). During the late Pliocene Gauss Epoch, it decreased to 69.4 m/m.y. A further, sharp diminution to 8.61 m/m.y. took place during the early Pleistocene Matuyama Epoch; in latest Matuyama time, sedimentation increased to 33.4 m/m.y. The lowest sedimentation rate, 2.3 m/m.y., occurred in the early Brunhes; during the late Brunhes, the rate did not exceed 17.5 m/m.y.

The average value of organic carbon in the sediments at Sites 514 is about 1.5 times greater than the average value (0.3%) in modern deep sea sediments. This high organic carbon content is attributed to the high sedimentation rates—fast burial protected the carbon from oxidation—and to the relatively constant environmental conditions in the area of Site 514.

In poorly consolidated, uniform sediments such as diatomaceous clays and oozes, physical properties are the most sensitive to subtle changes that may show up in seismic-reflection records but not in visual examination. Careful lithologic examination and description supply the tool by which variations in the physical properties may be interpreted. For example, evaluation of smear slide data from Site 514 clearly shows the relationship between fluctuations in bulk density and water content, on the one hand, and sediment composition, on the other. Clay content and diatom content are negatively correlated; clay maxima distinctly correspond with maxima in bulk density. In contrast, water content is reduced with increasing clay content, but increases parallel to the diatom content. Wide scattering of bulk density at one horizon at Site 514 is associated with the simultaneous occurrence of relatively high quartz content and a narrower period of clay-diatom alternation. Even when quartz content is very low, it seems in general to modulate the physical properties. Vane shear strengths also appear to decrease with an increase in clay content. Careful correlation of lithology and physical properties undoubtedly aids in the interpretation of the seismic records.

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Figure 9. Sedimentation rates of the Pliocene-Quaternary sediments at Site 514.

810	FOSSI	ANI L CH/	APHI	C ZC	I							
PLANKTONIC	BENTHIC FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SILICO. FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING	SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
Quaternary	RP		S Antarctissa denticulata	Z Z Coscinodiscus lentiginosus	FG	1	0.5	VOID			•	DIATOMACEOUS CLAY Olive brown (2.5Y 4(4) with 3-5 cm zones of olive (5Y 4/3) and gravith brown (2.5Y 4(2). Bioturbation minor to moderate, Diatom content increases downward. Mn nodule, 4 cm, in Core-Catcher. SMEAR SLIDE SUMMARY 1.70 1.101 1.139 D M M Cuartz 10 10 10 Feldspar - TR 5 Heavy minarats 1 45 44 Zeolites 5 3 3 Diatoms 30 40 45 Radiolezians 3 1 3

1.1	F	OSSI	L CH	ARA	CTER													
UNIT - HOC	PLANKTONIC FORAMINIFERS	BENTHIC FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY	STRUCTUBES	SAMPLES	u	THOL	ogić D	ESCRI	PTION	
Quaternary				Antarctissa denticulata	Coscinodiseus lentiginosus		2		<pre></pre>			• M M DOG M	MUDDY DIATOMAC Diatomaceous oazim grayih brown (2,5Y 4 is slight. Minor Mm m diseminated throughous Diatomaceous clay in 1 eoux, locally highly d burrows in Section 1. 39–99 cm in dark gra layering. Greenish bla 3, 0–10 cm. Bioturbati SMEAR SLIDE SUMM Quartz Faldspat Heavy minerals Cabonate untope. Diatoms Radiolarians Micronodules Volcanic glass CARBOMATE BOMB 1, 27–28 (0) MAGHATE DATA: Inclination Intensity (emu/cc) GRAIN SIZE: 1-50 (2, 20, 78) 3-50 (2, 17, 81)	EOUS Sectio incrono incrono salance bark y y mot k (56Y) 7 1 - 3 3 5 - 0 5 0 3 1 1 - - 2 2 8 8 0.6	OOZE n 1, 5-) with 1, 5-) with 1, 5- dules (s Sules (s Sules (s Construction d, Some or 2/1)" or. 22-60 D 7 1 - 46 40 3 1 1 - 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	AND 130 cc minor g set und is dark is dark is large, granula 3.70 D 10 TR - 46 TR - 40 3.71 10 - 5.0.4	DIATO m, is of rayish 2: ey and uisting, very d wry d Wn(?) n 102 cm 1 53 3.777 M 10 53 1 1 53 2 1 3.777 M 10 2 1 3.777 M 10 2 1 3.777 M 10 2 1 3.777 1 1 2 1 1 2 1 1 1 2 1 1 1 1 1 1 1 1 1 1 1 1 1	MACEOUS CLAY ire brown to dark oras. Bioturbation 5–7% quartz sand sharp. V 5/11, homogen- ark gray (SY 3/1) dula in Section 2, . Vague horizontal i-rich zone Section 4-66.3 225.1 5 0.415E–05

SITE 514 HO	LE	CORE	(HPC) 3 COI	REDIN	TERV	AL 5.6-10.0 m	SITE	514	HC	LE	, A	CORE	(HPC) 4 COF	ED I	TERV	AL 10.0-14.4 m
BIOSTRATIO	RAPHIC ZO	VE			П	П			BIOST	RATIC	RAPHIC	ZONE					П	
TIME - ROCK UNIT PLANKTONIC FORAMINIFERS	RADIOLAHIANS DIATOMS	FLAGELLATES	METERS	GRAPHIC LITHOLOGY	DESTURBANCE	SAMPLES	LITHOLOGIC DESCRIPTION	TIME - ROCK UNIT	FORAMINIFERS	FORAMINIFERS	RADIOLARIANS	BIATOMS StLICO-	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
Quatemary	B S Saturnalis circularis B ≳ Coacinodiscus eliptiporal/actinocyclus ingens	1 2 3 FG CC		V010		M • M M M M M M M M	MUDDY DIATOMACEOUS OOZE AND DIATOMACEOUS CLAY Diatomacous clay, Section 1, 14–71 cm, ery, (87 5/1) changing to oilve grap (64 7/2) at 32–39 cm, to greenial gray (56 6/1) at 34–42 cm, Layering at 30° dip. Muddy diatomacous court (Section 1, 71 cm-Section 2, 99 cm), greenial gray (56 Y 6/1-56 6/1), horizontally laminated (~ 1 cm) with cross of more inducated dark greenial ylaminated (~ 1 cm) with cross of more inducated dark greenial ylaminated (~ 1 cm) with cross of more inducated dark greenial ylaminated (~ 1 cm) with cross of more inducated dark greenial ylaminated (~ 1 cm) with cross of more inducated dark greenial ylaminated (~ 1 cm) moderately bioturbated Section 2, 0–51 cm. Gray core, moderately bioturbated Section 2, 0–51 cm. Gray core, moderately bioturbated Section 2, 0–51 cm. Gray core, moderately bioturbated Section 3, 2, 6–72 cm; clive gray 72–144 cm, bioturbated with dissentiated and grains. Mottled greenials gray (50' 6/1) at 121–131 cm. Angular, coansely crystalline acidic igneous pebble (1x2 cm) in a sardy targe Section 2, 90–102 cm. Vague horizontal stratification in Section 3. Large (2 cm) greenish black burtow Section 3, 33–40 cm; small (< 1 cm) pebble at 63 cm.	Outtemery		RP	5 Saturnalis circularis	Coscinodiacus elliptipora/Actimocyclus ingens o	2 3 6 <u>cc</u>		VOID	0 0000000	• • • • • • • • • • • • • •	MUDDY DIATOMACEOUS ODZE AND DIATOMACEOUS CLAY Muddy distomaceous ooze varies in color down section: Section 1, 31–116 cm – dark gray (6Y 4/1) to greenish gray, (6G 6/1); Section 2, 88 cm – Section 3, 1 cm – gray (6Y 5/10; Section 3, 1=44 cm – olve gray: Section 3, 44–01 cm – greenish gray. Bioturbation moderate throughout. Yages horizontal layering common. Dissominated sand common in Section 1, last in Section 3. In section 2, 88 cm – alve gray in the section 3. In section 2, Section 3, 44–01 cm – greenish gray. Bioturbation moderate throughout. Yages horizontal layering common. Inseminated sand common in Section 1, last in Section 3. In staining in upper part of Section 1; disseminated Mn micromodules in Section 2. Distomaceous clay. Section 3, 44–91 cm, greenish gray (5G 6/1), soft with 1–4 cm more inducated green zonse; dark laminae (0.5 cm) at 57 cm. Distomaceous clay. Section 3, 41–91 cm, greenish gray (5G 6/1), soft with 1–4 cm more inducated green zonse; dark laminae (0.5 cm) at 57 cm. Ouartz 7 Interminated quartz and Mn grains in oitwe gray zone at 76–80 cm. Steaminated quarts and Mn grains in oitwe gray zone at 76–80 cm. Ouartz 7 Intermity formacio 1 Ouartz 7 Stillooflagellates 1 Bardolobagins 5 <



SITE 514 HOLE CORE (HPC) 8 CORED INTERV	AL 27.6-32.0 m	SITE 514	HOLE	co	RE (H	C) 9 COREC	INTER	IVAL 32.0-36.4 m
	LITHOLOGIC DESCRIPTION	TIME - ROCK UNIT UNIT PLANKTONIC FORAMINIFERS 509 1000 1000 1000 1000 1000 1000 1000	AND ANNOFOSSILS	DIATOMS a 20	SECTION	GRAPHIC LITHOLOGY	SEDIMENTARY STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION
upper Plicocine 0 upper Plicocine 0 Build of the second	MUDDY DIATOMACEOUS OOZE Dark greenish pray (SGY 4/1), sparsely mottled throughout with very disky purple (SP 2/2) and, in Section 2 and 3, with greenish black (SGY 2/1), Dark gray (SY 4/1) or very dark gray layers or lenses in Section 1, 134–139 cm; Bection 2, 44–98 cm; Section 3, 28–29 cm. The latter instrual is analy, as is 1 cm lens in Section 3, 28–29 cm. The latter instrual is analy, as is 1 cm lens in Section 3, 28–29 cm. The latter instrual is analy, as is 1 cm lens in Section 3, 28–29 cm. The latter instrual is analy, as is 1 cm lens in Section 3, 28–29 cm. The latter instrual is analy, as is 1 cm lens in Section 3, 28–29 cm. The latter instrual is analy, as is 1 cm lens in Section 3, 28–29 cm. The latter instrual is analy, as is 1 cm lens in Section 3, 28–29 cm. The latter instrual is analy, as is 1 cm lens in Section 3, 28–29 cm. The latter instruction 2 is 2 in 1 in Nanofossiis – TR – Diatoms 60 58 60 Radiolariam 4 3 4 4 Silicoflagallates TR TR TR Ebridiami TR TR TR TR Ebridiami TR TR TR TR Ebridiami Socie 5 03.4 9.36 Intensity (emu/cc) 0.284E-06 0.770E-05 0.988E-05 CHAIN SIZE Part (1, 31, 68)	upper Plicocne	Europridium calvartense	COSCINDERSUS KOIDE//MILIOSOBENIA DA/DO/	0.5- 1 1.0 - 2 2 - 3		M - M - M M M M M M M M M M M M M M M M	$\begin{array}{c} \mbox{MUDDY DIATOMACEOUS OOZE} \\ \mbox{Atternating dive gray (SY 4/1) as in previous} \\ \mbox{core, with very dusky purple and greenish black horizontal laminae} \\ \mbox{throughout, except Section 1 where upper 90 cm are highly disturbed} \\ \mbox{and motion with these colors.} \\ \mbox{To T 4 cm of Section 3 has high clay consent and is olive gray.} \\ \mbox{SMEAR SLIDE SUMMARY} \\ & 1.105 2.20 2.40 3.5 \\ \mbox{D M D M} \\ \mbox{Ouartz} & 2 1 1 2 \\ \mbox{Fiddpar} TR & - & - \\ \mbox{Clay minerals} 23 25 22 4.3 \\ \mbox{Curbonste unspec} TR 1 TR TR \\ \mbox{Diatoms} 70 68 71 5.3 \\ \mbox{Rediotarians} 4 4 5 2 \\ \mbox{Slicolagelatarians} 1 1 TR TR \\ \mbox{Dotative STR 1 TR TR TR } \\ \mbox{Dotative STR 1 TR TR TR } \\ \mbox{Dotative STR 1 TR TR TR } \\ \mbox{Diatoms} TR & - & \mbox{TR 2 Slicolagelatarians} \\ \mbox{2, 106-107 (1.5)} \\ \mbox{MAGNETIC DATA: } 1.114 2.46 2.36 2.51 \\ \mbox{Infinition} 121.9 123.3 177.1 130.6 \\ \mbox{Infinition} 121.9 123.3 177.1 130.6 \\ \mbox{Infinition} 124.6 \\ \mbox{Lec Str 1 C DATA: } 2.415 2.446 \\ \mbox{Infinition} 124.6 \\ \mbox{Lec Str 2 O 6 0.83726-05 0.2506-06 0.4506-06 \\ \mbox{MAGNETIC DATA: } 2.35 3.66 \\ \mbox{3.424.7 120.7 \\ \mbox{Infinition} 112.6 \\ \mbox{Infinition} 112.6 \\ \mbox{Infinition} 124.6 \\ \mbox{Infinition} 124.6 \\ \mbox{Infinition} 124.6 \\ \mbox{Infinition} 125.6 \\ \mbox{Infinition} 124.7 120.7 \\ \mbox{Infinition} 126.6 \\ $

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Output of the second

TE 514 H	DLE	0	OR	E (HPC) 11 CO	RED	IN	TERV	/AL 40.8-45.2 m
BIOSTRATI	GRAPHIC Z	ONE						
PLANKTONIC FORAMINIFERS BENTHIC FORAMINIFERS	RADIOLARIANS	SILICO- FLAGELLATES	SECTION	GRAPHIC LITHOLOGY	DISTURBANCE	SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
upper ruccente	음 Helotholus verna 콜 ≥ upper Cosmiodiscus insignis	FG	1 2 CC				• M	MUDDY DIATOMACEOUS COZE As in previous core, predoministry gray (SY 5/1) with greenish black and very duck purple laminae distorted in most part into vertical stringers. Mottling and bioturbation sparse. SMEAR SLIDE SUMMARY 1 110 2:60 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 1100 0 0 0 0 0 0 0 0 0 0 0 1 0 0 0 2 0 33 33 Carbonate unspec. R 1 Diatoms 62 75 Radiolarian 4 2 Stordingeliates 7 7 TR Ebridian TR Ebridians - TR CARBONATE BOMBE: 2.21-22.0) MACHETIC DATA: 1-136 2.44 Inclination 46.2 52.9 Declination 2.52.6 71.4 Intensity (emu/cc) 0.29
		1		12				
BIOSTRATI	GRAPHIC Z	ONE	OR	E (HPC) 14 COI	RED	IN	TERV	AL 40.2-49.0 m
FOSSIL	HARACTE	n			11			
UNIT PLANKTONIC FORAMINIFERS BENTHIC FORAMINIFLRS	RADIOLAHIANS DIATOMS	SILICO- FLAGELLATES	SECTION	GRAPHIC LITHOLOGY	DISTURBANCI	SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
upper Pilocene	Helotholus verna Crasmicodicoras invidente	Antida and a state and a st	2	VOID	0	17 11 11	м.	$\begin{array}{l} \mbox{MUDDY DIATOMACEOUS OOZE} \\ \mbox{As in previous cores, gray, with blike zones in Section 2, 129–131 and 137–142 cm and Section 3, 5–12 cm. \\ \mbox{Green in black and very dusky purple laminae and stringers as before. \\ \mbox{Biurbaline sparse to moderate.} \\ \mbox{Green is dark gray (5Y 4/1).} \\ \mbox{Care-Catcher is dark gray (5Y 4/1).} \\ Care-Catcher is dark$

	BIO	STRA	AN	ARA	IC Z	ONE		e (ne	13
TIME - ROCI UNIT	LANKTONIC	SENTHIC PORAMINIFERS	VANNOFOSSILS	ADIOLARIANS	DIATOMS	SILICO-	SECTION	METERS	GRA LITHO

LE		COR	E (HF	PC) 13 CO	REDIN	TER	VAL 49.6-54.0 m	SITE	5	14 H	IOLE		CO	RE (HPC)	14 COR	ED I	VTERV	/AL 54.0-58.4 m
ND	ZONE				11				BIO	STRAT	AND	HIC Z	ONE					TT	
RADIOLARIANS	SILICONS SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION	TIME - ROCK	PLANKTONIC FORAMINIFERS	BENTHIC FORAMINIFERS	NANNOFOSSILS	SWOLVIG	SILICO- FLAGELLATES	METERS		GRAPHIC LITHOLOGY	DRILLING	STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION
Helotholis vema	Loumidatecus insigms	2	0.5			м	MUDDY DIATOMACEOUS OOZE As in previous cores, olive gray (5Y 4/2) in Section 1; gray (5Y 5/1) in states. astates. astates.	lower Pilocene			D Habitholite werea	Cosmiediscus insignis	4 80	0.5 1 1.0 2 2 3				M .	DIATOMACEOUS MUD Dark greenish pray (SGY 4/1), Greenish black and very dusky purple motiling where highly disturbed; as laminae where not. Black blotur- bations sparse to moderate in Section 2. Very sparse light olive pray (SY 6/2) motiling and sparse black (SY 2/1) mottling in Section 3 and Core-Catcher. SMEAR SLIDE SUMMARY 1-122 3-106 0 0 0 Cartz 3 3 Chy minerals 66 65 Carbons 27 30 Radiolation: 3 1 Silicottogolfates TH TR Sorong fisculat: TH TR Sorong fisculat: TH TR Sorong fisculat: 1-140 2-65 3-68 Infinition -15.1 -4.35 -48.4 Declination 124.2 118.8 137.0 Internity (emu/co) 0.133E-05 0.166E-05 0.180E-05 CRAIN SIZE 1-98 (1, 21, 78) 3-98 (1, 21, 78)
Icel	MIFG	100			1	1													

upper Pliocene

RP

SITE	514	HOLE	C	ORE	(HPC	:) 15 CO	RED	NTE	IVAL 58,4-62,8 m	SITE	51	4 1	IOL	É	C	DRE (H	PC) 16 CO	RED	INTE	RVAL 62.8-67.2 m
TIME - ROCK UNIT	PLANKTONIC FORAMINIFERS	ATIGRAPHIC AND SIL CHARACT SIL CHARACT BIL CHARACT BIL CHARACT BIL CHARACT	SILICO- SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY	STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION	TIME - ROCK UNIT	PLANKTONIC	BENTHIC FORMULTERS	NANNOFOSSILS UNAN	PHIC 2 RADIOLARIANS	SILICO.	SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	STRUCTURES	LITHOLOGIC DESCRIPTION
lower Pliceme	RP R	5 Helotholus verne	V FG	2 3 CCC				M	Diatomacceous MuD Dark greenish grav (BGY 4/1), fairly homogeneous with sparse black (SY 2/11 mottling, in Section 1 and Section 2, 0–66 cm. Bajanco of orer grav (SY 5/1), with sparse black mottling, and very and the grav (BB 5/1) and very ducky purple (B7 2/2) laminae in Section 2, 67–133 cm, Greenish black and very ducky purple laminae in 2–3 in intervals in Section 3, 10–104 cm. SMEAR SLIDE SUMMARY 1/110 2-87 0 Quartz 2 0 1 City minerati 27 66 Curbonate unspect 7R 2 Namofossili TR - Diatoma 20 25 Radiolarians 1 5 Silicoflagillates 1 1 Ebridians TR - Diatoma 20 25 Radiolarians 1 5 Silicoflagillates 2 1 Ebridians TR - Diatoma 20 25 Radiolarians 1 5 Silicoflagillates 2 1 Ebridians 1 5 Silicoflagillates 2 2 Maconetic UDATA: 1-133 2-80 3-64 Inclination -39.5 -60.4 -67.2 Diatoma 27.0 2-48.2 259.9 Intensity (smu/cc) 0.238E-05 0.132E-05 0.195E-05 Silicoflagillates 3 Silicoflagillates 3 Silicofl	lower Plocene	RI	RP		B Helotholus verna 3 ≥ 2		0.5 1 1 1.0 2 2 3		000000000000000000000000000000000000000	72 IL 7 N N N N	DIATOMACEOUS MUD As in Core 15; dark greenish gray with laminae and stringers of greenish black, and very dusky purples. Moderate bioturbation in black, and very dark purple in Sections 2 and 3. SMEAR SLIDE SUMMARY 260 2:120 D Duartz 1 4 Clay minerals 66 70 Carboarts unigers 1 TR Nannofoslis 1 – Distoms 27 23 Radiolarian 3 3 Silicolfagellates 1 TR Sponge spiculus TR TR MAGNETIC DATA: 287 2:121 3:24 Indination -65.1 -56.1 -57.9 Declination 251.3 268.6 269.8 Intensity (emu/cc) 0.202E-05 0.111E-05 0.458E-06 MAGNETIC DATA: 3:54 3:103 Indination -57.6 -61.0 Declination -57.6 -61.0 Declination -27.6 -61.0 Declination 293.4 303.9 Intensity (emu/cc) 0.375E-05 0.577E-05 GRAIN SIZE: 2.71 (2, 25, 73)

SITE 514 HOLE CORE (HPC) 17 CORED INTERVA	L 67.2–71.6 m	SITE 514 HOLE CORE (HPC) 18 CORED INTERVAL 71.6-76.0 m
BIOSTRATIGRAPHIC ZONE		BIOSTRATIGRAPHIC ZONE AND
TIME - ROCK UNIT - ROCK UNIT - ROCK UNIT - ROCK ROCKNING ROCKNIN ROCKNING ROCKNIN ROCKNIN ROCKNING ROCKNING ROCKNING ROC	LITHOLOGIC DESCRIPTION	
	$ \begin{array}{llllllllllllllllllllllllllllllllllll$	OUTOMACEOUS CLAY Data greenith gray (5C 6/1), firm; moderate biotratation, mainly solid portions, Frequent iminas of green, data green over data gray, with operating the greenith gray (5C 6/1), firm; moderate biotratation, mainly solid portions, Frequent iminas of green, data green over data gray, with operating the green layers. Complexicut ring burrens; 11.0–1.5 cm, Stetion 3, 72–80 cm). SMEAR SLIDE SUMMARY 10 M M 10 1 1 1 10 1 1 1 1 10 1 1 1 1 1 10 1 1 1 1 1 1 10 1 1 1 1 1 1 1 10 1 1 1 1 1 1 1 10 1 1 1 1 1 1 1 10 1.0 1 1 1 1 1 1 10 1.0 1 1 1 1 1 1 10 1 1 1 1 1 1 1

SITE	514	HO	LE	0	ORE	(HPC	:) 19 CO	REDI	TEP	IVAL 76.0-80.4 m	SITE	514	4 HC	DLE		CORE	(HPC)	20 CORI	ED II	TERVA	80.484.8 m
	BIOST	RATIG	RAPHIC	ZONE				TT	1			BIOS	STRATE	GRAPHIC	ZONE				T	TT	
×	FOS	SSIL C	ARACTI	ER		- 1		11	1.		×	F	OSSIL C	HARAC	TER		- 1	- 1	- 1		
8	52	BS BS	12	50	Z	\$2		1 mb			8	22	1	1 22	3	S S	22		-	11	
# E	SH	SSI IFE	AIL I	1 F	Ĕ	E I	GRAPHIC	NC NC	9	LITHOLOGIC DESCRIPTION	# E	올문	1 2 3	A		52	5	GRAPHIC	AR	2	LITHOLOGIC DESCRIPTION
μŝ	ON U	NIN S	5 9	1.1	S.	¥.	Lintocour	DNI NO	12		u S	101	UNI S	P	8	E E	E I	LITHOLOGY	NUTRA	2 2	
N.	ANA IL	NN NN	00	A 100	~	~ I		E PR	돌돌		E .	AWK	HAN	100	101	8 9	-	E		12	
	10 BE	NA P	RA	182				SEC DR	2 2			20	E01	A	SIL	21	- 1		DIS	E VE	
_		-		-		-			1			-	++	+-+	-	++	-	VOID	-	+++	
			1.1	1.2		1	VOID		1			1 1		1.1		11	-	Constant of the second	81	11	
			1.1			1		10		DIATOMACEOUS CLAY				1 1		11	-	MUD AND	ŏ	11	DIATOMACEOUS CLAY
			11		11	1	MUD AND	181	1	and a second stop of all and particular and the second stop of the		1 1	11	11		11	-	DIOF	21	11	Section 1, 5-63 cm: soupy mud and pipe rust,
1						-	PIPE	0	1	Section 1, 66 cm-Section 3, 38 cm: dark greenish gray (56 Y 4/1 to				1.1			-	PIPE	ŏ		Remainder plive gray (5Y 4/2) with minor greenish gray (5G 6/1)
1		1	11	1		0.5	RUST	181	1	cation not as prominent as in Core 18, from horizontal (Section 1,			11	1.1		0.	5-	RUST	8	11	in Section 1, becoming greenish gray to dark greenish gray in Sections
			11		11	-		o		125-150 cm) to 30°dip (Section 1, 66-125 cm). Large (2-3 cm)				11		11	-		ŏ	11	pronounced in Sections 2 and 3 with gray and green couplets at 3-5
		11		1		-	~			dark gray burrow Section 1, 144 cm.				11		1.	-		0,,	M	cm intervals in Section 3. More diatomaceous with depth.
1				1.2	11	-	\sim		M	MUDDY DIATOMACEOUS OOZE				1 1		1.1	15	2	1 13	1.1	
	1			1 1		-	~	111	100	Section 3, 38-126 cm and Core-Catcher: greenish gray (5GY 6/1) to				11		11	1-	~	11	1.1	SMEAR SLIDE SUMMARY
						1.0 -	~		1	dark greenish gray (5GY 4/1) with gray and green lamination couplets							0-K=	~	24		1.75 2.75 3.75
1	1.1					4	\sim		1."	as in Core 18. Moderate bioturbation. Very diatomaceous at 110-120						1.1.	1		11		D D D
1						-	~]	11	M	(GM)		1 1		11			1	~	1	M	Ouartz 3 3 10
				1.1	11	-	~	31		MIT AD ALLOS PUMMADY				11		11	F	2	21.		Clay minerals 58 60 39
1		4.4			11	-6	~	111		1.75 2.75 3.75		1 1		11		11	1		1 1	11	Nennofemile - TR TR
					\vdash	-6	~		1.4	D D D				1.1		H			11"		Diatoms 35 30 40
1					[[-6	~	11	1 "	Quartz 5 7 5		1 1		11			15		41	11	Radiolarians 3 5 7
1 1				1.0		-	~	11	I	Feldsper 1 1 -	e			1 1	11	11	15	~	11.	11	Micronodules 1 2 2
					11	-	~	11	1	Heavy minerals – TB TB				1 1		11	15	~	112	M	CARBONATE BOMB
2			1.3	1		-	~		M	Zeolites – 1 TR	G.				E		15	~	11.		2, 59-60 (3)
liao	1	1	da			-	~		M	Diatoms 30 30 45	ě			2	idi		-5	~	Ц.		
Plic			Tio		11	-	~	111	1	Radiolarians 2 3 1	<u>ط</u>			-en	E.	11	-5	-	ъĽ		MAGNETIC DATA: 1-83 1-124 2-32
10			in the			-	~	11	1	Micronadules 2 2 1	- NA			13	in the		15	~	1		Inclination -60.4 -40.4 -63.7
8		1.	in		4	-	~		1 in	CARRONATE ROMB.	2	1 1		101	-	2	5	~	11		Letentity (emular) 0.285E-05 0.439E-05 0.396E-05
-	1	1.1	the pile		11	-	~		1	2, 69-70 (0)		1 1		fot	sci	11	5	~	11.	M	
			elo		11	-	~							Ŧ	Vita	11	-5	~	1 8	1	MAGNETIC DATA: 2-88 3-32 3-89
			E N			1	~			MAGNETIC DATA: 1-80 1-108 1-133 2-10				1 1	-	11	-13	~	11"		Inclination -68.0 -67.6 -66.3
						-	~		1	Inclination -51.3 -31.0 -17.4 -18.4		1 1		1.1		11	-1-2	~	11	H	Declination 49.2 137.8 75.8
£			11		11	-	~		M	Lotensity (amu/or) 0.222E_05_0.430E_05_0.828E_05_0.793E_06				1 1			-1-	~	11		Intentity (encrec) 0.1802-00 0.2002-00 0.1272-00
1 1			11		11	-	~		1			1		1 1		11	-1-	×	1	OG	GRAIN SIZE:
1				1.3	++	-	~	11	M	MAGNETIC DATA: 2-32 2-55 2-80 2-132				11		H	-12	\sim	1 18	H	1-70 (1, 24, 75)
			11			+	~	1	1.00	Inclination -46.9 -70.7 -50.3 -36.7				11		11	-1-	~	11		Stor Sant
1			11	10	11	+	5	1	1	Declination 62.1 110.0 63.4 10.0 Interview (employ) 0.130E=06 0.818E=05 0.118E=04 0.109E=04		1 1	11	1.1	1	11	-1-	~	1	11	
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				1.1		1		1 F	1	MAGNETIC DATA: 3-10 3-32 3-55 3-80				11		11	-1-	~		11	
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					3	1		1 F	1 "	Interview (emulec) 0.973E-05 0.602E-05 0.319E-05 0.165E-05				11		3	12	~	1 1		
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		1				+		111	M	MAGNETIC DATA: 3-108				11		1.1	-77	~	1	1.1	
						-		7: I.	1	Inclination -49.4							17	~	1	M	
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0		1	11			1		TIF	1	intensity tentilities 0.2006.400		RP	RP	CG	AM F	G CC	12	~	1		
1			004	MIRC	-	-		1 i E	Ł.	GRAIN SIZE:											
	nr l'	" N	COA	m no	100	-		11	-	1-85 (1, 28, 71)											
1	1		1.1							3-40 (1, 47, 52)											

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BIOSTRATIC	GRAPHIC 2	ZONE		. War		T	T	Π		×	BIC	FOSSIL	AND	CTER		5	and the theory of			
UNIT ANKTONIC RAMINIFERS NTHIC RAMINIFERS	VDIOLARIANS ATOMS	LICO-	SECTION	METERS	GRAPHIC LITHOLOGY	SILLING .	STURBANCE	AMPLES	LITHOLOGIC DESCRIPTION	TIME - RO	PLANKTONIC FORAMINIFERS	BENTHIC	RADIOLARIAN	DIATOMS	FLAGELLATES	METER	GRAPHIC LITHOLOGY	DISTURBANCE	STRUCTURES SAMFLES	LITHOLOGIC DESCRIPTION
12 Participante de la construcción de la construcci	B Helotholus wama na halotholus wama na	ei ander a	1 2 3 6 6 6 7		VOID MUD ANI PIPE RUST			аа • м м	DIATOMACEOUS CLAY Section 1, 32–86 cm: source of mud, runt, and pebbles tone interface greenish gray (66 (1), firm; moderate to intense biotra- bacion; vague horizontal gray immus in Section 1, 140–150 cons color darker greenish and more prominent in Section 2, Clay color darker greenish gray (66 (21)) in Section 3. Methods and the section 3. Methods and t	Louve Difference			17-Jackbaline vientis	neuorous eene Nitzschie interfrigidaria	1	0.5			TTO: TTO: TTO: TTO: TTO: TTO: W + W I X W I	DIATOMACEOUS CLAY Dark greenish gray (BGY 4/1) to greenish gray (BG 6/1). Moderately to highly biotunbatel; dark gray burrows. Faint heruscontal stratifica- tion; ternines more prominent in Section 2 and 3 with gray/green ouplets as before. Bioturbation less with depth; ring burrow in Section 3, 66 cm. SMEAR SLIDE SUMMARY 1.75 2.75 3.75 D D D D Clay minerels 40 41 53 Zeolites TR 2 TR Diatoms 40 45 40 Radiolatina 7 5 3 Micronodules 2 2 1 CARBONATE BOMB: 1,59–60 (25) MAGNETIC DATA: 1.78 2.45 3.41 3.79 Teclination 100.2 256.4 0.8269.4 Intentity lemulce 0,100E-05 0.184E-05 0.459E-05 0.7 GRAIN SIZE: 1.25 (1, 26, 73) 3.35 (1, 21, 79)







SITE 514 HOLE CORE (HPC) 29	CORED INTERVAL	120.0-124.4 m	SITE	514	но	LE	cc	RE (HP	C) 30 CO	RED I	NTERVAL	124.4-128.8 m
BIOSTRATIGRAPHIC ZOME POSSIL ÉMARACTER VIOL 1990 1000 H - 2000 H -	OSSIL CHARACTER HOLLS STREED OF THE SHAPPHIC UTHOLOGY STREED OF THE SHAPPHIC UTHOLOGY STREED OF THE SHAPPHIC STREED OF THE STREED OF THE				FORAMINIFERS	RAPHIC ND IARACT SNVIUVIOIDAN	SULICO- FLAGELLATES	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENYARY	STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION
1 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10 10		DIATOMACEOUS MUD As in previous corres, dark greenish gray with sparse pale olive mottles; black bioturbations, and lamines sparse in Section 1, more prominent down section. Couplets Interest in Section 3. Pale olive (5Y 6/3) lens high in carbonate in Section 1, 32–37 cm. SMEAR SLIDE SUMMARY $\begin{array}{c} 1 & 30 & 1.100 & 3.56 \\ \hline M & D & D \\ \hline M & D & D \\ \hline Carbonate unspect. 60 & 1 & 1 \\ Olaroms & 8 & 322 & 26 \\ Faidoparts unspect. 60 & 1 & 1 \\ Olaroms & 8 & 322 & 26 \\ Faidoparts unspect. 7 & R & 1 & TR \\ Sponge sploaled T TR & 1 & TR \\ Sponge sploaled T TR & T & TR \\ Sponge sploaled T TR & T & TR \\ 0.56 & -66.3 & -66.0 & -60.3 \\ Declination & -60.8 & -56.5 & -61.0 & -60.3 \\ Declination & -60.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -66.8 & -56.5 & -61.0 & -60.3 \\ Declination & -64.8 & -51.5 & -49.4 \\ Declination & -56.2 & -64.8 & -51.5 \\ Declination & -56.2 & -64.8 & -51.5 \\ Declination & $	lower Pliacene		RP	B Helatholitat www.a		0.5- 1 1 1.0- - - - - - - - - - - - - -	VOID VOID VOID VOID VOID VOID VOID VOID			DIATOMACEOUS MUD As in previous cores, motiling sparse. Laminae sparse in Section 1: taminae couplest common in Sections 2 and 3. SMEAR SLIDE SUMMARY D D Ouartz 2 Chy minerals 68 Ouartz 2 Chy minerals 68 Ouartz 2 Chy minerals 68 Jatoma 75 Chy minerals 68 Jatoma 75 Radiolariams 78 Shicofingelistis 78 Inclimation -71.4 J20E-06 0.270E-06 Declination -11.4 J20E-06 0.270E-06 Declination 71.4 J20E-06 0.270E-06 Declination 70.0 Z5.5 20.2 MAGNETIC DATA: 1.115 L13.1 227 Z48 Inclimation Toto 0.255.5 Declination 70.0 Z5.5 20.2 MaONETIC DATA:<

SITE	514	HOLE		COF	E (HP	C) 31 CO	RED IN	TER	RVAL 128.8-133.2 m	SITE	Ξ	514	HOL	E		ORE (HPC) 32 CO	RED	INTE	RVAL 133.2-137.6 m
DCK	BIOSTF FOS	ATIGRA AND SIL CHA	PHIC ZC	INE Z	50		Π			DCK	-	FOSSIL CHARACTER				2 4				
TIME - RO	PLANKTONIC FORAMINIFEF	NANNOFOSSIL	RADIOLARIAN DIATOMS	FLAGELLATE SECTIO	METER	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY	SAMPLES SAMPLES	LITHOLOGIC DESCRIPTION	TIME - RO	at AMPTORIE	FORAMINIFER	NANNOFOSSIL	RADIOLARIAN	SILICO-	SECTIO	GRAPHIC LITHOLOGY	DRILLING	SEDIMENTAR STRUCTURES CAMPLES	LITHOLOGIC DESCRIPTION
lower Pilocene	RIM F	ис	Helocholus verma/Desmospyris sponglose undifferentiated	2 2 3 80	0.5			M M M M M M	DIATOMACEDUS MUD AND MUD As in previous cores, with Jaminas couplets common in all sections. Black motiles sparse ais pale olive (5V 6/3) mottles, increased carbon- ate content in Section 2 and locally in Section 3. MUD Diatom content markedly decreased in Sections 2 and 3, otherwise the sediment is the same as before. SMEAR SLIDE SUMMARY 1485 2.70 3.30 D D D Cuartz 3 1 1 Clay micerals 77 83 88 Carbonate unpox. 1 5 2. Foraminifers. 1 - 1 - Nanofostils TR 2 - Distoms 15 6 6 0 Radiolarians 4 2 3 Silicolargenetics TR TR TR Sponge sploulms TR TR TR Sponge sploulms TR TR TR CARBONATE BOMB: 2, 4-65 (3) MAGNETIC DATA: 1-45 1-69 1-131 2-45 Inclination 51.7 70.2 57.7 59.9 Declination 209.3 137.6 2.271.1 296.5 Internity (emu/cc) 0.315E-05 0.120E-06 0.300E-06 0.950E- MAGNETIC DATA: 2-89 2-131 3-45 3-89 Internity (emu/cc) 0.580E-06 0.102E-05 0.122E-05 0.340E-0 CRAIN SIZE: 1-131 14.2.3 77)	iower Pilocente		RM FI		A Helotholus venu/Detmospyris spongiosa undifferentiated	M CG	2 2 3	Image: set of the set			NANNOFOSSIL MUD AND MUD Namofossil mud in Section 1, 0–150 cm dark greenith gray (ISG 4/1) with faint, spare, greenith black and very durky purple laminae and must faint, spare, greenith black and very durky purple laminae and must faint but common. Spare black bioturbations. Mud (ISG 4/1) in Sections 2 and 3, with laminae couplers faint but common. Spare black bioturbations. Mud (ISG 4/1) in Sections 2 and 3, with laminae couplers faint but common. Spare black bioturbations. Mud (ISG 4/1) in Sections 2 and 3, with laminae couplers faint but common. Spare black bioturbations. SMEAR SLIDE SUMMARY 190 0 Quartz - - 2 City minerals 74 74 83 75 1 76 - 77 78 78 78 79 10 79 100 79 101 70 173 78 78 79 78 79 78 79 78 70 173 70 173 70 174 71 173 7

SITE 514 HOLE CORE (HPC) 33 CORED	TERVAL 137.6-142.0 m	SITE 514 HOLE CORE (HPC) 34 CORED INTERVAL 142.0-146.4 m
PORTUGATION CERTING AND CERTIFIC AND CERTIFICATION AND CERTIFICATI	LITHOLOGIC DESCRIPTION	
over Plocene 1 10 10 10 10 10 10 10 10 10	M DIATOMACEOUS MUD Dark greating any ISG 4/1); sparse greating have y duky purple laintics, termine acupter sommon. Black mottling throughout. Ring burrows common in Sections 2 and 3. Pale alive mottles sparse to moderate. M MUDDY NANNOFOSSIL CHALK Section 3, 65–100 cm; white ISY 8/11, with moderate pale olive mottling at 88–89 or . Lower contact at 100 cm hater, bioturbated. M SMEAR SLIDE SUMMARY 140 2.00 3.82 3.105 D M D Outrits 3 1 - 3 Clay minerals 76 75 26 66 Carbonate impoc. 1 1 5 1 Foraminifiem - 1 7 1 Nanofosils - TR 55 1 Diatoms 15 17 6 25 Radiolariam 55 1 3 Slicolagellates TR TR TR TR Diatoms 15 17 6 25 Radiolariam 59.6 56.7 64.1 66.5 Declination 59.6 56.7 64.1 66.5 Declination 59.6 56.7 64.1 66.5 Declination 59.6 0.450E-06 0.115E-05 0.640E-0 M MAGNETIC DATA: 1.37 1.53 1.87 1.93 Inclination 59.6 0.460E-06 0.115E-05 0.640E-0 M MAGNETIC DATA: 1.30 2.48 2.65 Intensity (emu/cc) 0.330E-06 0.460E-06 0.310E-06 0.550E-0 M MAGNETIC DATA: 3.10 3.28 3.48 3.51 Inclination 275.5 49.3 25.6 31.2 Declination 31.4 3.20 3.26.0 119.5 76.5 Intensity (emu/cc) 0.330E-06 0.760E-06 0.550E-00 M MAGNETIC DATA: 3.10 3.28 3.48 3.51 Inclination 41.1 33.8 Declination 41.1 33.8 Declination 40.3 75.4 Intensity (emu/cc) 0.340E-06 0.700E-06 0.330E-00 MAGNETIC DATA: 3.109 3.133 Inclination 40.3 75.4 Intensity (emu/cc) 0.3640E-06 0.940E-06 MAGNETIC DATA: 3.109 3.133 Inclination 40.3 75.4 Intensity (emu/cc) 0.3640E-06 0.940E-06 MAGNETIC DATA: 3.109 3.133 Inclination 40.3 75.4 Intensity (emu/cc) 0.3640E-06 0.940E-06 MAGNETIC DATA: 3.109 3.133 Inclination 40.3 75.4 MAGNETIC DATA: 3.109 3.133 Inclination 40.3 75.4 MAGNETIC DATA: 3.	Big Big

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DNIT	PLANKTONIC FORAMINIFERS	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SILICO- FLAGELLATES	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	EAMIN DE	SAMITLES	ינו	THOLOG	GIC DES	CRIPTION		
lower Pilocene				Helotholus vema/Desmospyris spongiosa undifferentiated	Nitzschia angulata		2					M M M M	DIATOMACEOUS MI As in previous core with sparse black (5) throughout. Greenid numerous couplets, in SMEAR SLIDE SUMM Culture Carbonate unspec. Nanofossila Diatoma Radiolarias Sinicoflagellares Sponge splucies CARBONATE BOMB 3, 48–49 (2,5) MACINETIC DATA: Inclination Declination Intensity (emu/cc) GRAIN SIZE: 1-60 (2, 6, 73) 3-50 (1, 29, 71)	UD s, dark Y 2/11 c h black 1800 D 3 77 17 2 77 77 77 77 77 77 77 77 77	greeniah mottling and ve ss 2 and 3 0 2 2 7 5 0 2 2 7 5 1 1 7 6 1 7 5 1 7 7 7 5 1 7 7 7 5 1 7 7 7 7 5 1 7 7 7 5 1 7 7 7 5 1 7 7 5 1 7 5 2 2 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 2 7 5 1 9 2 1 9 2 1 9 2 1 9 2 1 9 2 1 9 2 1 9 2 1 9 2 1 9 2 1 9 2 1 9 2 1 9 2 1 1 9 1 9	gray (5GY 4 throughout, 5 ry dusky pur 3, 2-84 42,4 199,3 0,910E-06 55,8 187,5 0,870E-06	 2.139 2.139 70.5 2.45.3 0.630E-06 3.43 15.9 0.130E-05 	3-12 63.3 222.0 0.0605-4 92.9 0.4806-



SITE 514 (HOLE 514)

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SITE 514 (HOLE 514)





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SITE 514 (HOLE 514)



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